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The impact of catchment source group classification on the accuracy of sediment fingerprinting outputs

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11	
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13	
14	Abstract
15	
16	The objective classification of sediment source groups is at present an under-investigated
17	aspect of source tracing studies, which has the potential to statistically improve
18	discrimination between sediment sources and reduce uncertainty. This paper investigates this
19	potential using three different source group classification schemes.
20	The first classification scheme was simple surface and subsurface groupings (scheme 1). The
21	tracer signatures were then used in a two-step cluster analysis to identify the sediment source
22	groupings naturally defined by the tracer signatures (scheme 2). The cluster source groups
23	were then modified by splitting each one into a surface and subsurface component to suit
24	catchment management goals (scheme 3). The schemes were tested using artificial mixtures

of sediment source samples. Controlled corruptions were made to some of the mixtures to

mimic the potential causes of tracer non-conservatism present when using tracers in natural
fluvial environments. It was determined how accurately the known proportions of sediment
sources in the mixtures were identified after unmixing modelling using the three
classification schemes.

The cluster analysis derived source groups (2) significantly increased tracer variably ratios 5 (inter-/intra-source group variability) (up to 2122%, median 194%) compared to the surface 6 and subsurface groupings (1). As a result, the composition of the artificial mixtures was 7 8 identified an average of 9.8% more accurately on the 0-100% contribution scale. It was found that the cluster groups could be reclassified into a surface and subsurface component (3) with 9 10 no significant increase in composite uncertainty (a 0.1% increase over scheme 2). The far smaller effects of simulated tracer non-conservatism for the cluster analysis based schemes (2 11 and 3) was primarily attributed to the increased inter-group variability producing a far larger 12 13 sediment source signal that the non-conservatism noise (1). Modified cluster analysis based classification methods have the potential to reduce composite uncertainty significantly in 14 15 future source tracing studies.

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17 Keywords: Sediment fingerprinting; Sediment sources; Discrimination, Tracing, Uncertainty18

19 Highlights:

20

Robust discrimination between sediment sources is essential for fingerprinting
Source groups were classified according to management goals and tracer signatures
Objective classification reduced intra- and increased inter-group variability
Objective classification significantly reduced uncertainty in unmixing model outputs
The impacts of tracer non-conservatism were reduced with objective classification

2 1. Introduction

3

4 Sediment fingerprinting has become a key method of determining the importance of the 5 sediment sources in a catchment (e.g. Collins et al. 2010a). However, several methodological 6 uncertainties associated with existing fingerprinting procedures have been highlighted in 7 recent publications (D'Haen et al. 2012; Koiter et al. 2013; Smith and Blake, 2014; Laceby 8 and Olley, 2015; Pulley et al. 2015a & b). Establishing a robust discrimination between 9 sediment sources using suitable tracers is a key recommendation for accurate source tracing (Collins and Walling, 2002), making it a goal of many sediment fingerprinting based studies. 10 11 A fairly robust discrimination between different land uses as well as subsurface (i.e. streambank) sources has been established using some tracers. For example, ¹³⁷Cs or excess 12 ²¹⁰Pb, where the mixing of tracer fallout through the soil profile during ploughing results in 13 lower activities in cultivated land in comparison to undisturbed grassland or woodland 14 (Walling and Woodward, 1992). Additionally, very low activities would be expected in 15 subsurface sources, which are not exposed to direct fallout (Collins and Walling, 2002; 16 Walling, 2004). However, fallout radionuclides (i.e. ¹³⁷Cs) my not produce ideal source 17 identification in many catchments. For example, robust discrimination might be limited in 18 many catchments if channel banks are composed of displaced surface material or floodplain 19 deposits (with ages >1950's), which has high activities (Walling, 2003). Alternatively, where 20 environmental factors only support shallow channel banks comprising surface soils rather 21 than well-developed vertical faces and processes of diffusion, bioturbation and eluviation 22 cause migration of ¹³⁷Cs down through the soil profile (Walling and Woodward 1992; 23 24 Walling, 2003; Mabit et al., 2008), or where agricultural rotation between arable crops and short-term ley or untilled grass reduces the distinction between cultivated and undisturbed 25

1 surface soils (Smith and Blake, 2014). As a result, there is often incomplete discrimination between sediment source using ¹³⁷Cs, or similarly, excess ²¹⁰Pb (e.g. Collins et al. 2001; 2 Collins et al. 2007; Smith and Blake 2014). Because of this, it is common practice that ¹³⁷Cs 3 4 and many other individual tracers are utilised in a composite fingerprint consisting of many tracers, to help avoid spurious source-sediment matches (e.g. Stanton et al. 1992; Collins et 5 6 al. 2013). Due to the complex nature of the dynamics of most tracers in the environment, the basis for source discrimination of many of the tracers utilised in composite fingerprints is 7 8 rarely understood, and instead, a 'black box' type methodology is commonly used, which has 9 been criticised (Koiter et al. 2013; Smith and Blake, 2014).

A review by Haddadchi et al. (2013) shows that it is currently practice to classify sediment 10 source groups by land use (including surface/ subsurface sources) in the vast majority of 11 sediment fingerprinting research (Slattery et al. 1995; Walling and Woodward, 1995; Collins 12 13 et al. 2010a; Collins et al. 2010b; Smith and Blake, 2014). This is despite the fact that the majority of catchments investigated in source tracing studies will contain heterogeneous 14 15 geology or soil types. For example, Pulley et al. (2015a), Collins et al. (2013), Smith and Blake (2014), Wilkinson et al. (2013), Palazón et al. (2015), Nosrati et al. (2014), Evrard et 16 al. (2013) and Gellis et al. (2009) have all recently published source tracing studies in 17 18 catchments with heterogeneous geology and/or soil. This potentially represents a problem if the signal of different land use is weakly expressed by the tracers used. Horowitz and 19 Stephens (2008) investigated the impact of land use on the chemistry of river sediment in a 20 large scale study of 51 river basins across the USA, with drainage areas ranging from 28 to 21 49,800 km². It was found that the only land use to have a significant effect on sediment 22 chemistry was urban areas. Therefore, the geochemical signal of land use (and subsurface 23 sources) in river sediments may possibly be very weak. In contrast, the signal of geology or 24 soil type may often be very strong. For example, with the dissolution of magnetic iron oxides 25

1 which can take place in anoxic and reducing soil conditions (Anderson and Rippey, 1988; 2 Roberts and Turner, 1993), or highly different tracer concentrations in different geological units (Collins et al., 1998; Owens et al. 1999; Pulley et al. 2015c) or spatially variable 3 4 anthroprogenic tracer inputs (Devereux et al. 2010; Rossini et al. 2010; Guieu et al. 2010). These factors are likely to result in land use classified source groups with a very large amount 5 6 of within-group variability. The effect of a large within-source group variability is to significantly increase uncertainty associated with source apportionment results (Small et al. 7 2002; Collins et al., 2010; Pulley et al. 2010a). 8

The impacts of tracer non-conservatism caused by factors such as organic matter (Wang et al. 9 2010; Carr et al. 2010; Nadeu et al. 2011) and particle size (Elrick 1987; Motha et al. 2003; 10 Pye et al. 2007; Pulley et al. 2015c) may also be increased when a small difference in tracer 11 concentration exists between source groups. For example, if there is only a 10% difference in 12 13 the mean tracer concentration of two source groups and non-conservatism causes a 5% change to a tracer during sediment transport, very large errors will be present in the final 14 15 outputs. Alternatively, if a 100% difference exists between tracer signatures in the source 16 groups, a 5% change caused by non-conservatism during sediment mobilisation, intermediate storage and delivery will only have a minor impact on source fingerprinting estimates. 17

A method which could potentially reduce within-source group variability and increase intra-18 group variability was developed by Walling et al. (1993) who used pre-selected tracers in a 19 cluster analysis to classify sediment source groups. It was found that land use was the 20 primary controlling factor on tracer signatures and classified 4 to 6 source groups. Walling 21 22 and Woodward (1995) also used this method and geology was identified as the major controlling factor controlling source group classification. Using this method of source 23 classification provides the benefit that the natural variability in tracer concentrations within a 24 25 catchment is used to define the source groups; therefore, each source group should have a low

within-group variability in tracer concentrations and be substantially different to other
groups. Despite these clear advantages, this method of source group classification has largely
been neglected in recent literature. It is likely that catchment management goals such as
identifying sediment inputs from a specific source such as eroding farm tracks (Collins et al.
2010b) have necessitated the prior selection of source groups without regard to the natural
variability in tracers within a catchment.

7 The overall question this paper aims to answer is: can the objective classification of sediment 8 source groups using an updated cluster analysis based method reduce gross uncertainty in 9 fingerprinting outputs? Additionally, can we modify the cluster analysis derived source 10 groups to suit management goals; in this example discriminating between surface and 11 subsurface sources, while maintaining the benefits of the cluster analysis method.

This study uses artificial mixtures of sediment source samples, some of which are
deliberately corrupted by numerous means to test the accuracy of unmixing model results
when the different source group classification methods are used. Error evaluation using
artificial mixtures has been increasingly adopted as a routine component of sediment source
tracing studies (e.g. Palazón et al. (2015).

17 2. Study area

18

The sediment source samples were retrieved from the largest tributary sub-catchment (4.3 km²) of Sywell Reservoir, which is located in the Nene river basin in the East Midlands of the UK. The catchment is composed of Jurassic age mudstones and sand and ironstones in the lower catchment as well as Quaternary diamicton in the upper catchment (Figure 1). Soils in the catchment are a combination of freely draining brown earths in the lower catchment over

1 the ironstone geology and poorly draining clayey soils in the upper catchment. The land use 2 is predominantly cultivated land (54.4%) used for wheat production with some areas of improved grassland (22.7%) which are used for sheep grazing, as well as woodland (22.7%) 3 4 (Figure 1; Morton et al. 2011). The River Nene basin has an average annual rainfall of 638 mm recorded at Althorp over the last 140 years according to records transcribed by the 5 authors from the UK Met Office archives. Construction of Sywell reservoir was completed in 6 1906, and an area of wetland has developed in alluvial deposits where the river enters the 7 8 reservoir close to sampling points 1 and 1b (Figure 1). Very little erosion of toposils was observed in the study catchment, with a single small area of cultivated land appearing to have 9 undergone some minor rill erosion. Channel banks were observed to have slumped and be 10 11 exposed to fluvial entrainment in many areas. A previously published fingerprinting 12 investigation in the River Nene basin by Pulley et al. (2015a) identified that there were large differences (24%) between the provenance predictions made by different sediment tracer 13 groups when tracing using land use source categories. Therefore, the Nene basin represents a 14 15 challenging environment for the successful application of robust sediment source 16 fingerprinting.



- 2 Figure 1: The geology and land use in the study catchment (after Morton et al 2011;
- 3 British Geological Survey 2011) and the locations of sediment source sampling points.

3. Methods

3.1.Sediment source sampling and laboratory analyses

Sediment source samples were collected from 11 locations along the rivers channel banks
 (Figure 1). Samples were only collected from the bank material and topsoils directly above it
 as this study aimed to investigate different source classification methods comprising two
 fundamental source categories (surface and subsurface) to simplify the interpretation of
 results.

At each sampling point, 2 to 10 samples of the channel bank material were collected at 10 to 6 7 15 cm intervals down the exposed channel bank face according to vertical stratigraphy; using 8 a non-metallic knife after 5 to 10 cm of superficial material had been removed in order to minimise contamination by mass failure surface drapes and flood deposits. The sampling 9 locations were selected primarily on the basis of the presence of exposed banks with a lack of 10 vegetation and accessibility and to be roughly evenly spaced along the entire channel network 11 length. An additional sample of topsoil was collected using a non-metallic trowel to a depth 12 13 of 5 cm from each sampling location in the cultivated or grass fields located past the riparian zone and outside of the limits of any floodplain (~10 m from the river channel). Each sample 14 15 was an individual sample and not a composite of multiple samples. A total of 58 subsurface 16 channel bank samples and 20 surface samples (the top 5 cm of the banks and nearby field topsoils) were collected. Each sample was oven dried at 40°C for 24 hours before being 17 sieved to $< 63 \mu m$ to conform to common practice in published fingerprinting studies (e.g. 18 19 Walling et al., 1993; Walling and Woodward, 1995; Collins et al., 1997, 2010a).

20 Mineral magnetic (Walden et al. 1997), geochemical (Collins et al. 2010a) and colour

signatures (Pulley and Rowntree, 2016) were measured as potential sediment source tracers.

22 Mineral magnetic signatures were measured using 8 to 10 g of each sample tightly packed

23 into 5 ml polystyrene sample pots. The properties shown in Table 1A were measured using

the methods described by Lees (1999). The repeat measurement of six samples for five

repetitions identified that a mean error (coefficient of variation; %) of 5.3% was associated
 with the measurement of magnetic tracers.

3	Geochemical tracers were measured using 0.8 g of each sample digested in 10 ml of aqua
4	regia at 180°C for 20 minutes in a CEM Mars 6 digestion unit. The concentrations of Al, B,
5	Ba, Ca, Cu, Fe, K, Li, Mg, Mn, Mo, Ni, P, Pb, S, Sr, V, Y, Zn and Zr were determined using
6	a Thermo Scientific iCAP 6500 dual view ICP-OES. The repeat measurement of samples
7	identified that a mean error (coefficient of variation; %) of 11.6% was associated with the
8	measurement of geochemical tracers.
9	Colour signatures were measured using the prepared samples packed into polythene bags.
10	Images of the source material were captured using a Lexmark x2650 colour scanner and were
11	imported into Gimp 2 open source image editing software. The mean intensity of reflected
12	red, green and blue light was recorded on the 0-255 scale of the RGB colour model. The
13	colouration indices shown in Table 1B were then calculated using the extracted RGB values.
14	The methods used for measurement are discussed in more detail by Pulley and Rowntree
15	(2016). The repeat measurement of samples identified that a mean error (coefficient of
16	variation; %) of 4.1% was associated with the measurement of colour signatures.

17

Table 1. The magnetic properties (Maher, 1988 Walden, 1999; Yang et al. 2010; Wang
et al. 2012) and colour signatures (Ray et al. 2004; Viscarra Rossel et al. 2006) used,
their calculation and the property they represent. All measurements were initially
performed on the <63 μm fraction.

Name	Calculation	Property	Instrument		
(A) Magnetic signatures					
Low frequency susceptibility (χ lf)	Raw data	All magnetic minerals	Bartington Instruments MS2b sensor		

Susceptibility of anhysteretic (zarm) remanance magnetisationARM x 3.14 x 10stable single domain terrinagnetic grains in to 0.02 to $0.4 \mu m$ rangeMolspin® anhysteretic remanent magnetiser; Molspin® pulse magnetiser Molspin@ slow-speed spinner magnetometerSaturation isothermal remanence magnetisation (1T) (SIRM)Raw dataAlmost all remanence carrying mineralsMolspin@ pulse magnetiser. Molspin@ slow-speed spinner magnetiser Molspin@ slow-speed spinner magnetometerBack isothermal remanence magnetisation (-100mT) (IRM- 100)Raw dataThe majority of single domain ferromagnetic grainsMolspin@ pulse magnetiser Molspin@ slow-speed spinner magnetometerHard isothermal remanence magnetisation (HIRM)IRM:rr/(1(1 x (IRM-100mT) / IRM rr/1))/2High coercivity canted antiferromagnetic grainsCalculated(B) Colour signaturesIRM:rr/(1(1 x (IRM-100mT) / IRM rr/1))/2High coercivity canted antiferromagnetic grainsCalculatedBack isothermal remanence magnetisation (HIRM)IRM:rr/(1(1 x (IRM-100mT) / IRM rr/(1(1 x) (IRM-100mT) / IRM-100mT) / IRM-100mT (IRM-100mT) / IRM-100mT)Molspin@ p	Frequency dependent susceptibility (χfd)	$((\chi_{\rm lf} - \chi_{\rm hf})/m) x 100$ (m = sample mass)	Ultrafine super paramagnetic grains (< 0.03 μm)	Bartington Instruments MS2b sensor
Saturation isothermal remanence magnetisation (1T) (SIRM)Raw dataAlmost all remanence carrying mineralsMolspin® pulse magnetiser, Molspin® slow-speed spinner magnetometerBack isothermal remanence magnetisation (-100mT) (IRM- 100)Raw dataThe majority of single domain ferromagnetic grainsMolspin® pulse magnetiser Molspin® slow-speed spinner magnetiser Molspin® slow-speed spinner magnetiserHard isothermal remanence 	Susceptibility of anhysteretic (χarm) remanance magnetisation	ARM x 3.14 x 10	stable single domain ferrimagnetic grains in the 0.02 to 0.4µm range	Molspin® anhysteretic remanent magnetiser; Molspin® slow-speed spinner magnetometer
Back isothermal remanence magnetisation (-100mT) (IRM- 100)Raw dataThe majority of single domain ferromagnetic grainsMolspin® pulse magnetiser Molspin® slow-speed spinner magnetometerHard isothermal remanence magnetisation (HIRM) $IRM_{TT}/(1-(1 x))^{T/IRM}$ $(IRM-100mT / IRM)^{TT}))/2$ High coercivity catted antiferromagnetic grainsCalculated(B) Colour signatures $IRM_{TT}/(1-(1 x))^{T/IRM}$ $(TT))/2$ High coercivity catted antiferromagnetic grainsCalculated(B) Colour signatures $Raw data$ Reflected red lightLexmark x2650GreenRaw dataReflected green lightLexmark x2650BlueRaw dataReflected blue lightLexmark x2650HRGB $(2xG)-R-B4$ HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R+G}{2}$ Chromatic informationCalculatedMateria(R-B) (R+B)Spectra slopeCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedColoration Index $(R-G)$ (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Saturation isothermal remanence magnetisation (1T) (SIRM)	Raw data	Almost all remanence carrying minerals	Molspin® pulse magnetiser, Molspin® slow-speed spinner magnetometer
Hard isothermal remanence magnetisation (HIRM)IRM1T/(1-(1 x (IRM-100mT / IRM (IR) (IR)High coercivity canted antiferromagnetic minerals or coarse multi- domain ferromagnetic grainsCalculated(B) Colour signaturesRedRaw dataReflected red lightLexmark x2650RedRaw dataReflected green lightLexmark x2650BlueRaw dataReflected blue lightLexmark x2650HRGB $(2xG)-R-B 4$ HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R+G}{2}$ Chromatic informationCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Back isothermal remanence magnetisation (-100mT) (IRM- 100)	Raw data	The majority of single domain ferromagnetic grains	Molspin® pulse magnetiser Molspin® slow-speed spinner magnetometer
(B) Colour signaturesRedRaw dataReflected red lightLexmark x2650GreenRaw dataReflected green lightLexmark x2650BlueRaw dataReflected blue lightLexmark x2650HRGB $(2xG)$ -R-B 4HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R-B}{2}$ Chromatic informationCalculatedMue Index $(2xR-G-B)$ (R+B)Spectra slopeCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedColoration Index $(R-G)$ (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Hard isothermal remanence magnetisation (HIRM)	IRM _{1T} /(1—(1 x (IRM- _{100mT} / IRM _{1T})))/2	High coercivity canted antiferromagnetic minerals or coarse multi- domain ferromagnetic grains	Calculated
RedRaw dataReflected red lightLexmark x2650GreenRaw dataReflected green lightLexmark x2650BlueRaw dataReflected blue lightLexmark x2650HRGB $(2xG)$ -R-B4HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R-B}{2}$ Chromatic informationCalculatedMue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedColoration Index $(R-G)$ (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	(B) Colour signatures			
GreenRaw dataReflected green lightLexmark x2650BlueRaw dataReflected blue lightLexmark x2650HRGB $(2xG)$ -R - B 4HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R-B}{2}$ Chromatic informationCalculatedSaturation Index $(R-B)$ (R+B)Spectra slopeCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedColoration Index $(R-G)$ (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Red	Raw data	Reflected red light	Lexmark x2650
BlueRaw dataReflected blue lightLexmark x2650HRGB $(2xG)$ -R - B 4HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R-B}{2}$ Chromatic informationCalculatedSaturation Index(R-B) (R+B)Spectra slopeCalculatedHue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedColoration Index(R-G) (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Green	Raw data	Reflected green light	Lexmark x2650
HRGB $(2xG)-R-B$ 4HueCalculatedIRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R-B}{2}$ Chromatic informationCalculatedSaturation Index $(R-B)(R+B)$ Spectra slopeCalculatedHue Index $(2xR-G-B)(G-B)$ Primary coloursCalculatedColoration Index $(R-G)(R+G)$ Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Blue	Raw data	Reflected blue light	Lexmark x2650
IRGB $\frac{R+G+B}{3}$ Light intensityCalculatedSRGB $\frac{R-B}{2}$ Chromatic informationCalculatedSaturation Index $(R-B)(R+B)$ Spectra slopeCalculatedHue Index $(2xR-G-B)(G-B)$ Primary coloursCalculatedColoration Index $(R-G)(R+G)$ Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	HRGB	<u>(2xG)-R –B</u> 4	Hue	Calculated
SRGB $\frac{R-B}{2}$ Chromatic informationCalculatedSaturation Index $(R-B)(R+B)$ Spectra slopeCalculatedHue Index $(2xR-G-B)(G-B)$ Primary coloursCalculatedColoration Index $(R-G)(R+G)$ Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	IRGB	<u>R+G+B</u> 3	Light intensity	Calculated
$\begin{array}{c c} Saturation \\ Index \end{array} & \underline{(R-B)}(R+B) & Spectra slope & Calculated \\ \\ Hue Index & \underline{(2xR-G-B)}(G-B) & Primary colours & Calculated \\ \\ Coloration Index & \underline{(R-G)}(R+G) & Soil colour & Calculated \\ \\ \\ \hline Redness Index & \underline{\frac{R^2}{(BxG^3)}} & Hematite content & Calculated \\ \end{array}$	SRGB	<u>R-B</u> 2	Chromatic information	Calculated
Hue Index $(2xR-G-B)$ (G-B)Primary coloursCalculatedColoration Index $(R-G)$ (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Saturation Index	<u>(R-B)</u> (R+B)	Spectra slope	Calculated
Coloration Index(R-G) (R+G)Soil colourCalculatedRedness Index $\frac{R^2}{(BxG^3)}$ Hematite contentCalculated	Hue Index	<u>(2xR-G-B)</u> (G-B)	Primary colours	Calculated
Redness Index $\frac{\underline{R}^2}{(BxG^3)}$ Hematite contentCalculated	Coloration Index	$(\underline{R-G})(R+G)$	Soil colour	Calculated
	Redness Index	(BxG^3)	Hematite content	Calculated

3.2. Sediment source group classification methods

The source samples collected were classified into different sediment source groups for tracing
 in the following three ways:

3 1: The simple method of classification into two fundamental groups; surface (topsoils 0-5 cm
4 depth) and subsurface (i.e. channel banks) sources.

2: A two-step cluster analysis based upon the methods of Walling et al. (1993) and Walling 5 6 and Woodward (1995) was used in SPSS 20 to determine the sediment source groups which 7 best fitted the measured tracer signatures. Prior to the cluster analysis, the tracer signatures were included in a principal component analysis with varimax rotation in SPSS 20 to identify 8 9 the tracers most strongly correlated with each principal component in the source samples. This was undertaken to simplify the variables input into the cluster analysis. The two-step 10 11 cluster analysis was repeated with between 2 and 6 clusters and the solution with the best silhouette coefficient (the smallest mean between-cluster distance minus the mean within-12 cluster distance, divided by the larger of the two distances) was used to define the catchment 13 14 source groups. This measure represents how well separated each cluster is from other clusters 15 and how closely related the data points in any individual cluster are.

16 3: The third method of source classification was to reclassify the surface and subsurface

17 sources of each cluster group (from method 2) into separate source groups. This classification

18 method was aimed at retaining the naturally present cluster groups while fully

19 accommodating the catchment management goal of discriminating between surface and

20 subsurface sources for the purpose of targeting sediment control strategies.

21

22 *3.3. Creation of the artificial mixtures of sediment sources*

1 The effectiveness of each of the three source group classification methods at reducing 2 uncertainty in the unmixing outputs was tested using artificial mixtures of the channel bank 3 and topsoil samples. Each mixture was created with known proportions of each sediment 4 source group derived using the three classification methods, these were generated only after applying the cluster analyses and discriminant analysis. An equal mass of sediment from each 5 6 source sample collected which was in each group was used unless otherwise specified. Some mixtures were deliberately corrupted in the ways shown in Table 2 in an attempt to replicate 7 8 potential ways in which the non-conservatism of tracers might occur in the natural 9 environment. Three repetitions for each mixture were unmixed, composed of the following proportions of surface and subsurface sources: 0.25:0.75, 0.50:0.50 and 0.75:0.25 (these were 10 11 converted into the proportions in the source groups derived using methods 2 and 3 by 12 knowing which group the individual source samples added to the mixtures were classified into). The overall question asked when fingerprinting each mixture is "how close to the actual 13 proportions of sediment in the artificial mixtures are the fingerprinting results derived using 14 15 the different classification methods?".

For the deliberately corrupted mixtures, the prepared mixtures were wet sieved through a 38 μ m stainless steel mesh using ultrapure distilled water and the 63-38 μ m and <38 μ m fractions retained for tracing. When organic matter was added, cotton wool (as organic matter of a uniform composition) was reduced to a powder using a blender and the appropriate mass added to each mixture.

22	Table 2: Th	e artificial	sediment	source mixtures	created a	nd their purpose
----	-------------	--------------	----------	-----------------	-----------	------------------

All channel bank and surface sources	bup classification change the 5 th to 95 th range of
with no corruption How does source grouncertainty produced	d by the Monte Carlo based unmixing model as well
as the error resulting	from measurement accuracy and the modelling
procedure? (i.e. how	close to the actual mixture composition are the
results derived using	the three classification methods?)

Only cluster 3 subsurface sources and all surface sources (this mixture was decided upon only after the use of the cluster analysis) Only cluster1 surface sources and all subsurface sources (this mixture was decided upon only after the use of the cluster analysis) Only a random 10% of samples from each source group	These three alterations to the mixtures investigate how much error can be caused by regional variability in sediment source inputs with each source group classification method. For example, if only a small part of channel bank composed of an unusual tracer signature undergoes mass failure contributing a disproportionally large amount of sediment.
All source samples with 10 - 30% of	How does the classification of source groups affect the error resulting
the sample mass added as organic	from the enrichment in sediment-associated organic matter during its
matter (cotton wool)	erosion, transport and storage?
All source samples sieved to <38µm	How does the classification of source groups affect the error that can
All source samples sieved to 63 -	result from particle size changes during sediment erosion, transport,
38µm	deposition and delivery.

- 1
- 2

3.4. Source group fingerprinting procedure

3 4

5 The key theory behind this paper is that the cluster analysis source group classification method will reduce the within-source group variability and increase the inter-source group 6 7 variability. To test if the classification methods achieve this aim, tracer variability ratios of the percentage difference in median tracer concentration between source groups divided by 8 9 the mean within-source group variability (coefficient of variation; %) were used (Pulley et al. 10 2015a). Prior to the identification of the composite fingerprints for tracing, any tracer with a maximum variability ratio lower than 1 in any pair of source groups was removed from 11 12 further analysis to help reduce the uncertainty present in the final results. A genetic algorithm driven linear discriminant analysis (GA-LDA; cf. Collins et al. 2012, 13 2013, 2014) was then used to identify the composite fingerprint of tracers best able to 14

15 discriminate between the sediment source groups. The GA-LDA was repeated for each of the

16 three sediment source group classification methods, to produce a unique composite

fingerprint for each. The percentage of source samples correctly classified into their
 respective group with the optimum fingerprint for each classification method was compared.

An unmixing model (Equation 1) was used to apportion the contributions of sediment in each 3 of the artificial mixtures (Table 2). Before inclusion in the model, all tracers were rescaled to 4 range between 0 and 1 by dividing each model value by the maximum value found in any 5 6 source group. The unmixing model incorporated Monte Carlo uncertainty analysis (Rowan et 7 al. 2000) which repeated the model for 3000 iterations, each iteration with a random tracer 8 value from within the range of the median +/- one median absolute deviation (MAD) of each source group. The model outputs were presented as the average median Monte Carlo result 9 with 5th and 95th percentile uncertainty error bars. No correction factors for organic matter or 10 particle size were used, and no weightings for within-source variability and discriminatory 11 efficiency were applied, as these may introduce additional uncertainty into the fingerprinting 12 13 process (Smith and Blake 2014; Laceby and Olley, 2014). The results of the modelling were compared to the known proportions of each source group present in the artificial mixtures. 14 15 The mean absolute difference (cf. Collins et al., 1997) between the median Monte Carlo source estimations and the known proportions of each source was calculated for the 3 16 repetitions of each mixture (Table 2). Using this method, the error present when unmixing the 17 18 uncorrupted and deliberately corrupted mixtures was quantified to find how source classification affects the accuracy of a hypothetical fingerprinting study. 19

20

21 Equation 1. The structure of the sediment source unmixing model.

$$\sum_{i=1}^{n} \left\{ \left(C_i - \left(\sum_{s=1}^{m} P_s S_{si} \right) \right) / C_i \right\}^2$$

Where C_i = concentration of fingerprint property (i) in sediment sample; P_s = the optimised
percentage contribution from source category (s); S_{si} = median concentration of fingerprint
property (i) in source category (s) n = number of fingerprint properties comprising the
optimum composite fingerprint; m = number of sediment source categories.

- 5 4. Results and discussion
 6
 7
 8 4.1.Source group classification
- 9

10 The first classification scheme separated samples into simple surface and subsurface sources (two source groups used in most published studies). The second source group classification 11 12 used the tracer signatures in a two-step cluster analysis. A total of 7 Principal Components 13 were identified in the tracer signatures accounting for 82.3% of the variance in the total data set. The tracer most strongly correlated with each Principal Component was identified (and 14 are those shown in Table 3) and included in the two-step cluster analysis. The cluster analysis 15 identified that a 3 cluster solution was optimal. The results presented in Table 3 suggest that 16 these clusters are based on catchment geology (Figure 1) with cluster 1 representing 17 18 ironstone-derived topsoils and channel banks (rich in vanadium and iron), cluster 2 representing mudstone and diamicton (rich in lithium) and the third component representing 19 limestone-derived channel bank material (rich in strontium and calcium). It is of note that 20 21 limestone is not marked on the geology map (Figure 1) highlighting a potential shortcoming 22 of a source classification schemes based only upon a geology or soil map as opposed to one which uses the tracer signatures for classification. Therefore, in the case of this study, the 23 24 tracer signatures are likely to be more naturally defined according to local geology rather than as simple surface and subsurface sources. In other catchments this may be different, 25

- 1 reflecting land use (Walling et al., 1993) or soil type. These groupings were used to form the
- 2 artificial mixtures used for testing in this paper.
- 3
- 4 Table 3: The locations of cluster centres in the two-step cluster analysis (group
- classification method 2 only). The cluster with the highest value for each tracer is
 highlighted in bold (See Table 1 for a description of colour signatures).

	C	Cluster	
	1	2	3
	Oordial Ironstone and	Mudstone and	
Interpretation	Sand and ironstone	diamicton	Limestone
Percentage of total samples (78) in group	9.0	70.5	20.5
Green (intensity)	90.67	107.05	127.94
SIRM (10 ⁻³ Am ³ kg ⁻¹)	9.65	2.06	0.91
Sr (mg kg ⁻¹)	58.48	36.65	109.2
V (mg kg ⁻¹)	114.75	50.31	61.28
Li (mg kg ⁻¹)	11.12	15.07	9.81
Mo (mg kg ⁻¹)	0.75	0.56	0.55

The third source group classification method used the three cluster groups as a starting point 8 9 and split the surface and subsurface samples in each cluster group into their own separate 10 groups. After doing this it was found that there was only one subsurface sample left in the 11 ironstone subsurface cluster; as a result the subsurface ironstone source group was removed from further analysis. The source groups identified with scheme 3 can be seen in (Figure 2, 3) 12 The source groups derived using each classification scheme were mapped on a diagram 13 representing the down bank profiles at each sampling point (Figure 2). The location of each 14 sampling point in the study catchment is shown in Figure 1. Mapping the two-step cluster 15 analysis source groups (classification scheme 2) shows that the majority of ironstone 16

17 classified samples are located in the centre of the catchment and on the surface (Figure 2).

18 The limestone group classified samples are all located in subsurface material at sampling

- 1 sites 3, 4, 7 and 8. In contrast, clay and diamicton derived material is present throughout the
- 2 entire study area.
- 3 From the viewpoint of a catchment manager wanting to know where to target mitigation
- 4 measures, classification scheme 2, using only the cluster analysis, is perhaps the least useful
- 5 since it could only identify how much sediment originated from some of the small outcrops
- 6 of limestone and ironstone in the centre of the study catchment. The most useful
- 7 classification scheme is number 3 with the greatest number of source groups.
- 8
- 9 Figure 2: A simplified map of the source samples derived using the different source
- 10 group classification schemes; each diagram represents the down bank profile with
- 11 samples over the 0 m line originating from adjacent agricultural fields and samples
- 12 immediately below the line being topsoil overlying the channel bank.



8 group variability).

1	The cluster analysis derived source groups (classification scheme 2) greatly increased (by up
2	to 2122%, median 194%) the variability ratios over the simple surface and subsurface source
3	groups (scheme 1), indicating a greater difference in tracer signatures between the source
4	groups and lower within-source group variability (Table 4). The variability ratios remained
5	substantially higher than the simple surface and subsurface source groupings (scheme 1)
6	when the cluster analysis based groups were split into a surface and subsurface component
7	(scheme 3). These results thereby indicate that the lowest uncertainty would be propagated
8	through to the unmixing model outputs with classification schemes 2 and 3 compared to
9	scheme 1.

Table 4: Source group median tracer concentrations, median absolute deviations (MAD) and tracer variability ratios for the different classification methods (the tracers shown are those selected in the PCA as representing 82.2% of variance in the total tracer dataset, not necessarily those used in the composite fingerprints).

	1: Surface and subsurface source groups								
Group	Name		Number of samples	SIRM	Li	Мо	Sr	V	Green
1	Surface	Median	20	3.05	13.39	0.56	39.81	55.84	99.8
		MAD		0.87	2.96	0.13	13.45	14.19	3.2
2	Channel banks	Median	58	1.42	12.27	0.52	38.55	51.88	111.5
		MAD		0.61	2.54	0.12	13.68	12.15	7.4
Surface	and channel banks	Variabil	ity Ratio	1.50	0.39	0.29	0.09	0.29	2.15
		2: T	wo - step	cluster s	ource g	roups			

	2. 1 wo step cluster source groups									
1	Ironstone	Median	7	10.20	11.89	0.79	51.50	113.20	94.6	
		MAD		5.69	3.22	0.20	5.90	41.65	3.3	
2	Clay	Median	55	1.88	13.43	0.53	32.83	47.76	105.6	
		MAD		0.58	2.96	0.10	7.93	9.79	4.2	
3	Limestone subsurface	Median	16	0.86	9.81	0.49	107.48	50.51	126.3	
		MAD		0.23	2.50	0.12	46.58	16.04	6.2	
Ironstone and mudstone + diamicton		Variabili	ty Ratio	1.89	0.47	1.51	2.04	2.02	2.79	

	Ironstone and limestone subsurface	Variabili	Variability Ratio		0.67	1.54	1.90	1.62	6.01	
	Mudstone + diamicton and limestone subsurface	Variabili	ty Ratio	1.87	1.13	0.38	2.06	0.21	3.71	
		Maximu	m ratio	1.89	1.13	1.54	2.06	1.62	3.71	
3: Two - step cluster source with only clay surface sources reclassified										
	1 Ironstone surface	Median	7	10.20	11.89	0.79	51.50	113.20	94.60	
		MAD		5.69	3.22	0.20	5.90	41.65	3.30	
	2 Clay subsurface	Median	41	1.65	13.43	0.53	33.76	51.60	108.10	
		MAD		0.65	2.96	0.10	7.30	10.44	4.20	
	3 Limestone subsurface	Median	16	0.86	9.81	0.49	107.48	50.51	126.30	
		MAD		0.23	2.50	0.12	46.58	16.04	6.15	
	4 Clay surface	Median	14	2.33	13.39	0.54	28.53	47.37	101.15	
		MAD		0.89	2.96	0.12	7.11	12.23	1.75	
	Ironstone surface and mudstone + diamicton subsurface	Variabili	ty Ratio	1.76	0.47	1.51	2.08	1.91	3.39	
	Ironstone surface and limestone subsurface	Variabili	ty Ratio	2.21	0.67	1.54	1.90	1.62	6.01	
	Ironstone surface and mudstone + diamicton surface	Variabili	ty Ratio	1.64	0.46	1.35	2.45	1.86	2.48	
	Mudstone + diamicton subsurface and limestone subsurface	Variabili	ty Ratio	1.45	1.13	0.38	2.11	0.08	3.29	
	Mudstone + diamicton subsurface and clay surface	Variabili	ty Ratio	0.75	0.01	0.07	0.67	0.36	2.29	
	Limestone subsurface and mudstone + diamicton surface	Variabili	ty Ratio	1.93	1.12	0.40	2.15	0.22	6.03	
		Maximu	m ratio	2.21	1.13	1.54	2.45	1.91	6.03	

4.2 Discriminant analysis

2 3

The GA-LDA produced composite fingerprints able to classify 100% of the source samples into their correct groups for all of the source group classification schemes. On this basis, all three classification schemes are suitable for achieving basic discrimination using the available tracers. The cluster analysis grouping method (scheme 2) required fewer tracers to achieve this discrimination than the other groups. Recent research by Sheriff et al. (2015) has

- 1 suggested that larger composite fingerprints may reduce uncertainty in some fingerprinting
- 2 methodologies. It was, however, found that including additional tracers to increase the size of
- 3 the fingerprint for source classification scheme 2 did not result in a significant change to
- 4 unmixing model accuracy in this study, and for this reason, the results derived using the
- 5 original smaller number of signatures are presented.
- 6

7 Table 5: The optimum composite fingerprint selected for each source group8 classification scheme.

	Discriminatory power	Tracers selected
1: Original groups	100%	SIRM, Red, Green, HRGB, IRGB, Al, K, Li, Mn, P, Sr, V, Zn
2: Cluster analysis groups	100%	SIRM, IRGB, Fe, P, Sr
3: Cluster groups with separate surface and subsurface components	100%	χlf, χam, Red, Green, Blue, HRGB, Colouration Index, Ba, Fe, Li, Mg, Sr

9

10

4.1.Unmixing model outputs

11 12

The artificial mixtures of known proportions of source samples (Table 2) were run through the unmixing model (Equation 1) using the composite fingerprints in Table 5, to assess how the different source classification schemes affected the accuracy of the source apportionment modelling results. Six of the seven mixtures were deliberately corrupted by sieving adding organic matter or using only a small number of samples from each source group (Table 2). This was done to mimic some of the key possible sources of tracer non-conservatism in the natural environment.

20 Figure 3 shows some examples of the actual and modelled proportions of sediment from each

source group in the artificial mixtures derived for the different source group classification

schemes. By way of summary, the results are only presented for one of the three samples
 unmixed and four of the seven mixture types. The full set of graphs are provided in the online
 supplementary material and the results are summarised in Table 6.

4 The error bars representing the 5th to 95th percentile range of uncertainty in model results

5 were very large with the simple surface and subsurface source groups (mean for all samples

6 analysed 71.31% on the 0 - 100% contribution scale, standard deviation 18.19%) (scheme 1).

7 The range of uncertainty was smallest with the cluster groups (mean 31.05%, standard

8 deviation 12.90%) (scheme 2) and cluster groups split into surface and subsurface

9 components (mean 38.63%, standard deviation 19.93%) (scheme 3).







methods, with 5th and 95th percentile uncertainty bars. The latter represent feasible
unmixing model solutions.

3

Table 6 summarises the mean differences between the actual and modelled contributions of 4 each source to the sediment mixtures for every unmixing model run. The outputs of nearly 5 6 every model run were statistically significantly different (P < 0.05) to those of other models. 7 For example, the source apportionment results of the mixtures with organic matter added were significantly different to those without organic matter added. The simple surface and 8 9 subsurface source groupings (scheme 1) resulted in large errors (mean 15.8%), even when no alterations are made to the mixtures, but the actual composition of the mixtures did, however, 10 mostly fall within the large range of model uncertainty when the mixtures were unaltered 11 (Table 6; Figure 3). The mean percentage differences between median tracer concentrations 12 in the source groups for the composite fingerprints used were 16.75% (scheme 1), 42.11% 13 (scheme 2) and 34.71% (scheme 3) (Table 4). The small difference in tracer signatures 14 between source groups using scheme 1 explains its poor performance, as the errors associated 15 with laboratory tracer measurement were quantified as between 4.1% and 11.6%, which 16 could remove much of the discrimination provided by the tracers used. Schemes 2 and 3 17 produce far lower errors with the unaltered mixtures (Mean 7.7 and 10.1%). 18

Using only part of each source group (either 10% of source samples or a samples only from specific cluster group in Table 3) in the mixtures to replicate sediment delivery from only a small part of the study catchment resulted in large errors in provenance apportionment when source group classification scheme 1 was used (mean 21.7%). Classification schemes 2 and 3 had much lower errors (mean 12.9%). Therefore, where sediment delivery to a river is highly localised, significant errors could be introduced if source groups are classified on the basis of

1	catchment-wide generic subsurface / surface sources alone in a catchment with a
2	heterogeneous geology or soil type when a composite fingerprinting approach is used.
3	Sieving the artificial mixtures to <38 μ m and 63 - 38 μ m to replicate changes to fine
4	sediment particle size during its transportation from source to river channel resulted in large
5	errors (mean 28.6%) when scheme 1 was used and lower errors (mean 13%) when
6	classification schemes 2 and 3 were used. The largest error resulting from any deliberate
7	corruption to the artificial source mixtures was caused when the samples were sieved to 63 -
8	38 μ m (mean 23.8%). This large error possibly suggests that the basis for source
9	discrimination may be significantly different between the $<38~\mu m$ and 63 - $38~\mu m$ fractions
10	of the source samples. Previous research has reported such a finding. For example, Motha et
11	al. (2003) and Pye et al. (2007) found higher concentrations of many tracers in fine, <20 $\mu m,$
12	fractions of catchment source material. Alternatively, Horowitz and Elrick (1987) found
13	anthropogenic pollutants such as Zn concentrated in coarser silts of stream sediments.
14	When cotton wool was added to the artificial mixtures to replicate the enrichment of organic
15	matter during sediment transport, this alteration counterintuitively slightly improved model
16	accuracy (by a mean of 1.5%). It may be that the sediment coated the organic matter meaning
17	that the sediment colour was not significantly changed by the organic addition. Alternatively,
18	it is possible that measurement error (of up to 11.6%) caused the tracer concentrations
19	measured in the mixtures to be too high. In this case dilution of the tracer signatures by
20	organic matter may well result in the observed improvement.
21	

Table 6: Mean absolute differences (%) between the actual and modelled proportions of
each sediment source in the artificial mixtures using the different sediment source
classification schemes. The lowest difference for each mixture is highlighted in bold.

Source classification scheme

	1 Surface and subsurface source groups	2 Two-step cluster groups	3 Two-step cluster groups with separate surface and subsurface components	Mean
No alteration	15.8	$10.1^{n/a}$	7.7*	11.2
Cluster 3 channel banks samples only	16.1*	12.1*	12.3*	13.5
Random 10% of each source group	26.8*	n/a	15.5*	21.2
Cluster 1 surface samples only	22.3*	11.1*	13.7*	15.7
10 – 30% organic matter added	14.8	9.0*	5.4	9.7
Mixtures sieved to 63 - 38 μm	36.6*	15.6*	19.3*	23.8
Mixtures sieved to <38 μm	20.6	15.7*	12.7*	16.3
Mean	21.8	12.3	12.4	

⁺ Significantly different model output distribution to the simple surface and subsurface classifications (for contributions from surface sources), Kruskal Wallis test P <0.05.

* Significantly different distribution to the unaltered mixture, Kruskal Wallis test P < 0.05.

1

2 5. Conclusions

3

The findings of this paper demonstrate how small differences in tracer signatures between 4 5 sediment source groups and a high within-source variability can introduce significant 6 uncertainty into unmixing model results. As a result, it was found that the simple classification of catchment sources as generic surface and subsurface sources in a catchment 7 with a heterogeneous geology resulted in large amount of error when using a composite 8 fingerprinting approach. This error was significantly reduced by the cluster analysis based 9 method, and was not significantly increased by splitting the cluster analysis source groups 10 into surface and subsurface components to suit catchment management goals. Therefore, a 11

cluster analysis based classification method with the modification of cluster groups appears to
 be the optimum method within the Sywell Reservoir catchment. This is likely to be the case
 for many other river catchments.

The effects of tracer non-conservatism were found to be substantially reduced by the high 4 5 tracer variability ratio associated with the cluster analysis based classification methods. The 6 reasoning behind this is that the source group signal of the tracers (inter-group variability) is 7 larger than the noise of tracer non-conservatism with these methods. An additional advantage 8 to the cluster analysis based methods is that far smaller errors are introduced by highly localised sediment inputs from only a small part of the catchment, which may have highly 9 distinctive tracer concentrations (e.g. from the ironstone geology in the case of the Sywell 10 study catchment). 11

Whilst this paper found that the sediment source groups in the cluster analysis were strongly controlled by catchment geology it should be emphasised that this method is likely applicable to catchments with homogenous soil types, channel bank composition and anthropogenic tracer inputs. As a result, we would recommend consideration of objective source classification schemes in combination with the modification of source groups to suit management goals. On the basis of our findings here, the optimum classification scheme for applying sediment source fingerprinting in the Sywell catchment is presented in Figure 4).



2	Figure 4: A flow diagram of the optimum source classification scheme identified for the
3	Sywell reservoir catchment.

- 4
- 5
- 6

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12

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3 6. Online supplementary data

4

Table S1: Loadings of the 7 largest principal components in the PCA analysis of tracer concentrations in the source samples.

							Rotation Sums of Squared	
	Initial Eigenvalues				Extraction Sums of Squared Loadings			
		% of	Cumulative		% of	Cumulative		
Component	Total	Variance	%	Total	Variance	%	Total	
1	10.990	26.805	26.805	10.990	26.805	26.805	7.978	
2	7.610	18.560	45.365	7.610	18.560	45.365	4.776	
3	5.228	12.751	58.117	5.228	12.751	58.117	6.534	
4	3.676	8.967	67.083	3.676	8.967	67.083	5.886	
5	2.794	6.814	73.898	2.794	6.814	73.898	7.662	
6	2.031	4.954	78.852	2.031	4.954	78.852	2.915	
7	1.402	3.419	82.271	1.402	3.419	82.271	5.286	

7

8 Table S2. The PCA structure matrix of tracer signatures in the source samples values

9 larger than 0.4 and smaller than -0.4 are highlighted.

Structure Matrix								
	Component							
	1	2	3	4	5	6	7	
Green	970	122	056	.079	444	.234	042	
IRGB	969	129	042	.060	433	.228	029	
Red	919	122	.254	184	324	.131	.064	
HRGB	828	028	191	.262	485	.254	173	
Blue	786	114	435	.368	463	.301	138	
Redness index	.775	.176	.489	350	.568	180	.158	
LOI %	.769	.468	302	160	.209	055	206	
Hue Index	.755	.408	170	004	.208	.167	.053	
Мо	.122	.826	.064	.007	.138	110	360	
Mn	.089	.710	.348	233	.339	181	356	
Ba	.188	.648	081	103	.003	558	443	
K	.040	.643	078	.615	164	129	336	
Cu	.033	.615	114	.445	299	014	480	

Zn	.320	.581	.526	061	.368	.081	176
Р	.428	.493	.310	160	.471	.346	.112
V	.127	.137	.875	263	.515	.137	.129
Fe	.152	030	.847	115	.433	.283	.189
Υ	.169	.079	.838	265	.622	.173	.125
Ni	170	.226	.793	.188	.061	.262	069
Saturation index	066	032	.771	597	.217	240	.212
colouration index	.277	.038	.720	618	.390	271	.221
Zr	118	003	.720	.450	051	.142	033
SRGB	446	051	.682	539	.001	098	.197
Li	116	086	.072	.898	183	153	096
Mg	233	.280	072	.864	381	102	347
S	058	271	156	.762	100	.326	.134
В	.088	.458	184	.701	179	148	318
D90	.450	060	.135	642	.369	.017	.525
SIRM	.397	.086	.254	169	.969	002	.013
Xlf	.389	.114	.215	149	.960	.057	057
Xarm	.262	.057	.148	145	.918	.048	061
Irm -100mT	.416	.029	.223	138	.867	.053	054
HIRM	.245	.165	.180	146	.718	075	.091
Pb	.482	.502	038	144	.505	312	268
Sr	371	137	.248	106	029	.840	.155
Ca	372	332	.002	251	138	.740	.352
Al	.030	012	154	063	037	244	.176
SSA	093	.309	086	.160	.044	060	964
D50	.079	333	.171	296	.160	.231	.943
D10	.020	345	025	089	172	032	.903
Span	.189	.328	138	178	.000	268	829

Classification scheme 1



Classification scheme 2





Classification scheme 3



Figure S1: Actual and modelled sediment source contributions to the artificial mixtures. Actual contributions are in dark grey and modelled contributions are in light grey.