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A laboratory experiment on the influence of aqueous alteration on noble gas compositions in the Allende meteorite

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Abstract: We have carried out a hydrothermal experiment on the Allende CV chondrite, in order to understand the observed noble gas loss from CM and CI chondrites, possibly caused by aqueous alteration on their parent bodies (L. B. Browning et al., Geochim. Cosmochim. Acta, 60, 2621, 1996; T. Nakamura et al., Geochim. Cosmochim. Acta, 63, 257, 1999). The Allende meteorite was disaggregated and kept with liquid water in a pressurized vessel under 15 atm of water vapor at 200°C for a week. The abundances of helium, neon, argon, krypton and xenon in the samples have been measured by stepwise heating technique. The abundances of noble gases were decreased by 24% (He), 27% (Ne), 38% (Ar), 23% (Kr) and 34% (Xe) after the hydrothermal treatment, suggesting that the hot water treatment had contributed to reduce the concentrations of noble gases in Allende. This gas loss also affected the isotopic composition of xenon. For the 1200°C fraction, the amount of ¹³²Xe-HL was decreased from the sample by about 80% during the hot water treatment, while that of 132 Xe-Q was decreased to a lesser extent (\sim 40%). This result seems to indicate that the carrier of Xe-HL is less resistant to the gas loss by the hot water alteration than phase Q for high temperature components. However, further studies are required as this trend becomes less obvious and rather similar if we take into consideration the xenon of 800°C fraction.

1. Introduction

Most carbonaceous chondrites have experienced aqueous alteration, resulting in changes in their texture, mineralogy and possibly chemical composition (Cohen et al., 1983; Tomeoka and Buseck, 1982, 1985; Zolensky, 1984; Zolensky et al., 1987). This aqueous alteration is considered to have occurred in the parent body of meteorite (Kerridge et al., 1979; McSween, 1979; Richardson, 1978; Armstrong et al., 1982; Tomeoka and Buseck, 1985). Water, presumably in the form of ice would have accreted along with lithic materials to form the meteorite parent bodies (Prinn and Fegley, 1987). The liquid water, which caused aqueous alteration, was derived from ice by heat generated within the parent body by the decay of short-lived radionuclides (DuFresne and Anders, 1962).

Recently, Nakamura et al. (1999) suggested that the aqueous alteration might

have played an important role in decreasing the abundances of noble gases in CM chondrite. They observed that the degrees of aqueous alterations in the CM chondrites were inversely correlated with the amounts of primordial noble gases trapped in phase Q (main noble gas carrier of carbonaceous material in chondrites; Lewis et al., 1975). Browning et al. (1996) also reported similar observation in eight CM chondrites. As the elemental abundances and isotopic compositions of noble gases in chondrites are very important tracers of the processes in the early solar system, the effects of such aqueous alterations on noble gas compositions should be clarified.

There have been a few experimental works investigating the artificial aqueous alteration effect on the mineralogy and the texture of the meteorite (Imae and Shinoda, 1998; Duke and Brearley, 1998). However, how the noble gas compositions are affected by the aqueous alteration is not known. It is also of great interest how the HL component (another primordial component carried in presolar diamonds; Lewis *et al.*, 1987) will be affected by such hydrothermal processes. Therefore, we have carried out the experiments in which we artificially altered the matrix of the Allende meteorite with hot water and examined its effects on noble gas compositions. We chose Allende because it has experienced a minimal aqueous alteration after its formation, and thus should be more sensitive to the artificial alterations than CM and CI that have been altered by natural processes.

2. Sample preparation and experimental procedures

We prepared two samples to examine possible change in their noble gas compositions before and after the artificial hydrothermal alteration. The starting material was the Allende matrix (AC-1) prepared by the freeze-thaw disaggregation. The great majority of AC-1 consists of matrix minerals, but some fragments of condrules and CAIs are also included. The AC-1 is the residual material after we partially removed the floating black fine material that was very similar to the chemical residue enriched in noble gases (Matsuda et al., 1999). Thus, it is likely that noble gas abundances in AC-1 are low compared to those in the bulk Allende. Parts of AC-1 were loaded into a 10 cm³ teflon vessel together with about 5 cm³ of pure water. This teflon vessel was then put inside the stainless-steel vessel. We heated up this sample container at 200°C for about one week. The rock to water ratio was about 1/50 by weight, where water would provide sufficient vapor to alter the sample under the pressure of about 15 atm. The altered materials (HM-1) were recovered from the vessel.

Both AC-1 and HM-1 were analyzed for the elemental abundances and the isotopic compositions of noble gases by stepwise heating gas extraction at 800, 1000, 1200, 1400 and 1600°C. The extracted gas was purified by two Ti-Zr getters heated at 700°C to remove any active species. After purification, the extracted noble gases were separated into four fractions (He+Ne, Ar, Kr and Xe) using two cold traps, and then each fraction was successively admitted to the sector-type mass spectrometer (VG5400, Osaka University) fitted with two collectors: Faraday for ⁴He and

⁴⁰Ar analyses and an electron multiplier with pulse counting system for other isotopes. Its sensitivities and mass discrimination factors for noble gases were calibrated by analyzing known amounts of atmospheric noble gases. Uncertainties for noble gas sensitivities are about 5% including all the errors of the volumes of line, standard air, etc. The blank correction was applied by subtracting the pre-sample blank levels measured at 800 and 1600°C (the hot blank at 800°C was corrected for the sample gas measured at 800°C, and that at 1600°C was applied for all other temperature steps). Typical hot blanks at 1600°C were as follows: ²²Ne= 2.7 to 5.0×10^{-13} , ${}^{36}\text{Ar} = 3.3$ to 12×10^{-12} , ${}^{84}\text{Kr} = 2.4$ to 11×10^{-13} and ${}^{132}\text{Xe} = 3.6$ to 9.0×10^{-14} cm³ STP. A hot blank for ⁴He was negligible by the Faraday detector, and we did not make the correction for He. We calculated the isotope ratios by assigning 100% error for the blank levels. In the case of neon isotopes, an isobaric interference on m/e = 22 from CO_2^{++} was subtracted to obtain the ²²Ne abundance. Our mass spectrometer can partially resolve the peak of ⁴⁰Ar⁺⁺ from the ²⁰Ne, so the conventional correction by using a production ratio of ${}^{40}Ar^{++}/{}^{40}Ar^{+}$ was not necessary.

3. Results and discussion

3.1. Elemental abundance

Tables 1 and 2 present the results of noble gas analyses by stepwise heating from AC-1 and HM-1, respectively. Figure 1a to d show the stepped release profiles of ²²Ne, ³⁶Ar, ⁸⁴Kr and ¹³²Xe, respectively. As shown in Fig. 1, the untreated sample released large portions of noble gases at 1200°C. Also in the treated sample, we found release peaks of ³⁶Ar, ⁸⁴Kr and ¹³²Xe at 1200°C. However, amounts of the released noble gases from the treated sample at 1200°C became smaller than those of the untreated sample. Total amounts of He, Ne, Ar, Kr and Xe decreased by 24%, 27%, 38%, 23% and 34%, respectively, after the hydrothermal treatment. These observations suggest that the artificial alteration is responsible for the loss of noble gases from the chondritic materials. This result confirms the observation of generally lower noble gas amounts in the heavily altered CM chondrites that might have lost their noble gases by aqueous alterations on their parent body (Browning *et al.*, 1996; Nakamura *et al.*, 1999).

3.2. Isotopic compositions

Hereafter, we will discuss the observed apparent gas loss from the samples by comparing the xenon isotopic ratios before and after the hot water treatment. In the present samples, we have recognized three distinct xenon isotope components: Air-Xe, Xe-Q and Xe-HL. A ¹³¹Xe/¹³²Xe vs. ¹³⁰Xe/¹³²Xe diagram (Fig. 2) is useful to detect the contribution of Air-Xe because Air-Xe, Xe-Q and Xe-HL do not fall on a single line. As shown in Fig. 2 where the data points of Xe released at 800°C are close to Air-Xe, absorbed air was released from the samples at the lowest temperature. Xenon isotopic ratios of higher temperature steps are clearly non-atmospheric and seem to be distributed around Xe-Q in Fig. 2. However, when

²²Ne ³He/ He ³⁶Ar ⁴He 84 Kr ²⁰Ne/²²Ne ²¹Ne/²²Ne $^{38}Ar/^{36}Ar$ 40 Ar/ 36 Ar 78 Kr/ 84 Kr 80 Kr/ 84 Kr 82Kr/84Kr 83Kr/84Kr 86Kr/84Kr 10⁻¹⁰cm³STP/g 10⁻⁶cm³STP/g 10-4 10 9 cm³STP/g 10⁻⁸cm³STP/g 20 0.2045 0.00422 0.1991 0.2043 0.3156 29.7 8.7 2.85 0.684 1.0 84.4 4.6 0.0399 ±0.4 ±0.29 ±0.070 ±0.0009 ±2.6 ±0.00061 ±0.0006 ±0.0010 ±0.0024 ±0.0041 5.7 25.4 5.2 1.97 0.0064 0.0437 0.1976 0.2023 0.806 3.2 0.2063 81.8 0.3101 3.0 ±0.1 ± 0.21 ±0.015 ±0.0009 ±2.9 ±0.0025 ±0.0015 ±0.0017 ±0.0022 ±0.0029 0.9 8.6 5.6 2.19 0.773 7.4 0.2033 11.2 7.0 0.00617 0.0429 0.19830.2047 0.3080 ±0.00037 ±0.4 ±0.23 ±0.079 ±0.0008 ±0.8 ±0.0007 ±0.0017 ±0.0017 ±0.0027 N.D. 0.831 0.0512 0.197 0.208 2.3 1.91 1.4 0.196327.0 1.9 N.D. 0.310 ±0.22 ±0.085 ±0.0009 ±4.3 ±0.0024 ±0.014 ±0.011 ±0.010 0.055 N.D. (0.2)7.4 0.089 0.1 0.1970 182 N.D. N.D. 0.220 0.341 (0.2)

±0.0050

0.2031

±0.0005

±17

23.0

±1.0

17

±0.024

±0.0007

0.0434

0.00560

±0.00051

±0.030

0.2049

±0.0017

0.1982

±0.0023

±0.049

0.3114

±0.0023

Table 1. Isotopic ratios and concentrations of AC-1 (18.01 mg).

132Xe 10 ⁻¹⁰ cm ³ STP/g	¹²⁴ Xe/ ¹³² Xe	¹²⁶ Xe/ ¹³² Xe	¹²⁸ Xe/ ¹³² Xe	¹²⁹ Xe/ ¹³² Xe	e ¹³⁰ Xe/ ¹³² Xe	¹³¹ Xe/ ¹³² Xe	e ¹³⁴ Xe/ ¹³² Xe	e ¹³⁶ Xe/ ¹³² Xe
3.0	0.00452	0.00403	0.07619	1.1039	0.1506	0.7994	0.4140	0.3697
	±0.00065	±0.00048	±0.00087	±0.0045	±0.0007	±0.0073	±0.0010	±0.0018
2.3	0.00523	0.00494	0.08376	1.3081	0.1634	0.8150	0.3895	0.3306
	±0.00045	±0.00052	±0.00097	±0.0069	±0.0012	±0.0083	±0.0035	±0.0029
6.1	0.00467	0.00381	0.08420	1.5448	0.1629	0.8174	0.3872	0.3261
	±0.00032	±0.00034	±0.00080	±0.0068	±0.0009	±0.0077	±0.0017	±0.0015
1.6	0.00403	0.00409	0.0839	2.041	0.161	0.8250	0.3930	0.3354
	±0.00055	±0.00080	±0.0014	±0.015	±0.013	±0.0081	±0.0050	±0.0040
0.1	N.D.	N.D.	0.053	2.55	0.095	0.820	0.47	0.400
			±0.016	±0.38	±0.012	±0.040	±0.11	±0.090
13	0.00466	0.00410	0.08207	1.4691	0.1596	0.8138	0.3949	0.3384
	±0.00024	±0.00024	±0.00051	±0.0078	±0.0017	±0.0044	±0.0014	±0.0013
	3.0 2.3 6.1 1.6 0.1	3.0 0.00452 ±0.00065 2.3 0.00523 ±0.00045 6.1 0.00467 ±0.00032 1.6 0.00403 ±0.00055 0.1 N.D.	3.0 0.00452 0.00403 ±0.00065 ±0.00048 2.3 0.00523 0.00494 ±0.00045 ±0.00052 6.1 0.00467 0.00381 ±0.00032 ±0.00034 1.6 0.00403 0.00409 ±0.00055 ±0.00080 0.1 N.D. N.D.	3.0 0.00452 0.00403 0.07619 ±0.00065 ±0.00048 ±0.00087 2.3 0.00523 0.00494 0.08376 ±0.00045 ±0.00052 ±0.00097 6.1 0.00467 0.00381 0.08420 ±0.00032 ±0.00034 ±0.00080 1.6 0.00403 0.00409 0.0839 ±0.00055 ±0.00080 ±0.0014 0.1 N.D. N.D. 0.053 ±0.016 13 0.00466 0.00410 0.08207	3.0	3.0	3.0	3.0

±0.032

0.747

±0.035

12

±1.4

2.41

±0.14

22

The datum with parenthesis means >20% contributions from the procedural blank

N.D. = not determined due to small amount of gas released.

Temp

°C

1000

1200

1400

1600

Total

27

28.1

±1.5

±0.0055

±0.0086

±0.0052

±0.0018

±0.013

Temp °C	⁴ He 10 ⁻⁶ cm ³ STP/g	³ He/ ⁴ He	²² Ne 10 ⁻⁹ cm ³ STP/g	²⁰ Ne/ ²² Ne	²¹ Ne/ ²² Ne	³⁶ Ar 10 ⁸ cm ³ STP/g	³⁸ Ar/ ³⁶ Ar	⁴⁰ Ar/ ³⁶ Ar	⁸⁴ Kr 10 ¹⁰ cm ³ STP/g	⁷⁸ Kr/ ⁸⁴ Kr	80Kr/84Kr	⁸² Kr/ ⁸⁴ Kr	⁸³ Kr/ ⁸⁴ Kr	⁸⁶ Kr/ ⁸⁴ Kr
800	14	26.7	6.4	3.21	0.657	1.1	0.1974	167.2	6.0	0.00589	0.03852	0.1988	0.2016	0.3071
		±0.7		±0.33	±0.067		±0.0009	±3.8		±0.00018	±0.00049	±0.0011	±0.0012	±0.0017
1000	4.9	20.6	3.8	3.08	0.680	1.7	0.1993	(*)8.6	(2.0)	0.00539	0.0377	0.1955	0.1964	0.3067
		±0.1		±0.35	±0.014		±0.0010			±0.00078	±0.0020	±0.0020	±0.0021	±0.0033
1200	0.7	9.5	3.1	2.18	0.787	3.4	0.1986	2.4	4.3	0.0063	0.04224	0.2014	0.2023	0.3113
		±0.5		±0.29	±0.081		±0.0009	±3.8		±0.0059	±0.00064	±0.0046	±0.0048	±0.0049
1400	0.2	N.D.	2.2	1.81	0.822	1.2	0.1946	10	(1.4)	0.00642	0.0489	0.2080	0.2026	0.3111
				±0.32	±0.084		±0.0011	±22		±0.00085	±0.0032	±0.0070	±0.0027	±0.0041
1600	N.D.		0.1	3.7	0.61	0.1	0.199	48	(0.2)	0.008	0.06	0.23	0.22	0.30
				±9.7	±0.11		±0.011	±112		±0.016	±0.12	±0.31	±0.20	±0.12
Total	20	20.8	16	2.78	0.711	7.5	0.1979	36.1	13	0.0061	0.0416	0.201	0.2017	0.3086

±0.0054

±4.6

Table 2. Isotopic ratios and concentrations of HM-1 (20.12 mg).

	132Xe 10 ⁻¹⁰ cm ³ STP/	/g ¹²⁴ Xe/ ¹³² Xe	¹²⁶ Xe/ ¹³² Xe	¹²⁸ Xe/ ¹³² Xe	¹²⁹ Xe/ ¹³² X	e ¹³⁰ Xe/ ¹³² Xe	¹³¹ Xe/ ¹³² X	e ¹³⁴ Xe/ ¹³² Xe	e ¹³⁶ Xe/ ¹³² Xe
800	1.5	0.00531	0.00366	0.0802	1.331	0.1493	0.8077	0.4260	0.3896
		±0.00033	±0.00027	±0.0014	±0.016	±0.0083	±0.0026	±0.0023	±0.0027
1000	2.0	0.00519	0.00392	0.0835	1.120	0.1623	0.828	0.3935	0.3366
		±0.00031	±0.00039	±0.0015	±0.005	±0.0033	±0.013	±0.0015	±0.0022
1200	3.4	0.00489	0.00471	0.0832	1.376	0.1648	0.8262	0.3811	0.3205
		±0.00017	±0.00039	±0.0014	±0.007	±0.0017	±0.0076	±0.0010	±0.0019
1400	1.4	0.00457	0.00455	0.0842	1.928	0.1676	0.837	0.3825	0.3243
		±0.00026	±0.00050	±0.0015	±0.025	±0.0038	±0.018	±0.0022	±0.0029
1600	0.2	0.0095	0.0086	0.0820	2.19	0.188	0.96	0.374	0.315
		±0.0037	±0.0037	±0.0050	±0.27	±0.036	±0.18	±0.012	±0.017
Total	8.6	0.00508	0.00436	0.08287	1.416	0.1624	0.8279	0.3922	0.3372
		±0.00015	±0.00022	±0.00074	±0.010	±0.0021	±0.0064	±0.0009	±0.0013

±0.034

±0.19

±2.1

The datum with parenthesis means >20% contributions from the procedural blank.

N.D. = not determined due to small amount of gas released.

^(*) means upper limit

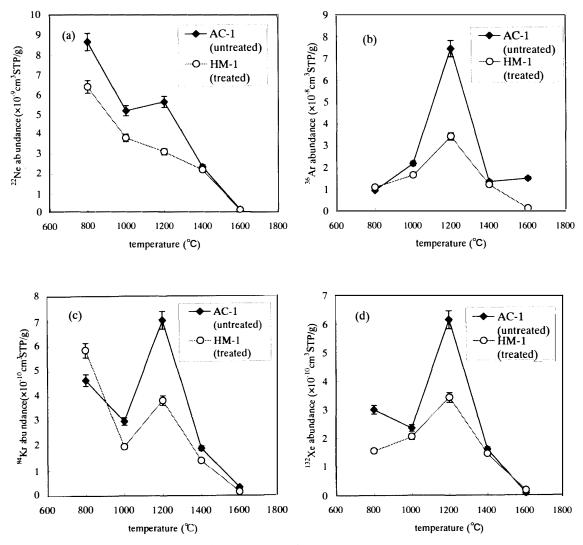


Fig. 1. Release profiles of ²²Ne, ³⁶Ar, ⁸⁴Kr and ¹³²Xe in the untreated and treated Allende matrix. Isotopic ratios of ²⁰Ne/²²Ne and ³⁸Ar/³⁶Ar indicate the contributions from spallogenic ²²Ne and ³⁶Ar in each temperature step (see Tables 1 and 2), but they probably do not affect the overall trends shown in (a) and (b); because spallogenic ²²Ne occurs in fairly constant proportion in each step and the contribution on ³⁶Ar should apparently be small in comparison to the original components (Q and HL).

these data are plotted in a ¹³⁴Xe/¹³²Xe vs. ¹³⁶Xe/¹³²Xe diagram (Fig. 3), the data points of high temperature fractions form a clear linear array extending from Xe-Q to Xe-HL, suggesting that xenon released at higher temperature steps are explained by variable mixtures of Xe-Q and Xe-HL. As these two components are very close each other in Fig. 2, such a mixing trend is less clear in Fig. 2 than in Fig. 3. Figure 3 is more useful than Fig. 2 to examine the relationship of Xe-Q and Xe-HL because these two components are more separable from each other. Note that, although there are some indications of Q and HL component for Ne, Ar and Kr in the samples (Tables 1 and 2), the presence of additional cosmogenic or atmospheric components makes it difficult to identify the Q and HL mixing relationship in those isotope systems.

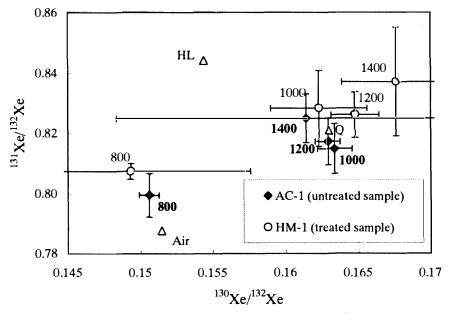


Fig. 2. A ¹³¹Xe/¹³²Xe versus ¹³⁰Xe/¹³²Xe diagram. The data of 1600°C are excluded because of larger uncertainties. The data source of Xe-Q is from Huss et al. (1996), HL from Huss and Lewis (1994) and Air from Nier (1950).

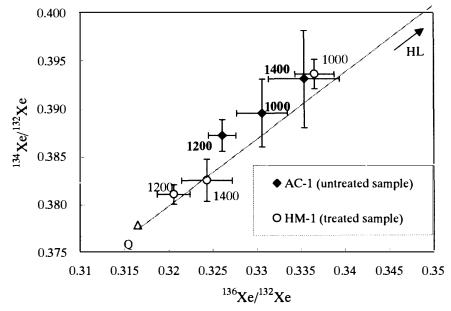


Fig. 3. A ¹³⁴Xe/¹³²Xe versus ¹³⁶Xe/¹³²Xe diagram. The data of 800°C are excluded because of air contamination. The data of 1600°C are excluded because of the large uncertainty. The data source of Q is from Huss et al. (1996), HL is from Huss and Lewis (1994).

3.3. Nature of the lost xenon by artificial hot water alteration

An interesting point to be noted in Fig. 3 is that the xenon isotopic ratios observed in the treated sample are systematically closer to Xe-Q compared to those of the untreated sample. This feature is most clearly seen in xenon released at 1200 and 1400°C where we observed the major gas loss during the hot water treatment (Fig. 1). The shifts of data points towards Xe-Q after the hot water treatment

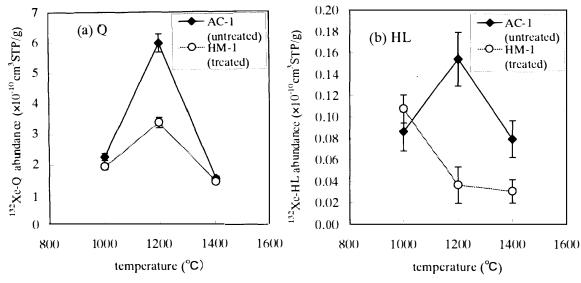


Fig. 4. The release profiles of 132 Xe-Q and 132 Xe-HL from AC-1 and HM-1. The data of 800 and 1600° C are excluded because of air contamination and large uncertainty, respectively. 132 Xe-Q and 132 Xe-HL concentrations are determined by the following equations; $\alpha(^{136}$ Xe/ 132 Xe)_Q+ $(1-\alpha)(^{136}$ Xe/ 132 Xe)_{HL}= $(^{136}$ Xe/ 132 Xe)_{measured} 132 Xe-Q= $\alpha(^{132}$ Xe)_{measured} 132 Xe-HL= $(1-\alpha)(^{132}$ Xe)_{measured}

Table 3. The percentages of the concentrations of \$^{132}Xe-Q\$ and \$^{132}Xe-HL\$ in HM-1 (the sample after the hydrothermal treatment) relative to those in AC-1 (the starting material).

¹³² Xe-Q	¹³² Xe-HL	Note
57±4	23±11	Only for the 1200°C fraction
69 ± 3	54 ± 10	For the 1000-1400°C fractions
	77± 7	Only for the 800°C fraction*
69±3	66± 6	Xenon-Q for the 1000-1400°C fractions and Xe-HL for the 800-1400°C fractions*

^{*} The concentration of ¹³²Xe-HL for the 800°C fraction was calculated on the assumption that Xe in the 800°C fraction was the mixture of Air and Xe-HL.

indicate that Xe-HL had been lost from the samples more significantly than Xe-Q for these temperatures. For the quantitative discussion, we have calculated the fractions of ¹³²Xe-Q and ¹³²Xe-HL components in each temperature step (see caption to Fig. 4 for the calculation). Figure 4 shows the release profiles of the deconvolved ¹³²Xe-Q and ¹³²Xe-HL for temperatures between 1000 to 1400°C. As shown in Fig. 4, the hot water treatment had removed 77% of ¹³²Xe-HL from AC-1, whereas this treatment had removed 43% of ¹³²Xe-Q at 1200°C fraction. This feature is also obvious when we compare the amount of ¹³²Xe-Q and that of ¹³²Xe-HL for the 1000 to 1400°C steps (Table 3). About 70% of ¹³²Xe-Q observed in AC-1 at these temperatures had been survived the hot water treatment while the HM-1 preserved 54% of the original ¹³²Xe-HL released between 1000 to 1400°C.

It has been reported that ordinary chondrites lost their Xe-HL component more significantly than Xe-Q by thermal metamorphism, which indicates that the carrier of Xe-HL is less resistant to thermal metamorphism than that of Xe-Q (Huss et al., 1996). Although the conditions of pressure, temperatures and duration time of hot water alteration experiment were quite different from those of thermal metamorphism, the present results indicate that the HL carrier is indeed less retentive than phase-Q during thermal and hydrothermal alterations as far as the results of high temperature steps are concerned.

We have excluded the data of 800°C fraction in the above discussion. They were largely air-contaminated, so the deconvolution with Xe-HL and Xe-Q was impossible for the 800°C data. However, it should also be noted that ¹³⁶Xe/¹³²Xe ratios in 800°C fractions are highest among those in other temperature steps (Tables 1 and 2), indicating that considerable amounts of ¹³²Xe-HL were released at 800°C. Thus, we calculated the amounts of ¹³²Xe-HL by assuming that Xe for the 800°C fractions was the mixture of only Air and Xe-HL. This would be reasonable as the 800°C data plot close to an Air and Xe-HL mixing trend in Fig. 2. The calculated amounts of ¹³²Xe-HL in the 800°C fractions are the largest among those estimated in all temperature steps in AC-1 and HM-1. We also observed the loss of ¹³²Xe-HL after the hot water treatment in the 800°C fraction, but its degree (23%) is apparently smaller than those observed in higher temperature fractions. As a result, when we include the 800°C fraction results for ¹³²Xe-HL, the suggested difference in retentivities of Xe-Q and Xe-HL became less clear (last row in Table 3). Therefore, although the high temperature fractions do suggest the higher gas retentivity for Xe-Q than Xe-HL upon the alteration as observed in the actual chondrites, further studies are necessary to clarify this feature because both components appeared to have been lost from the samples to a similar extent as a whole by the hot water treatment.

4. Summary

We have carried out a laboratory experiment that simulates the aqueous alteration of carbonaceous chondrites on their parent body. Elemental abundances and isotopic compositions of noble gas were measured in Allende before and after the alteration, and we found that:

- 1) The total abundance of each noble gas in Allende was decreased after the hot water treatment. This result confirms the observation of generally lower noble gas amounts in the CM chondrites that might have lost their noble gases by aqueous alterations on their parent body (Browning et al., 1996; Nakamura et al., 1999).
- 2) The xenon isotopic ratios suggest that the host phase of Xe-Q is less susceptive to gas loss by the hydrothermal alteration than that of Xe-HL for high temperature fractions. However, this feature is not clear when we compare the data for the total fraction. Further experiments should be necessary to examine the difference of the retentivity of Xe-Q and Xe-HL in more detail.

Acknowledgments

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References

- Armstrong, J. T., Meeker, G. P., Huneke, J. C. and Wasserburg, G. J. (1982): The Blue Angel: I. The mineralogy and petrogenesis of a hibonite inclusion from the Murchison meteorite. Geochim. Cosmochim. Acta, 46, 575-596.
- Browning, L. B., McSween, H. Y., Jr. and Zolensky, M. E. (1996): Correlated alteration effects in CM carbonaceous chondrites. Geochim. Cosmochim. Acta, 60, 2621–2633.
- Cohen, R. E., Kornacki, A. S. and Wood, J. A. (1983): Mineralogy and petrology of chondrules and inclusions in the Mokoia CV3 chondrite. Geochim. Cosmochim. Acta, 47, 1739–1757.
- DuFresne, E. R. and Anders, E. (1962): On the chemical evolution of the carbonaceous chondrites. Geochim. Cosmochim. Acta, 26, 1085-1114.
- Duke, C. L. and Brearley, A. J. (1998): Experimental aqueous alteration of Allende. Meteorit. Planet. Sci., 33, No. 4, Suppl., A43.
- Huss, G. R. and Lewis, R. S. (1994): Noble gases in presolar diamonds I: Three distinct components and their implications for diamond origins. Meteoritics, 29, 791–810.
- Huss, G. R., Lewis, R. S. and Hemkin, S. (1996): The "normal planetary" noble gas component in primitive chondrites: Compositions, carrier, and metamorphic history. Geochim. Cosmochim. Acta, 60, 3311-3340.
- Imae, N. and Shinoda, K. (1998): An Experimental study of hydrous mineral formation reaction between enstatite and water vapor. Meteorit. Planet. Sci., 33, No. 4, Suppl., A73-A74.
- Kerridge, J. F., Mackay, A. L. and Boynton, W. V. (1979): Magnetite in CI carbonaceous chondrites: Origin by aqueous activity on a planetesimal surface. Science, 205, 395–397.
- Lewis, R. S., Srinivasan, B. and Anders, E. (1975): Host phase of a strange xenon component in Allende. Science, 190, 1251-1262.
- Lewis, R. S., Tang, M., Wacker, J. F., Anders, E. and Steel, E. (1987): Interstellar diamonds in meteorites. Nature, 326, 160-162.
- Matsuda, J., Amari, S. and Nagao, K. (1999): Purely physical separation of a small fraction of the Allende meteorite that is highly enriched in noble gases. Meteorit. Planet. Sci., 34, 129–136.
- McSween, H. Y., Jr. (1979): Alteration in CM carbonaceous chondrites inferred from modal and chemical variations in matrix. Geochim. Cosmochim. Acta, 43, 1761–1770.
- Nakamura, T., Nagao, K., Metzler, K. and Takaoka, N. (1999): Heterogeneous distribution of solar and cosmogenic noble gases in CM chondrites and implications for the formation of CM parent bodies. Geochim. Cosmochim. Acta, 63, 257-273.
- Nier, A. O. (1950): A redetermination of the relative abundances of the isotopes of neon, krypton, rubidium, xenon, argon, and potassium. Phys. Rev., 79, 789–793.
- Prinn, R. G. and Fegley, B. J. (1987): The atmospheres of Venus, Earth and Mars: A critical comparison. Ann. Rev. Earth Planet. Sci., 15, 171-212.
- Richardson, S. M. (1978): Vein formation in the CI carbonaceous chondrites. Meteoritics, 13, 141–159. Tomeoka, K. and Buseck, P. R. (1982): Intergrown mica and montmorillonite in the Allende carbonaceous chondrite. Nature, 299, 326–327.
- Tomeoka, K. and Buseck, P. R. (1985): Indicators of aqueous alteration in Cm carbonaceous chondrites: Microtextures of a layered mineral containing Fe, S, O and Ni. Geochim. Cosmochim. Acta, 49, 2149–2163.
- Zolensky, M. E. (1984): Hydrothermal alteration of CM carbonaceous chondrites: Implications of the identification of tochilinite as one type of meteoritic PCP. Meteoritics, 19, 346-347.
- Zolensky, M. E., Barrett, R. A. and Gooding, J. L. (1987): Mineralogical variations within the matrices of CM carbonaceous chondrites. Meteoritics, 22, 544-545.

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