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Impact of upstream landslide on perialpine lake ecosystem: an assessment using multi-temporal satellite data

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Abstract

Monitoring freshwater and wetland systems and their response to stressors of natural or anthropogenic origin is critical for ecosystem conservation.

A multi-temporal set of 87 images acquired by Sentinel-2 satellites over three years (2016-2018) provided quantitative information for assessing the temporal evolution of key ecosystem variables in the perialpine Lake Mezzola (northern Italy), which was suffered from the impacts of a massive landslide that took place upstream of the lake basin in summer 2017.

Sentinel-2 derived products revealed an increase in lake turbidity triggered by the landslide that amounted to twice the average values scored in the years preceding and following the event. Hotspots of turbidity within the lake were in particular highlighted. Moreover both submerged and riparian vegetation showed harmful impacts due to sediment deposition. A partial loss of submerged macrophyte cover was found, with delayed growth and a possible community shift in favor of species adapted to inorganic substrates. Satellite-derived seasonal dynamics showed that exceptional sediment load can overwrite climatic factors in controlling phenology of riparian reed beds, resulting in two consecutive years with shorter than normal growing season, and roughly 20% drop in productivity according to spectral proxies: compared to 2016, senescence came earlier by around 20 days on average in 2017 season, and green-up was delayed by up to 50 days (20 days, on average) in 2018, following the landslide.

The approach presented could be easily implemented for continuous monitoring of similar ecosystems subject to external pressures with periods of high sediment loads.

Keywords: Sentinel-2, Lake Mezzola, helophytes, submerged macrophytes, phenology, turbidity

38 **1. Introduction**

39 Assessing the status of freshwater and wetland ecosystems and its temporal evolution in response to external
40 events of natural or anthropogenic origin occurring in the watershed is a key requirement towards ecosystem
41 conservation. In particular, monitoring aquatic plant communities in these systems is crucial because of their
42 role in biogeochemical processes in water column and sediments (Schindler and Scheuerell, 2002; Jeppesen
43 et al., 2012), their interactions with other autotrophic and heterotrophic organisms (Timms and Moss, 1984;
44 van Donk and van de Bund, 2002; Bolpagni et al., 2014), and the provision of food and shelter to animal
45 communities, such as fish and birds (Johnson and Montalbano, 1984; Carpenter and Lodge, 1986; Wang et
46 al., 2017).

47 Satellite data can provide valuable information on a range of ecosystem processes (Hestir et al., 2015; Zhang
48 et al., 2017; Murray et al., 2018). For more than five decades, Earth Observation (EO) techniques have been
49 used to study inland water quality and conditions (see Tyler et al., 2016), such as: changes in turbidity
50 (Olmanson et al., 2008), lake hydrodynamics (Pinardi et al., 2015), tidal effects on suspended solids (Eleveld
51 et al., 2014), bloom dynamics (Duan et al., 2014), submerged macrophytes cover (Bresciani et al., 2012). EO
52 techniques have been also used to investigate ecosystem degradation drivers, especially focusing on
53 vegetation status. Even though literature on this topic has principally dealt with terrestrial biomes (Smith et
54 al., 2014, and references therein), few works exist that, highlighting limitations due to satellite data features
55 (i.e. spatial, temporal, spectral resolutions, and data availability), have addressed applications in inland and
56 transitional environments, such as: loss and recovery of coastal vegetation after a tsunami (Villa et al., 2012),
57 disturbance on seagrass populations in Australia (Lyons et al., 2013; Kilminster et al., 2015), turbidity-driven
58 degradation of coral reef habitats (Fabricius et al., 2014), or aquatic vegetation changes with increased
59 eutrophication in China (Zhang et al., 2016).

60 New generation EO platforms, such as Landsat 8 and Sentinel-2, host mid-resolution (10-30 m pixel side)
61 optical multi-spectral sensors fine enough to monitor environmental phenomena at spatial and temporal
62 scales (see Pahlevan et al., 2017) not possible before they became operational, in 2013. Such technical
63 capabilities can improve existing EO based products and enable a number of new applications for remote
64 sensing of aquatic ecosystems (Murray et al., 2018), some of which were recently explored, focusing either
65 on aquatic vegetation (Stratoulas et al., 2015; Gao et al., 2017; Villa et al., 2018; Ghirardi et al., 2019), or on
66 water quality (Dörmhöfer et al., 2016; Pinardi et al., 2018; Fritz et al., 2019; Sòria-Perpinyà et al., 2020).

67 The aim of this work is to expand the range of EO applications in monitoring inland water systems by
68 providing a comprehensive framework that joins water quality and aquatic vegetation in a unique, real case
69 scenario, i.e. using Sentinel-2 data to study the effects of an upstream landslide event that altered sediment
70 inflow and deposition in a perialpine lake ecosystem (Lake Mezzola, northern Italy). Few works have been
71 carried out in the last decade on the impact of altered sediment deposition on perialpine lakes dynamics, e.g.
72 in response to upstream hydropower operations (Finger et al., 2006; 2007), or as reaction to increased flood

73 frequency (Fink et al., 2016), and EO data and techniques potentially satisfy the requirements for becoming
74 key tools towards further knowledge advancement on this topic.

75 The objective of this work is two-fold: i) to demonstrate the feasibility of monitoring the temporal evolution
76 of key ecosystem variables in a perialpine lake at fine spatial scale (10-20 m), using Sentinel-2 satellite data;
77 and ii) to use Sentinel-2 derived maps to quantitatively assess the impacts of an upstream landslide aftermath
78 on the aquatic and wetland plant communities in Lake Mezzola.

79

80 2. Study area

81 Lake Mezzola (46°13' N, 9°26' E, 199 m a.s.l.) is a deep (average depth: 26 m, maximum depth: 69 m)
82 perialpine lake located in northern Italy (Fig. 1). The lake, considered monomictic, covers an area of 5.85
83 km² and has the Mera River as main tributary, with its source in the Swiss Alps (2849 m a.s.l.), 40 km North-
84 East from the lake. The Mera River inflow, carrying sediments from the upstream Alpine basin, is
85 responsible for seasonal periods of high turbidity and suspended solids contribution (Secchi disk depth
86 varying between 1 and 5 m), while phytoplankton biomass is generally low (Chl-a concentration in the range
87 3-7 µg l⁻¹), with diatoms as dominant group (ARPA Lombardia, 2019). Lake Mezzola is connected
88 southwards to Lake Como by the Mera River, which in between the two forms a small shallow water basin,
89 called Lake Dascio (Fig. 1).

90 The nature reserve of “Pian di Spagna-Lago di Mezzola” encompasses the lake and around 16 km² of alluvial
91 floodplain, called Pian di Spagna, that lies between Lake Mezzola and the larger Lake Como. The nature
92 reserve, one of the oldest Ramsar sites in Italy and part of the European ecological network “Natura 2000”, is
93 an important migration and breeding area for more than 200 species of birds, including a variety of
94 waterbirds and raptors, and provides spawning and nursery areas for the rich fish fauna of Lake Mezzola.

95 This wetland system is characterized by lentic waters and marshes. In the southern portion of Lake Mezzola,
96 aquatic vegetation is mainly constituted by submerged vascular macrophytes, *Groenlandia densa* (L.) Fourr.,
97 *Potamogeton perfoliatus* L. and *P. lucens* L., and benthic macroalgae *Chara virgata* Kützing and *C.*
98 *globularis* Thuiller. The helophyte communities of Pian di Spagna wetland are dominated by common reed
99 (*Phragmites australis* (Cav.) Trin. Ex Steud), with some sparse presence of *Schoenoplectus lacustris* (L.)
100 Palla. More or less dense patches of isoetids - *Littorella uniflora* (L.) Asch, *Eleocharis acicularis* (L.) Roem.
101 & Schult, *Limosella aquatica* L., *Elatine hydropiper* L. and *Ranunculus reptans* L. - inhabit the lakeshore
102 area where seasonal water level fluctuations periodically leave the substrate uncovered, especially at the end
103 of summer. In the surroundings, wet meadows and sedges can be found together with some pastures and
104 stretches of open woodland (*Salix* spp., *Alnus glutinosa* (L.) Gaertn.).

105

106 **Figure 1 approximately here**

107

108 **2.1. Landslide event**

109 On 23 August 2017, a major rock landslide detached from the north face of Piz Cengalo towards Bondasca
110 valley, in South East Switzerland (Fig. 1), with a volume estimated at 3.1 million m³. After hitting a glacier,
111 the event evolved into a rock and ice avalanche, finally turning into a debris flow that reached the Bregaglia
112 valley in minutes (Mergili et al., 2019). The landslide was followed by a minor intensity event on 25 August.
113 Because of the joint events, 150 people were displaced from the village of Bondo (Switzerland), which was
114 partially destroyed. Resulting detrital material firstly reached the dammed river reservoir of Villa di
115 Chiavenna (Italy), which was emptied in July as a prevention measure. Afterwards, flowing downstream in
116 the Mera River, finer particles of landslide debris reached Lake Mezzola, triggering a massive increment in
117 water turbidity and suspended solids loads that lasted for some months.

118 Following the events, the Regional Environmental Protection Agency of Lombardy (ARPA Lombardia) has
119 put in place a specific monitoring program of the Mera River watershed and Lake Mezzola, in order to assess
120 the environmental effects of landslide aftermath, and in particular the impacts on water quality in Lake
121 Mezzola.

122

123 **3. Dataset**

124 **3.1. Satellite data**

125 Time series of Sentinel-2 data acquired over the study area (tile: T32TNS) covering three full years (2016-
126 2018) were gathered. The mission is managed by the European Space Agency (ESA) and the EU Copernicus
127 programme and data are distributed under a free and open data policy. Sentinel-2 is a constellation of two
128 satellites: Sentinel-2A (S-2A, launched on 23 June 2015) and Sentinel-2B (S-2B, launched on 7 March
129 2017). Both satellites carry on board the MultiSpectral Instrument (MSI), a mid-resolution (10-60 m) multi-
130 spectral (13 bands) push-broom imaging sensor (Drusch et al., 2012). Revisit time for Sentinel-2 is 10 days
131 with S-2A only (up to July 2017), and 5 days with both S-2A and S-2B satellites (i.e. from July 2017).

132 All the available S-2A and S-2B scenes acquired from 01 January 2016 to 31 December 2018, with cloud
133 cover less 50% on the area of Lake Mezzola and surroundings, were collected. The resulting time series (Fig.
134 2) was composed of 87 dates, acquired during the three years: 18 dates in 2016, 30 dates in 2017, and 39
135 dates in 2018.

136

137 **Figure 2 approximately here**

138

139 **3.2. Fieldworks and supporting data**

140 Time series of principal hydrological and meteorological parameters available for the last two decades were
141 retrieved from data collected by automatic sampling stations located nearby the study area and managed by

142 ARPA Lombardia. Daily average water level data of the Mera River collected in Lake Como (46°10'11" N,
143 9°23'37" E), 3 km downstream of Mera River outlet and available since 2013, were considered as good
144 proxy of Lake Mezzola water level variation. Meteorological daily data (1994-2018) were collected from a
145 weather station located 4 km upstream of the lake (46°14'01" N, 9°25'36" E), comprising: minimum,
146 average and maximum air temperature, in (°C), average and maximum global radiation ($W m^{-2}$), and
147 cumulated rainfall (mm).

148 An *in situ* floristic survey was performed at the beginning of August 2018, exploring in particular the
149 common reed and isoetid vegetation along the shoreline of the southern sector of Lake Mezzola. Afterwards,
150 a boat-based sampling campaign was carried out on 26 September 2018, for assessing the conditions of
151 submerged macrophytes and riparian reeds in the same area. Georeferenced photos (Camera Olympus Stylus
152 TG-4 Tough) and direct observations of vegetation conditions were collected in order to document plant
153 species present, reed density and phenological stage, as well as possible traces of damage.

154 The area covered by reed beds along the southern shore of Lake Mezzola and in Pian di Spagna wetland was
155 outlined by visual interpretation of high-resolution (<2 m pixel side) spaceborne images through the use of
156 geographic information software (QuantumGIS v3.4.3). Satellite images used were coming from Google
157 Earth (acquisition date: 10 April 2016) and WorldView-2 satellite (acquisition date: 01 July 2017).

158

159 **4. Methods**

160 Sentinel-2 data were downloaded as Top-Of-Atmosphere (TOA) reflectance data (Level-1C products). TOA
161 data were corrected for atmospheric effect using two different algorithms, depending on the target product.
162 For water quality retrieval, water-leaving radiance reflectance (ρ_w) was derived using ACOLITE, specifically
163 developed for coastal and turbid environments and adapted to Sentinel-2 data (Vanhellemont and Ruddick,
164 2016), selecting the per-pixel shortwave infrared-based aerosol correction. For aquatic vegetation mapping,
165 surface reflectance (ρ_0) was derived using SEN2COR (Louis et al., 2016), including embedded topographic
166 correction (SRTM DEM) and adjacency effect compensation (0.5 km radius).

167 Starting from the corrected Sentinel-2 reflectance dataset, three different mapping products were derived: i) a
168 time series of water turbidity for Lake Mezzola; ii) benthic substrate and submerged macrophytes cover
169 maps for the southern sector of Lake Mezzola; and iii) seasonal dynamics maps for reed beds in Pian di
170 Spagna wetland.

171 **4.1. Water turbidity maps**

172 A time series of water turbidity maps at the surface layers for Lake Mezzola (2016-2018), expressed in
173 formazin nephelometric units (FNU) at 20 m resolution, was produced from the ACOLITE-derived ρ_w ,
174 applying the single-band turbidity algorithm of Dogliotti et al. (2015), expressed in equation 1.

$$175 \quad \text{Turbidity [FNU]} = \frac{A_T^2 \rho_w(\lambda)}{1 - \rho_w(\lambda)/C^A} \quad (1);$$

176 where ρ_w refers to Sentinel-2 spectral band 4 (red) or band 8 (near infrared), and A_T and C were two
177 wavelength-dependent calibration coefficients taken from Nechad et al. (2009) and recalibrated to Sentinel-2
178 spectral ranges. The algorithm is reliable over a wide range of turbidity values (1-1500 FNU), avoiding
179 saturation issues by adopting a band-switching scheme between red and near infrared ranges (Dogliotti et al.,
180 2015).

181 **4.2. Benthic substrate maps**

182 Maps of benthic substrate and submerged macrophytes cover in the southern sector of Lake Mezzola were
183 derived at 10 m resolution from Sentinel-2 surface reflectance dataset, at key dates of different seasons (early
184 August in 2017 and 2018, and late September in 2016 and 2018), using the BOMBER tool (Giardino et al.,
185 2012). The tool implements a spectral inversion procedure of bio-optical models for both optically-deep and
186 optically-shallow waters, based on Lee et al. (1999). BOMBER was run for estimating bottom types and
187 cover fractions of bottom coverage, whilst setting water column properties as constant (e.g. Fritz et al.,
188 2019). In fact, for depths up to 10 m, as in the case of the shallow bank south of Lake Mezzola, the model
189 sensitivity to variations of water column properties is low and therefore bottom type estimation is less
190 dependent on possible errors due to water parameters setting (Giardino et al., 2016).

191 Optically-shallow water areas were detected when optimization error on spectral inversion for the optically-
192 deep model run surpassed 10% (Giardino et al., 2016 and reference herein). The bottom reflectance for the
193 optically-shallow model was then parametrised by using the reflectance values of sand and submerged
194 vegetation of a mix of macrophyte species (Bresciani et al., 2012), similar to those found in Lake Mezzola.
195 The specific water absorption and backscattering coefficients were the same used by Ghirardi et al. (2019)
196 for Lake Iseo. The concentrations of water constituents were set differently for each processed Sentinel-2
197 date and parameter, namely: i) suspended particulate matter values from ACOLITE turbidity outputs, by
198 assuming a unity conversion factor (Jafar-Sidik et al., 2017); ii) coloured dissolved organic matter absorption
199 values at 440 nm assumed fixed at 0.1 m^{-1} ; and iii) chlorophyll-a concentration varying according to the
200 month, i.e. 5 mg m^{-3} for early August and 2 mg m^{-3} for late September.

201 **4.3. Reed seasonal dynamics maps**

202 Starting from the spectral bands of the Sentinel-2 surface reflectance, the three-year time series of Water
203 Adjusted Vegetation Index (WAVI; Villa et al., 2014a) was calculated, according to equation 2. WAVI is a
204 spectral index developed as a proxy of density and biomass of aquatic vegetation, particularly effective for
205 emergent vegetation, i.e. helophytes (Villa et al., 2014a; 2014b).

$$206 \quad WAVI = 1.5 \frac{\rho_{0NIR}(b8_{MSI}) - \rho_{0BLUE}(b2_{MSI})}{\rho_{0NIR}(b8_{MSI}) + \rho_{0BLUE}(b2_{MSI}) + 0.5} \quad (2).$$

207 WAVI layers were produced, at 10 m spatial resolution, only for area covered by reeds in southern Lake
208 Mezzola and Pian di Spagna, making use of the reed beds map described in Section 3.2. Areas covered by
209 dense clouds, identified for each Sentinel-2 date as the pixels labelled with a cloud confidence level higher

210 than 20% (i.e. SEN2COR output layer “cloud_confidence” > 20) were masked out from the WAVI time
211 series (mask value = -1). Because S-2B data were available only from July 2017, a time series consistent
212 across the three years was built retaining only S-2A scenes. The final WAVI series covering the three years
213 (2016-2018) at nominal 10-day temporal resolution (revisit time of S-2A) were prepared by following Villa
214 et al. (2018), i.e. filling missing dates with void layers (value = -1).

215 Seasonal dynamics metrics of reed beds in Pian di Spagna wetland were then derived running TIMESAT
216 software (Jönsson and Eklundh, 2002; 2004) with 10-day WAVI series as input. TIMESAT was set for
217 running with no spike filtering, Asymmetric Gaussian curves as fitting method, and two iterations for the
218 envelope fitting (Gao et al., 2008; Villa et al., 2018). Maps of reed seasonal dynamics (10 m resolution) were
219 produced from TIMESAT outputs for: i) the start of the growing season timing, i.e. when WAVI curve
220 reached 0.5 of the maximum amplitude during green-up (SoS, expressed as the day of the year: DOY); ii) the
221 end of the growing season timing, i.e. when WAVI curve decreased below the 0.5 of the maximum
222 amplitude during senescence (EoS, as the DOY); iii) the peak WAVI reached during the season, as proxy of
223 maximum density (WAVI_max); and iv) the area under the WAVI curve, as proxy of reed seasonal
224 productivity (WAVI_integral).

225 **4.4. Assessment of landslide impacts on lake ecosystem**

226 Turbidity values for Lake Mezzola were extracted from Sentinel-2 derived time series. In particular, average
227 and maximum turbidity scores for lake centre were calculated for all dates falling within the aquatic
228 vegetation growing season, i.e. April to September (DOY 100-300), in years 2016, 2017 and 2018.

229 Changes in submerged macrophyte cover fraction in the optically-shallow southern sector of Lake Mezzola
230 and downstream Lake Dascio were assessed for evaluating the impacts due to increased turbidity in the
231 landslide aftermath. Impacts in terms of pixel-wise vegetation cover change was assessed in two different
232 moments of the growing season: at peak of growth (early August), calculating fractional cover difference
233 between 05 August 2018 (post-event) and 05 August 2017 (pre-event), and at the end of the growing season
234 (late September), calculating fractional cover difference between 29 September 2018 (post-event) and 29
235 September 2016 (pre-event). The assessment was made separately for southern Lake Mezzola and Lake
236 Dascio, for examining possible gradients in impact along the water flow direction.

237 In order to explore the impacts of landslide event on ecologically differentiated clusters, reed beds of Pian di
238 Spagna wetland were segmented into three groups, depending on their position within the ecosystem and
239 their ecological conditions, namely: island reeds, riparian reeds, and terrestrial reeds (see Supplementary Fig.
240 S1). Island reeds (Isl) were identified as the areas covered by reeds which are completely surrounded by
241 water, and are the more exposed to changes in lake level water quality. Riparian reeds (Rip) were identified
242 as the areas covered by reeds that are spatially connected to reed beds on the shoreline and lying less than 50
243 m inland, being directly influenced by lake level variation. Riparian reeds were further separated into two
244 groups: those located in Lake Mezzola (Rip_M), and those located in Lake Dascio (Rip_D). Terrestrial reeds
245 (Ter) were identified as reed covered areas that are not spatially connected to reed beds on the shoreline and

246 lying at least 200 m inland, being therefore considered as not directly influenced by changes in water quality
247 and sediment deposition and therefore not suffering direct impacts from landslide events. These spatial
248 thresholds were established based on the Lake Mezzola water level variation over the period investigated
249 (ranging from -0.2 to + 1.1 m with respect to the average level of the Mera River, 3 km downstream of the
250 lake; see par. 3.2). The number of 10 m pixels in Sentinel-2 derived maps and area covered by each of the
251 four groups are shown in Supplementary Fig. S1.

252 Seasonal dynamics metrics (SoS, EoS, WAVI_max, WAVI_integral) for the four reed groups were extracted
253 from satellite based maps for each of the three years (2016, 2017, 2018); from the extracted metrics data, the
254 differences in sample distribution (*p*-value and effect size) among years (2016, 2017, 2018) separately for
255 each reed group (Isl, Rip_M, Rip_D, Ter) were calculated.

256 **4.5. Statistical analysis**

257 Statistical analysis of differences within groups was performed using R v.3.6.1, with packages ggplot2 3.2.1,
258 ggpubr 0.2.2, FSA 0.8.25, rcompanion 2.3.0 (Mangiafico, 2016; Wickham, 2016; Ogle et al., 2019; R Core
259 Team, 2019). Because of non-normality of samples, multivariate differences among years were tested using
260 non-parametric methods, i.e. Kruskal-Wallis One Way Analysis of Variance on Ranks. Post-hoc pairwise
261 multiple comparison were performed using Dunn's test, and relevant *p*-value scores were calculated using
262 adjustment with the Benjamini-Hochberg method (Dunn, 1964). Effect size of pairwise sample differences
263 were calculated using Vargha and Delaney's A (VDA; Vargha and Delaney, 2000). VDA scores range from
264 0 to 1, with extreme values indicating stochastic dominance of one group over the other, and 0.5 value
265 indicating that two groups are stochastically equal (Vargha and Delaney, 2000). VDA values < 0.29 or >
266 0.71 indicate large effects.

267

268 **5. Results**

269 **5.1. Water turbidity patterns**

270 Turbidity values were extracted from Sentinel-2 based maps at the lake centre location (see Fig. 3a). As time
271 series in Fig. 3b shows, Piz Cengalo landslides of August 2017 (DOY 235-237) triggered a massive
272 increment in Lake Mezzola turbidity that lasted until the end of October, peaking around the second half of
273 September (DOY 267).

274 Average turbidity for lake centre during the 2017 growing season (DOY 100 to 300), when Piz Cengalo
275 landslides occurred, was 8.3 ± 0.2 FNU, 2.5 times the values of 2016 season, preceding the event, and 2.3
276 times the values of 2018 season, following it (Fig. 3c). In the same period, maximum turbidity for lake centre
277 reached 31.5 ± 2.2 FNU in 2017, respectively 3.4 and 3.7 times the maximum of 2016 and 2018 seasons.

278

279 **Figure 3 approximately here**

280

281 **Figure 4 approximately here**

282

283 **5.2. Impacts on submerged macrophytes**

284 Submerged macrophyte cover mapped from Sentinel-2 in the southern part of Lake Mezzola showed a heavy
285 loss in total area covered at peak of 2018 growing season, in early August, compared to the same period of
286 the previous year, i.e. before the landslide events took place. Fig. 4a shows that change in submerged
287 macrophyte fractional cover between 05 August 2018 and 05 August 2017 consisted in complete loss of
288 macrophyte coverage over 28.2 ha, and partial loss of coverage over 11.7 ha, out of total 70.3 ha (partially or
289 completely) covered by macrophytes in August 2017. Macrophyte cover at the end of growing season (29
290 September 2018) highlighted some signs of regain in total coverage happening in August and September
291 2018 (Fig. 4b): compared with pre-event conditions (29 September 2016), partial or full loss of macrophyte
292 cover adds up to 24.6 ha, while 23.4 ha of the area increased their macrophyte fractional cover. These
293 patterns indicate that submerged plants growth in 2018 may have been delayed compared to the previous
294 years (hence the highest loss at early August than late September) and/or a possible shift in community
295 composition may have happened (connected to localised partial coverage increment with respect to pre-event
296 conditions), favouring species characterized by late vegetative peak. Indeed, during the 2018 summer, the
297 aquatic vegetation turned out to be largely dominated by dense stands of *C. virgata*, with only a few spots of
298 *G. densa* and *P. perfoliatus*. Moreover, the broad bare sediments were largely loosely colonized by
299 charophyte seedlings with coverage rates below 5%.

300 Some impacts in terms of macrophyte cover loss is visible also for Lake Dascio, although with overall
301 magnitude far lower than what observed for upstream Lake Mezzola.

302 **5.3. Impacts on riparian reed beds**

303 Satellite-based maps of seasonal dynamics of reed beds in Pian di Spagna wetland revealed evidence of
304 strong stress conditions for riparian reeds, in terms of green-up (SoS, Fig. 5a), and green-down (EoS, Fig.
305 5b) timing, as well as seasonal productivity (WAVI_integral, Fig. 5d).

306 SoS of reed beds in Lake Mezzola came later in 2018 than in previous seasons by 30-33 days for island reeds
307 (Fig. 6a; $p < 0.0001$; $VDA > 0.916$) and by 20-22 days for riparian reeds (Fig. 6a; $p < 0.0001$; $VDA > 0.841$).

308 Right panel of Fig. 5a, relative to 2018 season, clearly show that such delay is particularly evident, i.e. up to
309 50 days later than in 2017, for reeds lying within the first 10-50 m from the waterfront and located in central
310 part of Lake Mezzola southern shore. SoS for terrestrial reeds (rightmost panel of Fig. 6a) was progressively
311 delayed from 2016 (DOY 128 ± 18) to 2017 (DOY 147 ± 11), and 2018 (DOY 155 ± 16).

312 EoS for island reeds in 2017, after landslide happened, came earlier by 24 and 35 days, respectively
313 compared to 2016 and 2018 (Fig. 6b; $p < 0.0001$; $VDA < 0.156$), while riparian reeds in both lakes Mezzola
314 and Dascio on average did not show notable differences in green-down timing across the three years (Fig.
315 6b; $0.371 < VDA < 0.517$). The spatial patterns of 2017 EoS depicted in central panel of Fig. 5b and violin

316 plots of Fig. 6b (2nd panel from the left), though, show that this average condition for Lake Mezzola reeds is a
317 composite effect of early senescence (around DOY 250) for riparian reeds located in central part and late
318 senescence (around DOY 320) for those growing in easternmost part of southern lake shore. An advance of
319 22 days on average for EoS of terrestrial reeds was observed in 2018, compared to 2017 season (Fig. 6b;
320 $p<0.0001$; VDA=0.156).

321 WAVI_integral of island reeds for both 2017 and 2018 seasons was lower than in 2016, respectively by 21%
322 and 19% (leftmost panel of Fig. 6d; $p<0.0001$; VDA<0.272). Riparian reeds in both lakes Mezzola and
323 Dascio did not show such highly sensitive differences in this proxy of seasonal productivity across the three
324 years, although some minor effect in terms of lowered WAVI_integral is visible for Mezzola when 2018
325 (post-event) and 2016 (pre-event) conditions are compared (2nd panel from the left of Fig. 6d; VDA=0.643).
326 Terrestrial reeds experienced a loss in WAVI_integral for 2018 with respect to 2017 by 11% (rightmost
327 panel of Fig. 6d; VDA=0.731). Out of the three seasons we considered, 2018 was generally the less
328 productive for helophytes of Pian di Spagna wetland (Fig. 5d and 6d).

329 No major difference was instead observed across the years for WAVI_max of any considered reed group
330 (Fig. 6c; $0.337<VDA<0.564$) even if some local patterns are visible in Fig. 5c. This indicates that maximum
331 canopy density reached during the growing season (surrogated by WAVI) was somewhat stable in 2016-
332 2018 in spite of changing meteorological drivers (see Supplementary Fig. S3) as well as landslide aftermaths.

333 The complete picture of results from statistical test of differences in seasonal dynamics metrics among years
334 carried out for each reed group as well as their effect size (VDA) is given in Supplementary Table S1.

335

336 **Figure 5 approximately here**

337

338 **Figure 6 approximately here**

339

340 **6. Discussion**

341 Satellite multi-temporal data provided a useful source of information for assessing ecosystem state in a
342 perialpine lake, in terms of water quality and macrophytes, both on cover and health conditions. Monitoring
343 capabilities demonstrated by the presented approach were amplified by the specific characteristics of
344 Sentinel-2, in terms of data revisit (5-10 days), spatial resolution (up to 10 m) and spectral content extending
345 from visible to shortwave infrared range, that meet the requirements for effective, operational ecosystem
346 monitoring (Kutser et al., 2006; Song et al., 2009).

347 Operational mid-resolution satellite data have recently been used for environmental applications on aquatic
348 systems focusing, for example on: riparian vegetation classification (Stratoulas et al., 2015), mapping water
349 quality and submerged macrophyte cover (Dömhöfer et al., 2016; Fritz et al., 2019), estimating aquatic
350 vegetation biomass (Gao et al., 2017), assessing the spatio-temporal evolution of primary producers (Pinarđi

351 et al., 2018; Ghirardi et al., 2019), analysing the seasonal dynamics of wetland plant communities (Villa et
352 al., 2018), and mapping cyanobacteria blooms (Sòria-Perpinyà et al., 2020).

353 Our work further shows the capabilities of mid-resolution satellite data and enlarges the scope with respect to
354 previous studies, by providing a comprehensive application that joins water quality (turbidity) and aquatic to
355 wetland vegetation (submerged macrophytes and helophytes) monitoring. In particular, the actual
356 contribution of satellite observations to freshwater ecosystem monitoring was demonstrated for assessing the
357 response to an extreme natural event on a real case. Sentinel-2 derived maps allowed us to highlight the
358 turbidity patterns in Lake Mezzola in space and time, the evolution of submerged macrophyte cover in
359 different years and the quantitative analysis of seasonal dynamics of helophytes communities living on the
360 lake shores. The approach we proposed could be easily implemented for continuous monitoring of Lake
361 Mezzola and surrounding areas during the current and next years, so to verify the evolution of aquatic
362 vegetation conditions and ecosystem equilibrium towards recovery or further degradation. Moreover, further
363 applications could be potentially extended to similar ecosystems - with water bodies larger than 2 ha and
364 hosting macrophyte communities with individual patches larger than 100 m² - susceptible to periodically
365 extreme sediment loads, which across the alpine and peri-alpine region could be due to upstream dam
366 regulations or landslides. The August 2017 landslides triggered an increase in turbidity that amounted to
367 twice the average and three times the maximum values scored in the years preceding and following the
368 events. Satellite-based maps revealed that river inflow in northern part of the lake and the lake central part
369 are hotspots of turbidity, highlighting surface circulation patterns. However, during 2017, high values of
370 turbidity were scored also for southern portion of Lake Mezzola, where a shallow bank (depth < 7 m) hosts
371 abundant macrophyte communities, and the helophyte beds of Pian di Spagna wetland reach the open waters.

372 Increased turbidity is a direct consequence of the massive load of detritus brought by the Mera River after
373 the Piz Cengalo landslides. Material brought by Mera River, almost completely inorganic (mainly made by a
374 mixture of silica and alumina, see Pettine et al., 2000), sedimented in Lake Mezzola and was the cause of
375 serious impacts on the lake ecosystem. The impacts shown by analysing the time series of Sentinel-2 derived
376 products were particularly intense on submerged macrophytes and riparian helophytes in southern part of the
377 lake, with cascade effects on the whole trophic chain and animal communities relying on them as primary
378 habitat.

379 Light availability is one of the key factors affecting submerged macrophyte establishment and productivity
380 (e.g. Barko et al., 1986, Lacoul and Freedman, 2006), with implications for growth rates and plant
381 morphology (e.g. Riis et al., 2012). High levels of turbidity and sediment deposition are therefore connected
382 with negative impacts on benthic vegetation, because they reduce light availability for photosynthesis and
383 can physically hinder germination of seeds or sprouting of new buds at the beginning of the season (Wood
384 and Armitage, 1997; Spencer and Ksander, 2002).

385 Our results documented a partial loss of total cover of submerged macrophytes (including charophyte beds)
386 in the season following the landslides, accompanied by signs of late growth, with the coverage peak moving

387 towards late September. Moreover, the analysis of spatial clustering of benthic vegetation dynamics shown
388 in Fig. 4 suggests the possibility of shifts in community dominance in favour of species less sensitive to the
389 effects of abundant deposition of inorganic, nutrient-poor sediments. Our 2018 surveys highlighted a clear
390 predominance of *C. virgata* over *G. densa* and *P. perfoliatus*. An annual or perennial species depending on
391 the growing depth, *C. virgata* can tolerate competition from other water plants (Guiry, 2019). Indeed, in
392 newly colonized or recently perturbed habitats charophytes may exhibit a strong pioneer behaviour and be
393 highly competitive towards vascular species (Bonis and Grillas, 2002; Brochet et al., 2010), despite of their
394 disadvantage in non-oligotrophic environments (Blindow, 1992). Our observations seem to go in this
395 direction, and we could explain the widespread presence of *C. virgata* in the southern banks of Lake Mezzola
396 referring to this pioneering habit. We therefore hypothesize that the submerged macrophyte cover mapped
397 after the events in 2018 could represent an initial stage of recolonization after a strong perturbation. This
398 process may have been strengthened by the physical and chemical features of particulate material brought in
399 the lake as a consequence of the landslides. The characteristics of this detrital material, rich in silica and
400 alumina, could support the widespread presence of oligotrophic species (i.e. charophytes and isoetids),
401 otherwise rare in areas tending to mesotrophic conditions (see Blindow, 1992; Sand-Jensen et al. 2008), such
402 as Lake Mezzola is. Such an effect of landslide deposits on submerged macrophyte community composition
403 could bring along interesting implications from the point of view of environmental management (e.g. of
404 mountain reservoirs and lakes), but requires further verification.

405 While temperature is considered the main factor controlling phenology of common reed (Irmak et al., 2013;
406 Petus et al., 2013; Anda et al., 2017), our results show that exceptional sediment load and deposition (e.g.
407 due to upstream landslides) can overwrite climatic factors in determining reed seasonal dynamics change.
408 Specifically, riparian reeds located south of Lake Mezzola suffered from an anticipated senescence in 2017
409 (late September, on average), initiated in the weeks after Piz Cengalo landslides, around 20 days before that
410 of 2016, as well as from a delayed green-up by more than 20 days in the following growing season, i.e. with
411 average SoS for island reeds in early July 2018.

412 Although 2018 was characterized by unusually cold early spring in North Italy, with growing degree-days
413 (GDD) cumulated up to mid-April lower than 1995-2014 average by 26°C d (-22% in relative terms, see
414 Supplementary Fig. S3), the late development of riparian reed beds of Lake Mezzola in 2018 (+33.2 and
415 +21.9 days on average in SoS compared to 2017 situation, for island and riparian stands respectively) was
416 not attributable only to meteorological anomaly. In fact, terrestrial reeds of Pian di Spagna wetland, which
417 we supposed not to be impacted by deposition of sediments in landslide aftermaths, due to their distance
418 from the shoreline (> 200 m), showed themselves some signs of delayed growth in 2018 (+9.2 days in SoS
419 on average, compared to 2017 conditions), but with far lower magnitude. Consistently with this, Anda et al.
420 (2017) found low temperatures in spring to delay reed growth by one week in Kis Balaton wetland, in
421 Hungary (46°37' N, 17°08' E). In addition to temperature in early spring, the extremely dry spring of 2018,
422 with -21% total cumulated rainfall from January to June, compared to 1995-2014 average (Supplementary

423 Table S2), might have contributed to slightly delayed SoS for terrestrial reeds, more sensitive to water stress
424 (Haslam, 1970).

425 The impact of landslide-triggered sediment deposition was evident already in autumn 2017 on island reed
426 communities, which suffered from senescence anticipated by three weeks with respect to pre-event
427 conditions (2016). In 2018, the situation went back to normality (average EoS at DOY 301), also favoured
428 by hot weather lasting longer than usual (mean temperature in October higher by 2.7°C compared to 1995-
429 2014 average, see Supplementary Fig. S3). On the other hand, the prolonged drought conditions of 2018
430 season, with -22% total cumulated rainfall from January to September compared to 1995-2014 average
431 (Supplementary Table S2), most probably amplified the water stress of terrestrial reed communities (Haslam,
432 1970), driving to an advanced green-down, about 10 days earlier than 2016 and 2017.

433 The anticipated senescence of 2017 and later green-up of 2018 resulted in two consecutive years of reduced
434 seasonal productivity (around -20% in terms of WAVI_integral scores) for riparian reed beds of Lake
435 Mezzola, with possible serious consequences in terms of reserves stored in rhizomes and capabilities to resist
436 further catastrophic events (Čížková et al., 2001).

437 Satellite-based maps are inevitably affected by a certain degree of uncertainty, due to the inherent
438 complexity of targets, i.e. biological and ecological systems. Water turbidity products derived with the
439 method implemented as in Section 4.1 have been assessed by Dogliotti et al. (2015) using *in situ* data
440 covering different sites and a wide range of turbidity (1-1000 FNU), scoring a mean relative error around
441 13.7%. Accuracy of benthic substrate cover derived with the method described in Section 4.2 has been
442 assessed by Ghirardi et al. (2019) over perialpine Lake Iseo against *in situ* reference data, resulting in per-
443 class accuracies (F-score) of 96.6% and 87.5% for submerged macrophytes and bare sediment cover classes,
444 respectively. Possible bias in key metrics of macrophyte seasonal dynamics derived from 10-day revisit
445 Sentinel-2 time series as delineated in Section 4.3 have been estimated with respect to the best temporal
446 resolution available for mid-resolution operational satellites by Villa et al. (2018) over emergent and floating
447 plants of Mantua lakes system, as under 2 days or 3 days for SoS and EoS, respectively.

448 The aforementioned uncertainty levels expected for the three ecosystem parameters at the basis of our
449 analysis do not substantially bias the primary findings described above, as they are generally around one
450 order of magnitude lower than the main effects of landslide impacts on Lake Mezzola ecosystem: i.e. water
451 turbidity increased by 2.3 to 3.7 times due to the landslide, and 20-30 day differences in SoS and EoS
452 measured for riparian reeds after the landslide aftermaths.

453 The impacts on riparian reeds are most visible in reed beds completely surrounded by water, thus more
454 sensitive to changes in water quality, as shown in Fig. 5 maps, i.e. representing SoS for 2018 (Fig. 5a, right
455 panel) and EoS for 2017 (Fig. 5b, central panel). Traces of dead, broken stems observed *in situ* in late
456 September 2018 (Supplementary Fig. S2) witnessed a retreat of the reed front with respect to previous
457 maximum coverage, as well as reed canopy thinning of some areas in central southern shore of Lake
458 Mezzola, particularly impacted according to Fig. 5 maps, thus corroborating satellite-based findings.

459 The decoupling between seasonal productivity proxy WAVI_integral, clearly showing the consequences of
460 landslides impacts on riparian reeds in both 2017 and 2018, and density-sensitive WAVI_max, not showing
461 sensitive differences among the three years for all reed types, could be explained by the growth of a turf
462 populated mainly by isoetids and other amphibious species - *L. uniflora*, *E. acicularis*, *L. aquatic* - below the
463 reeds in 2018 summer (documented during the *in situ* survey of 26 September 2018, see Supplementary Fig.
464 S2). The expansion of this understory was probably promoted by landslide impacts in a dual way, that is: i)
465 deposition of materials and increase in turbidity have weakened the riparian reed, as traces of broken reed
466 straws and low density reed patches were found during the *in situ* survey would support; and ii) the inorganic
467 sediments brought by the Mera river into the lake favoured the development of species adapted to organic-
468 poor substrates, as already observed for submerged macrophytes. Expansion of the understory and thinning
469 of riparian reeds probably balanced one each other in terms of integrated spectral response at canopy scale,
470 resulting in 2018 peak WAVI scores similar to those of previous years, while degradation of riparian reed
471 communities after 2017 summer landslides came out evidently from other seasonal dynamics metrics.

472 A possible explanation of the above described impacts of landslide aftermath on riparian reed belt of Lake
473 Mezzola can be found not only in the physical action of increased sediment loadings hampering rhizomes
474 oxygenation, but also in the quality of such new sediments, rich in silica and alumina. In fact, even if low to
475 moderate enrichment with silica can provide benefits for reed growth (Máthé et al., 2012; Schaller et al.,
476 2012a), high concentrations can compromise growth, possibly by inhibiting uptake of beneficial metals, i.e.
477 iron (Máthé et al., 2012; Schaller et al., 2012b). Literature on effects of alumina on reed plants are instead
478 not decisive, but tend to highlight some positive effects on phosphorous sequestration resulting into
479 competitive advantage towards other wetland species (Meyerson et al., 2002; Batty and Younger, 2004).

480

481 **Conclusions**

482 Monitoring the evolution of key lake ecosystem variables and their response to external events using remote
483 sensing is feasible, provided that satellite data with adequate spatial and temporal resolutions are available
484 and an approach is implemented based on products mapping different environmental parameters in a
485 comprehensive way.

486 Our work demonstrated the contribution of Sentinel-2 derived multi-temporal maps in assessing the impacts
487 of a landslide event on the ecosystem of Lake Mezzola (Northern Italy). We found a connection between the
488 landslide aftermath and lake ecosystem dynamics under different aspects, that are: i) water turbidity patterns
489 and their temporal evolution; ii) loss of biomass and possible shift in species compositions for submerged
490 macrophyte communities; iii) shortened growing season and reduced productivity for riparian reed beds on
491 the southern lake shore, due to early senescence in 2017 and delayed start of growth in 2018.

492 The results have shown that, although 2018 season was anomalously dry and hot compared to the previous
493 ones, the highlighted impact on riparian reeds is not attributable to meteorological drivers, as the terrestrial
494 reed beds in the area did not show the same delay in green-up and loss of productivity.

495 The utilization of Sentinel-2 satellite data in the framework of the presented approach would make possible
496 for environmental bodies and public authorities to carry on monitoring the state of Lake Mezzola ecosystem
497 equilibrium, in order to assess the evolution trend whether it goes towards recovery of further degradation for
498 both submerged macrophyte communities and helophytes, and also that of similar ecosystems subject to
499 external pressures with periods of high sediment loads.

500

501 **Author contributions**

502 **Paolo Villa**: Conceptualization, Methodology, Analysis, Investigation, Writing of original draft. **Mariano**
503 **Bresciani**: Conceptualization, Methodology, Investigation. **Rossano Bolpagni**: Analysis, Investigation.
504 **Federica Braga**: Methodology. **Dario Bellingeri**: Investigation. **Claudia Giardino**: Supervision. All
505 authors reviewed and edited the final manuscript.

506

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515

516 **Data availability**

517 The authors confirm that the main processed data supporting the findings of this study are available within
518 the article and its supplementary materials. Raw data and satellite-based products used for our analysis can
519 be made available from the corresponding author, P. Villa, upon reasonable request.

520

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Figure captions

Figure 1. Study area overview, including a near infrared colour composite of the area derived from the Sentinel-2 image acquired on 09 September 2018.

Figure 2. Satellite dataset, highlighting dates with cloud cover over more than 10% of the study area.

Figure 3. Evolution of water turbidity in Lake Mezzola derived from Sentinel-2 satellite data covering the period 2016-2018: a) maps of average water turbidity for in the macrophyte growing season (DOY 100-300) for the three years (shallow lake areas, i.e. depth < 1.5 m, are masked out); b) yearly time series of average turbidity at lake centre location extracted from satellite maps (landslide events in 2017 marked by grey vertical line); c) violin plots of average and maximum turbidity scores extracted from satellite maps during the macrophyte growing season (DOY 100-300); *p*-value scores for the differences among years (Kruskal-Wallis One Way ANOVA) and post-hoc pairwise multiple comparison (Dunn's test) are overlaid on panel c: < 0.0001 (****); < 0.001 (***); < 0.01 (*); < 0.05 (*); > 0.05 (ns).

Figure 4. Changes in submerged macrophytes cover for southern Lake Mezzola and Lake Dascio derived from selected Sentinel-2 satellite scenes: a) map of early August conditions, i.e. peak of vegetative season, in 2017 (pre-landslides) and 2018, and histograms of fractional cover difference extracted from these maps over the two lakes; b) map of late September conditions, i.e. end of growing season, in 2016 (pre-landslides) and 2018, and histograms of fractional cover difference extracted from these maps over the two lakes.

Figure 5. Maps of seasonal dynamics of reed beds in Pian di Spagna wetland derived from Sentinel-2 satellite time series for the three years from 2016 to 2018: a) start of season (SoS); b) end of season (EoS); c) a spectral proxy of peak canopy density (WAVI_max); d) a spectral proxy of seasonal productivity (WAVI_integral).

Figure 6. Violin plots (with encompassed box plots) of reed beds seasonal dynamics across the three years (2016-2018) for each separate reed group, i.e. Island reeds, Riparian reeds (in Lake Mezzola and Dascio), and Terrestrial reeds: a) start of season (SoS); b) end of season (EoS); c) a spectral proxy of peak canopy density (WAVI_max); d) a spectral proxy of seasonal productivity (WAVI_integral); *p*-value scores for the differences among years (Kruskal-Wallis One Way ANOVA) and post-hoc pairwise multiple comparison (Dunn's test) are overlaid on panel c: < 0.0001 (****); < 0.001 (***); < 0.01 (*); < 0.05 (*); > 0.05 (ns).

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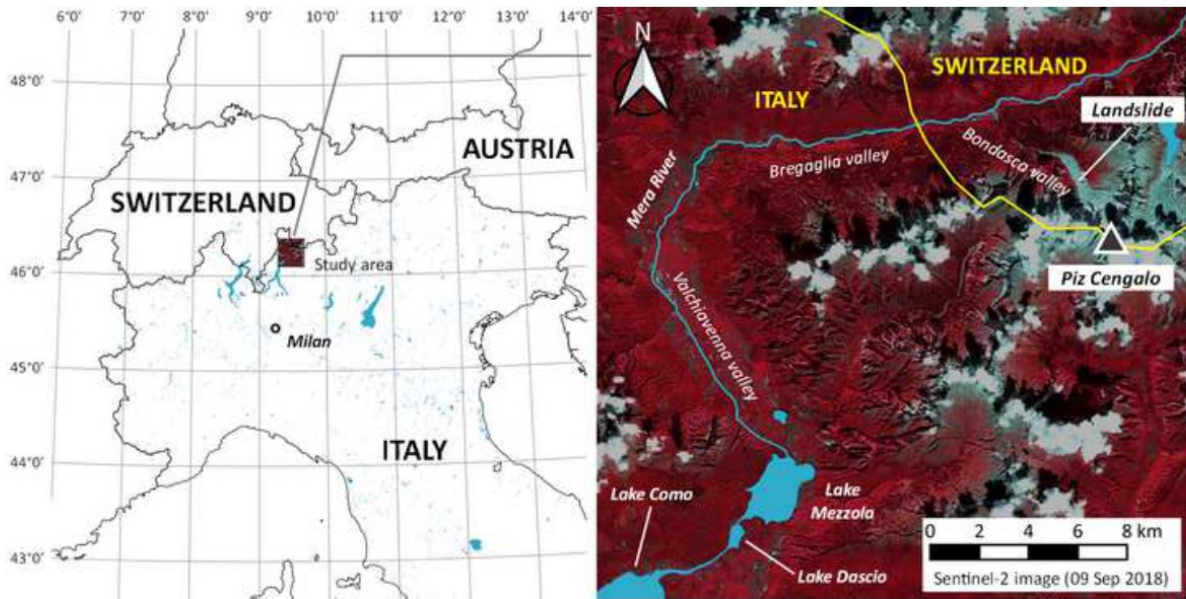


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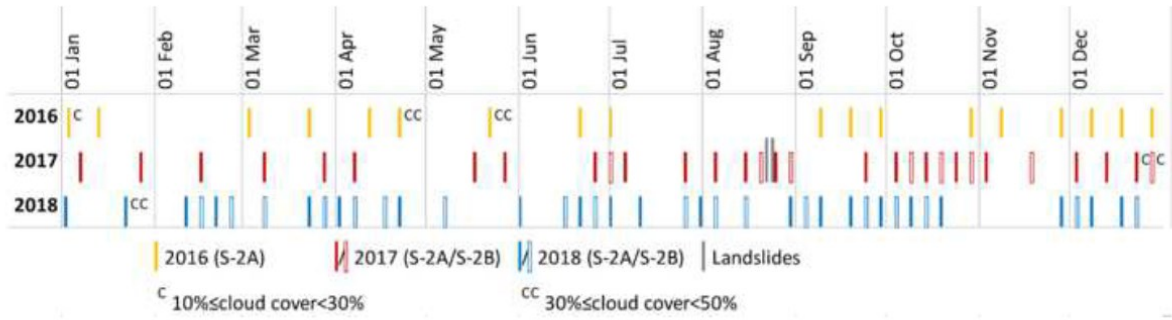


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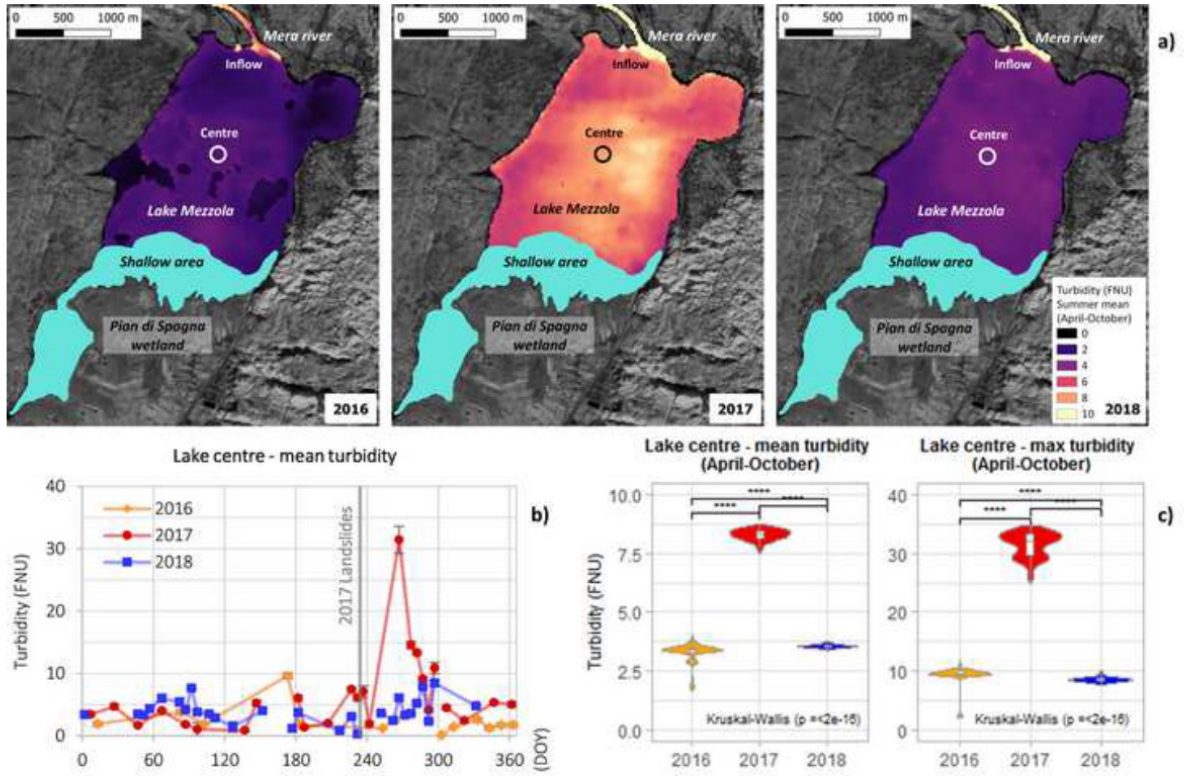


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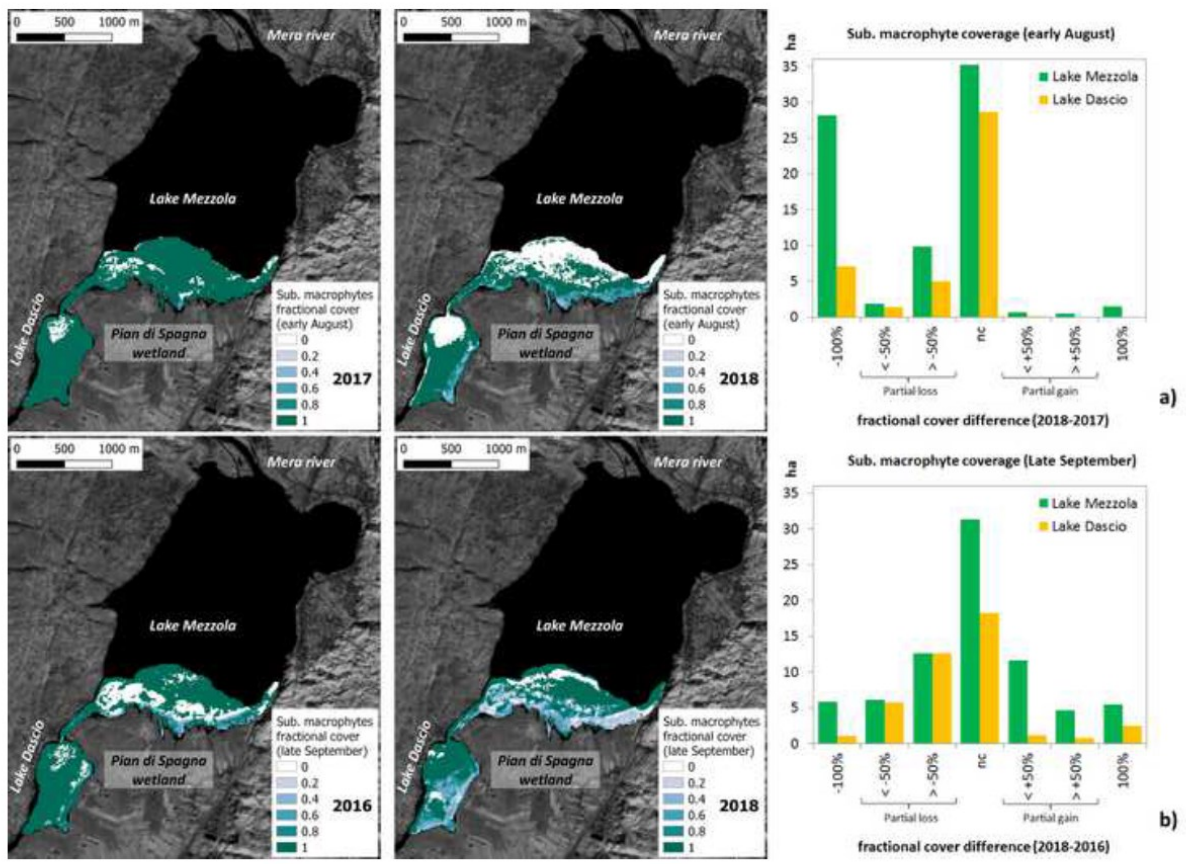


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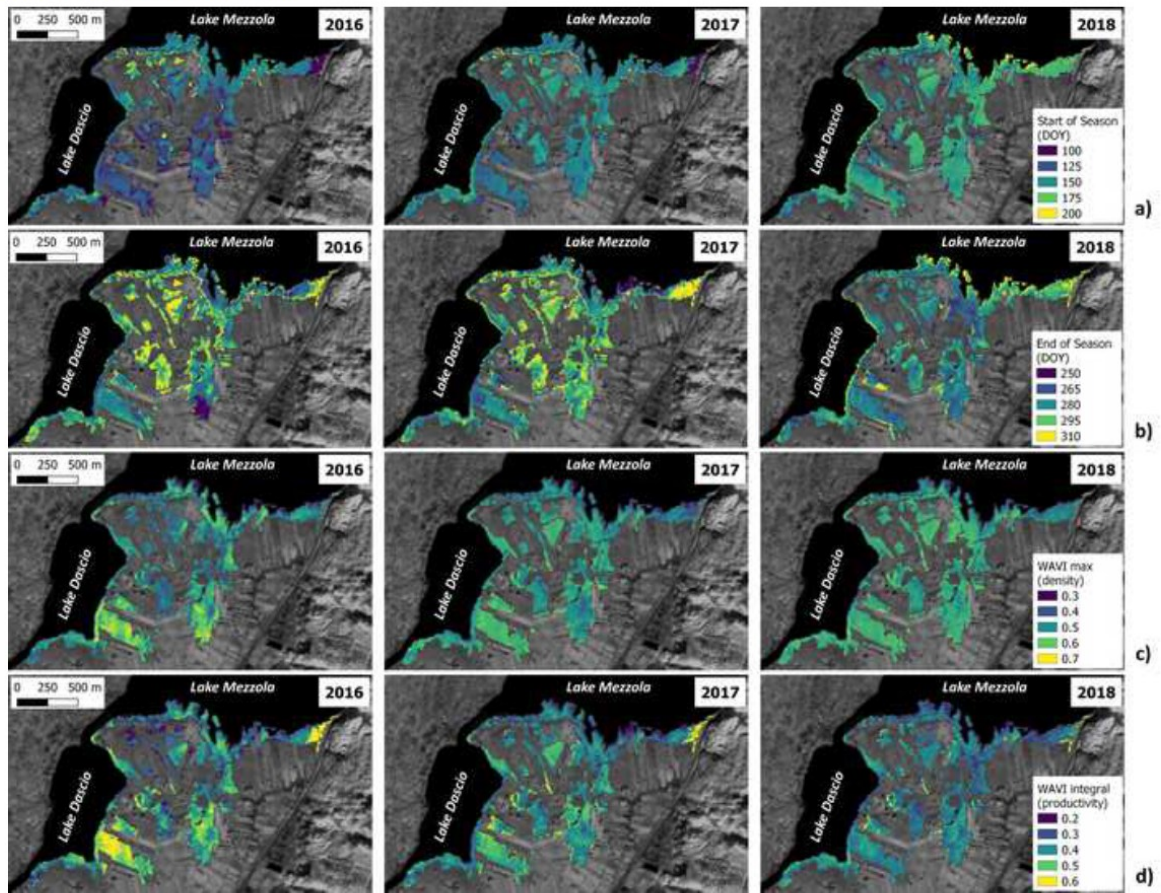


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