

# Structure of the 001 talc surface as seen by atomic force microscopy: Comparison with X-ray and electron diffraction results

Eric Ferrage, Gre&#769gory Seine, Anne-Claire Gaillot, Sabine Petit, Philippe De Parseval, Alain Boudet, Bruno Lanson, Jocelyne Ferret, Francois Marti

# ▶ To cite this version:

Eric Ferrage, Gre&#769gory Seine, Anne-Claire Gaillot, Sabine Petit, Philippe De Parseval, et al.. Structure of the 001 talc surface as seen by atomic force microscopy: Comparison with X-ray and electron diffraction results. European Journal of Mineralogy, E Schweizerbart Science Publishers, 2006, 18, pp.483-491. <10.1127/0935-1221/2006/0018-0483>. <hal-00192281>

HAL Id: hal-00192281 https://hal.archives-ouvertes.fr/hal-00192281

Submitted on 27 Nov 2007

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1		Structure of the {001} talc surface as seen by atomic force
2	]	microscopy: Comparison with X-ray and electron diffraction
3		results
4		
5	Eric	Ferrage <sup>1,*</sup> , Grégory Seine <sup>2</sup> , Anne-Claire Gaillot <sup>3</sup> , Sabine Petit <sup>4</sup> , Philippe de Parseval <sup>5</sup> ,
6	Alai	n Boudet <sup>2</sup> , Bruno Lanson <sup>6</sup> , Jocelyne Ferret <sup>7</sup> , François Martin <sup>8</sup>
7		
8	(1)	Department of Mineralogy, The Natural History Museum, Cromwell Road, London,
9		SW7 5BD, UK
10	(2)	Centre d'Elaboration des Matériaux et d'Etudes Structurales, CNRS, 29 rue Jeanne
11		Marvig, 31055 Toulouse Cedex 4, France
12	(3)	Department of Earth and Planetary Sciences, Johns Hopkins University, 34 and N.
13		Charles streets, Baltimore MD 21218, USA
14	(4)	Université de Poitiers, CNRS UMR 6532 'Hydr'A.S.A.', 40 avenue du Recteur Pineau,
15		F-86022 Poitiers Cedex, France
16	(5)	LMTG, CNRS UMR 5563, Observatoire Midi-Pyrénées, 14 Av. Edouard Belin, 31400
17		Toulouse, France
18	(6)	Environmental Geochemistry Group, LGIT – Maison des Géosciences, Joseph Fourier
19		University – CNRS, BP53, F-38041 Grenoble cedex 9, France
20	(7)	Talc de Luzenac S.A., BP 1162, F-31036 Toulouse Cedex, France
21	(8)	LASEH, UMR 6532 HYDR'ASA, Université de Limoges, 123 Av. Albert Thomas, Bât.
22		Les Dryades, 87000 Limoges, France
23		
24	* Aı	athor to whom correspondence should be addressed: e-mail : e.ferrage@nhm.ac.uk
25		

# 26 Abstract

27

28	In this study the surface structure of a centimetre sized crystal of talc from the
29	Trimouns deposit (Ariège, France) was imaged by atomic force microscopy. The direct image
30	shows detailed characteristics of clay tetrahedral surfaces. The unit-cell dimensions obtained
31	using atomic force microscopy ( $a_{or} = 5.47 \pm 0.28$ and $b_{or} = 9.48 \pm 0.28$ Å) are found to be
32	slightly higher, with an increased uncertainty, than those obtained using X-ray diffraction
33	$(a_{\rm or} = 5.288 \pm 0.007 \text{ and } b_{\rm or} = 9.159 \pm 0.010 \text{ Å})$ and selected-area electron diffraction
34	$(a_{\rm or} = 5.32 \pm 0.03 \text{ and } b_{\rm or} = 9.22 \pm 0.05 \text{ Å})$ . Talc has a quasi-ideal surface, clean of strong
35	structural distortion as compared to most of other clay minerals and unlikely surface
36	relaxation. The observation on the obtained image of apparent cell dimension enlargement is
37	then more likely attributed to instrumental artefacts, also responsible for scattered values of
38	unit-cell parameters rather than related to any surface structural features.
39	
40	Key-words : talc, atomic-force microscopy, surface structure, X-ray diffraction, selected-area
41	electron diffraction.
42	
43	1. Introduction
44	
45	Because of its beneficial properties to manufactured products, talc is widely used in
46	paper coating, paint, ceramics, and polymer industries. In the automotive industry, talc is
47	added to polymers to stabilize and harden automobile spare parts such as fenders, dashboards,
48	steering wheels, etc. However, because different talc varieties can exhibit a wide range of

49 physical and chemical properties it is necessary to optimize the use of these different types of

50 talc for specific industrial applications. In particular, it is necessary to take into account the

surface structure of talc which, for example, may or may not allow the epitaxial growth of
polymers (Ferrage et al., 2002).

53 Over the last decade, the availability of atomic force microscopes (AFM) has provided 54 new insights into mineral reactivity by providing images of crystal growth or dissolution at 55 the micro-to-nanometer scale (Hillner et al., 1992a and b; Kuwahara et al., 1998, 2001; 56 Bosbach et al., 2000; Lindgren, 2000; Bickmore et al., 2001; Brandt et al., 2003; Tournassat 57 et al., 2003; Aldushin et al., 2004; Yokoyama et al., 2005). AFM has also allowed the direct 58 visualization of the outer surfaces of minerals by imaging either groups of atoms at the 59 molecular scale (Drake et al., 1989; Weisenhorn et al., 1990; Hartman et al., 1990; Johnson et 60 al., 1991; Drake & Hellmann 1991; Wicks et al., 1992, 1993, 1998; Vrdoljak et al., 1994; 61 Kuwahara, 1999, 2001) or individual atoms (Lindgreen et al., 1991; Wicks et al., 1992; 62 Vrdoljak et al., 1994; Sokolov et al., 1997, 1999). Because of their perfect cleavage along the 63 {001} plane, which produces extremely flat external surfaces, phyllosilicates are especially 64 well-suited for an AFM study. 65 The structure of talc differs from that of mica, chlorite or lizardite which were 66 previously imaged by AFM (Hartman et al., 1990; Wicks et al., 1992, 1993, 1998; Vrdoljak 67 et al., 1994; Kuwahara, 1999, 2001; Lindgreen et al., 1991). Talc structure presents two 68 similar external tetrahedral faces which is suitable for microscopic observations. In addition, 69 the excess of octahedral charges is compensated for by deficiency in tetrahedral charges 70 (Martin et al., 1999), and the 2:1 sheet remains neutral, the interlayer of talc being thus devoid

of cations. Finally, the weak substitution in talc layers induce moderate structural distortions

such as tilts or tetrahedral rotations.

In this article we report the first atomic force microscope images of the surface of a
talc sample composed of very well-stacked layers from the Trimouns deposit (Ariège,
France). Images of the talc surface obtained from traditional AFM in repulsive mode are

compared with data in the literature and structural parameters of the bulk structure obtained
from X-ray diffraction (XRD) and selected-area electron diffraction (SAED).

78

79

80 2. Background

81

## 82 2.1. Geological setting and sampling

83 Talc (2:1 layer silicate) has a composition close to that of the  $[Mg_3Si_4O_{10}(OH)_2]$  end-84 member but with minor amounts of Fe, Al and F, and traces of Mn, Ti, Cr, Ni, Na and K 85 (Heller-Kallai & Rozenson, 1981; Noack et al., 1986; Abercrombie et al., 1987; Aramu et al., 1989; Coey et al., 1991; de Parseval et al., 1991, 1993; Martin et al., 1996, 1999). The sample 86 87 comes from the Trimouns talc and chlorite deposit, located in the French Pyrenees, ~100 km 88 south of Toulouse, at an altitude of 1,700 m. The origin of talc is well-established and is the 89 result of the hydrothermal alteration of wallrock in a zone of intense shearing between the 90 Saint Barthélémy dome and the low-grade Paleozoic metamorphic cover (Fortuné et al., 1980; 91 Moine et al., 1982, 1989; de Parseval et al., 1993). Dolostones of the Paleozoic cover were 92 transformed to talc, whereas the more siliceous and aluminous rocks (micaschists and granitic 93 pegmatites) evolved to a chlorite-dominant ore characterized by well-defined metasomatic 94 zones. The sampled talc of centimetric flakes that is commonly referred to as "talc flower" 95 when mined, was formed within geodes in the dolomite hanging wall of the ore body and is 96 associated with REE-minerals (de Parseval et al., 1997).

97

# 98 2.2. Crystallographic data

Talc is a 2:1 phyllosilicate structure (Fig. 1a) with two Si tetrahedral sheets (forming
an array of hexagonal rings of SiO<sub>4</sub> tetrahedra – Fig. 1b) each linked through the apical

101	oxygen to either, and leading to the presence in between of a sheet of octahedrally
102	coordinated Mg <sup>2+</sup> cations. Early structural characterizations of talc have described its
103	structure as monoclinic (C-1 space group – Gruner, 1934), but later publications have
104	described it as triclinic (C-1 space group - Rayner & Brown, 1973) or as pseudomonoclinic
105	( <i>Cc</i> ) but using a <i>P-1</i> space group (Perdikatsis & Burzlaff, 1981). In comparison to aluminous
106	phyllosilicates the misfit in lateral dimensions between tetrahedral and octahedral sheets is
107	smaller, leading to a better accommodation of the two sheets. As a result, the tetrahedral
108	rotation angle is diminished from $\sim$ 13-14° for muscovite to $\sim$ 3-4° for talc (Radoslovitch,
109	1961, 1962). The tetrahedral tilt angle out of the basal oxygen plane is very weak, giving an
110	essentially flat {001} surface. The moderate tetrahedral rotations induce a reduction of the
111	SiO <sub>4</sub> ring symmetry from hexagonal into ditrigonal (Fig. 1b).
112	
113	3. Experimental Methods
114	
115	3.1. Chemical analyses
116	The chemical composition of the talc sample used in this study was determined using a
117	Cameca SX50 electron microprobe. The sample was included in a resin and polished to obtain
118	perfect flat section. Operating conditions were 15 kV and 10 nA with beam size set to
119	$3 \times 3 \ \mu m$ . Standards used for calibration were: wollastonite for Si, corundum for Al, periclase
120	for Mg, haematite for Fe, pyrophanite for Mn and topaze for F.
121	
122	3.2. X-ray diffraction
123	To obtain information on the crystallographic parameters, particularly in the ab plane,
124	XRD data was collected on a randomly oriented sample. The centimetric talc flakes were
125	ground to a fine powder in an agate mortar. XRD data was collected with a Bruker D5000

126 diffractometer equipped with a Kevex Si(Li) solid-state detector and CuK $\alpha$  radiation.

Intensities were recorded at a  $0.02^{\circ}2\theta$  interval, from 2 to  $90^{\circ}$ , using a 50 sec counting time per 127 128 step. Accuracy on 20 positions was better than 0.001°. A rotating sample holder was used to 129 minimize the effect of preferential orientation. The openings of the divergence slit, the two Soller slits, the antiscatter slit, and the resolution slit were 0.5°, 2.3°, 2.3°, 0.5° and 0.06°, 130 respectively. Cell parameters were refined using U-Fit 1.3 software assuming a P-1 space 131 132 group (Perdikatsis & Burzlaff, 1981).

133

3.3. <sup>57</sup>Fe Mössbauer spectroscopy 134

A  ${}^{57}$ Fe Mössbauer absorption spectrum of the talc was collected over the range  $\pm 4$ 135 mm.s<sup>-1</sup> with a 512 channels analyzer. The Mössbauer spectrometer is composed of a compact 136 137 detector  $\gamma$ -system for high-counting rates and of a conventional constant-acceleration Mössbauer device (WISSEL). A <sup>57</sup>Co (in Rh) source with nominal activity of 50 mCi was 138 139 used. Talc flakes were finely ground under acetone (to minimize possible oxidation of Fe) and 140 the resulting powder was placed in a plexiglas sample holder. The spectrum was recorded 141 using a Canberra multichannel analyzer at 80°K to benefit from the second-order Doppler 142 effect. The isomer shift was recorded with respect to  $\alpha$ -Fe metal. As recommended by 143 Rancourt et al. (1992), the absorbtion thickness of the talc sample was calculated to minimize 144 the width of the absorption lines using data from the phlogopite-annite series as phlogopite 145 has a similar Fe-content to that of talc so as. The values are around 200 mg of mineral per cm<sup>2</sup>. Lorentzian lines shapes were assumed for decomposition, based on least-squares fitting 146 procedures. The  $\chi^2$  and misfit values were used to estimate the goodness of fit. 147 148

3.4. Selected-area electron diffraction 149

150 The electron diffraction study was performed on a Philips 420 transmission electron 151 microscope (TEM) equipped with a tungsten filament and operated at 120kV. Talc flakes 152 were gently crushing in distilled water and dispersed onto a 3mm holey carbon-coated Cu-153 mesh TEM grid. Selected-area electron diffraction (SAED) patterns were recorded after 154 orienting talc single crystals along the [001] zone axis using a double-tilt sample holder. The 155 camera constant of the talc SAED patterns was calibrated using a gold reference sample with 156 the crystals oriented along [100]. SAED patterns were recorded under the same parallel beam 157 illumination conditions for both gold and talc samples. The error on the camera constant was 158 estimated to be  $\sim 0.7\%$ . Distances between diffraction spots were measured on the negatives using a precision measurement table. 159 160

161 *3.5. Atomic force microscopy* 

162 The talc surface was investigated by atomic force microscopy (AFM) using a 163 Nanoscope II AFM from Digital Instruments. The talc sample was scanned in air using the 164 repulsive contact mode with constant deflection. The D-head piezoelectric scanner was 165 calibrated using HOP graphite and mica. The pyramidal shaped Si<sub>3</sub>N<sub>4</sub> cantilevers from Digital Instrument with 200 µm wide legs and a spring constant of 0.12 N.m<sup>-1</sup> were used. The talc 166 167 sample was cleaved along {001} plane surface prior to observation using adhesive tape. 168 Observations were performed using a vibration isolated platform and the thermal drift was 169 minimized by waiting for the AFM to equilibrate with ambient temperature. Contact forces, 170 scan speeds and scan directions were adjusted to optimize the resolution.

171

172 **4. Results** 

173

174 4.1. Mössbauer spectroscopy

175	Table 1 lists the iron site-occupancies of the talc spectrum fitting, taking into account
176	the quadrupole splitting distributions (Rancourt, 1994). The spectrum of the talc sample
177	obtained at 80°K (Fig. 2a) is similar to talc spectra reported by Martin et al. (1999). It shows a
178	difference in intensity between the two absorption bands, and a significant shoulder on the
179	side of the -0.5 mm.s <sup>-1</sup> peak. A good fit to the experimental data was obtained with four
180	Lorentzian doublets (Fig. 2b – Table 1). The first two, both with a large quadrupole splitting,
181	are due to $Fe^{2+}$ in octahedral sites. The third doublet with a small quadrupole splitting and a
182	low isomer shift is located within the first absorption band and attributed to $Fe^{3+}$ in the
183	tetrahedral site (Dyar & Burns, 1986; Rancourt et al., 1992; Rancourt, 1993; Martin et al.,
184	1999). The last doublet corresponds to Fe <sup>3+</sup> in octahedral sites (Blaauw <i>et al.</i> , 1980; Dyar &
185	Burns, 1986; Rancourt et al., 1992; Dyar, 1993; Rancourt, 1993; Martin et al., 1999).
186	According to this decomposition, Fe <sup>3+</sup> cations are thus present in both octahedral and
187	tetrahedral sites. However, $Fe^{2+}$ is the dominant iron species (81%), thus indicating reducing
188	conditions during the formation of talc.

## 190 4.2. Chemical composition

191 The structural formula of the talc sample was calculated on the basis of eleven192 oxygens from the average of ten bulk analyses obtained by electron microprobe:

193  $[Mg_{2.890}Fe^{2+}{}_{0.023}Mn^{2+}{}_{0.003}Fe^{3+}{}_{0.002}Al_{0.002}]_{\Sigma 2.917} [Si_{4.030}Fe^{3+}{}_{0.003}Al_{0.001}]_{\Sigma 4.034} O_{10} (OH_{1.975}F_{0.025})$ 

194 µThe distribution of iron between  ${}^{[6]}Fe^{2+}$ ,  ${}^{[4]}Fe^{3+}$  and  ${}^{[6]}Fe^{3+}$  was deduced from Mössbauer

195 spectroscopy results. Aluminum was split into <sup>[4]</sup>Al and <sup>[6]</sup>Al to insure the neutralization of

196 charges in the octahedral and the tetrahedral sheets, leading to the electroneutrality of the

197 layer. The composition of the talc sample is thus very close to that of the representative talc

198 determined by Martin et al. (1999) for the Trimouns deposit. Only minor substitutions are

present in the octahedral or tetrahedral sites leading to moderate local distortions in thetetrahedral sheet and thus to a nearly flat {001} surface.

201

# 202 *4.3. X-ray diffraction*

203 The powder diffraction pattern of the ground talc powder is characteristic of pure talc 204 with sharp 00 $\ell$  reflections that form a harmonic series ( $d_{001} = \ell \times d_{00\ell}$ ) thus indicating a well-205 ordered structure along the c\* axis without interstratification of chlorite layers (Fig. 3). Traces 206 of quartz and dolomite are also detected. Sample unit-cell parameters were refined in the 207 triclinic system (P-1 space group - Perdikatsis & Burzlaff, 1981) from the positions of all 208 measurable reflections between 2 and 90°20. Unit-cell parameters from the present study are 209 in agreement with those reported by Perdikatsis & Burzlaff (1981 – Table 2). It is possible to 210 deduce the  $a_{\rm or}$  and  $b_{\rm or}$  dimensions of the usual C-centered pseudo-orthogonal unit cell from 211 the refined abc unit-cell parameters defined in the triclinic system (Fig. 1b). As the  $\gamma$  angle is close to 120°,  $a_{\rm or}$  is similar to a and  $b_{\rm or}$  can be calculated as  $b_{\rm or} = a_{\rm or}\sqrt{3}$  (Table 2). 212

213

### 214 4.4. Selected-area electron diffraction

215 The SAED pattern of a talc single crystal observed along the [001] zone axis contains 216 very sharp *hk*0 diffraction spots arranged according to a pseudo-hexagonal pattern (Figure 4). In order to determine unit-cell dimensions in the  $\mathbf{a}^*\mathbf{b}^*$  plane,  $a_1^*$ ,  $a_2^*$  and  $a_3^*$  distances were 217 measured along [100], [010] and [110], respectively, and averaged to a single  $a^*$  dimension. 218 219 Measurements were repeated for twelve crystals perfectly aligned along [001], and all values 220 were then averaged to obtain a single  $a^*$  unit-cell dimension in the reciprocal space. However, 221 in order to determine the unit-cell dimension in the **ab** plane in the direct space, it is necessary 222 to take into account the 3D symmetry of the unit-cell:

223 
$$a = (\sin\alpha / a^*) / [1 + 2 \cos\alpha . \cos\beta . \cos\gamma - \cos^2\alpha - \cos^2\beta - \cos^2\gamma]^{-2}$$
(1)

224 As talc crystallizes in a triclinic system, unit-cell dimensions in the direct space cannot be 225 simply deduced from the distances measured on the SAED pattern without making some assumptions for the values of  $\alpha$ ,  $\beta$  and  $\gamma$  angles. Therefore,  $\alpha$ ,  $\beta$  and  $\gamma$  values from XRD 226 227 analysis (Table 2) were used in Eq. 1, and the following relationship was then obtained:  $a = 1.1545/a^*$ 228 (2) 229 Note that the same relationship can be obtained, without using results from XRD, if one assumes that (i) the  $\gamma$  angle is close to 120°, and that (ii) the  $\alpha$  and  $\beta$  angles are between 85 230 and 100° then the approximations  $\sin \alpha - 1$  and  $\cos^2 \alpha - \cos^2 \beta - \cos \alpha . \cos \beta - 0$  can be used. It 231 232 follows that from direct measurements of the SAED patterns and from the relation given above (Eq. 2), it is possible to deduce that  $a_{or} = a = 5.32(3)$  Å and that  $b_{or} = a_{or} \sqrt{3} = 9.22(5)$  Å 233 234 (Table 2).

- 235
- 236 4.5. Atomic force microscopy
- 237

### 238 Micro-topography of cleaved surface

239 A high-resolution image was recorded to observe the talc surface micro-topography  $(500 \times 500 \text{ nm}^2 - \text{Fig. 5})$ . Talc sample exhibits thick crystals in agreement with the 240 241 macrocrystalline character of the investigated sample. A perpendicular section of talc 242 topography is shown in Figure 5d and reveals that after cleavage these crystals are several tens of unit layers thick. This sample also has very large domains with flat surfaces which are 243 244 particularly suitable for AFM imaging. Note the variation in Z direction on these almost flat 245 regions that can be attributed to cantilever vibration which is greater when recording in air 246 condition as contrast to images recorded in liquid (Kuwahara et al., 1998, 2001). These 247 regions were chosen for recording molecular-scale image.

The raw image obtained at the molecular scale ( $5 \times 5 \text{ nm}^2$  - Fig. 6) shows an alignment 250 251 of triangles on the tetrahedral sheet surface. The of this raw image allows determining the m<sub>i</sub> 252 periodicity between adjacent triangles and the n<sub>i</sub> periodicities for the second neighbor 253 triangles along the perpendicular direction. Figure 7a shows the two-dimensional (2D) fast 254 Fourier transform (FFT) of the raw image. On this numerical diffraction pattern, spots are 255 distributed according to a pseudo-hexagonal pattern and different orders are visible that 256 indicate a well-defined periodic surface structure. It is then possible to index these spots 257 which correspond to m<sub>i</sub> and n<sub>i</sub> periodicities along two perpendicular directions in the raw 258 image (Fig. 7b). The regularity of the structure deduced from the sharpness of the diffraction 259 spots on 2D FFT is confirmed by direct periodicity measurements done on the raw image and reported in Figure 8, and leading to mean values of 5.47±0.28 Å and 9.48±0.28 Å for mi and 260 261 n<sub>i</sub> periodic distances, respectively.

262

#### 263 **5. Discussion**

264

### 265 5.1 Surface structure of talc

266 The raw image of talc surface (Fig. 6) is consistent with those found in the wealth of 267 literature devoted to AFM studies of tetrahedral sheet surfaces of phyllosilicates. If some 268 images present an arrangement of hexagonal rings corresponding to the basal oxygen plane of 269 tetrahedral sheet other images, obtained on micas, show a bright spot periodicity of about 5.1 - 5.3 Å (which could correspond to the layer unit-cell parameter of most phyllosilicates – 270 271 Hartman et al., 1990; Lindgreen et al., 1991; Wicks et al., 1992, 1993; Vrdoljak et al., 1994; 272 Kuwahara, 1999, 2001). This periodicity was initially interpreted as the hexagonal array of 273 interlayer cations but Kuwahara (1999) showed that it was possible to obtain both types of

274 image by varying the scan angle on the same sample. In addition, basal oxygen surface often 275 show the alternation of high and low positions for any tetrahedron defining a hexagonal ring. 276 This feature was first interpreted as resulting from the tetrahedral tilt of the basal oxygens but 277 is now commonly considered as a topography artefact induced by the interaction between an 278 asymmetric, double-atom, cantilever tip and the surface atoms (Gould et al., 1989; Vrdoljak 279 et al., 1994; Wicks et al., 1998; Kuwahara, 1999, 2001). According to these studies, triangles 280 on the unfiltered image (Fig. 6) represent a height exaggerated of half of a tetrahedra. Thus, 281 m- and n- periodicities between adjacent triangles correspond to  $a_{\rm or}$  and  $b_{\rm or}$  unit-cell 282 dimensions of talc. Although the AFM image was recorded in air, the image is most likely of 283 the basal oxygen surface rather than of some adsorbed species on the talc surface. Such 284 adsorption would indeed result in a blurry image with holes and/or aggregation of these 285 molecules. However, because cantilever drift easily removes interlayer cations from the mica 286 surface, in spite of the strong electrostatic binding forces (Kuwahara, 1999, 2001), a similar 287 "cleaning" of adsorbed species from the uncharged talc surface is most likely to occur. 288 The 2D FFT in which pseudo-hexagonal pattern was shown with different orders has 289 indicated a well-defined periodic structure (Fig. 7). Back 2D FFT is commonly applied to 290 remove extraneous signal from images but appears to be a controversial method. The two 291 main criticisms are: i) important features of the raw image are removed and/or features that 292 were not initially present are introduced, and ii) atomic positions are averaged. Wicks et al. 293 (1998) discussed thoroughly these potential pitfalls, pointing out that the first point is 294 essentially dependent on the competence of the user who should check carefully the FFT 295 spots. Then, if all the structure information is used in the back 2D FFT operation, only high-296 frequency noise is removed, and the overall image enhancement does not affect structure 297 details. Wicks et al. (1998) also demonstrated that following this image processing, atoms 298 displaced from their ideal positions were not averaged as initially supposed. Figure 9a shows

the filtered image from the back 2D FFT of areas outlined on the numerical diffraction pattern (Fig. 7a). The filtered image (Fig. 9a) is essentially similar to the raw image (Fig. 6) but shows a spectacular enhancement of structural features, mainly alternate tetrahedra. The individual tetrahedra positions can be drawn to reveal the surface structure of talc (Fig. 9b). Note the presence of a small "hump" in the siloxane cavity. However, because of evident artefacts due in particular to the interaction between the double-atom cantilever tip and the sample, extreme care should be taken in the interpretation of these features in structural terms.

307 5.2 Talc unit-cell parameters

To determine the unit-cell dimension of the talc surface layer, approximately 100 measurements in each of the m and n three directions (corresponding to  $\mathbf{a}_{or}$  and  $\mathbf{b}_{or}$  directions, respectively) were performed on the raw images (Fig. 8). The frequency distribution of m and n indicates a single-mode distribution for each which agrees with published distributions (Vrdoljak *et al.*, 1994; Kuwahara, 1999, 2001).

313 The  $a_{\rm or}$  and  $b_{\rm or}$  parameters measured on the talc surface by AFM are 5.47±0.28 Å and 314 9.48 $\pm$ 0.28 Å, respectively (Table 2). The uncertainty is quite high (~5% and 3% for  $a_{or}$  and 315  $b_{or}$ , respectively), but is consistent with other AFM studies on clay minerals (between ~2%) 316 and 8%, e.g., Vrdoljak et al., 1994; Kuwahara, 1999, 2001) regardless of the number of 317 measurements performed on a single image and of the number of images processed. The 318 obtained range for  $a_{\rm or}$  and  $b_{\rm or}$  parameters includes the values expected from the bulk structure 319 determination although the mean values are high. Similar large unit cell parameters have been 320 reported for other phyllosilicates, but the uncertainty systematically includes the ideal unitcell dimensions (Vrdoljak et al. (1994) and Kuwahara (1999) on chlorite and muscovite, 321 322 respectively). This enlargement of unit-cell parameters was attributed to surface relaxation. However, in contrast to micas or chlorite, no interlayer sheet or cation contributes to 323

interlayer cohesion in talc. As a result, surface relaxation is unlikely for talc and the observed
variation in unit-cell dimensions may rather be attributed to instrumental effects as evoked by
Vrdoljak *et al.* (1994). Note that improved unit-cell dimensions could be obtained by
recording images in liquid environment rather than in air conditions as done in the present
study (Kuwahara 1999, 2001; Sokolov *et al.*, 1997, 1999;).

329 Unit-cell dimensions obtained from XRD and SAED are consistent (Table 2) despite 330 difficulties intrinsic to SAED such as the calibration of the camera constant, the alignment of 331 a single crystal along [001] and the inability to precisely determine  $\alpha$ ,  $\beta$  and  $\gamma$  angles for a 332 triclinic system. Moreover, the uncertainty on the measured values is low for the two methods (~0.1% and 0.6% for XRD and SAED, respectively – Table 2). In contrast,  $a_{or}$  and  $b_{or}$ 333 334 parameters measured on the talc surface by AFM are scattered and the resulting uncertainty 335 on the unit-cell parameter is much higher as discussed above. XRD and SAED (providing a 336 good calibration of the camera constant) are more accurate methods to determine unit-cell 337 parameters because of the improved statistics (for XRD) and of the enhanced sensitivity to 338 crystal geometry.

Because the sample used in the present study presents a quasi-ideal surface structure with extremely limited tetrahedral tilts/rotations and limited, if any, surface relaxation, the variation in unit-cell dimensions can therefore rather be attributed to instrumental variability (calibration, tip-sample interaction). Extreme care should thus be used when interpreting the changes in unit-cell dimensions in terms of structural features for other clay minerals such as micas, kaolinites and chlorites in which such structural deformations are likely

345

```
6. Conclusion
```

348 In relation with the peculiar crystallinity of the studied talc sample from the Trimouns 349 deposit (Ariège, France), it has been possible using atomic force microscopy to obtain a 350 detailed image of the talc surface at the molecular scale. However, in spite of the high-quality 351 image obtained, artefacts that lead to the observation of only one tetrahedron out of two were 352 detected (double-atom cantilever tip interaction). In addition, the very peculiar crystal-353 chemistry of talc (very limited tetrahedral tilts and tetrahedral rotations, unlikely structure 354 relaxation) has revealed the existence of additional instrumental effects. These effects lead to 355 a slight distortion of the image that make a structural interpretation of these distortions, 356 described in previous studies for other clay minerals (micas, chlorites), ambiguous. Finally 357 the comparison of structure dimensions with data obtained by XRD and SAED shows AFM 358 limitations for unit-cell dimensions determination because of the increased uncertainty and of 359 the possible distortion (enlargement) by instrumental effects.

360

#### 361 Acknowledgements

362

EF and FM are grateful to the S.A. Talc de Luzenac for financial support. TEM
analyses were supported by NSF grant #EAR0409071. Fruitful discussions with Dirk
Bosbach (INE – Karlsruhe, Germany) and instructive comments on an earlier version of the
manuscript by Alain Baronnet (CRMCN Marseille, France) are also greatly acknowledged.
This study has beneficed of technical assistance from the Laboratoire de Chimie de
Coordination (Toulouse, France) for Mössbauer analysis. The manuscript was much improved
by the constructive reviews of Fred Wicks, Yoshihiro Kuwahara and AE Kenneth Livi.

370	References
371	
372	Abercrombie, H.J., Skippen, G.B. and Marshall, D.D. (1987): F-OH substitution in natural
373	tremolite, talc, and phlogopite. Contrib. Mineral. Petr., 97, 305-312.
374	Aldushin, K., Jordan, G., Fechtelkord, M., Schmahl, W.W., Becker, H.W., Rammensee, W.
375	(2004): On the mechanisms of apophyllite alteration in aqueous solutions. A combined
376	AFM, XPS and MAS NMR study. Clay. Clay Miner., 52, 432-442.
377	Aramu, F., Maxiav, G., Delunas, A. (1989): Mössbauer spectroscopy of talc minerals. Nuovo
378	<i>Cimento</i> , <b>11D</b> , 891-896.
379	Bickmore, B.R., Bosbach, D., Hochella, M.F., Charlet, L., Rufe E. (2001): In situ atomic
380	force microscopy study of hectorite and nontronite dissolution: Implications for
381	phyllosilicate edge surface structures and dissolution mechanisms. Am. Mineral., 86,
382	411-423.
383	Blaauw, C., Stroink, G., Leiper, W. (1980): Mössbauer analysis of talc and chlorite. J. Phys. I,
384	<b>41</b> , 411-412.
385	Bosbach, D., Charlet, L., Bickmore, B.R., Hochella, M.F.J. (2000): The dissolution of
386	hectorite: In-situ, real-time observations using Atomic Force Microscopy. Am.
387	Mineral., <b>85</b> , 1209-1216.
388	Brandt, F., Bosbach, D., Krawczyk-Barsch, E., Arnold, T., Bernhard, G. (2003): Chlorite
389	dissolution in the acid pH-range: A combined microscopic and macroscopic approach.
390	Geochim. Cosmochim. Ac., 67, 1451-1461.
391	Coey, J.M.D., Bakas, T., Guggenheim, S. (1991): Mössbauer spectra of minnesotaite and
392	ferrous talc. Am. Mineral., 76, 1905-1909.
393	De Parseval, P., Fontan, F., Aigouy, T. (1997): Composition chimique des minéraux de terres
394	rares de Trimouns (Ariège, France). C.R. Acad. Sci. II, 234, 625-630.

395	De Parseval, P., Fournes, L., Fortuné, J.P., Moine, B., Ferret, J. (1991): Distribution du fer
396	dans les chlorites par spectrométrie Mössbauer (57Fe) : Fe(III) dans les chlorites du
397	gisement de talc-chlorite de Trimouns (Pyrénées, France). C.R. Acad. Sci. II, 312,
398	1321-1326.
399	De Parseval, P., Moine, B., Fortuné, J.P., Ferret, J. (1993): Fluid-mineral interactions at the
400	origin of the Trimouns talc and chlorite deposit (Pyrénées, France). in "Current
401	Research in Geology Applied to Ore Deposits", P. Fenoll Hach-Ali, J. Torres-Ruiz, F.
402	Gervilla, eds. Univ. of Granada, Granada, 205-209.
403	Drake, B., Prater, C.B., Weisenhorn, A.L., Gould, S.A.C., Albrecht, T.R., Quate, C.F.,
404	Cannell, D.S., Hansma, H.G., and Hansma, P.K. (1989): Imaging crystals, polymers,
405	and process in water with atomic force microscope. Science, 243, 1586-1589.
406	Drake, B. & Hellmann, R. (1991): Atomic force microscopy imaging of the albite (010)
407	surface. Am. Mineral., 76, 1773-1776.
408	Dyar, M.D. (1993): Mössbauer spectroscopy of tetrahedral Fe(III) in trioctahedral micas.
409	Discussion. Am. Mineral., 78, 665-668.
410	Dyar, M.D. & Burns, R.G. (1986): Mössbauer spectral study of ferruginous one-layer
411	trioctahedral micas. Am. Mineral., 71, 955-965.
412	Ferrage, E., Martin, F., Boudet, A., Petit, S., Fourty, G., Jouffret, F., Micoud, P., De Parseval,
413	P., Salvi, S., Saint-Gérard, Y., Buratto, S, Ferret, J., Fortuné, J.P. (2002): Talc as
414	nucleating agent of polypropylene : morphology induced by lamellar particles addition
415	and interface mineral-matrix modelisation. J. Mater. Sci., 37, 1561-1573.
416	Fortuné, J.P., Gavoille, B., Thiebaut, J. (1980): Le gisement de talc de Trimouns près de
417	Luzenac (Ariège). Int. Geol. Congr., 26, E10, 43 p.

- 418 Gould, S.A.C., Burke, K., Hansma, P.K. (1989): Simple theory for the atomic force
- 419 microscope with a comparison of theoretical and experimental images of graphite.
  420 *Phys. Rev. B*, 40, 5363-5366.
- 421 Gruner, J.W. (1934): The crystal structure of talc and pyrophyllite. *Z. Kristallogr. Krist.*, 88,
  422 412-419.
- 423 Hartman, H., Sposito, G., Yang, A., Manne, S., Gould, S.A.C., Hansma, P.K. (1990):
- 424 Molecular scale imaging of clay mineral surfaces with the atomic force microscope.
  425 *Clay. Clay Miner.*, **38**, 337-342.
- Heller-Kallai, L. & Rozenson, I. (1981): The use of Mössbauer spectroscopy of iron in clay
  mineralogy. *Phys. Chem. Minerals*, 7, 223-238.
- Hillner, P.E., Gratz, A.J., Hansma, P.K. (1992a): Composite spiral growth kinetics of calcite
  revealed by AFM. *P. Soc. Photo-Opt. Inst.*, 1639, 160-170.
- Hillner, P.E., Manne, S., Hansma, P.K. (1992b): Atomic-scale imaging of calcite growth and
  dissolution in real time. *Geology*, 20, 359-362.
- 432 Johnsson, P.A., Eggleston, C.M., Hotchella, M.F. Jr. (1991): Imaging molecular-scale
- 433 structure and microtopography of hematite with the atomic force microscope. *Am.*
- 434 *Mineral.*, **76**, 1442-1445.
- Kuwahara, Y. (1999): Muscovite surface structure imaged by fluid contact mode AFM. *Phys. Chem. Minerals*, 26, 198-205.
- Kuwahara, Y. (2001): Comparison of the surface structure of tetrahedral sheets of muscovite
  and phlogopite by AFM. *Phys. Chem. Minerals*, 28, 1-8.
- 439 Kuwahara, Y. Uehara, S., Aoki, Y. (1998) Surface microtopography of lath-shaped
- 440 hydrothermal illite by tapping-mode and contact-mode AFM. *Clay. Clay Miner.*, **46**,
- 441 574-582.

442	Kuwahara, Y. Uehara, S., Aoki, Y. (2001) Atomic force microscopy study of hydrothermal
443	illite in Izumiyama pottery stone from Arita, Saga Prefecture, Japan. Clay. Clay
444	Miner., 49, 300-309.
445	Lindgreen, H. (2000): Electrical conduction in layer silicates investigated by combined
446	scanning tunnelling microscopy and atomic force microscopy. Clay Miner., 35, 643-
447	652.
448	Lindgreen, H., Garnæs, J., Hansen, P.L., Besenbacher, F., Lægsgaard, E., Stensgaard, I.,
449	Gould, S.A., Hansma, P.K. (1991): Ultrafine particles of North Sea illite/smectite clay
450	minerals investigated by STM and AFM. Am. Mineral., 76, 1218-1222.
451	Martin, F., Ildefonse, P., Hazemann, J.L., Grauby, O., Decarreau, A. (1996): Random
452	distribution of Ge and Si in synthetic talc: an EXAFS and FTIR study. Eur. J.
453	Mineral., <b>8</b> , 289-299.
454	Martin, F., Micoud, P., Delmotte, L., Marechal, C., Le Dred, R., De Parseval, P., Mari, A.,
455	Fortuné, J.P., Salvi, S., Beziat, D., Grauby, O., Ferret, J. (1999): The structural
456	formula of talc from the Trimouns deposit, Pyrénées, France. Can. Mineral., 37, 975-
457	984.
458	Moine, B., Fortuné, J.P., Moreau, P., Viguier, F. (1989): Comparative mineralogy,
459	geochemistry and conditions of formation of two metasomatic talc and chlorite
460	deposits : Trimouns (Pyrénées, France) and Rabenwald (Eastern Alps, Austria). Econ.
461	<i>Geol.</i> , <b>84</b> , 1398-1416.
462	Moine, B., Gavoille, B., Thiebault, J. (1982): Géochimie des transformations à l'origine du
463	gisement de talc et chlorite de Trimouns (Luzenac, Ariège, France). I - Mobilité des
464	éléments et zonalités. Bull. Mineral., 105, 62-75.
465	Noack, Y., Decarreau, A., Manceau, A. (1986): Spectroscopic and oxygen isotopic evidence
466	for low and high temperature origin of talc. Bull. Mineral., 109, 253-263.

- 467 Perdikatsis, B. & Burzlaff, H. (1981): Strukturverfeinerung am Talk Mg<sub>3</sub>Si<sub>4</sub>O<sub>10</sub>(OH)<sub>2</sub>. Z.
  468 *Kristallogr.*, **156**, 177-186.
- 469 Radoslovitch, E.W. (1961): Surface symmetry and cell dimensions of layer lattice silicates.
  470 *Nature*, **191**, 67-68.
- 471 Radoslovitch, E.W. (1962): The cell dimensions and symmetry of layer lattice silicates. I.
- 472 Some structural considerations. II. Regression relations. *Am. Mineral.*, **47**, 599-636.
- 473 Rancourt, D.G. (1993): Mössbauer spectroscopy of tetrahedral Fe(III) in trioctahedral micas.
  474 Reply. *Am. Mineral.*, **78**, 669-671.
- 475 Rancourt, D.G. (1994): I. Inadequency of lorentzian-line doublets in fitting spectra arising
  476 from quadrupole splitting distributions. *Am. Mineral.*, 21, 244-249.
- 477 Rancourt, D.G., Dang, M.Z., Lalonde, A.E. (1992): Mössbauer spectroscopy of tetrahedral
  478 Fe(III) in trioctahedral micas. *Am. Mineral.*, **77**, 34-43.
- 479 Rayner, J.H. & Brown, G. (1973): The crystal structure of talc. *Clay. Clay Miner.*, 21, 103480 114.
- Sokolov, I.Y., Henderson, G.S., Wicks, F.J. (1999): Theoretical and experimental evidence of
  "true" atomic resolution under non-vacuum conditions. *J. Appl. Phys.*, 86, 55375540.
- 484 Sokolov, I.Y., Henderson, G.S., Wicks, F.J., Ozin, G.A. (1997): Improved atomic force
- 485 microscopy resolution using an electric double layer. *Appl. Phys. Lett.*, **70**, 844-846.
- 486 Tournassat, C., Neaman, A., Villiéras, F., Bosbach, D., Charlet L. (2003): Nanomorphology
  487 of montmorillonite particles: Estimation of the clay edge sorption site density by low-
- 488 pressure gas adsorption and AFM observations. *Am. Mineral.*, **88**, 1989-1995.
- 489 Vrdoljak, G.A., Henderson, G.S., Fawcett, J.J., Wicks, F.J. (1994): Structural relaxation of the
- 490 chlorite surface imaged by the atomic force microscope. *Am. Mineral.*, **79**, 107-112.

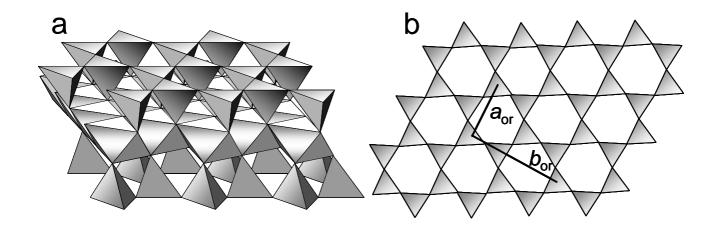
491	Weisenhorn, A.L., Mac Dougell, J.E., Gould, Cox, S.A.C., Cox, S.D., Wise, W.S., Massie, J.,
492	Maivald, P., Elings, V.B., Stucky, G.D., Hansma, P.K. (1990): Imaging and
493	Manipulating Molecules on a Zeolite Surface with an Atomic Force Microscope.
494	Science, 247, 1330-1333.
495	Wicks, F.J., Henderson, G.S., Hawthorne, F.C., Kjoller, K. (1998): Evidence for atomic-scale
496	resolution in atomic force microscopy of layer silicates. Can. Mineral., 36, 1607-1614.
497	Wicks, F.J., Kjoller, K., Eby, R.K., Hawthorne, F.C., Henderson, G.S., Vrdoljak, G.A. (1993)
498	Imaging the internal atomic structure of layer silicates using the atomic force
499	microscope. Can. Mineral., 31, 541-550.
500	Wicks, F.J., Kjoller, K., Henderson, G.S. (1992): Imaging the hydroxyl surface of lizardite at
501	atomic resolution with the atomic force microscope. Can. Mineral., 30, 83-91.
502	Yokoyama, S., Kuroda, M., Sato, T., (2005): Atomic force microscopy study of
503	montmorillonite dissolution under highly alkaline conditions. Clay. Clay Miner., 53,
504	147-154.

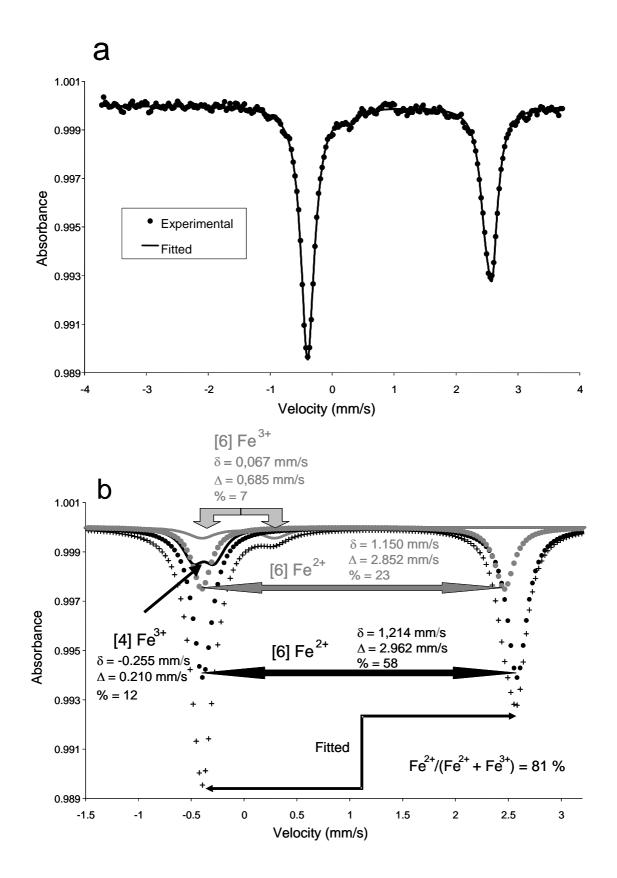
505 **Tables** 506 507 **Table 1.** Mössbauer parameters of talc sample at 80 K. Quadrupole splitting ( $\Delta$ ) and isomer shift ( $\delta$ ) are given in mm.s<sup>-1</sup>. 508 Fe<sup>2+</sup>  $\mathrm{Fe}^{3+}$  $\mathrm{Fe}^{2+}$  $(Fe^{2+}+Fe^{3+})$ % δ\* % Site Site Δ δ\* Δ 1.15 2.85 -0.25 0.21 12 [4] 81 81 [6] 2.96 0.07 1.21 0.68 7 [6] *Note:* δ values relative to Fe-metal. 509 510 Table 2. Comparison of cell parameters. Distances are given in Angstroms and angles in 511 degrees.  $a_{\rm or}$  and  $b_{\rm or}$  represent dimensions of the equivalent C-centered unit-cell with  $\gamma_{\rm or} =$ 512 90°). b a С β  $b_{\rm or}$ α  $a_{\rm or}$ γ Perdikatsis & 5.291(3) 5.290(3) 9.460(5) 98.68(5) 85.27(5) 119.90(5) 5.293(4) 9.168(8) Burzlaff (1981) XRD refinement 5.294(1) 5.281(1) 9.469(1) 98.71(1) 85.11(1) 119.98(1) 5.288(7) 9.159(10) (this study) Electron diffraction 5.32(3) 5.32(3) 9.22(5) (this study) AFM 5.47(28) 9.48(28) \_ \_ (this study) Note: Unit-cell parameters from Perdikatsis & Burzlaff (1981) are transformed to appear in the C-1 space group.

514

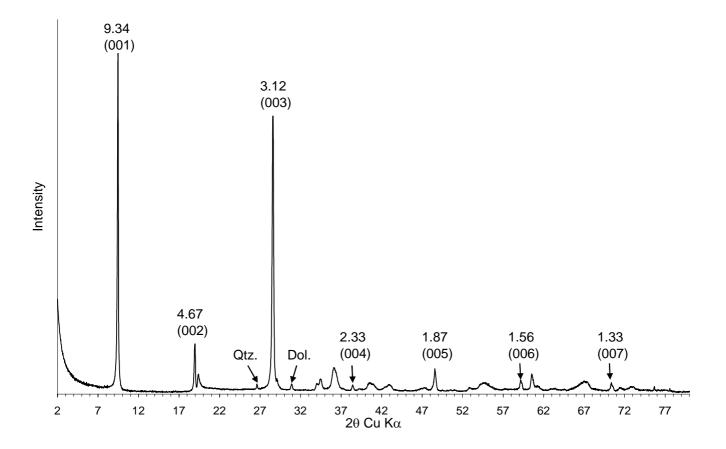
513

515	Figure Captions
516	
517	<b>Figure 1.</b> Talc structure along $c^*$ (a) and the tetrahedral surface plane where a and b cell
518	parameters in a quasi-orthogonal system $(a_{or}, b_{or})$ are indicated (b).
519	
520	Figure 2. Experimental and fitted Mössbauer spectrum of talc sample at 80 K (a) and
521	different components of the fitted spectrum (b).
522	
523	Figure 3. X-ray-diffraction pattern of talc sample. Dol. and Qtz. represent dolomite and
524	quartz traces, respectively.
525	
526	Figure 4. SAED pattern of talc sample recorded along the [001] zone axis.
527	
528	Figure 5. Micro-topography of talc surface (a) and the same image with isoaltitude curves
529	separated by 70 Å to enhance topography contrasts (b). (c) and (d) represent profiles
530	of cleaved surface morphology.
531	
532	<b>Figure 6.</b> AFM image of talc surface at molecular scale $5 \times 5 \text{ nm}^2$ .
533	
534	Figure 7. Two-Dimensional Fourier Transform applied to molecular scale AFM image (Fig. 6
535	- a) and indexation of spots in reciprocal space (b).
536	
537	<b>Figure 8.</b> Histogram of $a_{or}$ and $b_{or}$ cell parameters from AFM molecular scale image.
538	
539	Figure 9. Image shown in Fig. 5 after application of back 2D FFT to remove extraneous
540	signal (a). An enlarged view of a image with a redraw of individual tetrahedrons
541	(theoretical tetrahedral rotation not being taken into account).

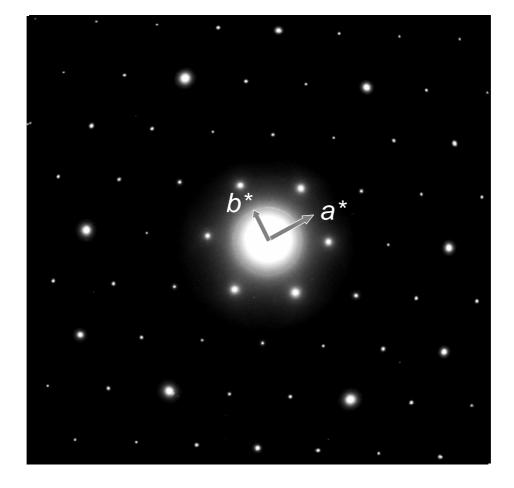


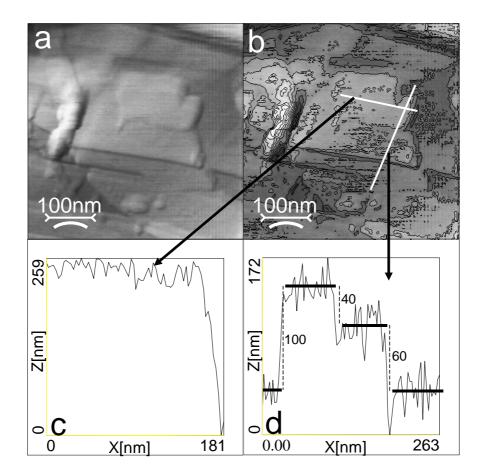


Ferrage et al., Fig. 02

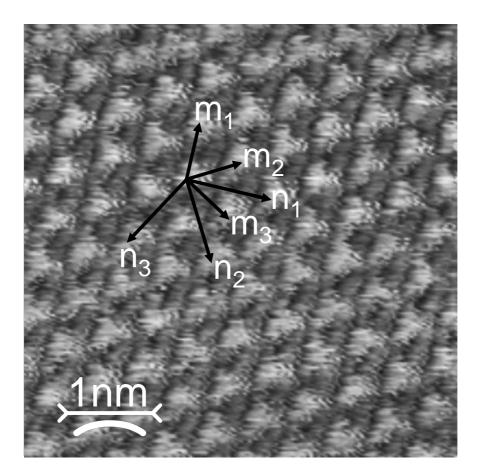


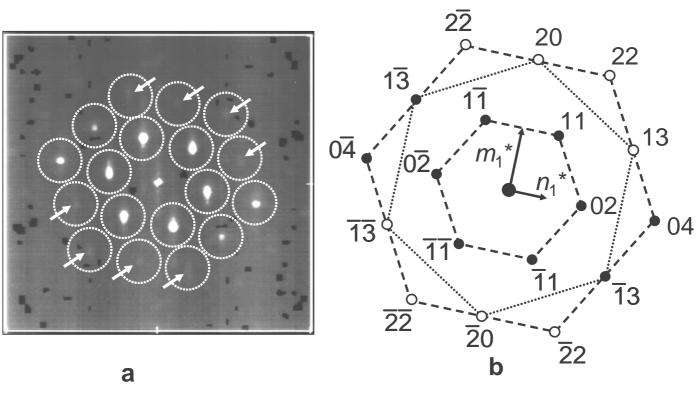
Ferrage et al., Fig. 03





Ferrage et al., Fig. 05





a

Ferrage et al., Fig. 07

