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## **Diachronous exhumation of HP-LT metamorphic rocks from southwestern Alps: evidence from fission-track analysis**

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## **Abstract**

New fission-track ages on zircon and apatite (ZFT and AFT) from the southwestern alpine paleo-accretionary wedge document a contrasting cooling history from east to west. In the eclogitic Monviso ophiolites, the ZFT ages are  $19.6 \pm 0.8$  Ma and the AFT ages are  $8.6 \pm 1.7$  Ma. In the HT-blueschist eastern Queyras, ZFT ages range from  $27.0 \pm 1.5$  Ma to  $21.7 \pm 1.6$  Ma and AFT ages from  $14.2 \pm 2.0$  to  $9.4 \pm 1.1$  Ma. In the LT-blueschist western Queyras, ZFT ages are between  $94.7 \pm 3.1$  Ma and  $63.1 \pm 2.9$  Ma and AFT ages are between  $22.2 \pm 1.6$  and  $22.6 \pm 1.5$  Ma. The Chenaillet ophiolite yields ages of  $118.1 \pm 3.7$  Ma on ZFT and of  $67.9 \pm 8.5$  Ma on AFT. These new FT data combined with petrological and geochronological constraints record a diachronous exhumation in the paleo-accretionary wedge during subduction and collision.

**Keywords:** Southwestern Alps, fission-tracks, cooling rates, diachronous exhumation

## **Introduction**

Geochronological data combined with P-T estimates provide the information related to exhumation of high pressure low temperature (HP-LT) metamorphic rocks whether their metamorphism and exhumation were synchronous. In the Alpine belt, fission-track (FT) data are mostly obtained in central, eastern and northwestern domains (Hunziker et al., 1992; Seward and Mancktelow, 1994; Fügenschuh and Schmid, 2003; Malusà et al., 2005). There are very few age data from the Piedmont zone of the southwestern Alps. Furthermore, available FT data (Carpena and Caby, 1984) are not in accord with geochronological constraints based on other techniques (Duchêne et al., 1997a; Agard et al., 2002). The aim of this paper is to present new apatite and zircon FT data from meta-ophiolites and related oceanic meta-sediments in the blueschist and eclogite units from the Piedmont zone of the southwestern Alps (Fig. 1). These new data combined with petrological data constrain the exhumation processes in the western Alps.

## **Geological setting**

The Piedmont zone of the southwestern Alps represents the juxtaposed domains during subduction and collision in late Cretaceous to Tertiary times (Tricart, 1984; Spalla et al., 1996). Three main ophiolite-bearing units are exposed from west to east (Fig. 1). At the top of the nappe pile, the Chenaillet massif represents an obducted portion of the Tethyan oceanic crust, escaping alpine metamorphism. This unit rests upon the Queyras "Schistes lustrés" complex composed of kilometric scale slices of meta-ophiolites embedded within Mesozoic meta-sedimentary rocks (Lemoine and Tricart, 1986). This unit shows increasing P-T conditions towards east, with LT-blueschist facies conditions in western Queyras up to the transitional condition between blueschist and eclogite facies in eastern Queyras (Caby, 1996; Agard et al., 2001; Schwartz et al., 2000). The blueschist units are structurally above the

eclogitic Monviso ophiolites and the HP to UHP Dora Maira internal crystalline massif. In the Monviso massif, the meta-sedimentary component is very small (< 20 vol%) relative to voluminous oceanic lithosphere (Fig. 2). Eclogitic meta-gabbros occurred as boudins in a matrix of serpentinites or highly deformed meta-basalts. The calculated P-T conditions are heterogeneous among different boudins, suggesting that different fragments correspond to oceanic slab subducted to different depths (Blake et al., 1995; Schwartz et al., 2001).

## **FT analysis and results**

### *The Chenaillet unit*

Amphibolite to prehnite-pumpellyite facies conditions, representing the metamorphism on the ocean floor has been recognized by Mével et al. (1978). Fourteen grains of zircons in albitite dyke cross-cutting the amphibolitized meta-gabbros (Table 1) give a central age of 118 +/- 3.7 Ma. The individual age dispersion is low (relative error of central age: RE < 1 %), from 86.6 to 142.5 Ma, with a chi square test ( $P\chi^2$ ) higher than 90 % (Fig. 3a). For apatites, a central age of 67.9 +/- 8.5 Ma was obtained on 9 grains. Despite their low uranium contents these apatites show a very low dispersion of individual ages (RE < 1 %) and a  $P\chi^2$  test higher than 85 %.

### *The eclogitic Monviso ophiolites*

Fourty three grains of zircon and apatite were investigated from the Fe-Ti meta-gabbros of the Passo Gallarino unit (Table 1), which records the peak metamorphic condition of 450 +/- 50°C and 13 - 15 kbar (Schwartz et al., 2000). The data show low relative errors (RE < 1 %) and high score of the  $P\chi^2$  test (value > 99 %) and yield an age of 19.6 +/- 0.8 Ma (Table 2 and Fig. 3d). In the case of apatites, 23 crystals have been dated with low uranium contents. However, all ages are in a range from 14.0 to 6.1 Ma, with a central age of 8.6 +/- 1.7 Ma (RE < 1 %), (Fig.2 and 3d).

### *The blueschist eastern Queyras domain*

The P-T conditions in the meta-gabbros (Table 1) have been estimated at 350-450 °C for a minimum pressure of 9 kbar (Caby, 1996; Schwartz, 2002), compatible with the development of chloritoid-phengite assemblage in the metapelites implying  $T < 450^{\circ}\text{C}$  (Agard et al., 2001). The zircon central ages are 25.7 +/- 1.2 Ma in the Casse de Peyroun area and 21.7 +/- 1.6 Ma in the Rocca Nera area (Fig. 2). The apatite central ages are younger than the zircon central ages with ages of 14.2 +/- 2.0 Ma in the Cirque de la Gavie area and of 9.4 +/- 1.1 Ma in the Rocca Nera area (Fig. 2). All zircons have a small dispersion of the individual crystal ages, with the RE always lower than one per cent and a high  $P\chi^2$  test (Fig.3 c). However, in the case of Rocca Nera, the  $P\chi^2$  is lower (value > 55 %) than in the other localities because of the low number of induced and spontaneous counted tracks (Table 2). The AFT results show a similar evolution, with a low dispersion of the individual crystal ages (RE < 1 %), but their central ages are always lower than the zircon central ages obtained in the same locality (Fig.3 c).

### *The blueschist western Queyras domain*

This domain corresponds to the westernmost part of the Piedmont zone (Fig. 2). P-T is estimated at 8-10 kbar for 300-350°C in the meta-gabbros (Caby, 1996; Schwartz, 2002), consistent with the development of carpholite in metapelites (Agard et al., 2001). Five ages have been obtained, two on apatites giving rise to homogeneous results around 22 Ma, and three on zircons with ages strongly varying between 63.1 +/- 2.9 Ma (Cervièrès area) and 94.7 +/- 3.1 Ma (Rocher Blanc area) (Fig. 3b).

The ZFT ages show a very low dispersion for the individual crystal ages and a  $P\chi^2$  test (Fig. 3b). In the sample from the Cervièrès area, ages were obtained only from two crystals of zircon, but the high density of induced and spontaneous tracks allows a central age of 63.1 +/-

2.9 Ma. These zircon ages are significantly older than the ages obtained by the same technique for the easternmost domains. Apatite ages have been obtained only for Rocher Blanc and Lago Nero areas. In both cases, the individual age dispersion is low (RE < 1 %) and the  $P\chi^2$  test is high (value > 90 %, Table 2). The central ages are compatible at about 22 Ma (Fig. 3b). In the case of the Lago Nero apatites, high uranium content allows to perform measurements of the confined track length (Fig. 3b). The track length distribution defines a symmetrical shape centered on 13.5 +/- 1.2  $\mu\text{m}$  compatible with a constant and progressive cooling of the sample through the partial retention zone (Dimitru, 1989).

### **FT interpretation**

The FT data from both zircons and apatites show a pattern for a single age for each sample (Fig. 3), suggesting that they record the time at which the rocks passed through their closure temperature. The closure temperature is a function of cooling rates (Dodson, 1981) and the FT closure temperature of zircons is 235-265°C (Zeitler et al, 1982) and 90-120°C for apatite (Hurford et al., 1991) in rocks with rapid cooling rate (> 10°C/Ma).

The FT age data are summarized as follows:

- the Chenaillet massif was not affected by alpine metamorphism. Even the low-temperature metamorphism (greenschist facies and prehnite-pumpellyite facies) must be older than 65 Ma, suggesting that the ophiolite was obducted before 65 Ma;
- Monviso and eastern Queyras show progressively older ages towards the west and tectonic upper levels (Fig.4). The eclogitic Monviso ophiolites cooled below 110°C only 9 Ma ago. Combining zircon and apatite ages yields a mean cooling rate of 10 to 12°C/Ma (Fig.4);
- in western Queyras, the AFT ages are similar, around 22 Ma, whereas zircon ages vary widely. The wide variation in zircon ages is explained by low temperature metamorphic condition ( $T < 350^\circ\text{C}$ ), which is close to the fission-track annealing temperature of zircon



(330°C; Yamada et al., 1995; Tagami et al., 1996; Tagami and Shimada, 1996). Thus these ages have no geological significance. The extrapolation of the trend obtained for the ZFT ages towards the west suggests that the western Queyras probably cooled below 240°C between 39 and 31 Ma (Fig. 4);

- the cooling histories of different ophiolite-bearing metamorphic units are diachronous. The easternmost eclogitic units cooled later and thus exhumed later than the westernmost blueschist units.

### **Alpine temperature-time paths**

The combination of our results with the available chronological constraints for the internal zones of the southwestern Alps (see Table 3) allows T-t paths to be constructed for both eclogitic and blueschist facies units (Fig. 5).

#### *The blueschist Queyras units*

In eastern Queyras, Ar-Ar ages of phengites are interpreted as the beginning of cooling while in western Queyras, Ar-Ar ages correspond to the crystallization ages of phengites, synchronous with the HP paragenesis (Table 3). The T-t paths show a diachroneity in the cooling history of a minimum of 5 Ma (Fig. 5). The eastern Queyras is first metamorphosed and partly cooled while western Queyras did not yet reached its thermal metamorphic peak condition. Since 43 Ma (at T ~ 300°C), cooling trends reversed and the western unit cooled before eastern unit, until the exhumation ended. Cooling rates varied with time, but remain low between 4 and 11°C/Ma. The similar cooling paths for western and eastern Queyras suggest that they were exhumed by the similar process but not at the same time.

#### *The eclogitic Monviso domain*

The Monviso is characterized by high cooling rates ( $> 35^{\circ}\text{C}/\text{Ma}$ ) in the initial period, but the cooling rate decreased to  $4.9^{\circ}\text{C}/\text{Ma}$  between  $400$  and  $300^{\circ}\text{C}$  and ca.  $12^{\circ}\text{C}/\text{Ma}$  below  $250^{\circ}\text{C}$  (Fig. 5). Immediately after the pressure peak, eclogites were exhumed rapidly up to a depth of about  $20$  km. In a second stage, the cooling rate decreases. During this final exhumation, the shape of the temperature-time path for the Monviso domain is similar to the Queyras units. However, the Monviso eclogites were cooled and thus exhumed later than the blueschist facies rocks.

### **Tectonic consequences and discussion**

#### *Arguments for a major tectonic contact in the Piedmont Schistes lustrés*

The Schistes lustrés and Monviso unit are separated by a ductile normal fault (Ballèvre et al., 1990). Our chronological data confirm that this must be a major tectonic contact inside the Piedmont zone. The proposed interpretation is supported by different lithologies, metamorphic conditions and cooling ages between the two units. The blueschist facies domain is mainly composed of meta-sediments with small and isolated ophiolitic bodies whereas the eclogitic domain corresponds to slices of oceanic lithosphere with minor meta-sediments. More generally, our results suggest that the alpine HP-LT metamorphic zoning cannot be regarded as progressive change in the metamorphic grade from west to east. Instead, each metamorphic zone likely corresponds to a juxtaposed tectonic unit with its own cooling history.

#### *Arguments for a Paleocene to Eocene intra-oceanic subduction zone.*

Our data are consistent with an intra-oceanic subduction zone for the Alpine Piedmont zone (Fig. 6). It is difficult to estimate the location of Chenaillet ophiolite in the Piedmont oceanic domain, but it remained in the upper tectonic position. Emplacement of the Chenaillet unit

onto the European margin could have occurred between upper Cretaceous and Eocene. During the Paleocene or Eocene, the Piedmont zone was metamorphosed under blueschist facies conditions in an accretionary wedge, while the eclogitized meta-ophiolites represent pieces of the subducted oceanic lithosphere. During active subduction, pieces of eclogites are exhumed in a serpentinitized subduction channel (Hermann et al., 2001; Schwartz et al., 2001).

#### *Continental subduction and decrease of the exhumation rates*

The UHP metamorphism in the Dora-Maira massif (Fig. 5) occurred at about 31-35 Ma (Duchêne et al., 1997b; Gebauer et al., 1997; Cliff et al., 1998). Considering this massif as remnant of the European continental margin, the continental subduction started prior Oligocene time. A low convergence rate, of about  $6-8\text{mm.yr}^{-1}$ , between Africa and Europe from Late Cretaceous to Eocene (Jolivet and Facenna, 2000; Schettino and Scotese, 2002) suggests a minimum of 17 Ma for the Dora Maira massif to reach the depth of 120 km recorded by the UHP metamorphism. This implies that the continental subduction probably initiated before 45-50 Ma in the southwestern Alps, i.e. at the time of the HP metamorphism in the Monviso ophiolites. Thus we propose that the onset of exhumation of the Piedmont zone is broadly contemporaneous with the initial continental subduction of the Dora Maira massif. The increased exhumation rate of the Queyras Schistes lustrés and the Monviso after 30 Ma is synchronous with the onset of exhumation of the Dora Maira massif (Fig. 5) and the initiation of the intracrustal Pennic frontal thrust, which mark the transition from subduction to collision dynamics.

#### **Conclusions**

Coupled FT dating on zircons and apatites from meta-ophiolites of the Piedmont zone documents diachronous cooling histories for four different tectono-metamorphic units.

1- the Chenaillet unit was cooled below 110°C before 65 Ma;

2- The exhumation of eclogite and blueschist facies rocks is diachronous, which confirms the existence of a major tectonic contact inside the Piedmont zone between the two units.

Three tectonic domains likely represent remnants of three different structural levels of the same paleo-subduction zone. (i) The highest structural level corresponds to the Chenaillet unit as obducted upon the European margin. (ii) The blueschist Queyras units (eastern and western Queyras) corresponds to deformed accretionary wedge, while (iii) the Monviso eclogitic units the subducted oceanic slab and exhumed in a serpentinite channel.

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## Figure captions.

Fig. 1 - Simplified tectonic map of the western Alps, with Argentera-Mercantour (AM), Pelvoux (P), Mont Blanc (MB), Sesia (SE), Dent Blanche (DB), Dora-Maira (DM), Gran-Paradiso (GP), Monte-Rosa (MR); a-Schistes lustrés; b-ophiolites with Chenaillet massif (CHE).

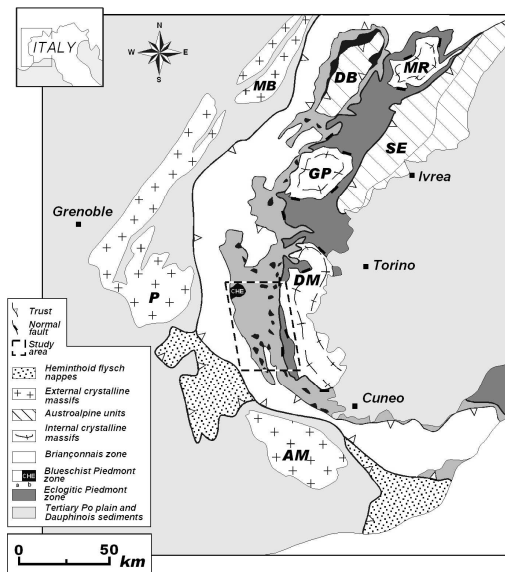


Fig. 2 - The new thermochronological data. 1-Chenaillet obducted ophiolite; 2-Briançonnais zone; 3-Blueschist western Queyras domain; 4-Blueschist eastern Queyras domain; 5-Eclogitic Monviso domain; 6-Acceglia zone; 7-Dora Maira crystalline massif; 8-Scattered blueschist ophiolites within the schistes lustrés complex. 9-Three main directions of lineaments, respectively N35E strike-slip faults and N60E and N160E normal faults (after Tricart et al., 2004). A-A' line locates the cross section in *figure 4*.

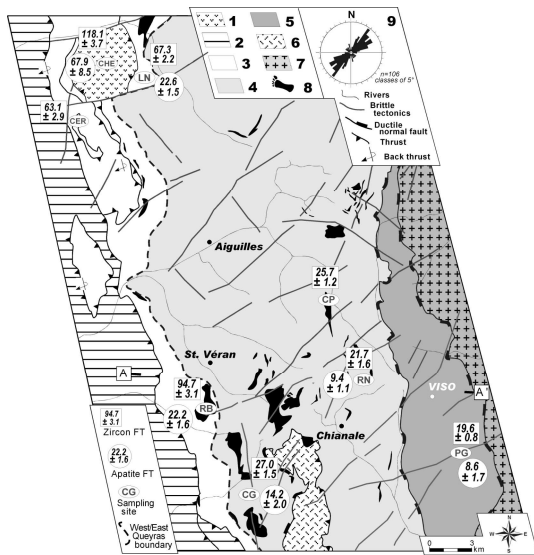


Fig. 3 - Radial plots of apatite and zircon ages for the different metamorphic domains of the Piedmont zone. The count data are presented using the radial plot of Galbraith (1990), corresponding to a graphical device for comparing crystals of differing ages and differing precisions. The position of the  $x$  scale records the uncertainty of individual grain estimates (dependant on the number of tracks), whilst each point has the same standard error in the  $y$  direction (indicated by the vertical  $+2$  to  $-2$  bar). The plot also includes a circular age scale such that the age of any crystal may be determined by extrapolating a line from the origin through the crystal's  $x,y$  coordinates to intercept the age scale. For the Lago Nero unit (LN) the apatite track-length measurements are presented. These data are displayed as a histogram with individual measurements grouped in  $1 \mu\text{m}$  bins. The number of tracks in each bin is represented on the  $y$  axis.

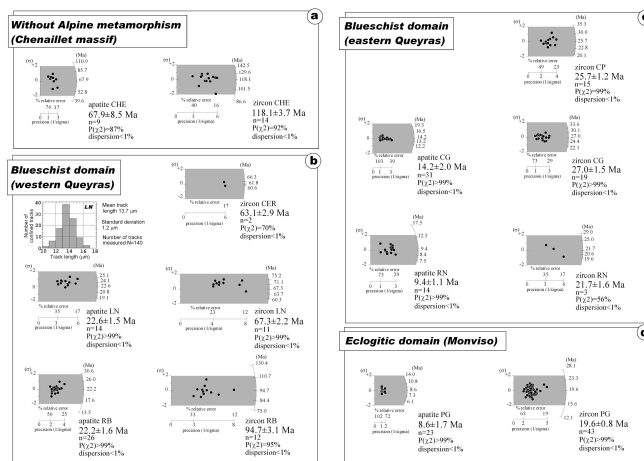


Fig. 4 - Cooling rates for the different tectono-metamorphic units presented on a cross section of the study area (A-A' transect in *figure 2*). The distribution of the cooling ages supports diachronous histories for the different units. In western Queyras, the non significant ZFT ages are related to the too low temperature conditions recorded during Alpine metamorphism (under 350°C). For the western Queyras the extrapolation of the trend obtained in the eastern part probably suggests that the western Queyras cooled below 240°C between 39 and 31 Ma. The ZFT and AFT ages are plotted in age / altitude diagram (we do not report ZFT data from the Chenaillet and the western Queyras). There is no clear correlation between the calculated ages and the altitudes, which suggests that the Piedmont zone was not exhumed as a single crustal unit. This is in agreement with the widespread occurrence of a post-metamorphic extensional brittle tectonics. As a consequence the Piedmont zone is composed of different blocks bounded by brittle faults.

Fig. 5 - Temperature-time paths for Monviso, western and eastern Queyras established by the compilation of geochronological data available in the Piedmont zone of the southwestern Alps. All radiometric data with references are presented in *Table 3*. Cooling rates are also indicated for each unit. These temperature-time paths support the diachronism of exhumation between blueschist and eclogitic units of the Piedmont zone.

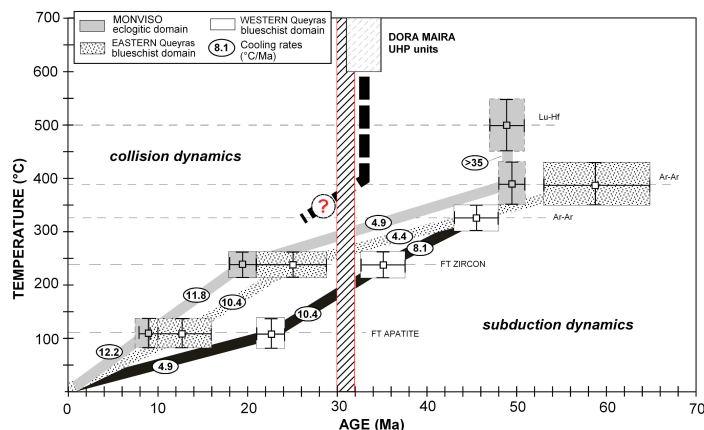
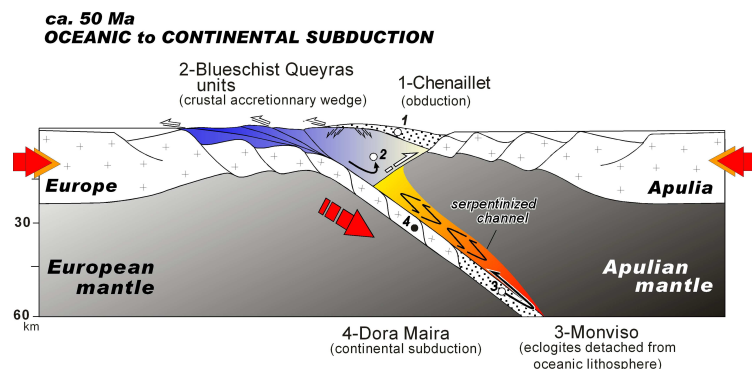


Fig. 6 - Tectonic sketch of the Piedmont zone during the Eocene. The different metamorphic domains that presently outcrop in the Piedmont zone “sample” different levels of a paleo-subduction zone.



Tab. 1 - Sample locations and petrological descriptions of rocks used for ZFT and AFT investigations. Elevation is in meters above sea level and geographic coordinates are latitude and longitude.

Sample	Location	geographic coordinates	Description	Elevation
PG	Passo Gallarino, Monviso Massif	44°38'N, 7°7'E	Centimetric to metric boudins of eclogitic Fe-Ti gabbro. The paragenesis consists of garnet, omphacite, glaucophane, rutile and K-quartz.	2621 m
RN	Rocca Nera, eastern Queyras	44°41'N, 6°59'E	Decimetric interstratified arkosic levels within the calcoschists. The rocks have suffered blueschist facies metamorphism with development of jadeite, glaucophane and zoisite.	2750 m
CP	Casse de Peyroun, eastern Queyras	44°44'N, 6°59'E	Cement of a magmatic breccia with doleritic elements. The blueschist paragenesis is characterized by development of jadeite, zoisite and glaucophane.	1956 m
CG	Cirque de la Gavie, eastern Queyras	44°36'N, 6°54'E	Albitic dykes boudinaged within serpentinites. The rock shows the association of albite, lawsonite (80%) and glaucophane developed during blueschist facies Alpine metamorphism.	2700 m
RB	Rocher Blanc, western Queyras	44°40'N, 6°52'E	Polygenic breccia with predominant granitoid elements, and subordinated metabasalt elements. The rock has recrystallized under blueschist facies conditions during Alpine metamorphism; the paragenesis consists of jadeite, lawsonite, phengite and glaucophane.	2700 m
LN	Lago Nero, western Queyras	44°53'N, 6°47'E	Breccia formed with detrital material of continental origin. The rock is composed of quartz, jadeite and phengite developed during blueschist facies Alpine metamorphism.	2030 m
CER	Cervièras, western Queyras	44°53'N, 6°44'E	Breccia formed with detrital material of continental origin. The rock consists of quartz, jadeite and phengite.	1630 m
CHE	Chenaillet massif	44°53'N, 6°44'E	Leucocratic zones cross-cutting a coarse-grained gabbro. Prehnite, epidote and calcite greenschist assemblage is related to the oceanic metamorphism.	2600 m

Tab. 2 - Global fission-track data measured for the Piedmont zone of the southwestern Alps. Zircons and apatites were separated using conventional rock-crushing, heavy liquids, Frantz magnetic separator and hand picking. Polishing and etching procedures were those recommended by Hurford (1990). All samples were analysed by the external detector method (Gleadow, 1981) which allowed the uranium concentration and age for individual crystals to be determined. Prior to irradiation, all zircons were etched using the KOH-NaOH eutectic at 220°C during 15 hours (Gleadow et al., 1976) and all apatites were etched with 1N HNO<sub>3</sub> at 20°C during 50 seconds. Samples were irradiated at the Orphée reactor of the Centre d'Etudes

Nucléaires of Saclay, France. Neutron fluences were monitored using uranium glass dosimeters (CN-1 for zircon irradiations and CN-5 for apatite irradiations). Ages were calculated using the zeta calibration approach, whereby a proportionality constant zeta is evaluated by multiple analysis of mineral standards of known age (Hurford and Green, 1983). Zeta values used in this study have been determined from repeated measurements of standard zircons from the Fish Canyon Tuff (USA) and standard apatites from the Durango deposit (Mexico) and the Fish Canyon Tuff (USA). Track densities ( $\rho$ ) ( $\times 10^5$  tracks  $\text{cm}^{-2}$ ), number of tracks counted (n) shown in brackets. All analyses by external detector method use 0.5 as  $4\pi/2\pi$  geometry correction factor. Apatite ages calculated using dosimeter glass CN-5; zircon ages calculated using dosimeter glass CN-1 (Hurford and Green, 1983).  $P\chi^2$  is the probability of obtaining  $\chi^2$  value for  $\nu$  degrees of freedom (where  $\nu$  = number of crystals - 1). RE is the relative error of central age. The ages correspond to the central ages following Galbraith (1992). These central ages are modal ages, weighted to account for the differing precision of the individual crystal age. Two uncertainties are related to the central age: the +/- indicates the analytical precision, while an estimate of the age dispersion (or the spread of the individual crystal data) is given by the relative standard error of the central age (RE). In our results the dispersion is always low ( $< 10\%$ ) leading to the recognition of a single age population. All uncertainties are quoted at  $1\sigma$  level.

Sample	Mineral	Number of Crystals	Spontaneous $P_s$ (Ns)	induced $P_i$ (Ni)	$P\chi^2$ %	RE %	Dosimeter $P_t$ (Nd)	Zeta ( $\pm\sigma$ ) $\zeta$	Age $\pm 1\sigma$ Ma
<b>Eclogitic domain (Monviso)</b>									
<b>PG</b>	apatite	23	0.049 (28)	1.59 (914)	>99	<1	19.39 (15899)	289 $\pm$ 13	8.6 $\pm$ 1.7
	zircon	43	2.43 (764)	2.25 (709)	>99	<1	2.90 (10734)	125 $\pm$ 3	19.6 $\pm$ 0.8
<b>Blueschist domain (eastern Queyras)</b>									
<b>RN</b>	apatite	14	0.810 (94)	24.03 (2786)	>99	<1	19.39 (15899)	289 $\pm$ 13	9.4 $\pm$ 1.1
	zircon	3	39.40 (164)	23.30 (97)	56	<1	2.05 (10854)	125 $\pm$ 3	21.7 $\pm$ 1.6
<b>CP</b>	zircon	15	5.27 (489)	2.64 (245)	>99	<1	2.05 (10854)	125 $\pm$ 3	25.7 $\pm$ 1.2
<b>CG</b>	apatite	31	0.255 (57)	5.01 (1120)	>99	<1	19.39 (15899)	289 $\pm$ 13	14.2 $\pm$ 2.0
	zircon	19	1.80 (285)	1.21 (192)	>99	<1	2.90 (10734)	125 $\pm$ 3	27.0 $\pm$ 1.5
<b>Blueschist domain (western Queyras)</b>									
<b>RB</b>	apatite	26	1.32 (263)	2.87 (574)	>99	<1	3.37 (9090)	289 $\pm$ 13	22.2 $\pm$ 1.6
	zircon	12	73.00 (2532)	9.86 (342)	95	<1	2.05 (10854)	125 $\pm$ 3	94.7 $\pm$ 3.1
<b>LN</b>	apatite	14	4.62 (374)	9.92 (803)	>99	<1	3.37 (9090)	289 $\pm$ 13	22.6 $\pm$ 1.5
	zircon	11	72.50 (2425)	13.80 (462)	>99	<1	2.05 (10854)	125 $\pm$ 3	67.3 $\pm$ 2.2
<b>CER</b>	zircon	2	165.00 (423)	33.60 (86)	70	<1	2.05 (10854)	125 $\pm$ 3	63.1 $\pm$ 2.9
<b>Without Alpine metamorphism (Chenaillet massif)</b>									
<b>CHE</b>	apatite	9	0.958 (57)	2.00 (119)	87	<1	9.88 (10454)	289 $\pm$ 13	67.9 $\pm$ 8.5
	zircon	14	80.50 (3385)	8.70 (366)	92	<1	2.05 (10854)	125 $\pm$ 3	118.1 $\pm$ 3.7