The Yamato 980459 olivine-phyric shergottite consortium

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Abstract: Among 4100 meteorite samples found around the Yamato Mountain area during the field season of 1998–1999, Yamato (Y) 980459, weight of 82 g, is identified as a new shergottite. In ordered to improve our knowledge of the shergottite parent body, possibly Mars, a coordinated consortium study of the Y980459 shergottite was organized. Petrological, mineralogical, chemical and isotopic studies revealed that Y980459 is related to the olivine-phyric shergottites and the most primitive Martian basalt so far reported. Y980459 has solidified much faster than other shergottites, as indicated by its lack of plagioclase or maskelynite. Y980459 shows a light REE depleted pattern similar to other "depleted shergottites" Dhofar 019, Dal al Gani (DaG) 476, Sayh al Uhamir (SaU) 005 and Queen Alexandra Range 94201. The Sm-Nd isotopic data suggest that Y980459 crystallized at ~470 Ma and derived from a highly depleted source region. Cosmic-ray exposure age of Y980459 is calculated to be ~1 Ma based on ¹⁰Be concentration, which is similar to those of olivine-phyric shergottites DaG 476 and SaU 005. I here present an outline of the consortium and implications for Martian geological history on the basis of ongoing studies.

The first petrographical description of Y980459 by Ansgar Greshake, Jörg Fritz and Dieter Stöffler was not submitted to *Antarctic Meteorite Research*, but to *Geochimica et Cosmochimica Acta* on August 15, 2003, and was published in the May 15, 2004 issue. The other papers concerning Y980459 appeared in the *Antarctic Meteorite Research*, **17** issue are of equal scientific importance and priority.

key words: shergottites, Martian meteorites

1. Introduction

The JARE-39 (1998–1999) wintering over party conducted a meteorite search around the Yamato Mountains area (Kojima *et al.*, 2000). Among 4100 meteorites found, Yamato (Y) 980459 was identified as a Martian meteorite, which resembles the olivine-phyric shergottites (Kojima and Imae, 2002; Misawa, 2003). Shergottites are subclassified basaltic, lherzolitic and olivine-phyric (*e.g.*, Meyer, 2003; Goodrich, 2003) and show extensive shock features due to impacts. Their crystalline ages are relatively young, 0.17–0.58 Ga (see Nyquist *et al.*, 2001). Recently, a total of 8 meteorites, Elephant Moraine 79001 lithology A (EETA 79001A), Dhofar (Dho) 019, Dar al Gani (DaG) 476/489/670/735/876/975/1037, Sayh al Uhamir (SaU) 005/008/051/060/

090/094/120/150, Northwest Africa (NWA) 1068/1110/1183/1775, NWA 1195, NWA 2046 and Y980459, have been thus far identified as olivine-phyric shergottites (Meyer, 2003; Irving *et al.*, 2004). Most of the olivine-phyric shergottites were recently recovered from the hot deserts of Oman, Libya and Northwest Africa and some of them were heavily affected by terrestrial contamination. Along with EETA 79001A, Y980459 was found on the bare ice field in Antarctica, and thus enable us to study this type of rock, which is considered to be less affected by terrestrial contamination. The DaG 476 and SaU 005 olivine-phyric shergottites are very similar in mineralogy, chemical compositions and cosmic-ray exposure (CRE) ages of ~1.2 Ma. Dhofar 019 has a significantly older CRE age of ~20 Ma. NWA 1068 has crystallization age similar to basaltic shergottites in the range of ~165–200 Ma and CRE age of ~3 Ma. Systematic studies of olivine-phyric shergottites will provide us new insights into differentiation processes of the Martian mantle and add important new information about Martian volcanism.

After the announcement of this new Antarctic Martian meteorite, more than fifteen sample requests for investigation were received. At a meeting held on 22nd November 2002, the Committee on Antarctic Meteorite Research recognized the need for a coordinated consortium study of the Yamato olivine-phyric shergottite. The goals of this consortium are to identify similarities and differences between Y980459 and other olivine-phyric shergottites, and to constrain chemical and isotopic signatures of the shergottite source(s). The consortium will facilitate a balanced and effective study of the mineralogy, petrology, and geochemistry of this small sample (Misawa, 2003).

2. Yamato 980459 overview

Y980459 was collected on the bare ice field near the Minami-Yamato Nunataks (Fig. 1) on 4th December 1998. The meteorite is a single stone, weighing 82.46 gram, $5.0 \times 4.2 \times 2.6$ cm in dimension, and is partially covered with shiny black fusion crust with ablation features (Fig. 2). The surface of the meteorite has been physically eroded and the fusion crust in large part removed (Fig. 2a, c). However, the interior of the rock seems to be fresh (Fig. 2c, d). Olivine is easily identified with the naked eye because of its yellow/orange color. Y980459 is composed of olivine phenocrysts, up to 2 mm in size, pyroxene and abundant interstitial glass (Fig. 3).

3. Curatorial processing of Yamato 980459

Our ongoing plan for the Y980459 consortium is summarized in Table 1. A 90 mg-sized sample (Y980459,70) was allocated to R.N. Clayton (Univ. Chicago) for measurements of oxygen isotopes. Even before the consortium commenced, Y980459,60 (184 mg) was allocated to K. Nagao (Univ. Tokyo) for noble gas study that confirmed Y980459 is a Martian meteorite. In addition to this sample, he received a 248-mg sample (Y980459,66) for trapped noble gas geochemistry. A homogenized powder sample (Y980459,80), weighing 2.585 g, was prepared and an aliquot of 1.4 g (Y980459,81) was used for wet chemical analysis at NIPR (Analyst: H. Haramura). Aliquots of this homogenized sample are available for bulk analysis upon request.

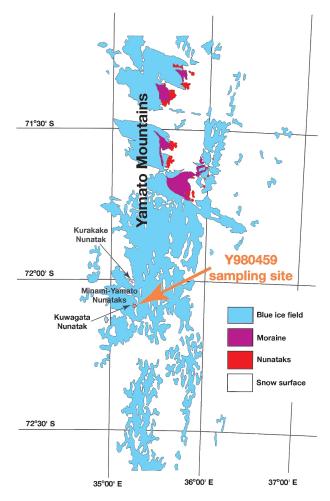


Fig. 1. The sampling site of the Yamato olivine-phyric shergottite. Yamato 980459 was collected at 72°5.881'S, 35°14.559'E, ~2 km south of the Kuwagata-yama Nunatak.

A mass of 308 mg (Y980459,64) was allocated to M. Grady for carbon, nitrogen and oxygen isotope studies. T. Mikouchi (Univ. Tokyo) received a 413-mg sample (Y980459,57) for transmitted electron microscopy (TEM) and for crystallization experiments. D. Stöffler (Humboldt Univ.) received Y980459,75 (231 mg) to study shock effects on the Yamato shergottite by TEM and X-ray diffraction analyses. M. Miyamoto (Univ. Tokyo) received a 108-mg (Y980459,69) sample for reflectance spectroscopy. A fragment weighing 523-mg (Y980459,55) for the Rb-Sr isotope systematics and a 120-mg powdered sample (Y980459,84) for lithophile trace element studies were allocated to N. Nakamura (Kobe Univ.). G. Dreibus (MPI, Mainz) received a 512-mg sample (Y980459,63) for the determination of Rb-Sr, Sm-Nd and U-Pb isotope systematics and a 252-mg powdered sample (Y980459,83) for the determination of trace element abundances with instrumental neutron activation analyses

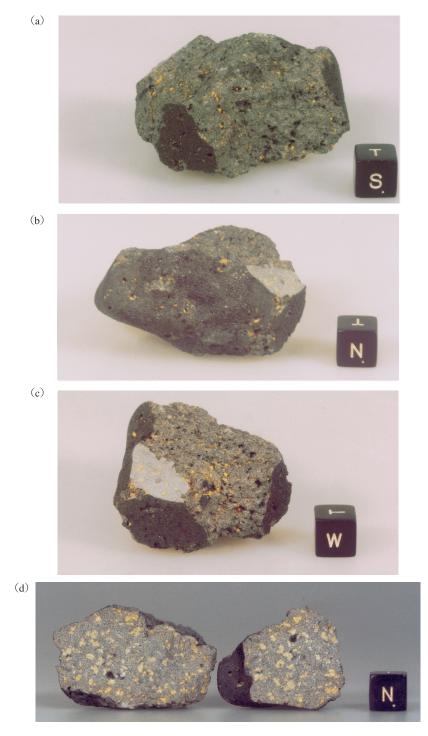


Fig. 2. Macroscopic features (a-c) and cross section (d) of Yamato 980459. The cube is 1 cm³.

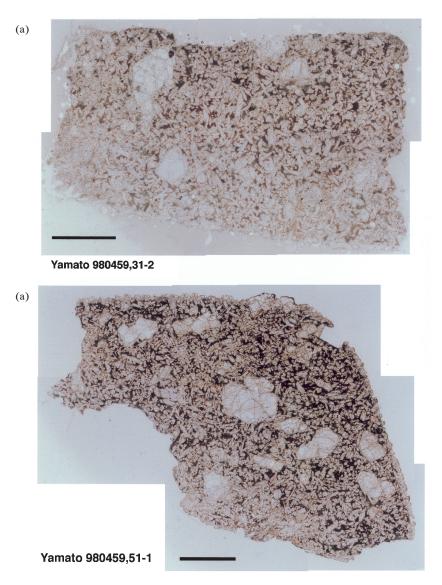


Fig. 3. Photomicrographs of (a) Yamato 980459,31-2 and (b) Yamato 980459,51-1. The meteorite is composed of olivine phenocrysts, up to 2 mm in size, pyroxene and glass with minor amounts of oxide minerals. Scale bar=2 mm.

(INAA) and radiochemical neutron activation analyses (RNAA). L. Nyquist (NASA-JSC) received fragments weighing a total of 1.487 g. Together with D. Bogard (NASA-JSC) and C.-Y. Shih (Lockheed-Martin), he will determine crystallization age by the Rb-Sr, ¹⁴⁷Sm-¹⁴³Nd and ³⁹Ar-⁴⁰Ar systems and mantle differention age by the ¹⁴⁶Sm-¹⁴²Nd system. G. McKay (NASA-JSC) received Y980459,54 (783 mg) for Re-Os and ¹⁴⁶Sm-¹⁴²Nd isotope systematics and for investigating shergottite petrogenesis. K. Marti (UCSD) received Y980459,62 (447 mg) for identification of fission, spallation

and indigenous components of nitrogen, argon, and xenon in the source region of shergottites. O. Eugster (Univ. Bern) received Y980459,59 (211 mg) for CRE age, parent body ejection time and terrestrial age determinations. In order to study in detail the cosmic-ray effects and pre-atmospheric sizes of Y980459 using ⁴¹Ca, ³⁶Cl, ²⁶Al, ¹⁰Be and ⁵³Mn, K. Nishiizumi (Space Sci. Lab., UCB) was allocated a 675 mg-sized sample (Y980459,56) containing fusion crust and a 457 mg-sized interior sample (Y980459,66). He will split the interior sample and send aliquots of chips to A. Jull (NSF Arizona AMS Lab., Arizona Univ.) for ¹⁴C measurement. V. Hoffman (Univ. Tübingen) received Y980459,65 (506 mg) for studying the magnetic properties of shergottites. Most samples were sent to investigators in February 2003.

Apart from those used for the initial classification (Y980459,31-1, 31-2, and 51-1), twelve polished thin sections (PTSs) were produced for the consortium, from three separate parent chips, Y980459,31, Y980459,41 and Y980459,51 (see Table 1). Thus far, PTSs have been loaned to Y. Ikeda (Ibaraki Univ.), T. Mikouchi (Univ. Tokyo), G. Dreibus (MPI, Mainz), H. McSween (Univ. Tennessee) and G. McKay (NASA-JSC) for mineralogical and petrological studies. Two PTSs, Y980459,41-2 and Y980459,51-2, loaned to G. McKay and to H. McSween, will be used for ion microprobe analysis. M.

Subnumber	Weight	Investigator	Institution	Type of investigation	
	(gram)	8			
31	1.115		NIPR	PTSs	
40	1.487	Nyquist, L.	NASA-JSC	chronology	
41	1.425		NIPR	PTSs	
51	0.950		NIPR	PTSs	
54	0.783	McKay, G	NASA-JSC	petrology & chronology	
55	0.523	Nakamura, N.	Kobe Univ.	chronology	
56	0.675	Nishiizumi, K.	Space Sci. Lab., UCB	CRE	
57	0.413	Mikouchi, T.	Univ. Tokyo	mineralogy & petrology	
58	0.457	Nishiizumi, K.	Space Sci. Lab., UCB	CRE	
59	0.211	Eugster, O.	Univ. Bern	CRE	
60	0.184	Nagao, K.	Univ. Tokyo	noble gases	
61	0.231	Stöffler, D.	Inst. Mineralogie, Berlin	shock features	
62	0.447	Marti, K.	UCSD	nitrogen & noble gas isotopes	
63	0.512	Dreibus, G	MPI, Mainz	chronology	
64	0.308	Grady, M.	NHM, London	stable isotopes	
65	0.506	Hoffman, V.	Univ. Tübingen	magnetic property	
66	0.248	Nagao, K.	Univ. Tokyo	noble gases	
69	0.108	Miyamoto, M.	Univ. Tokyo	reflectance spectroscopy	
70	0.092	Clayton, R.	Univ. Chicago	oxygen isotopes	
80	2.585		NIPR	powdered sample	
81	1.400		NIPR	bulk wet chemistry (from ,80)	
82	0.248	Ebihara, M.	Tokyo Metropolitan Univ.		
83	0.252	Dreibus, G.	MPI, Mainz	chemistry (from ,80)	
84	0.120	Nakamura, N.	Kobe Univ.	chemistry (from ,80)	

Table 1. Distribution of samples for Yamato 980459 Consortium, including PTS samples.

PTSs; Y980459,31-3: Dreibus, G (MPI, Mainz), Y980459,41-2: McKay, G (NASA-JSC), Y980459,41-3: Mikouchi, T. (Univ. Tokyo), Y980459,41-5: Ikeda, Y. (Ibaraki Univ.), Y980459,51-2: McSween, H. (Univ. Tennessee), Y980459,41-2: Wadhwa, M. (FMNH).

Wadhwa (FMNH) will study trace and minor element microdistributions of Y980459. In order to estimate the water content of Martian mantle, R. Lenz (Univ. Tennessee) will analyze Li, Be and B contents in pyroxene.

4. Results and discussion

4.1. Mineralogy and petrology

Petrographic features of Y980459 are similar to those of olivine-phyric shergottites EETA 79001A, DaG 476, Dho 019, SaU 005, NWA 1068, NWA 1195 and NWA 2046, although NWA 1195 and NWA 2046 also contain orthopyroxene phenocrysts. Mesostasis of Y980459 consists of dendritic olivine, pyroxene, chromite, ilmenite and sulfide but no phosphate (Mikouchi et al., 2003, 2004). A noteworthy feature observed in Y980459 is the absence of plagioclase or maskelynite. According to the classical definition, shergottites consist of a somewhat similar Ca-rich pyroxene to that in the eucrites, associated, however, with lath of isotropic maskelynite instead of anorthite (Prior, 1920). The absence of plagioclase is considered to be due to rapid cooling which suppressed nucleation of plagioclase (Greshake et al., 2003, 2004; Ikeda, 2003, 2004; McKay and Mikouchi, 2003; Koizumi et al., 2004; McKay et al., 2004; Mikouchi et al., 2004). Although mineral assemblage of Y980459 is different from those of shergottites, Y980459 could be classified as a member of the olivine-phyric shergottite subgroup. As described below, bulk chemical composition of Y980459 supports this. Olivine cores in Y980459 are the most magnesian (Fo₈₆) among shergottites and are almost same as olivine phenocrysts in NWA2046 (Irving et al., 2004). Thus, Y980459 is the most primitive olivine-phyric shergottite so far identified.

Shock features are observed in olivine and pyroxene. Shock melt pockets are heterogeneously distributed but they are rare (Ikeda, 2004; Mikouchi *et al.*, 2004). Although the lack of plagioclase prevents precise shock pressure estimation, the shock degree of Y980459 is estimated to be 20–25 GPa. (Ikeda, 2004; Greshake *et al.*, 2004; Mikouchi *et al.*, 2004). Shock pressures of shergottites range from 30–50 GPa (see Nyquist *et al.*, 2001). Thus, Y980459 is the least shocked shergottite so far reported.

4.2. Oxygen isotopic signature

Oxygen isotopic compositions of Y980459,70 ($\delta^{18}O = +4.31\%$ relative to SMOW, $\delta^{17}O = +2.52\%$, $\Delta^{17}O = +0.28\%$) are in agreement with a Martian origin of Y980459 (R.N. Clayton and T.K. Mayeda, written communication, 2001).

4.3. Major and trace element composition

Bulk chemical composition of Y980459 is given in Table 2 together with those of other olivine-phyric shergottites. Bulk mg (=100×molar Mg/[Mg+Fe]) values of Y980459, DaG 476 and SaU 005 are ~68, which is higher than those of EETA 79001A, Dho 019, NWA 1068 (mg=58-61).

The light REE (LREE) depleted pattern of Y980459 (Dreibus *et al.*, 2003; Nakamura *et al.*, 2003; Shirai and Ebihara, 2003) is similar to those of olivine-phyric shergottites EETA 79001A (Burghele *et al.*, 1983), Dho 019 (Neal *et al.*, 2001; Taylor *et al.*, 2002), DaG 476 (Ziepfel *et al.*, 2000; Barrat *et al.*, 2001a), SaU 005 (Dreibus *et al.*, 2001a), SaU 005 (Dreibus *et al.*, 2003), SaU 005 (Dreibus *et al.*, 2003; Shirai and SaU 005), SaU 005 (Dreibus *et al.*, 2003), SaU 005 (Dreibus *et al.*, 2003), SaU 005 (Dreibus *et al.*, 2003), SaU 005 (Dreibus *et al.*, 2004), SaU 005 (Dreibus *et al.*, 2005), SaU 005 (Dreibus *et al.*, 2005)

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	Y980459*	DaG476 [1, 2]	Dho019 [3]	SaU005 [4]	NWA1068 [5]	EETA 79001A [6]
SiO ₂	48.70	45.76		47.20		48.58
TiO ₂	0.54	0.39	0.49	0.42	0.77	0.64
Al_2O_3	5.27	4.37	6.65	4.53	5.75	5.37
Cr_2O_3	0.71	0.78	0.50	0.78	0.63	0.59
FeO	17.32	16.06	19.9	18.34	20.48	18.32
Fe_2O_3	0**					
MnO	0.52	0.45	0.48	0.46	0.46	0.47
MgO	19.64	19.41	14.6	20.49	16.50	16.31
CaO	6.37	<u>7.66</u>	<u>9.42</u>	5.74	7.91	7.05
Na_2O	0.48	0.51	0.89	0.60	1.14	0.82
K ₂ O	<0.02	<u>0.038</u>		0.022	0.16	0.03
H ₂ O(-)	0					
H ₂ O(+)	0					
P_2O_5	0.29	0.32	0.4	0.31		0.54
FeS	0.26					
Fe	0					
Ni	0.027	0.0230	0.0065	0.0310		0.015
Co	0.007	0.0051	0.0045	0.0055		0.004

Table 2. Bulk chemical compositions of olivine-phyric shergottites (wt%).

* An aliquot of 1.4 gram from the powdered sample (Y980459,81) was used for wet chemical analysis. Analyst: H. Haramura. ** Ferric iron was measured but not detected. *Italics*: total iron calculated as FeO. <u>Underline</u>: elemental abundances are significantly affected by terrestrial weathering.

Refs. [1] Barrat et al. (2001a), [2] Zipfel et al. (2000), [3] Neal et al. (2001), [4] Dreibus et al. (2000), [5] Barrat et al. (2001b), [6] Burghele et al. (1983).

al., 2000) and basaltic shergottite QUE 94201 (Dreibus *et al.*, 1996; Warren and Kallemeyn, 1997; Kring *et al.*, 2003). All olivine-phyric shergottites do not necessarily show the LREE depleted pattern: NWA 1068 shows a "flat" REE pattern (Barrat *et al.*, 2001b) similar to those of "enriched" shergottites Shergotty, Zagami, NWA 856 and Los Angeles. These trace element features observed in olivine-phyric shergottites could reflect differences among their source materials. Depleted shergottites (EETA 79001A, Dho 019, DaG 476, SaU 005, QUE 94201 and Y980459) could have derived from primitive, LREE-depleted, reduced source regions and could have been less affected by crustal contamination/assimilation (Shih *et al.*, 1982; Wadhwa, 2001; Herd *et al.*, 1997).

Terrestrial contamination affected trace elements, especially halogen, alkali and alkaline earth elements in Y980459 (Dreibus *et al.*, 2003; Shih *et al.*, 2003, 2004a). Even though the collected sample-size of QUE 94201 was about one seventh of Y980459, alkaline and alkaline earth elements in QUE 94201 were less affected compared with the case of Y980459. These alteration features observed in Antarctic shergottites imply complicated alteration processes during residence in Antarctica.

4.4. Noble gases and cosmic-ray exposure age

Isotopic ratios and concentrations of noble gases were determined for Y980459 (Nagao and Okazaki, 2003; Okazaki and Nagao, 2004). Noble gas abundances in

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Y980459 are similar to those of other shergottites. Contribution of Martian atmospheric components, ⁴⁰Ar and excess of ¹²⁹Xe, are small. Cosmic-ray exposure ages of Y980459 calculated based on cosmogenic ²¹Ne and ³⁸Ar are ~2.5 Ma (Okazaki and Nagao, 2004; Christen *et al.*, 2004), which is different from those of olivine-phyric shergottites Dho 019 (20 Ma, Shukolyukov *et al.*, 2002), EETA 79001A (0.5 Ma, Bogard *et al.*, 1984), DaG 476 (1.2 Ma, Zipfel *et al.*, 2000), and SaU 005 (1.2 Ma, Pätsch *et al.*, 2000) but rather similar to CRE ages of basaltic shergottites QUE 94201 (2.5 Ma, Eugster *et al.*, 1997), Shergotty (2.9 Ma, Eugster *et al.*, 2002), Zagami (2.85 Ma, Eugster *et al.*, 2002). The CRE age of Y980459 based on ¹⁰Be concentration is 1.1 ± 0.2 Ma (Nishiizumi and Hillegonds, 2004), similar to that of olivine-phyric shergottites DaG 476 and SaU 005 but shorter than the CRE ages (~2.5 Ma) based on cosmogenic ²¹Ne and ³⁸Ar, indicating a significant pre-exposure on the Martian surface (Nishiizumi and Hillegonds, 2004).

4.5. Rubidium-strontium and samarium-neodymium systematics

Shih *et al.* (2003, 2004a, b) studied the Rb-Sr and Sm-Nd isotope systematics of Y980459. It was hard to obtain fractions having wide ranges of Rb/Sr and Sm/Nd due to the lack of plagioclase or maskelynite in Y980459. The Rb-Sr system of Y980459 was "open" and has been severely disturbed by terrestrial weathering even though the sample apparently seems to be fresh. Six samples define a Sm-Nd isochron age of 472 ± 47 Ma with a high initial ε Nd value of $\pm 36.9\pm 2.2$ (Shih *et al.*, 2004b), identical to values reported for DaG 476 (Borg *et al.*, 2003). An initial ⁸⁷Sr/⁸⁶Sr ratio of 0.701384 \pm 0.000011 at 472 Ma is estimated from the weighted average of nine samples (Shih *et al.*, 2004b). The Rb-Sr and Sm-Nd isotopic systematics of Y980459, like those of DaG 476 and QUE 94201 (Borg *et al.*, 1997, 2003), are consistent with derivation from a source region that was strongly depleted in incompatible elements.

5. Summary

Although plagioclase or maskelynite is absent, Y980459 resembles olivine-phyric shergottites in its petrological features and in bulk chemistry. Y980459 is the most primitive Martian basalt so far identified and actually reflects its source region. Y980459 crystallized at 472 ± 47 Ma and derived from a low ${}^{87}\text{Sr}/{}^{86}\text{Sr}$ and a high ϵ Nd source region closely related to the source of DaG 476. The ${}^{10}\text{Be}$ exposure age of Y980459 is 1.1 ± 0.2 Ma, which is similar to those of olivine-phyric shergottites DaG 476 and SaU 005, suggesting a single launching of Y980459 with other olivine-phyric shergottites.

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reproduced from the scanned image taken by T. Yada. The author thanks N. Nakamura and G.A. McKay for their helpful comments on the manuscript.

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