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A sliced-3D approach to GPR FDTD modelling by optimising perfectly matched layers

(April 29, 2021)

Running head: **Sliced-3d FDTD**

ABSTRACT

Finite-Difference Time-Domain (FDTD) forward modelling is often used to gain a more quantitative understanding of the interactions between electromagnetic fields and targets. To undertake full 3D simulations the computational demands are challenging, so simulations are often undertaken in 2D where assumptions in the propagation of electromagnetic fields and source type can result in errors. Here we develop the concept of a sliced-3D simulation, wherein a thin slice of a 3D domain with strictly 2D geometry is used to minimise computational demands while obtaining synthetic waveforms that contain full 3D propagation effects. This approach requires optimisation of perfectly matched layer (PML) boundary condition parameters so as to minimise the errors associated with the source being located close to the boundary, and as a result of grazing-incident angle wave conversion to evanescent energy. We explore the frequency dependence of PML parameters, and establish a relationship between complex frequency stretching parameters and effective wavelength. The resultant parameter choice is shown to minimise propagation errors in the context of a simple radioglaciological model, where 3D domains may be prohibitively large, and for a near-surface cross-borehole survey configuration, a case where full waveform inversion may

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INTRODUCTION

Finite-Difference Time-Domain (FDTD) forward modelling has been used in many areas of exploration and near-surface geophysics to test the performance of novel processing algorithms and acquisition (Versteeg, 1993; Langhammer et al., 2017), in data processing directly for finite difference and reverse time migration (Fisher et al., 1992; Yilmaz, 2001; Leuschen and Plumb, 2001; Church et al., 2018), and as a part of inversion algorithms including full waveform inversion (FWI) (Virieux and Operto, 2009; Busch et al., 2012; Mozaffari et al., 2016). In electromagnetic applications, 2D formulations of the Yee algorithm (Yee, 1966) are generally used, which make the implicit assumption of lateral model invariance. The resultant synthetic 2D data have an incorrect amplitude scaling with travel time for which a correction must be made. Many studies have employed a Bleistein filter (Bleistein, 1986; Auer et al., 2013) in pre-processing of field data to enable comparison with 2D models (Mozaffari et al., 2016; Klotzsche et al., 2019), but it has been demonstrated that this can result in errors after the first break or in complex velocity models (Auer et al., 2013).

Reduction to 2D requires the operator to assume that the radar antennas are either cross-line or in-line, modes that are typically and hereafter denoted TM_z and TE_z respectively. The most commonly-used modelling platforms apply TM_z reduction from the principle that cross-line antennas are more widely used in many fields. However, the importance of source polarisation has been noted in several areas of the literature, including in glaciology (Langhammer et al., 2017), where the TE_z mode is more commonly applied in ground-based studies (e.g., Bingham et al. (2017)). To address the issues outlined above, 3D modelling

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4 must be developed, yet the computational demands are intense, and therefore there is a
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6 need to develop a computationally efficient approach to **modelling** 3D fields.
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9 In this paper, we seek to **minimise** the computational cost of full ~~polarisation~~ FDTD
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11 **modelling** of 2D geometries using a sliced-3D approach in gprMax, an open-source GPR
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13 **modelling** package (Warren et al., 2016). To do so we must **optimise** the boundary con-
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15 ditions, implemented by perfectly matched layers (PMLs) so as to attenuate noise due to
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17 grazing-wave interactions with the model boundaries. We investigate the frequency depen-
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19 dence of PML performance for the sliced-3D application, and demonstrate the effectiveness
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21 of the approach by applying the technique to two synthetic case studies where full 3D mod-
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23 els can be prohibitively large and where assumptions about the source and propagation
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25 mechanisms, that are implicit in 2D modelling, do not hold.
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31 THEORETICAL BACKGROUND

32 Approaches to **modelling** 2D geometries

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35 FDTD **modelling** is generally undertaken using Yee's algorithm (Yee, 1966; Taflove and
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37 Hagness, 2005). In brief, the algorithm involves a discretisation of Maxwell's equations of
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39 electrostatics, and an iterative propagation of a source term through time steps. The
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41 algorithm can be implemented in 3D or simplified to 2D in the TMz mode by assuming
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43 an infinitely long z-polarized dipole antenna (i.e. a line source) and cross-line geometry
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45 invariance to remove invariant E and H field components (Taflove and Hagness, 2005)
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47 (Figure 1). 2D simulations comprise a computationally quick method of **modelling** the
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49 response of a laterally invariant model. In practice, however, the assumption of an infinite
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51 z-polarised source is often violated due to the field logistics imposed on many GPR surveys.
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4 66 For example, due to the low frequencies often used in ground-based glaciological radio-echo
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6 67 sounding (Scott et al., 2010; Sevestre et al., 2015; King et al., 2016), lengthy dipole antennas
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8 68 are often towed in-line to the survey direction and as such cannot be modelled accurately
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10
11 69 using 2D FDTD algorithms.

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14 70 Additional issues with the 2D approach are encountered in the scaling of amplitude
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16 71 with travel time. In a 3D domain with a point source, $A \propto \frac{1}{r}$, where A is amplitude
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18 72 and r is distance, but in 2D the source becomes an infinite dipole and the relationship
19
20 73 becomes $A \propto \frac{1}{\sqrt{r}}$ (Bleistein, 1986; Auer et al., 2013). Because of this, when 2D modelling
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22 74 is employed the results need to be post-processed to obtain amplitudes that quantitatively
23
24 75 match field data. The 2D Green's function can be transformed between 2D and an equivalent
25
26 76 3D function through a $\frac{\pi}{4}$ phase shift and an amplitude scaling using the Bleistein filter
27
28 77 (Bleistein, 1986), expressed in the frequency domain as

$$G^{3D}(\omega) = G^{2D}(\omega) \sqrt{\frac{|\omega|}{2\pi\sigma}} \exp\left(-\text{sgn}(\omega) \frac{j\pi}{4}\right) \quad (1)$$

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36 78 where G^{2D} and G^{3D} are the 2D and 3D Green's functions, ω is angular frequency, $j = \sqrt{-1}$
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38 79 and $\text{sgn}(\omega)$ is the signum function of ω . σ is a scaling factor $\sigma = cr$, where r is distance
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40 80 (m) and c is velocity of propagation (ms^{-1}). This widely-used function (e.g. Deregowski
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42 81 and Brown, 1983; Vidale et al., 1985; Esmersoy and Oristaglio, 1988; Yang et al., 2013;
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44 82 Lomas and Curtis, 2019) is an asymptotic solution making the far-field assumption that
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46 83 distance $r \gg \lambda$, the wavelength of the signal, hence the near-field phase corrections are
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48 84 incorrect. The scaling function σ is commonly estimated for the first break arrival and is
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50 85 often inaccurate for the cases of (a) heterogeneous media, where c and r are uncertain or
51
52 86 complex, and (b) for later arrivals after the first break. Inaccurate amplitudes result in a
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54 87 degraded performance for FWI algorithms (Auer et al., 2013), resulting in more complex
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4 approaches requiring a good starting velocity model to be used (Van Vorst et al., 2014).
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7 To overcome issues of amplitude scaling, and to retrieve EM polarisations in the in-line
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9 survey orientation using a 2D **modelling** domain, several authors have used 2.5D implemen-
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11 tations of the Yee algorithm. These project the 3D algorithm onto a 2D plane by iterating
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13 over a series of constant wavenumbers k_z (e.g. Stoyer and Greenfield, 1976; Moghaddam
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15 et al., 1991; Xu and McMechan, 1997). This approach involves multiple easily **parallelis-**
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17 **94** **able** 2D syntheses, yet requires a reformulation of the Yee algorithm and post-processing
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19 of results, meaning that they have not, to date, been readily implemented in open-access
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21 FDTD software packages.
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27 **Sliced-3D FDTD **modelling****

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30 While the above approaches to data pre-processing are effective in converting processing
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32 to a 2D problem, full 3D FDTD **modelling** of 2D geometries remains the optimal solu-
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34 tion for generating full 3D **polarisation** and propagation effects (e.g. Mozaffari et al., 2016;
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36 Langhammer et al., 2017), although the computational demands of this approach can be
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38 significant. **Minimising** the width of a 3D domain is therefore desirable to **minimise** com-
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40 putational requirements, while retaining the benefits of 3D **modelling**. This we refer to
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42 as a sliced-3D approach, as it uses the 3D FDTD algorithm with a laterally-invariant 2D
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44 geometry, hence retaining the aforementioned correct amplitude scaling and source polar-
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46 ization capabilities. In the following we show that minimising the domain width can only be
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48 achieved through **optimisation** of boundary conditions, and that a such a sliced-3D approach
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53 can show improvements over 2D **modelling** for near-surface GPR **modelling**.
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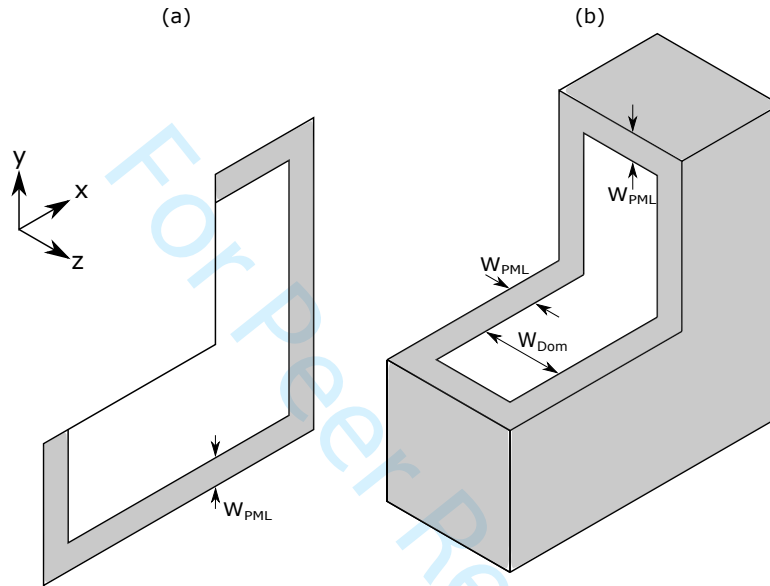


Figure 1: Schematic of (a) a 2D model and (b) a sliced-3D model, where $W_{Dom} > dx$; $W_{PML} = 15$ for both cases. Grey represents the PML region and white represents the model domain. The 2D model uses a 2D FDTD grid, while the sliced-3D model is a 3D FDTD domain with a minimised z domain width, bounded on all sides by a CFS-PML.

109 Perfectly Matched Layers

110 The boundaries of an FDTD grid are often terminated using a perfectly matched layer
 111 (PML) in which a complex stretching function s_u is used to both scale the model domain,
 112 and provide a mechanism for reflectionless signal attenuation. In the PML region, using
 113 cyclic notation $(i, j, k) \in (x, y, z), (y, z, x), (z, x, y)$ (Giannopoulos, 2018), Maxwell's equa-
 114 tions become

$$j\omega\tilde{D}_i = \frac{1}{s_j} \frac{\partial\tilde{H}_k}{\partial j} - \frac{1}{s_k} \frac{\partial\tilde{H}_j}{\partial z} \quad (2)$$

$$j\omega\tilde{B}_i = \frac{1}{s_k} \frac{\partial\tilde{E}_j}{\partial k} - \frac{1}{s_j} \frac{\partial\tilde{E}_k}{\partial j} \quad (3)$$

116 **Minimising** the z-dimension of a 3D model results in energy propagating within the
 117 model domain at grazing (low-incidence) angles to the PML boundary, hence we use a com-
 118 plex frequency stretched PML (hereafter, CFS-PML) (Roden and Gedney, 2000; Berenger,
 119 2002; Taflove and Hagness, 2005; Giannopoulos, 2008) where the stretching function s_u is
 120 of the form,

$$s_u = \kappa_u + \frac{\sigma_u}{\alpha_u + j\omega\epsilon_0} \quad (4)$$

121 where $u \in (i, j, k)$ is the orientation perpendicular to the model boundary, κ_u is a unitless
 122 quantity which dictates a real coordinate stretch in the PML region, α is a frequency shift
 123 factor, and $\frac{\sigma}{j\omega}$ introduces an imaginary spatial coordinate stretch mainly responsible for
 124 signal attenuation. In this paper we assume that the PML parameters are the same in each
 125 orientation, so we will refer to s_u, α_u, κ_u and σ_u as s, α, κ and σ , respectively.

126 The CFS-PML parameters can be tuned to improve performance over a frequency range
 127 and reduce non-physical reflections from the PML boundary. This is done by scaling pa-
 128 rameters α, κ and σ through the PML, usually using an integer polynomial m . σ is scaled

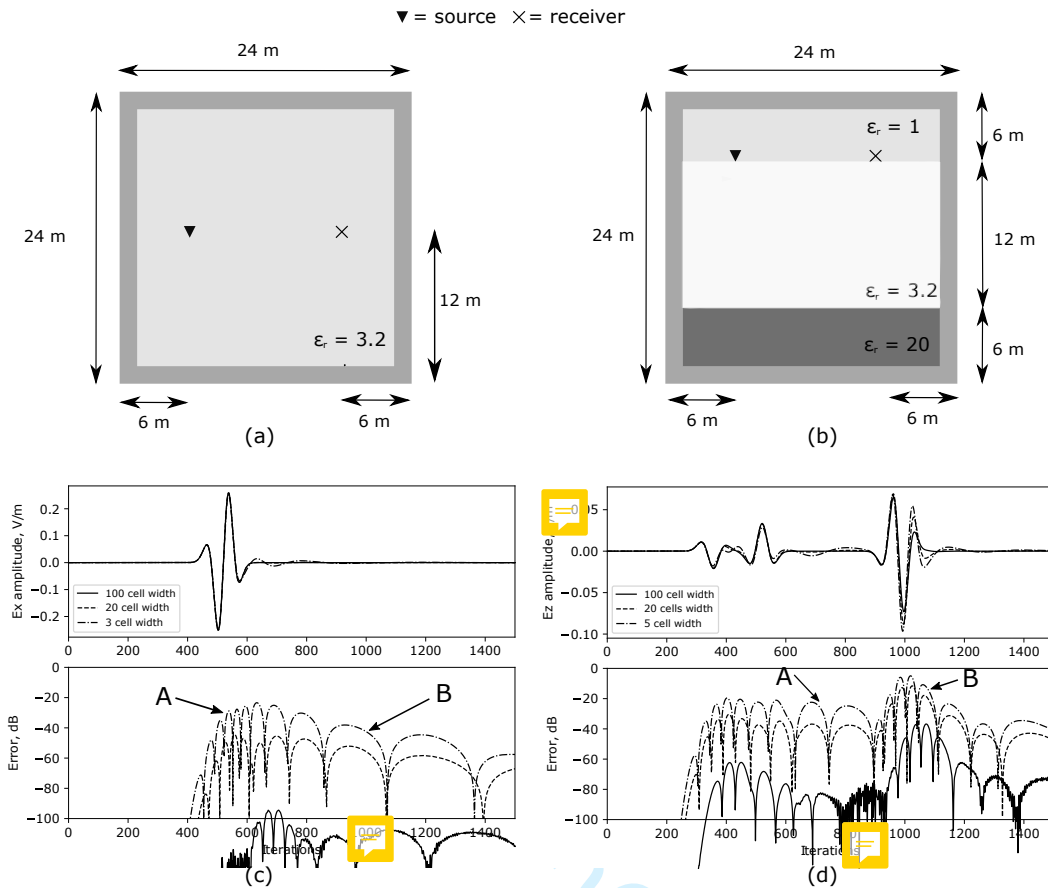


Figure 2: Comparison of the effects of cross-line domain size (in/out of the page), for (a and c) a homogeneous ice ($\epsilon = 3.2$) model with a Gaussian wavelet and a standard PML, and (b and d) a 3-layer model with a homogeneous ice layer overlying flat bedrock with a free-space layer above the surface. $\tau = 0.1m$ and PMLs are 10 cells thickness. Two sources of noise can be noted for each; 'A' shows high frequency noise as a result of normal incidence reflections through the PML. The arrival time of this noise is delayed in wider implementations, as the two way travel time between boundaries (out of the plane in (a) and (b)) increases. 'B' shows low-frequency, evanescent noise as a result of grazing wave interactions between the signal and PML boundary. A wider model results in minimisation of this noise, as the incidence angle increases with increasing width.

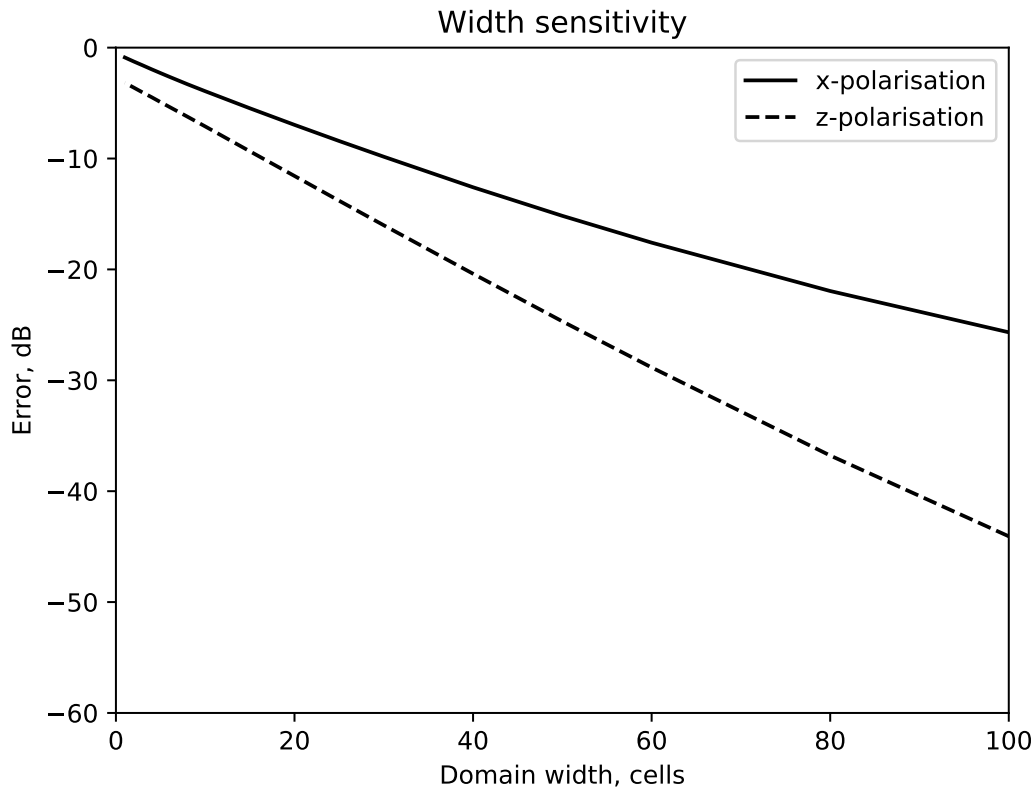


Figure 3: Plot of maximum error as a function of domain width for a homogeneous ice model shown in Figure 2(a). Decreasing model width results in an increased error as a result of interactions with grazing-angle incident energy.

129 from 0 to σ_{max} as

$$\sigma(x) = \sigma_{max} \left(\frac{x}{d} \right)^m, \quad (5)$$

130 where d is the depth of the PML in cells and $0 < x < d$ is the location within the PML

131 so as to avoid sudden changes in σ and associated non-physical reflections. We use the

132 commonly-used (Gedney and Zhao, 2010; Giannopoulos, 2012) estimate of optimum σ_{max}

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4 133 after Gedney (1996) whereby
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$$\sigma_{max} \approx \frac{m + 1}{150\pi dx \sqrt{\epsilon_r}}, \quad (6)$$

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9 134 where m is a polynomial scaling, dx is the spatial resolution, and ϵ_r is the relative dielectric
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11 135 constant. κ is similarly often scaled from 1 to κ_{max} by
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$$\kappa(x) = 1 + (\kappa_{max} - 1) \left(\frac{d - x}{d} \right)^m, \quad (7)$$

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18 136 such that $\kappa = 1$ (no coordinate stretch) at the model/PML interface and $\kappa = \kappa_{max}$ at the
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20 137 grid boundary.
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23 The frequency shift factor α is generally scaled from a maximum at the model/PML
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25 139 boundary to zero at the outermost grid boundary, to **minimise** the reflection coefficient at
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27 140 the PML/model boundary (Taflove and Hagness, 2005) and provide broadband attenuation
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30 141 within the PML. Hence,
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$$\alpha(x) = \alpha_{max} \left(\frac{1 - x}{d} \right)^m \quad (8)$$

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35 142 **Higher order** CFS-PMLs operate via a product of multiple contributions, by
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$$s = \prod_{i=1}^N s_i \quad (9)$$

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42 143 where N is the number of terms, i is the order, and s_i is defined in equation (4), with the aim
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44 144 of combining the characteristics of improved attenuation within the PML compared to the
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46 145 standard PML with the attenuation of evanescent energy of the CFS-PML. Typically, two
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48 146 terms ($N = 2$ in equation 9) are used for a **higher order** PML, but more terms are possible by
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51 147 introducing further terms of s_i . Feng et al. (2017) **undertook an optimisation of the higher**
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53 148 **order** PML for the application of broadband seismic **modelling** and showed a reduction
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56 149 in the error as a result. However, it is clear from inspection that such implementations
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150 introduce cross-terms in addition to the desired terms as, for a 2nd order CFS-PML

$$s = \left(\kappa_1 + \frac{\sigma_1}{\alpha_1 + j\omega\epsilon_0} \right) \left(\kappa_2 + \frac{\sigma_2}{\alpha_2 + j\omega\epsilon_0} \right) \quad (10)$$

(Giannopoulos, 2018). What remains unclear is what impact these additional cross-terms have in an optimisation process. Along with the higher number of degrees of freedom associated with multiple stretching functions, this results in the process becoming a cumbersome problem for the general case, and hence will not be considered in this study.

METHODOLOGY

We initially demonstrate the impact of using a small cross-line domain size on signal error as a result of the aforementioned evanescent energy. We demonstrate the effect of reducing the cross-line domain size for both a homogeneous ice ($\epsilon_r = 3.2$) model (Figure 2a), and a layered model of homogeneous ice overlying a bedrock layer ($\epsilon_r = 20$) (Figure 2b).

We then undertake a series of sensitivity experiments with uniform models to investigate the performance of PMLs in attenuating grazing wave energy on the boundary of the sliced-3D model for a sliced-3D model with a fixed domain size of 15 cells and a PML thickness of 15 cells. The experiments are performed at (a) 25 and (b) 50 MHz using a Ricker wavelet. We use a similar approach to Taflove and Hagness (2005) and Drossaert and Giannopoulos (2007) in testing parameter pairs over an expected range to derive the optimum values because, although this is a computationally intensive option, it allows a clear assessment of the sensitivity to different parameters. We initially do this using a κ scaling polynomial $m = 2$ and α polynomial $m = 1$ (see eqs. 7 and 8). The model was discretized at 0.1m to give a model domain size of 24 x 24 x 3.5 m. The PML thickness was extended compared to a typically used 10-cell implementation, with the intention of reducing errors due to normal

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4 170 incidence energy at the bounding edges, which may not be attenuated as effectively when
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6 171 optimisation is undertaken to reduce evanescent energy. We then repeat this approach
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9 172 to investigate the impact of polynomial order m for κ and α , running this test for all
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11 173 combinations between $m = 0$ (constant value) and $m = 6$.

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14 174 We then investigate the frequency dependence of optimal CFS-PML parameters, by do-
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16 175 ing a similar grid search parameter test as for the previous tests, but this time using an
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18 176 impulse source type followed by a convolution with a Ricker wavelet with central wavelength
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20 177 λ_c . We limit frequencies used to $20 < \lambda_c/dx < 100$, as this is the most commonly used range
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22 178 of λ/dx for efficient FDTD modelling, also noting the dispersion limit of $dx < \lambda_{min}/10$ (Gi-
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24 179 annopoulos, 1998) and that for a Gaussian waveform, the minimum significant wavelength
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26 180 considered for dispersion ($error < -40$ dB) is $\lambda_{min} \approx \frac{\lambda_c}{3}$. Using a grid size of 0.01 and 0.1
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28 181 m this allows testing in the range 100-700 and 10-70 MHz respectively. For this experiment,
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30 182 we use a 5-cell domain width with 15-cell PML.

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34 183 For each of the above sensitivity experiments a reference solution of a 3D model,
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36 184 $E(x, y, t)_{ref}$, is calculated using a large 3D model with an identical 2D geometry, to give
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38 185 the response where there is no interaction with bounding PMLs normal to the z orien-
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40 186 tation. The 3D model consists of identical geometry in the x - and y -orientation, with a
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42 187 120-cell model width in the z -orientation and a 10-cell PML using a constant $\kappa_{max} = 1$, and
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44 188 σ scaled linearly between 0 and σ_{max} after equation 6. As a result of this larger width there
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46 189 is no grazing-wave interaction with the model-PML interface, and we can assume this to
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48 190 be the best-case scenario with minimum error response. Errors are reported relative to this
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50 191 reference solution as in Roden and Gedney (2000); Berenger (2002); Giannopoulos (2008);
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4 192 Taflove and Hagness (2005); Feng et al. (2017) as
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$$7 \quad error(x, y, t) = 20 \log_{10} \frac{E(x, y, t) - E(x, y, t)_{ref}}{E_{ref_{max}}}, \quad (11)$$

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10 193 where $E(x, y, t)$ is the output electric field in time, $E(x, y, t)_{ref}$ is the reference solution in
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12 194 time, and $E_{ref_{max}}$ is the maximum value of the reference solution.

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15 195 To demonstrate the performance of the sliced-3D approach, we repeat our experiment of
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17 196 investigating domain width sensitivity to confirm that an improvement in error is observed
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19 197 using an optimised CFS-PML, before comparing the performance of an **optimised first order**
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21 198 PML with those previously published in the literature, including Feng et al. (2017), which
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23 199 was developed for computational seismology but follows similar theory, along with Gedney
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25 200 and Zhao (2010) and as outlined previously. For this case we use a model discretisation of
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27 201 0.005 m and free space ($\epsilon_r = 1$).
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33 NUMERICAL RESULTS

34 202 Impact of domain size

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37 203 Figure 2 shows the results of reducing domain size for homogeneous and layered models.
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39 204 Thin models (3 cell) model domain size show significant (> -20 dB) noise levels at signal
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41 205 arrival, followed by low-frequency ringing as a result of evanescent energy from the model-
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43 206 PML interface. Figure 3 shows the error for both x- and z- polarisations for the homogeneous
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45 207 ice model. We estimate an error of -40 dB (1%) to be a feasible target to reduce the errors
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47 208 below the signal-to-noise ratio of a typical radargram which, from **figure 3**, would require a
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49 209 domain width of 60 cells. At small domain sizes, the effect of evanescent energy is significant,
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51 210 whereby low frequency and high amplitude errors are introduced following the direct arrival
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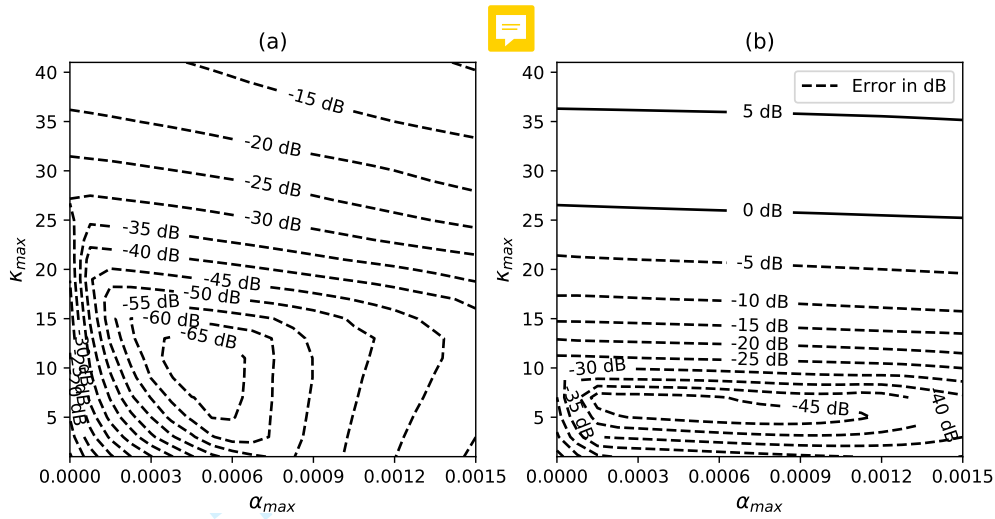


Figure 4: Contour plot of maximum error as a function of α_{max} (frequency shift factor, equation 4) and κ_{max} (stretching factor) for a homogeneous ice model with a 5 cell width model domain, 15 cell first-order PMLs, with $dx = 0.1\text{m}$. A Gaussian waveform with central frequency (a) 25 MHz and (b) 50 MHz is used.

211 (arrivals in Figure 2, marked 'B'). Thickening the PML has minimal impact on this error
 212 as it is induced by the model/PML boundary.

213 1st order PML optimisation

214 The optimum values for α and κ were estimated through a brute-force grid search approach,
 215 producing error contour plots exemplified by Figure 4. The grid search shows minimum
 216 error bounds of -65 dB and -45 dB for 25 and 50 MHz, respectively. A clear frequency
 217 dependence of the optimum parameters can be seen, indicating that optimum κ_{max} decreases
 218 with increasing frequency, and that the sensitivity of error to the α value decreases with
 219 increasing frequency. This is intuitive as κ dictates the real coordinate stretch of the PML

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4 220 - a higher value results in a higher stretch, such that the maximum $\frac{\lambda}{dx}$ within the stretched
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7 221 coordinates of the PML is minimised. The optimum value of α is approximately the same
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9 222 for both experiments, but has a much lower sensitivity **in the high frequency.**

223 **Impact of Polynomial Order**

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16 224 Figure 5 shows the minimum error for each grid search as a function of order of polynomial
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18 225 scaling. It is clear that, for this example, a constant α scaling function is the most efficient,
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20 226 with a maximum -80 dB error. Higher orders of α result in an error of at least -50 dB. A
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23 227 quadratic κ scaling function is shown to provide the optimum attenuation for all orders of
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25 228 α . This result contrasts with Taflove and Hagness (2005), where it is suggested that $\alpha = 0$
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27 229 at the outermost grid boundary to enable sufficient travelling wave energy attenuation. Our
28
29 230 optimal parameter setting is therefore **minimising** the effect of evanescent energy, with the
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31 231 remaining noise being primarily as a result of normal-incidence energy at the source point.
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36 232 **Frequency Dependence**

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40 233 Figure 6 shows minimum error, optimum α and optimum κ as a function of $\frac{\lambda}{dx}$, firstly
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42 234 demonstrating (Figure **6 (a)**) that error is relatively constant at approximately -70 dB for
43
44 235 all values of $\frac{\lambda}{dx}$ tested. Figure **6 (b)** shows the optimum selection of κ is linear with $\frac{\lambda}{dx}$,
45
46 236 with a linear relationship of

$$47 \kappa_{max} = 0.14 \frac{\lambda}{dx} - 1. \quad (12)$$

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51 237 In Figure **6 (c)**, α_{max} is plotted as $\log_{10} \alpha_{max} dx$ as a function of $\frac{\lambda}{dx}$. A negative linear
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53 238 relationship can be plotted for the range $20 < \frac{\lambda}{dx} < 70$, of form $y = mx + c$,

$$54 \log_{10}(dx\alpha) = -0.005 \frac{\lambda}{dx} - 4, \quad (13)$$

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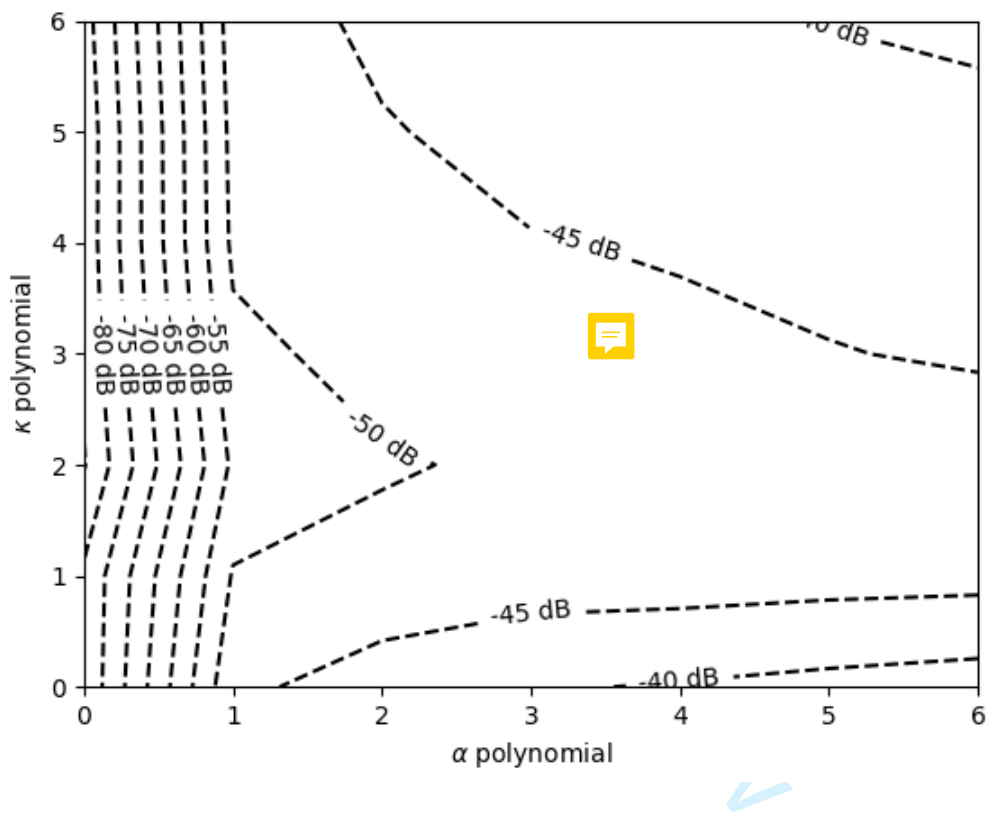


Figure 5: Minimum error for all combinations of polynomial scaling. This is found through repeating the results of Figure 4 for each combination of polynomials in α and κ . The optimum value is shown to be 0 for α and 2 for κ .

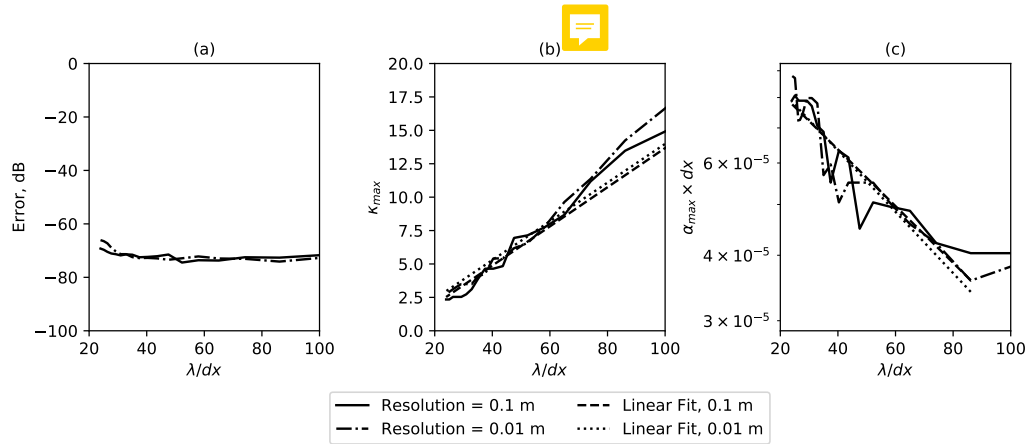


Figure 6: Frequency dependence of the 1st order CFS-PML parameters using a discretisation of 0.1 m and 0.01 m for the homogeneous model (Figure 2a). (a) shows the error as a result of the optimum parameters. (b) shows optimum κ_{max} as a function of $\frac{\lambda}{dx}$. A positive linear trend is observed as expected as a larger κ_{max} is expected for larger wavelengths. (c) shows $\alpha \times dx$ plotted as a function of $\frac{\lambda}{dx}$. This plot is scaled by discretisation on a lin-log plot, demonstrating that optimum α shows a slight negative trend with λ/dx and a scaling with resolution

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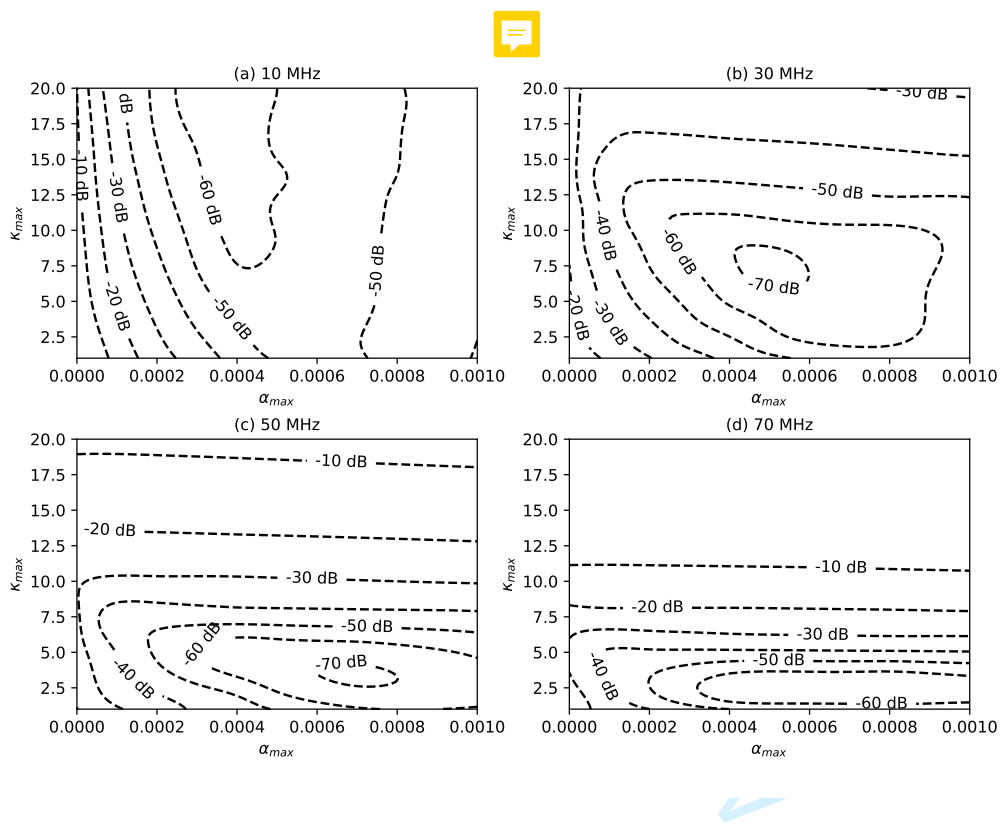


Figure 7: The error surface as a function of κ and α for a range of frequencies for $dx=0.01$ using the same experimental setup as in Figure 6 (a).

239 which reduces to

$$\alpha = \frac{10^{-4-0.005 \frac{\lambda}{dx}}}{dx}. \quad (14)$$

240 Together with equation 6, these values can be readily used as a guideline for 1st order
 241 CFS-PML parameters in the frequency range suggested as they only require calculation of
 242 a central wavelength λ and the discretisation. As such, they can be readily calculated in the
 243 FDTD implementation. Figure 7 additionally shows that at higher frequencies (i.e. lower
 244 values of λ/dx), the error is much more sensitive to the value of κ_{max} than to the value of
 245 α_{max} , and this provides a more stable linear regression result in Figure 6.

246 Domain width revisited

247 With our new understanding of optimum CFS-PML parameters, we now revisit signal error
 248 as a function of domain width (Figure 8). The optimised PML gives a consistent result of
 249 -38 dB for an x-polarised source type, and -45 dB for a z-polarised source type. The error
 250 increases slightly at a domain width of 10 cells for a z-polarised source, but remains under
 251 -40 dB down to a 3-cell domain width. In the following examples, we use a 5-cell domain
 252 width as a balance between computational requirements and accuracy.

253 Comparison of implementations

254 The performance of differing implementations is compared in Figure 9 using a sliced-3D
 255 homogeneous ice model as in Figure 2 (a), now using 60 receivers in the positive x direction,
 256 representing a common source point experiment. The difference between each result and
 257 the reference solution in Figure 9 (a) is shown in panels (b) to (e). This demonstrates that
 258 a correctly optimised 1st-order CFS-PML can produce synthetic data with no evanescent

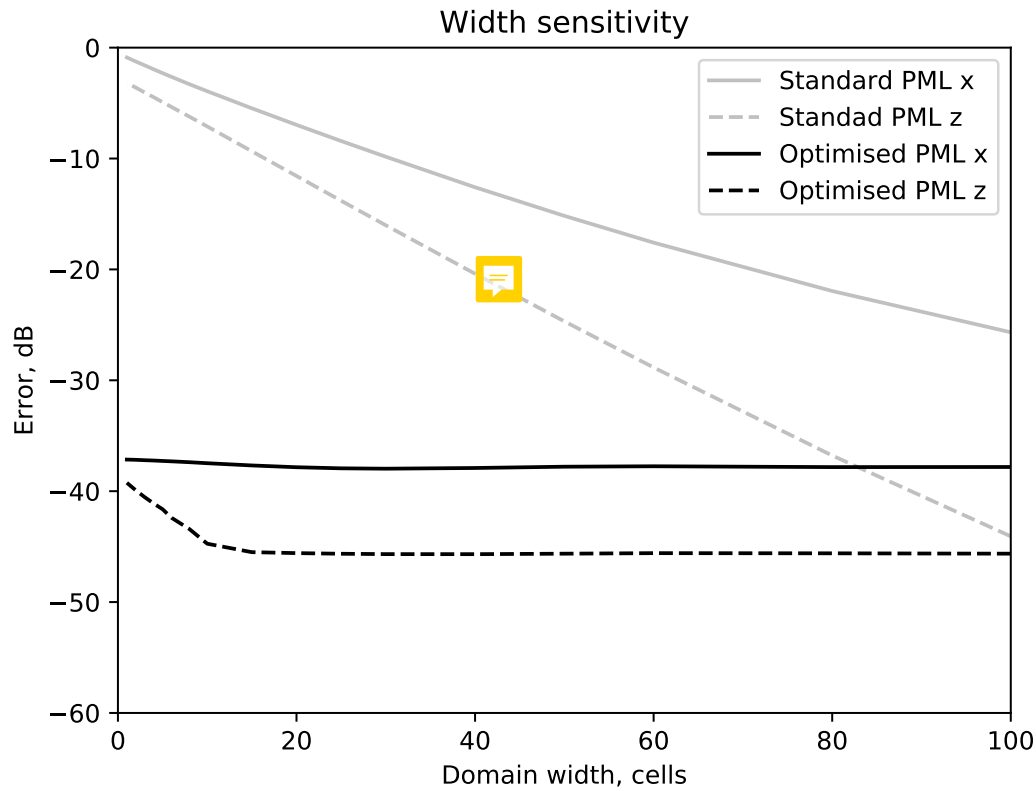


Figure 8: Comparison of the signal error as a function of domain width, as in Figure 3, but including results with an optimised CFS-PML, using the horizontally layered model as in Figure 2c) and d). A consistent -38 dB can be achieved for the x-polarised result and -45 dB for the z-polarisation using our recommendations for CFS PML parameters.

energy in a sliced-3D model domain. There is a slight error close to the source point in panel (e), as a result of the CFS-PML's reduced ability to attenuate normal-incidence energy. Other recommendations (panels (d) and (e)) for CFS-PML parameters show strong evanescent energy, showing that while these have been recommended for a general case for 1st-order and 2nd-order CFS-PMLs, they are not suitable in this application.

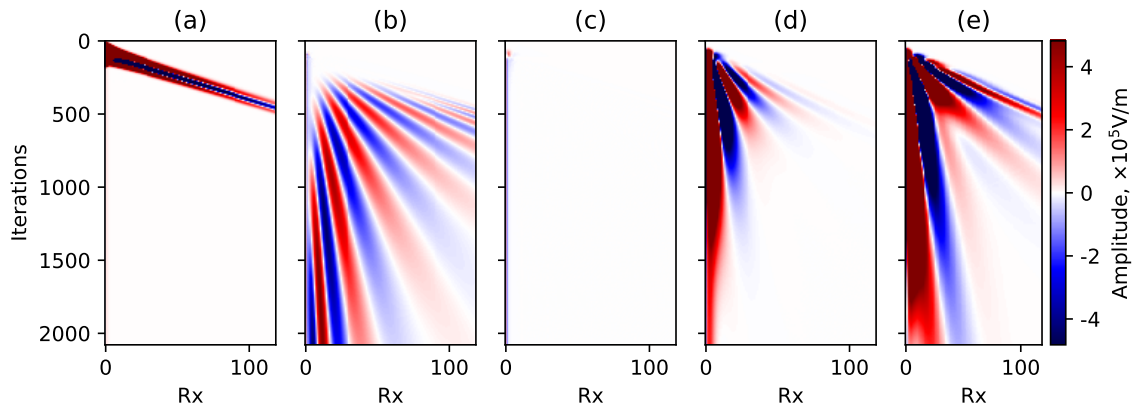


Figure 9: Error plots as a function of receiver offset and time for a 5-cell width sliced 3D domain with a z-polarised source at 50 MHz. Error is the difference between the result and a reference 3D solution. Colour scale is clipped at 1% of the maximum. (a) Reference solution from a 3D model showing direct arrival (b) a sliced-3D domain with no PML parameter optimisation (c) the same model with optimum parameters selected from equations 12 and 14, (d) with parameters selected from the results of Feng et al. (2017) and (e) with parameters recommended by Gedney and Zhao (2010). This comparison demonstrates that a well-optimised 1st order CFS-PML, using recommendations from this study, can show an improvement for grazing-wave interactions over generic parameters chosen for both 1st and 2nd order PMLs, which are often developed for different applications.

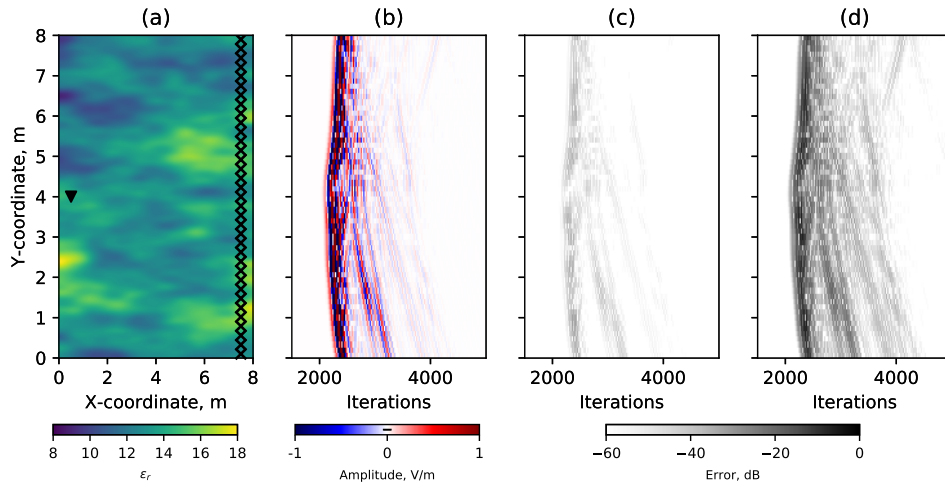


Figure 10: Cross-borehole GPR experiment example. (a) Model domain showing random variations in dielectric constant, overlaid with source point (triangle) and receiver locations (crosses). (b) Results of a full 3D modelling experiment using a z-polarised source. (c) Error plot (in dB) using a sliced-3D domain with parameters recommended in this paper. (d) Error plot (in dB) of 3D-to-2D transformed data using a Bleistein filter

EXAMPLES

264 Cross-borehole example

265 We now demonstrate the performance of sliced-3D FDTD modelling in two applications
 266 for which error levels and model computational demand are important considerations. We
 267 first use a cross-borehole survey configuration in the presence of a heterogeneous soil with
 268 ϵ_r ranging between 8 and 18. This is similar to the cross-borehole FWI experiment config-
 269 urations of Klotzsche et al. (2010) and the computational configuration of Mozaffari et al.
 270 (2016). We use a single z-polarised source point with a 200 MHz ~~central frequency~~ Ricker

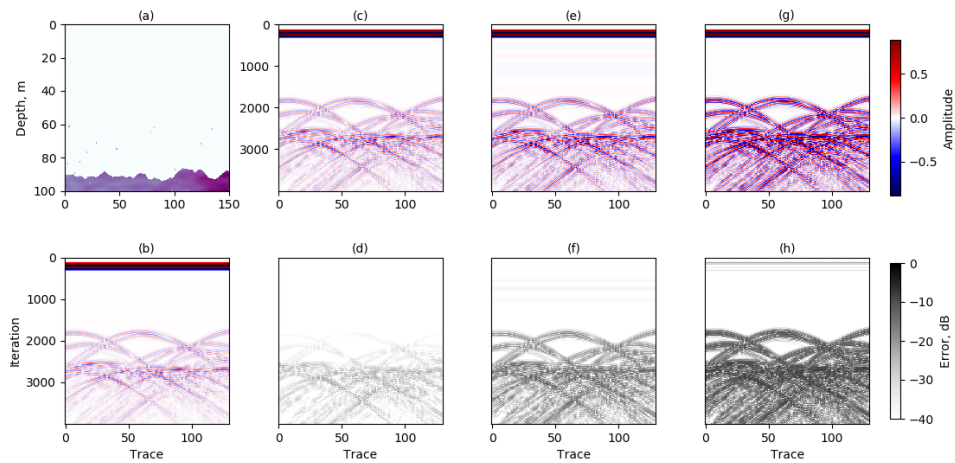


Figure 11: Synthetic representing a glacier bed with internal scattering points within the ice. (a) Initial model with homogeneous ice and a rough bed. (b) Reference model response from full 3D simulation. (c) Model response and (d) error with a sliced-3d domain and an optimised CFS-PML, using the [recommendations from section .](#) (e) Model response and (f) error for a sliced-3d domain with no optimised CFS-PML, using $\alpha_{max} = 0$, $\kappa_{max} = 1$ and $\sigma_{max} = \sigma_{opt}$. (g) Model response and (h) error for a 2D model followed by 2D-to-3D Bleistein filter transformation. A significant improvement in error can be observed when the correct source polarisation is used in a sliced-3D approach.



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4 271 wavelet, with an array of receivers located in a second borehole (see Figure 10a). The
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6 272 source and receiver boreholes are separated by 6 m. Using a discretisation of 0.02 m, the
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9 273 recommended parameters from equations 12 and 14 are $\alpha = 0.00397$ and $\kappa_{max} = 1.80$. We
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11 274 undertake the simulations in 3D, sliced-3D and in 2D. The sliced-3D model domain consisted
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13 275 of 1 cell width, with PMLs extended to 15 cells to minimise noise from normal-incidence
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18 277 All simulations are undertaken with a z-polarised source to enable like-for-like polar-
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20 278 isation comparison with the 2D implementation. We apply a frequency-domain Bleistein
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22 279 2D-to-3D filter to the 2D data (equation 1), with r equal to the straight line raypath be-
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24 280 tween source and receiver for each trace and c calculated from the RMS value of ϵ_r from the
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26 281 model. We compare the results in figure 10 (c) and (d), which shows a significantly lower
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28 282 error field for the sliced-3D approach.

283 **Common-offset glacier survey example**

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37 284 We now apply this approach to a model of a simple glacier with a rough bed and several
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39 285 internal scattering points. We use model dimensions of 150 x 100 x 3.5 m with a 15 cell
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41 286 PML thickness and resolution 0.1 m to demonstrate the low noise level achievable with our
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43 287 recommendations. A dipole source with a ~~thicker wavelet of central frequency of 25 MHz~~
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45 288 used. Given these model parameters, CFS-PML parameters are chosen to be $\alpha = 0.00046$
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47 289 and $\kappa_{max} = 3.70$, following Figure 6 and equations 12 and 14. We use a single-channel,
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49 290 common offset survey acquisition with source and receiver separated by 5 m to represent a
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51 291 typical survey with low-frequency dipole antennas. Several scattering points with $\epsilon_r = 80$
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53 292 are imposed to simulate scattering bodies found within polythermal ice (Barrett et al.,
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4 293 2008). We use 130 source/receiver locations along the surface of a freespace/ice interface.
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7 294 The results of the model are presented in Figure 11, along with error in dB in the second
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9 295 row ~~(d), (f) and (h)~~. Figure 11 (c) shows the solution for a sliced-3D model with optimised
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11 296 CFS-PML parameters with error compared to a 3D reference. This shows that optimisation
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13 297 of PML parameters can lower the error for scattering bodies to be consistently below -40
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15 298 dB, with only some later arrivals close to the bed with an error greater -40 dB. Figure 11
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17 299 (e) shows the response and (f) the error for a sliced-3D model with no PML optimisation. 
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19 300 Low-frequency noise is prevalent throughout and errors at the bed are significant. (g) and
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21 301 (h) show the response for a 2D model with 2D-to-3D transformation with the Bleistein filter
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23 302 assuming a first break time of $1\mu s$ for the bed return (2600 iterations Figure 11), and fails
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25 303 to replicate well the amplitudes for any of the scattering or bed returns 
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31 DISCUSSION

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34 304 The numerical results from the examples above show that the errors caused by near-grazing
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36 305 wave interactions with a bounding PML region can be significantly attenuated through
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38 306 optimisation of the first order CFS-PML parameters. We have suggested relationships
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40 307 between optimal parameters and model parameters to attenuate such low-frequency energy
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42 308 significantly as a function of λ/dx , which can be readily calculated using model parameters
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44 309 and source frequency used.
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49 310 In practice, the effect of κ in the CFS-PML formulation is a real stretching of the cells
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51 311 within the PML region. Higher values of κ result in increasing cell size within the PML
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53 312 region. As such, κ_{max} is a balance between larger stretch and non-attenuated dispersive
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55 313 effects. For larger stretch coefficients, low frequency energy is more effectively attenuated,
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4 314 although large cell sizes can result in numerical dispersion at the outermost bounds of the
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6 315 PML introducing high frequency noise that the PML is not effective at attenuating.
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9 316 While there have been significant developments in PML implementations through in-
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11 317 creasingly complex and higher order stretching functions, this study represents the first
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13 318 numerically-based approach to optimise 1st order CFS-PML parameters for a broad range
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16 319 of low frequency geophysical applications. We have compared our results to values pub-
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18 320 lished in the literature (Figure 9), although it must be noted that the previous values have
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20 321 been derived and estimated for different applications, and as such the performance cannot
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22 322 always be expected to match those derived for this application.
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25 323 We have suggested that our parameter choices can be applied for radioglaciological
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27 324 survey, but the effects of a wider range of dielectric materials have not been explored.
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29 325 Regions of higher ϵ_r result in increased numerical dispersion in the propagation, resulting
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31 326 in the requirement of a higher resolution model. In such a case, we require improved
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33 327 attenuation of lower $\frac{\lambda}{dx}$ values, shown in Figure 6 to result in a higher sensitivity of error
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35 328 on κ_{max} . This may be a limitation of the technique in applications to wider geoscientific
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37 329 applications of sliced-3D FDTD modelling.
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42 330 Further work in this area could explore the improvements that may be attained through
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44 331 optimisation of higher order CFS-PMLs, or through optimisation of recently developed
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46 332 multi-pole PML (Giannopoulos, 2018) However such approaches will necessarily be more
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CONCLUSIONS

334 We have shown through numerical modelling that optimisation of a 1st order CFS-PML can
335 be undertaken to minimise domain size to obtain full 3D polarisation synthetics in the case
336 of strictly 2D geometries. Such an approach is required to reduce the impact of grazing-
337 angle evanescent energy close to the model and PML boundary. For a 5-cell domain size
338 with a 15 cell PML, we can reach a maximum amplitude error of -70 dB (or 0.03%) over the
339 typical range of $\frac{\lambda_c}{dx}$ used for efficient numerical modelling. We have suggested relationships
340 between CFS-PML parameters α , κ and $\frac{\lambda}{dx}$ which demonstrate the suitability of such an
341 approach for wider applications of GPR FDTD modelling where consideration of waveform
342 polarisation is important. These recommendations mean this approach is readily applicable
343 in iterative processing algorithms, as parameters can be automatically estimated using the
344 defined model.

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