Latest progresses in ShUREX* (2015-2017) data analyses

(Estimation of kinetic energy dissipation rates from Pitot, CWT and radar data)

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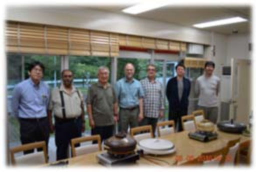
*ShUREX: Shigaraki- UAV – Radar Experiment

Introduction to ShUREX campaigns

P.I.: Prof. Lakshmi Kantha (Colorado University, USA)

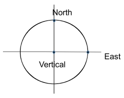


The ShUREX dream team!

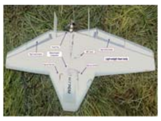
Instrumental set up (2016, 2017)

MU Radar

3 beams (0°,0°), (0°,10°), (90°,10°),



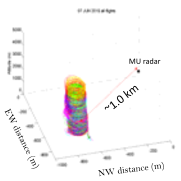
Acquisition time for one Doppler spectrum: 24.57 sec
Range sampling from 1.34 km to 20.35 km ($\Delta r = 150\ m$)



A **Datahawk UAV** equipped with mutiple sensors (**Kantha et al., 2017**).

Sensor type	CWT / Pitot
Sampling rate	800 Hz/ 400 Hz

Number of useable flights (on Sep 2018): 39



Maximum height (UAV): 4-5 km

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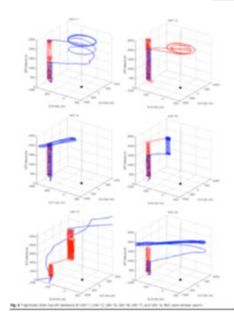
Examples of UAV trajectories



After Kantha et al. 2017

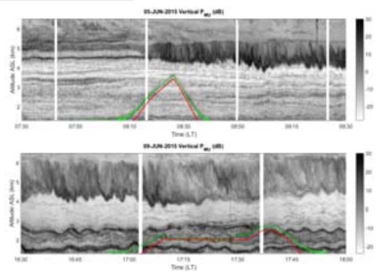


Figure 3: (a) Time-height cross-section of echo power P_{MU} (dB) at vertical incidence after doing the Capon processing in the height range 1.27-6.5 km from 07:30 LT to 09:30 LT on 05 June 2015. (b) Same as (a) from 16:30 LT to 18:00 LT on 09 June 2015.

After Luce et al. 2018

Focus of the presentation

Comparing **turbulence kinetic energy (TKE) dissipation rates** ε estimated from:

(1) relative air speeds measured by Pitot sensor (UAV)

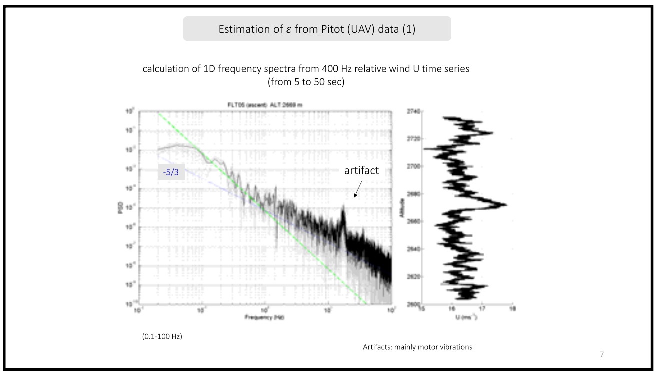
(2) MU radar Doppler spectral width using various models

(3) Temperature measurements by CWT sensors using a theoretical relationship between C_T^2 and ε

Part I: (1) – (2)

Part II: (3) – (1)

Part I: radar – Pitot comparisons


Estimation of ε from Pitot (UAV) data (1)
calculation of 1D frequency spectra from 400 Hz relative wind U time series
(from 5 to 50 sec)
-5/3
artifact
Frequency (Hz)
Altitude
(0.1-100 Hz)
Artifacts: mainly motor vibrations
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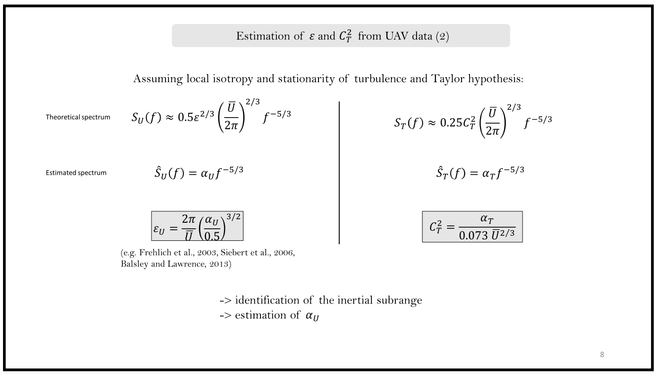
Estimation of ε and C_T^2 from UAV data (2)
Assuming local isotropy and stationarity of turbulence and Taylor hypothesis:
Theoretical spectrum
$S_U(f) \approx 0.5\varepsilon^{2/3}\left(\frac{\bar{U}}{2\pi}\right)^{2/3} f^{-5/3}$
$S_T(f) \approx 0.25C_T^2\left(\frac{\bar{U}}{2\pi}\right)^{2/3} f^{-5/3}$
Estimated spectrum
$\hat{S}_U(f) = \alpha_U f^{-5/3}$
$\hat{S}_T(f) = \alpha_T f^{-5/3}$
$\varepsilon_U = \frac{2\pi}{\bar{U}}\left(\frac{\alpha_U}{0.5}\right)^{3/2}$
$C_T^2 = \frac{\alpha_T}{0.073\ \bar{U}^{2/3}}$
(e.g. Frehlich et al., 2003, Siebert et al., 2006, Balsley and Lawrence, 2013)
-> identification of the inertial subrange
-> estimation of α_U
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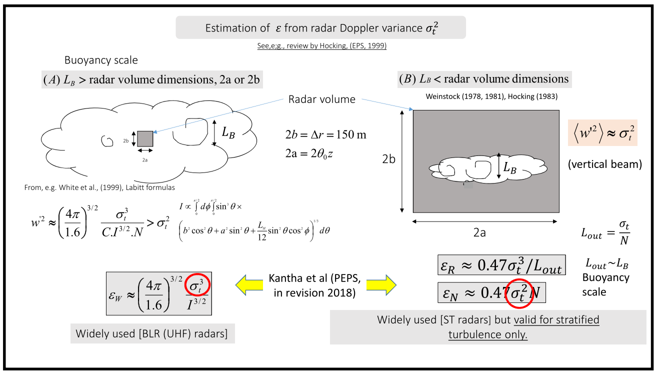
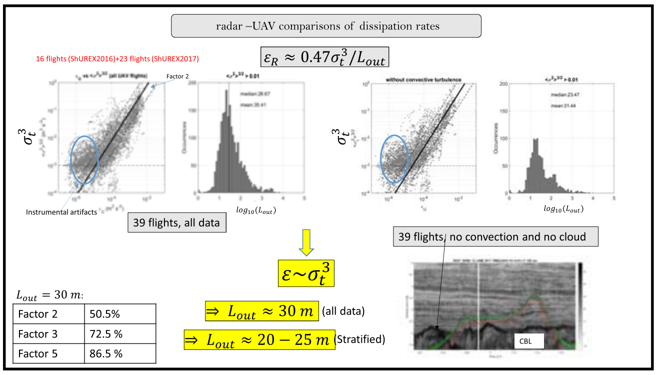
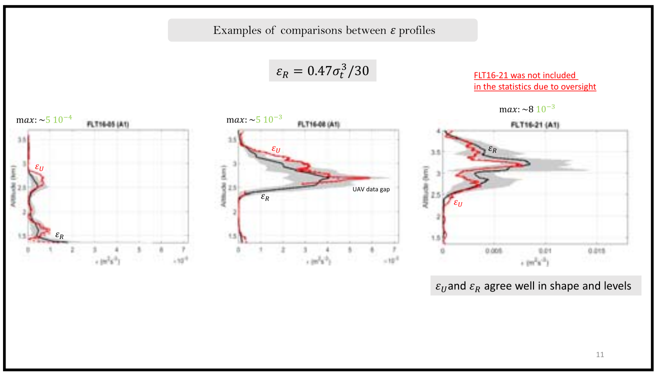

Estimation of ε from radar Doppler variance σ_t^2
See e.g., review by Hocking, (EPS, 1999)
Buoyancy scale
(A) L_B > radar volume dimensions, 2a or 2b
(B) L_B < radar volume dimensions
Weinstock (1978, 1981), Hocking (1983)
Radar volume
$2b = \Delta r = 150$ m
$2a = 2\theta_0 z$
2b
2a
$\langle w'^2 \rangle \approx \sigma_t^2$
(vertical beam)
From, e.g. White et al., (1999), Labitt formulas
$L_{out} = \frac{\sigma_t}{N}$
$\varepsilon_W \approx \left(\frac{4\pi}{1.6}\right)^{3/2} \frac{\sigma_t^3}{I^{3/2}}$
Kantha et al (PEPS, in revision 2018)
$\varepsilon_R \approx 0.47\sigma_t^3/L_{out}$
$\varepsilon_N \approx 0.47\sigma_t^2 N$
$L_{out} \sim L_B$
Buoyancy scale
Widely used [BLR (UHF) radars]
Widely used [ST radars] but valid for stratified turbulence only.


radar –UAV comparisons of dissipation rates
16 flights (ShUREX2016)+23 flights (ShUREX2017)
$\varepsilon_R \approx 0.47\sigma_t^3/L_{out}$
Factor 2
Instrumental artifacts
39 flights, all data
$log_{10}(L_{out})$
without convective turbulence
39 flights, no convection and no cloud
$\varepsilon \sim \sigma_t^3$
$L_{out} = 30$ m:
Factor 2 | 50.5%
Factor 3 | 72.5 %
Factor 5 | 86.5 %
$\Rightarrow L_{out} \approx 30$ m (all data)
$\Rightarrow L_{out} \approx 20 - 25$ m (Stratified)
CBL


Examples of comparisons between ε profiles
$\varepsilon_R = 0.47\sigma_t^3/30$
FLT16-21 was not included in the statistics due to oversight
max: ~5 10^{-4}
max: ~5 10^{-3}
max: ~8 10^{-3}
FLT16-05 (A1)
FLT16-08 (A1)
FLT16-21 (A1)
UAV data gap
ε_U and ε_R agree well in shape and levels
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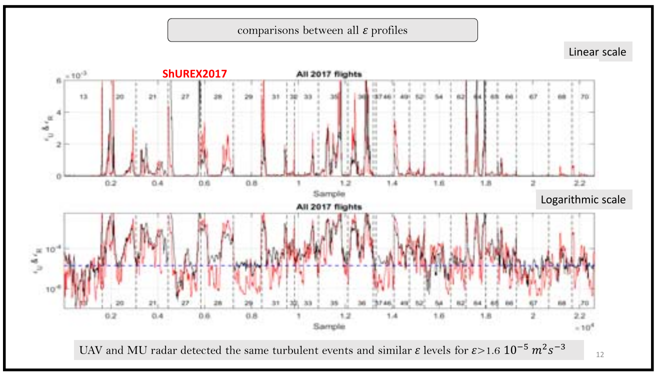
comparisons between all ε profiles
Linear scale
ShUREX2017
All 2017 flights
Sample
Logarithmic scale
UAV and MU radar detected the same turbulent events and similar ε levels for ε>1.6 10^{-5} m^2s^{-3}
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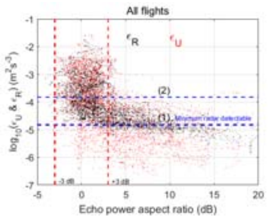
ε and aspect ratio (AR) of echo power



$AR = 10\,log10\left(\frac{P_V}{P_O}\right)$

	(1) 0.014 mW/kg	(2) 0.14 mW/kg
LR	$\varepsilon > 1.4\ 10^{-5}\ m^2s^{-3}$	$\varepsilon > 1.4\ 10^{-4}\ m^2s^{-3}$
AR < 3 dB	74 %	**95 %**
AR < 5 dB	86 %	**98 %**

ICAO: turbulence felt by medium-size aircraft > 1 mW/kg (light turbulence)

A large majority of ε_{UAV} values is associated with isotropic echoes (weak aspect ratios) => An a posteriori justification of the use of Doppler variance from the vertical beam.

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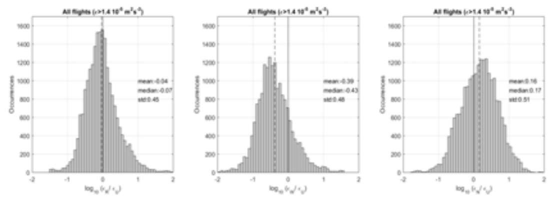
comparisons with ε_N and ε_W

Labitt/white formulation   Weinstock/Hocking formulation

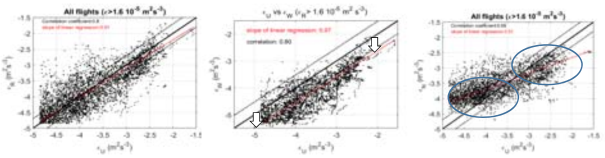


$\varepsilon_r \approx 0.47\sigma_t^3/L_{out}$ $\varepsilon_W \approx \left(\frac{4\pi}{1.6}\right)^{3/2}\frac{\sigma_t^3}{l^{3/2}}$ $\varepsilon_N \approx 0.47\sigma_t^2 N$

$<\varepsilon_r/\varepsilon_U> = 1$ (by definition) $<\varepsilon_W/\varepsilon_U> = 0.4$ $<\varepsilon_N/\varepsilon_U> = 1.45$

Comparisons between ε estimates

ε_R ε_W ε_N



Scatter plot of the various ε estimates and linear regression (with errors on x and y)

(1) Good agreement (regression slope ~0.9-1.0 and correlation ~0.8) between ε_U and ε_R $(or\ \varepsilon_W)$

(2) Disagreement between ε_{UAV} and ε_N: It is high for low values and becomes lower for high values.

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Conclusions of part I

a) The UAV and radar captured the same turbulent events with peaks of ε and σ_t at the same altitudes and times. => quantitative comparisons could be made.

b) $\varepsilon_{UAV} \sim \sigma^3/L_{out}$ with $L_{out} \approx 30\,m$. Energy dissipation rates can be estimated from the sole Doppler variance. at least in the lower troposphere

c) The asymptotic models ε_W and ε_N provide quite consistent levels:
$\varepsilon_W \sim 0.4\ \varepsilon_{UAV}$, in average, slight underestimation but no bias,
$\varepsilon_N \sim 1.45\ \varepsilon_{UAV}$, in average, *BUT* overestimates for low values and underestimates for high values (~>0.3 mW/kg)

d) In addition, since ε_R and ε_W are relevant for turbulence generated by convections or shear flow instabilities but ε_N is applicable to stratified turbulence only, the Weinstock model may not be suitable at least for tropospheric data.

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Part II: CWT – Pitot comparisons

Estimation of ε from C_T^2

For stratified turbulence (e.g. Ottersten, 1969; Gossard, 1982; Gavrilov et al. 2005):

$$\varepsilon_{CT2} = \left(\gamma\frac{g^2}{T^2}\frac{C_T^2}{N^2}\right)^{3/2}$$

Valid for dry or moist (unsaturated) air

$$\gamma = \frac{1}{\beta_\theta}\frac{P_r - Ri}{Ri} = \frac{1}{\beta_\theta}\frac{1-Rf}{Rf}$$

$\beta_\theta = 3.2$ (universal constant)

Ri: Richardson number, P_r= Turbulent Prandtl number
$Rf = {}^{Ri}/_{Pr}$: Flux Richardson number

mixing efficiency = $\frac{\text{Change in background potential energy due to mixing}}{\text{Energy expended}}$

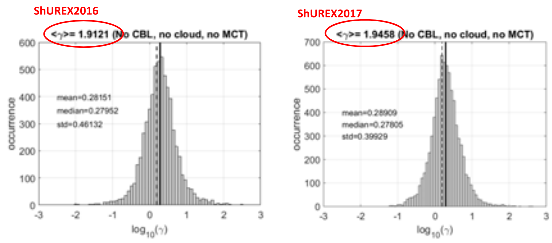
Mixing efficiency coefficient:
$\gamma_{R_f} \equiv R_f/(1-R_f)$

Gregg et al., 2018
Oceanic turbulence

1. It has long been understood that mixing efficiency is unlikely to be constant, but, because K_ρ from tracers and microstructure evaluated with $\gamma_{R_f} = \gamma_{\chi\epsilon} = 0.2$ agree better than do coefficients and efficiencies from simulations and experiments. Reasons for this agreement are not understood, particularly in view of the loss of po-

Statistical values of γ

$$\varepsilon = \left(\gamma \frac{g^2}{T^2}\frac{C_T^2}{N^2}\right)^{3/2}$$



Remarkable consistency between the results obtained from ShUREX2016 and ShUREX2017

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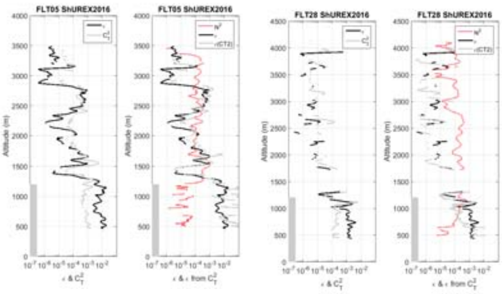
Examples of comparisons between ε_U and ε_{CT2}



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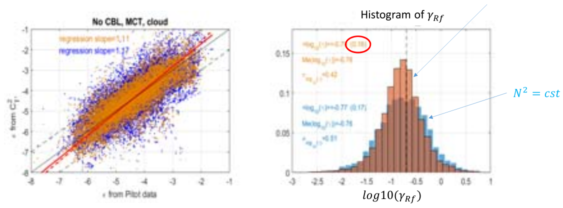
Statistical values of mixing efficiency coefficient γ_{Rf}



$$\gamma_{Rf} \approx 0.16 \cong 0.20$$

The mixing coefficient is found to be ~cst and similar to oceanic values

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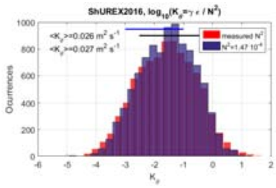
Eddy diffusivity $K_\theta = 0.16\frac{\varepsilon}{N^2}$

In the literature: $K_\theta \approx 0.33\frac{\varepsilon}{N^2}$
(e.g. Lilly, 1974;
Fukao et al., 1994,
Naström and Eaton, 2005)



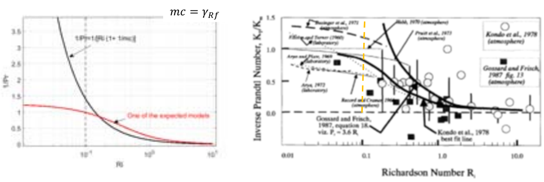
Black: After Alisse and Sidi (2000) for stratospheric turbulence (36 cases)
Blue: After Bertin et al. (1997), stratosphere (several cases)

The statistical values of K_θ are similar to prior estimates from HR balloon data at stratospheric heights

Conclusions of part II

1. TKE dissipation rates ε estimated from wind and temperature data compare well with a mixing efficiency coefficient γ_{Rf} of 0.16 (close to 0.2)
2. This result is not consistent with Prandtl number properties but…

$mc = \gamma_{Rf}$



For Ri >0.1, γ_{Rf}~0.16 provides reasonable agreements with laboratory experiments and theoretical expectations.
For Ri <0.1, large discrepancies. => **Is (effective) Ri <0.1 rare in lower tropospheric conditions (for stratified turbulence)?**

Perspectives

Since $\varepsilon = \left(\gamma \frac{g^2}{T^2}\frac{C_T^2}{N^2}\right)^{3/2} = \left(\gamma \frac{C_n^2}{M^2}N^2\right)^{3/2}$ (Ottersten, 1969)

C_n^2 can be estimated from ε making possible comparisons with C_n^2(radar)
⇒ Relevance of the radar estimates
⇒ Accurate calibration of C_n^2(radar)

(under processing)