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**A New Anisotropic Poroelasticity Model to Describe Damage Accumulation
During Cyclic Triaxial Loading of Rock.**

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29 **Summary**

30 Crustal rocks undergo repeated cycles of stress over time. In complex tectonic
31 environments where stresses may evolve both spatially and temporally, such as
32 volcanoes or active fault zones, these rocks may experience not only cyclic loading
33 and unloading, but also rotation and/or reorientation of stresses. In such situations,
34 any resulting crack distributions form sequentially and may therefore be highly
35 anisotropic. Thus, the tectonic history of the crust as recorded in deformed rocks may
36 include evidence for complex stress paths, encompassing different magnitudes and
37 orientations. Despite this, the ways in which variations in principal stresses influence
38 the evolution of anisotropic crack distributions remain poorly constrained. In this
39 work, we build on the previous non-linear anisotropic damage rheology model by
40 presenting a newly developed poroelastic rheological model which accounts for both
41 coupled anisotropic damage and porosity evolution. The new model shares the main
42 features of previously developed anisotropic damage and scalar poroelastic damage
43 models, including the ability to simulate the entire yield curve through a single
44 formulation. In the new model, the yield condition is defined in terms of invariants of
45 the strain tensor, and so the new formulation operates with directional yield
46 conditions (different values for each principal direction) depending on the damage
47 tensor and triaxial loading conditions. This allows us to discern evolving yield
48 conditions for each principal stress direction and fit the measured amounts of
49 accumulated damage from previous loading cycles. Coupling between anisotropic
50 damage and anisotropic compaction along with the damage-dependent yield condition
51 produces a reasonable fit to the experimentally obtained stress-strain curves.
52 Furthermore, the simulated time-dependent cumulative damage is well correlated with
53 experimentally observed acoustic emissions during cyclic loading in different
54 directions. As such, we are able to recreate many of the features of the experimentally
55 observed directional 3D Kaiser ‘damage memory’ effect.

56

57 **Keywords:** Mechanics, theory, and modelling; Elasticity and anelasticity; Creep and
58 deformation; Fracture and flow.

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61

62 1. Introduction

63 It is well-established that crack damage is generated in brittle rocks that are
64 subjected to a level of stress above some crack initiation threshold, and that this
65 cracking results in the output of elastic wave energy in the form of acoustic emissions
66 (AEs) (e.g., Meredith *et al.* 1990, Holcomb 1993, Lockner 1993a). During cyclic
67 loading, cracks close elastically during unloading and re-open elastically during re-
68 loading. If the level of stress during re-loading remains below the peak stress level
69 attained in any previous loading cycle, then no new cracking occurs and no further
70 cracking-related AE is generated. However, on any loading cycle where the previous
71 peak stress is reached or exceeded, new cracks are formed and are accompanied by
72 concomitant AE output (Kurita & Fujii 1979, Holcomb & Costin 1986, Li &
73 Nordlund 1993, Lockner 1993a, Pestman & Van Munster 1996, Lavrov 2001, 2003,
74 Browning *et al.* 2017, 2018). This observation of AE output only recommencing
75 when the previous maximum stress level is exceeded is known as the *Kaiser effect*
76 (Kaiser 1953) and is related to the ability of a material to accumulate and reproduce
77 information about previously experienced stress states. However, most experiments
78 that have probed aspects of the Kaiser effect to date have been conducted during
79 either uniaxial or conventional triaxial compression experiments and so have not been
80 able to probe fully for any directionality in crack damage accumulation related to the
81 orientation of principal stresses.

82 More recently, Browning *et al.* (2017, 2018) investigated the occurrence of a
83 Kaiser effect in samples of Darley Dale sandstone subjected to both conventional and
84 true triaxial stress conditions. Samples were loaded sequentially to increasing levels
85 of peak stress, both with the maximum principal stress maintained in the same
86 orientation and with the maximum principal stress rotated and applied sequentially in
87 three orthogonal orientations. Their results showed that, under true triaxial loading,
88 crack damage is a distinctly directional phenomenon, such that rocks can exhibit a
89 three-dimensional, directionally-dependent Kaiser effect, with AE only being
90 generated when the previous peak stress in any specific orientation was exceeded.
91 They therefore concluded that the Kaiser effect should more accurately be described
92 as a *damage memory effect* rather than a *stress memory effect*.

93 Traditionally, the analysis of rock deformation and failure criteria has been
94 formulated by, for example, a classical Coulomb-Mohr condition that defines brittle

95 failure, and by a yield cap criterion that defines cataclastic flow (e.g., Issen &
96 Rudnicki 2000). However, these formulations usually ignore any connection between
97 yield stress and the amount of inelastic damage in the form of microcracks, voids, or
98 other flaws that leave the yield stress unchanged, and in doing so, ignore the
99 underlying principle of the Kaiser effect. Laboratory experiments on porous rocks
100 demonstrate evidence of overall strain hardening and yield cap growth attributed to
101 plasticity and porosity loss (Baud *et al.* 2006, Tembe *et al.* 2008, Bedford *et al.* 2018).
102 Several models have been developed for elasto-plastic deformation of isotropic soils,
103 which are commonly formulated in a framework of continuum mechanics and can be
104 successfully applied to model rock behavior with complex yield conditions.

105 For example, the original Cam-Clay model (Roscoe & Burland 1968) provides
106 a description for the stress versus inelastic strain behavior for yield envelopes of any
107 shape defined in stress space (Muir Wood 1990). Modified Cam-Clay yield functions
108 were successively used in geo-mechanical modelling of hydrocarbon reservoirs (Chan
109 *et al.* 2004, Crawford *et al.* 2011) and in more generic studies of inelastic sandstone
110 deformation (Schultz & Siddharthan 2005, Skurtveit *et al.* 2013). In the modified
111 Cam-clay formulation, the stress conditions required for yield are described by the
112 elliptical function of differential and mean effective stress values. Grueschow &
113 Rudnicki (2005) discussed the various models that incorporate different shapes of the
114 evolving yield caps rather than the elliptic function and compared their model with
115 previous suggested by DiMaggio & Sandler (1971) and Carroll (1991). These studies
116 demonstrated that inelastic behavior of porous rocks are well described by various
117 plasticity models. Pijenburg *et al.* (2019) quantified the elastic and inelastic
118 contributions to the total deformation behavior of Slochteren sandstones and
119 concluded that not only the expanding yield envelopes, but also change in the elastic
120 moduli should be considered in order to obtain a proper fit to the experimental stress-
121 strain data. Damage rheology models are able to incorporate changes in both the local
122 elastic properties and the form of the porosity-induced yield cap such that deformation
123 patterns and modes of failure can be analyzed alongside the yield cap growth (e.g.,
124 Bercovici *et al.* 2001, Stefanov *et al.* 2011, Lyakhovsky *et al.* 2015, Vorobiev 2019).

125 It has been suggested that the observed Kaiser effect in rocks indicates strain
126 hardening in consecutive cycles such that the phenomena can be attributed to changes
127 in yield surface due to damage accumulation (Holcomb 1993). Damage and yield
128 surface growth are then likely coupled, and constraints on this coupling can aid

129 interpretation of damage localization patterns and the Kaiser effect (Gajst *et al.* 2020).
130 Damage evolution and time-dependent behavior in low porosity sandstones have been
131 investigated (Choens *et al.* 2021) through creep and conventional triaxial experiments,
132 and numerical analyses. Such quasi-static and creep experiments have been
133 successfully simulated using the modified poroelastic damage model of Lyakhovsky
134 *et al.* (2015). As damage accumulates in the samples, the yield cap evolves to keep
135 pace with the strain accumulation.

136 True triaxial experiments reported by Browning *et al.* (2017, 2018)
137 demonstrated that the orientation of distributed microcracks in Darley Dale sandstone
138 samples are essentially anisotropic and therefore require an extension of isotropic
139 damage models using a scalar damage parameter and a more complex formulation
140 that introduces a second-order damage tensor (Pantelev *et al.* 2021). The goal of this
141 paper is, therefore, to provide a complete quantitative description of the rheological
142 model with directional yield conditions (i.e., different values for each principal
143 direction) depending on the damage tensor and triaxial loading conditions. The new
144 model combines and extends the results of the previously developed anisotropic
145 damage model of Pantelev *et al.* (2021) and the scalar poroelastic damage model of
146 Lyakhovsky *et al.* (2015). The new analysis includes the ability to simulate yield
147 curves through a single formulation and recreates many of the features of the
148 experimentally observed directionally-dependent Kaiser damage memory effect
149 reported by Browning *et al.* (2018).

150

151 **2. Anisotropic Poroelastic Damage model**

152 2.1 Damage and porosity

153 Rock deformation is associated with the formation and growth of internal
154 flaws. From a mechanical point of view, these flaws can be divided into two classes:
155 1) microcracks (damage) contained in the matrix of a porous rock which act as
156 primary stress raisers or stress concentrators and hence contribute to brittle failure,
157 and 2) pores which can deform and before their collapse act to dissipate or
158 accommodate stress and hence contribute to distributed flow. For an isotropic rock
159 with a sufficiently large number of microcracks and pores, one can define a
160 representative volume in which the flaw density is uniform and described by two
161 scalar variables, damage (α) and porosity (φ). The damage variable is a mechanical

162 variable, which is responsible for the change in material stiffness and brittle failure at
163 a critical level of damage. For anisotropic rocks, we can consider a damage tensor,
164 Ω_{ik} , which represents not only the density of microcracks, but also their orientations.
165 The porosity variable is a geometrical property representing the volume fraction of
166 pores during and after deformation. As an alternative to porosity, we define a
167 compaction-strain tensor, ψ_{ij} , which is equal to the accumulated irreversible strain
168 resulting from loading and unloading. This tensor then represents not only the pore
169 volume change, but also deviations in the shape of the pores.

170 In the following sub-sections we describe the general thermodynamic
171 approach used to construct the scalar damage and poroelastic damage model, and
172 provide the main equations of the new anisotropic poroelastic damage model.
173 Detailed thermodynamic relations are provided in Appendix A and specific relation
174 for the isotropic and anisotropic model formulations are provided in Appendixes B
175 and C.

176

177 2.2 General thermodynamic approach

178 We derive the main equations of the poroelastic damage model using the basic
179 relations of irreversible thermodynamics, which provide constraints on the rates of
180 dissipative processes (e.g., (Onsager 1931, Biot 1955, Prigogine 1955, Truesdell &
181 Noll 2004, DeGroot & Mazur 2013). This approach has been applied successfully to
182 understand the kinetics of chemical reactions and phase transitions (e.g., Fitts 1962;
183 DeGroot and Mazur 2013), and as the basis for variational methods of continuous
184 media models (e.g., Sedov 1968, 1997, Malvern 1969, Berdichevsky 2009). The
185 constitutive behavior of the material, and flow rules controlling the kinetics of related
186 irreversible processes, is then entirely defined by specification of two potentials. The
187 first is the free energy, F , and the second is the dissipation function or local entropy
188 production, Γ . This approach has been used as the basis for other damage models (e.g.
189 Valanis 1990, Hansen & Schreyer 1994, Lyakhovskiy *et al.* 1997, 2015, Bercovici *et*
190 *al.* 2001, Hamiel *et al.* 2004a,b, Gaede *et al.* 2013). Following Onsager (1931), who
191 theoretically generalized the empirical laws of Fourier, Ohm, Fick, and Navier (see
192 review by Martyushev & Seleznev 2006) , we represent the specific local entropy
193 production as a product of thermodynamic fluxes and thermodynamic forces. For
194 small deviations from equilibrium, the Onsager principle can be obtained from the

195 maximum entropy production principle, the maximum dissipation rate of mechanical
 196 energy, or the von Mises principle (e.g., Martyushev & Seleznev 2006, Ziegler 2012).
 197 We now discuss the different forms of the energy function, beginning with the scalar
 198 isotropic damage formulation, then the coupling isotropic damage and porosity model,
 199 and finally we formulate the anisotropic model. The energy and entropy balance
 200 equations and general thermodynamic relations are provided in Appendix A.

201

202 2.3 Scalar damage and poroelastic damage model

203 The free energy of a solid (F) in the local damage model of Lyakhovsky *et al.*
 204 (1997) is assumed to be a function of the state variables, which are the temperature T ,
 205 the elastic strain tensor $\varepsilon_{ij} = g_{ij}^{(t)} - g_{ij}^{(0)}$ (the difference between the total strain
 206 tensor $g_{ij}^{(t)}$ and the irreversible strain tensor $g_{ij}^{(0)}$), and the scalar damage variable α :

$$207 \quad F = F(T, \varepsilon_{ij}, \alpha) \quad (1)$$

208 Using the balance equations for the energy and entropy, the Gibbs relation and the
 209 Murnaghan (1937) definition of the stress tensor, part of local entropy (Γ) production
 210 associated with evolving damage is:

$$211 \quad \Gamma = -\frac{\partial F}{\partial \alpha} \frac{d\alpha}{dt} \geq 0 \quad (2)$$

212 The complete thermodynamic derivations are presented in the Appendix A, where all
 213 the dissipation processes are discussed. Following the Onsager (1931) principle, the
 214 kinetic relation for damage evolution is:

$$215 \quad \frac{d\alpha}{dt} = -C \frac{\partial F}{\partial \alpha} \quad (3)$$

216 where C is the positive kinetic coefficient, which may be either constant or depend on
 217 the state variables.

218 Hamiel *et al.* (2004b) and then Lyakhovsky *et al.* (2015) and extended the
 219 scalar damage model to permit coupling of damage and porosity in their formulations.
 220 They followed Biot's theory of poroelasticity (Biot 1941, 1956) representing the free
 221 energy of a poroelastic medium as a sum of the elastic energy, and the Biot
 222 poroelastic coupling terms of the saturated medium. The free energy (1) is extended to
 223 be a function of two additional state variables, fluid volume content, ζ , and material
 224 porosity, φ :

$$225 \quad F = F(T, \varepsilon_{ij}, \alpha, \varphi, \zeta) \quad (4)$$

226 As both porosity and damage can evolve with time during deformation, their coupled
 227 kinetic equations are derived using a similar balance equation leading to the following
 228 local entropy production (Hamiel *et al.* 2004b, Lyakhovsky *et al.* 2015):

$$229 \quad \Gamma = -\frac{\partial F}{\partial \alpha} \frac{d\alpha}{dt} - \left(\frac{\partial F}{\partial \varphi} + \sigma_m \right) \frac{d\varphi}{dt} \geq 0 \quad (5)$$

230 where, $\sigma_m = -\sigma_{kk}/3$ is the mean stress. Once more, adopting the relations from
 231 Onsager (1931) gives a set of two coupled differential equations (Malvern 1969,
 232 DeGroot & Mazur 2013) which define the damage and porosity evolution:

$$233 \quad \frac{d\varphi}{dt} = -C_{\varphi\varphi} \left(\frac{\partial F}{\partial \varphi} + \sigma_m \right) - C_{\varphi\alpha} \frac{\partial F}{\partial \alpha} \quad (6a)$$

$$234 \quad \frac{d\alpha}{dt} = -C_{\alpha\varphi} \left(\frac{\partial F}{\partial \varphi} + \sigma_m \right) - C_{\alpha\alpha} \frac{\partial F}{\partial \alpha} \quad (6b)$$

235 These phenomenological kinetic equations guarantee the non-negative value of
 236 entropy production if the matrix of the kinetic coefficients;

$$237 \quad C_{ij} = \begin{vmatrix} C_{\varphi\varphi} & C_{\varphi\alpha} \\ C_{\alpha\varphi} & C_{\alpha\alpha} \end{vmatrix} \quad (7)$$

238 meets the following conditions (Malvern 1969, DeGroot & Mazur 2013): the diagonal
 239 cells ($C_{\varphi\varphi}$, $C_{\alpha\alpha}$) must be positive, and the off-diagonal terms are usually taken to be
 240 either symmetric or antisymmetric. Following the poroelastic damage model of
 241 Hamiel *et al.* (2004b) and Lyakhovsky *et al.* (2015) we adopt an antisymmetric
 242 structure ($C_{\varphi\alpha} = -C_{\alpha\varphi} = D$) of the kinetic matrix (7). These conditions assure
 243 positive dissipation, as in eq. 5. Larger D -values then lead to an earlier onset of
 244 damage and enhanced accumulation under the same confinement conditions. Hamiel
 245 *et al.* (2004b) and Lyakhovsky *et al.* (2015) discussed slightly different forms of the
 246 stress- or strain-dependent D -value and demonstrated how their scalar poroelastic
 247 model reproduces a yield cap and its evolution (see Appendix B for details). Recently
 248 Gajst *et al.* (2020) suggested a model with exponential damage-dependent D -value:

$$249 \quad D(\alpha) = D_1 e^{-D_2 \alpha} (-I_1)^N \sqrt{I_2} \quad (8)$$

250 where the first coefficient D_1 stands for the initial D -value of the damage-free
 251 material, and the second coefficient D_2 controls its decrease with increasing damage.
 252 The role of the exponent $N > 1$ is to control the shape of the yield cap; and this is
 253 further discussed in the supplementary materials. Gajst *et al.* (2020) demonstrated that
 254 the decrease in the D -value with accumulated damage shifts the yield condition or
 255 onset of damage to higher stress values and successfully reproduces the Kaiser effect.
 256 Since the model is formulated in terms of scalar damage and strain invariants, it

257 accurately reproduces the isotropic Kaiser damage-memory effect, but does not
 258 consider the effect of microcrack orientation and stress rotation. The experimentally
 259 observed directionally-dependent Kaiser damage-memory effect (Browning *et al.*
 260 2018) hence requires an anisotropic formulation.

261

262 2.4 Anisotropic poroelastic damage model

263 Recently Panteleev *et al.* (2021) extended the scalar isotropic damage model
 264 by developing a theoretical model for materials with orthotropic symmetry which
 265 describes the material damage using a second rank symmetric tensor, Ω_{ik} , in which
 266 the principal directions match the orientation of the principal loading axes. This
 267 assumption is supported by results of true triaxial experiments (Browning *et al.* 2017,
 268 2018) which demonstrated that the orientation of distributed microcracks was related
 269 to the level and orientation of the principal stresses. Therefore, most of the equations
 270 that follow are written with respect to the principal loading directions and stresses,
 271 while the complete three-dimensional formulation is presented in Appendix C.

272 The scalar damage variable α in the free energy form of equation (1) and the
 273 poroelastic model with the energy form from equation (4) is substituted by a damage
 274 tensor, Ω_{ik} . For the case of an isotropic material ($\Omega_{ij} = \Omega \delta_{ij}$), the anisotropic
 275 formulation reduces to the scalar model with the damage α equal to a squared value,
 276 $\alpha = \Omega^2$. In addition, (Lyakhovsky *et al.* (2022) showed that the deformation of pore
 277 space is inherently three dimensional and, as such, the compaction-strain strain tensor,
 278 ψ_{ij} , should replace porosity in the governing equations. The suggested energy
 279 function includes these two tensor state variables; the damage tensor Ω_{ik} , and the
 280 compaction-strain tensor ψ_{ij} :

$$281 \quad F = F(T, \varepsilon_{ij}, \Omega_{ij}, \psi_{ij}, \zeta) \quad (9)$$

282 The elastic strain tensor $\varepsilon_{ij} = g_{ij}^{(t)} - \psi_{ij}$ is now defined as a difference between the
 283 total strain tensor $g_{ij}^{(t)}$ and the tensor ψ_{ij} . The diagonal part of this tensor, $\varphi = \psi_{ij}\delta_{ij}$
 284 represents the material porosity, while the deviatoric components ($\psi_{ij} - \frac{1}{3}\varphi \delta_{ij}$) are
 285 associated with anisotropic compaction and other mechanisms related to the
 286 irreversible strain accumulation. Using the energy form (9), the dissipation associated
 287 with evolving tensors Ω_{ij}, ψ_{ij} , consists of two terms which are proportional to the
 288 their rate of change (see Appendix C for detailed derivations):

289
$$\Gamma = -\frac{\partial F}{\partial \Omega_{ij}} \frac{d\Omega_{ij}}{dt} + \left(\sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \right) \frac{d\psi_{ij}}{dt} > 0 \quad (10)$$

290 The phenomenological kinetic equations share the same structure with the poroelastic
291 formulation (6), but connect the tensor quantities as follows:

292
$$\frac{d\psi_{ij}}{dt} = C_{ijnm}^{\psi\psi} \left(\sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \right) - C_{ijnm}^{\psi\Omega} \frac{\partial F}{\partial \Omega_{ij}} \quad (11a)$$

293
$$\frac{d\Omega_{nm}}{dt} = C_{ijnm}^{\Omega\psi} \left(\sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \right) - C_{ijnm}^{\Omega\Omega} \frac{\partial F}{\partial \Omega_{ij}} \quad (11b)$$

294 Instead of the matrix (7) of the kinetic coefficients, every matrix term becomes a
295 fourth-rank tensor that can be written as:

296
$$C_{ijnm} = \begin{bmatrix} C_{ijnm}^{\psi\psi} & C_{ijnm}^{\psi\Omega} \\ C_{ijnm}^{\Omega\psi} & C_{ijnm}^{\Omega\Omega} \end{bmatrix} \quad (12)$$

297 The kinetic equations (11) guarantee a non-negative value of entropy production if the
298 cells of the matrix of the kinetic coefficients meet conditions like those of the
299 poroelastic model: 1) matrices of the diagonal cells ($C_{ijnm}^{\psi\psi}$, $C_{ijnm}^{\Omega\Omega}$) must be positively
300 defined; and 2) we also adopt an antisymmetric structure ($C_{ijnm}^{\Omega\psi} = -C_{ijnm}^{\psi\Omega}$) for the
301 off-diagonal terms, as was done previously for the poroelastic model.

302 In the next section, we specify the energy function (9) and kinetic coefficients
303 (12), and then demonstrate the main model features.

304

305 2.5 Energy function and kinetic equations, anisotropic model

306 The energy function for the anisotropic damage model includes a damage
307 tensor Ω_{ij} and so cannot be formulated only in terms of invariants of the strain tensor:

308
$$\begin{aligned} I_1 &= \varepsilon_{ij} \delta_{ij} \\ I_2 &= \varepsilon_{ij} \varepsilon_{ij} \end{aligned} \quad (13)$$

309 Following Murti et al., (1991) and Zhang and Cai, (2010), Panteleev *et al.* (2021)
310 incorporated invariants $I_1^{(\Omega)}$ and $I_2^{(\Omega)}$ of the tensor $\varepsilon_{ij}^{(\Omega)} = \frac{1}{2}(\varepsilon_{ik}\Omega_{kj} + \varepsilon_{jk}\Omega_{ki})$. In the
311 coordinate system of the principal damage values, these invariants are (see Appendix
312 C for the general case):

313
$$I_1^{(\Omega)} = \varepsilon_{11}\Omega_1 + \varepsilon_{22}\Omega_2 + \varepsilon_{33}\Omega_3$$

314
$$I_2^{(\Omega)} = (\varepsilon_{11}\Omega_1)^2 + (\varepsilon_{22}\Omega_2)^2 + (\varepsilon_{33}\Omega_3)^2 +$$

315
$$+ \frac{1}{2}\varepsilon_{12}^2(\Omega_1 + \Omega_2)^2 + \frac{1}{2}\varepsilon_{13}^2(\Omega_1 + \Omega_3)^2 + \frac{1}{2}\varepsilon_{23}^2(\Omega_2 + \Omega_3)^2 \quad (14)$$

316 We extend the energy function of Panteleev *et al.* (2021) using additional terms of
 317 Biot's theory of poroelasticity (Biot 1941, 1956); see also Hamiel *et al.* (2004b) and
 318 Lyakhovsky *et al.* (2015):

$$319 \quad F = \left[\frac{\lambda_0}{2} I_1^2 + \mu_0 I_2 \right] + \left[\mu_1 I_2^{(\Omega)} - \gamma I_1^{(\Omega)} \sqrt{I_2^{(\Omega)}} \right] + \frac{M}{2} [\beta I_1 - \zeta + \psi_{ij} \delta_{ij}]^2 + Ch \frac{\gamma}{2} \Omega_{ij} \Omega_{ij}$$

320 (15)

321 The energy function for nonlinear poroelastic damaged media includes two Hookean
 322 terms with the Lamé drained moduli of the intact (damage-free) rock λ_0, μ_0 and two
 323 second order terms with strain invariants $I_1^{(\Omega)}$ and $I_2^{(\Omega)}$. The modulus μ_1 controls the
 324 reduction of the effective shear modulus, and the coupling modulus γ is responsible
 325 for enhanced nonlinearity with damage accumulation (Panteleev *et al.* 2021). The
 326 third term in squared brackets, with Biot modulus M and coefficient β , differs from
 327 the classical poroelasticity only by the term $\psi_{ij} \delta_{ij}$, that represents the porosity.
 328 Similarly to the scalar model (Appendix B), we introduce the damage-dependent term
 329 with the coefficient Ch (Gajst 2020), which allows us to account for the cohesive
 330 forces that influence rock fracture under low confining pressures.

331 Following the definitions of the scalar damage model, we use $\mu_1 = \xi_0 \gamma$ with
 332 critical ratio $\xi_0 = I_1 / \sqrt{I_2}$ of the strain invariants (13). The ξ_0 value is related to the
 333 internal friction angle of the intact rock (Agnon & Lyakhovsky 1995) and controls the
 334 onset of damage accumulation in the scalar damage model as well as in the
 335 anisotropic model for the material with isotropic damage ($\Omega_{ij} = \Omega \delta_{ij}$). With this
 336 notation, the stress-strain constitutive relation for for k -component of the principal
 337 stress and damage values is (see Appendix C for complete relation):

$$338 \quad \sigma_k = \lambda_0 I_1 + 2\mu_0 \varepsilon_k - \gamma \Omega_k \sqrt{I_2^{(\Omega)}} + \gamma \left(2\xi_0 - \frac{I_1^{(\Omega)}}{\sqrt{I_2^{(\Omega)}}} \right) \varepsilon_k \Omega_k^2 + \beta M (\beta I_1 - \zeta + \psi_k)$$

339 (16)

340 and the fluid pressure is:

$$341 \quad p_f = \frac{\partial F}{\partial \zeta} = M (-\beta I_1 + \zeta - \psi_n) \quad (17)$$

342 Similarly to the Biot poroelasticity, the effective stress is defined as:

$$343 \quad \sigma_{ij}^{eff} = \sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \quad (18)$$

344 Kinetic coefficients (12) should be defined in order to provide the complete
 345 form of the kinetic equations (11). The matrix $C_{ijnm}^{\Omega\Omega}$ multiplied by $\partial F/\partial\Omega_{ij}$ defines
 346 the damage accumulation rate, driven by the thermodynamic force associated with the
 347 damage-dependent energy change. The most conservative assumption to define the
 348 components of this matrix is the absence of any interaction between different damage
 349 components on their kinetics:

$$350 \quad C_{ijnm}^{\Omega\Omega} = \frac{L}{\sqrt{\Omega_{ij}\Omega_{ij}}} (\delta_{ik} \delta_{jn} + \delta_{in} \delta_{jk}) \quad (19)$$

351 This form of damage kinetics was verified by Panteleev *et al.* (2021) using results
 352 from true triaxial rock mechanics experiments, and is therefore adopted here for the
 353 poroelastic model.

354 The off-diagonal antisymmetric coefficient, noted here as $D_{ijkn} = C_{ijnm}^{\Omega\psi} =$
 355 $-C_{ijnm}^{\psi\Omega}$, controls the coupling between irreversible strain (porosity) and damage
 356 accumulation. Extending the Gajst *et al.* (2020) model with an exponential damage-
 357 dependent D -value (8) to the tensor form and using the same type of strain
 358 dependency, we suggest the following form of the coupling kinetic coefficient D_{ijkn} :

$$359 \quad D_{ijkn} = D_1 \exp[-D_2 \Omega_{ij}] \delta_{kn} (-I_1)^N \sqrt{I_2} \quad (20)$$

360 where we use the standard definition of the exponent of the tensor \mathbf{X}_{ij} by means of its
 361 series representation (Hirsch *et al.* 1974):

$$362 \quad \exp(\mathbf{X}_{ij}) = \sum_{m=0}^{\infty} \frac{\mathbf{X}_{ij}^m}{m!} \quad (21)$$

363 Note that the principal values of the tensor $\exp(\mathbf{X}_{ij})$ are equal to the exponent of the
 364 principal value $\exp(\mathbf{X}_k)$.

365 Given the kinetic coefficients (19, 20), the equation for the damage evolution
 366 (11b) for principal components becomes (see Appendix C for complete relation):

$$367 \quad \frac{d\Omega_k}{dt} = K D_1 \exp[-D_2 \Omega_k] (-I_1)^{N+1} \sqrt{I_2} +$$

$$368 \quad + \frac{L}{\sqrt{\Omega_l \Omega_l}} \left[\varepsilon_k \sqrt{I_2^{(\Omega)}} + \left(\frac{I_1^{(\Omega)}}{\sqrt{I_2^{(\Omega)}}} - 2\xi_0 \right) \varepsilon_k^2 \Omega_k - Ch \Omega_k \right] \quad (22)$$

369 The effective mean stress ($\sigma_m^{eff} = -\sigma_{kk}^{eff}/3$) in (11b) was substituted here by the
 370 volumetric strain multiplied by the bulk modulus K ($\sigma_m^{eff} = -K I_1$) leading to the
 371 power index $N+1$.

372 The equation for the irreversible strain accumulation (11a) includes two terms.
373 The first term, $C_{ijnm}^{\psi\psi} \sigma_{nm}^{eff}$, describes compaction/dilation with the rate proportional to
374 the effective stress. The second is the damage-related coupling term multiplied to
375 D_{ijkn} (20), and describes the compaction or dilation associated with the formation and
376 growth of damage. Representing the effective stress as a superposition of the
377 volumetric (σ_m^{eff}) and deviatoric (τ_{ij}) components allows us to describe the different
378 mechanisms of the irreversible strain accumulation, or ψ_{ij} kinetics. The pressure
379 driven compaction in the isotropic case becomes a well-known porosity reduction to
380 its pressure-dependent equilibrium value, or Athy's (1930) law. (Lyakhovsky *et al.*
381 (2022) modified the scalar Athy relation and suggested that the 3-D equilibrium
382 compaction depends on both pressure and deviatoric stress components:

$$383 \quad \psi_{ij}^{(eq)} = B_0 \left[\delta_{ij} - \exp \left(-\frac{\sigma_m^{eff}}{B_1} \delta_{ij} - \frac{\tau_{ij}}{B_2} \right) \right] \quad (23)$$

384 which further suggested the kinetics of the pressure-driven 3-D compaction has the
385 form

$$386 \quad \frac{d\psi_{ij}}{dt} = A \left(\psi_{ij}^{(eq)} - \psi_{ij} \right) \sigma_m^{eff} \quad (24)$$

387 The equations (23, 24) consider not only the closure of voids or changes in the pore
388 space (isotropic porosity reduction), but also changes in void shape under non-
389 hydrostatic loading. Neglecting the term with deviatoric stress, or taking $B_2 \rightarrow \infty$ in
390 (23), reduces both the equilibrium compaction and kinetic equation to the traditional
391 scalar form formulated in terms of material porosity.

392 Experimental studies suggest that permanent inelastic deformation is not only
393 caused by pressure driven compaction, but also starts to accumulate at the onset of
394 microcracking (as evidenced by the output of AE) and increases all the way up to the
395 point of brittle failure (e.g., Lockner 1993, 1998, Martin & Chandler 1994). This
396 process is usually associated with the growth of microcracks and frictional sliding
397 between grains, rather than closure of voids or space between grains. For similar
398 reasons, Hamiel *et al.* (2004a) related the rate of irreversible strain accumulation with
399 the rate of their scalar damage growth. Keeping in mind that the scalar damage
400 variable is equivalent to the squared damage tensor, we extend their relation to:

$$401 \quad \frac{d}{dt} \varepsilon_{ij}^{(ir)} = \begin{cases} C_V \frac{d(\Omega_{in}\Omega_{jm})}{dt} \tau_{nm}, & \frac{d(\Omega_{in}\Omega_{jm})}{dt} > 0 \\ 0, & \frac{d(\Omega_{in}\Omega_{jm})}{dt} \leq 0 \end{cases} \quad (25)$$

402 The final kinetic equation (11) for the ψ_{ij} tensor incorporates all of the discussed
 403 mechanisms and to avoid a lengthy expression, is written for principal components:

$$\begin{aligned}
 404 \quad \frac{d\psi_k}{dt} = & A \left(\psi_k^{(eq)} - \psi_k \right) \sigma_m^{eff} + C_V \frac{d(\Omega_k^2)}{dt} \tau_k + \\
 405 \quad & + D_1 \exp[-D_2 \Omega_k] (-I_1)^N \sqrt{I_2} \left[\varepsilon_k \sqrt{I_2^{(\Omega)}} + \left(\frac{I_1^{(\Omega)}}{\sqrt{I_2^{(\Omega)}}} - 2\xi_0 \right) \varepsilon_k^2 \Omega_k - Ch \Omega_k \right] \quad (26)
 \end{aligned}$$

406 The first term of (26) represents the compaction prior to the onset of damage
 407 accumulation, which is a three-dimensional extension of the scalar Athy's compaction
 408 law. According to this term, the compaction approaches its stress-dependent
 409 equilibrium value with the rate proportional to the effective pressure. The second term
 410 is the three-dimensional equivalent of the damage-dependent irreversible strain
 411 accumulation with inverse of the effective viscosity or fluidity proportional to the rate
 412 of damage accumulation. This term describes extension or compaction depending on
 413 the sign of the deviatoric stress component. The last term represents the coupling
 414 between damage and porosity kinetics. Its sign, extension or compaction, is defined
 415 by the expression in the square brackets and depends on the loading and damage
 416 values.

417 The kinetic expressions for damage (22) and irreversible strain (26) provide
 418 the closed system of equations defining the three-dimensional evolution of the
 419 material properties.

420

421 2.6 Yield cap evolution

422 The new anisotropic poroelastic damage model shares the main features of the
 423 previously developed isotropic model of Lyakhovskiy *et al.* (2015), including the
 424 ability to simulate the entire yield curve through a single formulation. Their model
 425 addressed several different deformation regimes including elastic deformation and
 426 pressure-driven compaction, brittle failure, and cataclastic flow. Loading of a rock
 427 sample to a level of stress beyond the initial yield surface caused an accumulation of
 428 damage and resulted in a porosity change (causing either compaction or dilation). In
 429 this case, the modeled yield cap grows. Consequently, if the sample is unloaded and
 430 then reloaded, the new yield cap is found to occur at a higher stress state, which is in
 431 agreement with the Kaiser effect. These general features of the model yield cap are
 432 shown in strain space, i.e., differential strain ($\varepsilon_1 - \varepsilon_3$) versus volumetric strain (Fig.

433 1) instead more common stress space representation (differential stress vs. pressure).
 434 Several authors formulated yield conditions in terms of strains and demonstrated
 435 significant advantages of this approach for materials with evolving yield conditions as
 436 a function of material properties and loading history (Naghdi & Trapp 1975, Yoder &
 437 Iwan 1981, Han & Chen 1986, Puzrin & Houlsby 2001). Although there are some
 438 similarities between stress-space and strain-space formulations, they are not
 439 equivalent when material weakening is considered (Casey & Naghdi 1983, Einav
 440 2004, 2005). The yield surface (heavy black line in Fig. 1) represents the onset of the
 441 damage accumulation according to kinetic equation (22) reduced to the model version
 442 for the isotropic material ($\Omega_{ij} = \Omega \delta_{ij}$) and non-cohesive material ($Ch=0$):

$$443 \quad \frac{d\alpha}{dt} = C_d [D_1 e^{-D_2 \alpha} (-I_1)^{N+1} \sqrt{I_2} + I_2 (\xi - \xi_0)] \quad (27)$$

444 This form of the damage kinetic equation predicts the onset of damage accumulation
 445 ($d\alpha/dt = 0$), or the yield condition expressed in terms of strain invariants:

$$446 \quad \sqrt{I_2} = [D_1 e^{-D_2 \alpha} (-I_1)^N + I_1] / \xi_0 \quad (28)$$

447 The entire yield curve (heavy black line in Figure 1) is calculated for compactive
 448 volumetric strain (I_1) between zero and a certain critical value I_1^* , corresponding to the
 449 onset of damage under hydrostatic compaction. The differential strain ($\varepsilon_1 - \varepsilon_3$) were
 450 calculated using strain invariants and assuming triaxial loading conditions. The critical
 451 value (I_1^*) is defined from the damage onset or yield condition (28) under hydrostatic
 452 loading corresponding to the strain invariant ratio $I_1 / \sqrt{I_2} = -\sqrt{3}$:

$$453 \quad I_1^* = - \left(\frac{\sqrt{3} + \xi_0}{\sqrt{3} \cdot D_1 e^{-D_2 \alpha}} \right)^{1/N} \quad (29)$$

454 The detailed discussion of the size and shape of the yield envelope and the sensitivity
 455 to the model parameters, including material cohesion, is presented in Appendix B.
 456 The red line in Fig. 1 schematically represents the proportional load path with
 457 constant strain invariant ratio ($\xi = I_1 / \sqrt{I_2} = Const.$). At the initial stage of loading,
 458 when the stress level is beneath the yield cap (Regime I), quasi-elastic deformation is
 459 accompanied by material strengthening associated with crack closure and compaction
 460 (porosity decrease). When the load reaches the yield condition (red star), the damage
 461 accumulation starts. Distributed microcracking and grain crushing enable sliding
 462 along newly created internal surfaces, collapse of the pore space, and changes in grain
 463 packing arrangements leading to an overall porosity decrease. This deformational

464 Regime II is most prominent under high confining pressures and is usually treated as
 465 cataclastic flow associated with compaction. Damage accumulation as well as
 466 porosity reduction leads to the decrease of the coupling coefficient D , which in the
 467 presented version is considered as damage-dependent only (eq. 8). The damage
 468 kinetics affects the yield condition (28) and the yield cap evolves as strain is
 469 accumulated. This feature allows the Kaiser effect to be reproduced, such that the
 470 onset of damage (and its associated AE) occurs at increasingly higher stress levels, if
 471 the sample is unloaded and reloaded along the same loading path.

472 Under elevated values of the differential strain, the loading path crosses the
 473 compaction-dilation transition (dashed line in Fig. 1) separating the deformational
 474 Regime II with compaction and Regime III with inelastic porosity increase
 475 (dilatancy). This transition coincides with the Coulomb-Mohr failure criterion and
 476 meets the condition $\xi = \xi_0$. When the loading is pushed beyond the compaction-
 477 dilation transition, intensive damage accumulation, along with significant differential
 478 strains, lead to material dilation (porosity increase). Damage increase is then bounded
 479 by a certain critical value which eventually corresponds with macroscopic brittle
 480 failure, a dynamic stress drop, and a rapid conversion of the differential elastic strain
 481 into plastic strain components.

482 The damage kinetic equation (22) of the anisotropic model shares a similar
 483 structure to the scalar model (27). It also consists of two competing terms, but their
 484 values depend not only on the strain invariants, but also on the direction of loading
 485 relative to the principal values of the damage tensor. We re-write these terms
 486 assuming that the principal directions of the damage tensor (Ω_k) match the orientation
 487 of the principal loading axes with principal strain values ε_k :

$$488 \quad term_1 = D_1 \exp[-D_2 \Omega_k] (-I_1)^{N+1} \sqrt{I_2} \quad (30)$$

$$489 \quad term_2 = \frac{1}{\sqrt{\Omega_l \Omega_l}} \left[\varepsilon_k \sqrt{I_2^{(\Omega)}} + \left(\frac{I_1^{(\Omega)}}{\sqrt{I_2^{(\Omega)}}} - 2\xi_0 \right) \varepsilon_k^2 \Omega_k - Ch \Omega_k \right] \quad (31)$$

490 We take out two constants, K and L , and rescale D_1 to preserve the same proportion
 491 between these terms. Figure 2 demonstrates the evolution of these terms and their
 492 sums for the conventional loading, where Ω_1 is oriented in the direction of the axial
 493 load and ε_1 is the axial strain in the same direction. In this case, Ω_2, Ω_3 are oriented in
 494 the direction of the transverse loads with two equal values of strain $\varepsilon_2 = \varepsilon_3$. For the
 495 initially isotropic material ($\Omega_1 = \Omega_2 = \Omega_3$), the first term (30) is isotropic, since it

496 depends on the damage and strain invariants. Its value strongly increases with the
 497 volumetric deformation (confinement) and weakly depends on the differential strain
 498 (Fig. 2a). Values of the second term (31) are however different depending on whether
 499 the axial or transverse directions are observed. This is the case even for the isotropic
 500 material (Fig. 2 b1 and 2 b2 since the value explicitly depends on the strain
 501 components. The axial values (Fig. 2) are always negative and increase (i.e., become
 502 more negative) with both volumetric and differential components. Thus, the sum of
 503 the first (30) and second (31) terms for the axial direction is negative for most of the
 504 strain values except for a small area in the right bottom corner of the map Fig. 2 c1.
 505 This implies that damage is not likely to occur in the axial direction. For most of the
 506 loading cases, except under high volumetric stress (strain) above critical value, I_1^*
 507 (29), the axial damage component remains unchanged or even decreases.
 508 Mechanically, this implies that microcracks which are oriented normal to the axial
 509 load direction become closed.

510 The transverse values of the second term (31) are positive at relatively low
 511 volumetric strains and elevated differential strains (upper left corner of Fig. 2b1).
 512 They become negative with volumetric strain increase. This change is indicated by a
 513 heavy red line separating the negative and positive values in Figure 2b2. The
 514 summation of the positive and negative values of the first (Fig. 2a) and second (Fig.
 515 2b2) terms forms the yield cap (black heavy line in Fig. 2c2) which separates areas
 516 corresponding to damage accumulation (positive values) and healing or microcrack
 517 closure (negative values) for the transverse components Ω_2, Ω_3 .

518 The obtained shape of the yield cap for transverse damage components (red
 519 line in Fig. 3c2) is very similar to those predicted by the scalar model (see Figs B1-3
 520 in Appendix B). At low volumetric strains (with low level confinement) the yield cap
 521 is slightly shifted up depending on the cohesive force values (See Fig B3 in Appendix
 522 B for effect of cohesion). Loading above the yield cap leads to gradual accumulation
 523 of two transverse damage components, Ω_2, Ω_3 , that form at the same rate while the
 524 axial damage value, Ω_1 , remains unchanged. This situation creates a stress-induced
 525 cylindrical transverse damage anisotropy observed in triaxial experiments (Browning
 526 *et al.* 2018). The damage can be schematically shown as two families of microcracks
 527 (insert in Fig. 3). One family (blue cracks in Fig. 3) is oriented parallel to the axial
 528 loading direction and opening in the transverse directions. Another family (black

529 cracks in Fig. 3) normally oriented to the axial loading direction is closed-. The axial
530 damage value, Ω_1 , remains unchanged until large hydrostatic volumetric strains above
531 certain critical value, I_1^* close to 0.8%, are applied. This value is similar to that
532 predicted by the scalar model (29) and describes the onset of pore space collapse. The
533 anisotropic model predicts that the axial damage increases under the loading
534 conditions in the area on the right of the black dashed line in Fig. 3.

535 Accumulation of the transverse damage components increases the size of the
536 yield cap which is related to exponential damage-dependence in the first term (30). If
537 the sample is then reloaded, along the same loading path, the damage commences at
538 the point of the previous maximum stress value. This corresponds to the previously
539 mentioned enlargement of the yield cap. This Kaiser effect is schematically shown in
540 Fig. 4 as a “no rotation” cyclic loading. During the first loading cycle (blue path and
541 envelope), the onset of damage occurs at relatively low differential strain (strain
542 values and model parameters will be specified in the next section comparing with
543 experimental results). During the second cycle (purple path and envelope), the onset
544 of damage occurs at a significantly higher value of strain (or stress), close to the
545 maximum level of stress in the previous cycle. The quality of the Kaiser effect then
546 depends on how close the yield cap keeps pace with the strain accumulation. This
547 feature dramatically changes if the loading direction is rotated between consecutive
548 loading cycles. Only a small change between the blue, green, and red yield envelopes
549 is predicted for the case “rotation of the loading direction between cycles”. Similar to
550 the previous case, the starting material has an essentially isotropic distribution of
551 microcracks and flaws at the beginning of the first loading cycle (blue path and
552 envelope in Fig. 4). However, after unloading at the end of the first cycle, the sample
553 is no longer isotropic since new anisotropic microcrack damage has been formed,
554 with the cracks growing parallel to the maximum loading direction (Ω_2, Ω_3) and
555 closure in the plane normal to the maximum loading direction (Ω_1). For the second
556 loading cycle (red line), the loading orientations are rotated on the same sample that
557 contains the previously formed anisotropic damage, such that the maximum loading
558 direction is parallel to the previous Ω_2 . The low Ω_1 and one of the high Ω_3 damage
559 components are now associated with the transverse direction, while another high
560 component Ω_2 is axial. The onset of accumulation of the new Ω_1 damage component
561 occurs at almost exactly the same level of stress as in the first cycle. There is no

562 observation of any Kaiser effect, and this suggests that a completely new family of
563 microcracks are generated normal to those generated on during the first cycle. An
564 exactly similar scenario is observed during the third loading cycle when the sample is
565 re-loaded with the loading orientation again rotated such that the $-\Omega_3$ component
566 becomes parallel to the new maximum loading direction (the green path). The onset of
567 damage again occurs at essentially the same level of stress and strain as in cycle 1 and
568 cycle 2. Again, no Kaiser effect is observed. Analysis of Figure 4 explains why there
569 is apparently no Kaiser effect. The blue, red and green lines related to the first, second
570 and third sequential cycles exhibit very similar yield caps which suggests that the new
571 damage on each cycle is independent of the damage formed during earlier cycles. We
572 would therefore not expect to encounter a Kaiser effect under these conditions. It is
573 only when the sample is loaded to a higher level of stress in the same orientation, as
574 demonstrated by the purple line, that would expect to observe a manifestation of the
575 Kaiser effect.

576 To explain this model prediction further, Figure 5 shows the distribution of
577 damage kinetics terms for the two transverse low Ω_1 and high Ω_3 damage components
578 during the second load cycle for the sample with anisotropic damage $\Omega_1 < \Omega_2 = \Omega_3$. In
579 spite of the elevated damage in the axial direction, similar to the isotropic sample, the
580 axial damage component, Ω_2 , remains unchanged or even decreased or healed (maps
581 are not shown). Values of the first (Fig. 5a1 and a2) and the second (Fig. 5b1 and b2)
582 terms (30, 31) are similar. However, the first term for the Ω_1 component (Fig. 5a2) is
583 slightly larger than for the Ω_3 (Fig. 5a1). The values of the second term show the
584 opposite tendency (Fig. 5b1 and b2). Finally, the sum of these terms controlling the
585 damage kinetics (Fig. 5c1 and c2), gives significantly different yield cap (red heavy
586 lines in Fig. 5 c1,2). Figure 6 summarizes the shape and size of the yield envelope for
587 the anisotropic material (red lines) and compares them with yield caps for the initially
588 isotropic case (dotted lines). The yield cap of the low transverse damage component
589 (red line for Ω_1) is essentially coincident with the yield cap for the isotropic material.
590 However, the yield cap for the high transverse damage component Ω_3 is significantly
591 enlarged. If the loading is not large enough and it is only slightly above the red line
592 for Ω_1 component, the high transverse component Ω_3 remains unchanged and only
593 minor total damage is accumulated in the sample. The yield for the axial Ω_2
594 component is shifted to significantly higher values of volumetric strain, value, I_1^*

595 above 1.0%, but retaining the same shape as the yield for the axial component of the
596 isotropic material.

597

598 **3. Model verification**

599 3.1 Materials and experimental settings

600 In order to verify the ability of our new model to reproduce the features of the
601 experimentally observed three-dimensional Kaiser effect, we use results from two sets
602 of tests reported by Browning *et al.* (2018), namely; the sequential rotational
603 conventional triaxial (SRCT) loading test, and the cyclic sequential rotational
604 conventional triaxial (CSRCT) loading test. The experiments were conducted on dry
605 samples of the relatively homogeneous Darley Dale sandstone, which is a feldspathic
606 sandstone with a moderate porosity of ~13% and a grain size range from 0.08 to 0.8
607 mm (Wu *et al.* 2000, Heap *et al.* 2009). The deformation apparatus, based at the
608 laboratories of Koninklijke Shell Exploratie en Productie Laboratoriu (KSEPL),
609 Rijswijk, Netherlands, consists of a three-axis stressing frame constructed of flanged
610 steel beams, one of which was removable to allow the insertion of the cubic rock
611 samples (edge length of 50 mm). Loading was performed with three pairs of servo-
612 controlled hydraulic rams with a loading capacity of 300 kN. Hemispherical seatings
613 were used along orthogonal axes perpendicular to the faces of the cubic samples and
614 loading platens, with an edge length of 47.5 mm were interposed between the rams
615 and the sample faces to provide the contact surfaces.

616 In order to keep the sample centered within the apparatus and to ensure good
617 acoustic contact between the sample and the loading platens, a small pre-load of 4
618 MPa was applied along each of the three axes prior to testing. The load in each of the
619 three directions was measured using electronic load cells with an accuracy of $\pm 0.2\%$,
620 and the displacement in each direction was measured using linear variable differential
621 transformers (LVDTs) mounted between the loading platens. AE was monitored using
622 a piezo-electric transducer located in a recess within one of the platens.

623

624 3.2 Sequential rotational conventional triaxial (SRCT) loading

625 The SRCT loading test consisted of three loading cycles (Fig. 7a). Prior to the
626 first loading cycle, the sample was pre-loaded hydrostatically up to 4 MPa. Then, in
627 the first loading cycle, the differential stress (blue lines in Figure 7a) was increased in

628 the 1-direction to a maximum value of 80 MPa at a rate of 0.018 MPa/s and then
629 unloaded at the same rate, while both transverse stresses were held constant at 4 MPa.
630 To relate the experimental results to our model formulation, the axial loading
631 direction during the first cycle is associated with the first damage component Ω_1 ,
632 while the transverse damage components are Ω_2 and Ω_3 . In the second cycle, the
633 differential stress was rotated to the 2-direction (red lines in Figure 7a), and the new
634 damage then corresponds to damage component Ω_2 , while the transverse components
635 become Ω_1 and Ω_3 . The loading protocol was the same as in the first cycle such that
636 the differential load was increased at a rate of 0.018 MPa/s to a maximum value of 80
637 MPa and then unloaded at the same rate. Again, the two transverse stresses were held
638 constant at 4 MPa. Finally, the same loading protocol was applied in the third cycle
639 following a further rotation of the differential to the 3-direction (green lines in Figure
640 7a), such that Ω_3 now corresponds to the axial component and Ω_1 and Ω_2 correspond
641 to the transverse components.

642 Figure 7b (black line) demonstrates that the onset of AE occurs at
643 approximately the same level of stress in each cycle; between 35 and 45 MPa. This
644 suggests that no manifestation of the Kaiser effect is observed in this test. However,
645 while the onset of AE occurs at approximately the same level of stress in each cycle,
646 it is significant to note that the amount of AE (plotted as cumulative AE hits in Figure
647 7b), decreases with each new sequential loading cycle. The model-simulated
648 accumulated damage curves recreate this tendency for the three loading cycles
649 (colored lines for each cycle in Fig. 7b). The simulation commences with an initially
650 small amount of isotropic damage, $\Omega^2 = \Omega_1^2 + \Omega_2^2 + \Omega_3^2 \approx 1.5\%$. During the first
651 loading cycle, the two transverse damage components, Ω_2 and Ω_3 , grow at the same
652 rate (as shown in Fig. 7c) leading to a damage increase, $\Omega^2 \geq 4\%$. However, during
653 the second and third loading cycles, the two transverse damage components (Ω_1 and
654 Ω_3 in cycle 2, Ω_1 and Ω_2 in cycle 3) grow at different rates. This is entirely as
655 expected, because the rock is no longer isotropic following the anisotropic damage
656 formed during the first cycle. The overall amount of damage increase during cycles 2
657 and 3 together add only slightly more than 1% to the total amount of damage. Finally,
658 the simulated stress-strain curves from the model are shown in Figure 8 (colored
659 lines) and provide a very good fit to the experimental curves (black lines). However, it
660 should be noted that the fitting procedure is done by eye and the search for model

661 parameters is therefore non-unique. There are then several uncertainties and trade-offs
 662 between the model parameters. For example, the stresses and strains prior to the onset
 663 of damage accumulation (AE output) for the low damage starting material are
 664 expected to be linearly elastic, and hence permit the calculation of elastic moduli.
 665 However, for this specific test, on a relatively porous sandstone, a significant amount
 666 of the total strain is irreversible and associated with material compaction (eq. 23, 24).
 667 Panteleev et al., (2021), who processed results for similar experiments on samples of
 668 Darley Dale sandstone, also noted this feature. We therefore use Lamé moduli $\lambda_0 =$
 669 5 GPa , $\mu_0 = 7 \text{ GPa}$ for the starting material. These values are in agreement with the
 670 measured seismic wave velocities of $V_p \sim 3.4 \text{ km/s}$ and $V_s \sim 2.1 \text{ km/s}$ reported by
 671 Browning *et al.* (2017) for this material. Since all tests were performed under the
 672 same hydrostatic confining pressure of 4 MPa, it is impossible to constrain the shape
 673 of the yield curve. With this uncertainty we use the power index, $N=1$, in (20); $\xi_0 =$
 674 -0.1 and $Ch=2 \cdot 10^{-7}$ in (22) which corresponds to about 15 MPa of cohesive force
 675 (Appendix B). Scaling the range of the damage value, Ω^2 , from zero to 100% for total
 676 failure (e.g., Lyakhovsky *et al.* 1997) gives $\gamma = 10 \text{ GPa}$. Simultaneous fitting of the
 677 stresses and strains in the three cycles and the noted similarity between the
 678 accumulated damage and the cumulative acoustic emission output led to the following
 679 parameters: compaction equilibrium, $B_0=0.5\%$, $B_1=20 \text{ MPa}$, $B_2=15 \text{ MPa}$ in (23), $C_v=$
 680 $3 \cdot 10^{-2} (\text{MPa s})^{-1}$ in (25), rate coefficient in (24) $A=10^{-5} (\text{MPa s})^{-1}$, together with $L=7 \text{ s}^{-1}$,
 681 $K \times D_1=710^3 \text{ s}^{-1}$, and $D_2=30$ in (22). The relatively high D_2 value corresponds to the
 682 high sensitivity of the yield curve to the change in damage required to reproduce the
 683 Kaiser effect (e.g., Gajst *et al.* 2020).

684

685

686

687 3.3 Cyclic sequential rotational conventional triaxial (CSRCT) loading

688 CSRCT loading was performed using the same starting material and with the
 689 same initial pre-loading of 4 MPa along each of the three sample axes. In the first
 690 phase of this test, the differential stress was initially raised to 75 MPa along the 1-
 691 direction (i.e., the direction of the Ω_1 damage component) using the same loading rate
 692 as for the SRCT test, and the sample was then unloaded instantaneously (Figure 9a).
 693 The sample was subsequently reloaded in the same orientation but to a higher

694 differential stress of 80 MPa (blue lines in Fig. 9a), before again being unloaded
695 instantaneously. Then, during the second phase, the differential stress was rotated and
696 applied in the 2-direction (i.e., the direction of the Ω_2 damage component) to a level
697 of 65 MPa in the first cycle and 80 MPa in the second cycle (red lines in Fig. 9a).
698 Finally, in the third phase, the differential stress was rotated again and applied in the
699 3-direction (i.e., the direction of the Ω_3 damage component). During this phase, the
700 sample was loaded to 55 MPa in the first cycle and 80 MPa in the second cycle (green
701 lines in Fig. 9a). Once again, we see that the model simulated damage curves obtained
702 using the same material properties as for the SRCT test, are very similar in form to the
703 cumulative AE output recorded during the experiment (Fig. 9b). Similarly to the
704 SRCT test, the two transverse damage components, Ω_2 and Ω_3 , grow at the same rate
705 (Fig. 9c) during the loading in the 1-direction, as expected for the isotropic sample.
706 However, during the loading in 2- and 3-directions the rock is no longer isotropic
707 following the anisotropic damage formed during the loading in 1-direction. After
708 stress was rotated, the two transverse damage components Ω_1 and Ω_3 , and then, after
709 another rotation, components Ω_1 and Ω_2 grow at different rates. The modeled onsets
710 of damage for each loading cycle, are in very good agreement with the measured
711 onsets of AE output for each cycle during the experiment. The colored markers (stars)
712 on the loading curves of Figure 9a indicate the value of stress at the onset of AE, and
713 hence the onset of new damage. We note that both the experimental data and the
714 model simulations exhibit a distinct Kaiser effect when the samples were reloaded in
715 the same direction, but no such effect when the differential stress was rotated. Again,
716 this suggests that the onset of damage is, at best, only weakly affected by damage
717 accumulated during earlier phases when the differential stress was applied in different
718 orientations.

719

720 **4. Discussion and Concluding Remarks**

721 Here we build on the previous non-linear anisotropic damage rheology model
722 (Panteleev *et al.* 2021) by presenting a newly developed poroelastic rheological model
723 which accounts for both coupled anisotropic damage and porosity evolution. The new
724 model shares the main features of our previously developed anisotropic damage and
725 scalar poroelastic damage models, including the ability to simulate the entire yield
726 curve through a single formulation. In the new model, the yield condition is defined in

727 terms of invariants of the strain tensor, and so the new formulation operates with
728 directional yield conditions (different values for each principal direction) depending
729 on the damage tensor and triaxial loading conditions. This allows us to discern
730 evolving yield conditions for each principal stress direction and fit the measured
731 amounts of accumulated damage from previous loading cycles. Coupling between
732 anisotropic damage and anisotropic compaction along with the damage-dependent
733 yield condition produces a reasonable fit to the experimentally obtained stress-strain
734 curves. Furthermore, the simulated time-dependent cumulative damage is well
735 correlated with experimentally observed acoustic emissions during cyclic loading in
736 different directions. As such, we are able to recreate many of the features of the
737 experimentally observed directional 3D Kaiser ‘damage memory’ effect.

738 The main finding from this formulation is that each independent direction will
739 possess its own yield envelope. The state of each envelope then depends on the
740 direction, magnitude and history of loading. Yield in the transverse components are
741 similar in shape but not necessarily in value to the conventional scalar criterion and
742 the yield criterion in the axial direction has an entirely different shape. As such,
743 damage accumulation in the axial direction is possible only under high volumetric
744 stresses.

745 These results are important in nature since rocks in complex tectonic
746 environments, such as volcanoes or active fault zones, experience stresses that evolve
747 both spatially and temporally and experience not only cyclic loading and unloading,
748 but also rotation and/or reorientation of stresses. The resulting crack natural
749 distributions will then form sequentially and may be highly anisotropic. Thus, the
750 tectonic history of the crust as recorded in deformed rocks may include evidence for
751 complex stress paths, encompassing different magnitudes and orientations. Geodetic
752 and seismic data from periods of inflation and deflation at Krafla volcano in Iceland
753 demonstrates that the rate of seismicity increases only after the amount of inflation in
754 a previous cycle has been reached or exceeded (Heimisson *et al.* 2015). These, and
755 similar observations, point to a potential crustal scale Kaiser effect but the
756 directionality of fracture populations formed has received less attention. The data
757 from Heimisson *et al.* (2015) were interpreted assuming a conventional Kaiser effect
758 such that the orientation of loading and unloading of all of the episodes were assumed
759 to stress the crustal rocks in the same direction. It is noted that, prior to many cyclic
760 inflation episodes at active volcanoes, the level of inflation has often been at a higher

761 level the than the first cycle of a new episode. In such circumstances, the first cycle
762 should not produce seismicity, as the previous level of inflation has not been reached
763 or exceeded. This has traditionally been explained as related to crack healing
764 processes between cycles (Kim *et al.* 2014) but it can also be explained by our new
765 three-dimensional model. As the orientation and magnitude of crustal stresses can
766 vary enormously between different rock layers (Gudmundsson 2011) the new
767 inflation episode may have loaded the rocks under a slightly different axis and hence
768 triggered seismicity at lower stresses, similar to as demonstrated in Figure 9a.

769 Recent experimental results have shown that damage, under true triaxial
770 loading, is a distinctly directional phenomenon, and these results have also revealed a
771 3D directionally dependent Kaiser ‘damage memory’ effect. The developments of this
772 study provide an internally consistent framework for simulating evolving crustal rock
773 damage under repeated cycles of stress in complex tectonic environments where
774 stresses may evolve both spatially and temporally.

775

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785

786 **Data Availability**

787 Experimental data may be obtained from J.B. (e-mail: jbrowning@ing.puc.cl).

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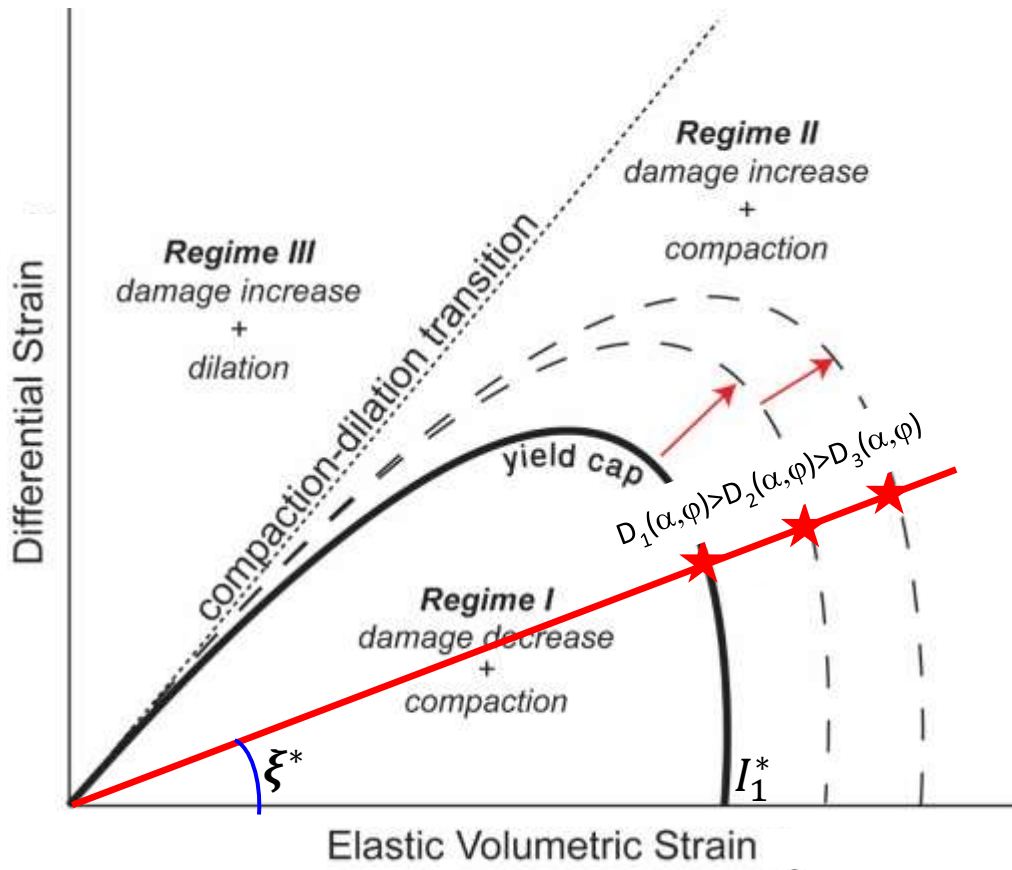
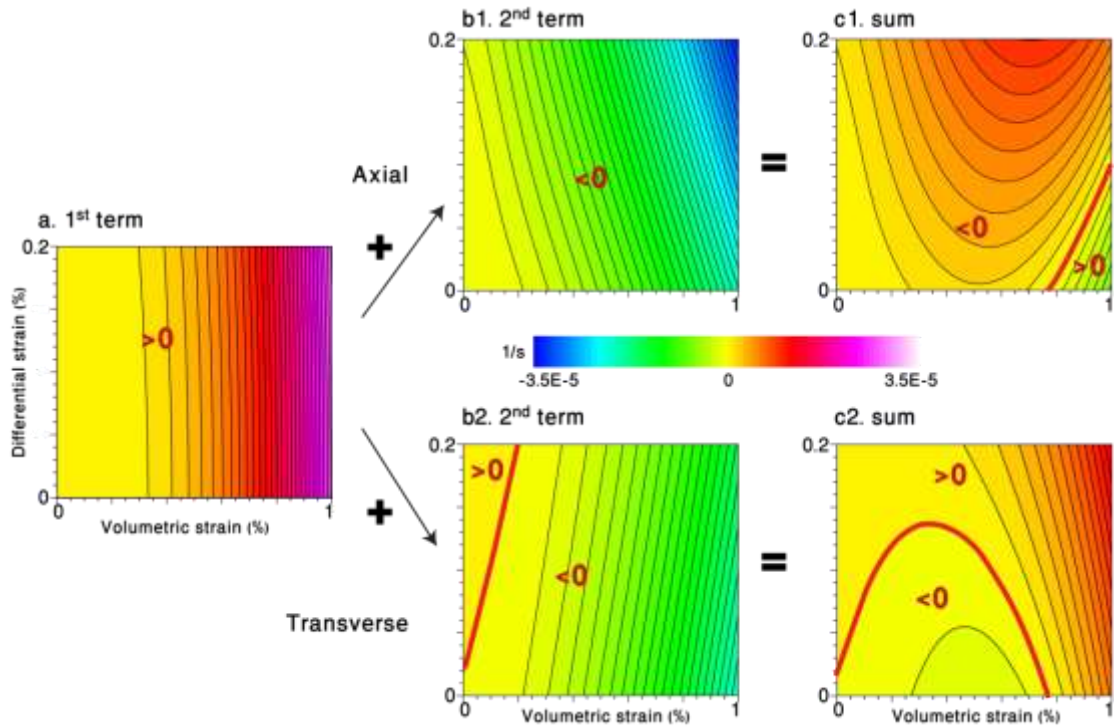


Fig. 1. Three different deformational regimes and growing yield surface neglecting the cohesive force

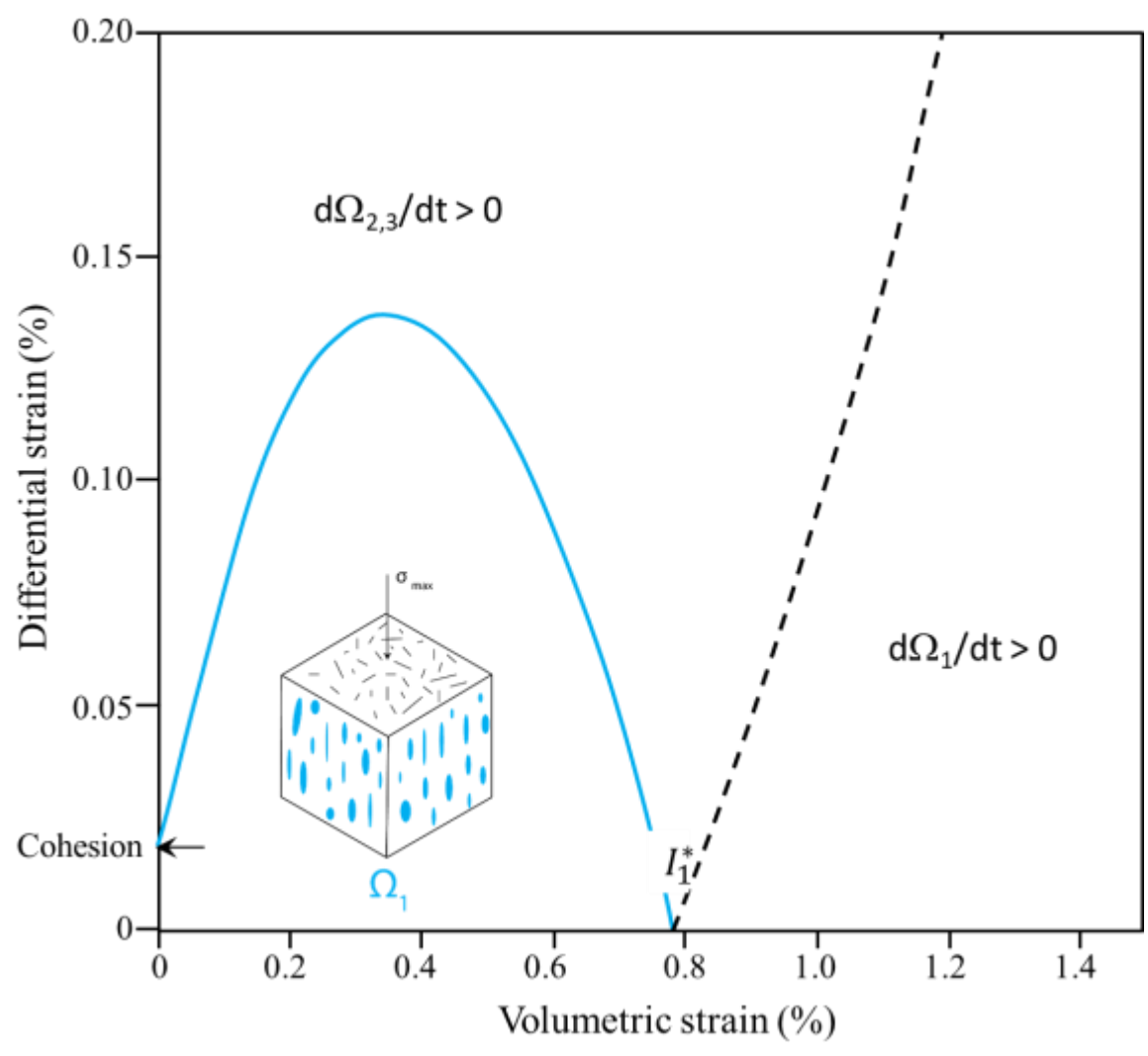
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Fig. 2 Damage kinetics for initially isotropic damage ($\Omega_1 = \Omega_2 = \Omega_3$) under conventional triaxial loading. (a) The 1st term is the same for all damage components. The 2nd term for axial component (b1) significantly differs from the values for the transversal components (b2). The damage kinetics in the axial direction (c1) differs from the transversal components (c2) which is similar in shape but not necessarily in values to the conventional scalar criterion.

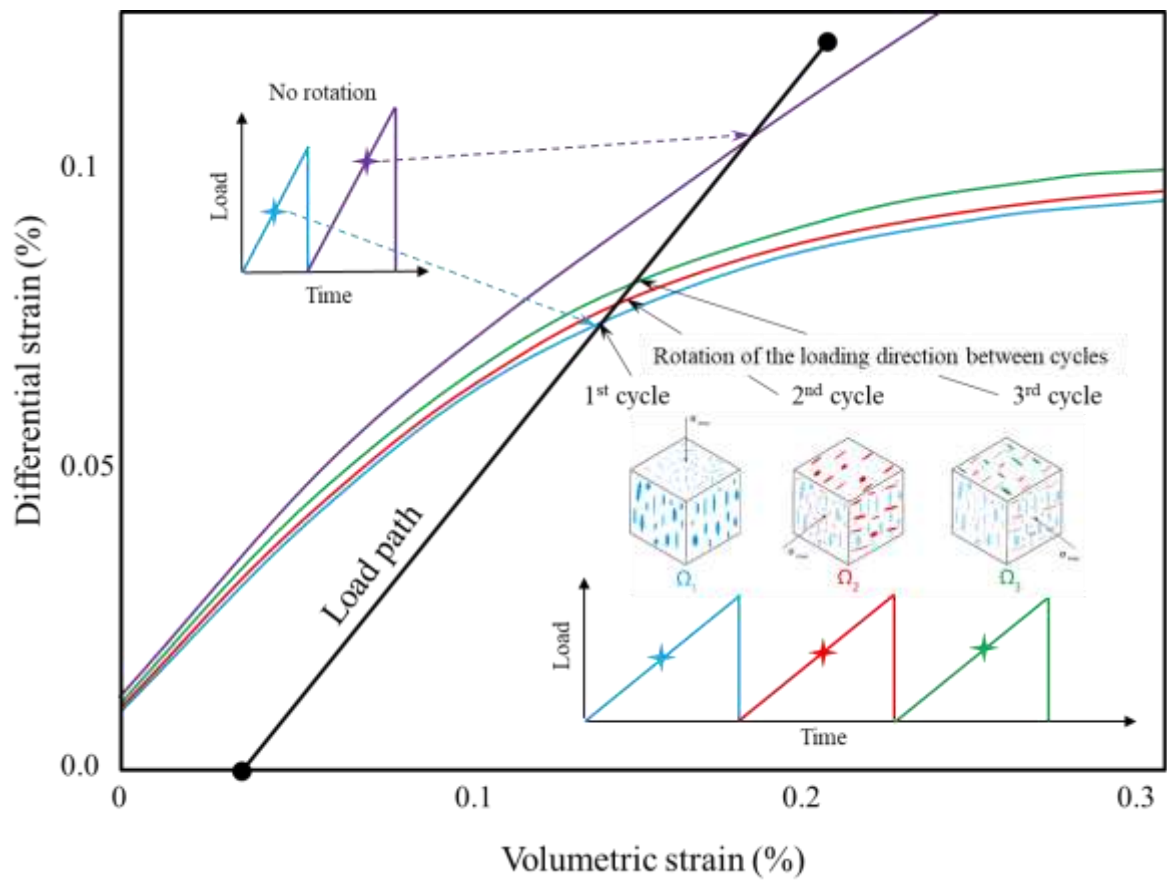
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Fig. 3. Yield envelope for initially isotropic material.

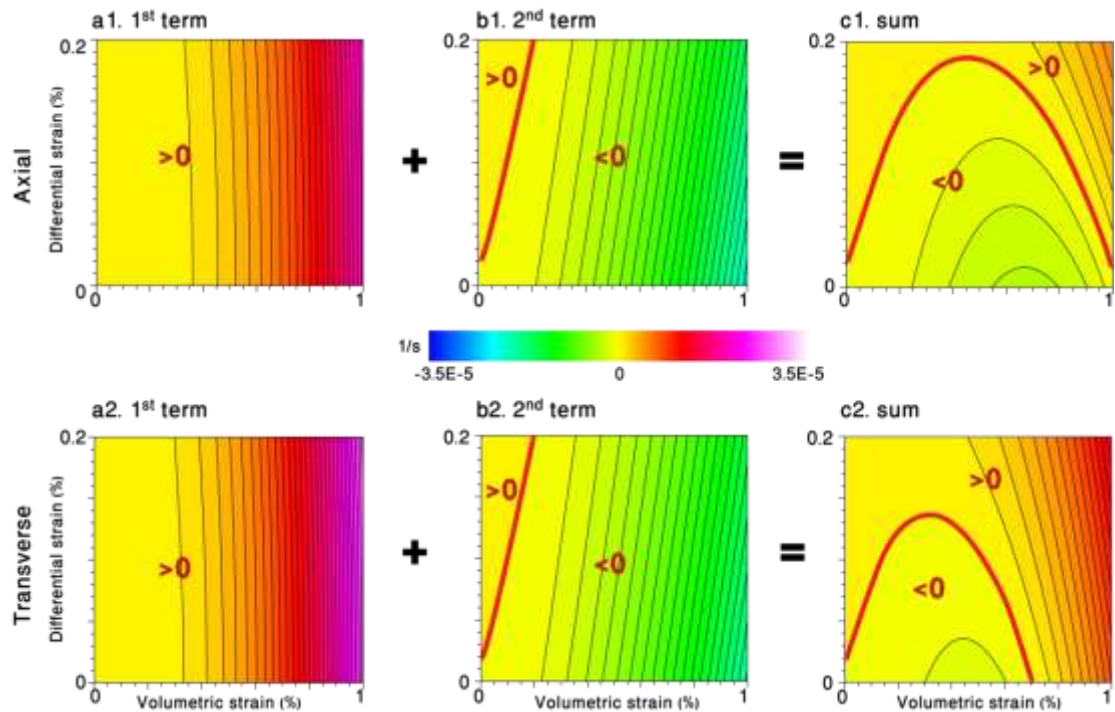
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Fig. 4. Evolving yield envelope for loading cycles in which the maximum load is applied in the same direction cyclically and in which the maximum loading direction is sequentially rotated with respect to the sample.

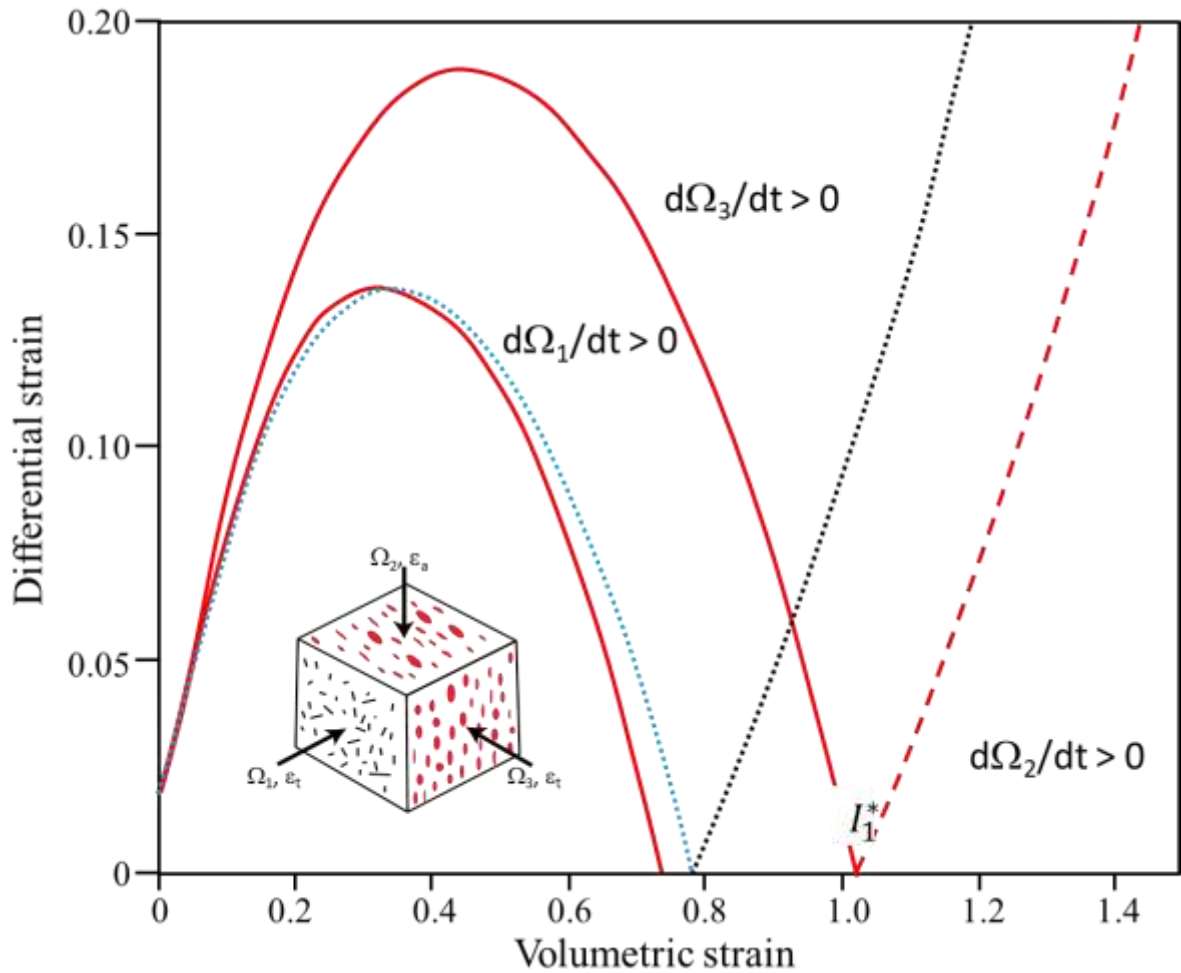
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Fig. 5. Kinetics of the transverse damage components (Ω_1 , Ω_3) for initially anisotropic initial damage ($\Omega_1 < \Omega_2 = \Omega_3$). Upper row represents the values for the larger (Ω_3) component and the lower row for the smaller (Ω_1) component. The pattern for both components, 1st (a1, a2) and 2nd- (b1, b2) terms are similar, but their different values leads to significantly different size of the yield envelope (c1, c2).

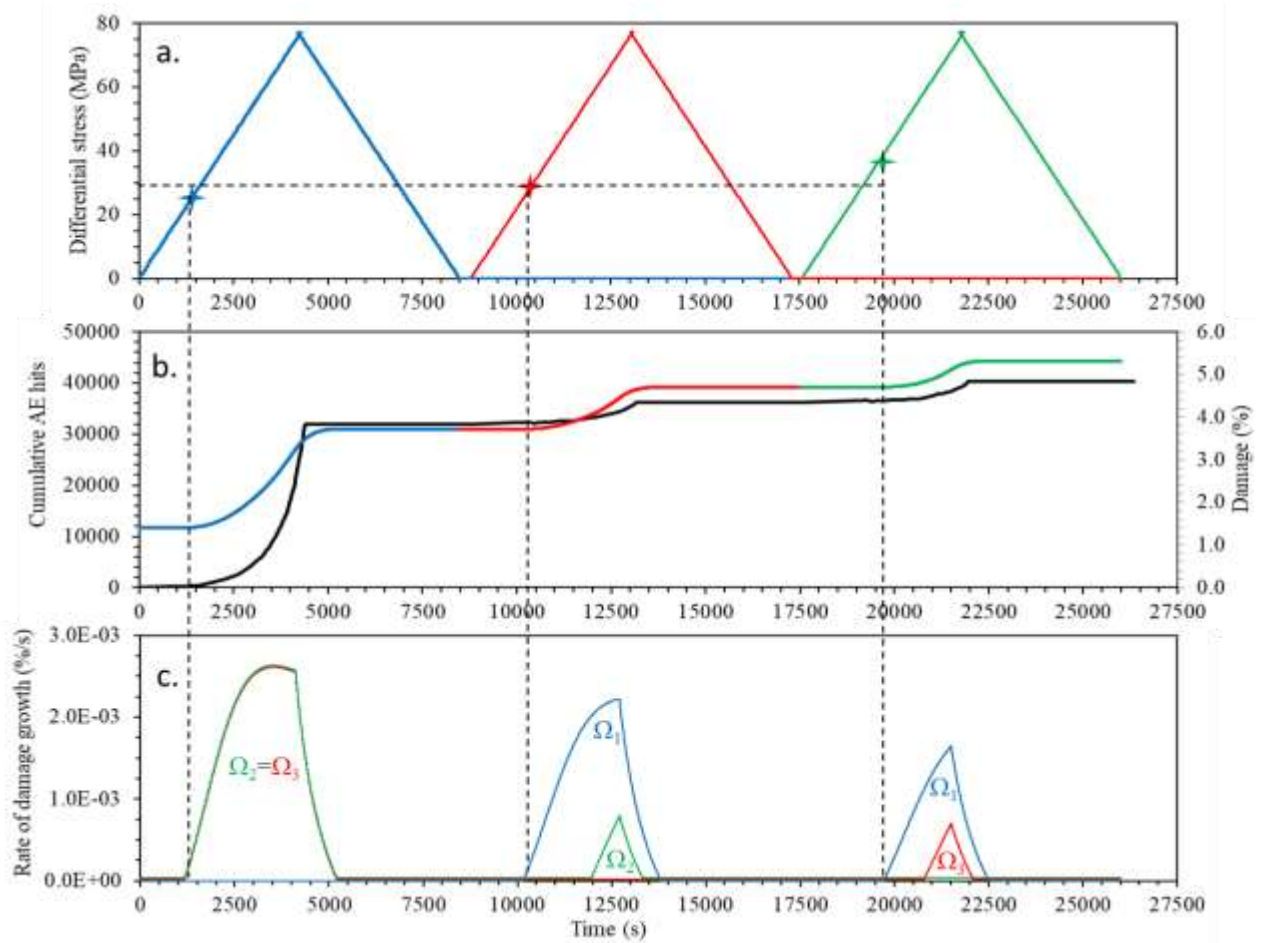
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Fig. 6. Yield envelope for initially anisotropic material ($\Omega_1 < \Omega_2 = \Omega_3$). The envelope for the smallest transverse (Ω_1) component (red line) is almost the same as for the initially isotropic material (dotted blue line). The envelope for larger transversal component (Ω_3) is significantly larger than for the component Ω_1 . The axial damage component (Ω_2) is accumulated only under high confining conditions.

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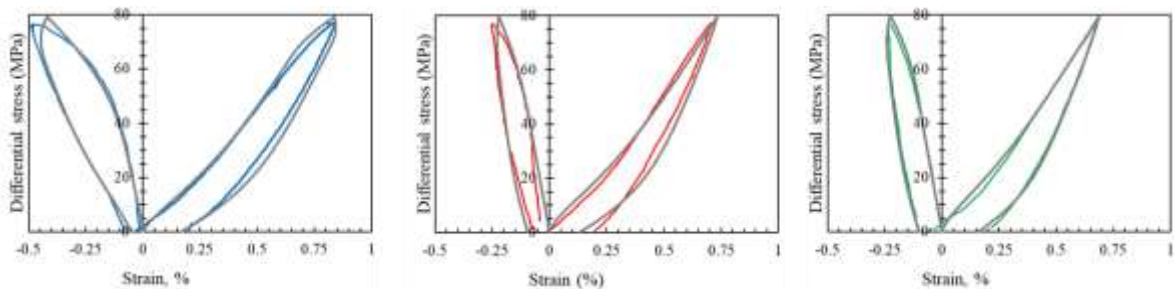


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1110 Fig. 7. Comparison between experimental results and modelling. (a) Sequential
1111 rotational conventional triaxial loading; markers show the onset of the observed AE
1112 (black line in (b)). Colored line in (b) shows the simulated damage accumulated
1113 during the loading. (c) rate of damage accumulation.

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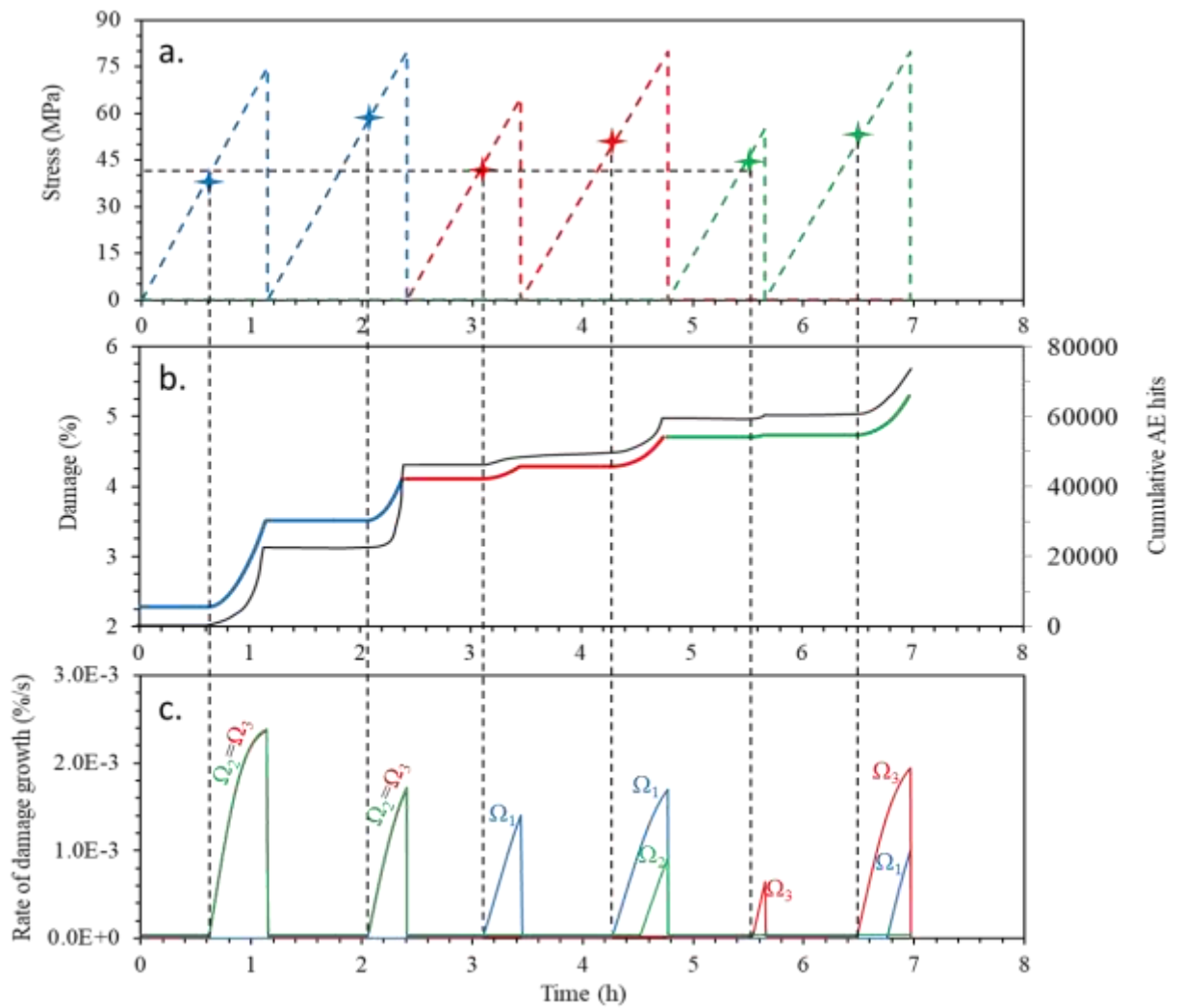
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1117 Fig. 8. Comparison between observed (colored lines) and simulated (grey lines)
1118 stress-strains for three loading cycles (blue, red, green).

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Fig. 9. Comparison between experimental results and modelling. (a) Cyclic sequential rotational conventional triaxial loading; the markers show the onset of the observed AE (black line in (b)). Colored line in (b) shows the simulated damage accumulated during the loading. (c) rate of damage accumulation.

1135 **Figure Captions**

1136 Figure 1. Three different deformational regimes and growing yield surface
1137 neglecting the cohesive force.

1138 Figure 2. Damage kinetics for initially isotropic damage ($\Omega_1 = \Omega_2 = \Omega_3$) under
1139 conventional triaxial loading. (a) The 1st term is the same for all damage components.
1140 The 2nd term for axial component (b1) significantly differs from the values for the
1141 transversal components (b2). The damage kinetics in the axial direction (c1) differs
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1143 values to the conventional scalar criterion.

1144 Figure 3. Yield envelope for initially isotropic material.

1145 Figure 4. Evolving yield envelope for loading cycles in which the maximum
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1148 Figure 5. Kinetics of the transverse damage components (Ω_1, Ω_3) for initially
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1150 larger (Ω_3) component and the lower row for the smaller (Ω_1) component. The pattern
1151 for both components, 1st (a1, a2) and 2nd (b1, b2) terms are similar, but their different
1152 values leads to significantly different size of the yield envelope (c1, c2).

1153 Figure 6. Yield envelope for initially anisotropic material ($\Omega_1 < \Omega_2 = \Omega_3$). The
1154 envelope for the smallest transverse (Ω_1) component (red line) is almost the same as
1155 for the initially isotropic material (dotted blue line). The envelope for larger
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1158 Figure 7. Comparison between experimental results and modelling. (a)
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1162 Figure 8. Comparison between observed (colored lines) and simulated (grey
1163 lines) stress-strains for three loading cycles (blue, red, green).

1164 Figure 9. Comparison between experimental results and modelling. (a) Cyclic
1165 sequential rotational conventional triaxial loading; the markers show the onset of the
1166 observed AE (black line in (b)). Colored line in (b) shows the simulated damage
1167 accumulated during the loading. (c) rate of damage accumulation.

1168 **Appendix A**

1169 Thermodynamic relations

1170 The presented formulation is based on balance equation of the irreversible
 1171 thermodynamics of the continuum media with internal variables (Malvern 1969,
 1172 Coussy 1995). We consider the system with density of the free energy, F , to be a
 1173 function of:

$$1174 \quad F = F(T, \varepsilon_{ij}, \Omega_{ij}, \psi_{ij}, \zeta) \quad (1)$$

1175 where T – temperature, ε_{ij} – elastic strain tensor, ψ_{ij} – compaction tensor, Ω_{ij} – the
 1176 damage tensor, and ζ is the change in volume fluid content defined by Biot (1941)
 1177 (see also Detournay & Cheng 1993). Since each variable can vary independently of
 1178 the other variables, Gibbs relation can be written as (Gibbs 1961):

$$1179 \quad dF = -SdT + \frac{\partial F}{\partial \varepsilon_{ij}} d\varepsilon_{ij} + \frac{\partial F}{\partial \Omega_{ij}} d\Omega_{ij} + \frac{\partial F}{\partial \psi_{ij}} d\psi_{ij} + \frac{\partial F}{\partial \zeta} d\zeta \quad (2)$$

1180 where $S = -\frac{\partial F}{\partial T}$ is entropy density (Einstein's summation convention is assumed).

1181 The balance equation for the density of the internal energy, U , includes three
 1182 source terms associated with shear heating (stress, σ_{ij} , times strain rate or time
 1183 derivative of the total strain rate tensor, e_{ij}), energy dissipation due to heat flux, Q_i ,
 1184 and the advective flux due to fluid flow (eq. 71 from Coussy *et al.* 1998):

$$1185 \quad \frac{dU}{dt} = \frac{d}{dt}(F + TS) = \sigma_{ij}e_{ij} - \nabla_i Q_i - \nabla_i [(h_f + Ts_f)q_i] \quad (3)$$

1186 where q_i is fluid flux, h_f is the enthalpy or Gibbs potential of the fluid, which is a
 1187 function of the fluid pressure, p_f , and temperature; s_f is the entropy of the fluid.
 1188 Similarly, the entropy balance equation includes positive entropy production, Γ ,
 1189 divergence of the heat flux, and similar advective term due to fluid flow (eq. 72 from
 1190 Coussy *et al.* 1998):

$$1191 \quad \frac{dS}{dt} = \frac{\Gamma}{T} - \nabla_i \left(\frac{Q_i}{T} \right) - \nabla_i (q_i s_f) \quad (4)$$

1192 The stress tensor and the fluid pressure are defined as (Malvern 1969, Coussy 1995):

$$1193 \quad \sigma_{ij} = \frac{\partial F}{\partial \varepsilon_{ij}} \quad (5)$$

$$1194 \quad p_f = \frac{\partial F}{\partial \zeta} \quad (6)$$

1195 Fluid mass conservation is:

$$1196 \quad \frac{d\zeta}{dt} = -\nabla_i (q_i) \quad (7)$$

1197 Combining (2, 3, 4) and using (5, 6, 7) we get a final expression for the local entropy
 1198 production in a form:

$$1199 \quad \Gamma = -\frac{Q_i}{T} \nabla_i T - q_i \nabla_i p_f + \sigma_{ij} \left(e_{ij} - \frac{d\varepsilon_{ij}}{dt} \right) - \frac{\partial F}{\partial \Omega_{ij}} \frac{d\Omega_{ij}}{dt} - \frac{\partial F}{\partial \psi_{ij}} \frac{d\psi_{ij}}{dt} \quad (8)$$

1200 First and second terms represent dissipation associated with heat transport and fluid
 1201 flow. The Fourier and Darcy laws establish linear relations between temperature and
 1202 fluid pressure gradients with corresponding fluxes:

$$1203 \quad Q_i = -K_{ij}^T \nabla_j T \quad (9a)$$

1204 and

$$1205 \quad q_i = -K_{ij}^F \nabla_j p_f \quad (9b)$$

1206 where K_{ij}^T and K_{ij}^F are positively defined thermal conductivity and permeability
 1207 tensors.

1208 The elastic strain tensor is the difference between total, g_{ij} , and irreversible
 1209 compaction, ψ_{ij} , tensors:

$$1210 \quad \varepsilon_{ij} = g_{ij} - \psi_{ij} \quad (10)$$

1211 Taking the time derivative of (10) and substituting the total strain rate tensor, $e_{ij} =$
 1212 dg_{ij}/dt , into (8), the part of the total dissipation, Γ_{DC} , associated with evolving
 1213 damage and compaction (three last term of Eq. 8) is expressed as a sum of two terms
 1214 proportional to the damage and compaction rates:

$$1215 \quad \Gamma_{DC} = -\frac{\partial F}{\partial \Omega_{ij}} \frac{d\Omega_{ij}}{dt} + \left(\sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \right) \frac{d\psi_{ij}}{dt} > 0 \quad (11)$$

1216 For small deviations from equilibrium, the entropy production or dissipation potential
 1217 may be approximated as a quadratic function of the rate of the internal variables. In
 1218 this case constitutive equations give the thermodynamic fluxes as a linear function of
 1219 the thermodynamic forces (Malvern 1969, DeGroot & Mazur 2013) also known as the
 1220 Onsager's (1931) relations:

$$1221 \quad \frac{d\psi_{ij}}{dt} = C_{ijnm}^{\psi\psi} \left(\sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \right) - C_{ijnm}^{\psi\Omega} \frac{\partial F}{\partial \Omega_{ij}} \quad (12a)$$

$$1222 \quad \frac{d\Omega_{nm}}{dt} = C_{ijnm}^{\Omega\psi} \left(\sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} \right) - C_{ijnm}^{\Omega\Omega} \frac{\partial F}{\partial \Omega_{ij}} \quad (12b)$$

1223 These phenomenological kinetic equations guarantee the non-negative entropy
 1224 production if the cells of the matrix of the kinetic coefficients

$$1225 \quad C_{ijnm} = \begin{vmatrix} C_{ijnm}^{\psi\psi} & C_{ijnm}^{\psi\Omega} \\ C_{ijnm}^{\Omega\psi} & C_{ijnm}^{\Omega\Omega} \end{vmatrix} \quad (13)$$

1226 meet several conditions (Malvern 1969, DeGroot & Mazur 2013). Matrixes of the
 1227 diagonal cells ($C_{ijnm}^{\psi\psi}$, $C_{ijnm}^{\Omega\Omega}$) must be positively defined. Off-diagonal terms are
 1228 usually taken to be either symmetric or antisymmetric. Following poroelastic damage
 1229 model of Hamiel *et al.* (2004b) and Lyakhovsky *et al.* (2015) we adopt antisymmetric
 1230 structure ($C_{ijnm}^{\Omega\psi} = -C_{ijnm}^{\psi\Omega}$) of the kinetic matrix (13). These conditions assure
 1231 positive dissipation (11).

1232 Further model formulation, demonstrating its main features as well as its
 1233 calibration, verification, and application to specific problems require definition of the
 1234 energy function (1) and kinetic coefficients (13). In the Appendix B we briefly discuss
 1235 the isotropic, scalar formulation, which instead of the tensorial damage-compaction
 1236 variables operates with the scalar damage, α , equal to the squared value, $\alpha = \Omega^2$,
 1237 derived by Panteleev *et al.* (2021), and porosity, φ . Scalar Ω and φ variables are
 1238 connected with the tensor variables as:

$$1239 \quad \Omega_{ij} = \Omega \delta_{ij} \quad (14a)$$

$$1240 \quad \Psi_{ij} = \varphi \delta_{ij} \quad (14b)$$

1241 The complete anisotropic formulation is presented in the Appendix C.

1242

1243 **Appendix B**

1244 Scalar poroelastic damage model

1245 Following Biot's theory of poroelasticity (Biot 1941, 1956) the free energy of
 1246 a poro-elastic medium, F , is a sum of the elastic energy and the poroelastic coupling
 1247 term of the saturated medium with the Biot modulus, M , and the Biot coefficient for
 1248 porous media, β :

$$1249 \quad F = \frac{\lambda(\alpha, \varphi)}{2} I_1^2 + \mu(\alpha, \varphi) I_2 - \gamma(\alpha, \varphi) I_1 \sqrt{I_2} + \frac{1}{2} M [\beta I_1 - \zeta + \varphi]^2 + Ch \gamma_m \alpha$$

1250 (15)

1251 In the literature discussing Biot poroelasticity, this coefficient is often noted as
 1252 α . To avoid a duplicate notation we use α for the scalar damage and change the
 1253 notation for the Biot coefficient to β .

1254 Since the target model is isotropic, the energy function may depend only on
 1255 invariants of the elastic strain tensor, ε_{ij} ($I_1 = \varepsilon_{ii}$, $I_2 = \varepsilon_{ij}\varepsilon_{ij}$). Following (Hamiel, *et*
 1256 *al.* 2004b, Lyakhovsky *et al.* 2015) the elastic energy for nonlinear damaged media
 1257 includes two Hookean terms with the Lamé drained moduli λ , μ , and an additional

1258 non-linear term with strain coupling modulus γ . Value of this additional modulus
 1259 varies from zero for damage free material with $\alpha = 0$ to $\gamma = \gamma_m$ at material failure
 1260 with $\alpha = 1$ (Lyakhovsky *et al.* 1997). Following Gajst (2020), we introduce the
 1261 damage-dependent term with the non-dimensional coefficient Ch . Below we will
 1262 show that this term allows accounting for the cohesive force which is important for
 1263 loading conditions with low confining pressures.

1264 Differentiation of the poroelastic energy (15) according to the constitutive
 1265 relations (5, 6), the stress tensor, σ_{ij} , and fluid pressure, p_f , are:

$$1266 \quad \sigma_{ij} = \frac{\partial F}{\partial \varepsilon_{ij}} = \left(\lambda - \frac{\gamma}{\xi} \right) I_1 \delta_{ij} + (2\mu - \gamma\xi) \varepsilon_{ij} + \beta M (\beta I_1 - \zeta + \varphi) \quad (16)$$

$$1267 \quad p_f = \frac{\partial F}{\partial \zeta} = M (-\beta I_1 + \zeta - \varphi) \quad (17)$$

1268 where $\xi = I_1/\sqrt{I_2}$ is the strain invariant ratio changing from $\xi = -\sqrt{3}$ for isotropic
 1269 compaction to $\xi = \sqrt{3}$ for isotropic dilation. Using the assumption that $\lambda = Const.$
 1270 and both μ, γ linearly depend on the damage (Agnon & Lyakhovsky 1995), the
 1271 derivatives of the energy function (15) are the effective stress:

$$1272 \quad \sigma_{ij} - \frac{\partial F}{\partial \varphi} \delta_{ij} = \sigma_{ij}^{eff} \quad (18)$$

1273 and damage induced energy change:

$$1274 \quad \frac{\partial F}{\partial \alpha} = \gamma_m (-I_2 (\xi - \xi_0) + Ch) \quad (19)$$

1275 where ξ_0 is the critical value of the strain invariant ratio corresponding to the
 1276 conditions of the Coulomb failure criteria (see Lyakhovsky *et al.* 1997 for details).
 1277 Substituting the derivatives (18, 19) into kinetic equations (12) and accounting to the
 1278 scalar nature of the damage, α , and porosity, φ , variables (14), and using the mean
 1279 stress ($\sigma_m = -\sigma_{kk}/3$), the coupled kinetic equations are reduced to (see Hamiel *et al.*
 1280 2004b, Lyakhovsky *et al.* 2015, and references therein):

$$1281 \quad \frac{d\varphi}{dt} = -A \sigma_m - D [I_2 (\xi - \xi_0) - Ch] \quad (20)$$

$$1282 \quad \frac{d\alpha}{dt} = D \sigma_m + C_d [I_2 (\xi - \xi_0) - Ch] \quad (21)$$

1283 Hamiel *et al.* (2004b) suggested that the coupling coefficients D is a power-law
 1284 expression of the effective pressure, $D \sim \sigma_m^N$. They demonstrated that the transition
 1285 from positive to negative values of the slope of the yield curve (yield cap) is a general
 1286 feature of the model. Here we modify the previous formulation using strain invariants

1287 I_1 , $\sqrt{I_2}$ and $\sigma_m = -K I_1$ (K is a bulk modulus); and also following Gajst *et al.* (2020)
 1288 D -value exponentially decreases with damage:

$$1289 \quad D = C_d D_1 e^{-D_2 \alpha} (-I_1)^N \sqrt{I_2}/K \quad (22)$$

1290 $N > 0$, D_1 and D_2 values define the shape of the yield envelope. The form (22) is
 1291 defined only for compaction ($I_1 \leq 0$); D_1 is either constant or function of the porosity
 1292 constrained using strain-defined yielding envelope. With this assumption, the kinetic
 1293 for damage accumulation (21) is:

$$1294 \quad \frac{d\alpha}{dt} = C_d [D_1 e^{-D_2 \alpha} (-I_1)^{N+1} \sqrt{I_2} + I_2 (\xi - \xi_0) - Ch] \quad (23)$$

1295 This form of the damage kinetic equation predicts the onset of damage accumulation
 1296 ($\frac{d\alpha}{dt} = 0$) for the intact material ($\alpha = 0$), or the yield condition:

$$1297 \quad D_1 (-I_1)^{N+1} \sqrt{I_2} + I_2 (\xi - \xi_0) - Ch = 0 \quad (24)$$

1298 Under shear load needed to overcome the cohesive force ($\tau_h = 2\mu_0 \varepsilon_h$) with zero
 1299 volumetric strain ($I_1 = 0$, $\xi = 0$, and $I_2 = 2\varepsilon_h^2$), the coefficient Ch is equal to:

$$1300 \quad Ch = -\xi_0 \frac{\tau_h^2}{2\mu_0^2} \quad (25a)$$

1301 Under large hydrostatic volumetric strains, neglecting the cohesive force, the onset of
 1302 damage occurs at:

$$1303 \quad I_1^* = -\left(\frac{\sqrt{3} + \xi_0}{\sqrt{3} D_1}\right)^{1/N} \quad (25b)$$

1304 The equations (25a, b) define two endmember yield values. The entire yield shape
 1305 may be calculated for any given volumetric strain between zero and I_1^* by solving (24)
 1306 as a quadratic equation for $\sqrt{I_2}$:

$$1307 \quad \sqrt{I_2} = \frac{D_1 (-I_1)^{N+1} + \sqrt{(D_1 (-I_1)^{N+1})^2 - 4 \xi_0 Ch}}{2 \xi_0} \quad (25c)$$

1308 The positive radical sign provides correct solution for $Ch = 0$. Figures B1-B3 show
 1309 sensitivity of the shape and size of the yield envelope to the change of the model
 1310 parameters D_1 , N , and Ch values calculated using (25c). For $N=1$ and non-cohesive
 1311 material ($Ch=0$) the envelope increases with decrease D_1 value keeping roughly self-
 1312 similar shape (Fig. B1). Changing power index value N affects the shape of the
 1313 envelope (Fig. B2). In these cases, D_1 values were rescaled to get the same volumetric
 1314 strain $I_1^* = -3.5\%$ according to Eq. (25b). Non-zero Ch value for the cohesive
 1315 material with $N=1$ and $D_1=15$ shifts envelope to larger differential strain values under
 1316 low volumetric strains (Fig. B3). The shown set of surfaces for $Ch = 0-1.0 \cdot 10^{-6}$

1317 corresponds to cohesive force values changing from zero to ~15 MPa for rock with
 1318 shear modulus $\mu_0 = 10$ GPa. For typical cohesive force values of a few MPa, this
 1319 model modification is important under relatively low confining pressures but may be
 1320 neglected when confining pressures are higher (of the order of tens of MPa).

1321

1322 **Appendix C**

1323 Anisotropic poroelastic damage model

1324 The energy function of the anisotropic material with the damage tensor, Ω_{ij} ,
 1325 cannot be formulated only in terms of strain invariants I_1, I_2 of the elastic strain
 1326 tensor, ε_{ij} ($I_1 = \varepsilon_{ii}$, $I_2 = \varepsilon_{ij}\varepsilon_{ij}$) used for the scalar model (eq. 15). Following Murti
 1327 *et al.* (1991) and Zhang & Cai (2010) the energy function should also depend on the
 1328 invariants of the tensor, $\varepsilon_{ij}^{(\Omega)}$, which is the symmetrized product of the elastic strain
 1329 and damage tensors:

$$1330 \quad \varepsilon_{ij}^{(\Omega)} = \frac{1}{2}(\varepsilon_{ik}\Omega_{kj} + \varepsilon_{jk}\Omega_{ki}) \quad (26)$$

1331 The invariants $I_1^{(\Omega)}$ and $I_2^{(\Omega)}$ of this tensor are:

$$1332 \quad I_1^{(\Omega)} = \varepsilon_{ik}\Omega_{kj}\delta_{ij} =$$

$$1333 \quad = \varepsilon_{11}\Omega_{11} + \varepsilon_{22}\Omega_{22} + \varepsilon_{33}\Omega_{33} + 2\varepsilon_{12}\Omega_{12} + 2\varepsilon_{23}\Omega_{23} + 2\varepsilon_{13}\Omega_{13}$$

$$1334 \quad (27)$$

$$1335 \quad I_2^{(\Omega)} = \varepsilon_{ij}^{(\Omega)}\varepsilon_{ij}^{(\Omega)} =$$

$$1336 \quad = (\varepsilon_{11}\Omega_{11} + \varepsilon_{12}\Omega_{12} + \varepsilon_{13}\Omega_{13})^2 +$$

$$1337 \quad + (\varepsilon_{12}\Omega_{12} + \varepsilon_{22}\Omega_{22} + \varepsilon_{23}\Omega_{23})^2 +$$

$$1338 \quad + (\varepsilon_{13}\Omega_{13} + \varepsilon_{23}\Omega_{23} + \varepsilon_{33}\Omega_{33})^2 +$$

$$1339 \quad + (\varepsilon_{11}\Omega_{12} + \varepsilon_{22}\Omega_{12} + \varepsilon_{23}\Omega_{13} + \varepsilon_{13}\Omega_{23} + \varepsilon_{12}(\Omega_{11} + \Omega_{22}))^2 +$$

$$1340 \quad + (\varepsilon_{23}\Omega_{12} + \varepsilon_{11}\Omega_{13} + \varepsilon_{33}\Omega_{13} + \varepsilon_{12}\Omega_{23} + \varepsilon_{13}(\Omega_{11} + \Omega_{33}))^2 +$$

$$1341 \quad + (\varepsilon_{13}\Omega_{12} + \varepsilon_{12}\Omega_{13} + \varepsilon_{22}\Omega_{23} + \varepsilon_{33}\Omega_{23} + \varepsilon_{23}(\Omega_{22} + \Omega_{23}))^2$$

1342

1343 We extend energy function of the Pantelev *et al.* (2021) non-linear anisotropic
 1344 damage rheology model by additional Biot terms similar to (15):

$$1345 \quad F = \left[\frac{\lambda_0}{2} I_1^2 + \mu_0 I_2 \right] + \gamma \left[\xi_0 I_2^{(\Omega)} - I_1^{(\Omega)} \sqrt{I_2^{(\Omega)}} \right] + \frac{1}{2} M [\beta I_1 - \zeta + \psi_{ij} \delta_{ij}]^2 + Ch \frac{\gamma}{2} \Omega_{ij} \Omega_{ij}$$

$$1346 \quad (28)$$

1347 The definitions (26-28) mean that the energy is the second order function of both
 1348 elastic strain and damage tensors. Substituting energy function (28) into (5) and (6)
 1349 the stress tensor is expressed as:

$$1350 \quad \sigma_{ij} = \frac{\partial F}{\partial \varepsilon_{ij}} = \lambda_0 I_1 \delta_{ij} + 2\mu_0 \varepsilon_{ij} - \gamma \frac{\partial I_1^{(\Omega)}}{\partial \varepsilon_{ij}} \sqrt{I_2^{(\Omega)}} + \gamma \left(\xi_0 - \frac{I_1^{(\Omega)}}{2\sqrt{I_2^{(\Omega)}}} \right) \frac{\partial I_2^{(\Omega)}}{\partial \varepsilon_{ij}} +$$

$$1351 \quad + \beta M (\beta I_1 - \zeta + \Psi_{kn} \delta_{kn}) \delta_{ij} \quad (29)$$

1352

1353 where

$$1354 \quad \frac{\partial I_1^{(\Omega)}}{\partial \varepsilon_{lm}} = \Omega_{lm},$$

$$1355 \quad \frac{\partial I_2^{(\Omega)}}{\partial \varepsilon_{lm}} = \frac{1}{2} (\Omega_{li} \varepsilon_{ik} \Omega_{km} + \Omega_{mi} \varepsilon_{ik} \Omega_{kl}) + \frac{1}{2} (\Omega_{lj} \varepsilon_{mk} \Omega_{kj} + \Omega_{mj} \varepsilon_{lk} \Omega_{kj}).$$

1356

1357 The fluid pressure is almost identical to (17)

$$1358 \quad p_f = \frac{\partial F}{\partial \zeta} = M (-\beta I_1 + \zeta - \Psi_{kn} \delta_{kn}) \quad (30)$$

1359 Similarly to (18) the effective stress is defined as:

$$1360 \quad \sigma_{ij} - \frac{\partial F}{\partial \psi_{ij}} = \sigma_{ij}^{eff} \quad (31)$$

1361 In order to evaluate the energy dissipation and write the couples damage-compaction
 1362 kinetic relations (12), the damage induced energy change should be calculated:

$$1363 \quad \frac{\partial F}{\partial \Omega_{ij}} = \gamma \left[\left(\xi_0 - \frac{I_1^{(\Omega)}}{2\sqrt{I_2^{(\Omega)}}} \right) \frac{\partial I_2^{(\Omega)}}{\partial \Omega_{ij}} - \varepsilon_{ij} \sqrt{I_2^{(\Omega)}} + Ch \Omega_{ij} \right] \quad (32)$$

1364

Formulating the tensor kinetic coefficients, we keep in mind that the coupling
 kinetic coefficient $D_{ijkn} = C_{ijnm}^{\Omega\psi}$ of eq. (13) should decrease with damage
 accumulation and expect that this will allow reproducing the directional Kaizer Effect
 observed under true triaxial loading (Browning *et al.* 2018). Similarly to the
 exponential relation (22), we suggest that the matrix D_{ijkn} of the kinetic coefficients
 exponentially decrease with damage accumulation. Extending this idea to the
 complete tensor form and using the same dependency of the coupling term as in the
 scalar model (22), we suggest the following form of the coupling kinetic coefficient
 D_{ijkn} :

$$1373 \quad D_{ijkn} = D_1 \exp[-D_2 \Omega_{ij}] \delta_{kn} (-I_1)^N \sqrt{I_2} \quad (33)$$

1374 where we use the standard definition of the exponent of the tensor \mathbf{X} by means of its
 1375 series representation (Hirsch *et al.* 1974)

$$1376 \quad \exp(\mathbf{X}) = \sum_{m=0}^{\infty} \frac{\mathbf{X}^m}{m!} \quad (34)$$

1377 Note that the principal values of the tensor $\exp(\mathbf{X})$ are equal to the $\exp(X_k)$.

1378 The most conservative assumption to define the components of the matrix
 1379 $C_{ijnm}^{\Omega\Omega}$ in (13) is the absence of the interaction between different components, i.e.,
 1380 $C_{ijnm}^{\Omega\Omega} = L (\delta_{ik} \delta_{jn} + \delta_{in} \delta_{jk})$, and L is proportional to $1/\sqrt{\Omega_{ij}\Omega_{ij}}$ as it was suggested
 1381 by verified by Pantelev *et al.* (2021) using results of true-triaxial rock mechanics
 1382 experiments ignoring effects of compaction. With these kinetic coefficients, the
 1383 equation (12b) for the damage evolution has the same structure as the damage kinetic
 1384 equation (23) of the scalar model.

$$1385 \quad \frac{d\Omega_{ij}}{dt} = D_1 \exp[-D_2 \Omega_{ij}] (-I_1)^N \sqrt{I_2} \sigma_m^{eff} +$$

$$1386 \quad + \frac{L}{\sqrt{\Omega_{kn}\Omega_{kn}}} \left[\varepsilon_{ij} \sqrt{I_2^{(\Omega)}} + \left(\frac{I_1^{(\Omega)}}{2\sqrt{I_2^{(\Omega)}}} - \xi_0 \right) \frac{\partial I_2^{(\Omega)}}{\partial \Omega_{ij}} - Ch \Omega_{ij} \right] \quad (35)$$

1387 The equation for the evolving compaction (12a) includes two terms. The first
 1388 term, equal to $C_{ijnm}^{\psi\psi} \sigma_{nm}^{eff}$ describe compaction/extension with the rate proportional to
 1389 the effective stress. The second one is the damage-related coupling term, which is the
 1390 $\frac{\partial F}{\partial \Omega_{ij}}$ multiplied to D_{ijkn} , and describes the compaction or dilation induced by the
 1391 damage growth. Representing the effective stress as a superposition of the volumetric
 1392 (P_e) and deviatoric (τ_{ij}) components allows to describe different mechanisms of the
 1393 irreversible strain accumulation or ψ_{ij} kinetics. The pressure driven compaction under
 1394 hydrostatic load is described by a well-known porosity reduction to its pressure-
 1395 dependent equilibrium value or Athy's (1930) law. Lyakhovsky *et al.* (2021) modified
 1396 the scalar Athy relation accounting not only for volumetric effects, but also
 1397 directional effects associated with changes of the shape of pore space and
 1398 accumulation of irreversible deviatoric strain components. The suggested 3-D
 1399 equilibrium compaction strain, $\psi_{ij}^{(eq)}$, depends on both pressure and deviatoric stress
 1400 components:

$$1401 \quad \psi_{ij}^{(eq)} = B_0 \left[\delta_{ij} - \exp \left(-\frac{\sigma_m^{eff}}{B_1} \delta_{ij} - \frac{\tau_{ij}}{B_2} \right) \right] \quad (36)$$

1402 and suggested kinetics of the pressure-driven 3-D compaction in the form

1403
$$\frac{d\psi_{ij}}{dt} = A \left(\psi_{ij}^{(eq)} - \psi_{ij} \right) \sigma_m^{eff} \quad (37)$$

1404 Neglecting the term with deviatoric stress or taking $B_2 \rightarrow \infty$ in (36) reduces both
 1405 equilibrium compaction (36) and kinetic equation (37) to the traditional scalar Athy's
 1406 (1930) law formulated in terms of material porosity.

1407 Experimental studies suggest that permanent inelastic deformation is
 1408 accumulated in high porosity rocks, but also is related to the damage accumulation. It
 1409 starts accumulating with the onset of the acoustic emission and increases all the way
 1410 up to brittle failure (e.g., Lockner 1993, 1998, Martin & Chandler 1994). This process
 1411 is usually associated with the growth of microcracks and frictional sliding between
 1412 grains, rather than closure of voids or open space between grains. For similar reasons,
 1413 Hamiel *et al.* (2004a) related the rate of irreversible strain accumulation with the rate
 1414 of their scalar damage growth. Keeping in mind that the scalar damage variable is
 1415 equivalent to the squared damage tensor, we extend their relation to the tensor form
 1416 keeping the same structure of functional relations:

1417
$$\frac{d}{dt} \varepsilon_{ij}^{(ir)} = \begin{cases} C_V \frac{d(\Omega_{in}\Omega_{jm})}{dt} \tau_{nm}, & \frac{d(\Omega_{in}\Omega_{jm})}{dt} > 0 \\ 0, & \frac{d(\Omega_{in}\Omega_{jm})}{dt} \leq 0 \end{cases} \quad (38)$$

1418 Here we combine both mechanisms and represent the static value of the tensor $\psi_{ij}^{(S)}$ as
 1419 a sum of

1420
$$\psi_{ij}^{(S)} = \psi_{ij}^{(eq)} + \varepsilon_{ij}^{(ir)} \quad (39)$$

1421 ~~The Final-final~~ kinetic equation (12a) for the ψ_{ij} tensor incorporates all the discussed
 1422 mechanisms (to avoid very length equation D_{ijkn} and $\frac{\partial F}{\partial \Omega_{kn}}$ are not substituted here):

1423
$$\frac{d\psi_{ij}}{dt} = A \left(\psi_{ij}^{(eq)} - \psi_{ij} \right) \sigma_m^{eff} + C_V \frac{d(\Omega_{in}\Omega_{jm})}{dt} \tau_{nm} - D_{ijkn} \frac{\partial F}{\partial \Omega_{kn}} \quad (40)$$

1424 The kinetic equation for damage (35) and compaction (40) provide the close system
 1425 of equations defining the 3-D evolution of the material properties. To study some
 1426 basic model properties, we follow the assumption of Panteleev *et al.* (2021) that the
 1427 principal directions of the damage tensor match the orientation of the principal
 1428 loading axes. This assumption is supported by results of true-triaxial experiments
 1429 (Browning *et al.* 2017, 2018) demonstrating orientation of distributed micro-cracks in
 1430 the sample volume under limited load, well before their localization into narrow fault
 1431 zones. Adopting the above assumptions, we re-write the stress-strain and damage
 1432 kinetic relations for principal values in k -direction (no summation):

1433
$$\sigma_k = \lambda_0 I_1 + 2\mu_0 \varepsilon_k + 2\mu_1 \Omega_k^2 \varepsilon_k - \gamma \Omega_k \sqrt{I_2^{(\Omega)}} - \gamma \xi^{(\Omega)} \Omega_k^2 \varepsilon_k + \beta p_f$$

1434 (41)

1435 and

1436
$$\frac{d\Omega_k}{dt} = D_1 \exp[-D_2 \Omega_k] (-I_1)^N \sqrt{I_2} \sigma_m^{eff} + \frac{L}{\sqrt{\Omega_l \Omega_l}} \left[\varepsilon_k \sqrt{I_2^{(\Omega)}} + \left(\frac{I_1^{(\Omega)}}{\sqrt{I_2^{(\Omega)}}} - 2\xi_0 \right) \varepsilon_k^2 \Omega_k - \right.$$

1437
$$\left. Ch \Omega_k \right] \quad (42)$$

1438 We note that for the isotropic damage ($\Omega_1 = \Omega_2 = \Omega_3$) summation of free damage
 1439 rate components gives the same equation as for the scalar damage with the same
 1440 damage-dependent yield envelope.

1441
$$\frac{d\psi_k}{dt} = A \left(\psi_k^{(eq)} - \psi_k \right) \sigma_m^{eff} + C_V \frac{d(\Omega_k^2)}{dt} \tau_k +$$

1442
$$+ D_1 \exp[-D_2 \Omega_k] (-I_1)^N \sqrt{I_2} \left[\varepsilon_k \sqrt{I_2^{(\Omega)}} + \left(\frac{I_1^{(\Omega)}}{\sqrt{I_2^{(\Omega)}}} - 2\xi_0 \right) \varepsilon_k^2 \Omega_k - Ch \Omega_k \right]$$

1443 (43)

1444 The first term of eq. (43) represents the compaction prior to the onset of the damage
 1445 accumulation, which is 3-D extension of the scalar Athy's law compaction. According
 1446 to this term the compaction approaches to its stress-dependent equilibrium value with
 1447 the rate proportional to the effective pressure. The second term is 3-D equivalent to
 1448 the damage-dependent irreversible strain accumulation with inverse of the effective
 1449 viscosity or fluidity proportional to the rate of the damage accumulation. This term
 1450 describes extension or compaction depending on the sign of the deviatoric stress
 1451 component. The last term represents the coupling between damage and porosity
 1452 kinetics. Its sign, extension or compaction, is defined by the expression in the square
 1453 brackets and depends on the loading and damage values.

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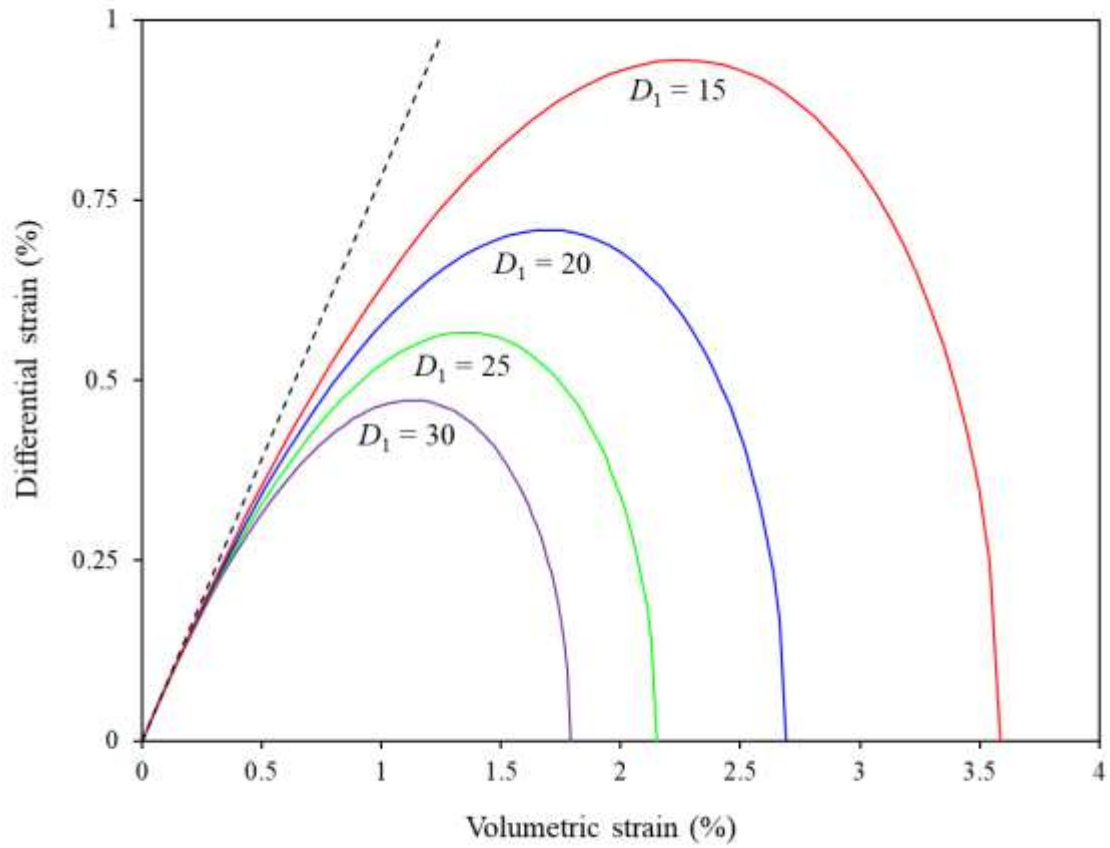
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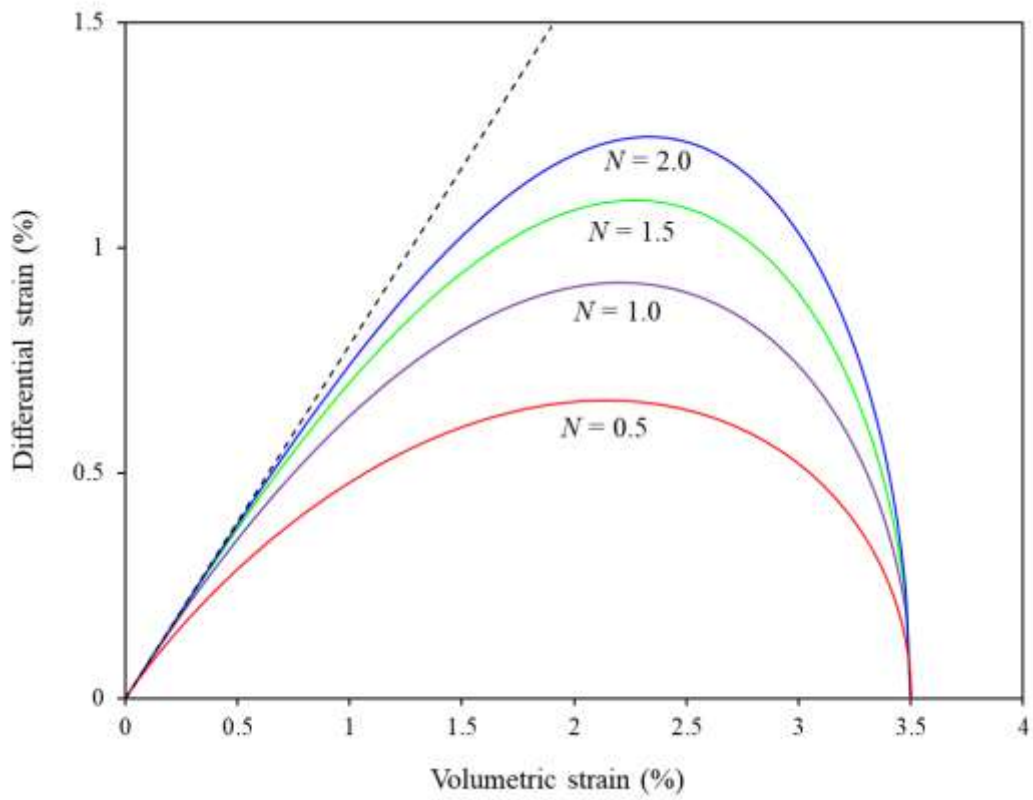
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Fig. B1. Yield envelope size increase with decrease in D -value ($N=1$)

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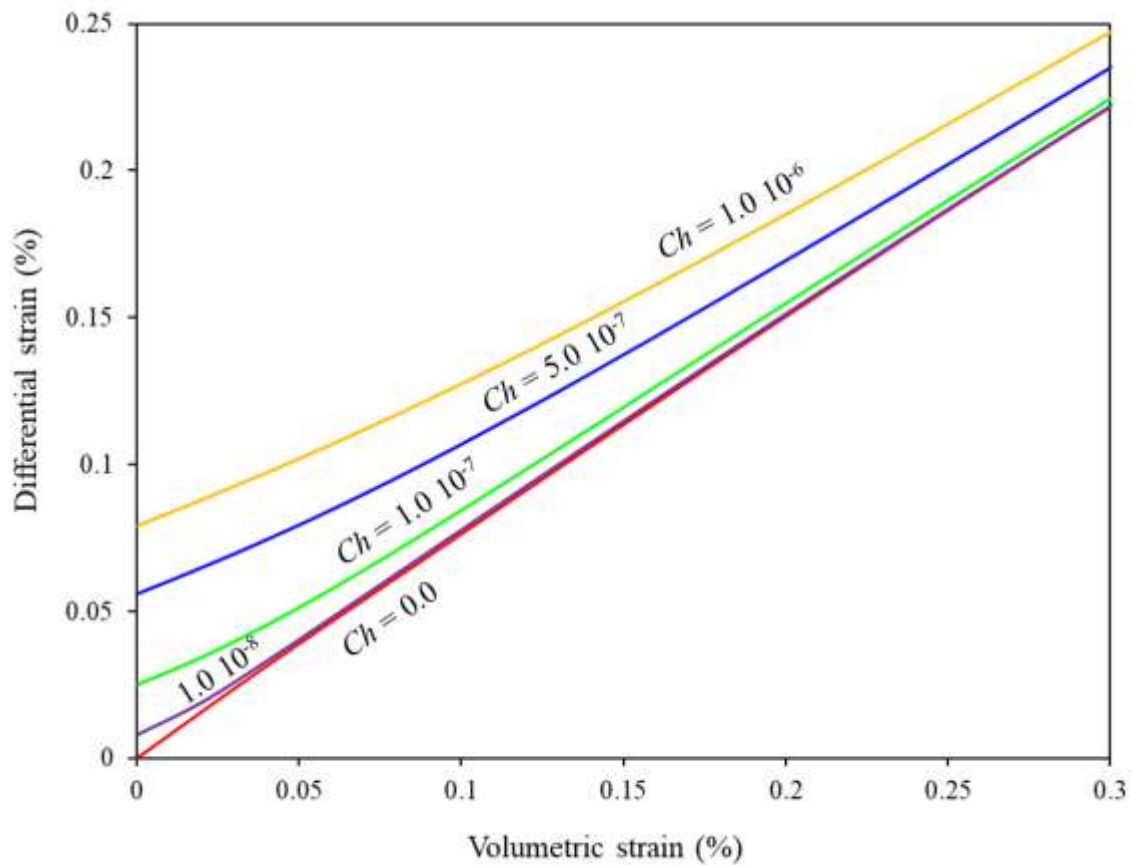


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Fig. B2. Shape of the yield envelope for N -values changing between 0.5 and 2.0

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1498 Fig. B3. The yield envelope is shifted up to higher differential strain values
1499 with increased Ch value ($N=1$ and $D=15$).

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1504 **Figure Captions (Appendix B)**

1505 Figure B1. Yield envelope size increase with decrease in D -value ($N=1$)

1506

1507 Figure B2. Shape of the yield envelope for N -values changing between 0.5 and 2.0

1508

1509 Figure B3. The yield envelope is shifted up to higher differential strain values with
1510 increased Ch value ($N=1$ and $D=15$)

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