

Did transit through the galactic spiral arms seed crust production on the early Earth?

Kirkland, C.L.¹, Sutton, P.J.², Erickson, T.³, Johnson, T.J.¹, Hartnady, M.¹, Smithies, H.^{1,4}, Prause, M.⁴, Gardiner, N.J.⁵

¹Timescales of Mineral Systems Group, School of Earth and Planetary Science, Curtin University; Perth, Australia

²University of Lincoln, School of Mathematics and Physics; Lincoln, LN6 7TS, UK

³Jacobs—JETS, Houston, Texas, 77058, USA

⁴Geological Survey of Western Australia; Perth, Australia

⁵School of Earth and Environmental Sciences, University of St. Andrews; St. Andrews KT16 9AL, UK

Abstract

Although there is evidence for periodic geological perturbations driven by regular or semi-regular extra-terrestrial bombardment, the production of Earth's continental crust is generally regarded as a function of planetary differentiation driven by internal processes. Here, we report time series analysis of the Hf isotopic composition of zircon grains from the North Atlantic and Pilbara cratons, the archetypes of Archean plate tectonic and non-plate tectonic settings, respectively. A ~170–200-million-year frequency is recognized in both cratons that matches the transit of the solar system through the galactic spiral arms, where the density of stars is high. An increase in stellar density is consistent with an enhanced rate of Earth bombardment by comets, the larger of which would have initiated crustal nuclei production via impact-driven decompression melting of the mantle. Hence, the production and preservation of continental crust on the early Earth may have been fundamentally influenced by exogenous processes. A test of this model using oxygen isotopes in zircon from the Pilbara Craton reveals correlations between crust with anomalously light isotopic signatures and exit from the Perseus spiral arm and entry into the Norma spiral arm, the latter of which well-matches the known age of terrestrial spherule beds. The data support

28 bolide impact, promoting the growth of crustal nuclei, on solar system transit *into* and *out of* the
29 galactic spiral arms.

30 Introduction

31 Earth is unique amongst the known planets in having continents, whose formation has
32 fundamentally influenced the composition of the mantle, hydrosphere, atmosphere, and biosphere.
33 Cycles in the production of continental crust have long been recognised (Condie, 1998) and
34 generally ascribed to the quasi-periodic aggregation and dispersal of Earth's continental crust as
35 part of the supercontinent cycle (Murphy and Nance, 2003). However, such cyclicity is also evident
36 in some of Earth's most ancient rocks that formed during the Archean (4000–2500 Ma) and
37 Hadean (>4000 Ma) eons, when plate tectonics may not have operated. These frequencies are
38 challenging to decipher in the early Earth given that the rock record becomes increasingly
39 fragmentary with age.

40 One mechanism to assess the cyclicity of crust production and its drivers is through the isotopic
41 records of rocks and constituent minerals (Puetz and Condie, 2019). Zircon U–Pb–Lu–Hf isotopic
42 datasets are particularly well-suited to this endeavor as they retain a time-encoded proxy for the
43 degree of source fractionation. Here, we investigate the cyclicity in the addition of new mantle-
44 derived (juvenile) crust and its subsequent reworking through the Hf isotopic record of dated
45 zircon grains from the Archean North Atlantic Craton (NAC) in West Greenland, and the Pilbara
46 Craton in Western Australia. We complement this Hf data set with oxygen isotope compositions
47 of dated Pilbara zircon.

48 The NAC is dominated by Mesoarchean (3200–2800 Ma) felsic (silica-rich) gneisses and ultramafic
49 complexes, along with younger supracrustal sequences and late-tectonic felsic to mafic intrusions
50 (Friend and Nutman, 2019). It is commonly considered to have formed through horizontal
51 (subduction–accretion) processes, and to provide evidence for the operation of plate tectonics (a

52 mobile-lid regime) by that stage in Earth history (Garde et al., 2020). By contrast, the older part of
53 the Pilbara Craton in Western Australia consists of variably-metamorphosed Palaeo- to
54 Mesoarchean ultramafic to felsic supracrustal successions surrounding 3600–3300 Ma (Smithies et
55 al., 2009) granite domes, and is generally regarded to have formed through vertical (non-plate
56 tectonic) processes in a single-plate (deformable stagnant lid) regime (Smithies et al., 2019).

57 Not all cratons were initiated at the same time, nor did they evolve at the same rate (Mole et al.,
58 2019). Consequently, rather than using a global dataset, we examine the isotopic time series on a
59 per craton basis. We fit quantiles (25th, 50th, 75th) to the NAC and Pilbara data sets using a moving
60 25 Ma bin. We interpret the 75th quantile to chart a more significant contribution from juvenile
61 magmatism and the 25th quantile to mark a greater degree of crustal recycling, whereas the 50th
62 quantile to reflect the average evolution. To further evaluate the zircon Hf time series, we use
63 various spectral analysis approaches including periodograms and continuous wavelet transforms
64 (CWT) (see DR1).

65 Zircon hafnium isotope time series

66 The 50th quantile of the NAC Hf data tracks along a $^{176}\text{Lu}/^{177}\text{Hf}$ ratio of 0.01 until 3200 Ma, when
67 average values deviate towards more radiogenic signatures (Fig. 1A). A periodogram highlights
68 frequency bands with >95% significance at ~198 Myr, ~113 Myr, several around ~74 Myr, and
69 54 Myr, and a band at ~680 Myr that is above the noise model (Fig. 2A). Wavelet analysis of the
70 various quantile time series reveal broadly similar frequencies including a quasi-continuous band
71 at ~200 Myr (Fig. 1B, C, D), and a ~170 Myr frequency considering only the >3200 Ma segment
72 (DR1). All quantiles show more frequency structure in the pre-3200 Ma component of the record,
73 with longer wavelengths dominating the post-3200 Ma segment.

74 By contrast, the Pilbara Hf time series (3800–2825 Ma) initially tracks values close to the chondritic
75 uniform reservoir (CHUR), then becomes more variable with some super-chondritic values after

76 3500 Ma (Fig. 1E). Spectral analysis of the 50th quartile fit highlights frequency bands at >95%
77 significance at ~114 Myr, ~80 Myr, and ~55 Myr, and bands at ~415 Myr and ~168 Myr that are
78 above the noise model (Fig. 2B). Wavelet analysis of the various quantile time series have
79 comparable bands with a continuous frequency at ~170 Myr (Fig. 1F, G, H). The frequency
80 response of the Pilbara zircon Hf timeseries is similar to the CWT pattern in the pre-3200 Ma
81 NAC.

82 Periodicities superimposed onto Earth systems

83 Various periodicities are thought to be superimposed onto the Earth system through processes
84 acting across varying temporal and spatial scales, including planetary, solar system, and even
85 galactic. These include a ~800 Myr resonance between tidal and free oscillations of the core, a
86 500–300 Myr supercontinent cycle, a ~200 Myr galactic year, and a ~30 Myr impact cycle
87 (Rampino et al., 2019). Other long period geological cycles have been proposed from studies of
88 large igneous provenances (LIPs) and mantle plumes (Prokoph et al., 2013).

89 The zircon Hf time series analysis for both the NAC and Pilbara cratons show frequency
90 components in the ~170–200 Myr range, which are similar to those reported from timeseries
91 analysis of LIPs, interpreted as a proxy for mantle plumes (Ernst and Buchan, 2002) (Fig. 2C).
92 However, the Milky Way Galaxy appears to have four major spiral arms, which implies spiral arm
93 passages also occur approximately every ~170–200 Myr (Rampino, 1997) as both arm and solar
94 system orbit the galactic center at different rates. During this passage the solar system will transit
95 through dense interstellar clouds and be subject to variable galactic tides (Fig. 3A). Perturbations
96 of the Oort cloud (Fig. 3B) due to oscillations of the solar system around its galactic plane, and
97 interactions between spiral arms with other areas of enhanced star formation during galactic
98 transit, have been directly linked by some to the flux of meteorite impacts on Earth (Goncharov
99 and Orlov, 2003).

100 Earth is subject to impacts from near Earth objects (NEOs), primarily derived from the main
101 asteroid belt, and comets (Ivanov, 2008). While the former is inferred to result in much more
102 frequent impacts (Granvik et al., 2018), the latter would result in two orders of magnitude more
103 energy released into the crust on collision for comparable size impactors (Nuth et al., 2018). More
104 energetic impacts excavate and uplift a greater volume of material, with impact melt production a
105 function of energy and momentum (Schmidt and Housen, 1987). Comets would be the most likely
106 impactor to carry a resolvable low frequency periodic signal as they are most susceptible to
107 perturbations from outside the solar system (Rampino, 1997). Independent of impact energies,
108 NEOs responsible for >20km craters occur every 0.75 Myr, which renders them unresolvable in
109 the Hf isotope record (Ivanov, 2008). The lower forcing frequency coupled with the higher impact
110 velocities of comets creates a detection bias in the Hf isotope record, despite NEOs as the
111 expected dominant impactors.

112 Impact-generated melt production is also affected by the target temperature, which would be
113 greater for an equivalent size impact on a warmer young Earth with steep thermal gradient and
114 high internal temperatures (Potter et al., 2012). We posit that melts from energetic comet impact
115 would form buoyant crustal nuclei that in turn would support crustal growth through hosting the
116 products of later differentiation. This process must have been independent of the prevailing
117 geodynamic settings since, for the NAC and Pilbara Cratons (e.g. subduction vs. stagnant lid
118 tectonics), similar frequencies in their zircon Hf isotope time series are preserved, suggesting a
119 common driving process. Additionally, the zircon Hf isotopic record tracks a similar frequency,
120 throughout the entire Archean, to that advocated for Palaeozoic impact induced extinctions
121 (Shoemaker, 1998). Arguably, the extension of this frequency into deep time points to a large-scale
122 and relatively constant driver over most if not all our planet's history.

123 Models of early Earth crust generation emphasize melting of hydrated basaltic rocks to produce
124 Tonalite–trondhjemite–granodiorites (TTGs). Smithies et al., (2021) suggested TTG magmas

125 could be generated, both, via a bottom up process facilitated by mantle-derived water, and later in
126 the Palaeoarchean, via a top down processes derived from sinking of greenstones. Models of
127 impacts on the early Earth suggest production of large felsic shallow impact melt pools (Grieve et
128 al., 2006). Hence, impacting critically fits into this formative history through providing buoyant
129 seeds that would act as nucleation points for later TTGs generated via either mantle or greenstone
130 extraction, simply as a function of having greater preservation potential.

131 A test using $\delta^{18}\text{O}$ in zircon

132 To better understand any link between the periodicities identified in the zircon Hf isotope time
133 series and external forcing we develop a model for the motion of the solar system within the Milky
134 Way Galaxy that we use to estimate mass distribution relative to our solar system. We then
135 compare this to the compiled zircon oxygen record of the Pilbara Craton (Johnson et al., 2022;
136 Fig. 4A, B).

137 Previous models have assumed a solar system orbit, relative to the spiral arms (i.e. galactic period),
138 of 752 Myr, based on the average time between Superchrons, the periods during which Earth's
139 magnetic field remained stable (Gillman and Erenler, 2019). The best fit galactic period to our data
140 is 748 Myr, resulting in a duration of 187 Myr between passages through the spiral arms.

141 Oxygen isotopes in zircon crystals formed within and inherited by igneous rocks in the Pilbara
142 Craton define a secular trend that further constrains crust production (Smithies et al.,
143 2021)(Johnson et al., 2022). Zircon oxygen isotopes start with, on average, $\delta^{18}\text{O}$ values lighter than
144 mantle values prior to 3400 Ma. Between 3400 and 3100 Ma zircon $\delta^{18}\text{O}$ values are mainly within
145 the mantle field, but thereafter extend to heavy $\delta^{18}\text{O}$ values consistent with widespread reworking
146 of supracrustal material (Fig. 4A). A continuous ~ 200 Myr frequency is evident in CWT analysis
147 of a LOWESS (Locally weighted scatterplot smoothing) fit to the zircon oxygen data (Fig. 4B),
148 with departures to light oxygen isotope values in zircon at ~ 3560 Ma and ~ 3430 Ma (Fig. 4A).

149 Zircon with isotopically-light oxygen is commonly associated with volcanic calderas and shallow
150 extensional systems. Impacts can also establish light signatures through generation of both large
151 volumes of hydrothermally altered crust and widespread post impact shallow melt pools (Grieve
152 et al., 2006). Giant impacts may also drive decompression melting of the mantle (Shibaïke et al.,
153 2016). Notably, whilst the record of preserved crust provides very little material (other than zircon)
154 to relate to the ~ 3560 Ma light oxygen isotope departure, the ~ 3430 Ma excursion corresponds to
155 the age of spherule beds in Australia and South Africa that are direct evidence of large impacts
156 (Byerly et al., 2002). Similar frequencies may also be expected for plume-driven models of early
157 continental growth (Reimink et al., 2014), however, the close temporal association with spherule
158 beds favours an impact-driver. The ~ 3560 Ma light zircon oxygen signature occurs on the
159 predicted exit from the Perseus arm, whereas the younger ~ 3430 Ma excursion corresponds to
160 entry into the Norma arm (Fig. 3A).

161 Links to vertical oscillations of the solar system to perturbations of Oort cloud comets by the
162 galactic tide or nearby passing stars have been made (Levison et al., 2004). The motion of the Sun
163 modulates the strength of these perturbations as they are influenced by the local stellar density.
164 However, the observed anisotropy in the perihelia of long-period comets does not support galactic
165 tide oscillations as the sole source of such perturbations (Delsemme, 1987), and single encounters
166 with a nearby star will not cause significant perturbation of the Oort cloud unless passing very
167 close (< 0.5 pc)(Torres et al., 2019). The higher densities of stars in spiral arms increases the
168 probability of multiple stellar encounters, which also increases across the time spent in the arm,
169 which may explain impacts around the *exit* of spiral arms.

170 A proposed mechanism for clustering of impacts when *entering* a spiral arm may be attributed to
171 perturbation of the Oort cloud induced by Giant Molecular Clouds (GMC). Encounters with
172 GMC are known to cause disturbance and, in extreme cases, cause comets to be lost to interstellar
173 space (Jakubík and Neslusan, 2008). Compression of gas passing across a galactic arm causes the

174 formation of a GMC, with subsequent gravitational collapse initiating star formation (Dobbs et
175 al., 2014). As GMC have lifetimes (10–50 Myr) less than the duration of passage through a spiral
176 arm, they are unlikely to perturb the Oort cloud during its *exit* of a spiral arm.

177 Galactic driver

178 Long period variations in the flux of comet impacts to Earth likely influenced crust production
179 through a variety of mechanisms. On the early Earth, giant impacts may have been the trigger for
180 production of Earth’s first crustal nuclei via impact-induced melting and magmatic differentiation
181 (Grieve, 1980). Younger impacts may have enhanced or modified processes of crust formation
182 and destruction operating in different modes, at different times and locations. Clearly, the role of
183 impacts in crust production will have changed with time owing to the exponential decline in
184 average projectile size and particle density in the early solar system (Marchi et al., 2014), but also
185 as the dominant tectonic mode on Earth transitioned into widespread subduction, as evident in
186 the increasing average crustal residence in Pilbara magmas after 3200 Ma (Fig. 1).

187 Isostatic adjustments triggered by giant meteorite impacts (Fig. 3C) are predicted to result in
188 extensive decompression melting of the mantle to produce thick plateau-like mafic–ultramafic pre-
189 cratonic nuclei (Jones et al., 2003). Meteorite impacts also cause intense fracturing and brecciation
190 of bedrock (Riller et al., 2018), and facilitate the development of prolonged hydrothermal
191 circulation (Kring et al., 2020) (Fig. 3D). Such circulation promotes enhanced fluid–rock
192 interaction, including significant shifts in $\delta^{18}\text{O}$ values, and supports partial melting, both in the
193 immediate aftermath of impacting and developing a fertile melt source long after the event. Once
194 initial mantle melt segregation occurred, it would be inevitable that such impact induced buoyant
195 felsic seeds would have had a lower propensity to be recycled than their dense mafic counterparts.
196 Whilst endogenous processes (e.g. plate tectonics) have been important in the establishment and
197 maintenance of Earth’s major geochemical reservoirs, the wider galactic environment may have
198 seeded cratonic nuclei via impacting, that ultimately grew to the continental crust we live on.

199 Acknowledgements

200 R.H.S. and M.P. publish with GSWA permission.

201 Captions

202 **Fig. 1:** (A) Zircon ϵHf evolution plot of the North Atlantic Craton with quantile fits. (B-D, F-H)
203 Continuous wavelet spectrums for 25th, 50th, 75th fits. Region of edge effects indicated by the
204 black line. (E) Zircon ϵHf evolution plot of the Pilbara Craton.

205 **Fig. 2:** (A) Periodogram for zircon ϵHf time series (50th quantile) from the: North Atlantic
206 Craton, (B) Pilbara Craton. Green line: 95% significance level above the noise model. Red line:
207 red noise model.

208 **Fig. 3:** (A) Milky Way Galaxy. Arm and solar system rotate clockwise at different rates. The +
209 number denotes the number of times Earth has passed through that arm. (B) Model of the Oort
210 cloud. (C) Schematic reconstruction of early Earth impact process (D) seeding crustal nuclei.

211 **Fig. 4:** (A) Continuous wavelet spectrum for LOWESS fit to Pilbara zircon oxygen isotope time
212 series (B) Igneous zircon oxygen isotope evolution plot for the Pilbara Craton. Spherule beds;
213 orange circles. Error bars at 2σ level. Dashed blue line; average crustal residence age.

214 References

215

- 216 Byerly, G. R., Lowe, D. R., Wooden, J. L., and Xie, X., 2002, An Archean Impact Layer from the Pilbara
217 and Kaapvaal Cratons: *Science*, v. 297, no. 5585, p. 1325.
- 218 Condie, K. C., 1998, Episodic continental growth and supercontinents: A mantle avalanche
219 connection?: *Earth and Planetary Science Letters*, v. 163, no. 1-4, p. 97-108.
- 220 Delsemme, A. H., 1987, Galactic Tides Affect the Oort Cloud - an Observational Confirmation:
221 *Astronomy and Astrophysics*, v. 187, no. 1+2, p. 913-918.
- 222 Dobbs, C. L., Krumholz, M. R., Ballesteros-Paredes, J., Bolatto, A. D., Fukui, Y., Heyer, M., Low, M. M.
223 M., Ostriker, E. C., and Vázquez-Semadeni, E., 2014, Formation of molecular clouds and global
224 conditions for star formation.: *Protostars and Planets VI*, v. 1312, p. 3-26.
- 225 Ernst, R. E., and Buchan, K. L., 2002, Maximum size and distribution in time and space of mantle
226 plumes: evidence from large igneous provinces, *in* Condie, K. C., Abbot, D., and Des Marais, D.
227 J., eds., *Superplume events in Earth's history: causes and effects*, Volume 34 *Journal of*
228 *Geodynamics*, p. 309 – 342.
- 229 Friend, C. R. L., and Nutman, A. P., 2019, Tectono-stratigraphic terranes in Archean gneiss complexes
230 as evidence for plate tectonics: The Nuuk region, southern West Greenland: *Gondwana*
231 *Research*, v. 72, p. 213-237.
- 232 Garde, A. A., Windley, B. F., Kokfelt, T. F., and Keulen, N., 2020, Archean Plate Tectonics in the North
233 Atlantic Craton of West Greenland Revealed by Well-Exposed Horizontal Crustal Tectonics,
234 Island Arcs and Tonalite-Trondhjemite-Granodiorite Complexes: *Frontiers in Earth Science*, v.
235 8, p. 526.
- 236 Gillman, M., and Erenler, H., 2019, Reconciling the Earth's stratigraphic record with the structure of
237 our galaxy: *Geoscience Frontiers*, v. 10, no. 6, p. 2147-2151.
- 238 Goncharov, G. N., and Orlov, V. V., 2003, Global repeating events in the history of the Earth and the
239 motion of the Sun in the Galaxy: *Astronomy Reports*, v. 47, no. 11, p. 925-933.
- 240 Granvik, M., Morbidelli, A., Jedicke, R., Bolin, B., Bottke, W. F., Beshore, E., Vokrouhlický, D., Nesvorný,
241 D., and Michel, P., 2018, Debiased orbit and absolute-magnitude distributions for near-Earth
242 objects: *Icarus*, v. 312, p. 181-207.
- 243 Grieve, R. A. F., 1980, Impact bombardment and its role in proto-continental growth on the early earth:
244 *Precambrian Research*, v. 10, no. 3, p. 217-247.
- 245 Grieve, R. A. F., Cintala, M. J., and Therriault, A. M., 2006, Large-scale impacts and the evolution of the
246 Earth's crust: The early years, *in* Reimold, W. U., and Gibson, R. L., eds., *Processes on the Early*
247 *Earth*, Volume 405, Geological Society of America, Special Paper, p. 23-31.
- 248 Ivanov, B., 2008, Size-Frequency Distribution Of Asteroids And Impact Craters: Estimates Of Impact
249 Rate, *in* Adushkin, V., and Nemchinov, I., eds., *Catastrophic Events Caused by Cosmic Objects:*
250 *Dordrecht*, Springer Netherlands, p. 91-116.
- 251 Jakubík, M., and Neslusan, L., 2008, The dynamics of the Oort cloud during a passage through a
252 spherical giant interstellar cloud with the Gaussian-density profile.: *Contributions of the*
253 *Astronomical Observatory Skalnaté Pleso*, v. 38, p. 33-46.
- 254 Jones, A. P., Price, D. G., DeCarli, P. S., Price, N., and Clegg, R., 2003, Impact Decompression Melting:
255 A Possible Trigger for Impact Induced Volcanism and Mantle Hotspots ?, *in* Koeberl, C., and
256 Martínez-Ruiz, F. C., eds., *Impact Markers in the Stratigraphic Record: Berlin, Heidelberg,*
257 *Springer Berlin Heidelberg*, p. 91-119.
- 258 Kring, D. A., Tikoo, S. M., Schmieder, M., Riller, U., Rebolledo-Vieyra, M., Simpson, S. L., Osinski, G. R.,
259 Gattacceca, J., Wittmann, A., Verhagen, C. M., Cockell, C. S., Coolen, M. J. L., Longstaffe, F. J.,
260 Gulick, S. P. S., Morgan, J. V., Bralower, T. J., Chenot, E., Christeson, G. L., Claeys, P., Ferrière,
261 L., Gebhardt, C., Goto, K., Green, S. L., Jones, H., Lofi, J., Lowery, C. M., Ocampo-Torres, R.,
262 Perez-Cruz, L., Pickersgill, A. E., Poelchau, M. H., Rae, A. S. P., Rasmussen, C., Sato, H., Smit, J.,
263 Tomioka, N., Urrutia-Fucugauchi, J., Whalen, M. T., Xiao, L., and Yamaguchi, K. E., 2020,

264 Probing the hydrothermal system of the Chicxulub impact crater: *Science Advances*, v. 6, no.
265 22, p. eaaz3053.

266 Levison, H. F., Morbidelli, A., and Dones, L., 2004, Sculpting the Kuiper belt by a stellar encounter:
267 Constraints from the Oort cloud and scattered disk.: *The Astronomical Journal*, v. 128, no. 5,
268 p. 2553.

269 Marchi, S., Bottke, W., Elkins-Tanton, L., Bierhaus, M., Wuennemann, K., Morbidelli, A., and Kring, D.,
270 2014, Widespread mixing and burial of Earth's Hadean crust by asteroid impacts: *Nature*, v.
271 511, no. 7511, p. 578-582.

272 Mole, D. R., Kirkland, C. L., Fiorentini, M. L., Barnes, S. J., Cassidy, K. F., Isaac, C., Belousova, E. A.,
273 Hartnady, M., and Thebaud, N., 2019, Time-space evolution of an Archean craton: A Hf-
274 isotope window into continent formation: *Earth-Science Reviews*, v. 196, p. 102831.

275 Murphy, J. B., and Nance, R. D., 2003, Do supercontinents introvert or extrovert?: Sm-Nd isotope
276 evidence: *Geology*, v. 31, no. 10, p. 873-876.

277 Nuth, J. A., Barbee, B., and Leung, R., 2018, Defending the earth from long-period comets and sneaky
278 asteroids: Short term threat response requires long term preparation: *Journal of Space Safety*
279 *Engineering*, v. 5, no. 3, p. 197-202.

280 Potter, R. W. K., Kring, D. A., Collins, G. S., Kiefer, W. S., and McGovern, P. J., 2012, Estimating transient
281 crater size using the crustal annular bulge: Insights from numerical modeling of lunar basin-
282 scale impacts: *Geophysical Research Letters*, v. 39, no. 18.

283 Prokoph, A., El Bilali, H., and Ernst, R., 2013, Periodicities in the emplacement of large igneous
284 provinces through the Phanerozoic: Relations to ocean chemistry and marine biodiversity
285 evolution: *Geoscience Frontiers*, v. 4, no. 3, p. 263-276.

286 Puetz, S. J., and Condie, K. C., 2019, Time series analysis of mantle cycles Part I: Periodicities and
287 correlations among seven global isotopic databases: *Geoscience Frontiers*, v. 10, no. 4, p.
288 1305-1326.

289 Rampino, M. R., 1997, Galactic cycle, *Encyclopedia of Planetary Science*: Dordrecht, Springer
290 Netherlands, p. 255-257.

291 Rampino, M. R., Caldeira, K., and Prokoph, A., 2019, What causes mass extinctions? Large
292 asteroid/comet impacts, flood-basalt volcanism, and ocean anoxia—Correlations and cycles:
293 *Special Paper of the Geological Society of America*, v. 542, p. 271-302.

294 Reimink, J. R., Chacko, T., Stern, R. A., and Heaman, L. M., 2014, Earth's earliest evolved crust
295 generated in an Iceland-like setting: *Nature Geoscience*, v. 7, no. 7, p. 529.

296 Riller, U., Poelchau, M. H., Rae, A. S. P., Schulte, F. M., Collins, G. S., Melosh, H. J., Grieve, R. A. F.,
297 Morgan, J. V., Gulick, S. P. S., Lofi, J., Diaw, A., McCall, N., Kring, D. A., Morgan, J. V., Gulick, S.
298 P. S., Green, S. L., Lofi, J., Chenot, E., Christeson, G. L., Claeys, P., Cockell, C. S., Coolen, M. J.
299 L., Ferrière, L., Gebhardt, C., Goto, K., Jones, H., Kring, D. A., Xiao, L., Lowery, C. M., Ocampo-
300 Torres, R., Perez-Cruz, L., Pickersgill, A. E., Poelchau, M. H., Rae, A. S. P., Rasmussen, C.,
301 Rebolledo-Vieyra, M., Riller, U., Sato, H., Smit, J., Tikoo-Schantz, S. M., Tomioka, N., Whalen,
302 M. T., Wittmann, A., Yamaguchi, K. E., Fucugauchi, J. U., Bralower, T. J., and Party, I. I. E. S.,
303 2018, Rock fluidization during peak-ring formation of large impact structures: *Nature*, v. 562,
304 no. 7728, p. 511-518.

305 Schmidt, R. M., and Housen, K. R., 1987, Some recent advances in the scaling of impact and explosion
306 cratering: *International Journal of Impact Engineering*, v. 5, no. 1, p. 543-560.

307 Shibaïke, Y., Sasaki, T., and Ida, S., 2016, Excavation and melting of the Hadean continental crust by
308 Late Heavy Bombardment: *Icarus*, v. 266, p. 189-203.

309 Shoemaker, E. M., 1998, Impact cratering through geologic time.: *Journal of the Royal Astronomical*
310 *Society of Canada*, v. 92, p. 297-309.

311 Smithies, R. H., Champion, D. C., and Van Kranendonk, M. J., 2009, Formation of Paleoproterozoic
312 continental crust through infracrustal melting of enriched basalt: *Earth and Planetary Science*
313 *Letters*, v. 281, no. 3-4, p. 298-306.

314 Smithies, R. H., Lu, Y., Johnson, T. E., Kirkland, C. L., Cassidy, K. F., Champion, D. C., Mole, D. R., Zibra,
315 I., Gessner, K., Sapkota, J., De Paoli, M. C., and Poujol, M., 2019, No evidence for high-pressure
316 melting of Earth's crust in the Archean: *Nature Communications*, v. 10, no. 1, p. 5559.
317 Smithies, R. H., Lu, Y., Kirkland, C. L., Johnson, T. E., Mole, D. R., Champion, D. C., Martin, L., Jeon, H.,
318 Wingate, M. T. D., and Johnson, S. P., 2021, Oxygen isotopes trace the origins of Earth's
319 earliest continental crust: *Nature*, v. 592, no. 7852, p. 70-75.
320 Torres, S., Cai, M.X., Brown, A. G. A., and Zwart, S. P., 2019, Galactic tide and local stellar perturbations
321 on the Oort cloud: creation of interstellar comets.: *Astronomy & Astrophysics*, v. 629, p. A139.

322

323 **Contributions**

324 CLK conceived the research and CLK and PS prepared the first draft of the manuscript with input from
325 TE and TJ. All co-authors assisted in reviewing and editing the revised manuscript.

326 **Competing interests**

327 The authors declare no competing interests.