



## Palynological Studies of Maastrichtian to Paleocene Sediments Exposed at Okpekepe, Western Flank of Anambra Basin, Edo State, Nigeria

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**ABSTRACT:** Outcrop samples were collected at a road cut along Okpekepe road in the western part of the Anambra Basin with a view to establishing the age of the sediments and their depositional environment. These consist mainly of shales, sandstone, siltstone and heterolith. The samples were processed and analyzed using standard palynological procedures. Index fossils of stratigraphic significance were used for dating of the sediments. An extremely poor recovery of palynomorphs characterized by *Acrostichum aureum*, *Laevigatosporites* sp, *Verrucatosporites* sp, *Cingulatisporites sornatus*, *Deltosporasp*, *Lycopodium* spp, *Cyathidites minor*, *Echitricolporites triangulatus*, *Retitricolporites irregularis*, *Retitricolporites* sp, *Liliacidites* sp, *Longapertites marginatus*, *Longapertites* sp, *Erecipites* sp., *Monosulcites perspinosus*, *Retistephanocolpites williamsi*, *Echiperiporites estalae*, *Dodoneaviscosa*, *Echitriporites strianguliformis*, *Periretipollis spinosus*, *Marginipollis concinnus*, *Cupaniedites* sp., *Gleicheniidites senonicus* *Andalusiella polymorpha* and *Andalusiella* sp was noticed. A Maastrichtian-Paleocene age and a continental depositional environment were assigned to the sediments.

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The study area is located at latitude N07° 11' 02.2" and longitude E006° 28' 06.2". It is a road cut exposure of the Mamu Formation within the Benin Flank (western part) of the Anambra Basin. It occurs along Okpekepe Road in Etsako East Local Government Area of Edo State, Nigeria. The first area where intensive hydrocarbon exploration took place in Nigeria was the Anambra Basin. The basin has a sediment thickness of about 12,000 metres (Agagu and Adighije, 1983). The roughly triangular Anambra Basin Southern Nigeria covers about 40,000 sq km. The Southern boundary of the basin coincides with the northern boundary of the Niger Delta basin (Nwajide and Reijgers, 1996). It is a symmetrical ridge stretching in sigmoid curve for over 500km from Idah around the River Niger to Arochukwu in the Cross River (Arua, 1988). The origin and evolution of the Anambra Basin can be link to the origin of the Benue Trough. The Santonian deformation shifted depositional axis to Anambra Basin and Afipko syncline. (Ajakaiye and Burke, 1972) associated the evolution of the Benue trough with the rifting and breakup of Gondwana Super Continent leading to the separation of the Africa from South America in the Late Jurassic. (Burke *et al.*, 1972) proposed a triple joint theory for the origin of the trough. They linked it to the ridge-ridge-ridge (r-r-r) triple junction. The three arms of the junction are the

South America, the Gulf of Guinea and the Benue trough. According to Hoque and Nwajide, (1984), the evolution of the Benue trough was in four stages: rifting, trough, deformation and platform stages. The Mamu Formation consists of sandstones, shales, siltstones, coal seams and heteroliths. The sandstones have fine-medium grains brown-yellow in colour. The shales and mudstones are grey to black in colour. The coal seams vary in thickness from a few inches to twelve feet. The coals are of medium quality and are rich in resins and waxes. They have high ash content and display a rhythmic pattern of deposition. The sediments are shallow water deposits (Cratchley and Jones, 1965). Open strand plain and coastal plain environments have been inferred for this formation by the (Agagu *et al.*, 1985) and (Nwajide and Reijgers, 1996) respectively.

This study aims at applying palynology to decipher the age and depositional environment of the studied section. Palynology as a an aspect of biostratigraphy has proved to play important roles in deducing age of sediments, building sequence stratigraphic framework and reconstructing paleoenvironment (Morley, 1995; Ola-Buraimo and Akaegbobi, 2013a; Ayinla *et al.*, 2013; Adeigbe *et al.*, 2013) towards improving

exploration and exploitation of oil and gas. The works of (Mebradu, 1988) is also relevant to this study.

AGE	ASUBULIC-ANAMBRA BASIN	AFIPO BASIN
30	Oligocene	Ogwashi-Azuba Formation
30		Ogwashi-Azuba Formation
54.5	Eocene	Ameki/Ameki Formation/ Ameki Formation
54.5		Ameki Formation
65	Paleocene	Imo Formation
65		Imo Formation
72	Maastrichtian	Nsukka Formation
72		Nsukka Formation
72		Ajali Formation
72		Ajali Formation
72		Mamu Formation
72		Mamu Formation
83	Campanian	Nkporo Owerri Formation/Enugu Shale
83		Nkporo shale/Enugu sandstone
83.5	Santonian	
83.5		Non-deposition/Amelion
86.5	Coniacian	Agwu sandstone/ Agwu shale
86.5		
86.5	Turonian	Eze-Aku Group
86.5		Eze-Aku Group (Include Ammel sandstone)
95	Concomitant-Albian	Asu River Group
95		Asu River Group
120	Apollin	
120	Santonian	
120	Maastrichtian	
		Unnamed Group
	PRECAMBRIAN	SACMERE COMPLEX

Fig 1. Correlation Chart for Early Cretaceous strata in southeastern Nigeria (After Nwajide, 1990)

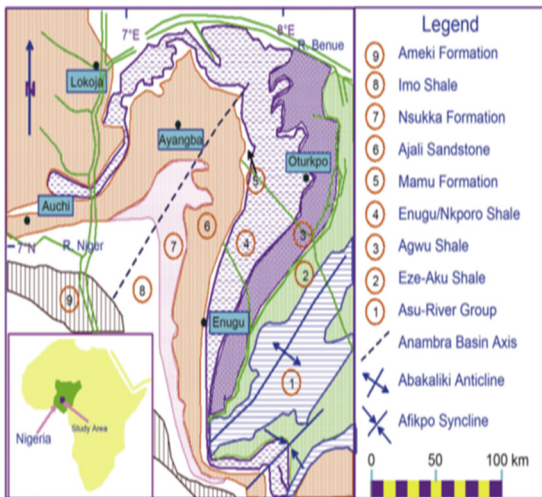


Fig 2. Geological map of the Anambra Basin showing location of study area (From Nton and Bankole, 2013).

## MATERIALS AND METHOD

Standard palynological processing procedures (Faegri and Iversen, 1989; Wood *et al.*, 1996) were employed for this purpose. Ten grams of each dry sample was crushed into small fractions. The samples were treated with dil HCl and HF to remove carbonates and silicates respectively. Concentrated HNO<sub>3</sub> was used for oxidation and heated over Bunsen burner in a fume cupboard. KOH was introduced, allowed to stay overnight and stirred until a neutral solution was achieved. Ammonium hydroxide concentrate was

added and diluted with water. At this stage separation of organic matter from inorganic was done by floating zinc bromide, ZnBr. Slides were mounted using Norland adhesive mounting medium and dried. One slide per sample was analyzed under an optical microscope and the photomicrographs of well-preserved palynomorph specimens were taken. Palynomorph identification was done using the works of (Germeraad *et al.*, 1968) and (Evamy *et al.*, 1978). The data was eventually plotted using *StrataBugs* software.

## RESULTS AND DISCUSSION

The results of the five outcrop samples processed and analyzed for palynological studies are outlined and discussed in the following sections.

*Sample 1:* The palynomorphs recorded here are characterized by predominance of Cretaceous-Tertiary spores and pollens assemblage such as *Deltospora sp.*, *Cingulatisporites ornatus*, *Cyathidites minor*, *Lycopodium sp.*, *Erecipites sp.*, *Longapertites sp.*, *Liliacidites sp.*, *Echitriporites trianguliformis*, *Marginipollis concinnus*, *Cupanioidites sp.*, *Dododnea viscosa*, *Retistephanocolpites williamsi*, *Periretipollis spinosus*, *Monosulcites perspinosus*. The quantitative top occurrence of *Cyathidites minor* marks the top of Paleocene boundary and the presence of *Echitriporites trianguliformis*, *Retistephanocolpites williamsi* and *Longapertites sp* indicate Maastrichtian age.

*Sample 2:* The presence of *Echitricolporites triangulatus*, *Deltospora sp.*, *Cingulatisporites ornatus*, *Erecipites sp.*, *Cupanioidites*, *Gleicheniidites senonicus*, and *Leoitriletesp* with spot occurrence of pteridophyte spores such *Laevigatosporites sp* indicate Maastrichtian to Paleocene. The age confirmation was based on the presence of *Deltospora sp* and *Gleicheniidites senonicus* which are palynomorphs of Late Cretaceous.

*Sample 3:* Paucity of marker palynomorphs was observed here. Forms like *Longapertites marginatus*, *Cyathidites sp.*, *Verrucatosporites sp* and *Zonocostites ramonae* were however recorded.

*Sample 4:* The presence of *Andalusiella polymorpha*, *Andalusiella sp.*, *Leoitrietes sp* with minor occurrences of pteridophyte spores such as *Verrucatosporites sp* and *Laevigatosporites sp* indicate Maastrichtian to Paleocene age.

*Sample 5:* The presence of *Cyathidites sp.*, *Verrucatosporites sp* and *Retitricolporites sp* without any distinct marker species.



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