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Geochemical study of the early cretaceous Fahliyan oil reservoir in the northwest Persian Gulf

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Abstract

Three crude oil samples from the Fahliyan Formation in ‘KG’ and ‘F’ fields in the northwest Persian Gulf, namely KG-031, F9A-3H and F15-3H for the geochemical study. In this study, the physicochemical properties, gas chromatography (GC, GC Mass) and (Detailed Hydrocarbon Analysis) DHA analyses for the collected Fahliyan oils were carried out. The API, Trace Element (Ni, V) and S% parameters indicated that the Fahliyan oil was generated from a source rock which deposited in reducing environment condition with a carbonate-shale compound lithology. Moreover, low pour point, higher S% and low viscosity parameters of “KG” sample confirmed the existence of medium oil characteristics in this field. In addition, the geochemical outcomes of GC, GC–MS and DHA analyses indicated that the ‘KG’ oils are more aromatic compared with ‘F’ oil; while biomarkers revealed that Fahliyan reservoir oil is highly mature and was formed from a carbonate source rock containing types II, III kerogen. Thus, sterane/hopane biomarkers (C_{24}/C_{23} and C_{22}/C_{21} ratios) revealed that Fahliyan oil originated from carbonate source rocks deposited in an anoxic to dysoxic environment, which is consistent with the above analyses. It was identified that the source rock age is early Cretaceous to late Jurassic. It can be reported that the Fahliyan oils from both fields were generated in the same source rock and have almost the same physical properties, and will have the same production strategy.

Keywords Fahliyan formation · Oil–oil correlation · Oil–source rock correlation · Geochemistry · Crude oils

Introduction

The Persian Gulf, a small sea adjacent to the Indian Ocean, lies between Iran and the Arabian Peninsula (see Fig. 1). The Persian Gulf and its coastal parts contain the major

quantities of crude oil and natural gas, approximately 715 billion barrels of crude oil (2/3 of the world’s proven oil reserves) and 2462 tcf of gas (1/3 of the total verified world gas reserves) (Rabbani 2014; Baniasad et al. 2017). In this area, the presence of many significant regional caprocks, considerable source rock beds, great carbonate reservoirs, huge anticlinal traps and continuous sedimentation are the main factors turning this gulf into an excellent region for hydrocarbon accumulations (Rabbani et al. 2014). One part of this region is the Fahliyan Formation in the NW of the Persian Gulf, and many significant oil reservoirs were discovered there in the last 50–60 years (Soleimani et al. 2017). Thus, the presence of widespread reservoir development in combination with the organic-rich deepwater Garau formation among other source rocks turned this formation into an important target for hydrocarbon exploration (Adabi et al. 2010).

The stratigraphic and sedimentary environments of the Fahliyan Formation were reported by several authors including James and Wynd (1965), Shakib (1994), Moallemi

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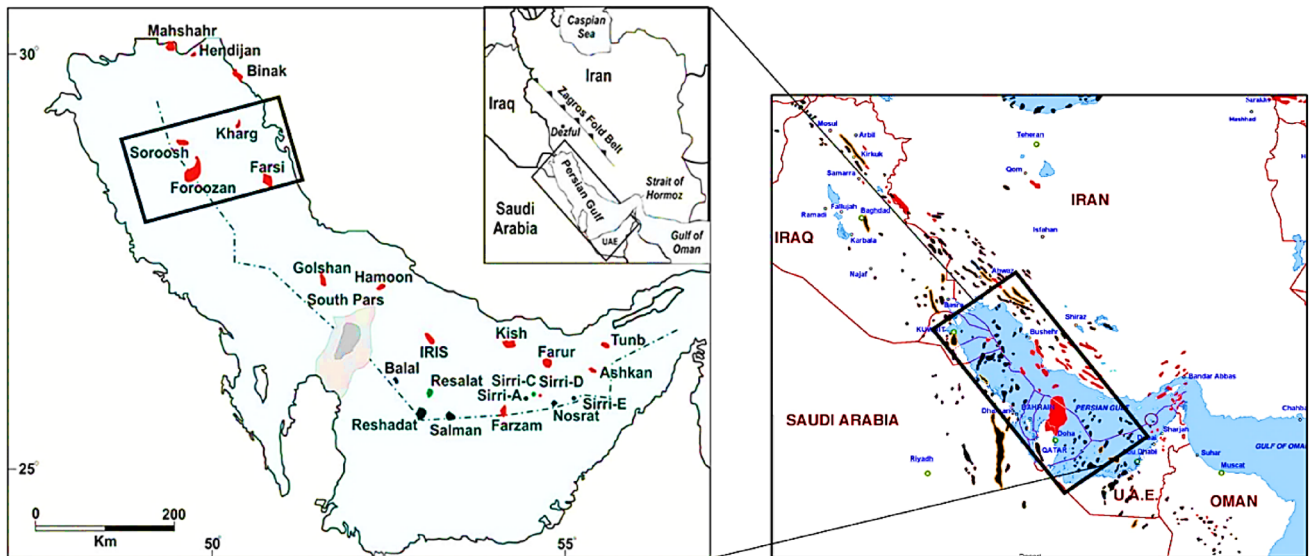


Fig. 1 Map showing the Northwest Persian Gulf area (Southwest Iran) studied

(2000). In recent facies analysis, it was stated that the Fahliyan Formation in some sections of Southwest Iran is composed of both, shallow and deep marine facies; restricted marine barriers, and open marine shallow facies (deposited on a carbonate shelf platform), and pelagic and calciturbidite deep marine facies (deposited on the intrashelf basin) (Konroud and Lasemi 2008).

However, the geology of this formation is well described (see Table 1) (Kavoosi et al. 2006; Jamalians et al. 2010; Fegghi et al. 2012; Abyat et al. 2012, 2014; Mahbaz et al. 2012; Shakeri and Parham 2013; Jamalians and Adabi 2014; Soleimani et al. 2017), considerable uncertainty related to the maturity and quality of source rocks still exists which may affect on the reliability of geochemical models. Consequently, this study focuses on characterizing the Fahliyan Formation with regards to the depositional environment, organic matter, source rock, and thermal maturity in two fields in the Iranian Fars sector. Hence, a detailed organic geochemical analysis with bulk geochemical parameters for the Fahliyan Formation oils was performed here.

Geological setting

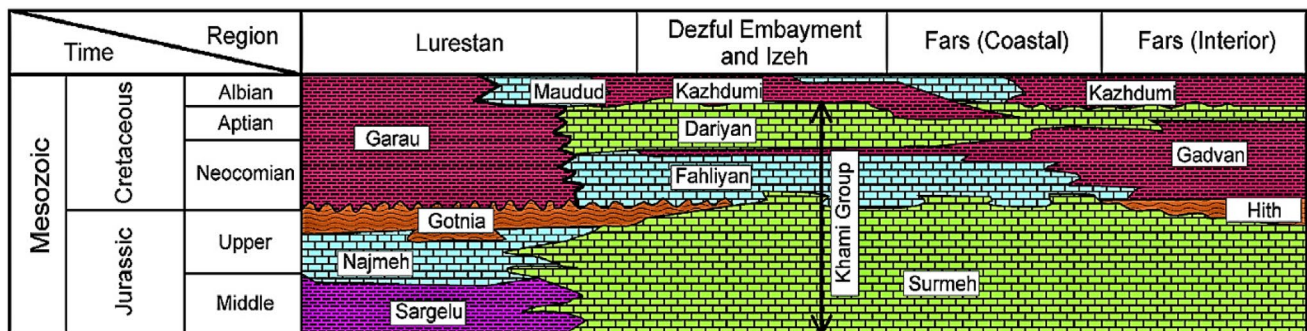
Fahliyan Formation within the surface section of the Kuh-e Siah was overlaid by the Lower Cretaceous Gadvan Formation with a conformable contact but it unconformably covers the Upper Jurassic Hith Formation (see Fig. 2) (Soleimani et al. 2017). The Kuh-e Siah anticline can be seen within the Fars Salient in the south of the Zagros fold-thrust belt (ZFTB) and is composed of a sequence of heterogeneous recent Neoproterozoic–Phanerozoic sedimentary cover strata

with a thickness of about 7–12 km (Aubourg 2004). ZFTB is the re-shaped form of the Zagros sedimentary basin. According to the current positions, this basin was stretched out over the northeast of the Afro-Arabian continental margin which develops the Zagros orogeny from the closure of the Neo-Tethys due to three geotectonic events including the sliver emplacement of oceanic crust over the Afro-Arabian continental margin, Afro-Arabian continental collision with the Iranian plates, and subduction of the Neo-Tethyan oceanic crust beneath the Iranian plates (Sherkati et al. 2005; Bahroudi and Koyi 2003). Generally, the ZFTB is extended by about 2000 km from the southeast of Turkey through the north of Syria and northeast of Iraq to southwestern Iran. Meanwhile, it is one of the most essential tectonic units in Iran which has a length of greater than 1500 km and 100–300 km in width (Abyat et al. 2012).

Zagros Orogen is composed of three separate parallel tectonic zones, such as Zagros imbricate zone (ZIZ), Urumieh Dokhtar magmatic assemblage (UDMA) and Zagros fold-thrust belt (ZFTB). ZFTB as an external part of the orogen extends parallel to the southwest of ZIZ by 300 km in width which contains some salients and recesses from northwest to southeast, such as Kirkuk recess in northern Iraq, Lorestan salient, Dezful recess and Fars salient in southwestern Iran (Alavi 2007, Kamali et al. 2013; Rbbani et al. 2014). From west, east, northeast and southeast, the Fars zone has been surrounded by Kazerun Fault, Zendan Fault, High Zagros Fault and Persian Gulf coastal line, respectively (Sabouhi et al. 2009). This zone is mainly characterized by the presence of 15–20 km anticlines which was developed from bending mechanism on the Hormuz Salt. However, these anticlines

Table 1 Summary of some previous works about the geochemistry of the Fahliyan Formation

Author	Year	Function	Outcome
Feghhi et al.	(2012)	Sedimentological analysis	The Fahliyan Formation was distinguished as a high-quality reservoir in the Southwest Iran due to occurring many effective diagenetic processes, such as stylolization, micritization, silicification, dissolution, bioturbation, and cementation
Shakeri and Parham	(2013)	Characterization and quality control	They stated that there were various diagenetic factors which have impacts on the quality of this reservoir in shallow and deep burial, meteoric, and marine environments, such as dissolution, dolomitization, compaction, fracturing, and cementation
Sabouhi et al.	(2009)	Facies analysis	It consists of one shelly facies and 13 carbonate microfacies, and subdivided into four facies groups according to the depositional settings, such as a lagoon, open marine, tidal-flat, and bar
Adabi et al.	(2010)	Geochemistry analysis	10 microfacies were identified by an analysis of the sequence stratigraphy, depositional environment, and geochemistry of Fahliyan formation
Jamalian et al.	(2010)	Characterization analysis	Identifying 25 microfacies which were divided into four deposition belts include the lagoon, tidal flat, shoal, and open-marine
Abyat et al.	(2012)	Lithostratigraphy & microbiostratigraphy	According to the identified foraminifers and algae fossil species, they identified two biozones, such as Orbitolina-Choffattella Salpingoporella and Pseudocyclammina lituus Arabica
Abyat et al.	(2014)	Microfacies analysis	They recognized 18 microfacies, which subdivided into pelagic, hemipelagic, inner and middle shelf, and coastal deposits, and outlined three biozones according to the discovered microfossils, such as Beriasian, Barremian, and Valanginain-Hauterivian
Mahbaz et al.	(2012)	Rock physics modeling	The developed model showed that the shale and limestone have various trends but marl was always laid stable between them
Abedpour et al.	(2016)	Biostratigraphic & lithostratigraphic analysis	Depending on the distribution of recognized algae and foraminifers, only 3 biozones were recognized, such as pseudochrysalidina conica, calpionella alpine, and pseudocyclammina lituus-Trocholina sp. assemblage zones
Jamalian and Adabi	(2014)	Geochemistry and microfacies analysis	Fahliyan Formation was build up in a shallow warm-water environment in the age of early Cretaceous; its porosity and original texture were mainly influenced by dissolution, dolomitization and cementation diagenetic processes

**Fig. 2** Lithostratigraphic chart of the Mesozoic of the Zagros Basin (Alavi 2007)

have been divided into five different trends by thrust faults with the presence of salt domes (Bordenave 2003). The impact of these salt domes on the anticline can be used as the most important structural features in this region.

Even, the trapping style of hydrocarbon in this region varies to that formed in Embayments and regions because its source rock developed more gas and less oil accumulations (Schwarzbauer et al. 2000; Amijaya et al. 2006).

Samples and methods

Three crude oil samples were collected from KG-031 well in 'KG' oilfield and F9A-3H and F15-3H wells in 'F' oilfield in the northwest of the Parisian gulf for analysis in this study (see Fig. 1). The density (API) of oil samples was measured at 15.6 °C (60 °F) using an Anton Paar density meter (Anton Paar DMA5300M). In addition, the kinematic viscosity of the crude oils was measured at three different temperatures (10 °C, 20 °C and 40 °C) using VST-2400 viscometer.

Furthermore, selected crude oils were subjected to a series of analytical tests (Sulfur content, nickel and vanadium ratios, pour points, cut points and 'SARA' = Saturate, Aromatic, Resin and Asphaltene fractionation). The sulfur, carbon, hydrogen and nitrogen contents of oil samples were measured using a LECO CHNS-932 analyzer at 1000 °C burning temperature in the presence of oxygen. The nickel and vanadium ratios were measured using atomic absorption spectrophotometry. Thus, 2–3 g of each sample were treated overnight with an equal weight of concentrated sulfuric acid with the purity of 95–99% total mass. After recovering this acid by heating on a hot plate with gradual temperature increase, the samples were ashed with gradually increased temperature up to 450 °C and were dried overnight. The ashed samples were then heated on a hot plate for 10–20 min and 1 ml of concentrated aqueous HCl was added to the solution until it was clear and yellow, and subsequently filtered. The resulting solution was diluted and analyzed for Ni and Va on a Varian spectrum 14.4–30 atomic absorption spectrophotometer.

In the SARA analysis, the asphaltene fraction was precipitated by addition of *n*-heptane with a purity of 98%. The asphaltene-free samples and the saturate fractions were analyzed via gas chromatography in order to obtain the distribution of alkanes and Pristane (Pri), Phytane (Phy), C17, C18, etc. For conducting the gas

chromatography (GC) analysis, we followed the procedures drawn by Schwarzbauer et al. (2000), Amijaya et al. (2006). The asphaltene-free sample was injected into the GC in a heated port linked with a fused-silica cross-linked methyl silicone gum with the instrument specifications and test conditions shown in Table 2. Finally, the complementary geochemical analysis and biomarker finger printing were carried out on of saturate and aromatic fractions of oil samples using gas chromatography–mass spectrometry (GC–Mass) analysis with the instrument specifications and test conditions shown in Table 3.

Results and discussion

Hydrocarbon source

Table 4 shows the various measured physical properties of crude oils. The sulfur content and V/(Va + Ni) ratio of the analyzed samples were ranged between 1.32 to 2.24 mol% and 62 to 73 mol%, respectively. According to Lewan (1984), when the Va/(Va + Ni) ratio is greater than 0.5%, it is an oxygen-free environment. Va/(Va + Ni) ratios for the oil samples from the KG-031, F15-3H, and F9-3H wells were 72.5%, 67.7% and 72.3%, respectively. When considering these ratios in combination with the sulfur contents, it can be concluded that all oil samples were formed from a single source rock which deposited in a reducing environment with calcareous lithology as shown in Fig. 3 (Lewan 1984).

When plotting sulfur content versus API gravity of Fahliyan oils, in accordance with Mashhadi and Rabbani (2015),

Table 2 Instrument specifications and test conditions of gas chromatography (GC) analysis

Parameter	Range
GC model	Varian-Cp-3800
Capillary column	CP-Sil 5 CB (25 m × 25 μm)
carrier gas	Helium (5 ml/min)
Detector	FID
Sample size	0.2 μl
Injector temp	310 °C
Oven temp	40 °C (4 °C/m increasing rate)
Final temp	290 °C
Detector temp	310 °C

Table 3 Instrument specifications and test conditions of gas chromatography–mass spectroscopy (GC–MS) analysis

Parameter	Range
GC model	Varian-Cp-3800
Capillary column	DB-5 (20 m × 25 μm)
carrier gas	Helium (5 ml/min)
Detector	FID
Sample size	0.1 μl
Injector temp	280 °C
Oven temp	50 °C (5 °C/m increasing rate)
Final temp	280 °C
Detector temp	310 °C
Mass spectrometer	DSQ model
Ione source	70 eV
scanning mode	SIM
Scanning rate	0.4 s/500 amu
Vacuum system	Turbo mol. pump (1 × 10 ⁻⁷ mil torr)

Table 4 Physical properties of oil samples

Well	Viscosity (cP)			Sulfur (wt.%)	Pour point (°C)	Va (ppm)	NI (ppm)	V/(V+NI) (wt.%)	API°	Density (kg/m ³)
	10 °C	20 °C	40 °C							
KG-031	12	8	6.5	2.24	- 40	19	7.2	72.5	30.4	876
F15-3H	41	21	12.2	1.95	- 17	21	10	67.7	33.6	857
F9-3H	18	11	8	1.32	- 23	23	8.8	72.3	31.3	869

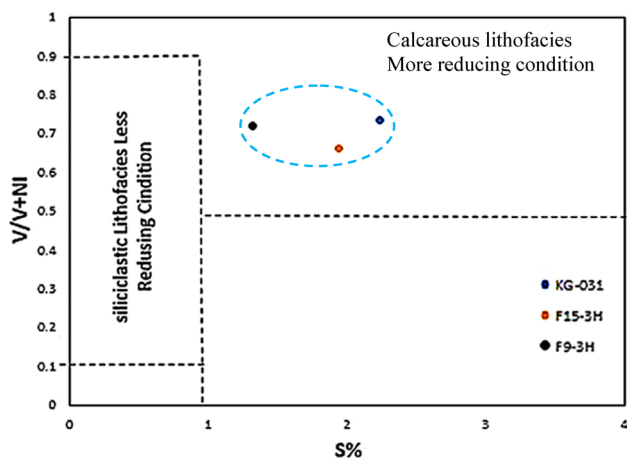


Fig. 3 V/(V + Ni) ratio versus Sulfur content of analyzed crude oils

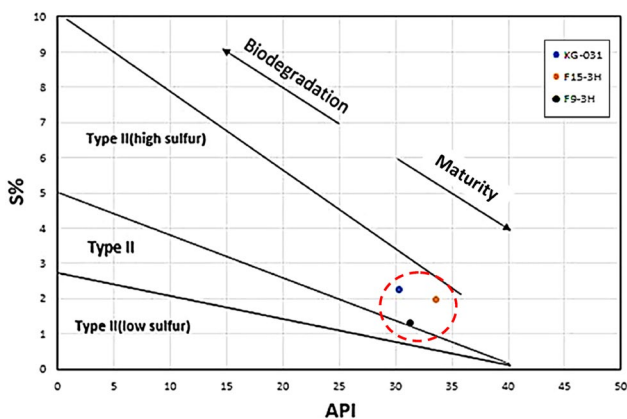


Fig. 4 Cross-plot showing the relation between API gravities and Sulfur contents of the crude oils tested

it is clear that these oils were generated from a matured source rock containing kerogen type II (see Fig. 4). According to the API standard ratios, the Fahliyan oils can be classified as light to medium oils with the presence of high sulfur contents ranging from 1.32 to 2.24 wt.%.

Oil-oil correlations

Bulk properties of the analyzed oil samples indicate that the ‘KG’ oil is a medium oil with an API gravity of 30.4° and the ‘F’ oils are light oils with API gravities ranging from 31.3 to 33.6°, as shown in Table 4. Furthermore, it is clear that the Fahliyan oils consist of different components. Thus, the bulk composition of studied oils is generally dominated by higher concentrations of saturated components (43.39–60.09 mol%) compared to aromatic, resin, and asphaltene compounds (with various ratios from 1.24 to 38.62 mol%) (see Table 5). The ratio of saturated to aromatic hydrocarbon is higher than 1.0 for all samples, and the saturate, aromatic, resin and asphaltene fractions are small with ratios < 70 mol%. Hence, they were buried at shallow to medium depth (Liu et al. 2017).

Fahliyan oils extracted from two extracted oil reservoirs were further investigated using cut point analysis. In this way, they were distilled 5% intervals with respect to yield percent of boiling points. When comparing the cut points of these oils, it can be seen that there is a close resemblance in the distribution patterns between the oil samples from F15-3H and F9-3H wells (see Fig. 5). However, the KG-031 oil shows a clear difference from 15% yield onward, which indicates the presence of heavier oil components when compared to the ‘F’ oil. This is true with respect to the weight percent of hydrocarbon fractions and API degrees discussed above.

Table 5 Geochemical data and gross chemical composition of crude oils tested

Oil Sample	Oilfield	Saturates (mol%)	Aromatics (mol%)	Sat/Aro	Resins (mol%)	Asphaltenes (mol%)
KG-031	KG	49.05	38.62	1.27	8.69	3.64
F9A-3H	F	60.09	20.43	2.94	18.24	1.24
F15-3H	F	43.39	33.51	1.28	17.64	5.46

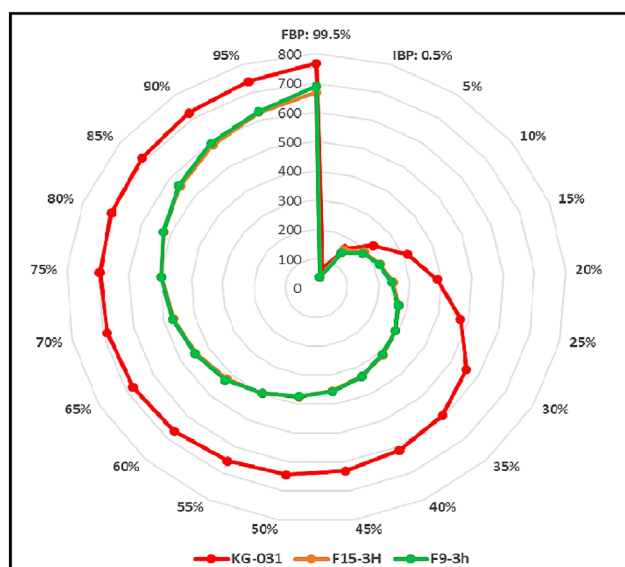


Fig. 5 Comparison distribution of cut points of crude oils tested

DHA results were divided into five various groups including paraffins, aromatics, iso-paraffins, naphthenes and olefins (see Fig. 6). As it can be seen, the percentage of paraffinic hydrocarbon from both F15-3H and F9-3H wells was higher than what measured from the ‘KG’ oil. However, the aromatic ratio in ‘KG’ oil is higher compared with the ‘F’ oils. In summary, it can be stated that the Fahliyan oils are mainly composed of paraffinic and iso-paraffinic hydrocarbons with low carbon content.

Organic matter and depositional environment

Gas chromatographic (GC) results of the analyzed crude oils are shown in Fig. 7 and summarized in Table 6. Generally, the normal alkane distribution patterns can be seen in the range between C_8 and C_{31} while maximum peaks are basically identified between C_{10} and C_{14} . According to

Mobarakabad et al. (2011), it was remarked that alkanes smaller than C_8 might be lost during oil sample handling. Specifically, a normal alkane distribution was identified for F9-3H oil (Fig. 7a), which has a Pr/Ph ratio of 0.72; while the pattern distribution of F15-3H oil indicates that the oil is mainly composed of n-alkanes with a low amount of isoprenoids alkane, Fig. 7b. The third crude oil from KG-031 well, however, has a complete distribution of the normal alkanes (Fig. 7c), but its Pr/Ph ratio (0.89) is slightly higher. This normal alkane distribution pattern is a key feature of source rocks containing a high content of marine organic matter input (Peters et al. 2005).

The depositional environment and organic matter source are commonly identified using the Pr/Ph ratio; Pr/Ph ratios higher than 3 indicate oxic deposition with terrigenous organic matter input, and low Pr/Ph ratios (< 1) indicate a highly saline and reducing depositional setting (Hunt 1996; Li et al. 1999). As the Pr/Ph ratios of the Fahliyan oils are less than 0.9, we conclude that the ‘F’ and ‘KG’ oils have been generated in a saline and reducing environment.

From Pr/ nC_{17} versus Pr/ $n-C_{18}$ ratio cross-plots shown on Fig. 8, the maturity, organic matter input and dependent on Tissot and Welte (1978), Peters et al. (2005), depositional environment of the Fahliyan oils were identified. Thus, in combination with the normal alkane distribution discussed above, it can be concluded that non-biodegraded hydrocarbons are present in the Fahliyan Formation. On the basis of having Pr/ nC_{17} and Pr/ $n-C_{18}$ ratios relatively low (ranging from 0.21 to 0.41), the analyzed oils might be originated from marine carbonate source rocks, and deposited in reducing environments with marine organic matter. This is true with what Abdullah and Connan (2002) was reported. Considering the small variance in these ratios, it is clear that these crude oils were generated from a single source rock with a high maturity level.

Based on the concept of Seifert and Moldowan (1980), the ternary C_{27} , C_{28} , and C_{29} regular steranes diagram was also considered to determine and confirm the depositional

Fig. 6 DHA group type distribution based on light oil fractions

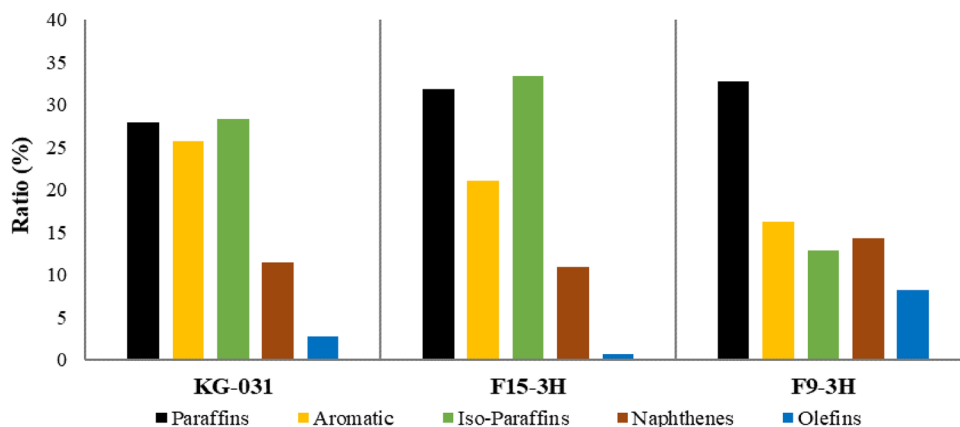


Fig. 7 Gas chromatograms of Fahliyan crude oils from ‘F’ and ‘KG’ fields

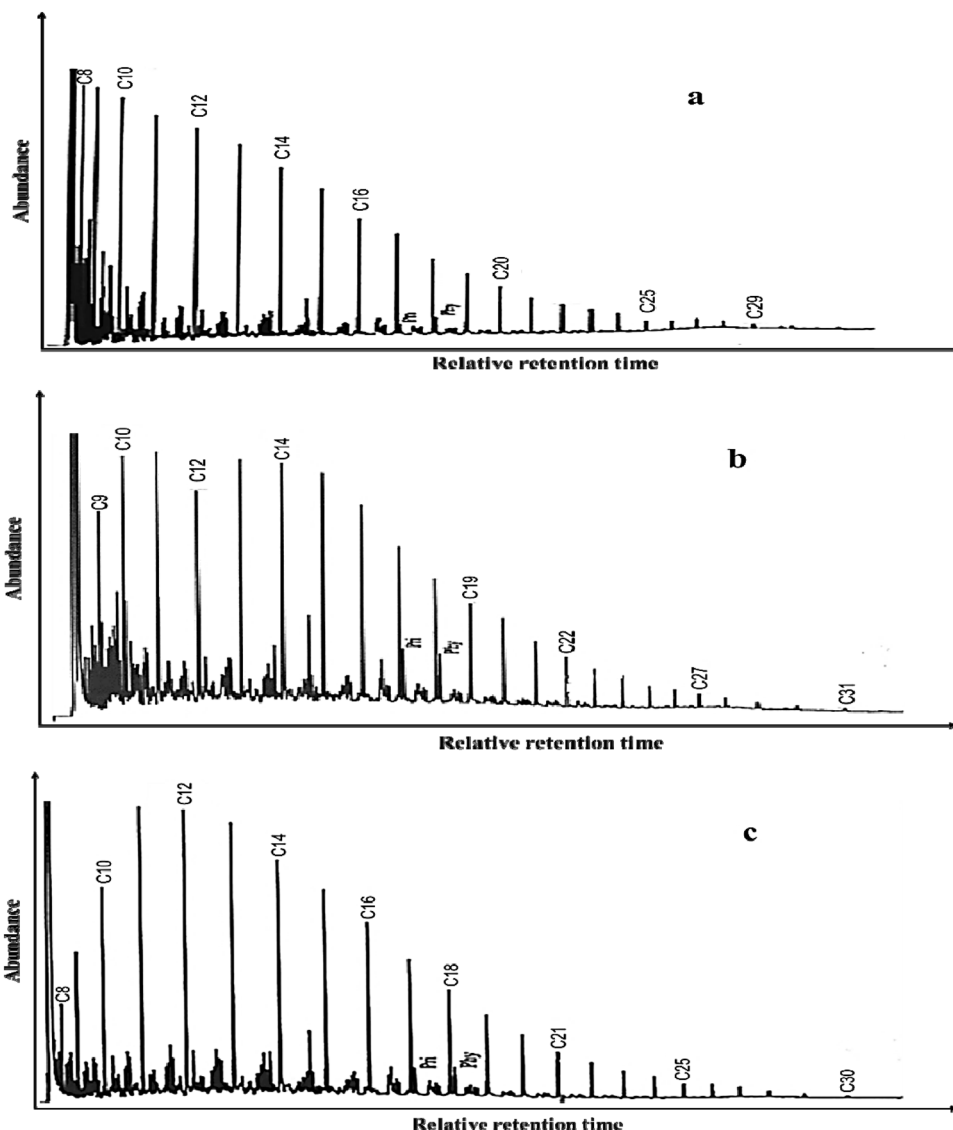


Table 6 GC analysis results for saturate fraction of oils

Well	$\left[\frac{\text{Pri}}{\text{C17}} \right]$	$\left[\frac{\text{Phy}}{\text{C18}} \right]$	$\left[\frac{\text{Pri}}{\text{Phy}} \right]$	$\left[\frac{\text{C19}}{\text{C20}} \right]$	$\left[\frac{\text{C21}}{\text{C22}} \right]$	$\left[\frac{\text{C23}}{\text{C24}} \right]$	$\left[\frac{\text{C25}}{\text{C26}} \right]$	$\left[\frac{\text{C27}}{\text{C28}} \right]$	$\left[\frac{\text{C29}}{\text{C30}} \right]$	$\left[\frac{\text{C31}}{\text{C32}} \right]$	CPI
KG-031	0.31	0.41	0.89	1.08	1.13	1.08	1.07	1.14	1.13	1.05	1.00
F9-3H	0.21	0.35	0.72	1.23	1.14	1.13	1.09	1.14	1.11	1.01	0.94
F15-3H	0.31	0.32	0.71	1.17	1.11	1.10	1.09	1.11	1.09	0.99	0.93

environment of the crude oil source rocks (Fig. 9). Clearly, the oil samples from both fields were derived from a single source rock, which was deposited in an open marine environment. The C_{27} , C_{28} , and C_{29} steranes have approximately the same concentrations in the ‘F’ oils ($C_{27} \approx 23$, $C_{28} \approx 77$ and $C_{29} \approx 52$), and varied concentrations in ‘KG’ oil ($C_{27} \approx 19$, $C_{28} \approx 81$ and $C_{29} \approx 50$). Meanwhile, the C_{27} , C_{28} and C_{29} in

all analyzed oils have the same concentrations trends, i.e., $C_{28} > C_{29} > C_{27}$.

In addition, dibenzothiophene/phenanthrene (DBT/P) versus Pr/Ph ratios indicate the lithology of the source rock (Hughes et al. 1995), for instance carbonate rocks have a ratio higher than 1 and shales have ratios less than 1. The analyzed crude oils fall in Zone-1B of the cross-plot, which indicates that the lithology of the source rock is marine carbonate and marl (Fig. 10). However, DBT/P

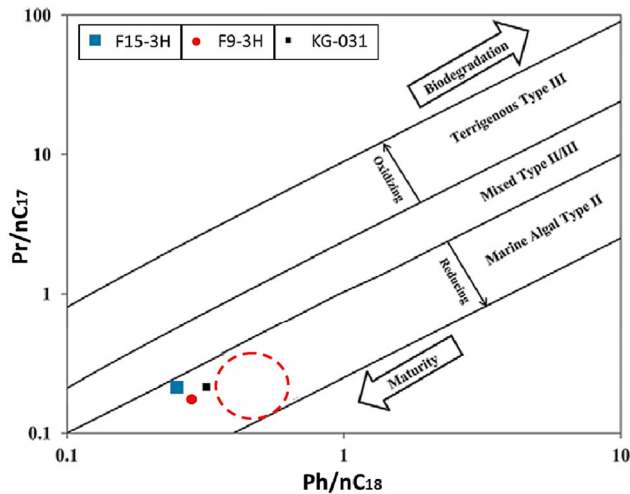


Fig. 8 Pr/n-C₁₇ versus Ph/n-C₁₈ ratio cross-plots for the studied crude oils

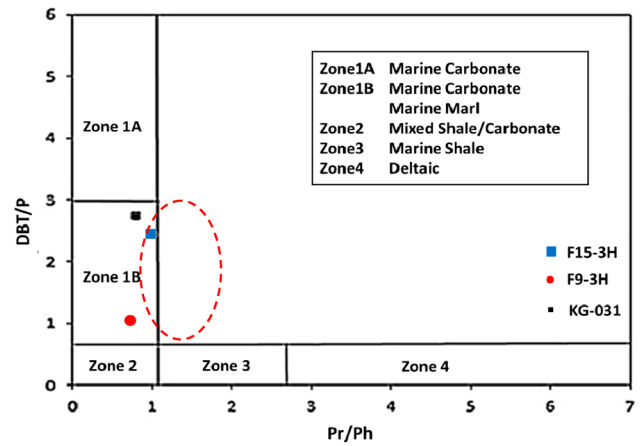


Fig. 10 Pr/Ph ratio versus DBT/P ratio cross-plot for studied oil samples (Seifert and Moldowan 1980)

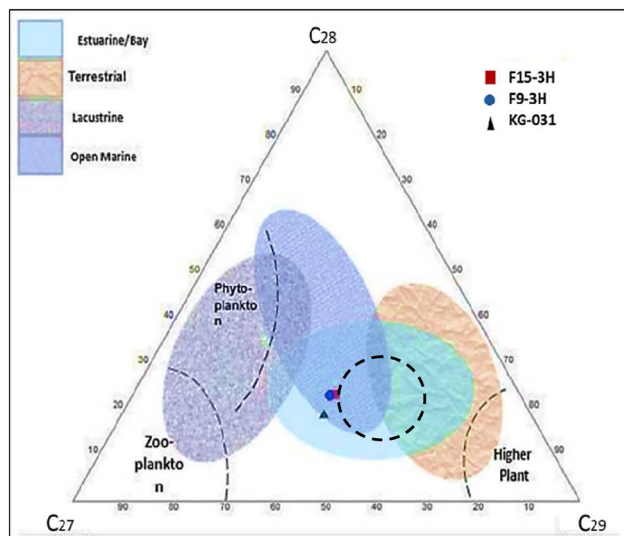


Fig. 9 Ternary diagram illustrating the relative C₂₇, C₂₈, and C₂₉ sterane contents of the studied oil samples

is greater than 1 for the Fahliyan crude oils, this ratio ranges from 1.2 to 2.8, representing the various source rock facies. Higher DBT/P ratios for F15-3H and KG-031 oils thus show higher carbonate content in their source rocks than in case of F9-3H oil.

Thermal maturity

For evaluation of oil maturity levels, steranes and hopanes concentrations were analyzed (Table 7). Thus, two plots were drawn, first the C₂₉ $\alpha\alpha$ 20S/(20S + 20R) versus C₂₉ $\alpha\beta\beta$ /($\alpha\alpha\alpha$ + $\alpha\beta\beta$) ratios, and

second the methyl dibenzothiophene (MDBT) versus C₃₂ 22S/(22S + 22R) ratios as shown in Fig. 11 (Peters and Moldowan 1991; Peters et al. 2005). All oil samples are in the mature oil interval as the C₂₉ $\alpha\alpha$ 20S/(20S + 20R) and C₂₉ $\alpha\beta\beta$ /($\alpha\alpha\alpha$ + $\alpha\beta\beta$) ratios range from 0.47–0.48 to 0.52–0.57, respectively (Fig. 11a). Although the methyl dibenzothiophene (MDBT) ratio of F15-3H oil is lower when compared to the other oils, Fig. 11b also indicates that the source rock of the Fahliyan crude oils is well matured (Peters et al. 2005). Figure 12 expresses the relationship between 20S/(20S + 20R), $\beta\beta$ /($\alpha\alpha$ + $\beta\beta$) and 22S/(22S + 22R) for C₂₉ and C₃₂ steranes and oil maturity levels. With increasing maturity, the 20S/(20S + 20R) and $\beta\beta$ /($\alpha\alpha$ + $\beta\beta$) C₂₉ sterane ratios increase until equilibrium ratios are reached (Peters et al. 2005). Additionally, the 22S/(22S + 22R) ratio of C₃₂ homohopanes increases during maturation up to 0.6 (Mashhadi and Rabbani 2015). As the measured 20S/(20S + 20R) and $\beta\beta$ /($\alpha\alpha$ + $\beta\beta$) for C₂₉ sterane and 22S/(22S + 22R) for C₃₂ homohopanes from the Fahliyan Formation vary from 0.47 to 0.6, it confirms that the analyzed oil are highly mature (Fig. 12).

Furthermore, the Hopane parameters were analyzed to further identify the maturity level (Peters and Fowler 2002). We found that the variation of C30Ts/C30H content in 'F' and 'KG' field oil is small (ranging from 0.22 to 0.26), Table 7a. The Gammacerane/C₃₀₋₃₁ ratio varies widely (from 0.16 to 0.69) and it is less than 1 for all samples studied. Moreover, the Ts/(Ts + Tm) ratio lies between 0.29 and 0.45, which indicates that Fahliyan is a mature oil (Peters et al. 2005).

Table 7 Steranes and Hopanes biomarkers obtained (m/z 217) and Terpanes (m/z 191) measured for the selected oils

Well/Reservoir	C19/C23t	C22/C21t	C22/C24t	C24/C23t	C24Te/C23t	C24Te/C26t	C24Te/C23t	C24Te/C26t	C24Te/C23t	C24Te/C26t	C24Te/C23t	C24Te/C26t	C24Te/C23t	C24Te/C26t	C24Te/C23t	C24Te/C26t	C24Te/C23t	C24Te/C26t	Ole/C30H	C30Ts/C30H	Gam/C30H	Gam/C30HR	C35HS/C34HSC35	Homo-hopane Index	Ts/(Ts+Tm)	Mor/C30H	C32S/(S+R)							
<i>(a) Hopanes</i>																																		
F15-3H	0.35	0.93	0.52	0.47	0.47	1.57	16.9	0.15	0.24	1.58	0.01	0.00	0.26	0.29	0.69	0.64	0.2	0.45	0.07	0.58														
F9-3H	0.42	0.75	0.75	0.42	0.54	1.42	7.67	0.18	0.25	1.53	0.01	0.00	0.25	0.17	0.39	0.86	0.17	0.37	0.07	0.58														
KG-031	0.23	0.95	0.76	0.29	0.75	1.4	7.82	0.14	0.19	1.37	0.01	0.00	0.22	0.16	0.32	1.24	0.15	0.29	0.08	0.59														
<i>(b) Steranes</i>																																		
F15-3H	36.5	23.7	39.8	0.4	0.28	0.88	0.477	43.4	21.9	43.6	1.27	0.12	0.12	0.12	0.12	0.12	0.12	0.12	9.6	79.8	10.6													
F9-3H	37.6	23.4	39	0.35	0.28	0.92	0.479	35	23	42	1.2	0.16	0.16	0.16	0.16	0.16	0.16	0.16	11.7	71.6	16.8													
KG-031	41	19	40	0.25	0.15	0.93	0.481	32.25	23	44.7	1.39	0.12	0.12	0.12	0.12	0.12	0.16	9.2	76.84	13.96														

Oil-source rock correlation

As discussed above, and based on the GC hopane/sterane analysis (see Fig. 13), the oil samples were identified to stem from a carbonate-rich source rock with marine organic matter input, which deposited under marine reducing conditions to at high maturity level. In addition, the relationships between Pr/Ph, C₂₉/C₃₀ hopane, DBT/P, Pr/nC₁₇ and Ph/nC₁₈ illustrate that the carbonate source rock was deposited in a reducing marine environment. Considering the hopane/sterane parameters on the star diagrams, the Fahliyan oils from the two different fields are very similar and show almost the same hopane/sterane ratios (Fig. 13). This also suggests that the analyzed crude oils were formed from the same single source rock.

Diasterane/sterane ratios are commonly used to distinguish the type of source rock from which the hydrocarbons were generated (i.e., carbonate or clastic), Seifert and Moldowan (1980), Gratz et al. (2011); low diasterane/sterane ratios indicate anoxic clay-poor or carbonate source rock, while high diasterane/sterane ratios suggest that the hydrocarbon derived from clay-rich source rocks. Furthermore, in some crude oils, high thermal maturity and/or heavy biodegradation results in a high diasterane/sterane ratio (Zumberge 1984; Alavi 2007). In this work, the C₂₇ Dia/(Dia + Reg) versus C₂₉/C₃₀ hopane ratios have been plotted to identify the depositional environment of the oil source rocks (Fig. 14a). It can be seen that the Fahliyan oils have low diasterane/sterane ratios, suggesting that they were formed in carbonate source rocks. C₂₂/C₂₁ and C₂₄/C₂₃ tricyclic terpane ratios can also be used to identify the type of source rock; for shaly source rock the C₂₂/C₂₁ tricyclic terpane ratio is less than 0.5, whereas ratios higher than 0.5 indicate carbonate source rock (Grantham and Wakefield 1988). Thus, the C₂₂t/C₂₁t ratios (ranging from 0.75 to 0.95) versus the C₂₄t/C₂₃t ratios (ranging from 0.29 to 0.47) were plotted in Fig. 14b. Clearly, all oil samples were generated from marl-carbonate source rocks, consistent with the hopane analysis.

In order to determine the age of the source rock, C₂₈/C₂₉ regular sterane ratios were examined; ratios < 0.7 indicate Paleozoic and Jurassic ages, while ratios > 0.7 indicate Upper Jurassic to Cretaceous-Tertiary ages (Zumberge 1984). Here, %C₂₈ regular sterane versus C₂₈/C₂₉ sterane ratios ranged from 0.38–0.61 and 18.55–25.87, respectively. Thus, these crude oils accumulated in deeper reservoirs, and originating from late Jurassic or early Cretaceous rocks (Fig. 15).

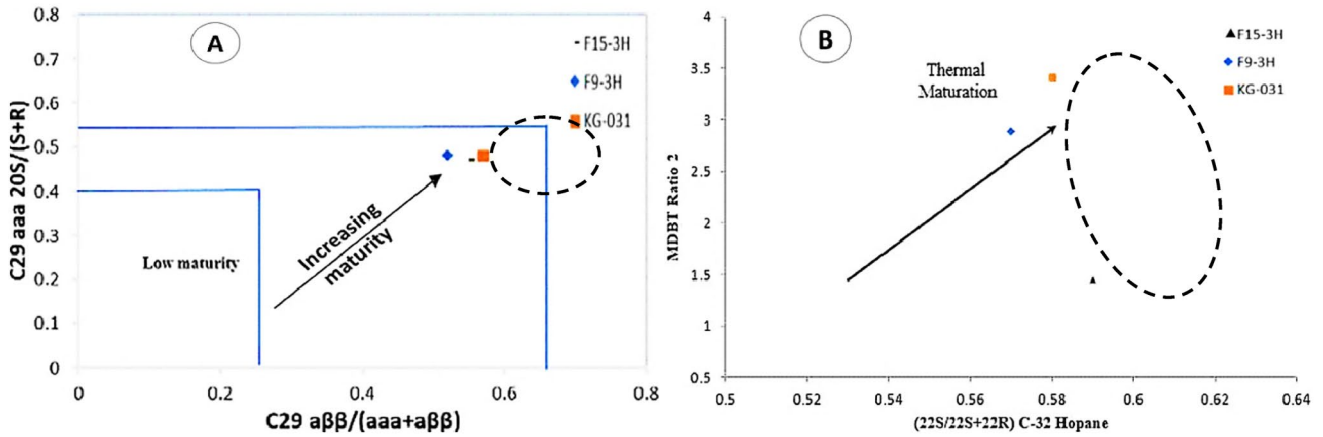


Fig. 11 Cross-plots showing the maturity level of the oil samples studied based on a C_{29} $aaa/(aaa + \beta\beta)$ versus C_{29} $\beta\beta/(\beta\beta + \alpha\alpha)$ steranes ratio and b MDBT 2 versus $C_{22S}/(22S + 22R)$ C_{32} Hopanes ratio (Peters et al. 2005)

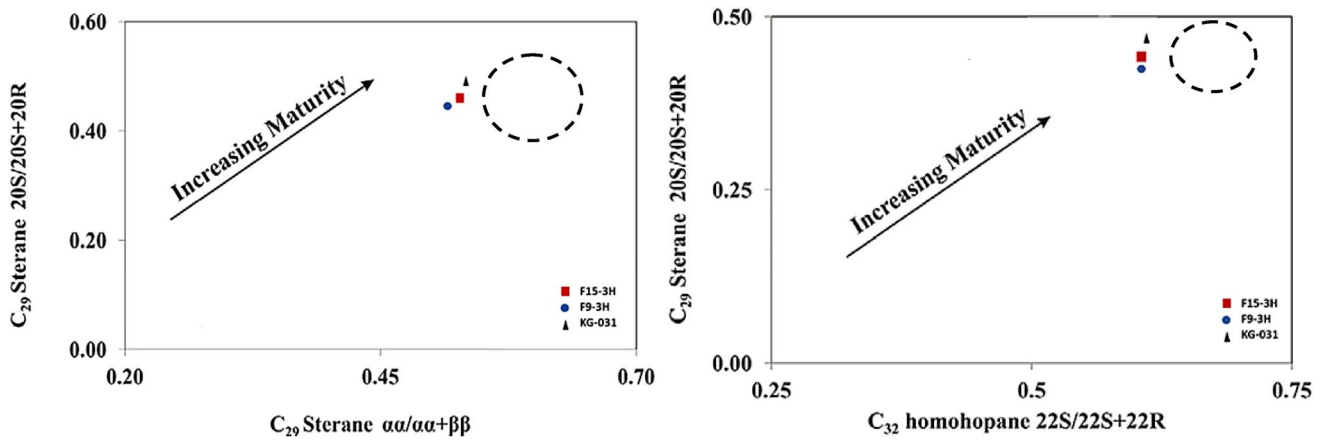


Fig. 12 Comparison of the thermal maturity of the studied Cretaceous source rocks and Fahliyan oils

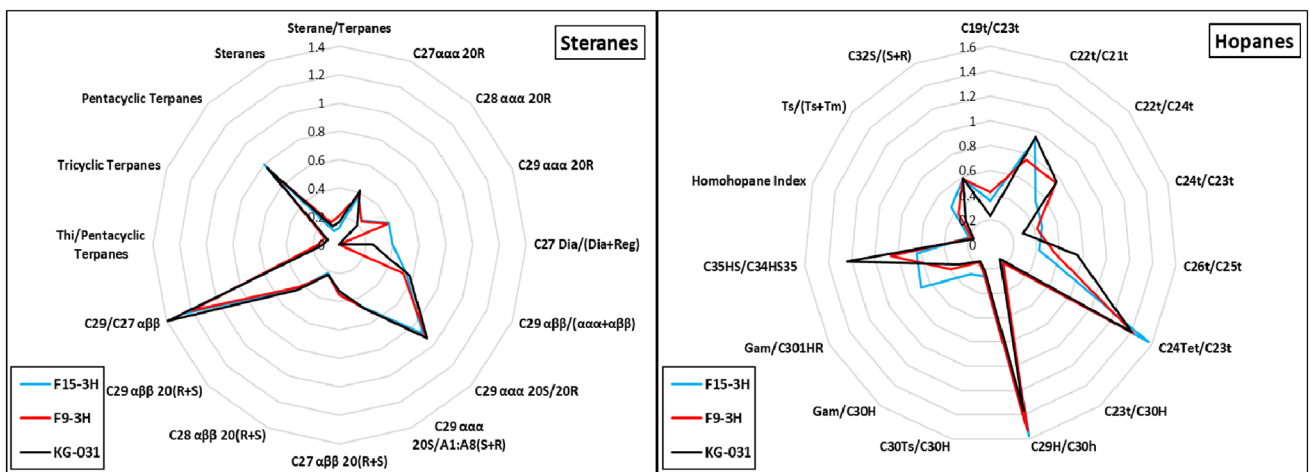


Fig. 13 Steranes (left) and Hopanes (right) parameters of the studied oils

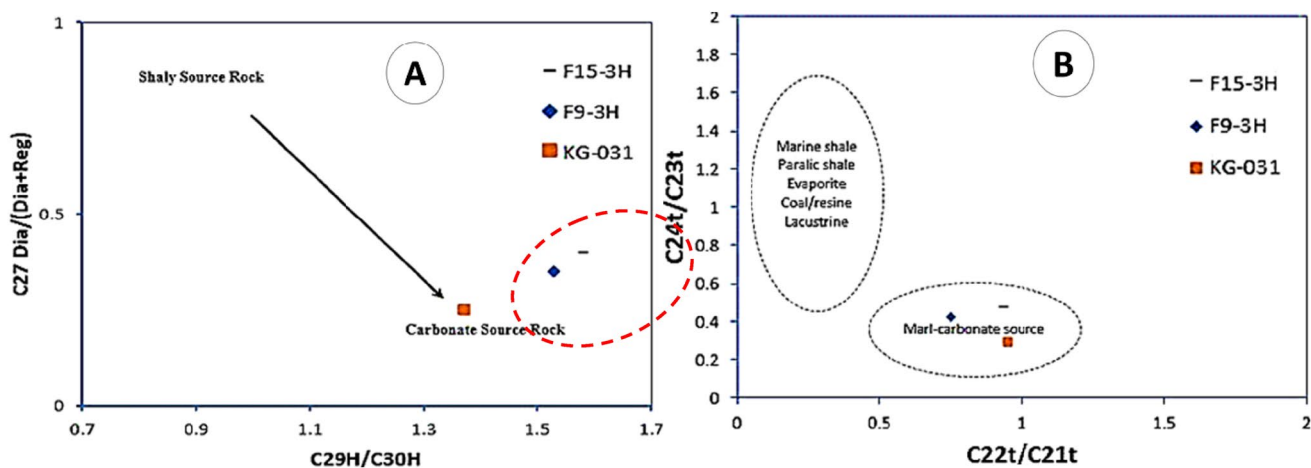


Fig. 14 Biomarkers cross-plot of a C_{27} diasterane/ C_{27} diasteranes versus C_{29} norhopane/ C_{30} hopane ratio (29H/30H), and b $C_{24}t/C_{23}t$ versus $C_{22}t/C_{21}t$ ratios of tricyclic terpanes

Conclusions

Multiple major Fahliyan reservoirs have been identified in the NW region of the Persian Gulf (Adabi et al. 2010; Abedpour et al. 2016; Feghi et al. 2012). However, the origin, source rock and maturity of the oil within this formation are only poorly understood. Thus, in this study, a detailed petroleum geochemical analysis of three selected crude oils from Fahliyan reservoirs in SW Iran has been performed. In this way, it was verified that the hydrocarbon in these reservoirs was generated from the late Jurassic

to early Cretaceous. In addition, the oil bulk properties and biomarkers were characterized, and an oil–oil and oil–source rock correlations study was carried out between the three crude oils and the source rock. From this extensive analysis, it is clear that all oils are nonbiodegraded with high maturity, have a high sulfur content and were generated in carbonate source rocks containing kerogen type II. The source rock again was deposited under marine reducing conditions with marine organic matter as the dominant input.

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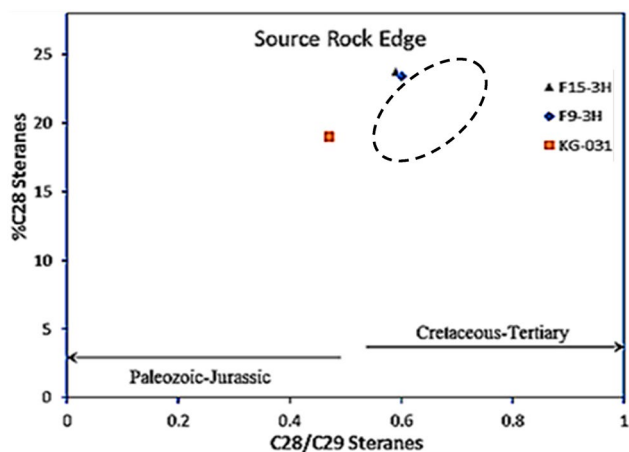


Fig. 15 C_{28}/C_{29} versus $\%C_{28}$ regular sterane ratios indicating the age of the source rock

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