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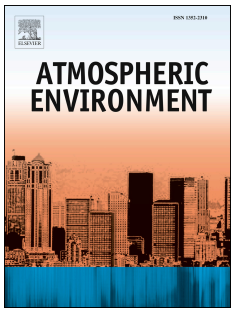

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PII: S1352-2310(19)30442-X

DOI: https://doi.org/10.1016/j.atmosenv.2019.06.051

Reference: AEA 16813

To appear in: *Atmospheric Environment*

Received Date: 30 January 2019

Revised Date: 31 May 2019

Accepted Date: 24 June 2019

Please cite this article as: Nataliya, Z., Olga, Muñ., Evgenij, Z., Maria, G., Jesús, E.-C., Matthew J, B., Jouni, P., Light scattering from volcanic-sand particles in deposited and aerosol form, *Atmospheric Environment* (2019), doi: https://doi.org/10.1016/j.atmosenv.2019.06.051.

Light scattering from volcanic-sand particles in deposited and aerosol form

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Abstract

The light-scattering properties of volcanic sand collected in Iceland are studied here to characterize the sand particles and develop a reference for future remote-sensing observations. While such sand is common in Iceland, the smaller-size fraction can be readily transported by winds and found in the atmosphere at distant locations. The sand appears dark when deposited on a surface due to the high optical absorption of the material. Therefore, atmospheric regions containing such particles during a dust storm may absorb sunlight considerably, causing redistribution of solar energy. Here, we measure the angular scattered-light intensity and degree of linear polarization from the sand. This is done with two experimental apparatuses, the Cosmic Dust Laboratory (CoDuLab) at the Institute de Astrofísica de Andalucía (IAA) and the goniospectropolarimeter (FIGIFIGO) at the Finnish Geospatial Research Institute (FGI). Two scattering-scenarios of practical interest for remote-sensing applications are

considered: (1) single sand-particles suspended in aerosol as an optically thin cloud, and (2) the same particles deposited on a substrate. We also model the measurements with the discrete dipole approximation to estimate the complex-valued refractive index m, where we find that $m \approx 1.6 + 0.01i$ at $\lambda = 647$ nm. Lastly, we present a comparative analysis of the polarimetric response of the sand particles with that reported in the literature for carbon-soot, another highly absorbing atmospheric contaminant.

Key words: volcanic sand; remote sensing; polarimetry; radiometry; photometry; particulate surface; aerosols; light scattering; discrete dipole approximation; refractive index; soot

1. Introduction

One of the largest uncertainties with regard to the interaction of solar radiation between the atmosphere and the Earth-surface, i.e., the Earth's radiative energy budget, is associated with aerosols including dust (Boucher *et al*., 2013). This work studies Icelandic volcanic sand, a significant source of dust in Northern Europe that is close to arctic glacier (Prospero *et al*., 2012). In particular, our study reveals a degree of similarity between the optical properties of Icelandic volcanic sand and black carbon particles; specifically, both particle types strongly absorb solar radiation thus heating the atmosphere and reducing the amount of solar energy received at the Earth surface. The majority of black carbon in the atmosphere has anthropogenic origin, while the volcanic sand discussed here is natural in origin. Given its history of volcanic activity, Iceland has experienced an extended period of absorbing-aerosol effects as evidenced by the abundance of volcanic sand, which has led to climatic influences both locally and regionally across the northern latitudes.

Volcanic sand is one of the main dust-sources in Iceland due to the abundance of rock of volcanic origin. The weather conditions are favorable for active sand-formation via the erosion of solidified lava flows, and thus, about 20% of Iceland is covered with volcanic sand (Arnalds, 2015). Sandy deserts cover large portions of the south coast and glacial margins of the active volcanic zone from the Mýrdalsjökull glacier to areas northeast of the Vatnajökull glacier (Fig. 1). The desert area near the Mýrdalsjökull glacier can be seen in the enlarged MODIS satellite image along with a visible portion of the contaminated glacier. For a more detailed map of the sand dessert distribution, see Arnalds, 2015 (chapter 11).

Volcanic sand in Iceland consist mostly of basaltic glass (Arnalds, 2015). Such basaltic volcanic materials can be found in other volcanically active areas such as Hawaii and other states of the USA and in New Zealand (Edgett & Lancaster, 1993), however the composition, particle size distribution, and microphysical properties vary with the place of origin and sand formation factors. It is also worth noting that Iceland has one of the largest volcaniclastic sand-fields (Arnalds *et al*., 2001).

Frequent dust-storms lift small volcanic-sand particles off the surface and transport them over great distances. For instance, they can be deposited in high latitude regions like Svalbard and Greenland (Groot Zwaaftink, 2017). Unlike desert sand, which is typically a quartz-mineral (e.g., Volten *et al.,* 2001; Nousiainen *et al.,* 2009), the composition of volcanic sand is less well known. Nevertheless, what is known about the composition of suspended volcanic sand suggests a significant impact on the atmosphere, specifically across the Northern latitudes (Dagsson-Waldhauserova *et al.*, 2016) in addition to its contribution to accelerating glacier melt (Wittmann, 2017).

The transport of volcanic sand by wind contaminates both the atmosphere and ground-surface along its transport path, where suspended particles efficiently absorb solar radiation leading to simultaneous heating of the atmosphere and surface. Because the suspended particles eventually settle-out, for

instance, on a pure ice and/or snow surface, the surface albedo is altered leading to enhanced heating and an increase in the surface's density (Meinander *et al.*, 2014; Peltoniemi *et al.*, 2015). This in turn triggers melting or evaporation of ice and snow surfaces (Qian, 2009).

Contamination of the atmosphere or icy/snow terrestrial surfaces by volcanic sand can be detected with remote-sensing techniques through ground-based and satellite observations (e.g. AERONET, Holben *et al.*, 1998, Sinyuk, et al, 2007, GOME-2: Munro *et al*. 2016, CALIOP/CALIPSO: Winker *et al.*, 2009). The reflectance and polarization of sunlight scattered by atmospheric aerosols contain important information about the microphysical properties of the particles. Indeed, polarimetry is a powerful and promising tool for the retrieval and characterization of these microphysical properties. Presently, several space instruments have polarization sensors and provide Earth observational data (e.g. Herman, 2005, Munro *et al*. 2016). Moreover, a number of new space missions are planned, which will perform airborne polarimetry (Dubovik *et al*., 2019). Interpretation of such measurements, however, remains difficult primarily because the measurements are simultaneously affected by the particles' shape, size distribution, and chemical composition. Comparison of satellite data with ground-based measurements may show significant differences. An example is Tao *et al*., 2017, who demonstrate an evaluation of the MODIS Deep Blue aerosol algorithm in the desert region of East Asia and compare to retrievals with ground-based observations obtained with China Aerosol Remote Sensing Network. They find that the MODIS-based retrievals of aerosol optical depth can be significantly underestimated. Besides the aerosol optical depth characteristics, aerosol-type classification is another challenge (Kahn & Gaitley, 2015). For instance, a set of dust mixtures is used to define the aerosol type in Multiangle Imaging Spectroradiometer (MISR) retrievals algorithms. Thus, knowledge of the optical properties of various dust types is critical for the success of such retrievals. Laboratory studies of the light scattering

properties of dust particles could meaningfully improve the algorithm inputs and interpretation of *in situ* measurements.

In this work, we investigate the reflectance and degree of linear polarization of sunlight scattered by volcanic-sand particles suspended in air and the same particles deposited on a surface. As such, our experiment reproduces both of measurement scenarios relevant for remote-sensing observations of volcanic sand, i.e., in the atmosphere or deposited on an ice/snow surface. Previous related measurements mainly focus on either single-particles or deposited-particles (Muñoz et al. 2004; 2015; Hadamcik, 2002; Sun 2014; Peltoniemi *et al.*, 2009, Wilkman *et al.*, 2016). There are only few examples where both light-scattering scenarios are simultaneously studied (e.g., Shkuratov *et al.*, 2004, 2006; Francis *et al.*, 2011) and Icelandic volcanic sand is not encompassed in that work. The measurements in our work are conducted at two experimental facilities: the goniospectropolariphotometer (FIGIFIGO) located at the Finnish Geospatial Research Institute (FGI) and Cosmic Dust Laboratory (CoDuLab) at the Instituto de Astrofísica de Andalucía (IAA). The FIGIFIGO facility (Fig. 2) is designed to measure the light-scattering response from a particle-coated surface (Peltoniemi et al., 2014), whereas the CoDuLab facility (Fig. 3) is used to measure the full scattering-matrix of particles suspended in air (Muñoz et al., 2012). We also complement our study with mass spectrometry to infer the elemental composition of the volcanic sand samples used.

Light scattering properties of surface-deposited volcanic sand have, in part, been studied before. In Peltoniemi *et al.* (2015), the sand is used as a highly absorbing contaminant for a snow surface, where that study focus on how the contamination affects the reflection and polarization properties of the snow. In work by Zubko *et al.* (2016), the optical properties of high-contrast two-component mixtures involving volcanic sand are studied, where the sand serves as a dark component among two types of bright components, salt (NaCl) and ferric sulfate ($Fe_2(SO_4)_3$). Note, however, that the light-scattering

122 behavior of volcanic sand is investigated in these studies only when the particles are deposited on a
123 substrate (particulate surface). The single-particle regime of light scattering is not yet investigated.
124

125 **2. Sample description**

126 Volcanic-sand particles mainly consist of poorly crystallized glasses of basaltic to andesitic origin. The
127 samples we consider are a mixture of glaciofluvial volcanic ash originating from beneath the
128 Mýrdalsjökull glacier mixed with ash from the Eyjafjallajökull and Grímsvötn eruptions of 2010 and
129 2011, respectively. We choose this sample because it is representative of materials that are typical of
130 aerosol-dust sources in Southern Iceland and the particles deposited on glaciers or snow in that area
131 (Arnalds *et al.,* 2013, Arnalds *et al.,* 2016). Specifically, our samples were collected from the
132 Mýrdalssandur area in Iceland. The large black-colored area in Fig. 1 corresponds to the field of
133 volcanic sand and the upper layer of this sand (about 10 cm thickness) was collected with a shovel.
134 Wind erosion in the area contributes to the redistribution of loose surface material, and according to
135 Arnalds et al. 2016, the rates of surface transport of aeolian materials is between 500 and 3,000 kg m−1
136 year−1. This means that about 0.5–3 tons are blown over a 1 m wide transect each year. The relevant
137 volcanic sand formation and erosion processes in Iceland are outlined in more detail, in e.g., (Baratoux
138 *et al.* 2011, Arnalds *et al.* 2013).

139 Our sand samples generally divide into the following categories:

140 1) Natural volcanic sand without processing (except for drying);
141 2) Sieved volcanic sand where the size of the particles is less than 250 μm, including:
142 a) Dry sand;
143 b) Wet sand, where moisture is provided by an atomizer;
144 3) Milled volcanic sand where the particles are ground to produce a fine-grained powder.

Samples 1–2 are studied with the FIGIFIGO experimental apparatus only. The natural sample (1) is abundant with coarse, millimeter-sized particles. Unfortunately, with the CoDuLab apparatus, such particles are too large for a feasible study of their light-scattering behavior at the single-particle level. The problem arises primarily from the aerosol generator, which becomes jammed by such large particles. The sieved volcanic sand (2) consists of particle sizes < 170 μm is used in the CoDuLab experiments, although the signal-to-noise ratios (SNRs) are low. The poor SNRs can be explained as follows: First, many of the sub-millimeter sieved samples remain too large for optimal operation of the aerosol generator. Although the generator operates considerably better with sample (2) than sample (1), the amount of suspended dust remains low leading to weak scattering-signals. Second, due to the large size of the constituent particles, they are much darker in appearance compared to the smaller, micron-sized particles and this can be seen in Fig. 4. Indeed, the milled sample (3) exhibits a brighter appearance in Fig. 4 compared to samples (1) and (2). Optical-microscope images of the samples are included in Fig. 4 to highlight the variability of particle sizes and shapes.

To obtain a more detailed view of the particle morphology in the natural sand sample (1), the particles are also examined by scanning electron microscopy (SEM) at various magnifications, see Fig. 5. As seen, the particles exhibit a highly irregular and somewhat vesicular morphology. The elemental composition of these particles is analyzed with X-ray spectrometry and the results are presented in Fig. 6. The analysis is repeated for particles of different sizes and it is notable that no significant variation in the chemical composition is found.

The size distribution of the milled volcanic sand sample (3) is shown in Fig. 7. The distribution is measured at the IAA CoDuLab with the *MasterSizer2000* instrument by *Malvern Scientific*. Note that the *MasterSizer2000* measures the flux of laser light scattered at a several scattering angles θ near forward-scattering direction. To retrieve information about the particles, it is customary to fit such

168 measurements with the two methods: Fraunhofer-diffraction theory or Mie theory (Bohren & Huffman,
169 1983). As neither of these theories applies well to highly irregular particle shapes in general, the fitting
170 procedure can yield results with error, e.g., with respect to particle size.

171 What emerges from Fig. 7 is that for particle radii > 0.25–0.3 μm, both approaches reveal a power-
172 law size distribution r^{-n} that can be fit by n = 3.5 in the Fraunhofer framework and at n = 3.2 in the Mie
173 approach. However, we stress that both approaches assume idealistic targets (spherical particles) and
174 therefore, the results should be taken with caution when applied to the highly irregular particles here.
175 This point will cause some uncertainty in the retrievals of the size distribution that should be taken into
176 account.

177

178 **3. Experimental facilities and measurement details.**

179 **3.1 FIGIFIGO**

180 The Finnish Geodetic Institute goniospectropolariphotometer FIGIFIGO (Fig. 2) is designed to measure the
181 reflectance and degree of linear polarization of various surfaces, both in the laboratory and in the field conditions
182 (e.g. Peltoniemi *et al.*, 2015b). A detailed description of the FIGIFIGO can be found in Peltoniemi *et al.,* 2014.
183 In our study, the sample is deposited on the surface by uniform sprinkling particles on a black substrate with a
184 layer of 0.8 - 1 cm thick. The measurements are taken in the principal plane, i.e., when the surface normal lies
185 within the scattering plane and to improve the SNR we repeat the measurements 25 times. To compare phase-
186 angle dependences of the reflectance and the degree of linear polarization of FIGIFIGO with CoDuLab
187 measurements, we present FIGIFIGO results obtained in the waveband λ = 642 – 652 nm (hereafter λ = 647
188 nm). The maximum uncertainty, ~3%, appears in the polarimetric measurements at some phase angles whereas
189 the average uncertainty in the polarimetric response is ~2%. The uncertainty in the measurements of reflectance
190 is noticeably lower compared to that of the polarimetric measurements.

3.2 CODULAB

The IAA Cosmic Dust Laboratory (CoDuLab) is designed to measure the light scattering response from aerosol particles, see Fig. 3. We notice that in the case of irregularly shaped particles, the light-scattering response can be described by the so-called (4×4) *scattering matrix* or the *Mueller matrix* consisting of six non-zero elements (see, e.g., Bohren & Huffman 1983):

$$F = \begin{pmatrix} F_{11} & F_{12} & 0 & 0 \\ F_{12} & F_{22} & 0 & 0 \\ 0 & 0 & F_{33} & F_{34} \\ 0 & 0 & -F_{34} & F_{44} \end{pmatrix}. \quad (1)$$

Using the CoDuLab facility, one can measure all six non-zero elements over the *scattering angle* θ range from 3° to 177° at several wavelengths. In our study, we investigate the upper block of non-zero elements. For a detailed description of the CoDuLab facility refer to (Muñoz *et al.*, 2012).

The results presented here are obtained at λ = 647 nm. In particular, we measure the F_{11} and F_{12} elements of the scattering matrix for the milled volcanic sand sample (3). The intensity of the scattered sunlight I and its degree of linear polarization P are defined via the elements of the scattering matrix as follows: $I \propto F_{11}$, $P = -F_{12} / F_{11}$.

4. Results and discussion

4.1 Reflectance and degree of linear polarization

We first investigate the difference in light scattering behavior for the sieved sample deposited on a surface in dry and wet conditions by measuring the reflectance and degree of linear polarization as a function of λ. Spectra for the wet and dry samples at θ=170° and θ=140° are shown in Fig. 8. The reflectance of the wet sample is found, on average, to be nearly half of that for the dry sample. Moreover, the shape of the reflectance curve clearly depends on presence of water, which is especially noticeable for λ in the range 500-1000 nm. The reflectance plot demonstrates the highly absorbing property of volcanic sand evidenced by the maximum value being only ~ 0.044 at λ=650 nm. The strongest reflectance response is observed for the dry sample at θ = 170°. Note, that while the reflectance spectra of this sample differs unambiguously at the two scattering angles (θ=170° and θ=140°), a similar difference for the wet sample is less in magnitude.

In contrast to the reflectance spectra, the curve shapes and scattering response of the polarization spectra are not as sensitive to water content. As shown later, however, a discrimination between wet and dry particles is easier at smaller θ. On the other hand, the degree of linear polarization decreases as the scattering angle grows. This is better seen in the angular dependence of the polarization in the wet and dry samples shown in Fig. 9, where the measurements are taken at λ = 647 nm. We see that the reflectance of the wet sample (blue curve) decreases considerably whereas the polarimetric response becomes noisier. In the latter case, however, the polarimetric response for the wet sample (blue curve) is higher compared to the dry sample (black curve). The difference between the dry and wet samples is best seen at side scattering, θ ~ 90°. Also, the increase of uncertainty in our measurements can be explained by the continuous evaporation of water during the measurements. To minimize the effect due to evaporation, we replenish the water content in the sample several times during the course of the experiment with water atomizer.

Particle size is a dominate factor affecting the light-scattering behavior of a sample, which is illustrated in Fig. 10 for reflectance and polarization at λ = 647 nm. Here, the natural and milled volcanic sand samples (1) and (3), respectively, are illuminated at θ=148°. Near the backscattering direction, i.e., $\theta \rightarrow 180°$, the reflectance of the milled sample appears to be considerably higher than that of the natural sample. This change is consistent with the so-called *Umov effect* or *Umov law*, which is the inverse correlation between reflectance near backscattering, i.e., geometric albedo, and the maximum value that the degree of linear polarization may acquire, e.g., see Shkuratov & Opanasenko 1992; Zubko *et al.* 2016. According to the Umov effect, greater polarization is expected from the natural sand compared to the milled sand. Indeed, the maximum degree of linear polarization of the milled sample is ~2.5 times less than that for the natural sample. Also note, that the value of negative polarization near the backscattering direction is greater for the milled sample. The Umov effect can be seen for the near-backscattering reflectance and polarization maximum for the dry and wet sieved samples presented in Fig. 9.

Comparing the reflectance of the dry sieved samples shown in Fig. 9 with the reflectance for the natural and milled samples in Fig. 10, one finds that the reflectance of the sieved sample is greater than the natural sand and less than the milled sand. Correspondingly, the polarization of the sieved sample is smaller than for the natural sand and greater than the milled sand.

Comparative analysis of the light-scattering behavior for deposited particles and the same type of particles suspended in air as an “optically thin cloud,” i.e., as an aerosol, can provide important information needed for the interpretation of observations of atmosphere and underlying terrain. The milled volcanic sand is measured in both scenarios, i.e., deposited on a surface or as an aerosol, and the results are shown in Fig. 11. Note that for the aerosol, it is not feasible to measure an absolute flux of the scattered light. Therefore, we normalize the reflectance data at θ = 90°. In the FIGIFIGO

255 measurements, the deposited particles are illuminated at $\theta = 52°$ with respect to the surface normal.
256 Due to measurement constraints on the θ range and the increased intensity of scattered light from single
257 particles near the forward-scattering direction, our analysis is limited to the range $60° < \theta < 177°$.

258 In Fig. 11, one can see that the normalized reflectance for the aerosol is noticeably stronger for
259 small θ compared to the data for deposited particles. However, the aerosol data also show a decrease
260 and become weaker for $\theta > 90°$. The maximum value of the degree of linear polarization appears
261 greater for deposited particles compared to the aerosol, i.e., $P_{max} \approx (22.07 \pm 0.93)\%$ at $\theta_{max} \approx 60°$ and
262 $P_{max} \approx (18.7 \pm 1.1)\%$ at $\theta_{max} \approx 100°$, respectively. Although this difference is not large, it is detected
263 with confidence in Fig. 11. For $\theta < 105°$, this finding qualitatively differs from that reported in
264 Shkuratov *et al.* (2007), for example, where ten different samples are investigated including clay,
265 olivine, feldspar, and volcanic ash. All those samples reveal a systematically lower polarization for
266 deposited samples. Our results, however, are in a good agreement with soot measurements by Francis
267 et al., 2011, where polarization measured from surface packed particles considerably exceeds that from
268 the aerosol particles. Notice in Fig. 11 that the negative polarization branch near backscattering for
269 deposited particles appears deeper than the aerosol, which is opposite to the conclusions drawn in
270 Shkuratov et al. (2004). This difference could result from specific features of our sample of volcanic
271 sand. It is significant, for example, that our sand is much darker in appearance than any sample used in
272 Shkuratov et al. (2004; 2007).

273 The light scattering response for volcanic sand differs considerably from other types of sand
274 such as desert sand, which is obvious from their visual appearance. While volcanic sand is dark in
275 appearance (its geometric albedo is ~0.05), desert sands are considerably brighter.

In Fig. 12 we plot the normalized reflectance (left) and degree of linear polarization (right) as a function of scattering angle for three different samples: white-clay, milled volcanic-sand, and the ash particles from the Eyjafjallajokull volcano. The light scattering properties of white clay and volcanic ash were studied with CODULAB by Munoz *et al*. in 2011 and Merikallio *et al*. in 2015. Here we present a comparison of the light scattering properties of volcanic sand in this study to the white clay and Eyjafjallajokull volcanic ash samples. White clay mainly consists of illite, kaolinite, montmorillonite, quartz and is an important component of aerosols in the atmosphere. Eyjafjallajokull volcanic ash was collected at 5 km from the source after the April 2010 eruption, where the main constituent is silica, SiO_2. These particle also contain Al_2O_3, CaO, TiO_2, FeO, MgO and Na_2O.

As one can see from Fig. 12 there is similarity between the three polarization curves. The refractive index m of the Eyjafjallajokull ash particles could be similar to that in the milled volcanic sand. However, they should differ significantly from that of the white clay, at least with regard to the imaginary part, Im(m). Nevertheless, the light scattering response of white-clay particles resembles that of the volcanic sand and ash particles. The resemblance could be explained by a difference in size distribution of the white-clay and volcanic sand and ash particles that, by coincidence, compensates the difference in refractive index. The angular profiles of the reflectance at large scattering-angles (>90°) clearly differs for the dark samples (volcanic sand and ash) and light samples, i.e., the white clay. Such a feature could be useful in passive remote-sensing of aerosol particles.

Fig. 13 compares our results to the polarization measurements obtained for levitated soot-particles by Francis *et al.*(2011) The data adopted from Francis *et al.* is measured at λ = 632.8 nm from a dense cloud of levitated agglomerated particles of Polymethyl Methacrylate (PMMA), hereafter called the soot sample. The soot particles are micrometer-sized aggregates, where the constituent grains have a diameter of several tens of nm. The maximum polarization of the volcanic sand is about 19%, while the

299 soot is about 25%. Based on this, and considering the Umov law, we can conclude that the near-
300 backscattering reflectance of the soot sample is lower than that of milled volcanic sand.
301
302 **4.2 Constraint of the complex refractive index of volcanic sand**
303 An advantage of our study with single-scattering particles is that interpretation data does not involve
304 common complications due multiple scattering. We can develop a quantitative model describing the
305 single-scattering particles based on a numerically exact solution of the Maxwell equations, although
306 this cannot be done here for the deposited particles. The ultimate goal of such modeling is retrieval of
307 the microphysical properties of the particles.
308 In general, light scattering by submicron and micron-sized particles is dependent on their size
309 distribution, shape, and complex refractive index (*m*). However, Zubko *et al.* (2015) show that in the
310 case of highly irregular particles, the effect of size distribution and *m* on the light scattering behavior
311 dominates the effects of particle shape. Furthermore, the size distribution of the milled volcanic sand is
312 constrained by the *MasterSizer2000* measurements (Fig. 7) and this allows us to estimate *m* for the
313 volcanic sand with the light-scattering model.
314 Specifically, we model the angular profiles of the scattered intensity *I* and degree of linear
315 polarization *P* for the milled volcanic-sand particles using the so-called *agglomerated debris particles*
316 method (Zubko, 2015, Zubko *et al*. 2015b). Such particles have a disordered morphology with a
317 packing density of the constituent material being ~ 0.236. Six examples of the agglomerated debris
318 particles are shown in Fig. 14, which are generated by systematically damaging a perfect sphere as
319 described in Zubko *et al.* (2013). An notable feature of the model particles is that they reproduce
320 analogous laboratory measurements of a variety of samples similar to those considered here, such as
321 feldspar (Zubko *et al.*, 2013), forsterite (Zubko, 2015), and olivine (Videen *et al.*, 2018). In particular,

the model parameters, i.e., the size distribution and *m*, applied to these analogous measurements closely match the actual microphysical characteristics of the samples.

Analysis of satellite data is typically done with spheroidal model-particles, see for example Dubovik et al. However, Dubovik et al. 2006 shows that the use of spheroidal particles could not satisfactorily reproduce laboratory measurements when multiple wavelengths are involved. Furthermore, the microphysical properties retrieved with spheroidal particles do not match the true properties of the feldspar particles (refractive index, size, and aspect-ratio distributions), while agglomerated debris particles is in good quantitative agreement with the true microphysical characteristics of feldspar (Zubko *et al.*, 2013).

A particle's light-scattering behavior depends, in part, on the ratio of its radius r to wavelength λ, is commonly called the size parameter $x = 2\pi r/\lambda$ (Bohren & Huffman 1983). In application to an irregularly shaped particle, we assign r to a sphere that circumscribes the model particle used. Then, we compute the relevant light-scattering quantities using agglomerated debris particles for $1 < x < 32$ with the *discrete dipole approximation* (DDA). The DDA is a flexible technique designed for numerical simulation of light scattering by particles with an arbitrary shape (Yurkin *et al.* 2007). In the DDA framework, the particle is replaced by a set of cubic cells that reproduces the shape and internal structure, where the size of the cells d is sufficiently small compared to λ. As demonstrated in Zubko *et al.* (2010), the DDA yields robust numerical results for $2\pi d|m|/\lambda \leq 1$. Each cell is then approximated by an electric dipole and, thus, the integral equation describing interaction of an electromagnetic wave with the particle is transformed into a system of linear algebraic equations. This system of equations is then solved via an iterative method. In order to comply with the discretization criterion for d above, we consider two cases for each agglomerated debris particle: $64 \times 64 \times 64$ cells and $128 \times 128 \times 128$ cells. The former is used when $x \leq 15$ and the latter when $x > 15$. The size parameter is varied in steps of Δx

345 = 1 for $1 < x < 15$ and $\Delta x = 2$ for $16 < x < 32$. Given that $\lambda = 647$ nm in our study, this size-parameter
346 range corresponds to a particle-size range $0.1\ \mu m < r < 3.3\ \mu m$. As demonstrated in Zubko *et al.* (2013),
347 Zubko (2015), and Videen *et al.* (2018), such a range is sufficient to reproduce the light-scattering
348 response in a polydisperse system exhibiting a power-law size distribution with an exponent of $n \geq 2.9$.

349 To investigate the dependence on m, we apply the DDA to agglomerated debris particles at 46
350 values of m, with real and imaginary parts spanning the ranges $1.1 < \mathrm{Re}(m) < 2.43$ and $0 < \mathrm{Im}(m)$ 1,
351 respectively. For every pair (x, m) we average the light-scattering response over a minimum of 500
352 random shapes to ensure a statistically reliable result. We also perform size averaging using a power-
353 law distribution r^{-n} over the full range of r considered. Initial values for n are inferred from Fig. 7, i.e.,
354 n = 3.2 and 3.5. However, we do not consider these as precisely known values. Instead, we assume a
355 degree of variation of ±0.5, which recognizes the uncertainty inherent to the size distribution
356 measurements with the *MasterSizer 2000* instrument.

357 We investigate all the available refractive indices, searching for the best fit to the maximum value of
358 the degree of linear polarization $P_{max} \approx (18.7 \pm 1.1)\%$ found for the milled-sand sample at $\theta = 100°$.
359 When such fit is possible, we then compare the entire angular profiles of I and P measured for the
360 given sample with that from the DDA model applied to agglomerated debris particles (Zubko *et al.*,
361 2013; Zubko, 2015; Videen *et al.*, 2018). As Fig. 14 shows, the best fit for $\lambda = 647$ nm is obtained for
362 $m = 1.6 + 0.01i$ and $n = 2.85$. As one can see, the intensity is reproduced well for all θ, whereas the
363 degree of linear polarization tends to agree less near backscattering ($\theta > 140°$) where the phenomenon
364 of negative polarization is observed (i.e., $I_\perp < I_\parallel$). The same qualitative behavior is seen for a feldspar
365 particle (Zubko *et al.*, 2013), although the difference is smaller. Overall, the agglomerated debris
366 particle model in other work closely matches the true microphysical properties of the measured particle

(Zubko *et al.,* 2013; Zubko, 2015; Videen *et al.,* 2018). Such performance of the model lends confidence to our conclusion in Fig. 14 that $m = 1.6 + 0.01i$ for the volcanic sand at $\lambda = 647$ nm.

Finally, consider the question of how the inferred material absorption, i.e., Im(m) = 0.01, corresponds to the dark appearance of the milled volcanic sand, having a reflectance ~ 0.1 near backscattering $\theta = 175°$ and presumably ~ 0.12 at $\theta = 180°$ (see Fig. 11). We draw attention to previous laboratory measurements of m in powdered kerogen type-II reported by Khare *et al.* (1990). In particular, for the red part of the spectrum, a similar value for Im(m) is found, Im(m) ≈ 0.012, whereas the powder is described as having the dark appearance. Thus, our finding that Im(m) = 0.01 for the volcanic sand agrees well with the visual appearance of the sand deposited on a surface.

5. Conclusions

Our study of the light-scattering behavior of Icelandic volcanic sand achieves a quantitative characterization of this important material. The reflectance and polarization deposited sand strongly depends on the particle size-distribution. According to the Umov law, the maximum of the polarization degree encodes information on the material's optical absorption and reflectance. This effect is demonstrated for the three samples of volcanic sand containing particles with different size distributions. The degree of linear polarization differs by nearly a factor of three between the milled and natural-sand samples whereas the polarization seen for the sieved sample is smaller than the natural sand and greater than the milled sand. However, the water added to the sieved sand increases its polarimetric response considerably, so that at some scattering angles the response is nearly the same as the natural sand, which consists of larger particles. This observation may have important implications for remote-sensing observations of regions with such sand present.

Our comparative analysis of the reflectance and polarization response of particles suspended as an aerosol with those deposited on a surface reveal:

(1) The normalized reflectance of light scattered by the aerosols is noticeably stronger at small scattering-angles compared to that of the deposited particles.

(2) The positive degree of linear polarization for the aerosols is greater than for the deposited particles for $90° < \theta < 160°$.

(3) The maximum polarization is as large as $P_{max} \approx 22\%$ occurring at $\theta_{max} \approx 60°$ for deposited particles, and $P_{max} \approx 19\%$ at $\theta_{max} \approx 100°$ for the aerosol particles.

(4) The polarization response from the aerosol and deposited particles becomes similar at $\theta = 95°$.

(5) The negative polarization branch of deposited particles is deeper than that of the aerosol.

Finally, based on discrete-dipole modeling of the reflectance and degree of linear polarization we estimate the refractive index of the Icelandic volcanic sand to be $m = 1.6 + 0.01i$ at $\lambda = 647$ nm.

Acknowledgments

This research was partially supported by the Academy of Finland Project no. 260027 and the COST Action MP1104 “Polarization as a tool to study the Solar System and beyond”. NZ acknowledges Magnus Ehrnrooth Foundation for the research travel support. We thank P. Dagsson Waldhauserová, O. Arnalds, A. Virkkula, O. Meinander, and J. Svensson for their help obtaining the samples and for relevant discussions. We acknowledge the use of imagery provided by services from NASA's Global Imagery Browse Services (GIBS), part of NASA's Earth Observing System Data and Information System (EOSDIS). We also would like to thank reviewers for their constructive reviews.

411 **References**

412

413 **Arnalds**, O., Gisladottir, F.O., Sigurjonsson, H., 2001. Sandy deserts of Iceland: an overview. Journal
414 of Arid Environments 47:359–371. https://doi.org/10.1006/jare.2000.0680

415 **Arnalds**, O**.**, Thorarinsdottir, E.F., Thorsson, J., Waldhauserova, P.D., Agustsdottir, A.M., 2013. An
416 extreme wind erosion event of the fresh Eyjafjallajökull 2010 volcanic ash. Scientific Reports 3.
417 https://doi.org/10.1038/srep01257

418 **Arnalds**, O., 2015. The Soils of Iceland, World Soils Book Series. Springer Netherlands.
419 https://doi.org/10.1007/978-94-017-9621-7Kok, J.F., Parteli, E.J.R., Michaels, T.I., Karam, D.B., 2012.
420 The physics of wind-blown sand and dust. Reports on Progress in Physics 75, 106901.
421 https://doi.org/10.1088/0034-4885/75/10/106901

422 **Baratoux**, D**.**, Mangold, N., Arnalds, O., Bardintzeff, J.-M., Platevoët, B., Grégoire, M., Pinet, P.,
423 2011. Volcanic sands of Iceland – Diverse origins of aeolian sand deposits revealed at Dyngjusandur
424 and Lambahraun. Earth Surface Processes and Landforms 36, 1789–1808.
425 https://doi.org/10.1002/esp.2201

426 **Boucher**, O., D. Randall, P. Artaxo, C. Bretherton, G. Feingold, P. Forster, V.-M. Kerminen, Y. Kondo,
427 H. Liao, U. Lohmann, P. Rasch, S.K. Satheesh, S. Sherwood, B. Stevens and X.Y. Zhang, 2013:
428 Clouds and Aerosols. In: Climate Change 2013: The Physical Science Basis. Contribution of Working
429 Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change [Stocker,
430 T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M.
431 Midgley (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.

Bohren, C. F., Huffman, D. R., 1983. Absorption and Scattering of Light by Small Particles. Wiley, New York.

Dagsson-Waldhauserova, P., Magnusdottir, A., Olafsson, H., Arnalds, O., 2016. The Spatial Variation of Dust Particulate Matter Concentrations during Two Icelandic Dust Storms in 2015. Atmosphere 7, 77. https://doi.org/10.3390/atmos7060077

Dubovik, O., A. Sinyuk, T. Lapyonok, B.N. Holben, M. Mishchenko, P. Yang, T.F. Eck, H. Volten, O. Muñoz, B. Veihelmann, W.J. van der Zande, J.-F. Leon, M. Sorokin, and I. Slutsker, 2006: Application of spheroid models to account for aerosol particle nonsphericity in remote sensing of desert dust. J. Geophys. Res., 111, D11208, doi:10.1029/2005JD006619.

Dubovik, O., Li, Z. Mishchenko, M.I. et al., 2019. Journal of Quantitative Spectroscopy & Radiative Transfer, 224, 474–511

Edgett, K.S. & Lancaster, N. , 1993. Volcaniclastic aeolian dunes: terrestrial examples and application to martian sands. Journal of Arid Environments, 25: 271–297.

Groot Zwaaftink, C.D., Arnalds, Ó., Dagsson-Waldhauserova, P., Eckhardt, S., Prospero, J.M., Stohl, A., 2017. Temporal and spatial variability of Icelandic dust emissions and atmospheric transport. Atmospheric Chemistry and Physics 17, 10865–10878. https://doi.org/10.5194/acp-17-10865-2017

Hadamcik, E., 2002. Polarization of Light Scattered by Fluffy Particles (PROGRA2 Experiment). Icarus 155, 497–508. https://doi.org/10.1006/icar.2001.6732

Herman M., 2005. Aerosol remote sensing from POLDER/ADEOS over the ocean: improved retrieval using a nonspherical particle model. J. Geophys. Res., 110:D10S02. doi: 10.1029/2004JD004798 .

Kahn, R. A., and Gaitley, B. J. (2015), An analysis of global aerosol type as retrieved by MISR. J. Geophys. Res. Atmos., 120, 4248– 4281. doi: 10.1002/2015JD023322.

Khare, B. N., Thompson, W. R., Sagan, C., Arakawa, E. T., Meisse C., & Gilmour, I., 1990. Optical constants of kerogen from 0.15 to 40 mm: comparison with meteoritic organics. In: Proceedings of Lunar and Planetary Science Conference XXI, 627–628.

Meinander, O., Kontu, A., Virkkula, A., Arola, A., Backman, L., Dagsson-Waldhauserová, P., Järvinen, O., Manninen, T., Svensson, J., de Leeuw, G., Leppäranta, M., 2014. Brief communication: Light-absorbing impurities can reduce the density of melting snow. The Cryosphere 8, 991–995. https://doi.org/10.5194/tc-8-991-2014

Munro, R., Lang, R., Klaes, D., Poli, G., Retscher, C., Lindstrot, R., et al., 2016. The GOME-2 instrument on the Metop series of satellites: instrument design, calibration, and level 1 data processing –an overview. Atmos. Meas. Tech., 9:1279–301. doi: 10.5194/amt- 9- 1279- 2016 .

Merekallio, S., Muñoz, O., Sundström, A.-M., Virtanenm T.H. Horttanainen, M., de Leeuw, G., Nousiainen, T., 2015. Optical modeling of volcanic ash particles using ellipsoids. Journal of Geophysical Research, 120(9), p. 4102-4116.

Muñoz, O., Volten H., Hoveneir, J.W., veihelmann, B., van der Zande, W.J. Waters, L.B.F.M., Rose, W.I., 2004. Scattering matrices of volcanic ash particles of Mount St. Helens, Redoubt, and Mount Spurr Volcanoes. Journal of Geophysical Research, 109, Issue D16, Cite ID D16201.

Muñoz, O., Moreno, F., Guirado, D., Ramos, J.L., Volten. H., Hovenier, J.W. The IAA cosmic dust laboratory: Experimental scattering matrices of clay particles. Icarus, 211 (2011) 894-900.

Muñoz, O., Moreno, F., Guirado, D., Dabrowska, D.D., Volten, H., Hovenier, J.W., 2012. The Amsterdam–Granada Light Scattering Database. Journal of Quantitative Spectroscopy and Radiative Transfer 113, 565–574. https://doi.org/10.1016/j.jqsrt.2012.01.014

Peltoniemi, J., Hakala, T., Suomalainen, J., Puttonen, E., 2009. Polarised bidirectional reflectance factor measurements from soil, stones, and snow. Journal of Quantitative Spectroscopy and Radiative Transfer 110, 1940–1953. https://doi.org/10.1016/j.jqsrt.2009.04.008

Peltoniemi, J.I., Hakala, T., Suomalainen, J., Honkavaara, E., Markelin, L., Gritsevich, M., Eskelinen, J., Jaanson, P., Ikonen, E., 2014. Technical notes: A detailed study for the provision of measurement uncertainty and traceability for goniospectrometers. Journal of Quantitative Spectroscopy and Radiative Transfer 146, 376–390. https://doi.org/10.1016/j.jqsrt.2014.04.011

Peltoniemi, J.I., Gritsevich, M., Hakala, T., Dagsson-Waldhauserová, P., Arnalds, Ó., Anttila, K., Hannula, H.-R., Kivekäs, N., Lihavainen, H., Meinander, O., Svensson, J., Virkkula, A., de Leeuw, G., 2015a. Soot on Snow experiment: bidirectional reflectance factor measurements of contaminated snow. The Cryosphere 9, 2323–2337. https://doi.org/10.5194/tc-9-2323-2015

Peltoniemi, J.I., Gritsevich, M., Puttonen, E., 2015b. Reflectance and polarization characteristics of various vegetation types // Springer Praxis Books, Light Scattering and Radiative Transfer, Light Scattering Reviews 9, pp. 257-294, http://dx.doi.org/10.1007/978-3-642-37985-7_7

Prospero, J. M., Bullard, J. N., Hodgkins, R. 2012. High-latitude dust over the north atlantic: inputs from icelandic proglacial dust storms. Science, 335, issue 6072, pp. 1078-1082 https://doi.org/10.1126/science.1217447

Qian, Y., Gustafson, W.I., Jr., Leung, L.R., Ghan, S.J., 2009. Effects of soot-induced snow albedo change on snowpack and hydrological cycle in western United States based on Weather Research and

494 Forecasting chemistry and regional climate simulations. Journal of Geophysical Research 114.
495 https://doi.org/10.1029/2008jd011039

496 **Shkuratov**, Yu.G., Opanasenko, N.V., 1992. Polarimetric and photometric properties of the moon:
497 Telescope observation and laboratory simulation. II – The positive polarization. Icarus 99, 468–484;
498 https://doi.org/10.1016/0019-1035(92)90161-Y

499 **Shkuratov**, Y., Ovcharenko, A., Zubko, E., Volten, H., Muñoz, O., Videen, G., 2004. The negative
500 polarization of light scattered from particulate surfaces and of independently scattering particles.
501 Journal of Quantitative Spectroscopy and Radiative Transfer 88, 267–284.
502 https://doi.org/10.1016/j.jqsrt.2004.03.029

503 **Shkuratov**, Yu., Bondarenko, S., Kaydash, V., Videen, G., Muñoz, O., Volten, H., 2007. Photometry
504 and polarimetry of particulate surfaces and aerosol particles over a wide range of phase angles. J. Quant.
505 Spectrosc. Radiat. Transfer 106, 487–508; https://doi.org/10.1016/j.jqsrt.2007.01.031

506 **Sinyuk**, A., Dubovik, O., Holben, B., Eck, T.F., Bréon, F-M, Martonchik, J., et al., 2007. Simultaneous
507 retrieval of aerosol and surface properties from a combination of AERONET and satellite data. Remote
508 Sens. Environ.;107:90–108.

509 https://doi.org/10.1016/j.rse.2006.07.022

510 **Sun**, Z., Zhang, J., Tong, Z., Zhao, Y., 2014. Particle size effects on the reflectance and negative
511 polarization of light backscattered from natural surface particulate medium: Soil and sand. Journal of
512 Quantitative Spectroscopy and Radiative Transfer 133, 1–12.
513 https://doi.org/10.1016/j.jqsrt.2013.03.013

Tao, M., Chen, L., Wang, Z., Wang, J., Che, H., Xu, X., … Hou, C. (2017). Evaluation of MODIS Deep Blue aerosol algorithm in desert region of East Asia: Ground validation and intercomparison. Journal Geophysical Research: Atmospheres, 122, 10, 357–10, 368.

https://doi.org/10.1002/2017JD026976

Videen, G., Zubko, E., Arnold, J.A., MacCall, B., Weinberger, A.J., Shkuratov, Yu., & Muñoz, O., 2018. On the interpolation of light-scattering responses from irregularly shaped particles. J. Quant. Spectrosc. Radiat. Transfer, 211, 123–128 ; https ://doi.org/10.1016/j.jqsrt.2018.03.009

Wilkman O., Gritsevich M., Zubko N., Peltoniemi J., Muinonen K., 2016. Photometric modelling for laboratory measurements of dark volcanic sand // Journal of Quantitative Spectroscopy and Radiative Transfer 185, 37–47 http ://dx.doi.org/10.1016/j.jqsrt.2016.08.013

Winker, D.M. Vaughan, M.A., Omar, A., Hu, Y., Powell, K.A., Liu, Z., et al., 2009. Overview of the CALIPSO mission and CALIOP data processing algorithms. J Atmos Ocean Technol, 26 :2310–23. Doi : 10.1175/2009JTECHA1281.1 .

Wittmann, M., Groot Zwaaftink, C.D., Steffensen Schmidt, L., Guðmundsson, S., Pálsson, F., Arnalds, O., Björnsson, H., Thorsteinsson, T., Stohl, A., 2017. Impact of dust deposition on the albedo of Vatnajökull ice cap, Iceland. The Cryosphere 11, 741–754. https://doi.org/10.5194/tc-11-741-2017

Zubko, E., Petrov, D., Grynko, Ye., Shkuratov, Yu., Okamoto, H., Muinonen, K., Nousiainen, T., Kimura, H., Yamamoto, T., & Videen, G., 2010. Validity criteria of the discrete dipole approximation. Appl. Opt., 49, 1267–1279; https://doi.org/10.1364/AO.49.001267

Zubko, E., Muinonen, K., Muñoz, O., Nousiainen, T., Shkuratov, Yu., Sun, W., & Videen, G., 2013. Light scattering by feldspar particles: comparison of model agglomerate debris particles with

laboratory samples. J. Quant. Spectrosc. Radiat. Transfer, 131, 175–187; https://doi.org/10.1016/j.jqsrt.2013.01.017

Zubko, E., 2015. Modeling light scattering by forsterite particles. Opt. Lett., 40, 1204–1207; https://doi.org/10.1364/OL.40.001204

Zubko, E., Shkuratov, Yu., & Videen, G., 2015b. Effect of morphology on light scattering by agglomerates. J. Quant. Spectrosc. Radiat. Transfer, 150, 42–54 ; https ://doi.org/10.1016/j.jqsrt.2014.06.023

Zubko, N., Gritsevich, M., Zubko, E., Hakala, T., Peltoniemi, J.I., 2016. Optical measurements of chemically heterogeneous particulate surfaces. Journal of Quantitative Spectroscopy and Radiative Transfer 178, 422–431. https://doi.org/10.1016/j.jqsrt.2015.12.010

546 **Figures**

547

548 **Figure 1**

549 Satellite image of Iceland obtained with MODIS (Blue Marble from August 2004). Enlarged image of
550 the selected area is taken from MODIS 08.07.2009).

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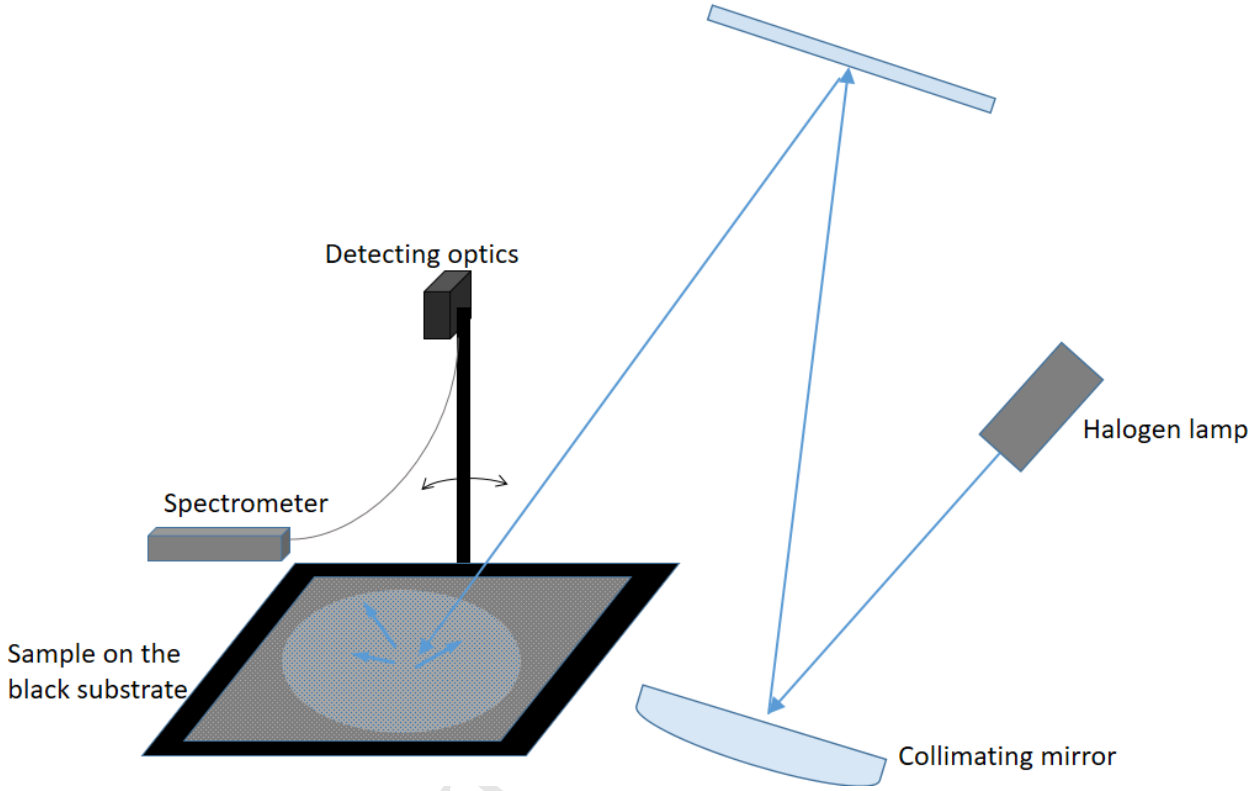
Figure 2

Scheme of Finnish Geospatial Research Institute goniospectropolariphotometer FIGIFIGO setup



565 **Figure 3**

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567 Scheme of IAA Cosmic Dust Laboratory setup

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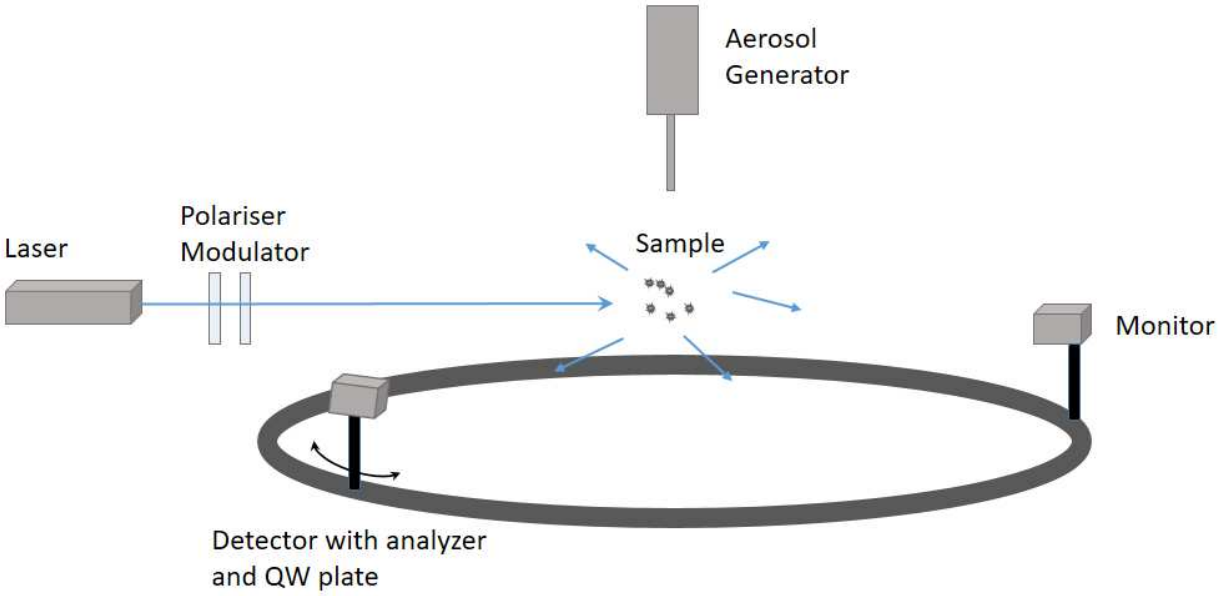

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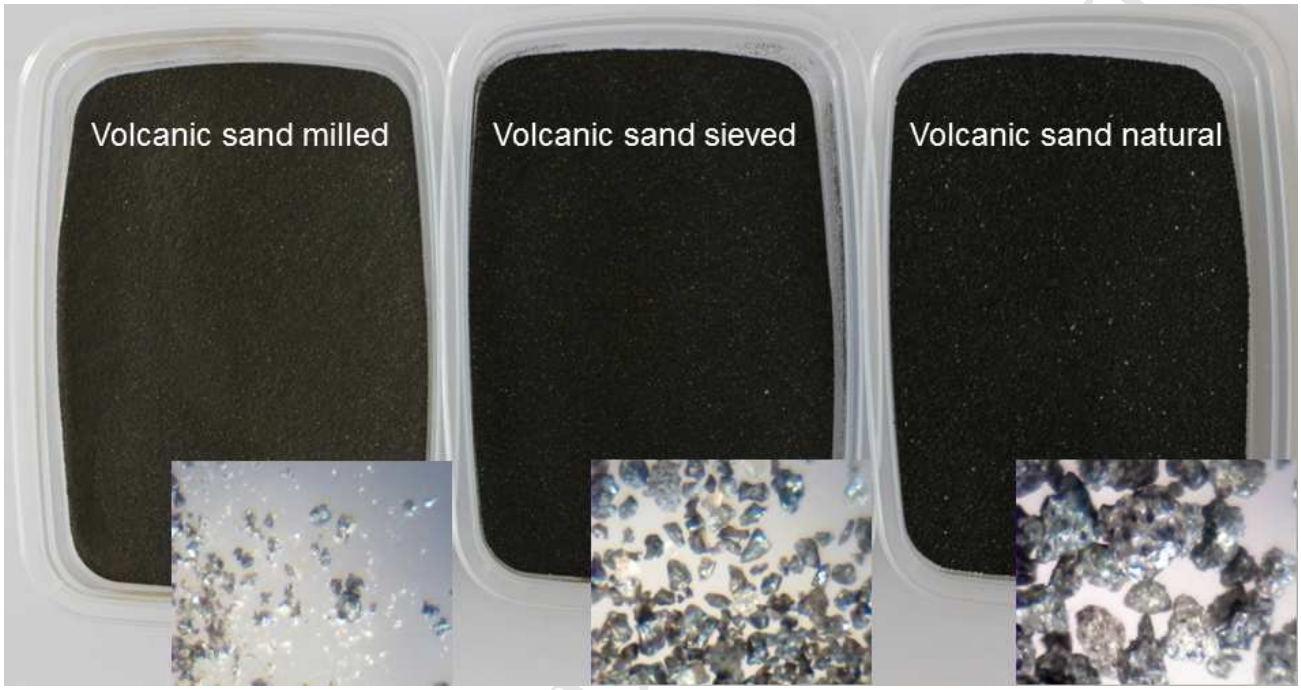
Figure 4

Appearance of milled, sieved, and natural volcanic sand samples along with optical-microscope images of the same samples.



584 **Figure 5**

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586 Scanning electron microscopic images of the natural sample

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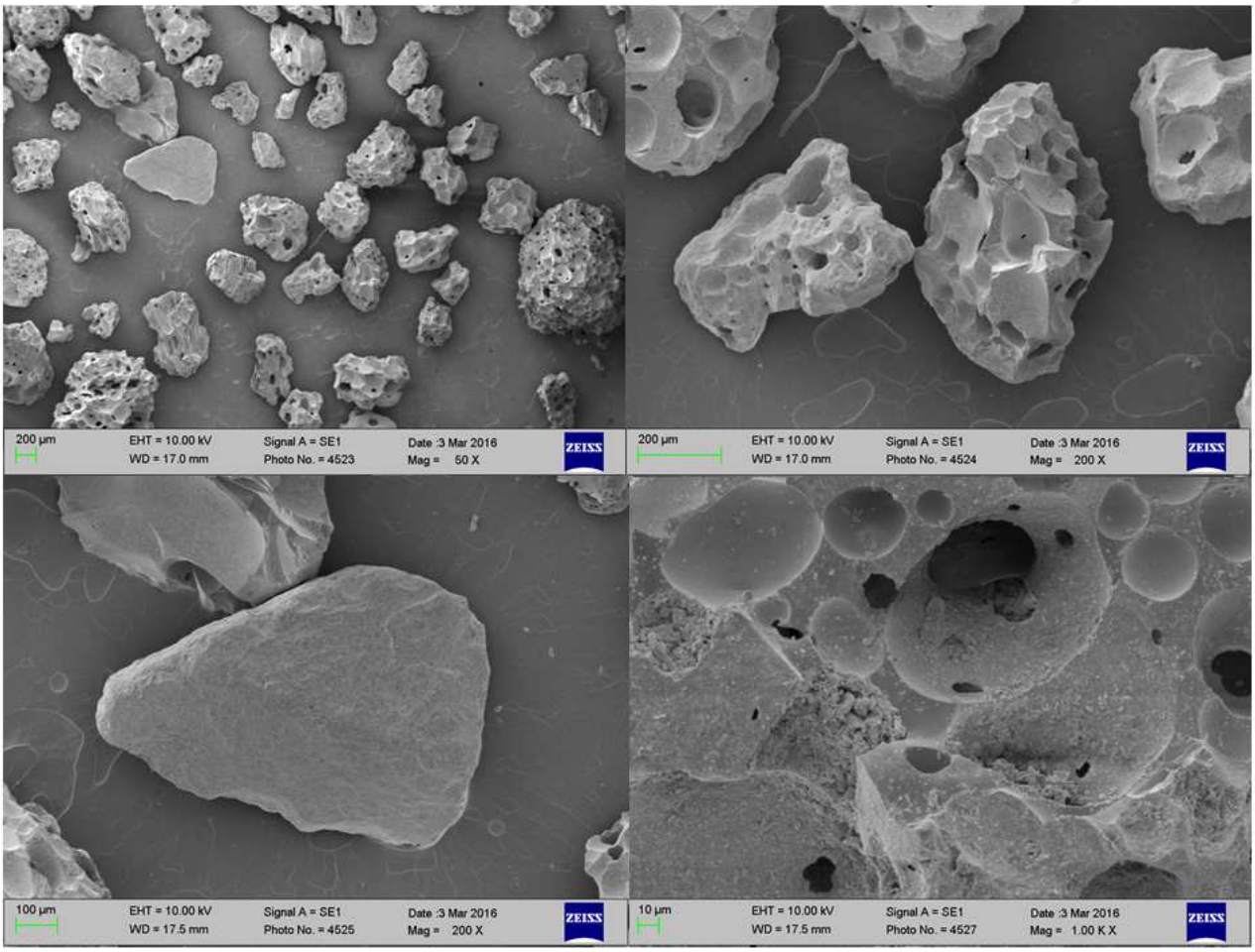

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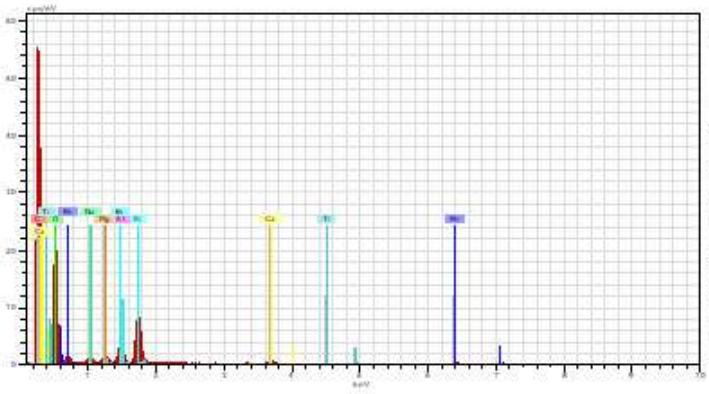
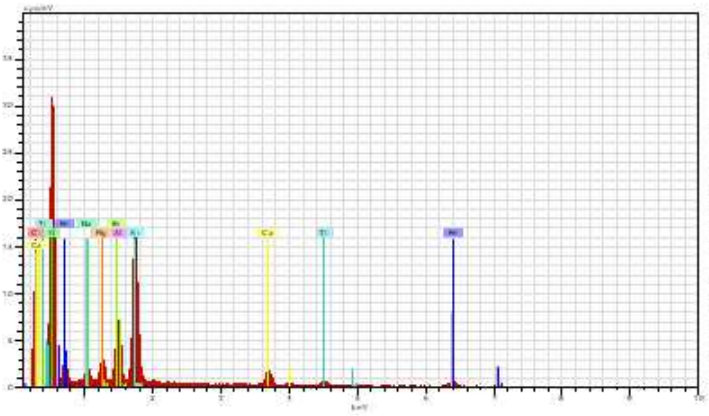
Figure 6

Element composition of natural volcanic sand particles analyzed with the X-ray spectrometry



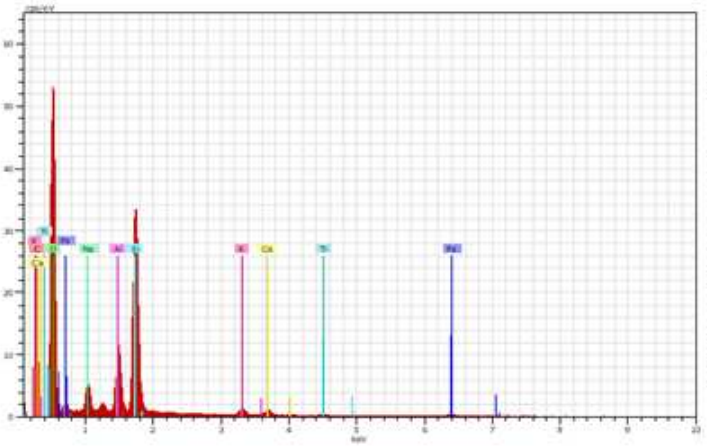
Spectrum: volcanic sand 80

Element	Series	unn. C [wt.%]	norm. C [wt.%]	Atom. C [at.%]	Error (1 Sigma) [wt.%]
Oxygen	K-series	47.80	27.32	25.33	5.23
Iron	K-series	10.74	6.14	1.63	0.51
Silicon	K-series	9.61	5.49	2.90	0.42
Aluminium	K-series	1.23	0.70	0.39	0.08
Calcium	K-series	2.93	1.68	0.62	0.13
Titanium	K-series	1.52	0.87	0.27	0.08
Magnesium	K-series	0.88	0.50	0.31	0.07
Sodium	K-series	0.81	0.46	0.30	0.07
Bromine	L-series	3.21	1.84	0.34	0.17
Carbon	K-series	96.23	55.00	67.92	10.19
Total:		174.97	100.00	100.00	



Spectrum: volcanic sand 81

Element	Series	unn. C [wt.%]	norm. C [wt.%]	Atom. C [at.%]	Error (1 Sigma) [wt.%]
Carbon	K-series	25.94	19.11	30.99	3.11
Oxygen	K-series	53.60	39.49	48.07	5.94
Iron	K-series	19.06	14.04	4.90	0.88
Silicon	K-series	17.68	13.02	9.03	0.75
Aluminium	K-series	2.40	1.77	1.28	0.14
Calcium	K-series	5.01	3.69	1.79	0.20
Titanium	K-series	3.13	2.30	0.94	0.14
Magnesium	K-series	1.79	1.32	1.06	0.12
Sodium	K-series	1.50	1.11	0.94	0.12
Bromine	L-series	5.62	4.14	1.01	0.28
Total:		135.72	100.00	100.00	



Spectrum: volcanic sand 82

Element	Series	unn. C [wt.%]	norm. C [wt.%]	Atom. C [at.%]	Error (1 Sigma) [wt.%]
Carbon	K-series	32.63	24.14	34.68	3.70
Oxygen	K-series	59.74	44.20	47.67	6.44
Iron	K-series	6.94	5.14	1.59	0.35
Silicon	K-series	23.91	17.69	10.87	1.00
Aluminium	K-series	5.39	3.99	2.55	0.27
Calcium	K-series	1.82	1.35	0.58	0.09
Sodium	K-series	2.52	1.87	1.40	0.18
Potassium	K-series	1.29	0.96	0.42	0.07
Titanium	K-series	0.90	0.67	0.24	0.06
Total:		135.15	100.00	100.00	

598 **Figure 7**

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600 Size distribution of the milled volcanic sand particles on linear scale on the right and log scale on the
601 left.

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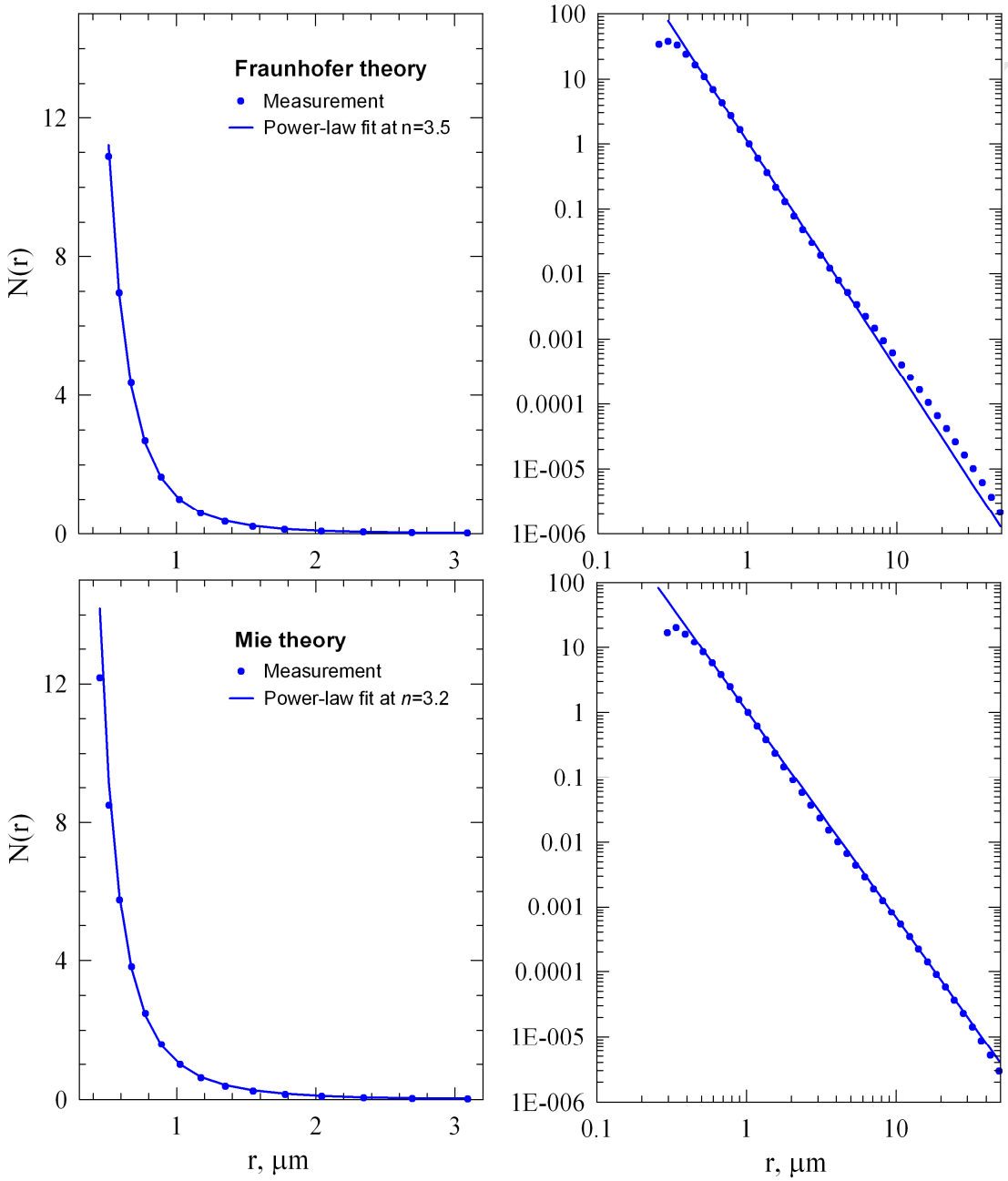

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606 **Figure 8**

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608 Reflectance and degree of linear polarization as function of wavelength of sieved wet and dry volcanic
609 sand particles deposited on the surface.

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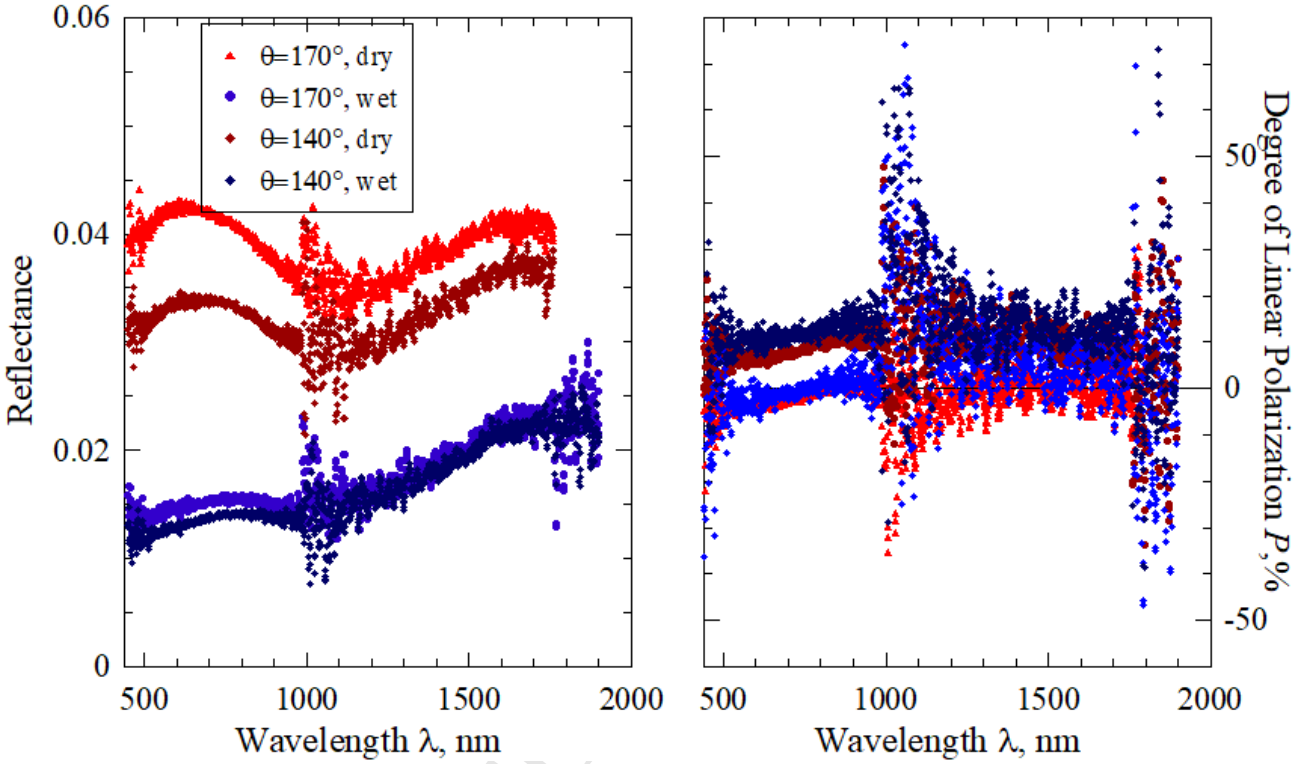

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Figure 9

Reflectance and degree of linear polarization of the dry and wet sieved volcanic sand as a function of scattering angle at wavelength of 647 nm

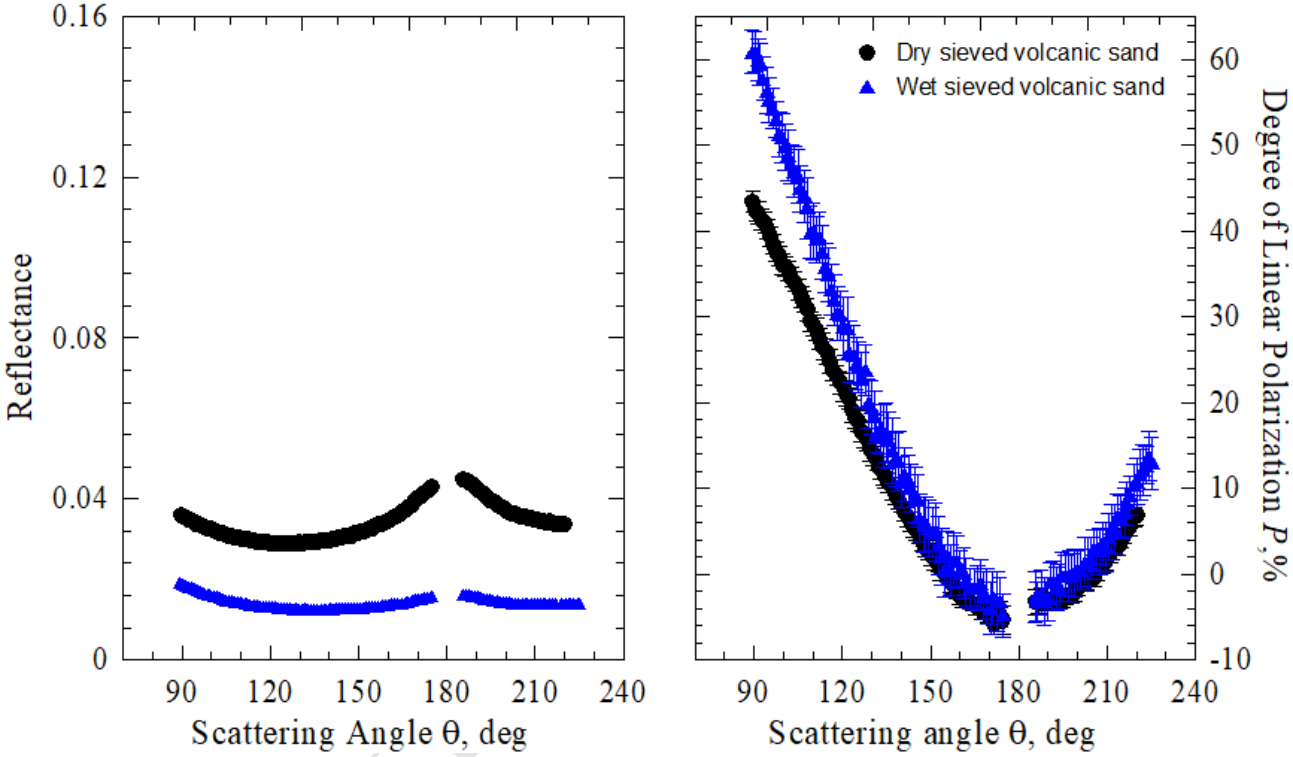
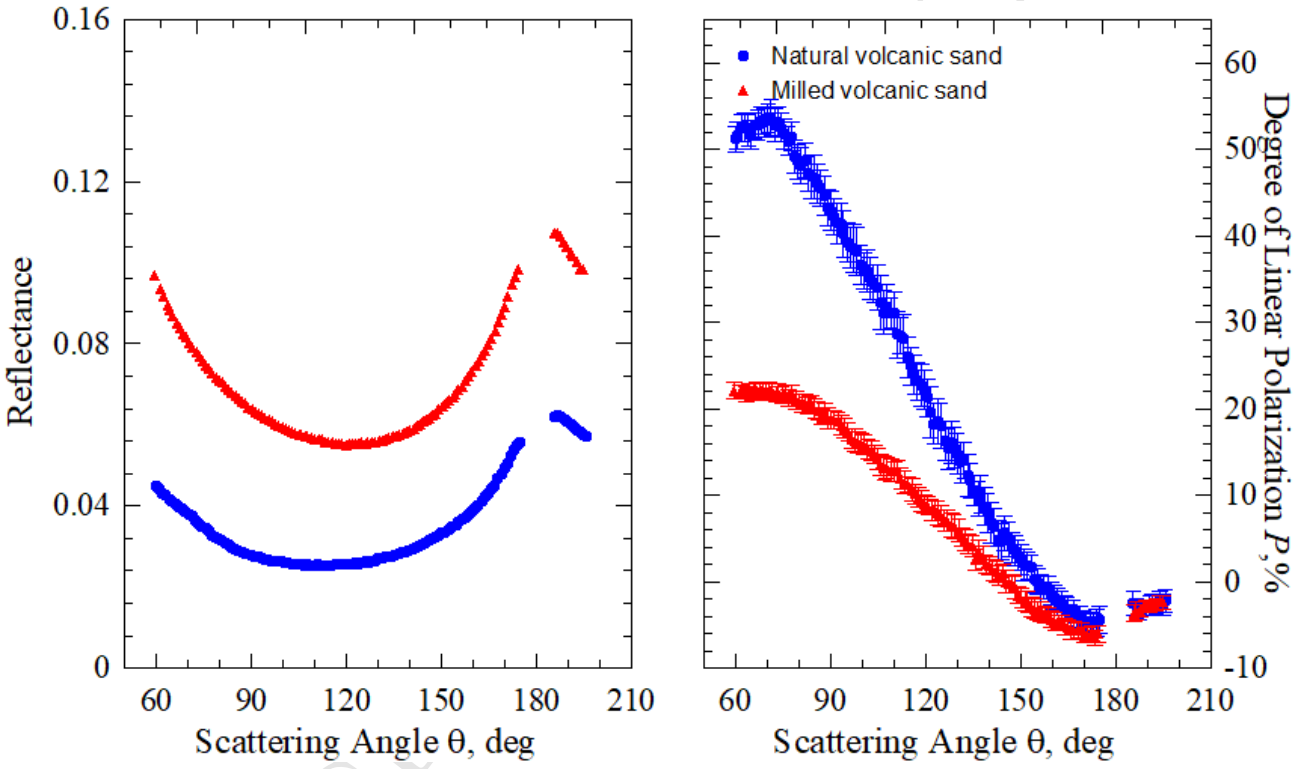


Figure 10

Reflectance and degree of linear polarization of the dry natural and milled volcanic sand as function on phase angle at wavelength of 647 nm



635 **Figure 11**

636 The scattering angle dependence of the normalized reflectance and degree of linear polarization
637 obtained for the surface and aerosol of the milled volcanic sand.

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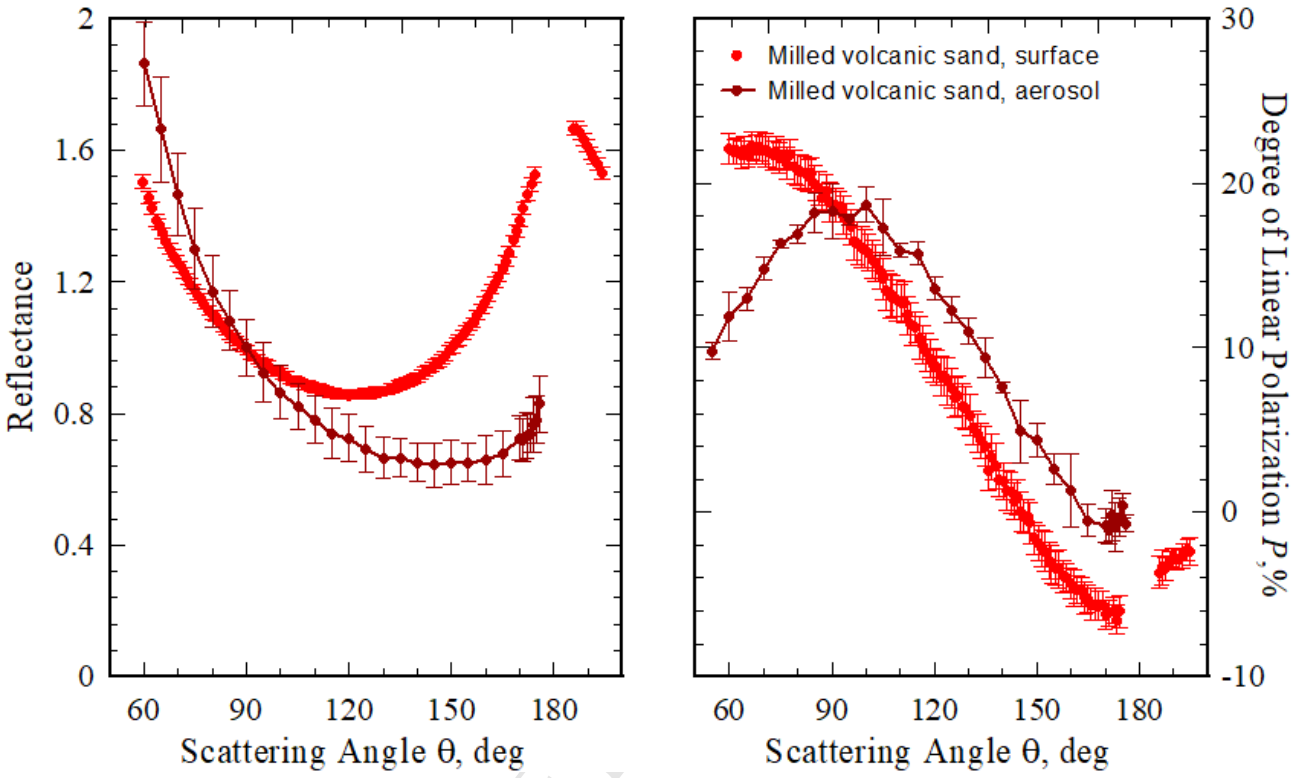

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642 **Figure 12**

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644 The normalized reflectance and degree of linear polarization as a function of scattering angle for three
645 different samples: white-clay, milled volcanic-sand, and the ash particles from the Eyjafjallajokull
646 volcano.

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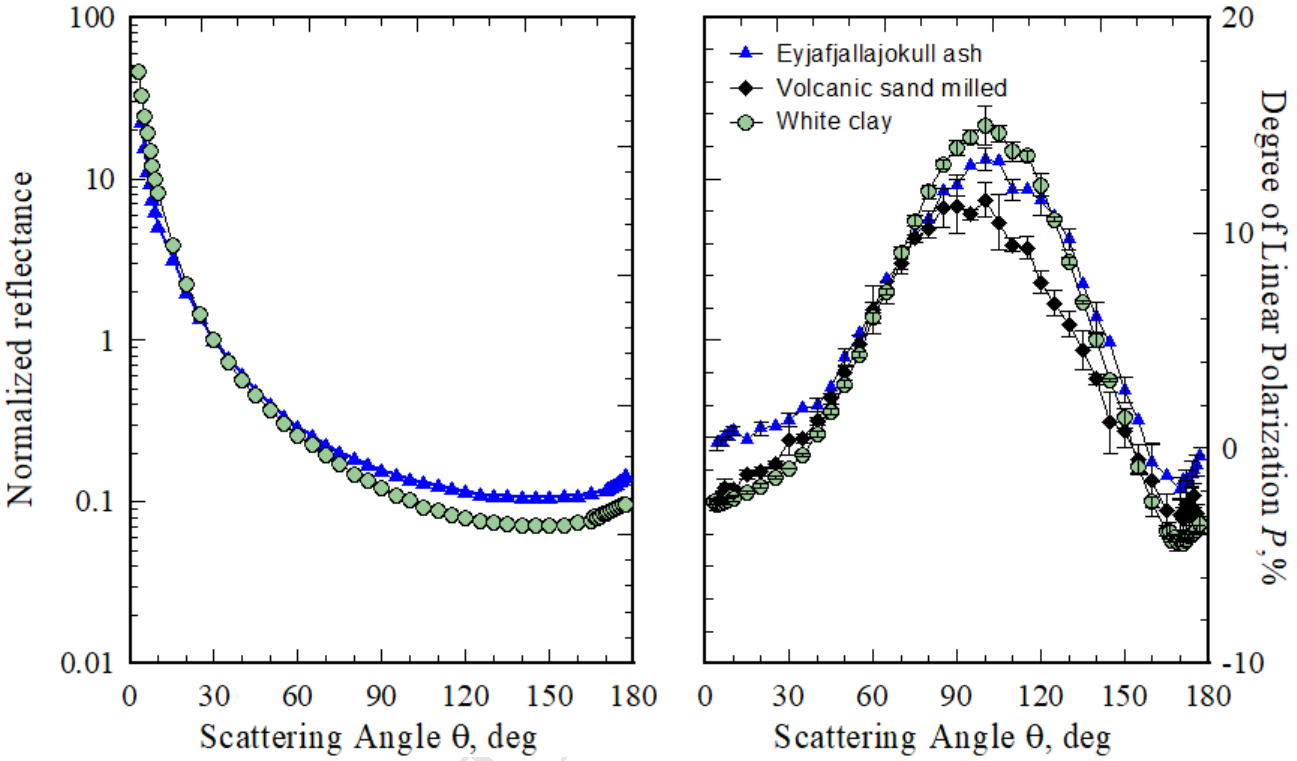

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651 **Figure 13**

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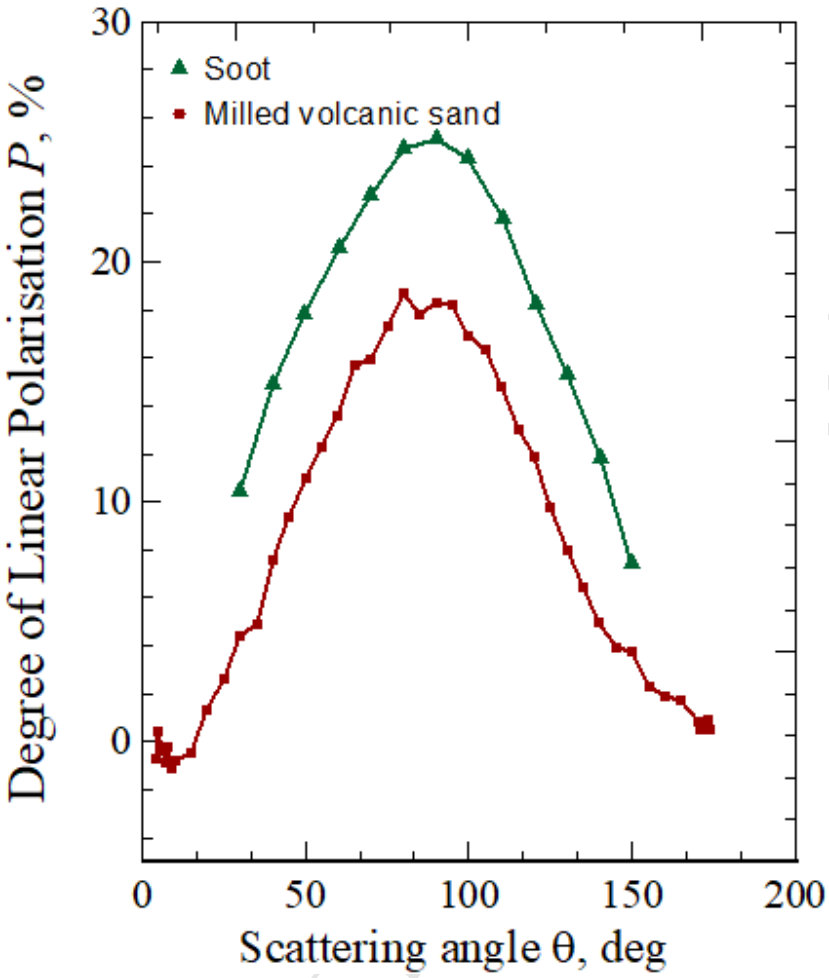
653 Comparison of the polarization response of volcanic sand λ = 647 nm and soot particles at λ = 632.8
654 nm (data adopted from Francis *et al. 2011*) suspended into the air.

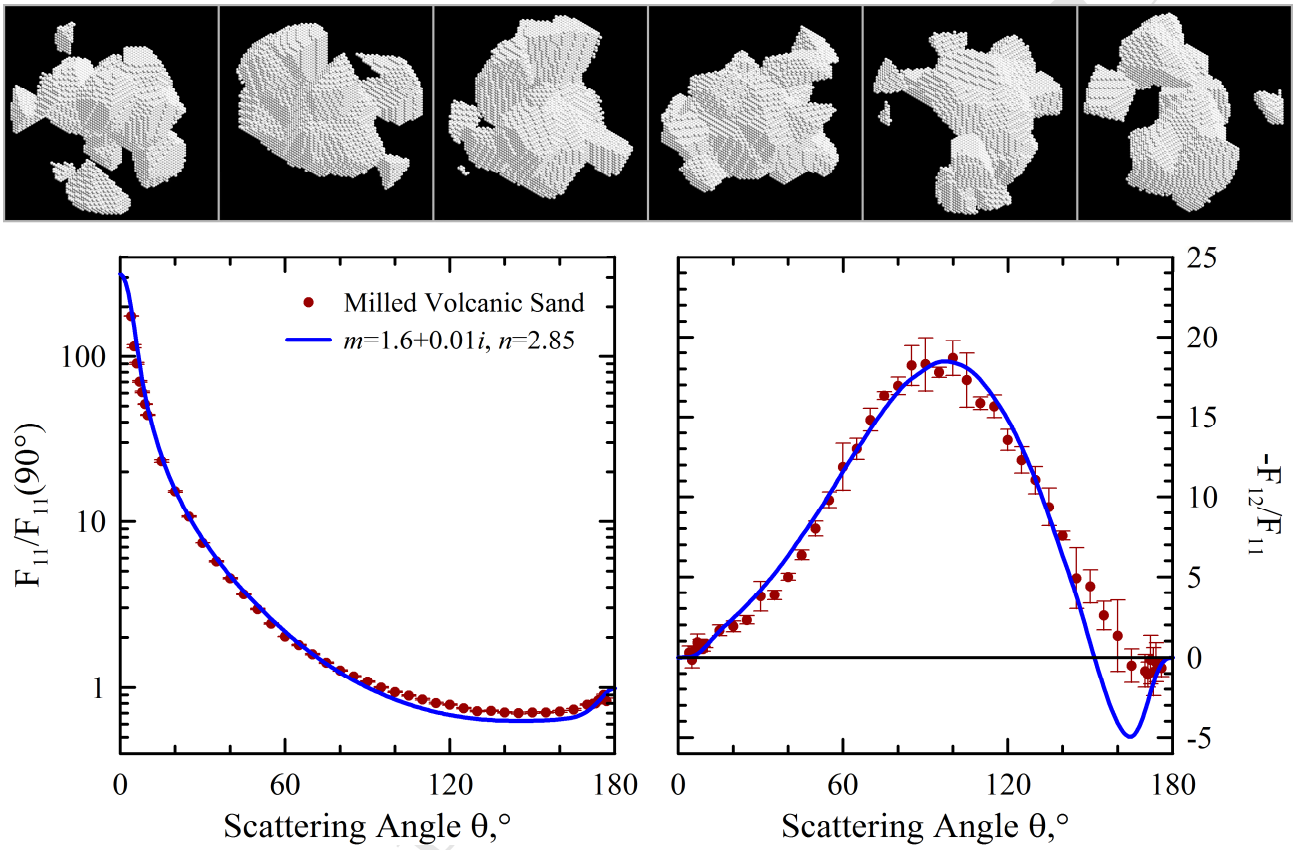
655



656

Figure 14

On the top: Six examples of the modeled agglomerated debris particles. On the bottom: Intensity *I* (left) and degree of linear polarization *P* (right) as a function of the scattering angle θ in the milled volcanic sand and their model (blue curve) at $n = 2.85$ and $m = 1.6 + 0.01i$.



Research Highlights

- We report angular scattered-light intensity and polarization from the volcanic sand
- Light scattering by volcanic sand is studied in single-particle and deposited modes
- Refractive index of volcanic sand is estimated
- Polarimetric response of volcanic sand is compared to that in carbon-soot