

Full title:

EFFECTS OF TOPOGRAPHY AND SURFACE SOIL COVER ON EROSION FOR
MINING RECLAMATION. THE EXPERIMENTAL SPOIL HEAP AT EL
MACHORRO MINE (CENTRAL SPAIN)

Short title:

TOPOGRAPHY AND SURFACE SOIL COVER IN MINING RECLAMATION

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ABSTRACT

Mining reclamation tries to reduce environmental impacts, including accelerated runoff, erosion and sediment load in the nearby fluvial networks and their ecosystems. This study compares the effects of topography and surface soil cover on erosion on man-made slopes coming from surface mining reclamation in Central Spain. Two topographic profiles, linear and concave, with two surface soil covers, subsoil and topsoil, were monitored for two hydrologic years. Sediment load, rill development, and plant colonization from the four profiles were measured under field conditions. The results show that, in the case of this experiment, a thick and non-compacted topsoil cover on a linear slope yielded less sediment than carbonate colluvium or topsoil cover on a concave slope. This study also shows that vegetation establishment, which plays an important role in erosion control, depends on topography. Plant cover was more widespread and more homogeneous on linear profiles with topsoil cover. On concave slopes, plant establishment was severely limited on the steepest upper part and favoured in the bottom. This

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4 37 study suggests that management of topography and surface soil cover should be approached
5 38 systematically, taking three outcomes into consideration: i) topsoil can lead to a successful
6 39 mining reclamation regardless of topography, ii) created concave slopes can lead to a successful
7 40 mining reclamation, and iii) topography determines the vegetation colonization pattern.
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12 42 **Key words:** topographical design, topsoil, constructed slopes, concave slopes, water erosion,
13 43 vegetation.
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16 45 INTRODUCTION

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19 47 Mining, which supplies materials thought essential for our society, has serious environmental
20 48 impacts. Opencast mining impacts all ecosystem components: substrata, topography, hydrology
21 49 (surface and groundwater), soil, vegetation, fauna, atmosphere, and landscapes (Osterkamp &
22 50 Morton, 1996; Evans, 2000; Rivas *et al.*, 2006). Often, mining impacts also have adverse effects
23 51 on nearby ecosystems. Among these off-site effects, the hydrologic impact of mines on
24 52 downstream fluvial ecosystems is one of the most detrimental (Toy & Hadley, 1987; Nicolau &
25 53 Asensio, 2000).
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31 55 Theoretically, mining reclamation should reduce these impacts. However, in spite of the
32 56 significant development of mining reclamation techniques over the years, failures on mining
33 57 reclamation are common (Haigh, 2000). Inadequate management of landform design at many
34 58 reclaimed mining sites has been identified as the main reason for reclamation failures because of
35 59 accelerated water erosion (Loch, 1997; Nicolau & Asensio, 2000).
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41 61 To achieve effective control of water erosion, an integrated management of topography, surface
42 62 soil cover, and vegetation is required (Nicolau, 2003). Of these three factors, the management of
43 63 topography and surface soil cover is considered an essential component of mining reclamation
44 64 practices by many (e.g., Evans & Willgoose, 2000; Toy & Black, 2001; Moliere *et al.*, 2002;
45 65 Toy & Chuse, 2005).
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51 67 For mine reclamation to be successful, efforts also must be directed towards the creation of
52 68 biologically functional and stable soils that reduce soil erosion and facilitate the rehabilitation of
53 69 post-mined lands (Bradshaw & Chadwick, 1980; Whisenant *et al.*, 1995). Soil erosion
54 70 negatively affects vegetation growth through several mechanisms: the removal of seeds and
55 71 nutrients from surface soil, direct plant removal, and the loss of water through surface runoff
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4 72 (Pimentel *et al.*, 1995; Espigares *et al.*, 2011). Indeed, seeds removal is sometimes a negligible
5 73 reason to explain the lack of vegetation even in bare surfaces (see Cerdá & García-Fayos, 1997).
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7 74 The most common soil surface used is topsoil (coversoil) spread on the slope surface; this
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9 75 approach is considered essential in most cases (Power *et al.*, 1981; Kapolka & Dollhopf, 2001).
10 76 Additionally, a wide range of modifications can be applied to improve physical and chemical
11 77 soil properties (Bradshaw & Chadwick, 1980). Armoring surface with rocks is a convenient and
12 78 cost-effective measure to decrease soil erodibility (Toy *et al.*, 2002).
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16 80 The most common approach of topography management consists of terraced landforms, graded
17 81 spoil banks comprising alternating short constant-gradient slopes and benches. Artificial ditches
18 82 commonly drain off the concentrated runoff (Bugosh, 2006). Without maintenance, many
19 83 terraced landforms succumb to water erosion in the long term (Loch, 1997). Linear slopes can
20 84 be unstable, especially if the base level is continuously changing by ditch incision, which causes
21 85 the slopes to respond by eroding or mass failure (e.g. Haigh, 1980, 1985). Erosion problems
22 86 also arise due to ponding or exceeding the storage capacity of the terraces (Sawatsky *et al.*,
23 87 2000). According to Hancock *et al.* (2003), linear slopes erode and increase sediment loss until
24 88 achieving a stable profile, which is usually concave. Additionally, we have reported how
25 89 terraced spoil heaps in this physiographic setting of the Upper Tagus are not stable within a
26 90 decadal span time, and they evolve to gullied landforms (see Sanz *et al.*, 2008).
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34 92 Arguments have frequently been raised in favour of topographic designs that replicate 'natural'
35 93 landscapes. This geomorphic approach is based on knowledge of geomorphic processes, mostly
36 94 fluvial processes operating for an extended period of time. The objective of these designs is the
37 95 construction of steady-state landscapes (Riley, 1995; Schor & Gray, 2007).
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42 97 Application of truly geomorphic approaches (Sawatsky & Beckstead, 1996; Toy & Chuse,
43 98 2005) depends very much on the exploitation method and timing. Implementing a geomorphic
44 99 approach is more difficult and expensive in active mines which already have terraced landforms.
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46 100 Often, only basic modifications of individual slopes (contour berm or contour linear steep slope)
47 101 can be cost-effective. Geomorphic approaches are easier to implement before mining activities
48 102 start or at abandoned mines. These two situations highlight the success of Bugosh's approach, a
49 103 computerized method (GeoFluv) of mining reclamation based on fluvial geomorphic principles
50 104 (Bugosh, 2004). His approach seeks hydrologic balance in reclaimed minescapes and is
51 105 perfectly tuned with the approach of Toy & Chuse (2005) who suggested that constructed
52 106 landscapes should include hydrologic basins, composed of slopes and watercourses. When basic
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4 107 modification of individual slopes is the only possibility, the GeoFluv method plays an important
5 108 role to decrease the slope length factor. This is carried out by building first and second order
6 109 channel drainage density, so that frequent small subwatersheds transform long slopes in shorter
7 110 ones, making the resultant landforms more resistant to erosion.
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11 112 The topographic profile of individual constructed slopes has been discussed for long in the field
12 113 of mining reclamation (Haigh, 1985; Toy *et al.*, 2002; Hancock *et al.*, 2003). Many studies have
13 114 reported a relationship between soil erosion and slope shape. These include the first studies in
14 115 geomorphology related to soil erosion on individual slopes (Meyer & Kramer, 1969), laboratory
15 116 experiments (D'Souza & Morgan, 1976), and the application of erosion models. For example,
16 117 Hancock *et al.* (2003) and Priyashanta *et al.* (2009) applied the SIBERIA model to demonstrate
17 118 the greater stability of concave slopes compared to linear ones. However, no field experimental
18 119 studies have been conducted to assess the reclamation benefits of concave slopes compared to
19 120 linear slopes.
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27 122 Because less sediment exportation occurs on concave slopes compared to other shapes (linear,
28 123 convex or S-shape) (Meyer & Kramer, 1969), these studies have led to the belief that concave
29 124 slopes are very stable. While watershed size and runoff increase downslope, the slope gradient
30 125 decreases, and this reduces runoff velocity and erosion ability (Toy *et al.*, 2002).
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35 127 Martín-Duque *et al.* (2010) explained how a holistic geomorphic approach to mining
36 128 reclamation, using both topographic and surface soil cover management, led to a successful
37 129 mining reclamation in a quarry of Central Spain. The current study is based on that work and
38 130 describes a field experiment carried out at the El Machorro kaolin mine of Central Spain. The
39 131 objective of this study was to compare the erosion response of two constructed slopes, linear
40 132 and concave, with two different surface soil covers. These soil covers were: i) subsoil
41 133 (carbonate colluvium), a natural superficial sediment that drapes the sandy sedimentary rocks
42 134 underlying the original slopes around the mine, and ii) topsoil, soils developed originally on top
43 135 of the carbonate colluvium. A linear slope of overburden material with no cover was used as a
44 136 control for linear slopes. A concave slope of overburden material with no cover could not be
45 137 constructed, because the experimental layout had to be adapted to pre-existing topographic
46 138 conditions. Therefore, a total of four different combinations of topography and surface soil
47 139 cover, that we call 'reclamation treatments', and one control (overburden linear slope), were
48 140 monitored in this study. A core objective of this study was to compare the response of both
49 141 topographies and both surface soil covers, to acquire know-how for efficient mining reclamation
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4 142 at similar sites. Our working hypothesis was that concave slopes would yield less sediment than
5 143 linear slopes. We also expected a dramatic reduction in soil loss from topsoil and carbonate
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7 144 colluvium compared to overburden material.
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10 146 MATERIALS AND METHODS

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12 148 *Study area*

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16 150 El Machorro is an active contour mine with an ongoing terraced reclamation approach. It is
17 151 located in the buffer zone of the Upper Tagus Natural Park (UTNP, *Parque Natural del Alto*
18 152 *Tajo*, in Spanish) in Central Spain (40° 39' 29" N, 2° 2' 26" W, datum World Geodetic System
19 153 1984, WGS84) (Figure 1). This protected area was established in 2000 by a regional law
20 154 (DOCM, 2000) because of its outstanding biodiversity, specifically regarding aquatic
21 155 ecosystems. It is also very diverse geologically (Carcavilla *et al.*, 2008) and biologically
22 156 (DOCM, 2000).
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28 158 The Upper Tagus landscape is characterized by plateaus and mesas capped by Cretaceous
29 159 carbonates, with their slopes and canyon scarps underlain by sandy sediment that hold the
30 160 kaolin (*Arenas de Utrillas* Formation) exploited in several mines (Olmo & Álvaro, 1989;
31 161 González Amuchastegui, 1993).
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36 163 On mesa tops, the soils are chromic luvisols, calcareous cambisols, mollic leptosols, and rendzic
37 164 leptosol. On slopes, carbonate colluvia with calcareous cambisols are common (IUSS Working
38 165 Group WRB, 2007). The vegetation is representative of mediterranean-continental
39 166 environments, with communities dominated by *Juniperus thurifera* on the highest plateaus, and
40 167 pine (*Pinus nigra* subsp. *salzmanii*) and gall oak (*Quercus faginea*) in valleys (MARM, 1997–
41 168 2006).
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47 170 The climate of this area is temperate mediterranean with dry and mild summers (Csb, according
48 171 to Köppen, 1918), but with a noticeable continental influence. The moisture regime is dry
49 172 mediterranean (Papadakis classification) (CNIG, 2004). Mean annual precipitation is 780 mm
50 173 and mean annual temperature is 10°C (AEMET, 2012). Seasonally, this area is characterized by
51 174 long and cold winters with snow common and short, dry summers with high intensity
52 175 rainstorms. The spring and fall are usually wet. The rainfall erosive factor, R (equivalent to the
53 176 R factor of RUSLE), is estimated to be about 80 (ICONA, 1988).
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5 178 *Rainfall and temperature monitoring*

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9 180 To measure rainfall quantity and intensity, a tipping-bucket automatic raingauge (0.2 mm/pulse)
10 181 (Davis Instruments, 2005) with a HOBO Event data logger was installed 100 m away from the
11 182 experimental spoil heap, at 1230 m.a.s.l. Raingauge data were downloaded at the same time as
12 183 the sediment collection. Total rainfall volume (mm) and maximum rainfall volume in 24 hours
13 184 (mm) were calculated. In addition, the return period of annual precipitation for each year was
14 185 estimated using the CHAC software (CEDEX, 2004). Each year, temperature data were
15 186 obtained from a nearby weather station (AEMET, 2012).

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21 188 *Experimental design*

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24 190 An experimental spoil heap was built by the mining operator company of El Machorro mine,
25 191 CAOBAR, in the summer of 2008, on the foundations of an existing spoil heap. Two different
26 192 topographic slope shapes, linear and concave, were constructed with spoils (overburden
27 193 materials) and covered with two surface soil covers: subsoil (carbonate colluvium), and topsoil
28 194 (Table I). Additionally to these four reclamation treatments, one linear slope of the spoil heap
29 195 with overburden material (spoils) was left uncovered as a control (Figure 2). The four
30 196 'reclamation treatments' and the control were monitored for two hydrologic years (2009 and
31 197 2010) starting from November 6, 2008.

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37 199 At the experimental spoil heap, articulated dump trucks built the terraced spoil heaps by directly
38 200 unloading materials, and a bulldozer compacted and finished the benches. The dump trucks
39 201 could not drive on the linear slopes because of their high slope gradient, so the trucks drove on
40 202 the benches and unloaded the two surface soil covers directly downslope. The concave slope
41 203 was built by a bulldozer that drove on the concave slope reshaping it and spreading the surface
42 204 soil covers at the same time. Summing up, the experimental spoil heap had two parts. The first
43 205 one was a terraced system with two linear slopes and one intermediate bench. Each linear slope
44 206 had the two surface soil covers (carbonate colluvium and topsoil) and the exposed overburden
45 207 material (control); the second part was a concave slope with the two surface soil covers,
46 208 therefore five different slopes were monitored (see Figure 2).

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54 210 Mining and reclamation operations within the mine prevented the construction of the upper part
55 211 of the concave slope during the first hydrologic year of the study. During this period, the

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4 212 concave slope consisted of its half-lower part, equivalent in height to a single linear slope plus
5 213 its bench. Additionally, run-off from the upper slope formed an alluvial fan on the concave
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7 214 slope covered with carbonate colluvium. Therefore, data could not be collected on this treatment
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9 215 during the first year. The concave slope was fully constructed in the second year to have the
10 216 same width and length as a set of two linear slopes with an intermediate bench. This
11 217 modification could be considered a limitation of this study.
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14 219 Linear slopes had a mean length of 11 m (standard deviation 0.6), with a slope gradient of 32°.
15 220 The bench was 5 m wide with a reversed-slope gradient of 14° in cross section and 2° in
16 221 longitudinal section. Concave slopes had a slope length of 25 to 30 m during the first year and
17 222 35 to 40 m during the second year. Their gradient increased from bottom to top from 4° to 26°
18 223 (first year) and from 4° to 32° (second year) (See table II for details). The concave slope
19 224 curvature was described using the equation proposed by Stefano *et al.* (2000):
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$$y = H \left(1 - \frac{x}{\lambda} \right)^n$$

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28 227 where x = horizontal abscissa and y = the corresponding elevation

29 228 H= difference of level

30 229 λ = slope length measured along the horizontal axis

31 230 n= exponent that varies according slope shape, following Stefano *et al.* (2000)

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33 232 Short concave slopes (first year) had an n value between 1.34 and 1.32, whereas long concave
34 233 slope values (second year) were between 1.40 and 1.47 (Figure 3). A differential Global
35 234 Positioning System (GPS, model number Leica 1200) was used to survey the concave slope
36 235 profiles. Slope surveys were conducted once a year (12 May 2009 and 1 July 2010).
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45 237 Three composite samples were taken from each soil cover to characterize their physical
46 238 properties (shown in Table III). The thickness of both carbonate colluvium and topsoil ranged
47 239 between 30 and 75 cm on linear slopes. This wide range resulted from directly unloading
48 240 material from upslope without spreading it. Carbonate colluvium and topsoil on concave slopes
49 241 were 20-30 cm thick, and were spread by a bulldozer.
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52 242
53 243 The core of this study is based on the field measurement of the sediment amount yielded by
54 244 each reclamation treatment and the control. Three open plots were set up for every slope.
55 245 Sediment amount was recorded using silt fences (Robichaud & Brown, 2002), with a width of 3
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4 246 m, placed across the toe of the slopes. Silt fences trap sediment while allowing water to pass
5 247 through. According to Robichaud & Brown (2002), the trap efficiency of silt fences is 68 to
6 248 98%. Because sediment could fill and overload silt fences, possibly resulting in a loss of
7 249 sediment, periodic cleaning of silt fences was necessary (Robichaud & Brown, 2002).

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11 251 Sediment yield was measured at the toe of the concave slope and at the toe of the lower single
12 252 linear slope of the set of two linear slopes (Figure 2). Sediment from the upper linear slope were
13 253 not measured, but they did not run onto the monitored lower linear slope, as they were deposited
14 254 on the intermediate reversed sloped bench and drained out of the monitored lower linear slope
15 255 (Figure 2). The short reversed slope of the terrace bench was not counted in the balance, as it
16 256 was observed that it did not yield any sediment.

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21 258 Therefore, a total of 12 (first year) and 15 (second year) sets of 'open' plots (plots without
22 259 artificial boundaries) with silt fences were monitored. Since the plots were open, there were
23 260 differences in plot size due to different drainage areas. The area of each open plot, measured
24 261 using differential GPS, ranged between 23.5 and 83.7 m² (first year) and between 23.5 and 124
25 262 m² (second year) (Table II).

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31 264 *Sediment yield*

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33 266 The protocol for monitoring the open plots consisted of collecting the sediment trapped by the
34 267 silt fences and weighing the sediment in the field, using a portable weight scale. The sediment
35 268 from a single plot was mixed and a portion of the mixed sediment was taken to calculate the
36 269 percentage of moisture, using the method by Ramos-Scharrón & McDonald (2007). The erosion
37 270 rate was calculated and the results were expressed as Mg ha⁻¹yr⁻¹. Annual sediment yields and
38 271 standard deviations were also calculated for each treatment.

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44 273 *Rill development*

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46 275 Overburden materials at El Machorro mine are mainly sandy, with very low clay content. The
47 276 very low cohesion makes the overburden material vulnerable to detachment by runoff, so that
48 277 gully formation is common.

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52 279 To monitor the landform evolution of the four different reclamation treatments and the control,
53 280 photographs were taken of each open plot before sediment was collected. Rill networks were

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4 281 measured after they formed. Width and depth were measured in at least 80% of all rills in three
5 282 slope positions (top, middle, and bottom).
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8 284 The length, width, and depth of rills and gullies were measured with a tape, following the
9 285 method described by Morgan (Morgan, 2005). Rill volume was estimated by multiplying the rill
10 286 cross-sectional area —“U” shape, for carbonate colluvium and “V” shape, for overburden
11 287 material— by their mean rill length. This rill volume was then divided by the treatment area, to
12 288 obtain estimated values for sediment removed by rill erosion ($\text{m}^3 \text{m}^{-2}$). This value was then
13 289 transformed to sediment weight per area (Mg ha^{-1}) by multiplying the volume by the mean bulk
14 290 density of each surface soil cover that was calculated by the core method (Sobek *et al.*, 1978).
15 291 Three soil core samples were taken from each slope treatment for bulk density calculations. The
16 292 sediment amounts resulting from rill measurements and from the silt fences were then
17 293 compared.
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19 295 *Vegetation colonization*
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26 297 Vegetation cover was measured using digital photographs and a point-frequency method
27 298 (Brakenhielm & Liu, 1995; Vanha-Majamaa *et al.*, 2000) one year after the end of the second
28 299 year of the study (October 2011). Because no seeding was applied in any of the reclamation
29 300 treatments, we therefore measured spontaneous vegetation colonization.
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36 302 *Statistical analysis*
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40 304 To compare the effects of topography and surface soil cover on sediment yield, paired t-tests
41 305 were conducted comparing sediment yield from treatments with the same topography but with
42 306 different surface soil cover (i.e. linear slope with carbonate colluvium vs linear slope with
43 307 topsoil) and sediment yield from treatments with same surface soil cover but with different
44 308 topography (i.e. concave slope with topsoil vs linear slope with topsoil). Analyses were
45 309 conducted separately for each study year. For linear slopes, data were also analyzed for both
46 310 years combined, because the plots were not modified during the second year. Statistical analyses
47 311 were made using Statgraphics Centurion XVI.I software, version 16.1.17 (StatPoint
48 312 Technologies Inc., 2012). The significance level was $\alpha=0.05$.
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RESULTS

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4 316 *Rainfall and temperature*
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7 318 A total of 324 rain days were registered during the study period, accounting for a total rainfall of
8 319 1426 mm. Annual rainfall for the second year (992 mm) was approximately twice that of the
9 320 first year (434 mm), with return periods of 5 and <2 years, respectively. Climatic characteristics
10 321 of each study year are shown in Table IV. Monthly rainfall ranged from 1 mm (July 2009) to
11 322 290 mm (December 2009). The maximum rainfall recorded in 24 hours was 49 mm. Frost-free
12 323 days were slightly more common in the second year.
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18 325 *Sediment yield*
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21 327 During the two years studied, open plots were sampled approximately once a month, resulting
22 328 in a total of 21 samples: 10 samples during the first year and 11 samples during the second year.
23 329 Mean sediment yield and standard deviation of each reclamation treatment are shown in Table
24 330 V, along with rainfall characteristics for the period between two consecutive sediment
25 331 collections. The sediment yield rates for the three plots within the same treatment did not differ
26 332 significantly ($p>0.05$, paired t-test).
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32 334 Significant differences were found when sediment yield rates from reclamation treatments with
33 335 the same topography but different surface cover were compared (Table VI). For the first year,
34 336 the comparison between linear slope with topsoil (LS-TS) and linear slope with overburden
35 337 material (LS-OM) showed a significant difference ($p=0.01$, t-test). For the second year, the
36 338 comparison of these two treatments also showed a significant difference ($p=0.003$). Regarding
37 339 the two-year data analyses, significant differences were found between all tested pairwise
38 340 treatments on linear slopes ($p<0.05$, paired t-test). When slopes with the same surface cover but
39 341 different topography were compared, no meaningful significant differences were found.
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45 343 Regarding annual sediment yield rates, the short concave slope with topsoil (SCS-TS) had lower
46 344 sediment yield values than any linear slope during the first year, regardless of surface soil cover
47 345 (Figure 4 and Table V). The sediment yield rates of linear slopes depended on the surface soil
48 346 cover: the slope with topsoil had the lowest rate ($12 \text{ Mg ha}^{-1}\text{yr}^{-1}$), one order of magnitude less
49 347 than that with carbonate colluvium ($120 \text{ Mg ha}^{-1}\text{yr}^{-1}$) or overburden material ($282 \text{ Mg ha}^{-1}\text{yr}^{-1}$).
50 348 In the second year, the linear slope with topsoil (LS-TS) produced the lowest erosion rate (3 Mg
51 349 $\text{ha}^{-1}\text{yr}^{-1}$). The other two linear slopes had the higher values: $126 \text{ Mg ha}^{-1}\text{yr}^{-1}$ with carbonate
52 350 colluvium and $347 \text{ Mg ha}^{-1}\text{yr}^{-1}$ with just overburden. The effect of surface soil cover was not
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4 351 found for the long concave slopes. The slope with topsoil (LCS-TS) yielded 20 Mg ha⁻¹yr⁻¹ of
5 352 sediment and the slope with carbonate colluvium (LCS-CC) yielded 16 Mg ha⁻¹yr⁻¹ (Figure 4).
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8 354 *Rill development*
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11 356 Rill development was different on concave and linear slopes. Concave slopes developed a rill
12 network in the upper part, lacking rills in its lower part. Linear topography allowed a continuous
13 357 rill network along the slope length. In both cases, rill development depends on the surface soil
14 358 rill network along the slope length. In both cases, rill development depends on the surface soil
15 359 cover characteristics.
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18 361 Rill development on concave slopes
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21 363 The concave slope covered with topsoil (SCS-TS) did not develop rills during the first year,
22 364 which was dryer than the second one. Indeed, this treatment resisted the most intense rainfall in
23 365 24 hours of the first year (38.4 mm), which occurred just after building the experimental spoil
24 366 heap and spreading the surface soil cover, but before the silt fences were installed. During the
25 367 second year, small rills formed in the steepest area of the concavity, near the top of the slope,
26 368 but they were small and disappeared downslope. These rills were not measured, because we
27 369 assumed the sediment eroded from them was deposited within the slope.
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30 371 Plots on the concave slope with carbonate colluvium surface soil cover (SCS-CC) could not be
31 372 monitored during the first year, because run-on from upslope formed noticeable alluvial cones
32 373 within the open plots. In the second year, the upper parts of both concave slopes were
33 374 reconstructed, making them longer. During the second year, the concave slopes behaved
34 375 similarly, regardless of their surface soil cover: rills were formed at the top of the slope and
35 376 disappeared downslope. On the long concave slope with carbonate colluvium, these rills were
36 377 discontinuous, with a “U” shape, and mean length of 6 m. The estimated sediment volume
37 378 eroded from these rills over the two-year period was 1.4 m³, or 0.004 m³ m⁻², based an area of
38 379 330 m² on the LCS-CC. No mass movements, such as mudflows, occurred on the concave slope
39 380 with carbonate colluvium. The calculated bulk density for carbonate colluvium was 1.26 g cm⁻³,
40 381 so the estimated weight of sediment from the concave slope with carbonate colluvium was 50
41 382 Mg ha⁻¹. Since 80% of rills were measured, the estimated total mass of sediment was 63 Mg ha⁻¹.
42 383 Two-year sediment yield measured in the open plots of this same slope was 16 Mg ha⁻¹. The
43 384 estimated amount of sediment determined from rill development has the same order of
44 385 magnitude as that measured at the silt fences, for the two-year period (Figure 6).
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5 387 Rill development on linear slopes

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8 389 The linear slope with topsoil (LS-TS) did not develop perceptible erosive forms during the two
9 390 years. The linear slope covered with carbonate colluvium (LS-CC) was subject to small
10 391 mudflows in the first year. Additionally, an incipient rill network developed. After this initial
11 392 geomorphic evolution, the plots remained very stable throughout the two-year period, with only
12 393 small mudflows and minor rills. At the end of the second year, rills were discontinuous, with a
13 394 “U” shape, with an average width of 30 to 40 and depth of 10. The estimated average length was
14 395 7 m, and the estimated sediment volume eroded from rills was 0.4 m^3 . The estimated sediment
15 396 removed by rill erosion was $0.004 \text{ m}^3 \text{ m}^{-2}$. Considering the corresponding bulk density (1.27 g
16 397 cm^{-3}), the estimated sediment yield was 51 Mg ha^{-1} (from 80% of rills), corresponding to a total
17 398 sediment of 64 Mg ha^{-1} (for 100%). This estimated sediment yield is one order of magnitude
18 399 lower than that measured at the silt fences (246 Mg ha^{-1} for the two-year period) (Figure 6).

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22 401 The linear slope covered with overburden material (LS-OM) developed an evenly defined rill
23 402 network. These rills were deeper and much more numerous than those formed on the carbonate
24 403 colluvium. The rills were 20 cm-wide on average, and had an average depth of 20 to 30 cm,
25 404 maximum 50 cm, at the end of the first year (Figure 5). Small alluvial cones were formed at the
26 405 bottom of the slopes. A progressive disintegration of sand clods on the linear slope surface was
27 406 also observed during the two years. During the second year, the rill-erosion process continued,
28 407 leading to the formation of gullies, being these landforms defined in the same way that Brice
29 408 (1966, p. 290): “a recently extended drainage channel that transmits ephemeral flow, has steep
30 409 sides, a steeply sloping or vertical head scarp, a width greater than about 1 foot, and a depth
31 410 greater than about 2 feet”. At the end of the second year, the rills were continuous, “V”-shaped,
32 411 with an average width and depth of 45 cm and 25 cm, respectively. Gullies with a maximum
33 412 width of 200 cm and depth of 150 cm were also measured. Rill length was the same as on the
34 413 linear slope, 11 m. The estimated sediment volume eroded from rills was 4.75 m^3 , and 0.045 m^3
35 414 m^{-2} , the highest of the slopes monitored (Figure 6). The estimated sediment eroded by rill
36 415 processes, calculated using the bulk density of 1.41 g cm^{-3} , was 793 Mg ha^{-1} (considering 100%
37 416 of rills). The estimated sediment yield quantified from rill development was higher than that
38 417 measured at the silt fences (629 Mg ha^{-1} for the two-year period).

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4 421 At the start of the study period all plots were bare, without any vegetation. As geomorphic
5 422 evolution progressed, natural plant colonization occurred. Concave and linear slopes covered
6 423 with topsoil showed plant establishment in the following spring (spring of 2009). In October
7 424 2011, plants covered 30% of the concave slope and 50% of the linear slope (table III). Plants
8 425 spatial pattern was not homogeneous on the concave slope with topsoil, so that plants were not
9 426 evenly distributed along the slope, but the linear slope showed a uniform vegetation distribution.
10 427 On the concave slope, vegetation cover was more extensive in the lower part of the slope than at
11 428 the top. Table VII shows the plant species identified in each topsoil-covered slope. Although
12 429 species richness is similar in both slopes (14), species composition is quiet different (being only
13 430 5 species common among to the two slopes). No vegetation was observed on carbonate
14 431 colluvium or overburden material.
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433 DISCUSSION

434 *Sediment yield*

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27 436 Our results suggest that surface soil cover controls sediment yield on linear slopes more than on
28 437 concave ones. This is supported by the fact that linear topography has no mechanisms to control
29 438 sediment fluxes, while concave topography is able to store sediment at the toe (Stefano *et al.*,
30 439 2000; Toy *et al.*, 2002). On linear slopes, control of erosion could be improved by using a
31 440 different surface soil cover. Our results are consistent with previous findings: topsoil was the
32 441 best surface soil cover, providing better conditions for soil development and plant establishment
33 442 than other materials (Power *et al.*, 1981; Haigh, 2000).
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39 444 Similar erosive response was observed in the first year for the topsoiled slopes, whether short
40 445 concave (SCS-TS) or linear (LS-TS), indicating that, under favorable soil conditions, the role of
41 446 topography was not evident. During the second year, topsoiled slopes behaved differently.
42 447 While sediment yield from the linear slope with topsoil (LS-TS) was reduced, sediment yield
43 448 from the long concave slope (LCS-TS) was greater than the yield from the short concave slope
44 449 (SCS-TS). The increased length and drainage area could explain the increase in sediment yield.
45 450 In agreement with this, several authors have reported that, under the same environmental
46 451 conditions, shorter slopes produce less sediment than longer ones (Toy & Foster, 1998; Liu *et*
47 452 *al.*, 2000; Toy *et al.*, 2002; Toy & Chuse, 2005).
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54 454 Another aspect must be considered: constraints existed for combining soil surface covers and
55 455 topography. The depth, uniformity, and quality of surface soil cover were determined by
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4 456 reclamation operations. On linear slopes, the surface soil cover was spread out by direct
5 457 unloading of trucks, which provided a more homogeneous and less compacted layer. However,
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7 458 on concave slopes the spreading out process had to be carried out with a bulldozer, which
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9 459 compacted the soil (Barber & Romero, 1994; Chong & Cowser, 1997). Soil compaction has
10 460 been reported to reduce the land's capacity to absorb rainwater, accelerating runoff and erosion
11 461 (Haigh & Sansom, 1999). The greater thickness and porosity of linear slopes with topsoil, as
12 462 well as a better spatial distribution of surface soil cover, could explain lower rates of sediment
13 463 yield than for the concave slope. This means that slope topography affects surface soil cover
14 464 depth and quality in reclaimed landscapes (Hancock *et al.*, 2003; Priyashanta *et al.*, 2009) (see
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16 465 table VIII).
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21 467 The smaller second-year sediment yield from the long concave slope ($16 \text{ Mg ha}^{-1}\text{yr}^{-1}$) compared
22 468 with linear slope with carbonate colluvium ($126 \text{ Mg ha}^{-1}\text{yr}^{-1}$) suggests that concave topography
23 469 helps to reduce sediment yield. The yield was smaller even though the concave slope was longer
24 470 than the corresponding linear slope, and even though the concave slopes had been recently
25 471 constructed.
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30 473 To assess the validity of the sediment yield measurements, it is important to take into account
31 474 that, although the plots were open, the length and area of the linear slopes were similar. Because
32 475 of this, we consider that converting sediment yield to per unit area, and comparing them, was
33 476 justified. However, the long concave slopes had larger open plots. A larger contributing area
34 477 implies a higher erosive power, but, the fact that the slope was concave implies a lower erosive
35 478 power. The combined consequence of these effects could not be separated and quantified.
36 479 Therefore, converting sediment yield to per unit area for concave slopes, and comparing them
37 480 with linear slopes, has an evident uncertainty. Despite of that, the comparison was made
38 481 because they are real alternatives of reclamation, both for this site and elsewhere: concave
39 482 slopes or terraced ones as a topographic possibility of regarding spoil heaps.
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46 484 *Rill development*
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50 486 In our experiment, rill development on linear slopes showed clear differences depending on the
51 487 surface soil cover. Whereas no rills were formed on the linear slope with topsoil (LS-TS), a
52 488 widespread rill network was developed on overburden material (LS-OM), and only few rills and
53 489 mudflows occurred on carbonate colluvium (LS-CC). This very different geomorphic behavior
54 490 indicates that soil cover is dominant in controlling erosion processes on linear slopes. Topsoil
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4 491 resists erosion (Sawastky *et al.*, 1996), because its higher infiltration rate decreases runoff and,
5 492 therefore, soil detachment (Haigh & Samson, 1999). On the other hand, rill erosion is very
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7 493 common in overburden materials, because higher bulk density promotes overland flow
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9 494 (Soulliere & Toy, 1986; Moreno-de las Heras *et al.*, 2010). Two additional factors favored rill
10 495 formation in overburden material: the low rock cover and the sandy texture (Quansah, 1981;
11 496 Porta *et al.*, 1989) as described in table III.
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13 497
14 498 Generally speaking, rills grow by incision and by side-wall sliding (Nicolau, 2002). The
15 499 different cross sections —V vs U shape— and size could be explained as a consequence of
16 500 different surface soil covers. Rills developed on overburden material were V-shaped and larger
17 501 than those on carbonate colluvium. This was likely due to the sandy texture and lower cohesion
18 502 of overburden, favoring more effective incision and side-wall collapse, and causing rill
19 503 widening. Rills developed on carbonate colluvium were observed to be U-shaped and smaller.
20 504 This could be interpreted as a result of higher cohesion in carbonate colluvium because of lower
21 505 sand and higher silt content than in overburden material. The carbonate colluvium also has a
22 506 higher surface roughness (due to the abundance of rock fragments), which would also contribute
23 507 to a smaller rill size development. Roughness decreases overland flow and runoff because of
24 508 surface ponding and increased hydraulic roughness that reduces the effective flow shear stress
25 509 (Darboux *et al.*, 2002; Toy *et al.*, 2002; Gómez & Nearing, 2005).
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34 511 Sediment yield estimated to have been eroded from rills differed from sediment yield measured
35 512 in silt fences. At least two factors affect the interpretation of the results. Sediment yield
36 513 estimated from rills assessment represented only rill erosion. For all comparisons it is important
37 514 consider that rill assessment has some limitations, and it is an estimation. At the same time, silt
38 515 fences trap sediment from rill, inter-rill erosion and mudflows, and it is necessary to take into
39 516 account how efficiently the silt fences trap sediment. According to Robichaud & Brown (2002)
40 517 the total values for sediment yield could be 2% to 32% higher. One might expect then that rill
41 518 erosion estimates were probably low and silt fence measurements could be higher.
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49 520 For the linear slope with carbonate colluvium (LS-CC), sediment yield estimated from rills
50 521 assessment was one order of magnitude lower than sediment yield measured at silt fences (64
51 522 Mg ha⁻¹ and 246 Mg ha⁻¹ respectively, figure 6). This difference could be explained by the fact
52 523 that small mudflows occurred on this slope. For the linear slope with overburden material (LS-
53 524 OM), the estimated sediment yield from rills was 164 Mg ha⁻¹ (21 %) higher than the sediment
54 525 yield measured in silt fences. This could be explained by the fact that small alluvial cones were
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4 526 formed at the bottom of the slope and also because sediment overloaded the silt fences on some
5 527 occasions. For the concave slope with carbonate colluvium (LCS-CC) the difference between
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7 528 the two values was 47 Mg ha⁻¹, being 75 % higher the sediment yield estimated from rills. This
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9 529 was likely due to some sediment that was deposited downslope and did not fill the silt fences.

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11 531 *Vegetation colonization*

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14 533 In our study, the plant establishment pattern was quite different on the linear vs the concave
15 534 slope (always regarding topsoiled treatments).

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19 536 The linear profiles allowed more widespread and homogeneous plant cover. This could be
20 537 because their abiotic characteristics: slope angle and surface soil cover depth and compaction
21 538 which were very homogeneous, so that its environmental heterogeneity is nor remarkable. In
22 539 fact, species associated to worse soil conditions —i.e. *Thymus vulgaris*, *Brachypodium*
23 540 *phoenicoides*, or *Aphyllanthes monspeliensis*— appear only in the linear slope.

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25 542 The concave profile includes two very different environments (upper steepest part and lower
26 543 flatter part). Plant colonization occurred mainly in the lower and flatter one, where water
27 544 availability as well as the seed bank richness should be higher. In addition, woody species have
28 545 been identified here (*Genista scorpius* and *Sideritis hirsute*).

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30 547 These facts are interpreted as the development of a more ‘structured’ plant community in the
31 548 concave slope than in the linear one. In turn, we consider this as a result of a more
32 549 heterogeneous environment on the concave slope. Of course, given the very few years of
33 550 vegetation colonization, these are preliminary results, and a larger time-span is needed for more
34 551 conclusive results, as far as the vegetation development is concerned.

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36 553 The greater amount of continuous vegetation cover on the linear slope could be another
37 554 explanation for the lower sediment yield rates for linear vs concave slopes. In this respect, the
38 555 value of 50 % of vegetation cover reached by this linear slope with topsoil and the decrease of
39 556 sediment yield amount seems to be in agreement with the literature. Indeed, the role of
40 557 vegetation cover in sediment yield control is well known. Several authors have observed that, in
41 558 mediterranean environments, erosion rates are greatly reduced when vegetation cover rises up
42 559 above 30% (Thornes, 2004; de Luis *et al.*, 2001; Gimeno-García *et al.*, 2007). Andres & Jorba
43 560 (2000) and Moreno-de las Heras *et al.* (2009) confirmed empirically the drastic reduction of soil

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4 561 loss with a 30% plant cover for slopes constructed for mining reclamation in central and
5 562 northeast Spain. They recommend a 50% plant cover in practice as a conservative target. For
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7 563 man-made slopes there is considerable evidence that the restoration of 50% cover with
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9 564 herbaceous vegetation is decisive for site stabilization. And this is what our experiment seems to
10 565 show. The literature reflects, however, that it is not only a question of cover, but also a matter of
11 566 how the vegetation cover is distributed, such as in natural ecosystems (Cerdá *et al.*, 2010).
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14 568 CONCLUSIONS

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17 570 These conclusions are addressed for mining scenarios similar to the one described, active mines
18 571 which already have terraced landforms, with possibility of being improved either by limited
19 572 topographic modifications (concave slopes) or by different use of surface soil covers. However,
20 573 the long term instability of terraced spoil heaps has been proved, with special emphasis in arid
21 574 and semi-arid climates, as the mediterranean one (see Introduction for references). Therefore,
22 575 wherever mining reclamation is less conditioned by previous mining works, we recommend a
23 576 mining reclamation based in a geomorphic approach, instead of in terraced slopes.
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26 578 The effect of topography (linear or concave) on soil erosion was prominent when slopes were
27 579 covered by carbonate colluvium. Without topsoil, concave slopes yielded much less sediment
28 580 than linear slopes, with deposition occurring primarily at the flatter bottom part of the slope,
29 581 reducing off-site sediment exportation. Therefore, building concave topographies could be
30 582 considered advisable when no topsoil is available.
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33 584 The interaction between vegetation establishment and topography is complex. Natural plant
34 585 cover was more widespread and more homogeneous on linear slopes than on concave ones. In
35 586 the latter, natural plant colonization on the steepest part of the concavity was severely limited.
36 587 The bottom of the concavity provided more favorable conditions for plant growth.
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39 589 The three main activities involved in mining reclamation (slope construction, use of surface soil
40 590 cover, and plant establishment) did not operate independently in reducing sediment yield and
41 591 erosion. This study suggests that the debate about the management of topography and surface
42 592 soil cover, and their relationship with vegetation, should be approached under a systemic
43 593 perspective. The main trade-offs between major variables should be considered: i) topsoil can
44 594 lead to a successful mining reclamation regardless of the two types of topography considered in
45 595 our experiment; ii) managing topography by creating concave slopes can lead to a successful
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4 596 mining reclamation when the use of topsoil is limited; and iii) topsoil and topography determine
5 597 the plant colonization pattern.
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30 805 FIGURE CAPTIONS

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34 807 Figure 1 Location of the study area within the Iberian Peninsula and within the province
35 808 of Guadalajara. The experimental spoil heap is located at El Machorro mine.
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39 810 Figure 2 Experimental spoil heap of El Machorro mine, during the second study year,
40 811 after conversion of the short concave slopes to long concave slopes. Top, treatment
41 812 scheme; bottom, photograph taken October 2011, one year after experiment finished.
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43 813 LCS-TS = long concave slope with topsoil, LCS-CC = long concave slope with
44 814 carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-TS = linear
45 815 slope with topsoil, LS-OM = linear slope with overburden material. The long concave
46 816 slope with overburden material (LCS-OM) could not be constructed.
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50 818 Figure 3 Concave slope shapes and their n values. The n value is an exponent that varies
51 819 according to slope shape, following the equation of Stefano *et al.* (2000). The original,
52 820 short concave slopes were converted to long concave slopes at the end of the first year.
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5 822 Figure 4 Mean annual sediment yield (Mg ha^{-1}) by treatment and study year. The error
6 bars represent the standard deviation. SCS-TS = short concave slope with topsoil,
7 LCS-TS = long concave slope with topsoil, LCS-CC = long concave slope with
8 carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-TS = linear
9 slope with topsoil, LS-OM = linear slope with overburden material. The short
10 concave slope with carbonate colluvium (SCS-CC) was not monitored during the
11 first year.
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19 830 Figure 5 Photographs showing geomorphic evolution and vegetation colonization at the
20 experimental spoil heap (see text for explanation).
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24 833 Figure 6 Comparison of sediment yield measured from silt fences with sediment yield
25 estimated from rill erosion, for the two year study period. LCS-CC = long concave
26 slope with carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-
27 OM = linear slope with overburden material.
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31 837 TABLE CAPTIONS
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34 839 Table I Slope code and starting month and year of measurements for each treatment.
35 Measurements did not start in October 2008 because the spoil heap was built that month
36 840
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38 842 Table II Experimental treatments and their characteristics
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40 844 Table III Surface soil cover characteristics and vegetation cover. Values are means.
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43 845 Vegetation survey was carried out in May 2010
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46 847 Table IV Climate characteristics of each study year
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49 849 Table V Rainfall characteristics and sediment yield on sampling dates. Total values are
50 also included by hydrologic year. S/LCS-TS = short/long concave slope with topsoil,
51 LCS-CC = long concave slope with carbonate colluvium, LS-CC = linear slope with
52 carbonate colluvium, LS-TS = linear slope with topsoil, LS-OM = linear slope with
53 overburden material
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855 Table VI Results of paired t-test. Statistical significance level: $*\alpha=0.05$. S/LCS-TS =
 856 short/long concave slope with topsoil, LCS-CC = long concave slope with carbonate
 857 colluvium, LS-CC = linear slope with carbonate colluvium, LS-TS = linear slope with
 858 topsoil, LS-OM = linear slope with overburden material

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860 Table VII Plant species established in the slopes with topsoil.

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862 Table VIII Concave and linear profile characteristics related to sediment yield, rill
 863 development, and establishment of vegetation

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TABLES

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Table I

Code	Treatment		Month	Calendar year
	Topographic profile	Surface soil cover		
SCS-TS	Short concave slope	topsoil	November	2008
LCS-TS	Long concave slope	topsoil	October	2009
LCS-CC	Long concave slope	carbonate colluvium	October	2009
LS-TS	Linear slope	topsoil	November	2008
LS-CC	Linear slope	carbonate colluvium	November	2008
			October	2009
LS-OM	Linear slope	overburden material	November	2008
			October	2009

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Table II

Treatment	Open plot number	Topographic profile	Surface soil cover	Surface soil cover thickness (cm)	Slope length (m)		Slope gradient (°)		Area (m ²)	
					Yr 1	Yr 2	Yr 1	Yr 2	Yr 1	Yr 2
SCS-TS and LCS-TS	1								83.7	91
	2		topsoil						82.9	104
	3	concave		20 to 30	33	40	4 to 26	4 to 32	73.2	100
SCS-CC and LCS-CC	4								58.7	106
	5		carbonate colluvium						70.3	124
	6								61.5	100
LS-TS	7								30.9	
	8		topsoil						35.5	
	9								45.7	
	10								27.5	
LS-CC	11	linear	carbonate colluvium	30 to 75	11		32		23.5	
	12								43.2	
	13								31.3	
LS-OM	14		overburden						43.5	
	15								31.3	

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885 **Table III**

Treatment	Sand (%) 2–0.05 mm	Silt (%) 0.05 – 0.002 mm	Clay (%) <0.002 mm	Organic matter (%)	Bulk density (g cm ⁻³)	Textural classification (USDA)	Soil structure	Rock cover (%)	Vegetation cover (%)
LCS-TS	49.7	29.8	20.5	2.3	1.06	sandy clay loam	medium or coarse granular 2–5mm	20	30
LCS-CC	39.8	47.2	13.1	0.6	1.26	loam	medium or coarse granular 2–5mm	40	0
LS-TS	39.2	40.8	20.0	3.3	1.09	loam	fine granular 1–2mm	20	50
LS-CC	51.1	36.9	12.1	0.6	1.27	loam	fine granular 1–2mm	25	0
LS-OM	68.4	16.1	15.5	0.2	1.41	sandy-loam	fine granular 1–2mm	10 to 5	0

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888 **Table IV**

Year	Year	
	First year	Second year
Annual rainfall (mm)	434	992
Maximum rainfall (month/mm)	Dec 08/125	Dec 09/290
Minimum rainfall (month/mm)	Jul 09/1.00	Aug 10/4.20
Max. rainfall in 24h (mm)	38.4	49.0
Average annual temperature (°C)	10.1	10.3
Maximum average temperature (month/°C)	Aug 09/21.0	Jul 10/20.5
Minimum average temperature (month/°C)	Dec 08/2.00	Jan 10/1.60
Frost free days per year	223	267

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891 **Table V**

Sampling date	# rain days	Total rainfall (mm)	Max. rainfall 24h (mm)	Mean sediment yield (Mg ha ⁻¹)/Standard deviation (SD)									
				S/LCS-TS	LCS-CC	LS-TS	LS-CC	LS-OM					
01Oct2008 - 6Nov2008	9	82.6	38.4	Open plots were not yet installed									
19Dec2008	16	138	19.2	0.003 /0.003	0.02 /0.03	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.00 /0.00	
23Jan2009	14	17.4	4.80	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.00 /0.00	6.38 /9.02	3.27 /4.62				
30Jan2009	3	10.6	6.20	0.00 /0.00	-	-	0.00 /0.00	29.5 /41.7	44.2 /62.6				
12Feb2009	6	11.6	7.60	0.00 /0.00	-	-	10.0 /14.2	60.4 /40.3	102 /42.2				
13Mar2009	6	3.20	1.00	0.00 /0.00	-	-	0.39 /0.29	3.48 /4.14	14.9 /12.3				
21Apr2009	16	75.6	12.6	0.00 /0.00	-	-	0.32 /0.29	0.78 /0.46	35.4 /45.8				
09Jun2009	14	52.8	28.8	1.25 /1.23	-	-	0.31 /0.03	14.7 /2.81	43.3 /4.80				
24Jun2009	6	6.40	3.60	0.00 /0.00	-	-	0.10 /0.05	0.14 /0.02	0.49 /0.14				
12Aug2009	6	8.20	6.40	1.30 /0.89	-	-	0.34 /0.13	4.90 /1.47	30.0 /2.18				
01Oct2009	16	27.6	10.6	0.04 /0.03	-	-	0.07 /0.03	0.12 /0.01	8.56 /0.83				
1st year total	112	434	-	3	-	12	120	282					
Mean	10.2	39.5	12.7	0.26	-	1.16	12.0	28.2					
Median	9.00	17.4	7.60	0.00	-	0.20	4.19	22.4					
SD	5.04	43.2	11.6	0.54	-	3.12	19.3	31.1					
07Oct2009	3	5.80	4.00	0.03 /0.02	0.00 /0.00	0.00 /0.00	0.23 /0.12	3.13 /0.47					
29Oct2009	8	43.0	25.0	0.47 /0.44	0.00 /0.00	0.14 /0.10	1.20 /0.56	15.2 /11.8					
12Nov2009	9	6.80	2.80	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.00 /0.00	0.20 /0.05					
10Dec2009	13	51.6	20.2	0.50 /0.43	0.00 /0.00	0.00 /0.00	0.73 /0.12	19.4 /16.0					
18Jan2010	29	328	49.0	8.39 /1.56	7.38 /5.62	0.61 /0.43	23.8 /17.4	102 /12.5					
02Mar2010	30	153	24.8	1.29 /0.80	2.35 /1.97	1.67 /1.15	74.6 /34.0	38.1 /7.18					
05Apr2010	25	79.4	24.4	0.18 /0.15	1.82 /2.19	0.07 /0.08	2.56 /2.48	16.4 /2.27					
19May2010	27	156	26.0	2.04 /1.05	1.97 /2.81	0.09 /0.13	1.44 /0.83	43.0 /16.4					
01Jul2010	23	56.8	20.4	2.10 /1.18	2.27 /3.93	0.18 /0.19	5.50 /1.54	67.8 /30.5					
30Sep2010	28	35.8	4.20	3.28 /1.06	0.00 /0.00	0.32 /0.11	14.1 /2.94	29.7 /11.2					
03Nov2010	17	76.6	22.4	1.90 /0.77	0.00 /0.00	0.00 /0.00	1.42 /0.66	11.8 /3.36					
2nd year total	212	992	-	20	16	3	126	347					
Mean	19.3	90.2	20.3	1.84	1.44	0.28	11.4	31.6					
Median	23.0	56.8	22.4	1.29	0.00	0.09	1.44	19.4					
SD	9.67	93.2	13.2	2.42	2.23	0.50	22.2	30.5					
2 year total	324	1426	-	23	16	15	246	629					

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894 Table VI

Study year	Treatments compared	T-test results P value
2009	SCS-TS vs LS-TS	0.38
	LS-TS vs LS-CC	0.09
	LS-TS vs LS-OM	0.01*
	LS-CC vs LS-OM	0.18
2010	LCS-TS vs LCS-CC	0.69
	LCS-TS vs LS-TS	0.05*
	LCS-CC vs LS-CC	0.15
	LS-TS vs LS-CC	0.11
	LS-TS vs LS-OM	<0.01*
2009+2010	LS-CC vs LS-OM	0.09
	LS-TS vs LS-CC	0.02*
	LS-TS vs LS-OM	<0.01*
	LS-CC vs LS-OM	0.03*

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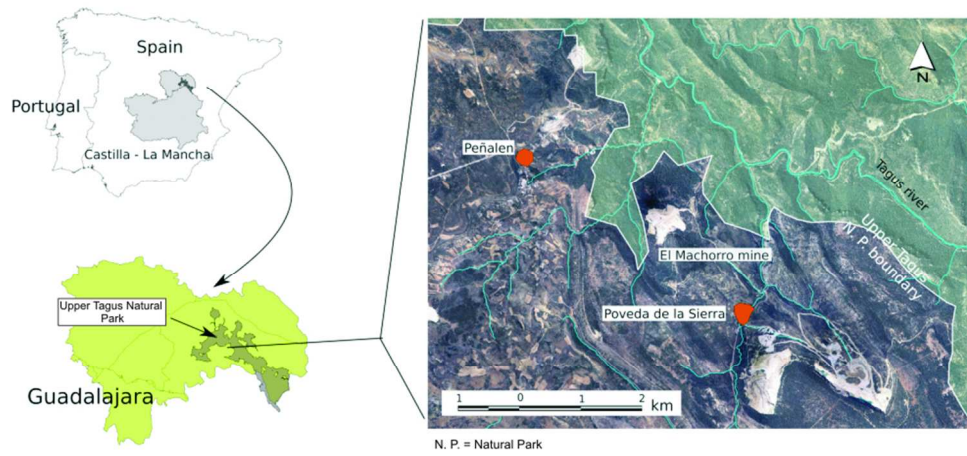
919 Table VII

Concave slope with topsoil	Linear slope with topsoil
Family Compositae	
<i>Cuprina crupinastrum</i>	<i>Hieracium pilosella</i>
<i>Leucanthemum vulgare</i>	
Family Euphorbiaceae	
<i>Euphorbia sp.</i>	<i>Euphorbia sp.</i>
Family Gramineae (=Poaceae)	
<i>Arrhenatherum elatius subsp. bulbosum</i>	<i>Brachypodium phoenicoides</i>
<i>Festuca gr. rubra</i>	<i>Bromus erectus</i>
Family Lamiaceae	
<i>Sideritis hirsuta</i>	<i>Thymus vulgaris</i>
Family Leguminosae (=Fabaceae)	
<i>Coronilla repanda</i>	<i>Coronilla repanda</i>
<i>Genista scorpius</i>	<i>Lotus corniculatus</i>
<i>Medicago lupulina</i>	<i>Medicago lupulina</i>
Family Liliaceae	
	<i>Aphyllanthes monspeliensis</i>
Family Plantaginaceae	
<i>Plantago sp.</i>	
Family Rosaceae	
<i>Filipendula vulgaris</i>	<i>Rosa sp.</i>
<i>Sanguisorba minor</i>	<i>Sanguisorba minor</i>
Family Rubiaceae	
<i>Asperula montana</i>	
<i>Galium lucidum</i>	<i>Galium lucidum</i>
Family Resedaceae	
	<i>Reseda alba</i>
	<i>Reseda phyteuma</i>

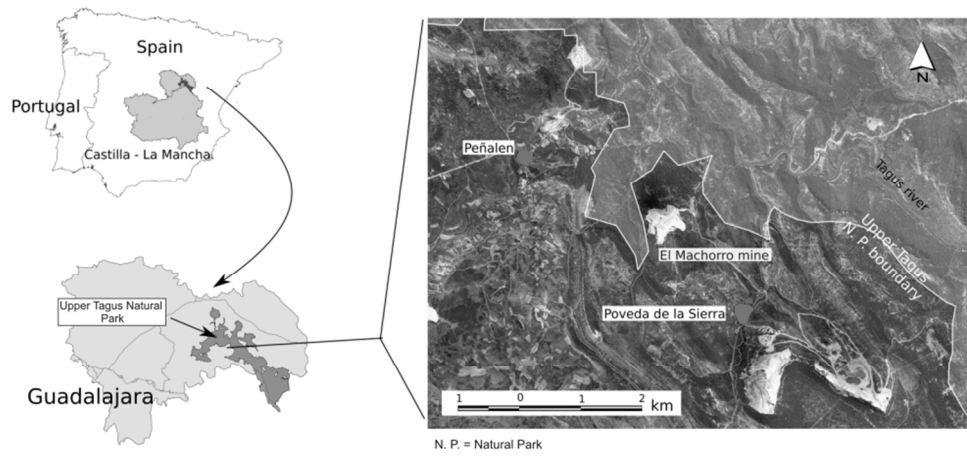
920 Table VIII

Topographic profile	Runoff control	Sediment yield control	Soil surface cover	Natural plant colonization
Concave slope	Watershed size and runoff increase downslope, while slope gradient decreases. Decrease of energy downslope.	Sediment accumulates at lower, flat part of slope	↑ compaction ↓ thickness heterogeneous distribution	heterogeneous distribution plant colonization more difficult in steep upper part of slope than in the lower part
Linear slope	Watershed size and runoff increase downslope, while slope gradient is constant. Increase of energy downslope.	None	↓ compaction ↑ thickness homogeneous distribution	homogeneous distribution

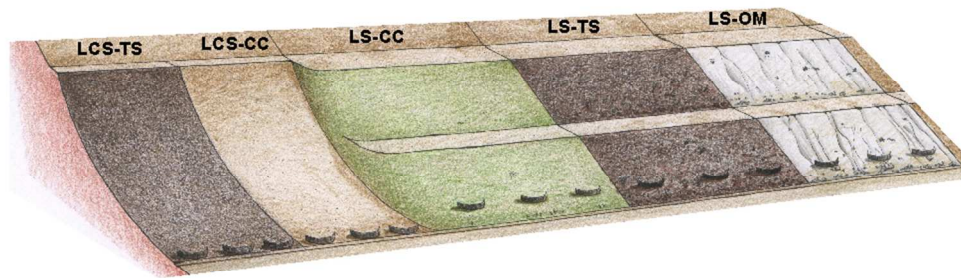
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Location of the study area within the Iberian Peninsula and within the province of Guadalajara. The experimental spoil heap is located at El Machorro mine.



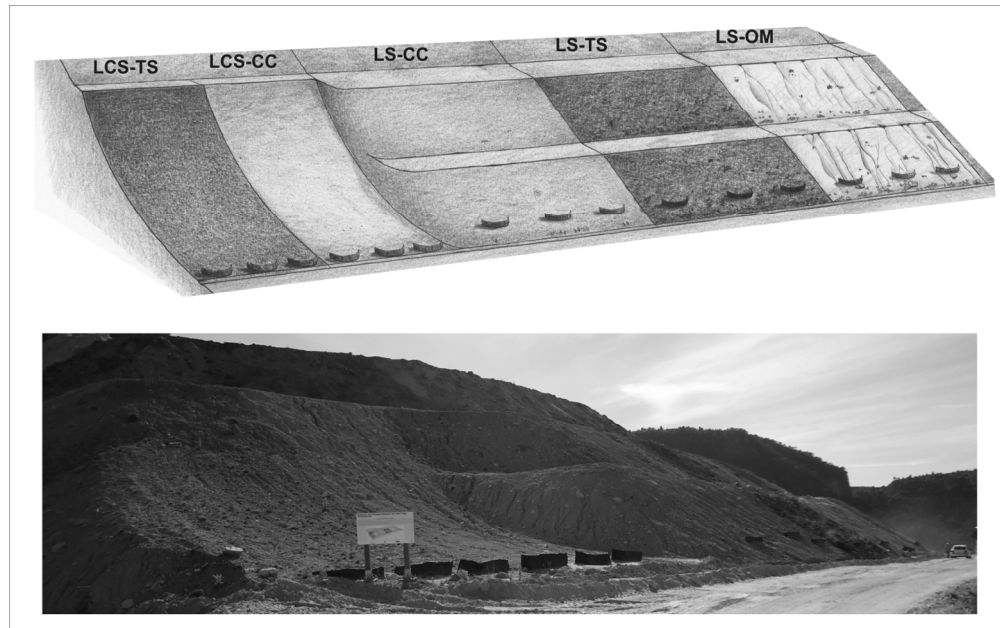
23 Location of the study area within the Iberian Peninsula and within the province of Guadalajara. The
24 experimental spoil heap is located at El Machorro mine.
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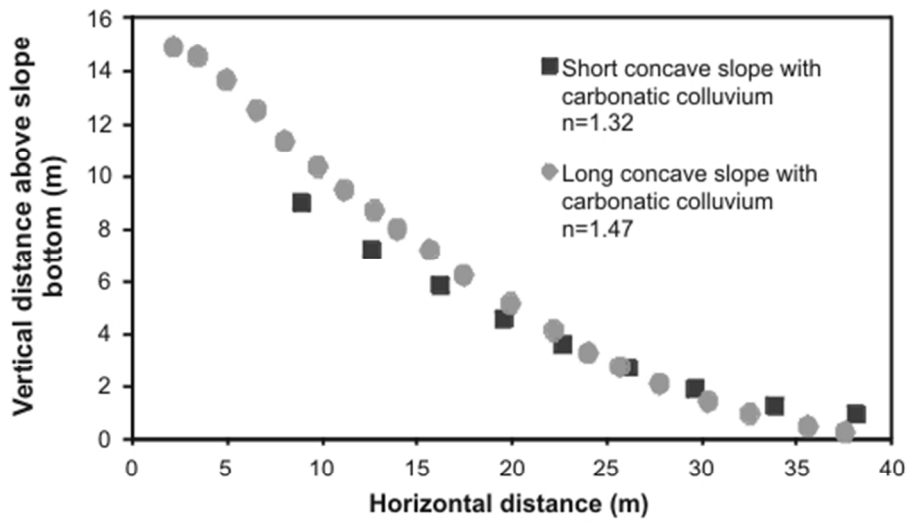
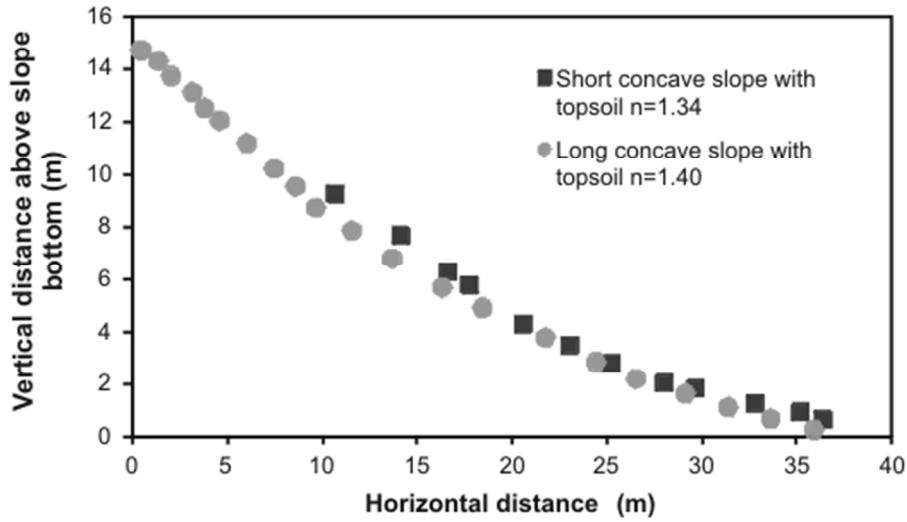
Experimental spoil heap of El Machorro mine, during the second study year, after conversion of the short concave slopes to long concave slopes. Top, treatment scheme; bottom, photograph taken October 2011, one year after experiment finished. LCS-TS = long concave slope with topsoil, LCS-CC = long concave slope with carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-TS = linear slope with topsoil, LS-OM = linear slope with overburden material. The long concave slope with overburden material (LCS-OM) could not be constructed.

150x100mm (170 x 170 DPI)

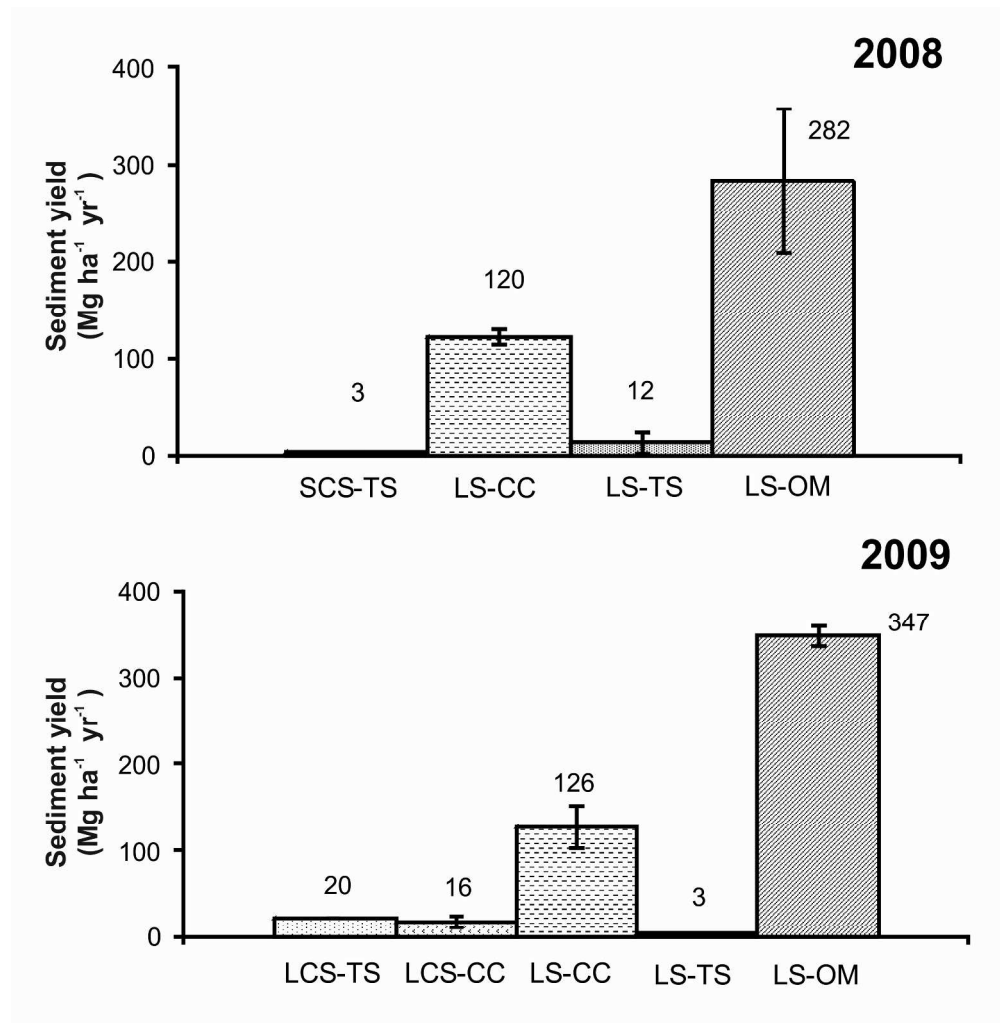


Experimental spoil heap of El Machorro mine, during the second study year, after conversion of the short concave slopes to long concave slopes. Top, treatment scheme; bottom, photograph taken October 2011, one year after experiment finished. LCS-TS = long concave slope with topsoil, LCS-CC = long concave slope with carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-TS = linear slope with topsoil, LS-OM = linear slope with overburden material. The long concave slope with overburden material (LCS-OM) could not be constructed.

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Concave slope shapes and their n values. The n value is an exponent that varies according to slope shape, following the equation of Stefano et al. (2000). The original, short concave slopes were converted to long concave slopes at the end of the first year.



Mean annual sediment yield (Mg ha⁻¹) by treatment and study year. The error bars represent the standard deviation. SCS-TS = short concave slope with topsoil, LCS-TS = long concave slope with topsoil, LCS-CC = long concave slope with carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-TS = linear slope with topsoil, LS-OM = linear slope with overburden material. The short concave slope with carbonate colluvium (SCS-CC) was not monitored during the first year.













364x371mm (200 x 200 DPI)

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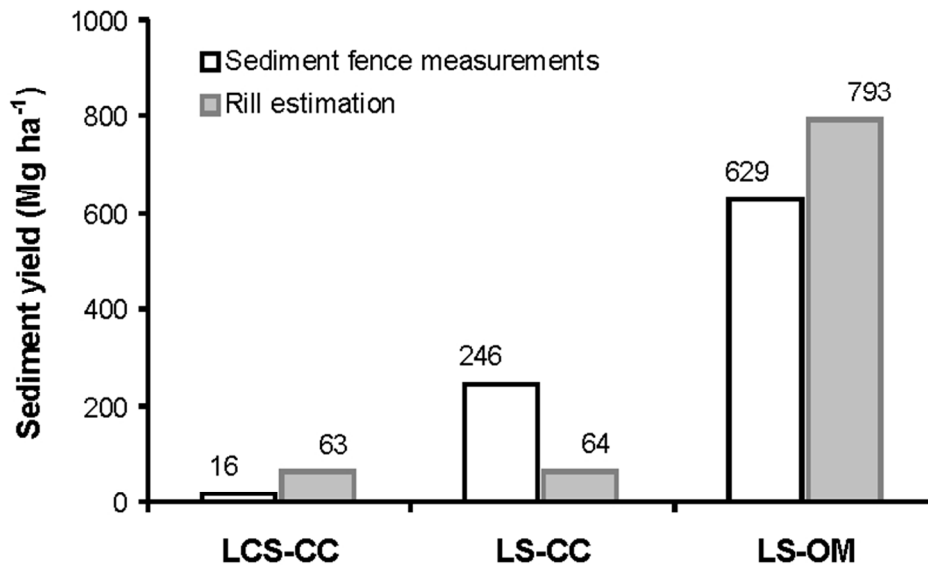
	Prior to installation of sediment fences	End of first study year, through end of study	One year after the experiment ended
Short / long concave slopes with topsoil and carbonate colluvium			
Linear slope with carbonate colluvium			
Linear slope with topsoil			
Linear slope with overburden material			

Photographs showing geomorphic evolution and vegetation colonization.

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	Prior to installation of sediment fences	End of first study year, through end of study	One year after the experiment ended
Short / long concave slopes with topsoil and carbonate colluvium			
Linear slope with carbonate colluvium			
Linear slope with topsoil			
Linear slope with overburden material			

Photographs showing geomorphic evolution and vegetation colonization.



Comparison of sediment yield measured from sediment fences with sediment yield estimated from rill erosion, for the two year study period. LCS-CC = long concave slope with carbonate colluvium, LS-CC = linear slope with carbonate colluvium, LS-OM = linear slope with overburden material.

127x76mm (170 x 170 DPI)