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## EDITED BY

Wenting Feng,  
Institute of Agricultural Resources and  
Regional Planning (CAAS), China

## REVIEWED BY

Xiangjin Shen,  
Northeast Institute of Geography and  
Agroecology, (CAS), China

## \*CORRESPONDENCE

Tom Sizmur  
t.sizmur@reading.ac.uk

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# Importance of Diurnal Temperature Range (DTR) for predicting the temperature sensitivity of soil respiration

Adetunji Alex Adekanmbi<sup>1,2</sup> and Tom Sizmur<sup>1\*</sup>

<sup>1</sup>Department of Geography and Environmental Science, School of Archaeology, Geography and Environmental Science, University of Reading, Reading, United Kingdom, <sup>2</sup>Department of Soil Science and Land Management, School of Agriculture and Agricultural Technology, Federal University of Technology, Minna, Nigeria

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## Introduction

During the 21st century, global mean temperature is expected to rise by 1.5°C to 5.7°C (1). Climate change has already resulted in an overall decrease in the number of cold days and nights, and an increase in the number of warm days and hot nights, across most land areas globally (2). Our changing climate will influence soil ecosystems because soils have a complex interaction with the atmosphere through carbon, nitrogen, and hydrological cycles (3). Soil is the largest terrestrial carbon pool (4–6), but it also provides a habitat for diverse and complex communities of organisms (7). Soil represents a huge potential source of volatile carbon and a potential sink for additional carbon. Soil can therefore buffer CO<sub>2</sub> losses into the atmosphere, depending on the balance between photosynthesis, autotrophic respiration, and heterotrophic respiration (2, 8). This balance exerts major controls on the biogeochemical interactions between land and atmosphere leading to the exchange of greenhouse gases like CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O (2), the emissions of which could cause positive feedbacks that warm our climate system (9, 10). While the response of autotrophic respiration to changing climates is relatively well understood, predicting changes to the soil carbon sink due to climate change has been a major source of uncertainty in projections. Although it is known that increasing temperature can stimulate microbial degradation of soil organic carbon and increase the atmospheric concentration of CO<sub>2</sub> (10–12), the magnitude of this positive feedback is unclear.

## Changes to the Diurnal Temperature Range (DTR)

Much research on the way that ecosystems respond to changing climate is based on the assumption that the future global warming will arise from symmetrical increases in daily maximum (daytime) and daily minimum (night-time) temperature, reflecting a commensurate increase in daily mean temperature. However, evidence from recent studies shows that daily mean temperature increases arise disproportionately due to increases in daily minimum (night-time) temperatures (13). Over the last 50 years, negative trends in the Diurnal Temperature Range (DTR) have been observed due to an approximately 0.9°C increase in daily minimum and only 0.6°C increase in daily maximum temperature (14). Therefore, while the climate warms, we have a reduction (i.e., narrowing or dampening) of the DTR. The decreasing DTR is attributed to decreases in sunshine duration and increases in the amount of cloud (due to the effects of aerosols), precipitation, and water vapour. However, local deviations from this global trend are expected as different regions will experience different changes in cloud cover, precipitation, and water vapour (15–18). Climate models have predicted that this trend may continue throughout the 21<sup>st</sup> century (13, 19).

The understanding that increased global annual mean temperature will increase global soil respiration rates is well documented in the literature (20). However, the impact of narrowing or decreasing the DTR on soil microbial community structure and physiological functions (such as respiration) remains largely unknown. Laboratory studies that examine how community level interactions will be influenced by climate change usually incubate mesocosms continuously at constant temperatures that mimic a possible future average temperature at the study location (13, 21). Most of these studies do not account for the influence of diurnal temperature oscillations on the temperature sensitivity of soil respiration, especially in the upper few centimetres of soils (e.g. 22). It is important to consider such temperature oscillations because short term temperature fluctuations can have important influences on processes that are driven by microorganisms like bacteria and fungi (23).

## Soil respiration under oscillating temperatures and asymmetric warming

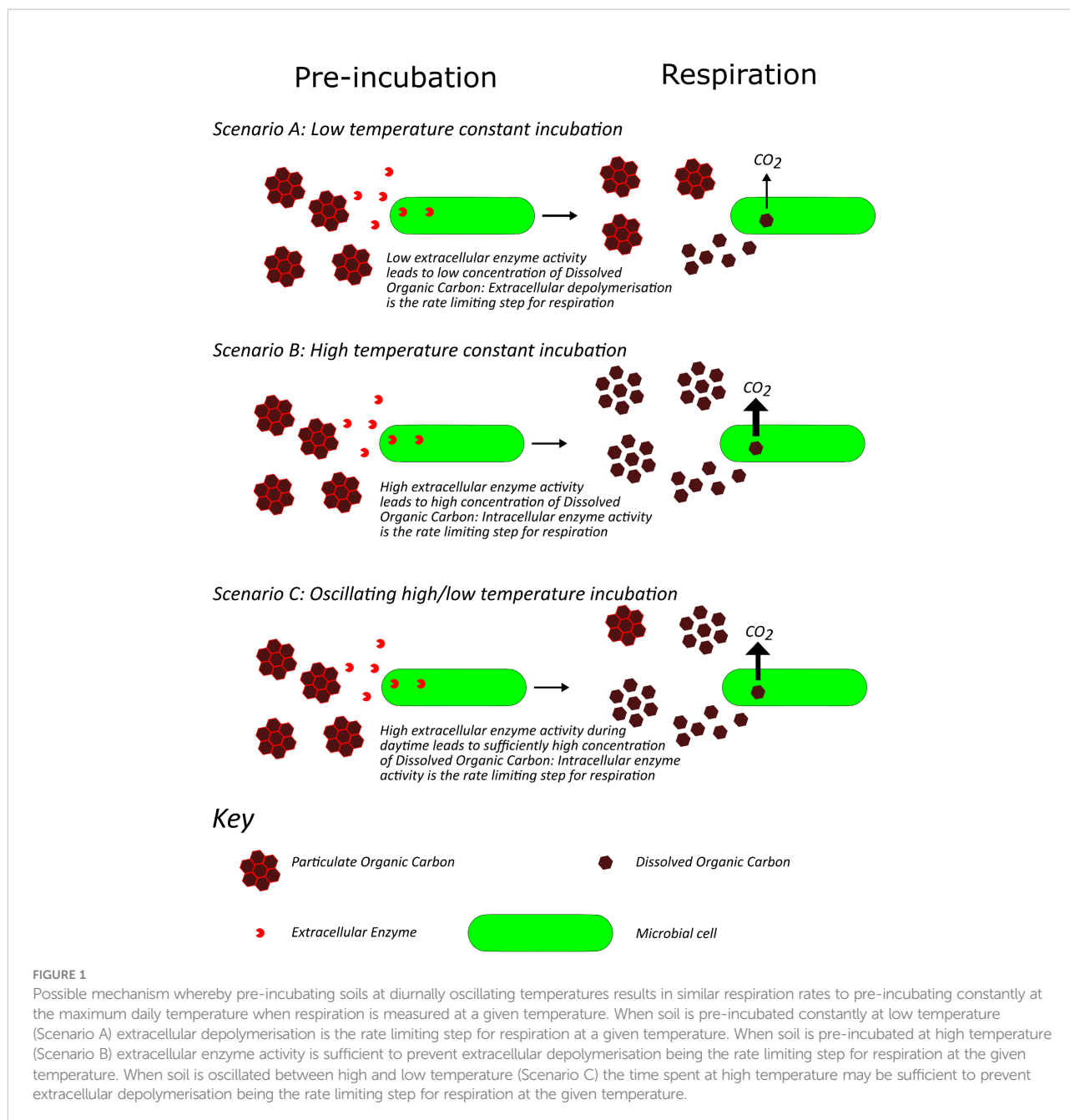
Studies examining the effects of diurnal temperature oscillations on soil respiration are rare. Zhu and Cheng (24) assessed the temperature sensitivity of soil organic carbon decomposition under a constant temperature regime and a diurnally oscillating

temperature regime. Their results showed that temperature sensitivity under constant temperatures were consistently higher than those under diurnally oscillating temperatures. A possible mechanistic explanation for this observation is that microorganisms incubated under constant temperatures adapt to optimise metabolism at a particular temperature and this is exponentially faster at higher temperatures, whereas microorganisms incubated at oscillating temperatures are unable to optimise metabolism at the daily maximum temperature and thus exhibit lower temperature sensitivity. Conversely, Uvarov et al. (25) observed similar cumulative respiration between constant and oscillating temperatures in both the field and laboratory. In this case, the mechanistic explanation for this result is that the bioavailability of substrates to microorganisms (which was similar between the constant and oscillating treatments) may have mediated the respiration rate rather than the temperature regime.

In a field experiment, Xia et al. (26) reported that the effect of diurnal warming on soil respiration was not equal to the summed effects of day and night warming, even though night warming showed greater warming-induced respiration than day warming. They also demonstrated that day and night warming, using infrared radiators, influenced the daily mean soil temperature differently; the daily mean soil temperature increased more under night warming than day warming treatments. Adekanmbi et al. (27) undertook a laboratory incubation experiment and found that the legacy effect of a pre-incubation oscillating between 5°C and 15°C was a significantly greater respiration rate than soil pre-incubated constantly at 5°C or 10°C and was similar to soil pre-incubated constantly at 15°C, even when respiration was measured at the same temperature in all treatments. A possible mechanistic explanation for this finding is that, assuming that extracellular depolymerisation is typically the rate limiting step in soil organic matter decomposition (28), the time spent at 15°C in the fluctuating treatment was sufficient to supply sufficient substrate for intracellular decomposition (Figure 1).

## Impact of land use change on soil Diurnal Temperature Range (DTR) and respiration

Although a reduction in DTR has been forecasted, due to climate change, soil temperature oscillations can also be influenced by land use and altered by land use change. Wei et al. (29) observed an overall decline in DTR between 2003 and 2013 in the agricultural pastoral ecotone of Northern China across croplands, forests, and land that changed between croplands and grasslands. However, they observed an increase in DTR when land use changed from grassland to forest. The lower DTR under grassland was as a result of both decreasing maximum (daytime) temperature and increasing minimum (night-time) temperatures (29). It has been demonstrated that



soil microbes adapt to the temperature of their original site, thus expressing a legacy effect (30). Therefore, while land use change may have (well understood) direct impacts on soil respiration, it also has a poorly understood indirect impact on the temperature sensitivity of respiration due to a change in the temperature regime that soil microorganisms are exposed to. Thus, understanding the impact of diurnal temperature oscillations on the temperature sensitivity of respiration is a pre-requisite for fully understanding the influence of land use change on the soil carbon cycle.

## Soil depth as an analogy for soils with different diurnal temperature ranges

The dampening of the DTR with soil depth represents an interesting analogue for the expected pattern in air or surface temperatures forecasted in future climate scenarios. For instance, de Farias et al. (31) assessed hourly, daily and monthly soil temperature from 5, 10, and 20 cm depths and found that,

while soil temperatures oscillate diurnally down the profile, a decreasing amplitude in DTR was observed with depth (from 5cm to 20cm) due to higher diurnal temperature fluctuations at the surface, compared to lower layers. Whereas soil respiration is usually measured in the laboratory from samples of soil collected at a defined soil depth range, or in the field as the integrated net flux from the entire soil profile (32), there could be a variation in the temperature sensitivity of soil CO<sub>2</sub> flux as a result of thermal diffusivity down the soil profile (33) which is not accounted for in our current understanding. However, the reasons for differential temperature sensitivity with depth also include confounding changes in biotic and abiotic conditions with soil depth. For example, the amplitude of DTR (31), texture, organic carbon (34), abundance and distribution of heterotrophic soil microbial groups (35), and enzyme activities (36) all vary with soil depth. Hicks Pries et al. (22) measured the carbon flux through the whole soil profile in response to 4°C warming and observed similar temperature sensitivity down the profile. Topsoil is where most of the organic carbon and soil microorganisms are concentrated and where the variations in DTR have the greatest impact (37). Nevertheless, soil depth may be a useful analogy for studying the influence of DTR on the temperature sensitivity of soil respiration.

## Discussion

It is clear that a lack of experimental observations severely limits our ability to reach a consensus regarding the influence that diurnal oscillations have on the temperature sensitivity of soil respiration. These limitations to our current understanding mean that we are not yet in a position to predict with certainty the influence that a narrowing of the DTR will have on soil respiration. Nevertheless, predictions of the response of soil carbon fluxes to temperature change clearly do need to account for daily temperature oscillations and future changes to the DTR so that they can accurately predict the impact of climate change on positive soil carbon feedbacks to the atmosphere.

Current projections that use only the size of the soil carbon stock and mean daily temperature to predict the response of carbon loss from soil due to warming may not accurately represent the actual mechanisms and may overestimate total carbon losses [e.g. (38)]. This overestimation is because soil organic matter decomposition, carbon stocks, and persistence, are the products of processes that occur locally and there is need for more robust soil biogeochemical models that better represent how historic temperature condition and short term temperature oscillations shape the response of microbial communities and soil organic matter to warming (39). It is imperative to ensure that the next generation of land-surface models adequately simulate the impact of asymmetric warming, including daily

temperature extremes, on soil microbial activity and soil heterotrophic respiration. We recommend that future projections and soil biogeochemical model parameterisation should only adopt laboratory data to estimate temperature sensitivity when the laboratory data is collected at temperatures within the typical diurnal range (or predicted diurnal range) in the region where predictions are being made. DTR could be used as an input parameter, alongside daily mean temperature, for parameterising soil carbon models.

To satisfy the next generation of soil carbon models, and their requirement to incorporate the influence of a changing DTR, there is an urgent need for laboratory data on the temperature sensitivity of soil respiration that is collected under realistic environmental conditions, including diurnal oscillations. It is understandable why laboratory incubations are currently undertaken at constant temperatures. Laboratory equipment that allows easy establishment of consistent and controllable oscillation of temperatures in the laboratory is not widely available. There will typically be a lag between a change in the temperature in an incubation chamber and the temperature in the soil incubated in a mesocosm inside the chamber, which means that measurements cannot be directly attributed to the temperature in the mesocosm at the time of the measurement. However, for laboratory experiments to represent useful analogies of phenomena that occur under field conditions, efforts need to be made to bridge the gap between uniform laboratory incubations and the stochasticity of the real-world environment.

## Author contributions

TS conceived the idea for the paper. AAA wrote the first draft of the paper. TS provided edits and comments. AAA created subsequent drafts and the final draft of the paper. Both authors contributed to the article and approved the submitted version.

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## Conflict of interest

The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

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