

Lattice deformation of blue halite from Zechstein evaporite basin: Kłodawa Salt Mine, Central Poland

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Abstract Outcrops of natural blue and purple halite crystals have been found in Kłodawa (Poland) salt deposit originating from the Permian (Zechstein) salt formation within tectonic zones. Field works carried out on various levels of Kłodawa Salt Mine indicated differences both in intensity of the hue and in the size of the outcrops. Their occurrence was connected with the presence of epigenetic sediments rich in potassium. For the samples of blue halite, an optical anisotropy (birefringence) was observed for both standard mineralogical thin sections and thick plates, indicating a deviation from cubic symmetry. The blue colouration of the halites, described by ultraviolet–visible spectroscopy, is due to the presence of colour centres. The colour centres have been recently identified as: F, R₁ (F₃), R₂ (F₃), M and *plasmons* (Weselucha-Birczyńska et al., Vib Spectrosc 60:124–128, 2012). The trace amount of impurities detected in the blue halites by means of scanning electron microscopy–energy dispersive X-ray spectroscopy and micro-X-ray fluorescence techniques were connected mainly with solid inclusions such as K_xNa_(1-x)Cl, KMgCl₃·6H₂O, KCl, orthorhombic sulphur, quartz and some other phases like pyrite. Crystallographic data obtained by X-

ray diffraction experiments for the single crystals of the halite from Kłodawa with different saturation of blue or purple colours, as well as for natural colourless halite, revealed lowering of space-group symmetry to monoclinic, orthorhombic, trigonal, tetragonal or even triclinic systems while the space group typical for pure NaCl is *Fm-3m*.

Introduction

The blue colour of halite has been a subject of interest for many scientists. For example, there are several papers describing radiation damage of halite mainly for the usage of exploited salt mines as waste storages (Groote et al. 1991a,b, 1994; Seinen et al. 1992, 1994a,b; Weerkamp et al. 1994; den Hartog 1999; den Hartog and Vainshtein 1997; den Hartog et al. 2002, 2007). A review article by Sonnenfeld (1995) summarised the state of the art of the colour of rock salt. He pointed out factors responsible for the blue colour of halite crystals such as irradiation, soaking in sodium vapours, introducing impurities, inclusions of interstitial matter or structural defects in the crystal lattice. Blue halite occurs frequently in tectonically disturbed zones (tectonites) accompanied by sylvite veins (Davidge and Pratt 1963; Sonnenfeld 1995) and is considered as mineralogical rarity. Especially important for understanding the blue colouration of the halite from Kłodawa Salt Mine seems to be the report on blue halite by Howard and Kerr (1960). They had attributed the blue colouration of natural salt to the bombardment of gamma-ray from the potassium-40 present in sylvite, which accompanied the halite bed. Blue and purple halite variations associated with potassium minerals are typical for Zechstein salt formation found in Kłodawa salt dome belonging to the Polish Zechstein Basin.

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The field work on all mining levels of Kłodawa Salt Mine, prior to laboratory experiments, were summarised by Toboła et al. (2007). In this paper, we present the results of investigations using optical microscopy, ultraviolet–visible (UV–vis) and Raman spectroscopies, scanning electron microscopy–energy dispersive X-ray spectroscopy (SEM-EDS) and crystal structure analyses of blue halite samples from Kłodawa Salt Mine.

Geological setting

The development of the Polish Zechstein Basin as the eastern part of the extensive Central European Basin was strongly controlled by faults that penetrate pre-Permian basement (Ziegler 1990). In the main depocentral axis of the Polish Basin, called the Mid-Polish Trough (MPT), the thickness of Permian sediments was completed with full development of the Zechstein evaporitic basin. These sediments were covered by thick Mesozoic and Quaternary beds. The geological features of MPT including tectonics, lithostratigraphy, correlation between tectonics of the basement and the development of MPT were described by numerous authors (e.g. Marek and Znosko 1972a, b; Wagner 1994; Dadlez et al. 1995; Wagner et al. 2002; Dadlez 2003; Krzywiec et al. 2003; Krzywiec 2006a, b; and references cited therein).

The Zechstein Basin was filled mostly with evaporites which are about 1,500 m thick (Wagner 1994). The lithostratigraphic classification of the basin sediments is based on evaporite sedimentary cycles. Four cyclothems were distinguished in the sediments. Three of them, PZ1, PZ2 and PZ3, were built of carbonates, anhydrites, rock-salt and potassium–magnesium salts. The cyclothems followed transgressive–regressive cycles of the Zechstein Sea. In the cyclothem PZ4, five subcyclothems (PZ4a–PZ4e) were distinguished. These subcyclothems were formed as a result of cyclic climatic changes (Wagner 1994; Pokorski and Wagner 1995) and showed terrigenous–evaporite sequences. These evaporites were covered by younger sedimentary rocks up to 8,000 m thick. The cover and active basement tectonics caused the development of numerous salt-bed structures, most of them associated with the MPT region (Krzywiec 2004a, 2006a, b; Ollier and Koziar 2007). Various aspects of the salt tectonics, such as the development of salts domes in relation to sub-Zechstein basement activity and the influence of the upraising salt structures on surrounding sediments, were analysed by many authors including recent papers by Krzywiec (2004a, b, 2006a, b) and references cited therein.

The Kłodawa salt dome is considered to be the largest diapir in Poland. It is located in Central Poland and belongs to Izbica Kujawska–Łeczyca salt ridge (Fig. 1a, acc. Garlicki and Szybist 1986). The Kłodawa dome, first discovered by gravimetric survey, was finally recognised by several boreholes made in the period of 1946–1950 (Werner et al. 1960). Studies of the

diapir structure (Charysz 1973 and references cited therein; Burliga et al. 1995) indicated that it is about 26 km long, 0.5–2 km wide, elongated in NW–SE direction, and it reaches the depth of 7,000–8,000 m. The cross-section of the upper part of the diapir, shown in Fig. 1b (acc. Szybist 1995), revealed its asymmetric shape with the NE side dipping moderately (55–70°) eastward and with the SW side—almost vertical. The diapir, surrounded by deformed Mesozoic (Triassic–Jurassic) and partly Tertiary deposits, is covered by Tertiary and Quaternary sediments (Garlicki and Szybist 2008). The Kłodawa salt deposit is mainly composed of rocks which belong to fully developed PZ2–PZ4 cyclothems (Fig. 1c, acc. Garlicki and Szybist 1991). The lowermost unit (PZ1 cycle)—consisting of Kupferschiefer (T1), Zechstein Limestone (Ca1) and Lower Anhydrite (A1d)—is known exclusively as tectonically transported blocks (Werner et al. 1960; Burliga et al. 1995). The occurrence of the Oldest Halite (Na1) is questionable due to its petrological similarity to the Older Halite (Na2). In the case of the Older Halite (Na2), high purity of the rock salt and its common occurrence resulted in extensive exploitation of the deposit. The salts of PZ3–PZ4 cyclothems contain impurities, such as the laminae of anhydrite and the intercalations of K–Mg minerals. In the lithological profile of these cycles, the clayey salts (zuber) are also present (Wagner 1994). Moreover, some kinds of pure salt such as lined salt (part of Na3a–PZ3) and pink salt (PZ4) have also been noticed (Charysz 1973; Garlicki 1993). In PZ3 cyclothem, a thick complex (20–30 m) of Younger Potash Salt (K3) was documented (Garlicki and Szybist 2008). The complex is built of kieseritic carnallite.

Due to salt strata deformations, the internal structure of the Kłodawa diapir was found to be complicated. The deformation effects, such as piercing of Older Halite (Na2) through Younger and Youngest Halite, tectonic tailing out and an occurrence of narrow folds with high amplitude, resulted from the various rheological properties of the rocks. The character and style of the deformations vary throughout the Kłodawa diapir (Tarka 1992). In general, two anticlines elongated in NW–SE direction were found to predominate in the SW and NE cross-sections of the diapir (Fig. 1b). The anticlines are separated by a deep central syncline composed of the youngest salt layers.

The blue halite outcrops are localised in KS39 chamber (on the level 600 m) and in the ventilating gallery between chambers KS39 and KS38 (on the level 562 m). Both outcrops are situated within the Older Halite (Na2). Another small outcrop was documented within the lower part of the Younger Halite (Na3a) on the level 600 m. The close contact with sylvinitic veins containing KCl and NaCl mixture is the common feature of all these associations (Natkaniec-Nowak and Toboła 2003; Toboła et al. 2007; Janiów et al. 2008). Substantial outcrop of blue halite was also found on the mining level 750 m near intersection of GTP2 and GTP2a galleries. In contrast to the associations mentioned above, in this outcrop, blue halite is accompanied by epigenetic carnallite.

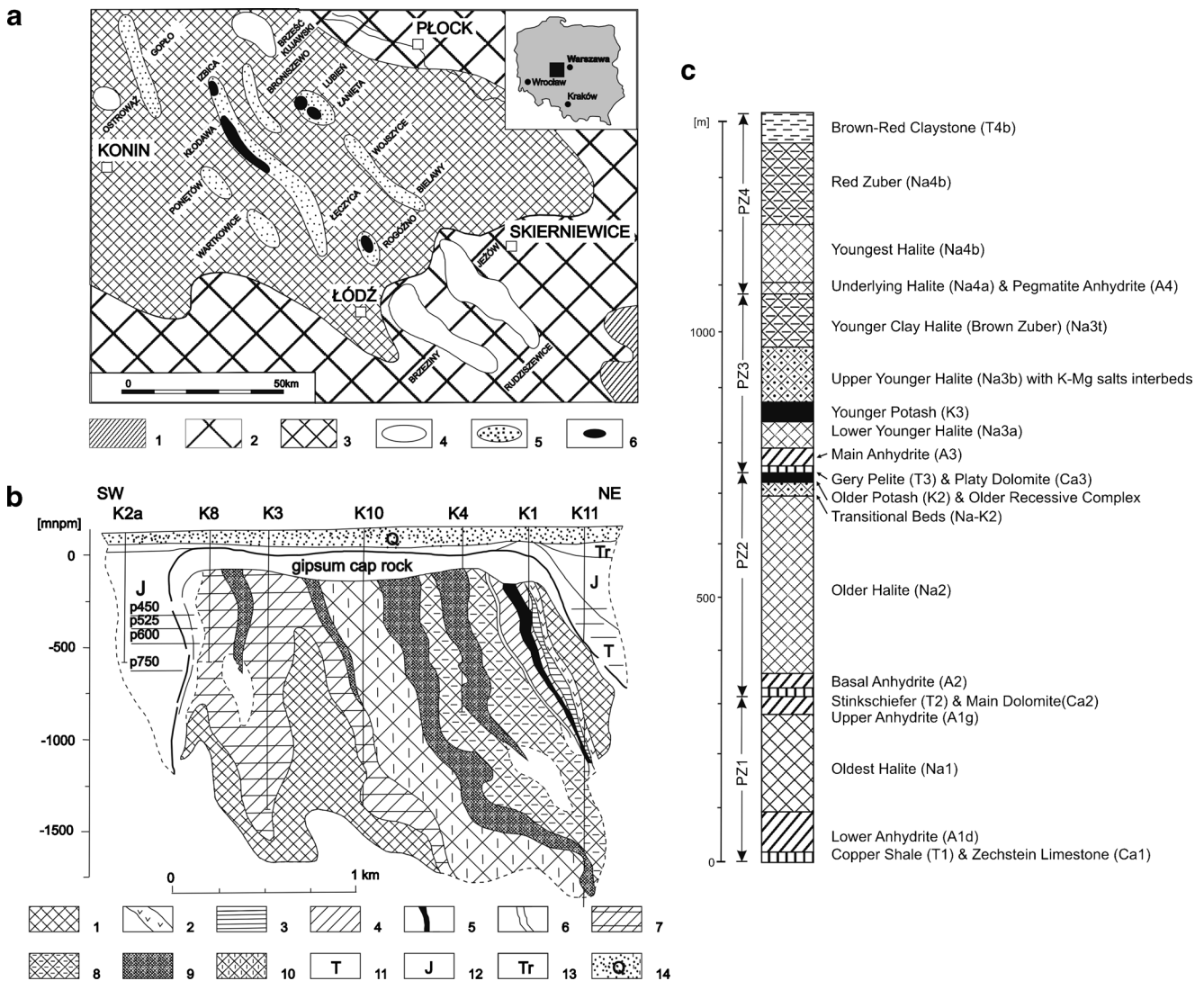


Fig. 1 **a** Part of Zechstein lithofacies map showing the localization of salt structures in the Łódź–Konin–Płock area (acc. Garlicki and Szybist 1986). 1 sulphate-carbonate-littoral facies, 2 chloride facies, 3 chloride facies with K–Mg salts, 4 salt pillows and shafts, 5 salt domes structure partly breaking through the Mesozoic formation, 6 salt domes structure completely breaking through the Mesozoic formation. **b** Geological cross-section through the upper part of Kłodawa salt dome (acc. Szybist

1995): 1 Older Halite, 2 Older Potassium Salt, 3 Grey Salt Clay and Main Anhydrite, 4 Younger Halite, 5 younger potassium salt (kieseritic carnallite), 6 carnallite-bearing association, 7 younger salts association, 8 Brown Zuber and clayey salts, 9 youngest halite (pink), 10 Red Zuber with clayey salts, 11 Triassic formation, 12 Jurassic formation, 13 Tertiary formation, 14 Quaternary formation. **c** Lithostratigraphic profile of Zechstein deposit in Kłodawa salt diapir (acc. Garlicki and Szybist 1991)

Samples and methods

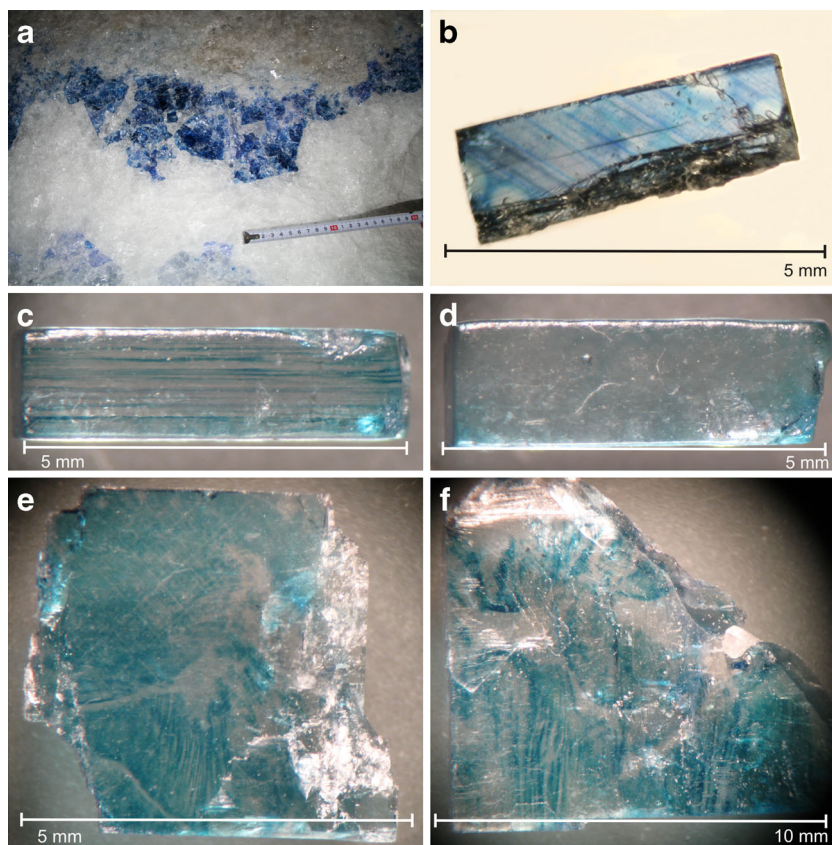
Blue and purple colours of halite have been frequently found in Kłodawa Salt Mine. The coloured halite was recognised on all mining levels (e.g. 450, 525, 562, 600, 750 m). The blue-halite aggregates in the deposit were differentiated with respect to their size and the intensity of the hue. Contacts between the aggregates and surrounding sediments as well as the associations of accompanying minerals were diversified.

The specimens for our studies were sampled from KS39 chamber (on the level 600 m), from the ventilating gallery between chambers KS39 and KS38 (on the level 562 m)—the

blue halite outcrops were accompanied by KCl veins—and from the level 750 m where they are accompanied by the epigenetic carnallite.

In prior of our studies, the contents of radioactive isotopes in the samples of sylvinite and blue halite from Kłodawa Salt Mine were investigated (Kozak, 2007, private information). The specific radioactivity of sylvinite and blue halite was found to be 13,930(630)Bq/kg (K-40) and 1,582(72)Bq/kg (K-40), respectively, whereas there were no Ra-226 and Th-228 isotopes. These values indicate the contents of KCl as 845(41)g per 1,000 g of sylvinite and only 96(4)g KCl per 1,000 g of the blue halite, assuming 0.0117 mol% of potassium-40 isotope in the samples of blue halite.

Fig. 2 **a** Aggregates of the blue halite between sylvine vein (*white salts*) and halite (level 562 m, the ventilation-cumulative gallery). **b, c** Parallel belts of blue colour in halite single crystals. **d** Side view of the halite crystal from **c**. **e, f** Irregular blue zones in single crystals of halite



The saturation of blue hue in the halite aggregates was not homogeneous (Natkaniec-Nowak and Toboła 2003; Toboła et al. 2007; Janiów et al. 2008). It was found that intensive blue or even navy blue colour was characteristic for relatively big halite single crystals 2–6 cm in size (Fig. 2a). These crystals were mostly either euhedral or subhedral with clearly visible cleavage. In macroscopic observations, they displayed zones of dark blue hue or light blue hue alternate with colourless zones (Fig. 2b). Small crystals were usually either anhedral or subhedral. Their linear size ranged from 2 mm to about 2 cm but crystals above 5 mm predominated. The hue of these crystals was also diverse and changed from blue or light blue to light purple. Diversity of colour in single crystals was observed under stereoscopic microscope (Fig. 2c–f). Comparing Fig. 2b–d, it is worth noting that the belts of blue colour in halite single crystals formed layers not always parallel to the cleavage planes and sometimes they might have irregular shapes (Fig. 2e, f).

The mineralogical analyses were performed on standard thin sections in polarised, transmitted light using an OLYMPUS BX51 microscope equipped with DP12 camera. Blue and purple samples of halite were characterised by reflection and absorption UV–vis spectra registered in the range of 200–900 nm using Shimadzu 2101PC spectrophotometer

equipped with ISR 240 device. Additionally, absorption spectra on small areas (0.16 mm²) were obtained using an Ocean Optics “diode array” PC2000 spectrometer. The halite samples of various saturation of blue colour were subjected SEM-EDS analysis. Thick sections were prepared by cracking the selected single crystal along cleavage planes. Analyses were performed using FEI Quanta 200 FEG scanning microscope with an EDS detector. The system was operated at 15 kV accelerating voltage under high vacuum. The different chemical compositions for the single-crystal halite fragments were also examined using the micro-X-ray fluorescence (μ -XRF) spectrometer Bruker M4 TORNADO equipped with an SDD detector. Operating parameters during the measurements were as follows: X-ray tube with Rh anode 30 W with working voltage of 50 kV, the intensity of 600 μ A, spot diameter of 30 μ m (for line Mo–K), 100 s counting time per pixel; SDD detector: the surface area of 30 mm², energy resolution <145 eV (Mn–K). Single-crystal X-ray diffraction was performed for a set of crystals representing all types of observed halite hue: navy blue, pale blue, purple and natural colourless. X-ray intensity measurements of diffracted beams were carried out on a Nonius KappaCCD diffractometer and Agilent Technologies SuperNova, using graphite monochromated MoK α radiation (λ =0.71073 Å) in both cases.

Results

Optical microscopy

For optical microscopy investigation, the crystals representing all observed hue of blue-coloured halite with euhedral shape and cubic {100} cleavage were sampled mainly from mining level 562 m. Borders between single crystal domains were clear. However, sometimes, the spaces between single crystals were filled with fine crystallites showing optical properties similar to anhydrite. In some fragments of halite crystals, an optical anisotropy was observed under polarising microscopy, and the typical examples are shown in Fig. 3. Optical anisotropy was also observed in thick sections (0.5 mm) prepared for fluid inclusion analysis (Fig. 4a, b; compare Weselucha-Birczyńska et al. 2008). Observations under the transmitted-light microscope revealed that the blue halite contains many various solid-phase inclusions well separated from the halite matrix. The most frequent inclusions were those with the oval shape atypical for evaporates (Fig. 4a, c, d). These inclusions usually were birefringent except those of KCl. The size of the inclusions ranged from a few dozen micrometres to few millimetres. Smaller inclusions were almost spherical (Fig. 4c), whereas the bigger ones displayed irregular contours (Fig. 4d). It is worth noting that around solid inclusions in blue halite the radial colourless areas, of 0.1–0.5 mm in thickness, were often observed (compare Figs. 2b, 3 and 4a, c, d). Most of the mineral phases of the solid inclusions in dark blue halite were characterised by corrosion of their borders and changes in optical properties. The results were confirmed by chemical analysis using μ -XRF method for the solid inclusions and indicated the presence of following elements in trace amount: Br, Ca, Cl, Fe, K, S and Ti (Haschke 2010 private information). It is worth pointing out that X-ray single crystal

diffraction performed for the selected solid state inclusions from blue halite crystals led us to distinguish mineral phases such as sylvite, carnallite, orthorhombic sulfur, quartz, anhydrite, polyhalite and pyrite (Stadnicka and Zelek 2008). Additionally, in halite matrix, the trace amount of K, Mg and Cu (0.25 ppm in navy blue, pale blue and purple halite, and 0.42 ppm in natural colourless halite) were detected by atomic absorption spectrometry (AAS) method (Reczyński, 2005 and 2011, private information), and the trace amount of Cu (0.33–0.44 ppm) was confirmed by anodic stripping voltammetry (ASV) method (Baś, 2011, private information).

UV–vis and Raman spectroscopy

Typical absorption spectra for navy blue (A), pale blue (B), purple (C) and natural colourless (D) halite are shown in Fig. 5a. In general, for navy blue samples, two distinct absorption bands were found, at 630 and 385 nm. The spectra for pale blue halite samples had the characteristic absorption band at 610 nm, whereas the band at 580 nm was typical for the purple samples. At the high-energy side of the main broad band, a weak band at about 380 nm has always been observed. It was worth noting that, among blue salt samples from Kłodawa Salt Mine, one could also find amber-coloured fragments. These fragments, while exposed to sunlight, showed an immediate colour change, at first from amber to grey followed by the development of blue hue after a few years (1–5) storage at ambient pressure and temperature ca. 5 °C.

For navy blue NaCl halite (samples A) a very intensive Raman band (Fig. 5b), observed at about 200 cm^{-1} , was successfully resolved into three constituents at 194, 219 and 254 cm^{-1} . The similar band set (Fig. 5b) was also noticed for pale blue samples (B), while it was not seen for purple samples (C) as well as for natural colourless crystals (D). The bands

Fig. 3 Anisotropic zones in the blue halite single crystal observed using polarising microscopy: PX at 0 and 45°

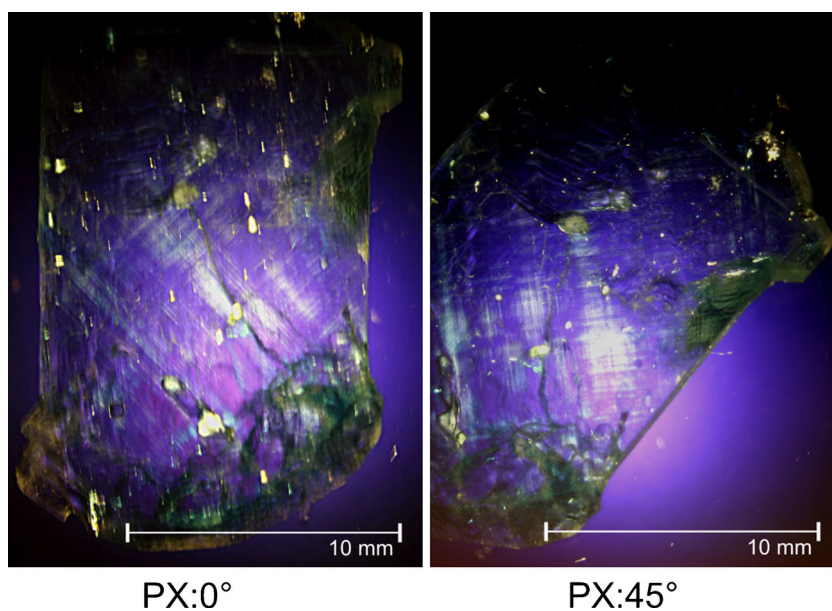
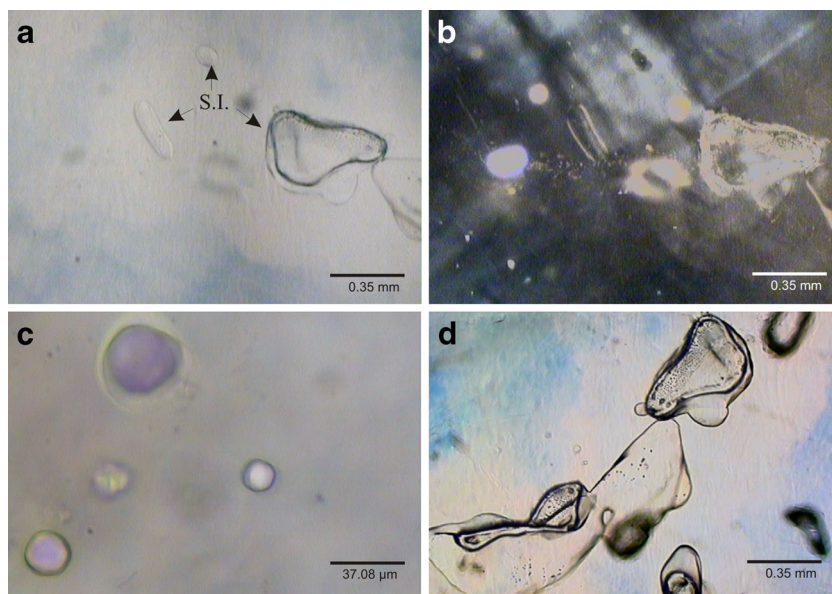


Fig. 4 **a** Arrangement of blue colour in 0.5-mm thick plate of halite (P1), S.I. denotes solid inclusions. **b** The same area between crossed polars (PX)—the birefringence can be observed. **c** Aggregates of small, spherical solid inclusions in blue halite (thick plate, P1). **d** Large solid inclusions with irregular rounded shape (thick plate, P1)



were destroyed while samples were heated from room temperature to 450 °C.

SEM-EDS investigations

The samples for SEM-EDS data acquisition were separated from halite with different saturations of blue hue. Some examples of backscattered electron microscopy images of halite samples are presented in Fig. 6 (for other examples, see Toboła and Natkaniec-Nowak 2008). The concentration of other elements coexisting with sodium and chloride in halite samples as found from semiquantitative EDS analysis are given in Table 1. In all investigated samples, the main rock was determined as pure halite, whereas other phases were well separated in the forms of solid state inclusions. EDS of navy blue single

crystals from level 562 m indicated pure halite matrix (NaCl) with KCl as solid state inclusions. It is worth noting that there is no difference between the composition of single crystal of blue halite and single crystal of natural colourless halite from the level 562 m except that the colourless halite contained more solid state inclusions with trace amounts of K, Mg, Al, Ca, Fe and S, Si. The presence of K, Mg, Cl and O indicated carnallite. The source of Fe and S could be pyrite, whereas that of Ca, S and O could be anhydrite. In addition to quartz, Al, Si and O, together with K, Ca, Fe and Mg, could suggest the presence of some aluminosilicates. Oxygen accompanied Na, K and Cl might be from the water molecules. Two-dimensional EDS elemental distribution maps for blue halite sample of Fig. 6b from level 562 m revealed pure NaCl matrix and well-separated KCl solid state inclusion (Fig. 7).

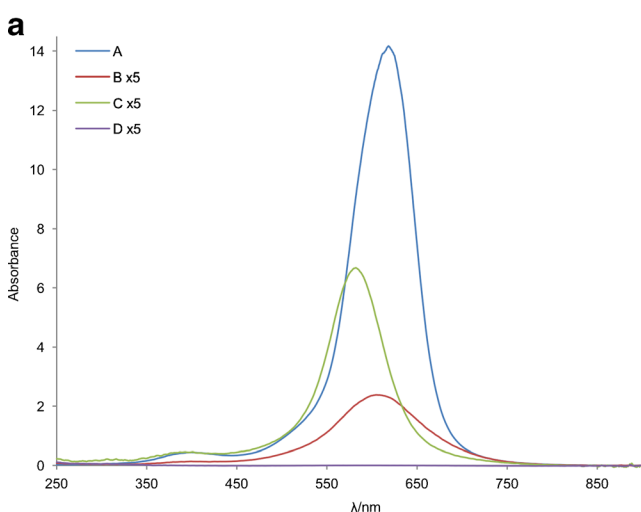
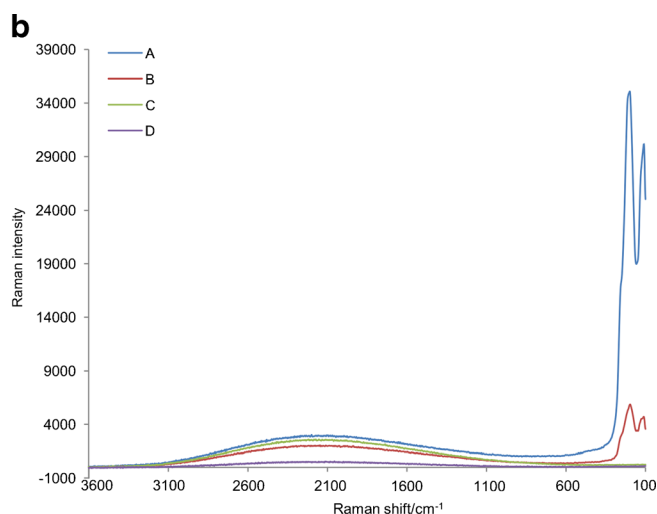
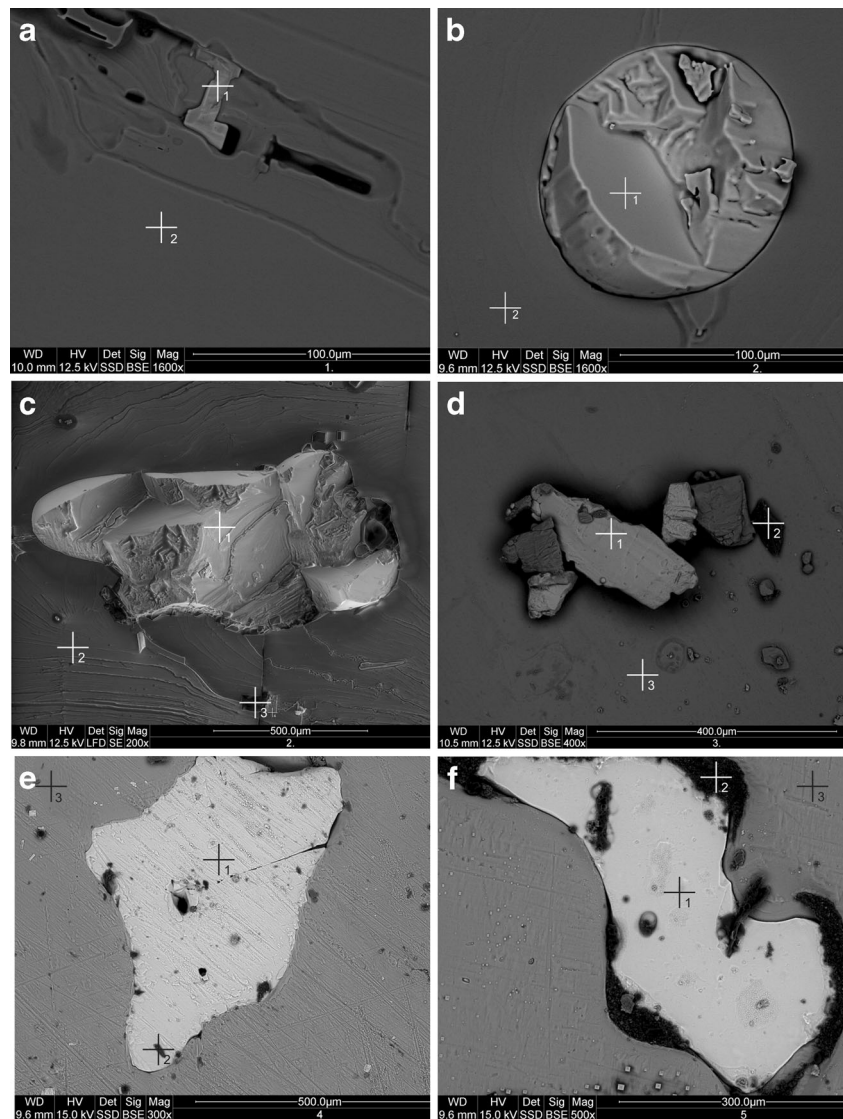


Fig. 5 **a** UV-vis spectra for halite samples from Klodawa Salt Mine: navy blue (A)—blue line, pale blue (B)—red line, purple (C)—green line and colourless (D)—violet line (the values for B, C and D were multiplied



by 5 for the sake of clarity). **b** μ-Raman spectra for single-crystal samples A, B, C and D

Fig. 6 Backscattered electron microscopy (BSE) images of halite single crystals containing: navy blue halite (**a, b**) from level 562 m, colourless halite (**d**), and light blue (**e, f**) from level 750 m. For navy blue halite from level 562 m, secondary electron image (SE) is shown in **c**. Chemical analysis was performed for each point marked by cross (for other examples, see Toboła and Natkaniec-Nowak 2008)



Single crystal X-ray diffraction for halite matrix

The analyses of crystallographic data, such as lattice parameters, unit-cell volume and various space group assignment were derived from the diffraction patterns. The results are given in Table 2 (data from KappaCCD diffractometer) and Table 3 (data from SuperNova diffractometer).

The intensity data collected with the KappaCCD diffractometer revealed significant deviations of intensity distribution from that characteristic for the cubic NaCl (Table 2). The halite samples of navy blue hue (A) predominantly showed the deformation of symmetry to monoclinic space group $C2/m$, rarely to tetragonal one ($I4/mmm$). For the samples of pale blue hue (B), the deformations to rhombohedral $R\bar{3}m$ (for crystals with greenish blue shade) and to orthorhombic $Fmmm$ or $Immm$ space group (for crystals with light blue shade) were typical except for one case of monoclinic deformation for the

sample B2 originating from level 750 m. In the purple samples (C), cubic space group $Fm\bar{3}m$ was usually preserved.

Single crystal X-ray diffraction for solid inclusions in halite matrix

The selection of single crystals of solid inclusions found in different types of halite (A–C) was done with polarising microscopy. X-ray diffraction measurements using monochromated $MoK\alpha$ radiation were performed only for those single crystals that were stable in ambient conditions after their isolation from inclusions. The crystallographic data for identified phases are listed in Table S1. The crystal structures of solid state inclusions were solved using SIR92 (Altomare et al. 1994). Tables S2 presents residual factors obtained during refinement procedure (SHELXL97, Sheldrick 2008).

Table 1 Element concentration (at.%) from EDS analysis for halite samples from Kłodawa Salt Mine

Elements	At.%															
	A		B		C			D			E			F		
Samples	1	2	1	2	1	2	3	1	2	3	1	2	3	1	2	3
Na	28.81	48.55	21.95	49.70	9.98	48.21	16.05	18.06	30.51	48.49	8.90	13.34	50.26	12.15	25.14	49.26
K	20.21	–	29.28	0.50	37.24	0.86	5.90	31.47	1.84	0.54	39.23	25.16	0.89	33.96	5.03	0.67
Ca	–	0.32	–	–	–	–	–	–	2.84	–	–	–	–	–	1.44	–
Al	–	–	–	–	–	–	–	–	2.27	–	–	–	–	–	14.03	–
Si	–	–	–	–	–	–	–	–	12.55	–	–	–	–	–	0.23	–
Fe	–	–	–	–	–	–	–	–	2.38	–	–	–	–	–	–	–
Mg	–	–	–	–	0.44	0.31	8.51	–	0.66	–	–	2.15	–	–	–	–
Cl	49.56	50.01	48.77	49.80	48.30	49.36	30.76	50.47	30.52	48.90	46.28	35.23	46.35	46.29	28.94	47.80
O	1.42	1.13	–	–	4.04	1.26	38.78	–	15.40	2.07	5.59	24.13	2.51	7.60	23.96	2.28
S	–	–	–	–	–	–	–	–	1.03	–	–	–	–	–	1.24	–
Total	100	100.01	100	100	100	100	100	100	100	100	100	100.01	100.01	100	100.01	100.01

Standardless ZAF corrections for semi-quantitative element normalised data were used

Discussion

The halite single crystals from Kłodawa Salt Mine, both with blue hue and colourless ones, indicated optical anisotropy. All birefringent fragments were found in the halite crystals of pronounced cleavage. According to Shlichta (1968), the birefringence could be considered as a result of stress and mechanical deformations of halite beds. Most of the mineral phases of the solid inclusions in dark blue halite were characterised by corrosion of their borders and changes in optical properties resulting from the secondary alterations of rocks (Heflik et al.

2008). It is worth noting that, around colourless solid-state inclusions in blue halites, there was usually observed colourless surrounding in close vicinity of the inclusion, whereas around blue halite inclusion in colourless matrix the close surrounding was blue (compare Figs. 2, 3 and 4).

The results of chemical composition analyses indicated that the blue halite samples consist of pure NaCl, whereas the source of trace amounts of other elements are the inclusions. Despite of the investigation of gaseous and liquid inclusions in the halite from Kłodawa Salt Mine by Raman spectroscopy, which was published recently by Weselucha-Birczyńska et al.

Fig. 7 Examples of EDS elemental distribution maps for blue halite single crystals from level 562 m containing rounded small solid inclusions presented in the Fig. 6b

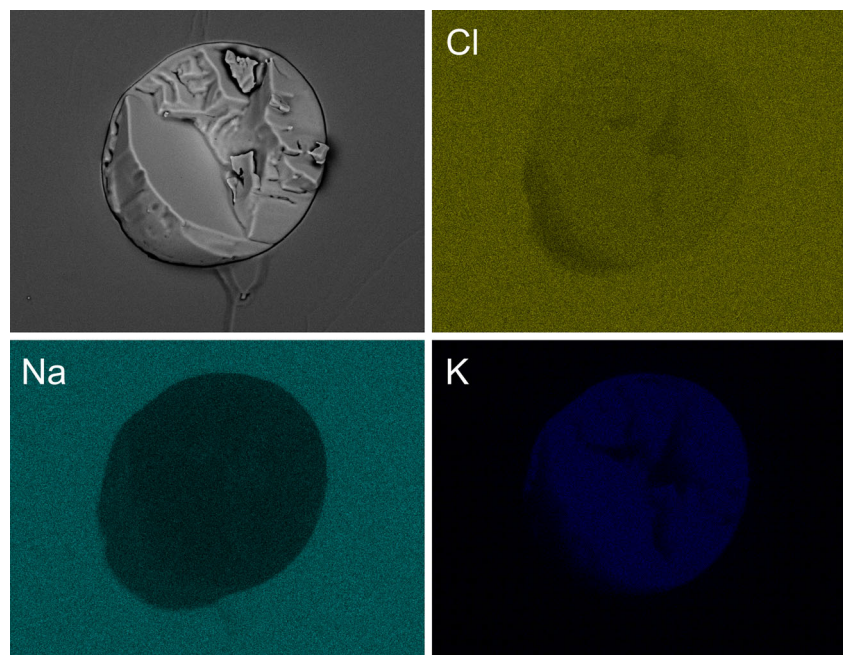


Table 2 Space group symmetry and lattice parameters for representative samples of halite from Klodawa Salt Mine

Single-crystal fragment	Preliminary space group	θ_{\max}	Space group*	Number of unique reflections	a (Å)	b (Å)	c (Å)	α (°)	β (°)	γ (°)	$V_{k.cl}$ (Å ³)
A1—navy blue	<i>I2/m</i>	34.35	<i>Fm-3m</i> ^a	35	5.6373(4)	5.6384(2)	5.6384(2)	90.022(4)	89.894(4)	89.894(4)	179.22(2)
A1 ₂ —colourless	<i>Fm-3m</i>	34.31	<i>Fm-3m</i>	35	5.6384(3)	5.6384(3)	5.6384(3)	90.000	90.000	90.000	179.25(2)
A1 ₃ —navy blue	<i>I4/mmm</i>	36.22	<i>Fm-3m</i> ^a	39	5.6415(3)	5.6421(2)	5.6421(2)	90.000	90.000	90.000	179.59(1)
A1 ₄ —navy blue	<i>C2</i>	38.16	<i>F23</i> ^b	89	5.6350(2)	5.6350(9)	5.6350(9)	90.000	90.000(9)	90.000	178.90(6)
A2 ₁ —navy blue	<i>C2/m</i>	34.97	<i>Fm-3m</i>	35	5.6410(3)	5.6406(4)	5.6410(3)	89.995(5)	90.023(6)	90.005(6)	179.49(6)
A2 ₂ —colourless ^c	<i>Fm-3m</i>	35.27	<i>Fm-3m</i>	38	5.6408(3)	5.6408(3)	5.6408(3)	90.000	90.000	90.000	179.48(3)
B1—blue	<i>Fmmm</i>	34.31	<i>Fmmm</i>	121	5.6382(3)	5.6394(3)	5.6522(3)	90.000	90.000	90.000	179.72(2)
B2—blue	<i>C2/m</i>	25.03	<i>Fm-3m</i>	22	5.6639(4)	5.6649(3)	5.6649(3)	90.181(6)	90.234(6)	90.234(6)	181.76(4)
C1—purple	<i>Fm-3m</i>	34.29	<i>Fm-3m</i>	35	5.6415(2)	5.6415(2)	5.6415(2)	90.000	90.000	90.000	179.55(1)
C2—purple	<i>Fm-3m</i>	35.25	<i>Fm-3m</i>	37	5.6435(4)	5.6435(4)	5.6435(6)	90.000	90.000	90.000	179.74(3)
D1—colourless	<i>Fm-3m</i>	34.29	<i>Fm-3m</i>	35	5.6416(2)	5.6416(2)	5.6416(2)	90.000	90.000	90.000	179.56(1)
D2—colourless	<i>C2/m</i>	33.04	<i>Fm-3m</i>	33	5.6440(2)	5.6440(2)	5.6440(2)	90.000	90.000(10)	90.000	179.8(1)

Diffractionmeter Nonius KappaCCD, 20 °C, $\lambda(\text{MoK}\alpha)=0.71073$ Å (55 kV, 30 mA)

^a Space group *Fm-3m* imposed, calculations using the program Platon (Spek 2009)

^b Refinement on *Fm-3m* and *Fm-3* unstable

^c After the recrystallization of the navy blue fragment of sample A1 from aqueous solution

(2008), there are many solid state inclusions in the form of single crystals.

Two distinct absorption bands at 630 and 385 nm observed for navy blue halite are responsible for superimposed complementary colours: blue and yellowish green, respectively. The characteristic absorption band at 610 nm for pale blue halite samples is responsible for clear blue colour, and the band at 580 nm for the purple colour. The findings are consistent with the earlier work by Zelek et al. (2008). Similar spectroscopic properties were also observed for Stassfurt blue halite samples (also Zechstein deposit) by Phipps and Brode as early as in 1926. The blue hue of halites is due to the colour centre aggregates. The centres were identified as: F, R₁ (F₃), R₂ (F₃) bands, *plasmon* band and M-band. The populations of these colour centres were recently estimated by Wesołucha-Birczyńska et al. (2012). Usually in the studied samples, the dispersed discrete F-centres with characteristic band at

465 nm (compare Howard and Kerr 1960) display very low population. The detailed interpretation of the colour centres behaviour under heating is going to be published elsewhere.

The observed Raman spectra with the bands in the region 100–300 cm⁻¹ (194, 219 and 254 cm⁻¹, acc. Wesołucha-Birczyńska et al. 2012) confirmed the presence of colour centres and also metal colloid particles in the studied halite samples. The bands were destroyed while samples were heated, which indicated the appropriate assignment of the bands.

The diffraction results obtained with a KappaCCD diffractometer prompt us to constrain the experimental data to the ideal structure (six Cl⁻ ions in the corners of Na⁺ octahedron and vice versa) and to space group *Fm-3m* typical for pure, perfect NaCl. In this procedure, the intensities which failed to fulfil *Fm-3m* space-group conditions were omitted and averaged lattice parameters were compared to those obtained from measurements performed with SuperNova diffractometer.

Table 3 Space group symmetry and lattice parameters for representative samples of halite from Klodawa Salt Mine

Single-crystal fragment	Space group	θ_{\max}	Number of reflections	a (Å)	b (Å)	c (Å)	α (°)	β (°)	γ (°)	$V_{k.cl}$ (Å ³)
A3—navy blue	<i>Fm-3m</i>	40.65	50	5.6469(2)	5.6444(2)	5.6442(2)	90.022(3)	90.027(3)	90.009(3)	179.90(1)
A4—navy blue	<i>Fm-3m</i>	40.57	48	5.6520(3)	5.6512(4)	5.6524(4)	90.018(5)	90.013(5)	89.993(5)	180.54(2)
A5—navy blue	<i>Fm-3m</i>	40.64	45	5.6434(8)	5.6364(8)	5.6440(8)	90.081(12)	89.971(11)	89.983(11)	179.53(4)
B3—blue	<i>Fm-3m</i>	40.65	49	5.6422(3)	5.6420(3)	5.6424(2)	90.020(4)	90.001(4)	90.003(4)	179.62(2)
C3—purple	<i>Fm-3m</i>	45.42	59	5.6453(2)	5.6454(2)	5.6446(2)	90.010(3)	89.995(3)	90.025(3)	179.89(1)
D3—colourless	<i>Fm-3m</i>	40.85	51	5.6500(2)	5.6498(2)	5.6454(3)	90.022(4)	90.015(4)	89.980(4)	180.21(1)

Diffractionmeter Agilent Technologies SuperNova, 25 °C, $\lambda(\text{MoK}\alpha)=0.71073$ Å (50 kV, 0.80 mA)

Table 4 Lattice parameters of synthetic solid solution $K_xNa_{1-x}Cl$ taken from PDF-2 database (Faber and Fawcett 2002) and used for the derivation of relationship between the lattice parameters and chemical compositions: $y=0.6513x+5.6407$ ($R^2=1$)

Synthetic solid solution	x	a (Å)
NaCl	0	5.6402
$K_xNa_{1-x}Cl$	0.2	5.7712
$K_xNa_{1-x}Cl$	0.4	5.9018
KCl	1	6.2917

The averaged lattice parameters were then used to estimate the potassium contents in blue halite samples. For several samples (B1, C2, D2, C3, A3, D3, A4 and B2), despite their various characteristic hue, significant enlargement of lattice parameters was observed when compared to the value typical for the colourless one, i.e. $a=5.6408(3)\text{Å}$ found for $A2_2$ crystals obtained through the recrystallisation of the navy blue $A1$ sample from aqueous solution. The enlargement of lattice parameters might indicate the presence of small contents of potassium ions in these crystals. For some samples, e.g. $A1_2$ and $A1_4$ (navy blue) or $A1_1$ (a natural colourless part of $A1$), the significant contraction of lattice parameters was observed, when compared to $a=$

$5.6408(3)\text{Å}$. The presence of vacancies could be considered as a reason of the contraction. For the purpose of the estimation of potassium contents, a relationship between the values of cubic lattice parameter and possible chemical composition was derived assuming the formation of KCl/NaCl solid solution and using the literature data for synthetic $K_xNa_{(1-x)}Cl$, where $x=0.0, 0.2, 0.4$ and 1.0 (PDF-2 database, Faber and Fawcett 2002). The lattice parameters determined for the solid solutions are given in Table 4. The equation obtained was then applied to the averaged lattice parameters of the studied samples (Fig. 8), indicating the highest potassium contents in the pale blue sample B2 ($x=0.0366$). The relatively high potassium contents were also found for the pale blue sample B1 ($x=0.0039$), in which it was similar to the value of 3,024 ppm obtained by AAS method for the pale blue sample B8 (Reczyński, 2005 and 2011, private information).

The solid state inclusions identified by single-crystal diffraction methods were the following: those isolated from the navy blue halite of type A were found to be sylvine (KCl), orthorhombic sulphur (S_8) and α -quartz ($\alpha\text{-SiO}_2$); for the pale blue halite of type B, mainly carnallite ($KMgCl_3\cdot 6H_2O$) was identified, and the most common solid inclusions in the halite type C were pyrite (FeS_2) and carnallite.

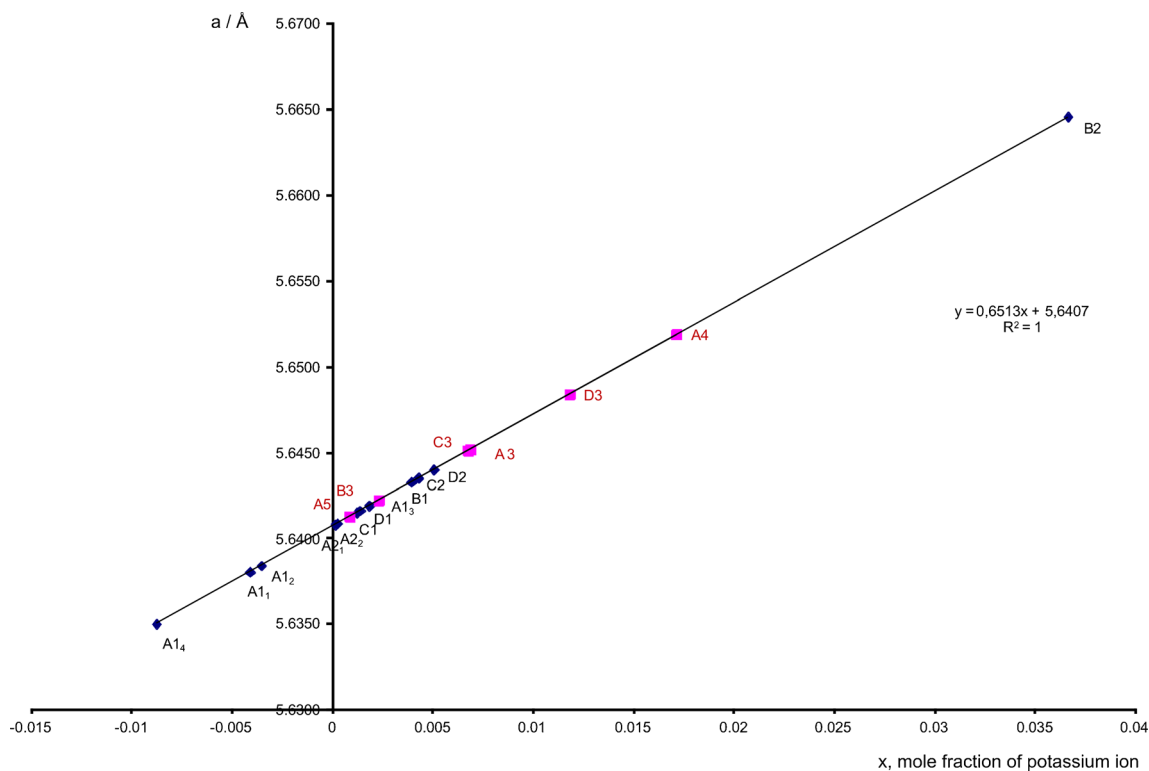


Fig. 8 Lattice parameters versus chemical composition parameter x for single crystals of halite from Kłodawa Salt Mine assuming solid solution $K_xNa_{(1-x)}Cl$ (the line represents the relationship according to Table 4).

The sample numbering is consistent with that used in Table 2 (marked by navy blue diamonds) and Table 3 (marked by pink squares)

Conclusions

Fieldwork observations carried out in Kłodawa Salt Mine indicated that blue halite occurrence was mainly connected with the epigenetic sediments, such as sylvinite rich in potassium. The high contents of potassium-40 could be considered as the source of gamma-ray necessary for the production of scattered F-centres (compare Howard and Kerr 1960). In particular thermodynamic conditions, the dispersed F-centres might possibly aggregate. The formation of such aggregates (F-centre clusters) was pointed out by Farge and Fontana (1979), Georgiev (1988) and also by Seinen et al. (1994a). It should be emphasised that the chemical impurities are not responsible for the blue colouration of the halites.

The presence of epigenetic sediments does not always guarantee the occurrence of blue halite, though the co-occurrence of blue halite with the primary potassium-magnesium salts, such as carnallite and polyhalite, was rarely observed. Additionally, the blue halite associations together with the epigenetic sediments were related to tectonites (Sonnenfeld 1995; Davidge and Pratt 1963). The tectonic deformation of the salt strata observed in macroscopic scale has been manifested in microscopic scale as a specific birefringence in thin samples not only of blue halite but also of natural colourless halite from the Kłodawa Salt Mine. The influence of geological environment seems to be a significant condition for the formation and stabilisation of blue halite deposits.

From single crystal X-ray diffraction study of the blue halites, it could be concluded that, in Kłodawa Salt Mine, not only halite beds are highly deformed but also the crystal structure of blue halite is deformed at atomic level. The deformation degree of the blue halite structure from that characteristic for the pure cubic NaCl could be roughly correlated with the observed hue of the blue colour which is a result of a particular colour centre aggregation.

In our opinion, the clusters of various types of colour centres seem to be stabilised by different sort of already existing lattice defects. The size of colour centre clusters and the type of structure deformation at atomic level would have an effect on the stabilisation of the whole system and thus could be regarded as responsible for the observed variety of absorption bands and consequently for stable, at ambient conditions, blue and/or purple hue of the halites from the Kłodawa Salt Mine.

Main deformation mechanism under the stress conditions concerned dislocation creep and some processes of static recrystallisation, which however did not eliminate the primary numerous defects (such as dislocation debris) in halite crystals (compare Schléder et al. 2007). After irradiation, the discrete F-centres should be randomly distributed in halite crystals; however, the deformed crystals contain defects that can act as electron traps and assist clustering processes of colour centres (Davidge and Pratt 1963). The clusters of colour centres are stabilised at the defects.

Our conclusions seems to be consistent with the results of the analysis of microstructures in halite samples from the Kłodawa Salt Mine published recently by Schléder et al. (2007) but different, to some extent, from the statement by Howard and Kerr (1960) that the gamma-ray irradiation is exclusively causing the deformation of the halite.

The main issue derived from the experiments performed concerning the generation of the colour centre clusters and their stabilisation by structural deformation still remains to be verified. In the future, it would be necessary to find a qualitative relationship between the hue of halite and the size of colour centre clusters, in order to consider the influence of the energy and intensity of isotope radiation on the type of the produced single, scattered colour F-centres, and to discuss the possibility of colour centres aggregation along with the degree of structure deformations of the halite from the Kłodawa Salt Mine.

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