

## Effects of increased $p\text{CO}_2$ and iron availability on phytoplankton assemblages in the Southern Ocean

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A risk of ocean acidification (i.e., an increase in seawater  $p\text{CO}_2$  and a concomitant decrease in pH) will be the most severe in the Southern Ocean (Orr et al. 2005), which is also known to be a high-nitrate, low-chlorophyll (HNLC) region. In addition, climate change might increase atmospheric iron deposition to some oceanic regimes and its availability for the phytoplankton assemblages (Woodward *et al.*, 2005). However, little is known about the impacts of ocean acidification, as well as iron enrichment on phytoplankton assemblage in this oceanic Southern Ocean. We therefore conducted simulated *in situ* incubation experiments during the austral summer of 2011/2012 on board the TR/V *Umitaka-Maru* to examine the effects of  $\text{CO}_2$  and iron availability on the community composition and photosynthetic physiology of phytoplankton. Seawater samples were collected from ~15 m at stations C02, C07, D07, and D13 (Fig. 1) with an acid-clean Teflon pump system. Prior to incubation,  $\text{FeCl}_3$  solutions were added into acid-clean incubation bottles in order to reduce iron limitation for phytoplankton growth. Ambient  $\text{CO}_2$  (Fe-added) and high- $\text{CO}_2$  (ca. 750  $\mu\text{atm}$ ; Fe-and- $\text{CO}_2$ -added) treatments (duplicate each) were constructed with or without the addition of  $\text{CO}_2$ -saturated seawater to the bottles. Duplicated non-iron-added (control) treatments were also prepared to assess the effects of iron enrichment on phytoplankton assemblages. Incubations were conducted for 3 or 4 days in a light- and temperature-controlled laboratory incubator (100  $\mu\text{mol photons m}^{-2} \text{s}^{-1}$  and *in situ* temperature). Initial phytoplankton community was mainly dominated by haptophytes at station C02, while diatoms were predominant at the other stations in terms of chlorophyll *a* (Chl-*a*) biomass (Fig. 2). At the end of incubation, concentrations of Chl-*a* were significantly higher in the Fe-added treatments relative to the controls at stations C02, C07, and D13 (*t*-test,  $p < 0.05$ ). In photosynthesis-irradiance experiments, remarkable increases in the maximum photosynthetic rate ( $P^{\text{B}}_{\text{max}}$ ) were observed in response to iron additions at these stations. However, effects of iron enrichment on Chl-*a* and  $P^{\text{B}}_{\text{max}}$  were not clear at station D07. A biomarker of haptophytes (19'-hexanoyloxyfucoxanthin) decreased in response to a rise in  $\text{CO}_2$  level at the end of incubation at stations C02 and D13. On the other hand, at the stations C07 and D07, a biomarker of diatoms (fucoxanthin) decreased in the high  $\text{CO}_2$  treatments relative to ambient  $\text{CO}_2$  treatments. At all stations,  $P^{\text{B}}_{\text{max}}$  values were little affected with  $\text{CO}_2$  availability. Our results suggest that the progression of ocean acidification and iron enrichment can alter the community composition in the Southern Ocean, although the response would differ among geographical locations due to the differences in environmental conditions.

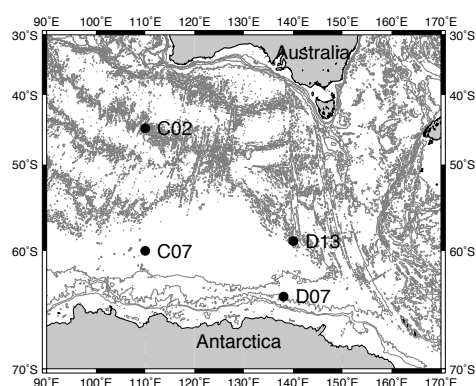


Fig. 1. Sampling sites of seawater for our incubation experiments.

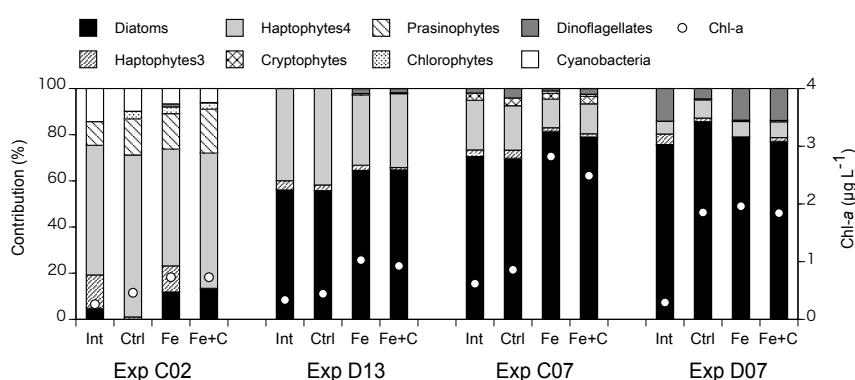


Fig. 2. Mean contributions of each phytoplankton class to total Chl-*a* biomass as estimated by CHEMTAX and mean Chl-*a* biomass. Int, Ctrl, Fe, and Fe+C indicate initial, control, Fe-added, and Fe-and- $\text{CO}_2$ -added treatments, respectively.

### References

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