



Trends of Planetary Boundary Layer Height Over Urban Cities of China From 1980–2018

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Huo Y, Wang Y, Paasonen P, Liu Q, Tang G, Ma Y, Petaja T, Kerminen V-M and Kulmala M (2021) Trends of Planetary Boundary Layer Height Over Urban Cities of China From 1980–2018. Front. Environ. Sci. 9:744255. doi: 10.3389/fenvs.2021.744255 Boundary layer height (BLH) is an important parameter in climatology and air pollution research, especially in urban city. We calculated the BLH with a bulk Richardson number (Ri) method over urban cities of China during 1980–2018 using European Centre for Medium-Range Weather Forecasts (ECMWF) ERA-interim data after carefully validation with sounding data obtained from two meteorology stations in eastern China during 2010–2018. The values of BLH between these two types of data have correlation coefficients in the range of 0.65–0.87, which indicates that it is reasonable to analyze long-term trends of the BLH from ERA data sets. Using ERA-interim calculated BLH, we found that there is an increasing trend in the daytime BLH in most cities of eastern China, particularly during the spring season. A correlation analysis between the BLH and temperature, wind speed, relative humidity and visibility revealed that the variability in meteorological parameters, as well as in aerosol concentrations over highly polluted eastern China, play important roles in the development of the BLH.

Keywords: boundary layer height, aerosol, visibility, trend, meteorology parameter

INTRODUCTION

Planetary boundary layer (PBL) is the turbulent layer of the troposphere and plays an important role influencing the concentration, transport and diffusion of atmospheric compounds (Stull, 1988). The PBL height is most commonly identified as an inversion in the potential temperature and dewpoint, or as a peak value in the low-level wind speed. Air pollutants are concentrated in the PBL, and the daily variation of the PBL is crucial in interpreting the diurnal variation of air pollutants in areas with anthropogenic emissions (Kaser et al., 2015; Petäjä et al., 2016; Sun et al., 2013; Ma et al., 2020; Xian et al., 2021). In general, the height of planetary boundary layer varies spatially and temporally, typically ranging from about a hundred meters to some kilometres. The PBL height can be as high as 5 km over a desert in mid-summer due to strong surface heating, while being as low as 50–100 m during a night over land under clear sky and light wind conditions (Zhang et al., 2013; Chen et al., 2016).

Radio sounding measurements, providing the profiles of wind, humidity and temperature, are typically used in weather stations to determine the height of PBL (Seidel et al., 2010). In general, such sounding is conducted twice a day at 08:00 and 20:00. However, the method needs a lot of man power to be conducted in long-term measurements, in addition to which it cannot provide information about the diurnal evolution of the boundary layer height due to the limited time resolution. Seidel et al (2010) compared seven methods of calculating the PBL height from a 505-station global radiosonde data set, and also attempted to quantify aspects of structural and parametric uncertainty in the values of the PBL height. Zhang et al (2013) estimated trends of the PBL height over 25 stations in Europe using radiosonde data from 1973-2010, and found significant increases in the daytime PBL height at most of the stations. Guo et al (2016) compared the PBL height between the radiosonde and reanalysis data from 2011-2015 in China, and found a good agreement between these two data sets. Remote sensing, as an alternative method, is widely used to determine the PBL height from a backscattering signal during the past decades. Wang et al (2012) derived the PBL height at two locations of Lanzhou, China, using a micro-pulse lidar and microwave

radiometer, and found that these two results were consistent with each other during strong convective situations. Zhang et al (2016) compared the PBL height from CALIOP and radiosonde data in Beijing and Jinhua, and obtained correlation coefficients of 0.59 and 0.65 at Beijing and Jinhua, respectively, between these two kinds of data.

Very recently, Guo et al (2019) found an increasing trends of the PBL height over China during 1976–2003, but a non-uniform decrease after 2004. Capital cities of China suffer the most from air pollution, but the variability and trends of the PBL height are unknown in those regions.

With the rapid economic growth in China, air pollution has become an issue with large influences on climate, human health and visibility degradation during the last decades (Wang et al., 2015). Wang et al (2020), Ding et al (2016) and Petäjä et al (2016) and other work (Ma et al., 2020) found that atmospheric air pollution can be enhanced by boundary layer-aerosol interactions, characterised by higher aerosol concentrations and shallower boundary layer heights. Therefore, a better knowledge of the long-time variation of the PBL height in China is essential when evaluating variations in aerosol concentrations and effects of regional climate change. However, only a few studies reported a long-



Station(°N,°E)	DJF					MAM					JJA					SON			
	z	ws	т	сс	RH	z	ws	т	сс	RH	z	ws	т	сс	RH	z	ws	т	сс
Beijing(40,116)	0.7	4.8	4.5	0.7	-0.1	7.0	4.8	6.9	-2.2	-1.9	1.7	-9.2	4.4	-2.3	-1.4	-0.5	-1.7	3.2	0.2
Shanghai(31,121)	-0.1	-4.7	4.1	-1.8	-1.0	4.7	0.5	10.2	-3.8	-3.3	1.3	-7.0	6.7	-1.5	-1.9	2.8	-5.5	6.8	-3.1
Guangzhou(23,113)	1.5	-2.3	2.1	-0.6	-1.1	4.5	-8.6	4.5	-2.9	-2.1	1.2	-8.5	1.7	-0.9	-1.3	0.7	-16.6	3.0	0.5
Shenzhen(22,114)	-1.0	-6.8	2.6	-0.0	-1.1	0.1	-13.0	4.4	-2.6	-1.7	-1.2	-17.5	1.9	0.1	-1.2	-0.4	-22.8	3.5	1.0
Tianjin(39,117)	1.3	3.3	4.3	0.6	0.0	2.2	3.1	5.3	-1.8	-1.5	-0.1	-6.2	3.5	-2.1	-1.2	-2.5	-3.6	2.5	0.2
Chongqing(29,106)	0.9	0.5	1.8	-0.6	-0.8	2.1	1.1	6.1	-1.2	-1.6	4.3	-1.2	4.7	-0.8	-2.0	3.0	0.7	5.1	-1.9
Haerbin(45,126)	-0.8	-14.2	0.8	-3.2	-1.3	0.4	-0.8	1.4	0.9	0.6	-0.1	-7.9	4.0	-2.3	-0.9	4.1	-10.9	6.9	-1.0
Changchun(44,125)	-1.5	-11.7	0.1	-1.4	-0.9	3.4	-0.6	2.8	-0.3	0.0	0.5	-6.5	4.3	- 2.0	-1.1	4.6	-11.2	6.6	-0.5
Shenyang(41,123)	1.1	-5.5	3.5	-0.7	-1.6	2.2	-2.5	3.3	-0.3	-0.3	-0.2	-13.0	2.6	-1.4	-0.1	0.6	-6.5	5.1	-0.2
Taiyuan(37,112)	0.2	-3.0	3.1	0.8	0.0	1.4	8.0	5.5	-1.9	-1.7	9.4	-10.5	3.5	-2.4	-1.5	1.8	-2.7	1.8	0.8
Shijiazhuang(38,114)	0.6	5.8	2.9	0.4	0.4	5.3	6.8	5.8	-2.0	-1.9	1.2	-1.6	2.6	-2.2	-0.8	-0.9	0.6	1.4.	0.3
Lanzhou(36,103)	-6.1	2.8	2.1	0.4	3.6	1.2	2.2	5.6	-0.3	1.3	-2.7	1.7	5.1	0.7	1.0	-8.6	3.0	3.4	4.6
Xian(34,108)	-1.5	-3.2	2.2	0.3	0.2	2.1	1.1	7.4	-2.3	-1.8	1.8	-3.5	4.9	-2.5	-1.7	-1.5	-0.8	4.0	-0.6
Jinan(36,120)	-2.1	-10.7	4.0	0.0	-0.3	2.5	6.3	4.4	-0.6	-0.9	2.6	-9.9	2.4	-1.4	-0.6	0.3	-8.6	3.8	-0.4
Zhengzhou(34,113)	-0.7	-1.7	2.6	-0.1	-0.8	0.5	6.0	6.5	-2.2	-2.0	4.3	0.0	4.3	-2.5	-2.0	2.2	0.9	2.8	-1.5
Chengdu(30,104)	3.1	-0.0	2.9	-1.7	-1.5	4.5	-4.3	6.9	-1.0	-1.9	1.3	-2.7	3.6	-0.6	-1.2	-0.0	-1.6	3.4	-1.0
Wuhan(30,114)	-0.1	-1.2	2.6	-0.4	-1.5	1.8	-0.6	7.8	-2.0	-2.3	1.7	1.8	3.3	-0.5	-1.4	1.3	-5.7	4.4	-1.4
Hefei(31,117)	-1.1	-2.4	2.5	-0.8	-0.9	2.6	1.3	7.8	-2.5	-2.6	0.8	-4.5	2.6	-0.6	-0.9	0.4	-3.7	4.0	-1.8
Nanjing(32,118)	-0.1	-3.5	3.5	-0.8	-1.1	3.2	2.3	8.0	-3.2	-2.8	-0.1	-8.7	3.4	-1.0	-1.2	1.5	-1.9	4.9	-2.4
Hangzhou(30,120)	0.5	-2.4	4.7	-1.7	-0.9	5.7	-1.3	9.7	-3.6	-3.0	1.5	-9.0	4.8	-0.7	-1.2	1.5	-6.8	6.1	-2.1
Guiyang(26,106)	2.4	-0.2	0.4	-2.1	-1.3	2.3	-7.3	3.8	-1.1	-1.1	1.1	-7.5	0.2	-0.1	-0.9	2.5	-0.2	2.2	-1.9
Changsha(28,113)	0.3	-1.2	2.5	-0.5	-1.4	2.5	-2.9	7.3	-3.0	-2.0	2.9	-1.1	2.0	0.2	-1.4	1.5	-8.5	4.3	-2.3
Nanchang(28,115)	-0.6	2.6	3.7	-0.1	-0.9	1.8	-1.2	8.1	-3.4	-2.2	-0.7	-4.2	1.6	0.4	-0.7	0.9	-3.2	5.8	-2.4
Fuzhou(26,119)	0.6	-0.2	4.8	-1.7	-1.2	2.0	-3.3	6.4	-3.2	-2.2	-0.9	-8.2	2.6	-0.1	0.2	-1.1	-11.0	5.0	-0.6
Kunming(25,102)	-1.4	-1.1	2.2	-0.2	-0.1	-3.6	-14.2	1.5	1.1	0.3	2.7	-4.4	2.8	-0.8	-1.8	3.6	2.1	3.8	-3.7
Nanning(22,108)	1.8	3.5	1.2	-3.9	-1.4	2.6	-4.0	2.8	-1.4	-1.2	-0.7	-4.7	1.8	0.6	-1.7	1.1	-6.8	3.4	-1.3

TABLE 1 Seasonal trends in 14 BJT BLH, wind speed, temperature, cloud cover, relative humidity (z, WS, T, CC, RH with units of m, 10⁻³ m/s, 10⁻² K, 10⁻³, ‰, respectively) during 1980–2018. Boldfaces are significant at 95% confidence interval or greater.

The boldfaces are significant at 95% confidence interval or greater.

RH -0.038 -**16.6** -0.5 0.1 -**1.5** -1.8 -1.3 -0.2 0.1 -0.0

> **4.1** -0.6 -0.3 -1.4 -0.1 -1.3 -0.9 -1.5 -1.1 -0.8 -1.4 -1.4 -0.1

-**2.2** -1.4





term variation of the PBL height over highly polluted cities in China (Guo et al., 2016; Tang et al., 2016).

In this study, we used ERA-reanalysis data combined with radio sounding measurements in determining long-term trends in the BLH in the capital cities of each province in eastern China. Our results provide insights into the decadal variation of the BLH in polluted eastern China during the urbanization period. We also explored the spatial variability in the BLH over China during these years.

DATA AND METHODOLOGY

We chose 26 capital cities of each province in eastern China to analyze the boundary layer heights (BLHs) variation and made a comprehensive analysis. The detailed city list and locations are listed in **Table 1**.

Dataset

The ERA-Interim reanalysis are assimilated results, including model products and various measurements (Dee et al., 2011). Its





model-layers data contain 60 vertical layers (starting with about 25 m near surface, decreasing to about 500 m around 500 hPa), which has been used to calculate the boundary layer height (Seidel et al., 2010; Guo et al., 2016). In this work, the reanalysis data with a horizontal resolution of $0.75^{\circ} \times 0.75^{\circ}$ and time resolution of 6 h were used for the BLHs calculation, after which we interpolated the BLHs to 26 city sites (**Figure 1**). In addition, the surface temperature, relative humidity and wind speed obtained from reanalysis products were used in our analysis to constrain the factors influencing the boundary layer. Visibility data from the strictly quality-controlled Chinese observational data set were also used to analyze influences of aerosols on the boundary layer height.

We applied radiosonde data at two sounding stations in central China during 2010–2018 to validate our results from the ERA-Interim reanalysis calculation. The two stations are a part of China Meteorological Administration L-band secondresolution ground-based sounding network. The instrument used for validation is a digital radiosonde sensor (GTS1), which is developed by the Shanghai Changwang Meteorological Science and Technology Company. The instrument has been used widely in China meteorology stations as routine measurements, and a comparison between the GTS1 and Vaisala RS80 shows that the instrument performance good (Bian et al., 2010). For simplicity, we hereafter call the reanalysis results and sounding results as ERA-BLH and CMA-BLH, respectively.

The Method Used to Estimate BLH

A bulk Richardson number (Ri) method was carefully applied to calculate the BLHs in this work. The Ri method was first put forward by (Vogelezang and Holtslag, 1996). The method is suitable for both stable and convective boundary layer, in which the value of Ri can be expressed as:

$$Ri(z) = \frac{(g/\theta_{vs})(\theta_{vz} - \theta_{vs})(z - z_s)}{(u_z - u_s)^2 + (v_z - v_s)^2 + bu_*^2}$$
(1)



Here the subscript *s* and *z* indicate the surface and specific height, respectively, the variables *g* and θ are gravity acceleration and potential temperature, respectively, *u* and *v* are the horizontal wind velocity components, *u*_{*} is the surface friction velocity, and *b* is a constant. In general, *bu*_{*} is much smaller than the bulk wind shear term (the two other terms together in the denominator) and can be neglected. In our calculations, the surface height was assumed to be 2 m and the surface wind was approximated to be zero, and the first level at which the interpolated Ri was equal to 0.25 was interpreted as the BLH. This method has been proven to be one of the best methods for the climatological analysis of the BLH because of its applicability for both stable and convective boundary layers (Seidel et al., 2010).

RESULTS AND DISCUSSION

Validation of ERA-BLH With CMA-BLH

This section presents an inter-comparison of the ERA-BLH with CMA-BLH data at the stations of Fuyang and Anqing in eastern

China. The soundings were usually launched at 08:00 and 20:00 Beijing time (BJT), so the value of BLH were mainly compared at these two moments. Figures 2, 3 show scatter plots between the ERA-BLH with CMA-BLH in Fuyang and Anging, respectively, the upper panels correspond to 08:00 and the lower panels to 20:00 BJT. The correlation coefficients between two kinds of BLH data were larger than 0.7° at 08:00 for most of the time, which is consistent with earlier results in the Beijing station (Guo et al., 2016). These high correlation coefficients give us confidence that ERA-BLH is an appropriate alternative approach to analyze changes in the boundary layer height. It is interesting to note that comparison results at 08:00 BJT were better than those at 20:00 BJT. A possible reason for this is that the convective boundary layer transmits to a nocturnal stable boundary layer after sunset, so that the existence of a residual layer at 20:00 may increase differences between two kinds of BLH data.

To further validate our analysis method, we compared BLH with meteorological parameters for both ERA data and measurement data from the Anqing and Fuyang stations (**Figure 4**). We found a good consistency between the



observed and reanalysis data in terms of the couplings of meteorological conditions and BLH. The validation presented here gives confidence in applying the ERA data to the analysis of long-term interactions between the BLH and meteorological variability.

Trends of Planetary Boundary Layer Height in Urban Cities of China

The development of the day-time boundary layer height is closely related to solar radiation reaching the surface, so the surface temperature, wind speed, relative humidity, cloud cover and aerosol concentration may be factors that directly or indirectly influence the BLH. A typical daily variation of the BLH showed a maximum at around noon due to the convective activity, whereas low values occurred during nighttime, termed the nocturnal boundary layer. Early mornings and later afternoons were transition periods between a stable boundary layer and convective boundary layer. Therefore, we used the BLH at 14: 00 local time when the convective activity is the main factor influencing the BLH.

A statistical analysis of the BLH, temperature, wind speed, relative humidity and cloud cover in 26 cities in eastern China are shown in **Table 1**. The BLH showed an increasing trend at most of the cities, and especially so in spring and summer. The increasing trend was the highest in the city of Taiyuan, being equal to 11 m/year in spring and 9.4 m/year in summer. The variation of the BLH during the winter season was complex compared with the other seasons, showing an equal number of increasing and decreasing trends between the 26 cities. In every season, there were also cities having a decreasing trend of the BLH. The largest decreasing trend of the BLH was observed in Lanzhou, with a value of -8.6 m/year during the autumn season. It is also worth note that the decreasing trend of the BLH in Lanzhou, significant in summer, autumn and winter seasons, is very likely related to the significant increase of the relative humidity. Using 6.5 years of lidar measurements in Hong Kong, Yang et al. (2013) found a slight decreasing trend of the daily maximum mixing layer height from 2003 to 2009. We could also clearly see that the surface temperature at most of the cities increased during 1980-2015. The increase



of temperature in urban cities is a combined effect of urban heat island effect and regional warming (Cao et al., 2016).

Figure 5 shows the spatial variability in the long-term trend of the BLH across eastern China. There are both positive or negative trends, as one can see, but many of the trends are not statistically significant. Based on **Table 1**, the BLH has higher increasing trends during the spring season than during the other seasons, whereas decreasing trends can be seen for nine stations during the summer seasons and for eight stations during the autumn season. The differences in the values of the BLH between the different seasons are related to differences in meteorological parameters and aerosol concentrations.

Influence of Meteorology

A boundary layer height is closely related to the atmospheric stability, which in turn is driven by the surface temperature, wind speed, relative humidity and aerosol loading (Zhang et al., 2013). **Figure 6** depicts the correlation coefficient between the BLH and surface temperate at 14:00 in 26 cities of eastern China. A positive correlation was obtained for most of the stations during the four

seasons. In particular, the correlation is significant during the summer season; for example, the correlation between the BLH and surface temperature could as high as 0.6° in North China Plain and eastern China, while being lower at around 0.3° at the stations in the south area. In winter, the correlation between the BLH and surface temperature showed a different pattern, being higher in the south area and having even negative values in the north part of China. From **Table 1** we can see that the surface temperature in most of the cities have increased, which is consistent with the fast urbanization and global warming (Cao et al., 2016). Elevated surface temperatures lead to increased sensitive heat flux, and as a result, the boundary layer will develop higher during daytime. In general, higher increases in the surface temperature were associated with larger increases in the boundary layer height in China (**Figure 6**).

A high surface temperature results in a larger sensible heat flux, which lead to high boundary layer height (Stull, 1988). As presented in **Figure 7**, the relative humidity shows negative correlation with the BLH in all the urban stations considered here. **Figure 8** shows the correlation between the wind speed and BLH, with high values over the north region of China indicating



that larger wind speeds favor the development of BLH. However, a negative correlation coefficient was observed in the stations of Guangzhou and HK. A possible explanation for this is that high wind speeds usually correspond to precipitation events in Guangdong Province, as a result of which the development of BLH is suppressed due to lack of upward turbulent kinetic energy (Back and Bretherton, 2005).

Aerosol Effect on BLH

Anthropogenic emissions of aerosol particles and their precursors have decreased the atmospheric visibility in China (Che et al., 2007; Wang et al., 2009, Wang et al., 2019). Since aerosol particles also decrease the amount of solar radiation reaching the surface, they might suppress the development of the boundary layer. The degradation of surface visibility could be a good proxy of light extinction due to atmospheric aerosol. We investigated long-term trends of visibility at 14:00 local time. **Figure 9** shows the correlation between the BLH and visibility during 1980–2015 in eastern China. In general, the correlation coefficient showed positive values, especially in the north area stations, with a clear variability between the different seasons. In particular, the boundary layer height seemed to be influenced more by aerosol particle during the cold season compared with the warm season. The highest correlation between the boundary layer height with visibility were observed at the stations in North China Plain. This indicates that aerosol pollution may have had a significant influence on the development of boundary layer height in North China Plain (Wang et al., 2020), even though it is also possible that the high correlation is due to aerosols being diluted more efficiently under high boundary layer conditions. The lowest correlations were observed in the south areas of China, suggesting more important role of meteorology parameters in determining the boundary layer height compared with aerosol particles.

CONCLUSION

In this study, we analyzed long-term trends of the boundary layer height in eastern China during the 1980–2018 period from ERAreanalysis data, after carefully validating the boundary layer height data from ERA with sounding results. The results showed a slight increasing trends of the boundary layer height in most urban stations of eastern China, especially in the spring season. The correlations of the BLH with the surface temperature, relative humidity, wind speed and visibility were presented and discussed. The results suggest that the increased temperature is the main driving force of the variation in the BLH, while aerosol particles also play important role. Our results revealed an increased trend of daytime boundary layer height under the warming climate, despite an increased loading of atmospheric pollutants.

DATA AVAILABILITY STATEMENT

The raw data supporting the conclusions of this article will be made available by the author, without undue reservation.

AUTHOR CONTRIBUTIONS

YW initiated the study. YH, YW, PP, GT, TP, VK and MK interpreted the data. YH, YW and VK wrote the article.

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