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STATISTICAL APPROACHES TO  
CERTAIN PROBLEMS IN GEOPHYSICS

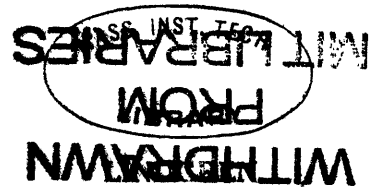
by

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## ABSTRACT

### STATISTICAL APPROACHES TO CERTAIN PROBLEMS IN GEOPHYSICS

by

Stephen Milton Simpson, Jr.

Submitted to the Department of Geology and Geophysics on August 14, 1953, in partial fulfillment of the requirements for the degree of Doctor of Philosophy.

Several specific problems in seismic and gravitational data interpretation are considered from the statistical viewpoint. Least squares techniques are applied to the two types of interpretation, and, for seismic records, other approaches are discussed.

The fitting of an  $n$ th order polynomial in  $x$  and  $y$  to gravity data by the method of least squares is investigated as a method for approximating regional gravitational anomalies. The normal equations for the general case are derived and simplification considered. It is shown that, with a symmetrical rectangular distribution of gravity readings, each set of these equations breaks up into smaller subsets. The resulting simplification brings fairly high order polynomials into the range of practical computation. For a particular gridwork the polynomial coefficients may be expressed explicitly as linear combinations of the right hand members of the normal equations. Once this is done, the least squares fitting of any data taken over such a gridwork may be effected relatively easily. The explicit expressions for the coefficients are derived for a square gridwork of 121 points and for polynomials of order 2, 3, and 4. A set of actual gravity readings is analysed in this fashion. The gravity residuals are determined and contoured. The comparison of these contours with each other (for various order polynomials) and with contours derived by a standard, much more involved process, is favorable. This consistency, despite certain detrimental features of the data used indicates that the method may deserve to find practical application as a routine, first step, gravity reduction procedure. The problem is pursued with regard to different gridworks, and a table is derived which contains, in effect, the normal equations for representative grids up to a size containing 2601 points, and for polynomials through order four.

As an approach to the understanding of linear operators, as they apply to the analysis of seismic records, a simple form of linear operator is studied. For this form of operator, the so-called "cosine operator", certain properties are derived in the general case, and interpreted geometrically. These include relationships between the exact

form of cosine operator chosen, the correlation properties of the series which the operator is to predict, the individual errors of prediction, and the sums of squared errors of prediction. The results are applied to two classes of time series in connection with spectrum analysis, and, for one class, filter characteristics are computed for a specific cosine operator.

An iterative method for determining least squares fits of linear operators to multiple time series is discussed geometrically. An argument is presented, based on the geometry of the two term operator, to show that, in the case of near singularity where many solutions will almost satisfy the least squares criterion, the exact solution is necessary for the purposes under consideration.

Several interpretive procedures are devised for finding information from seismic records. The first deals with discriminating an unknown velocity in a two velocity system. An adaption is made for detecting reflections, and practical example are given of the two uses. The second employs a form of testing phase, between seismic traces and their predictions by linear operators, to determine reflection times, and is illustrated with an example. The third combines the concept of ensemble averages with linear operators to determine step-out times of reflections. The last procedure suggests a special experimental arrangement, coupled with a certain type of correlation analysis, for detecting reflections.

Included as appendixes are descriptions of four programs written by the author for the Whirlwind I Digital Computer. These permit high speed computation of: two dimensional polynomial residuals; linear operator prediction errors and their running averages; least squares linear operator coefficients (by an iterative method); and auto-correlations, cross-correlations and "traveling" auto- or cross-correlations.

Thesis Supervisor: Patrick M. Hurley  
Title: Professor of Geology

## TABLE OF CONTENTS

Acknowledgments .....	1
Introduction .....	11
 <b>PART I LEAST SQUARES RESIDUAL GRAVITY</b>	
Introduction .....	I-1
Theory .....	I-2
Simplified Solutions .....	I-5
Case n = 1 .....	I-6
Case n = 2 .....	I-7
Case n = 3 .....	I-8
Case n = 4 .....	I-11
Discussion .....	I-15
Applications .....	I-16
Example .....	I-17
Setting up the Normal Equations .....	I-20
 <b>PART II SEISMIC RECORD ANALYSIS BY LINEAR OPERATORS</b>	
Introduction .....	II-1
Statistical Methods .....	II-2
Single Frequency Cosine Operators .....	II-4
Geometry of Cosine Operators .....	II-6
Conclusions .....	II-8
Example .....	II-9
Cyclical Nature of Cosine Operators .....	II-10
Cosine Operator Predicting an Autoregressive Series .	II-11
Cosine Operators on Series with Gaussian	
Spectrum Distribution .....	II-14
Computational Example .....	II-15
A Method of Finding Linear Operators	
for Least Squares Fitting Procedures .....	II-18
Accuracy .....	II-22
 <b>PART III SOME INTERPRETIVE PROCEDURES</b>	
Introduction .....	III-1
Velocity Separation .....	III-2
Tests of the Method .....	III-6
Conclusions .....	III-9
Phase Test .....	III-10
Ensemble Average .....	III-13
Travelling Correlations .....	III-15
 Conclusions and Suggestions for Further Work	
References	
APPENDIX A	Polynomial I Description and Use Of
APPENDIX B	Prediction XV Description and Use Of
APPENDIX C	Iteration I Description and Use Of
APPENDIX D	Auto Cross-Correlation I Description and Use Of
APPENDIX E	Biographical Note



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## INTRODUCTION

It is well known that experimental data taken in Geophysical studies surpasses in accuracy the interpretation that must be made on the data. The reason is that the problems are very complex. For one thing, it can be shown, in the treatment of certain types of problems, that no unique solution exists. An example is the infinity of possible mass distribution corresponding to a given gravity profile. In other problems the physical situation dealt with is so inhomogeneous and anisotropic that exact solution is impossible. It would be hopeless to attempt to explain rigorously the presence of any particular oscillation on a seismogram.

Data such as this, subject to a certain amount of randomness, and on which "best" estimates must be made, is well suited to statistical evaluation. The numerical data taken in gravity surveys does undergo evaluation of this type. The least squares approach, however, is not being utilized on a large scale. This is probably due to practical limitations, and it is one of the problems of this paper to see if these limitations may be minimized.

On the other hand, the raw data of seismology occurs in analogue or curve form. Standard procedures of interpretation consist mainly of rules of thumb, learned by long experience, and still largely dependent on the qualification of the individual interpreter. There is a need to put these

procedures on a more rigorous basis. This basis may be found in the concepts of time series as developed in economics, meteorology, and other fields. Much work must be done to determine the best means of applying these concepts to seismic data, since, in certain ways, both the data and the desired information are unique. Another purpose of this paper, then, is to propose several special methods of application, and to develop certain theory necessary for a better understanding of time series concepts as they apply to seismology.

Statistical methods, in general, require computation, and often on a large scale. A "program" written for a digital computer is a tool which will do this work automatically. The author has written several programs for the Whirlwind I Digital Computer to perform computations related to the above discussed problems, and includes these programs as appendixes, with the feeling that other investigators may find them useful.

**PART I**

## LEAST SQUARES RESIDUAL GRAVITY

### Introduction

Variations in the attraction of gravity over the surface of the earth are due to many causes, but these often fall into two general categories. Phenomena such as the thickening or thinning of the crust cause relatively slow, smooth and widespread gravity fluctuations. We call these regional effects. On the other hand, such things as ore body emplacements, caverns, and local density heterogeneities cause more rapid irregular changes, and these are termed local effects.

The actual gravity values measured over an area usually represent a combination of regional and local effects. The separation of these effects is of primary importance in interpretation, and many mathematical methods have been devised to eliminate guesswork in the problem. Essentially, most of the methods represent an averaging process which gives at each point an approximate value of the regional effect alone. The local effect is then found simply by subtraction from the measured values.

Many of these methods possess two undesirable aspects. First of all the averaging must be done at each point individually. Secondly, the averaging includes only gravity values in the vicinity of the point considered. It is hard to say just how serious these drawbacks are, but it seems worth-while to investigate a method which does not

encounter them. In a least squares approximation all values are averaged simultaneously. Moreover, the resulting approximation is not merely a set of discrete points but a continuous surface of values over the area, a property which is sometimes of value.

The purpose of Part I is to consider in some detail how the method of least squares may be applied to this problem, and how a simplified method of procedure may be set up for practical application.

Part I represents an extension of the work done by W.B. Agocs † . Agocs approximates the regional anomaly by a plane surface derived from least squares criteria. He shows that, in an artificial example, the residual anomaly is better derived from least squares procedures than by the use of the "arithmetic mean regional" procedure. For higher order polynomials than a plane surface the algebra rapidly becomes more involved.

### Theory

It is easiest to illustrate the method for an idealized geologic example in two dimensions. Fig. 1.1 shows a wave in the bottom of the crust and a single ore body emplacement. The points on the graph would then be the measured values of gravity across the area. Fitting a fairly low order polynomial to these values by least squares gives us the curve AB which best fits all the points.

† Ref. 1

This curve will approximate the regional effect more closely than the local effect, and it is apparent that the fit will be closest at some distance from the ore body. Thus the dashed line of Fig. 1.1, representing the difference between the polynomial and the observed values, gives a good indication of the location of the anomalous mass.

In the two-dimensional problem the approximating polynomial is a surface, and interpretation is made from contours of the residuals.

Let us approximate the regional gravity by a polynomial of order  $n$  in  $x$  and  $y$ .

$$G(xy) = \sum_{i=0}^n \sum_{j=0}^{n-1} c_{ij} x^i y^j \quad 1.1$$

Thus for  $n=2$

$$G(xy) = c_{00} + c_{10}x + c_{20}x^2 + c_{11}xy + c_{01}y + c_{02}y^2$$

The  $c$ 's are unknown coefficients to be determined in accordance with the condition that the sum of the squares of the residuals is to be minimized. Let  $g(xy)$  be the measured values of gravity. Then the residuals are

$$\begin{aligned} R(xy) &= g(xy) - G(xy) \\ \text{and} \quad R^2(xy) &= [g(xy) - G(xy)]^2 \end{aligned}$$

Hence

$$\begin{aligned} \sum_{xy} R^2(xy) &= \sum_{xy} \left[ \sum_{i=0}^n \sum_{j=0}^{n-1} (c_{ij} x^i y^j - \sum_{k=0}^n \sum_{l=0}^{n-k} c_{kl} x^k y^l) \right]^2 \\ &= -2 \sum_{xy} g(xy) \sum_{i=0}^n \sum_{j=0}^{n-1} c_{ij} x^i y^j + \sum_{xy} g^2(xy) \end{aligned}$$

or

$$\sum_{xy} R^2(xy) = \sum_{xy} \left[ \sum_{i=0}^n \sum_{j=0}^{n-1} \sum_{k=0}^n \sum_{l=0}^{n-k} c_{1j} c_{kl} x^{k+1} y^{l+j} \right] - 2 \sum_{xy} [g(xy) \sum_{i=0}^n \sum_{j=0}^{n-1} c_{ij} x^i y^j] + \sum_{xy} g^2(xy)$$

Differentiating this expression with respect to  $c_{1j}$ , and setting each derivative equal to zero for minimization, we obtain  $(n+1)(n+2)/2$  linear equations for the same number of unknown coefficients

$$\sum_{k=0}^n \sum_{l=0}^{n-k} c_{kl} \sum_{xy} x^{k+1} y^{l+j} = \sum_{xy} g(xy) x^i y^j \quad 1.2$$

where  $j = 0, 1, \dots, (n-1)$

$i = 0, 1, \dots, n$

There are really three variables, or sets of variables, in equations 1.2 — the order of the polynomial  $n$ , the set of points  $xy$ , and the set of gravity values at these points. The first two of these variables determine the coefficient matrix of the  $c_{kl}$ 's. Once these two are chosen, a unique inverse matrix exists, which, if found, may be used to compute the  $c_{kl}$ 's for all sets of gravity values taken over the same  $xy$  pattern. This alone would be a major simplification if the method were to be used on a production basis. But we shall also see that, using a simple reasonable restriction, both the problem of finding the inverse and the form of the inverse itself will be greatly simplified.



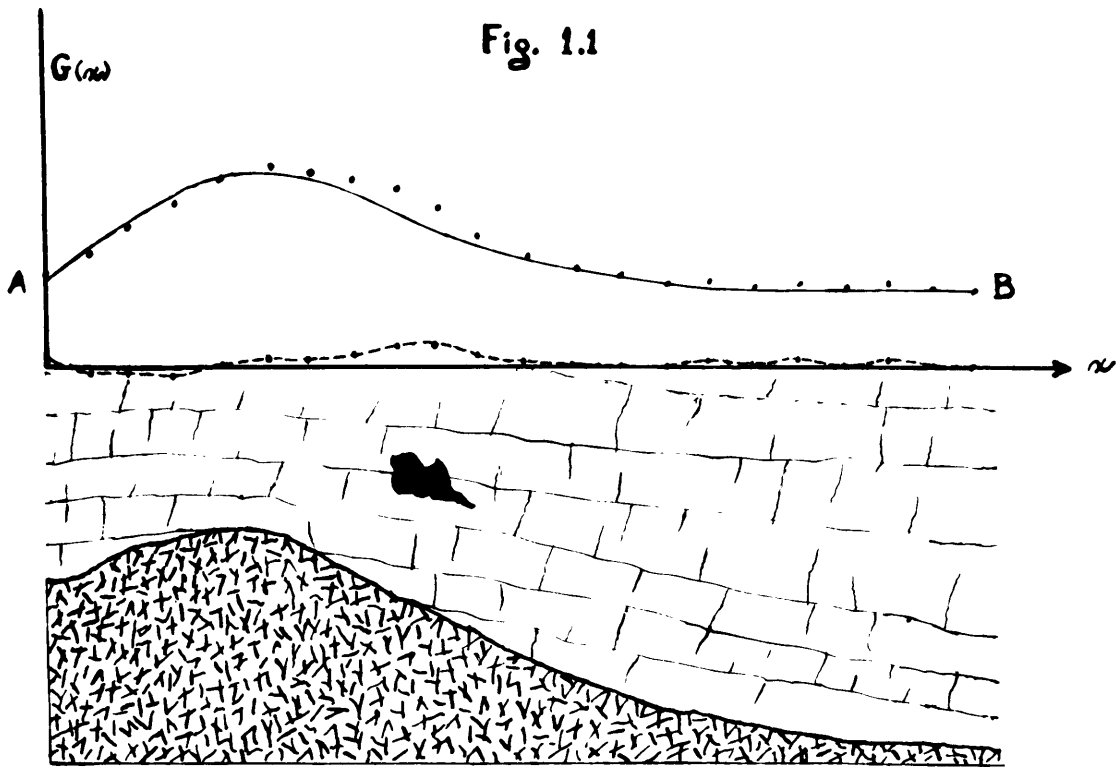
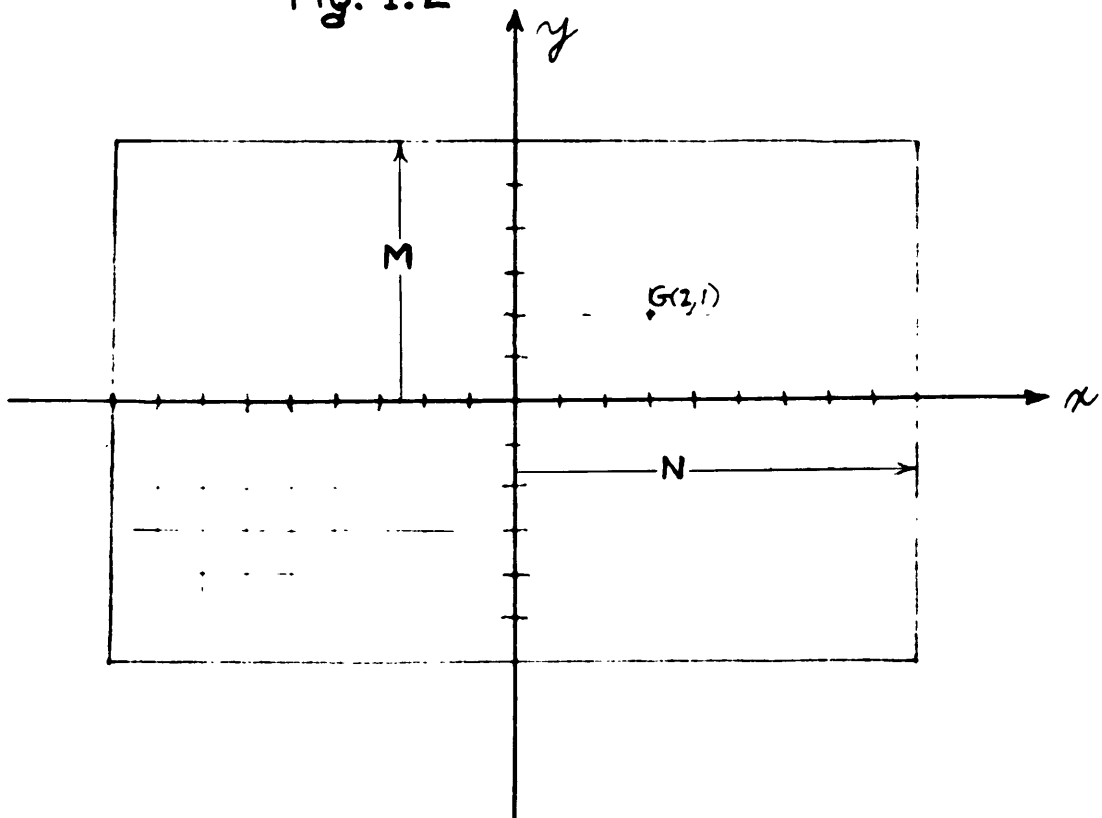


Fig. 1.2



Simplified Solutions

In many cases gravity readings are taken over a square, or at least rectangular, network. When this is so we may take the axes so that the rectangle is symmetrical about them, and number our ordinates and abscissae in integers, as in Fig. 1.2 . It is then easy to see that over such a network summations of the form  $\sum_{xy} x^i y^j$  will vanish whenever  $i$  or  $j$  is odd. Thus many of the coefficients of the  $c_{kl}$ 's in equations 1.2 will drop out. This leads to considerable simplification, with the bigger systems breaking up into several smaller ones. Furthermore, if we take a definite network we may solve the equations explicitly for the  $c_{kl}$ 's in terms of the summations  $\sum_{xy} (xy) x^i y^j$ .

To demonstrate how this is done we shall solve the equations for  $n = 1, 2, 3,$  and  $4,$  over a square network of 121 points. The systems are positive definite and symmetric, well adapted to solution by the matrix method of P.D. Crout† .

The non-vanishing summations over this network are

$\sum x^0 y^0 = M =$	121		
$\sum x^2 = \sum y^2 =$	1210	$\sum x^2 y^2 =$	12100
$\sum x^4 = \sum y^4 =$	21538	$\sum x^2 y^4 = \sum x^4 y^2 =$	215380
$\sum x^6 = \sum y^6 =$	451330	$\sum x^2 y^6 = \sum x^6 y^2 =$	4513300
$\sum x^8 = \sum y^8 =$	10185538	$\sum x^4 y^4 =$	3833764

† Ref. 2

Case n = 1

The normal equations are

$$\begin{aligned}c_{00}M + c_{10}\Sigma x + c_{01}\Sigma y &= \Sigma g(xy) \\c_{00}\Sigma x + c_{10}\Sigma x^2 + c_{01}\Sigma xy &= \Sigma g(xy)x \\c_{00}\Sigma y + c_{10}\Sigma xy + c_{01}\Sigma y^2 &= \Sigma g(xy)y\end{aligned}$$

Reducing immediately to

$$c_{00} = \frac{\Sigma g(xy)}{M} \quad c_{10} = \frac{\Sigma g(xy)x}{\Sigma x^2} \quad c_{01} = \frac{\Sigma g(xy)y}{\Sigma y^2}$$

giving

$$G(xy) = \frac{1}{121}\Sigma g(xy) + \frac{x}{1210}\Sigma g(xy)x + \frac{y}{1210}\Sigma g(xy)y$$

or, to six places

$$G(xy) = 8.26448 \cdot 10^{-4} [10\Sigma g(xy) + x\Sigma g(xy)x + y\Sigma g(xy)y]$$

Case n = 2

The normal equations are

$c_{00}$	$c_{01}$	$c_{02}$	$c_{11}$	$c_{10}$	$c_{20}$	
$M$	$\Sigma y$	$\Sigma y^2$	$\Sigma xy$	$\Sigma x$	$\Sigma x^2$	$= \Sigma g(xy)$
$\Sigma y$	$\Sigma y^2$	$\Sigma y^3$	$\Sigma xy^2$	$\Sigma xy$	$\Sigma x^2 y$	$= \Sigma g(xy)y$
$\Sigma y^2$	$\Sigma y^3$	$\Sigma y^4$	$\Sigma xy^3$	$\Sigma xy^2$	$\Sigma x^2 y^2$	$= \Sigma g(xy)y^2$
$\Sigma xy$	$\Sigma xy^2$	$\Sigma xy^3$	$\Sigma x^2 y^2$	$\Sigma x^2 y$	$\Sigma x^3 y$	$= \Sigma g(xy)xy$
$\Sigma x$	$\Sigma xy$	$\Sigma xy^2$	$\Sigma x^2 y$	$\Sigma x^2$	$\Sigma x^3$	$= \Sigma g(xy)x$
$\Sigma x^2$	$\Sigma x^2 y$	$\Sigma x^2 y^2$	$\Sigma x^3 y$	$\Sigma x^3$	$\Sigma x^4$	$= \Sigma g(xy)x^2$

which reduce to

$$c_{01} = \frac{1}{1210}\Sigma g(xy)y \quad c_{11} = \frac{1}{12100}\Sigma g(xy)xy \quad c_{10} = \frac{1}{1210}\Sigma g(xy)x$$

and three equations for  $c_{00}$ ,  $c_{02}$ , and  $c_{20}$

$$1210c_{00} + 1210c_{02} + 1210c_{20} = \Sigma g(xy)$$

$$1210c_{00} + 21538c_{02} + 12100c_{20} = \Sigma g(xy)y^2$$

$$1210c_{00} + 12100c_{02} + 21538c_{20} = \Sigma g(xy)x^2$$

The solutions are

$$c_{00} = \frac{1}{9438} [278\Sigma g(xy) - 10(\Sigma g(xy)x^2 + \Sigma g(xy)y^2)]$$

$$c_{02} = \frac{1}{9438} [\Sigma g(xy)y^2 - 10\Sigma g(xy)]$$

$$c_{20} = \frac{1}{9438} [\Sigma g(xy)x^2 - 10\Sigma g(xy)]$$

Thus

$$\begin{aligned} G(xy) = & \frac{1}{9438} [278\Sigma g(xy) - 10(\Sigma g(xy)x^2 + \Sigma g(xy)y^2)] \\ & + y \left[ \frac{1}{1210} \Sigma g(xy)y \right] \\ & + y^2 \left[ \frac{1}{9438} (\Sigma g(xy)y^2 - 10\Sigma g(xy)) \right] \\ & + xy \left[ \frac{1}{12100} \Sigma g(xy)xy \right] \\ & + x \left[ \frac{1}{1210} \Sigma g(xy)x \right] \\ & + x^2 \left[ \frac{1}{9438} \Sigma g(xy)x^2 - 10\Sigma g(xy) \right] \end{aligned}$$

Or to six places

$$\begin{aligned} G(xy) = & [.0294554\Sigma g(xy) - 1.05955 \cdot 10^{-3}(\Sigma g(xy)x^2 + \Sigma g(xy)y^2)] \\ & + y[8.26448 \cdot 10^{-4}\Sigma g(xy)y] \\ & + y^2[1.05955 \cdot 10^{-4}(\Sigma g(xy)y^2 - 10\Sigma g(xy))] \\ & + xy[8.26448 \cdot 10^{-5}\Sigma g(xy)xy] \\ & + x[8.26448 \cdot 10^{-4}\Sigma g(xy)x] \\ & + x^2[1.05955 \cdot 10^{-4}(\Sigma g(xy)x^2 - 10\Sigma g(xy))] \end{aligned}$$

Case n = 3

The normal equations are

	$c_{00}$	$c_{01}$	$c_{02}$	$c_{03}$	$c_{11}$	$c_{12}$	$c_{21}$	$c_{10}$	$c_{20}$	$c_{30}$	
1) M	y	y <sup>2</sup>	y <sup>3</sup>	xy	xy <sup>2</sup>	x <sup>2</sup> y	x	x <sup>2</sup>	x <sup>3</sup>	= g	
2) y	y <sup>2</sup>	y <sup>3</sup>	y <sup>4</sup>	xy <sup>2</sup>	xy <sup>3</sup>	x <sup>2</sup> y <sup>2</sup>	xy	x <sup>2</sup> y	x <sup>3</sup> y	= gy	
3) y <sup>2</sup>	y <sup>3</sup>	y <sup>4</sup>	y <sup>5</sup>	xy <sup>3</sup>	xy <sup>4</sup>	x <sup>2</sup> y <sup>3</sup>	xy <sup>2</sup>	x <sup>2</sup> y <sup>2</sup>	x <sup>3</sup> y <sup>2</sup>	= gy <sup>2</sup>	
4) y <sup>3</sup>	y <sup>4</sup>	y <sup>5</sup>	y <sup>6</sup>	xy <sup>4</sup>	xy <sup>5</sup>	x <sup>2</sup> y <sup>4</sup>	xy <sup>3</sup>	x <sup>2</sup> y <sup>3</sup>	x <sup>3</sup> y <sup>3</sup>	= gy <sup>3</sup>	
5) xy	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	x <sup>2</sup> y <sup>2</sup>	x <sup>2</sup> y <sup>3</sup>	x <sup>3</sup> y <sup>2</sup>	x <sup>2</sup> y	x <sup>3</sup> y	x <sup>4</sup> y	= gxy	
6) xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	x <sup>2</sup> y <sup>3</sup>	x <sup>2</sup> y <sup>4</sup>	x <sup>3</sup> y <sup>3</sup>	x <sup>2</sup> y <sup>2</sup>	x <sup>3</sup> y <sup>2</sup>	x <sup>4</sup> y <sup>2</sup>	= gxy <sup>2</sup>	
7) x <sup>2</sup> y	x <sup>2</sup> y <sup>2</sup>	x <sup>2</sup> y <sup>3</sup>	x <sup>2</sup> y <sup>4</sup>	x <sup>3</sup> y <sup>2</sup>	x <sup>3</sup> y <sup>3</sup>	x <sup>4</sup> y <sup>2</sup>	x <sup>3</sup> y	x <sup>4</sup> y	x <sup>5</sup> y	= gx <sup>2</sup> y	
8) x	xy	xy <sup>2</sup>	xy <sup>3</sup>	x <sup>2</sup> y	x <sup>2</sup> y <sup>2</sup>	x <sup>3</sup> y	x <sup>2</sup>	x <sup>3</sup>	x <sup>4</sup>	= gx	
9) x <sup>2</sup>	x <sup>2</sup> y	x <sup>2</sup> y <sup>2</sup>	x <sup>2</sup> y <sup>3</sup>	x <sup>3</sup> y	x <sup>3</sup> y <sup>2</sup>	x <sup>4</sup> y	x <sup>3</sup>	x <sup>4</sup>	x <sup>5</sup>	= gx <sup>2</sup>	
10) x <sup>3</sup>	x <sup>3</sup> y	x <sup>3</sup> y <sup>2</sup>	x <sup>3</sup> y <sup>3</sup>	x <sup>4</sup> y	x <sup>4</sup> y <sup>2</sup>	x <sup>5</sup> y	x <sup>4</sup>	x <sup>5</sup>	x <sup>6</sup>	= gx <sup>3</sup>	

Summations are assumed for all these quantities and g is written for g(xy). The equations reduce considerably.

Equation 5 gives us

$$c_{11} = \frac{1}{12100} \Sigma gxy$$

1, 3, and 9, combine to give three equations for  $c_{00}$ ,  $c_{02}$ , and  $c_{20}$ , which have the same solutions as for the case n = 2. 2, 4, and 7, and 6, 8, and 10, combine to give two independent systems which have identical coefficients. Thus 2, 4, 7, are

$$\begin{aligned} 1210c_{01} + 21538c_{03} + 12100c_{21} &= gy \\ 21538c_{01} + 451330c_{03} + 215380c_{21} &= gy^3 \\ 12100c_{01} + 215380c_{03} + 215380c_{21} &= gx^2y \end{aligned}$$

With solutions

$$c_{01} = \frac{1}{679536} [4450\Sigma gy - 178\Sigma gy^3 - 72\Sigma gx^2 y]$$

$$c_{03} = \frac{1}{679536} [10\Sigma gy^3 - 178\Sigma gy]$$

$$c_{21} = \frac{1}{94380} [\Sigma gx^2 y - 10\Sigma gy]$$

also

$$c_{10} = \frac{1}{679536} [4450\Sigma gx - 178\Sigma gx^3 - 72\Sigma gxy^2]$$

$$c_{30} = \frac{1}{679536} [10\Sigma gx^3 - 178\Sigma gx]$$

$$c_{12} = \frac{1}{94380} [\Sigma gxy^2 - 10\Sigma gx]$$

and from the case  $n = 2$

$$c_{00} = \frac{1}{9438} [278\Sigma g - 10(\Sigma gx^2 + \Sigma gy^2)]$$

$$c_{02} = \frac{1}{9438} [\Sigma gy^2 - 10\Sigma g]$$

$$c_{20} = \frac{1}{9438} [\Sigma gx^2 - 10\Sigma g]$$

We also have

$$c_{11} = \frac{1}{12100} \Sigma gxy$$

To six places

$G(xy) =$

$$\begin{aligned} &+ [.0294554\Sigma g - 1.05955 \cdot 10^{-3}(\Sigma gx^2 + \Sigma gy^2)] \\ &+ y[6.54859 \cdot 10^{-3}\Sigma gy - 2.61943 \cdot 10^{-4}\Sigma gy^3 - 1.05955 \cdot 10^{-3}\Sigma gx^2y \\ &+ y^2[1.05955 \cdot 10^{-4}(\Sigma gy^2 - 10\Sigma g)] \\ &+ y^3[1.47159 \cdot 10^{-5}\Sigma gy^3 - 2.61943 \cdot 10^{-4}\Sigma gy] \\ &+ xy[8.26445 \cdot 10^{-5}\Sigma gxy] \\ &+ xy^2[1.05955 \cdot 10^{-5}\Sigma gxy^2 - 1.05955 \cdot 10^{-4}\Sigma gx] \\ &+ x^2y[1.05955 \cdot 10^{-5}\Sigma gx^2y - 1.05955 \cdot 10^{-4}\Sigma gy] \\ &+ x[6.54859 \cdot 10^{-3}\Sigma gx - 2.61943 \cdot 10^{-4}\Sigma gx^3 - 1.05955 \cdot 10^{-3}\Sigma gxy^2] \\ &+ x^2[1.05955 \cdot 10^{-4}(\Sigma gx^2 - 10\Sigma g)] \\ &+ x^3[1.47159 \cdot 10^{-5}\Sigma gx^3 - 2.61943 \cdot 10^{-4}\Sigma gx] \end{aligned}$$

Case n = 4

The normal equations are

	c <sub>00</sub>	c <sub>01</sub>	c <sub>02</sub>	c <sub>03</sub>	c <sub>04</sub>	c <sub>11</sub>	c <sub>12</sub>	c <sub>13</sub>	c <sub>21</sub>	c <sub>22</sub>	c <sub>31</sub>	c <sub>10</sub>	c <sub>20</sub>	c <sub>30</sub>	c <sub>40</sub>	
1)	M	y	y <sup>2</sup>	y <sup>3</sup>	y <sup>4</sup>	xy	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>2</sup>	xy <sup>22</sup>	xy <sup>3</sup>	x	x <sup>2</sup>	x <sup>3</sup>	x <sup>4</sup>	= g
2)	y	y <sup>2</sup>	y <sup>3</sup>	y <sup>4</sup>	y <sup>5</sup>	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>22</sup>	xy <sup>23</sup>	xy <sup>32</sup>	xy	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	= gy
3)	y <sup>2</sup>	y <sup>3</sup>	y <sup>4</sup>	y <sup>5</sup>	y <sup>6</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>23</sup>	xy <sup>24</sup>	xy <sup>33</sup>	xy <sup>2</sup>	xy <sup>22</sup>	xy <sup>32</sup>	xy <sup>42</sup>	= gy <sup>2</sup>
4)	y <sup>3</sup>	y <sup>4</sup>	y <sup>5</sup>	y <sup>6</sup>	y <sup>7</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>6</sup>	xy <sup>24</sup>	xy <sup>25</sup>	xy <sup>34</sup>	xy <sup>3</sup>	xy <sup>23</sup>	xy <sup>33</sup>	xy <sup>43</sup>	= gy <sup>3</sup>
5)	y <sup>4</sup>	y <sup>5</sup>	y <sup>6</sup>	y <sup>7</sup>	y <sup>8</sup>	xy <sup>5</sup>	xy <sup>6</sup>	xy <sup>7</sup>	xy <sup>25</sup>	xy <sup>26</sup>	xy <sup>35</sup>	xy <sup>4</sup>	xy <sup>24</sup>	xy <sup>34</sup>	xy <sup>44</sup>	= gy <sup>4</sup>
6)	xy	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>22</sup>	xy <sup>23</sup>	xy <sup>24</sup>	xy <sup>32</sup>	xy <sup>33</sup>	xy <sup>42</sup>	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	= gxy
7)	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>6</sup>	xy <sup>23</sup>	xy <sup>24</sup>	xy <sup>25</sup>	xy <sup>33</sup>	xy <sup>34</sup>	xy <sup>43</sup>	xy <sup>22</sup>	xy <sup>32</sup>	xy <sup>42</sup>	xy <sup>52</sup>	= gxy <sup>2</sup>
8)	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>6</sup>	xy <sup>7</sup>	xy <sup>24</sup>	xy <sup>25</sup>	xy <sup>26</sup>	xy <sup>34</sup>	xy <sup>35</sup>	xy <sup>44</sup>	xy <sup>23</sup>	xy <sup>33</sup>	xy <sup>43</sup>	xy <sup>53</sup>	= gxy <sup>3</sup>
9)	xy <sup>2</sup>	xy <sup>22</sup>	xy <sup>23</sup>	xy <sup>24</sup>	xy <sup>25</sup>	xy <sup>32</sup>	xy <sup>33</sup>	xy <sup>34</sup>	xy <sup>42</sup>	xy <sup>43</sup>	xy <sup>52</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>6</sup>	= gxy <sup>2</sup>
10)	xy <sup>22</sup>	xy <sup>23</sup>	xy <sup>24</sup>	xy <sup>25</sup>	xy <sup>26</sup>	xy <sup>33</sup>	xy <sup>34</sup>	xy <sup>35</sup>	xy <sup>43</sup>	xy <sup>44</sup>	xy <sup>53</sup>	xy <sup>32</sup>	xy <sup>42</sup>	xy <sup>52</sup>	xy <sup>62</sup>	= gxy <sup>22</sup>
11)	xy <sup>3</sup>	xy <sup>32</sup>	xy <sup>33</sup>	xy <sup>34</sup>	xy <sup>35</sup>	xy <sup>42</sup>	xy <sup>43</sup>	xy <sup>44</sup>	xy <sup>52</sup>	xy <sup>53</sup>	xy <sup>62</sup>	xy <sup>4</sup>	xy <sup>5</sup>	xy <sup>6</sup>	xy <sup>7</sup>	= gxy <sup>3</sup>
12)	x	xy	xy <sup>2</sup>	xy <sup>3</sup>	xy <sup>4</sup>	xy <sup>2</sup>	xy <sup>22</sup>	xy <sup>23</sup>	xy <sup>3</sup>	xy <sup>32</sup>	xy <sup>4</sup>	x <sup>2</sup>	x <sup>3</sup>	x <sup>4</sup>	x <sup>5</sup>	= gx
13)	x <sup>2</sup>	x <sup>2</sup>	xy <sup>22</sup>	xy <sup>23</sup>	xy <sup>24</sup>	xy <sup>3</sup>	xy <sup>32</sup>	xy <sup>33</sup>	xy <sup>4</sup>	xy <sup>42</sup>	xy <sup>5</sup>	x <sup>3</sup>	x <sup>4</sup>	x <sup>5</sup>	x <sup>6</sup>	= gx <sup>2</sup>
14)	x <sup>3</sup>	x <sup>3</sup>	xy <sup>32</sup>	xy <sup>33</sup>	xy <sup>34</sup>	xy <sup>4</sup>	xy <sup>42</sup>	xy <sup>43</sup>	xy <sup>5</sup>	xy <sup>52</sup>	xy <sup>6</sup>	x <sup>4</sup>	x <sup>5</sup>	x <sup>6</sup>	x <sup>7</sup>	= gx <sup>3</sup>
15)	x <sup>4</sup>	x <sup>4</sup>	xy <sup>42</sup>	xy <sup>43</sup>	xy <sup>44</sup>	xy <sup>5</sup>	xy <sup>52</sup>	xy <sup>53</sup>	xy <sup>6</sup>	xy <sup>62</sup>	xy <sup>7</sup>	x <sup>5</sup>	x <sup>6</sup>	x <sup>7</sup>	x <sup>8</sup>	= gx <sup>4</sup>

Equations 1, 3, 5, 10, 13, 15, reduce to give a system of six equations for c<sub>00</sub>, c<sub>02</sub>, c<sub>04</sub>, c<sub>22</sub>, c<sub>20</sub>, and c<sub>40</sub>. 2, 4, 9, and 7, 12, 14, give two sets of equations for c<sub>01</sub>, c<sub>03</sub>, c<sub>21</sub>, and c<sub>10</sub>, c<sub>30</sub>, c<sub>12</sub>, respectively, which are equivalent to the corresponding equations for the case n = 3.



The new equations to be solved are

$$\begin{array}{rcccccc}
 c_{00} & c_{02} & c_{04} & c_{22} & c_{20} & c_{40} & \\
 M & y^2 & y^4 & x^2 y^2 & x^2 & x^4 & = \Sigma g \\
 y^2 & y^4 & y^6 & x^2 y^4 & x^2 y^2 & x^4 y^2 & = \Sigma g y^2 \\
 y^4 & y^6 & y^8 & x^2 y^6 & x^2 y^4 & x^4 y^4 & = \Sigma g y^4 \\
 x^2 y^2 & x^2 y^4 & x^2 y^6 & x^4 y^4 & x^4 y^2 & x^6 y^2 & = \Sigma g x^2 y^2 \\
 x^2 & x^2 y^2 & x^2 y^4 & x^4 y^2 & x^4 & x^6 & = \Sigma g x^2 \\
 x^4 & x^4 y^2 & x^4 y^4 & x^6 y^2 & x^6 & x^8 & = \Sigma g x^4
 \end{array}$$

and

$$\begin{array}{rcc}
 c_{11} & c_{13} & c_{31} \\
 x^2 y^2 & x^2 y^4 & x^4 y^2 & = \Sigma g x y \\
 x^2 y^4 & x^2 y^6 & x^4 y^4 & = \Sigma g x y^3 \\
 x^4 y^2 & x^4 y^4 & x^6 y^2 & = \Sigma g x^3 y
 \end{array}$$

The last set has solutions

$$\begin{aligned}
 c_{11} &= \frac{1}{679536} [689.84 \Sigma g x y - 17.8 (\Sigma g x y^3 + \Sigma g x^3 y)] \\
 c_{13} &= \frac{1}{679536} [\Sigma g x y^3 - 17.8 \Sigma g x y] \\
 c_{31} &= \frac{1}{679536} [\Sigma g x^3 y - 17.8 \Sigma g x y]
 \end{aligned}$$

The solution of the first set to six places is

$$c_{00} = 4.53280 \cdot 10^{-2} \Sigma g + 1.58932 \cdot 10^{-4} (\Sigma g y^4 + \Sigma g x^4) \\ 1.35839 \cdot 10^{-4} \Sigma g x^2 y^2 - 6.399122 \cdot 10^{-3} (\Sigma g y^2 + \Sigma g x^2)$$

$$c_{02} = 1.62141 \cdot 10^{-3} \Sigma g y^2 - 2.81527 \cdot 10^{-3} \Sigma g \\ -5.51847 \cdot 10^{-5} \Sigma g y^4 - 1.35839 \cdot 10^{-5} (\Sigma g x^2 y^2 - 10 \Sigma g x^2)$$

$$c_{04} = 2.20739 \cdot 10^{-6} \Sigma g y^4 - 5.51847 \cdot 10^{-5} \Sigma g y^2 + 1.58932 \cdot 10^{-5} \Sigma g$$

$$c_{22} = 1.35839 \cdot 10^{-6} (\Sigma g x^2 y^2 - 10 (\Sigma g x^2 + \Sigma g y^2)) + 100 \Sigma g$$

$$c_{20} = 1.62141 \cdot 10^{-3} \Sigma g x^2 - 2.81527 \cdot 10^{-3} \Sigma g \\ -5.51847 \cdot 10^{-5} \Sigma g x^2 - 1.35839 \cdot 10^{-5} (\Sigma g x^2 y^2 - 10 \Sigma g y^2)$$

$$c_{40} = 2.20739 \cdot 10^{-6} \Sigma g x^4 - 5.51847 \cdot 10^{-5} \Sigma g x^2 + 1.58932 \cdot 10^{-5} \Sigma g$$

To simplify writing  $G(xy)$  we introduce the abbreviations

A = $\Sigma g$	H = $\Sigma g x^3 y$
B = $\Sigma g x$	I = $\Sigma g x y^2$
C = $\Sigma g x^2$	J = $\Sigma g x^2 y^2$
D = $\Sigma g x^3$	K = $\Sigma g x y^3$
E = $\Sigma g x^4$	L = $\Sigma g y$
F = $\Sigma g x y$	M = $\Sigma g y^2$
G = $\Sigma g x^2 y$	N = $\Sigma g y^3$
	P = $\Sigma g y^4$

Hence

$$\begin{aligned}G(xy) = & [4.53280 \cdot 10^{-2}A + 1.58932 \cdot 10^{-4}(P + E) \\& + 1.35839 \cdot 10^{-4}J - 6.39122 \cdot 10^{-3}(M + C)] \\& + y[6.54859 \cdot 10^{-3}L - 2.61943 \cdot 10^{-4}N - 1.05955 \cdot 10^{-3}G] \\& + y^2[1.62141 \cdot 10^{-3}M - 2.81527 \cdot 10^{-3}A - 5.51847 \cdot 10^{-5}P \\& - 1.35839 \cdot 10^{-5}(J - 10C)] \\& + y^3[1.47159 \cdot 10^{-5}N - 2.61943 \cdot 10^{-4}L] \\& + y^4[2.20739 \cdot 10^{-6}P - 5.51847 \cdot 10^{-5}M + 1.58932 \cdot 10^{-5}A] \\& + xy[1.01516 \cdot 10^{-3}F - 2.61943 \cdot 10^{-5}(K + H)] \\& + xy^2[1.05955 \cdot 10^{-5}I - 1.05955 \cdot 10^{-4}B] \\& + xy^3[1.47159 \cdot 10^{-6}K - 2.61943 \cdot 10^{-5}F] \\& + x^2y[1.05955 \cdot 10^{-5}G - 1.05955 \cdot 10^{-4}B] \\& + x^2y^2[1.35839 \cdot 10^{-6}(J - 10(C + M) + 100A)] \\& + x^3y[1.47159 \cdot 10^{-6}H - 2.61943 \cdot 10^{-5}F] \\& + x[6.54859 \cdot 10^{-3}B - 2.61943 \cdot 10^{-4}D - 1.05955 \cdot 10^{-3}I] \\& + x^2[1.62141 \cdot 10^{-3}C - 2.81527 \cdot 10^{-3}A - 5.51847 \cdot 10^{-5}C \\& - 1.35839 \cdot 10^{-5}(J - 10M)] \\& + x^3[1.47159 \cdot 10^{-5}D - 2.61943 \cdot 10^{-4}B] \\& + x^4[2.20739 \cdot 10^{-6}E - 5.51847 \cdot 10^{-5}C - 1.58932 \cdot 10^{-5}A]\end{aligned}$$

## Discussion

An interesting property which has developed in these four cases makes the extension to higher order approximations somewhat simpler. If  $n$  is odd then all the coefficients  $c_{ij}$  whose subscripts add to an even number are the same as the corresponding coefficients for the  $(n - 1)$ st case. If  $n$  is even the coefficients with subscripts adding to an odd number are the same as for the preceding case. This property can be shown directly from equation 1.2.

Thus for the case  $n = 5$  we expect nine of the coefficients ( $c_{00}, c_{02}, c_{04}, c_{40}, c_{20}, c_{11}, c_{13}, c_{31}$ ) to be the same as for the case  $n = 4$ , and we need only write the twelve remaining equations for  $i + j$  odd.

A polynomial of order equal to the number of points taken will exactly fit the data. However it is practically impossible to use polynomials even approaching such a high order for reasonably-sized gridworks, and this danger seems slight. There is still a real problem in the choice of  $n$ . If the regional effect is in reality a fairly low order effect, polynomials of high  $n$  will begin to approximate the local anomalies too closely. Other systems, however, run into the same problem, and this point would be best settled by experience with the data.

Another important practical consideration is the amount of work to be done, i.e., the determination of the

summations  $\Sigma g x^i y^j$ , of the  $c_{ij}$ 's and the solution of  $G(xy)$  at each point. We devote the next section to this problem.

### Applications

To illustrate the work necessary we discuss a convenient scheme for application to the case  $n = 3$ . The use of a computing machine with cumulative multiplication is desirable.

Assume that the grid has been determined and the gravity values written at each intersection as shown. This is done on tracing paper as shown in Fig. 1.3.

The numbers  $\alpha_1$  and  $\beta_1$  above each vertical line and to the left of each horizontal line represent the sums of  $g(xy)$  along those lines. Then as we may easily compute the sums  $\Sigma g$ ,  $\Sigma gx$ ,  $\Sigma gx^2$ ,  $\Sigma gx^3$ ,  $\Sigma gy$ ,  $\Sigma gy^2$ ,  $\Sigma gy^3$ , from the relations

$$\begin{aligned} \Sigma g &= \Sigma_1 \alpha_1 & \Sigma gx^3 &= \Sigma_1 \alpha_1 (1)^3 & gy^2 &= \Sigma_1 \beta_1 (1)^2 \\ \Sigma gx &= \Sigma_1 \alpha_1 (1) & \Sigma gy &= \Sigma_1 \beta_1 (1) & gy^3 &= \Sigma_1 \beta_1 (1)^3 \\ \Sigma gx^2 &= \Sigma_1 \alpha_1 (1)^2 & & & & \end{aligned}$$

Each of these computations involves one machine operation of eleven cumulative multiplications.

For the remaining summations  $\Sigma gxy$ ,  $\Sigma gx^2y$ ,  $\Sigma gxy^2$  it is convenient to have a similar grid which can be placed under the original one. This second grid has the values of  $xy$ ,  $x^2y$ ,  $xy^2$ , at each point as shown in Fig. 1.4 and can be used for each application.

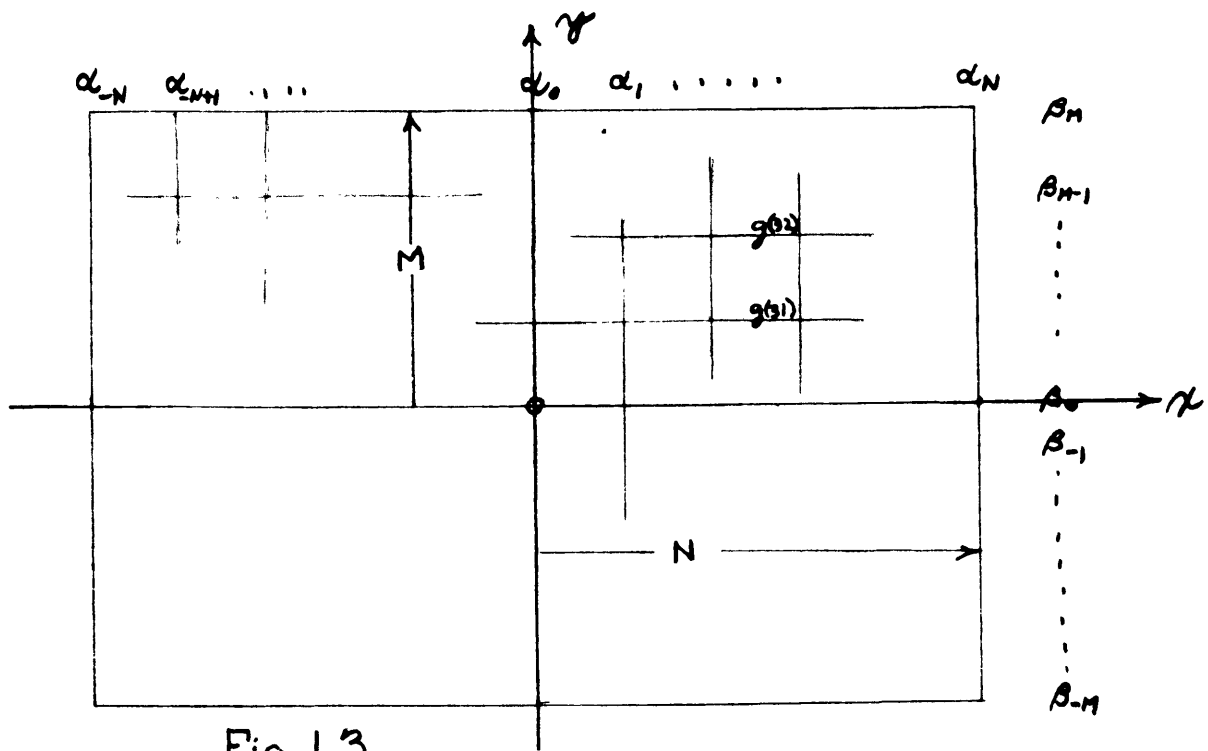


Fig. 1.3

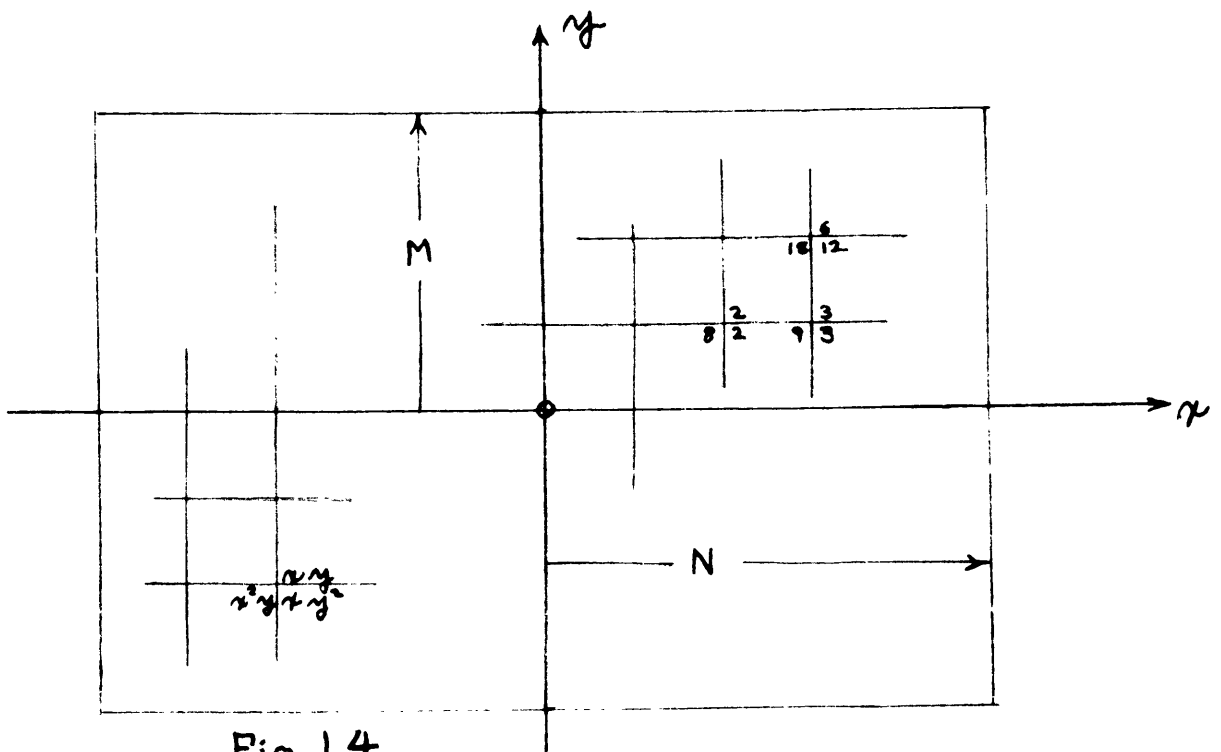


Fig. 1.4

The values of  $g(xy)$  then appear in the vacant upper left hand corner of each point, making the multiplications apparent. Each of these three summations then involves a cumulative addition of 100 multiplications.

The  $c_{ij}$ 's are then found as ten cumulative additions of two of three multiplications each.

$G(xy)$  is now completely determined with the writing down of less than 50 numbers and it remains to solve this equation for each point. This involves ten cumulative multiplications at each of the 121 points with a final subtraction to determine the residuals. A second tracing paper grid laid over both of the others would simplify this and the residuals could be written down in a form ready to be contoured.

A nice feature of this scheme is the absence of any tabulation of data. It may be extended fairly simply to higher degrees.

#### Example

As a test of this method residuals were computed on gravity readings supplied by a mining company. This data was not in a convenient form for use since the readings were taken in mine tunnels and not over a grid. To get them in grid form, the readings were first contoured as shown in Fig. 1.5 and then values extrapolated to the grid. This involves several inaccuracies. First of all, where readings are sparse, the extrapolated values are bound to contain

considerable error. Secondly, the computations weight all values equally so that the accurately extrapolated points, in areas of dense readings, suffer from the inaccuracies in the less dense areas.

Three sets of residuals were computed, one for second, third, and fourth order polynomials. This was done to test the effect of polynomial degree on the residuals. The polynomials were fitted directly to the raw gravity, without making the usual topographic corrections. The justification for neglecting to do this is seen in Fig. 1.11. This figure shows values of regional minus topographic corrections computed by the mining company, and contours of these values. The contours demonstrate that this correction is a low order effect (in this particular situation) and can obviously be easily absorbed into a polynomial as low as degree two.

Figs. 1.6, 1.7, and 1.8 then show residuals for second, third, and fourth order polynomials respectively, and were computed as described previously. Once the reading had been contoured and extrapolated, it took about a day to compute each set of residuals. The computed polynomials appear in the upper right-hand corner of Figs. 1.6, 1.7, and 1.8. Fig. 1.10 shows the contours of the polynomials themselves, and shows that the similarity of the residuals is due to the similarity of the polynomials used.



In Fig. 1.9 is contoured the residual gravity as computed by the mining company which supplied the data. Their computational procedure required several months to produce this diagram, which, in important respects, is quite similar to the contours of Figs. 1.6, 1.7, and 1.8. Part of the difference is due to the fact that the least squares residuals are forced to oscillate around a mean of zero, so that many negative contours appear. Other differences may well be attributable to the inaccuracies of contouring as mentioned above. It seems clear however, that the similarity is sufficiently great to justify the use of the least squares procedure, at least for a first evaluation of gravity data. This seems particularly true in view of the relative speed with which this procedure may be carried out.

These results were encouraging enough so that a program was written for the WWI Digital Computer to perform the majority of the computations automatically. This program finds the residuals for a polynomial up to the sixth order over an arbitrary grid shape, once the polynomial is known. A description of this program appears in Appendix A.

In the next section, we take up the problem of setting up the normal equations for various sizes and shapes of grids.

RAW GRAVITY -1300.0 CONTOURED

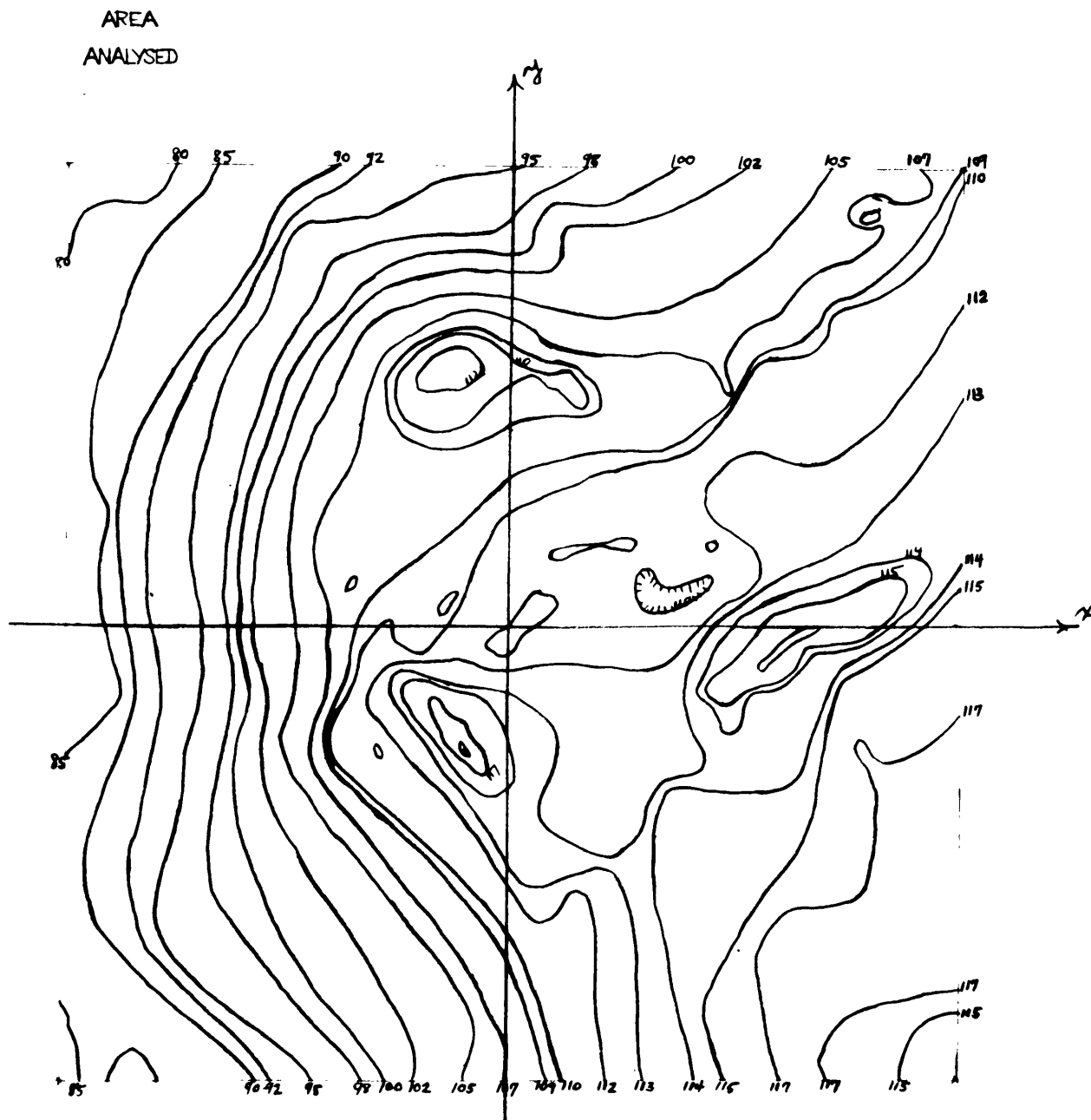


Fig. 1.5

# RESIDUAL GRAVITY WITH SECOND ORDER POLYNOMIAL

$$G(x,y) = 41.37 - .932y - .341y^2 - .0557xy + 2.904x - .385x^2$$

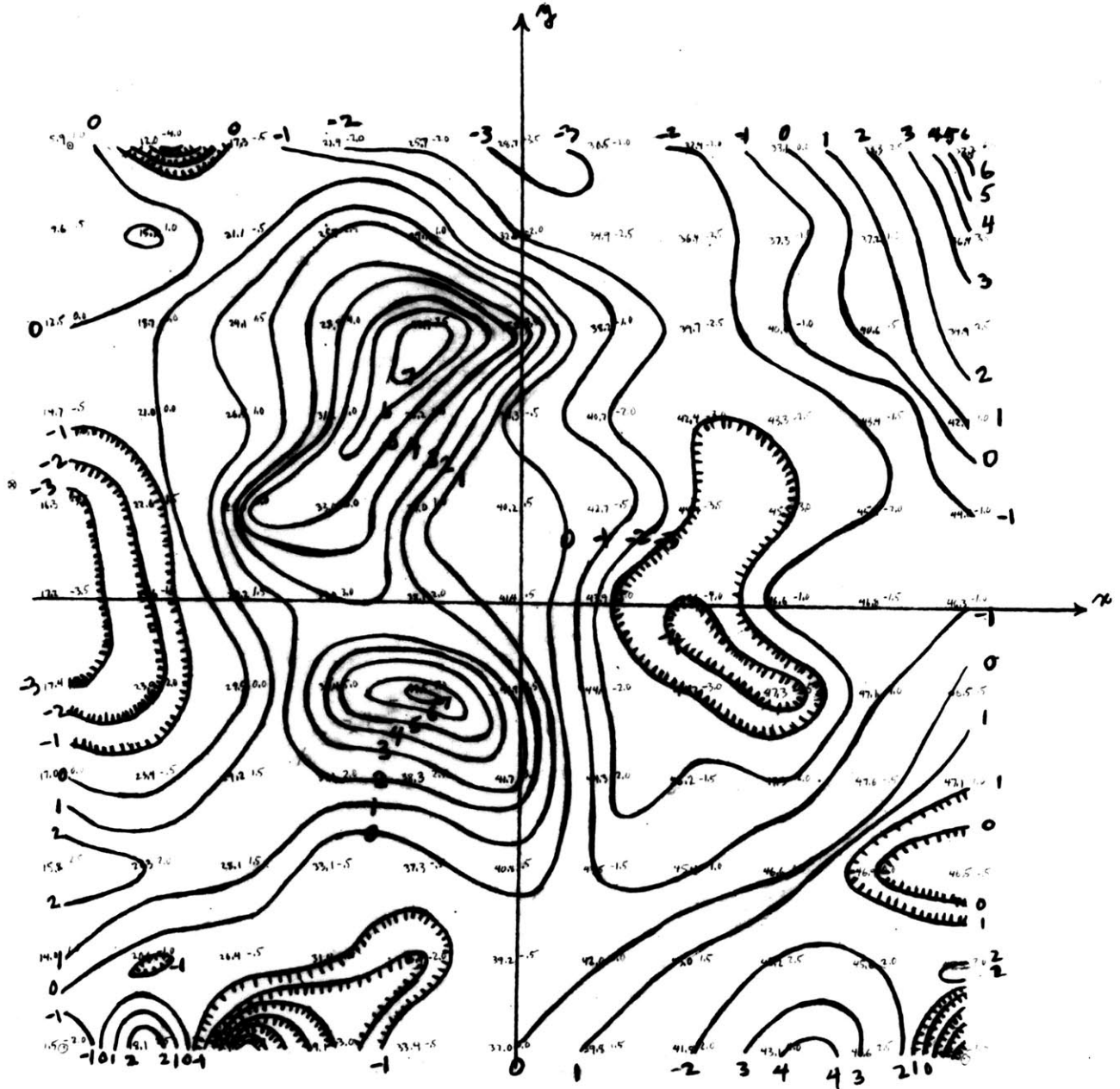


Fig. 1.6

$$G(x,y) = 41.37 - .98x - .391y^2 - .0034y^3 - .056xy + .011x^2y + .0263xy^2 + 1.949x - .385x^2 + .0389x^3$$

RESIDUAL GRAVITY WITH  
THIRD ORDER POLYNOMIAL

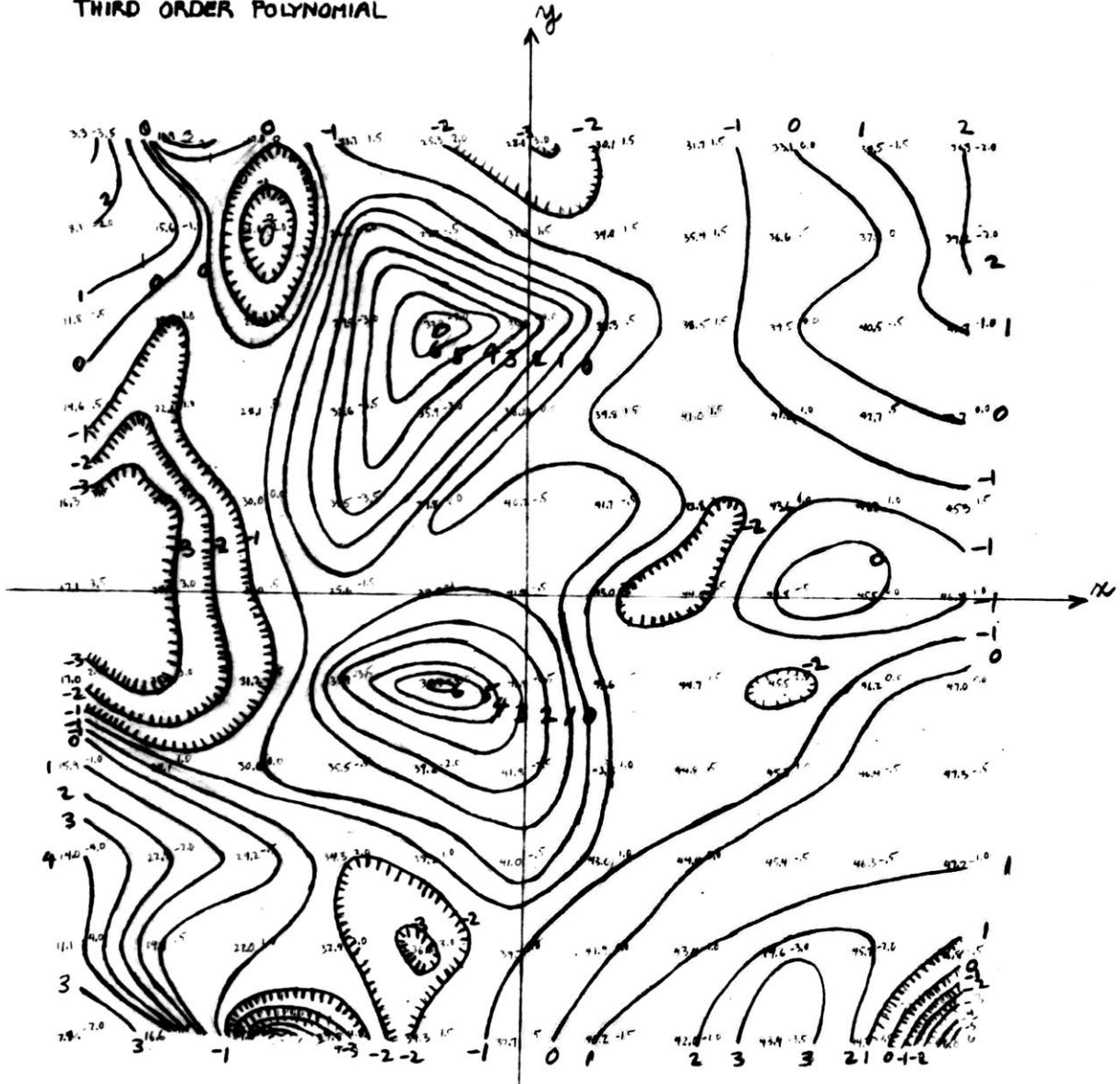


Fig. 1.7

# RESIDUAL GRAVITY WITH FOURTH ORDER POLYNOMIAL

$$G(x,y) = 42.04 - 892y - 251y^2 - 0.0338xy^3 - 0.00737y^4$$

$$- 245xy + 0.263xy^2 + 0.0015xy^3 + 0.0012xy^4 + 0.0918x^2y^2$$

$$+ 0.0947x^2y + 1.75x - 5.88x^2 + 0.389x^3 + 0.0444x^4$$

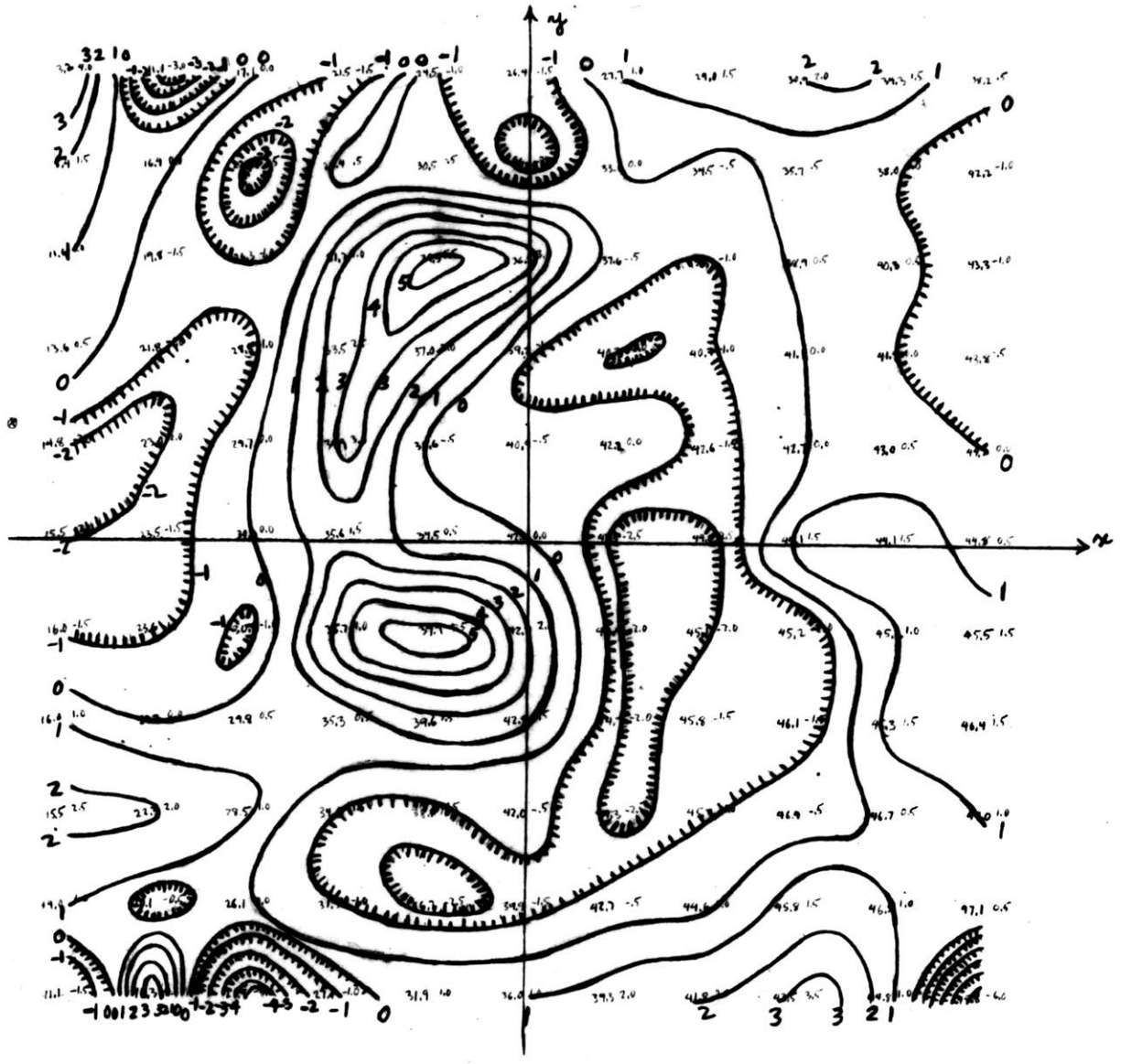


Fig. 1.8

# RESIDUAL GRAVITY COMPANY DATA

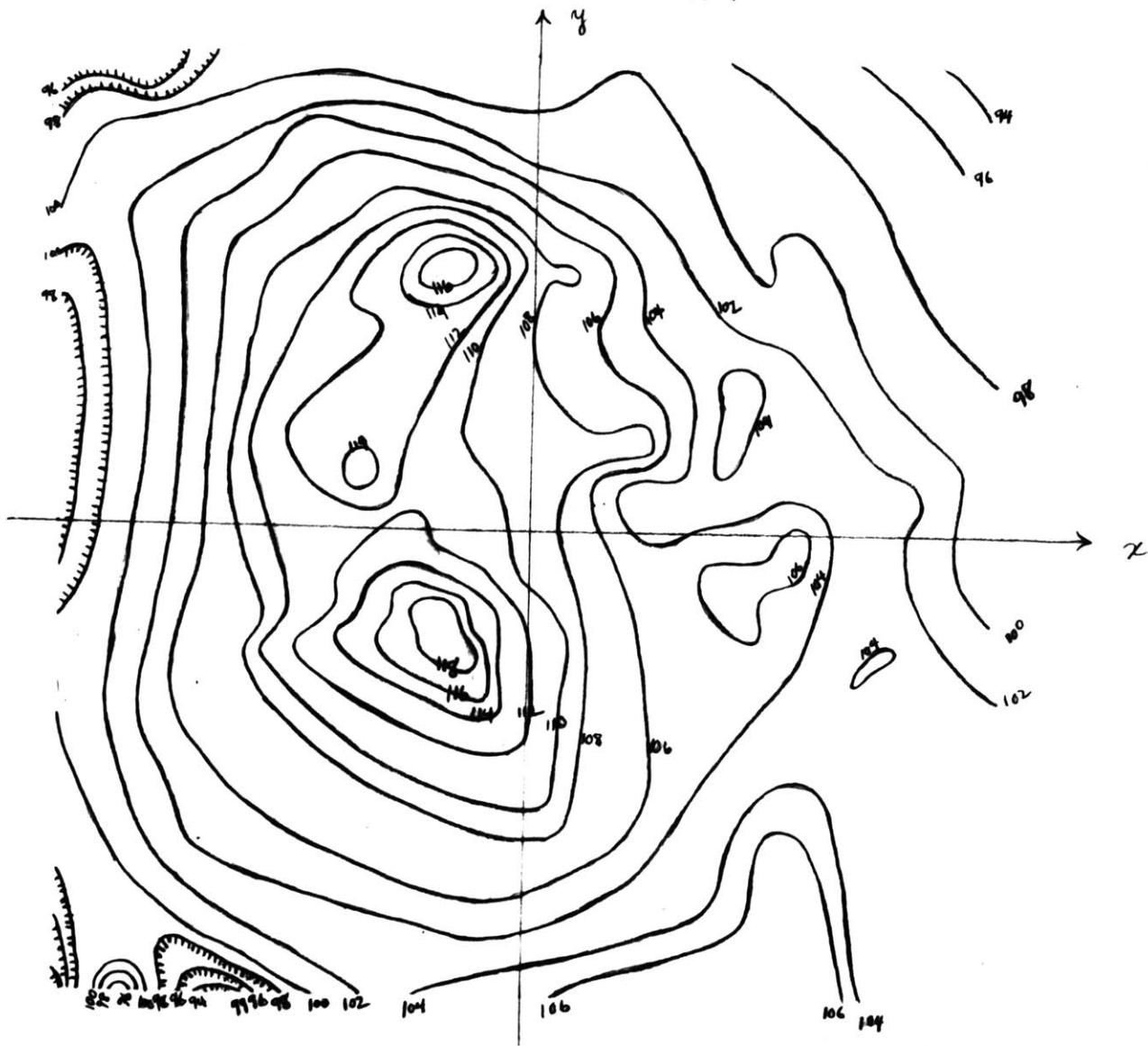


Fig. 1.9

# CONTOURS OF $G(x,y)$

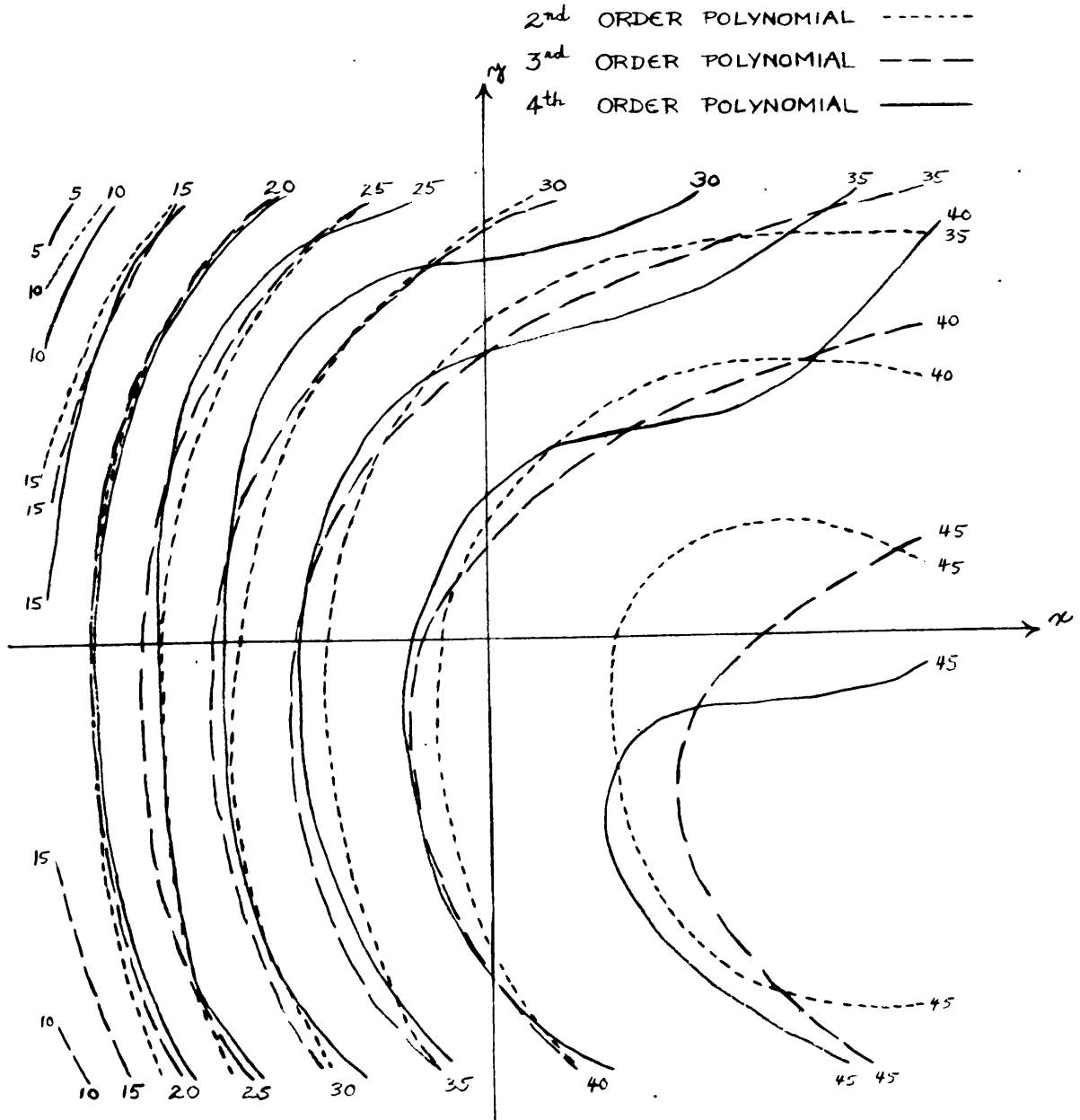


Fig. 1.10

# REGIONAL CORRECTION MINUS TOPOGRAPHIC CORRECTION

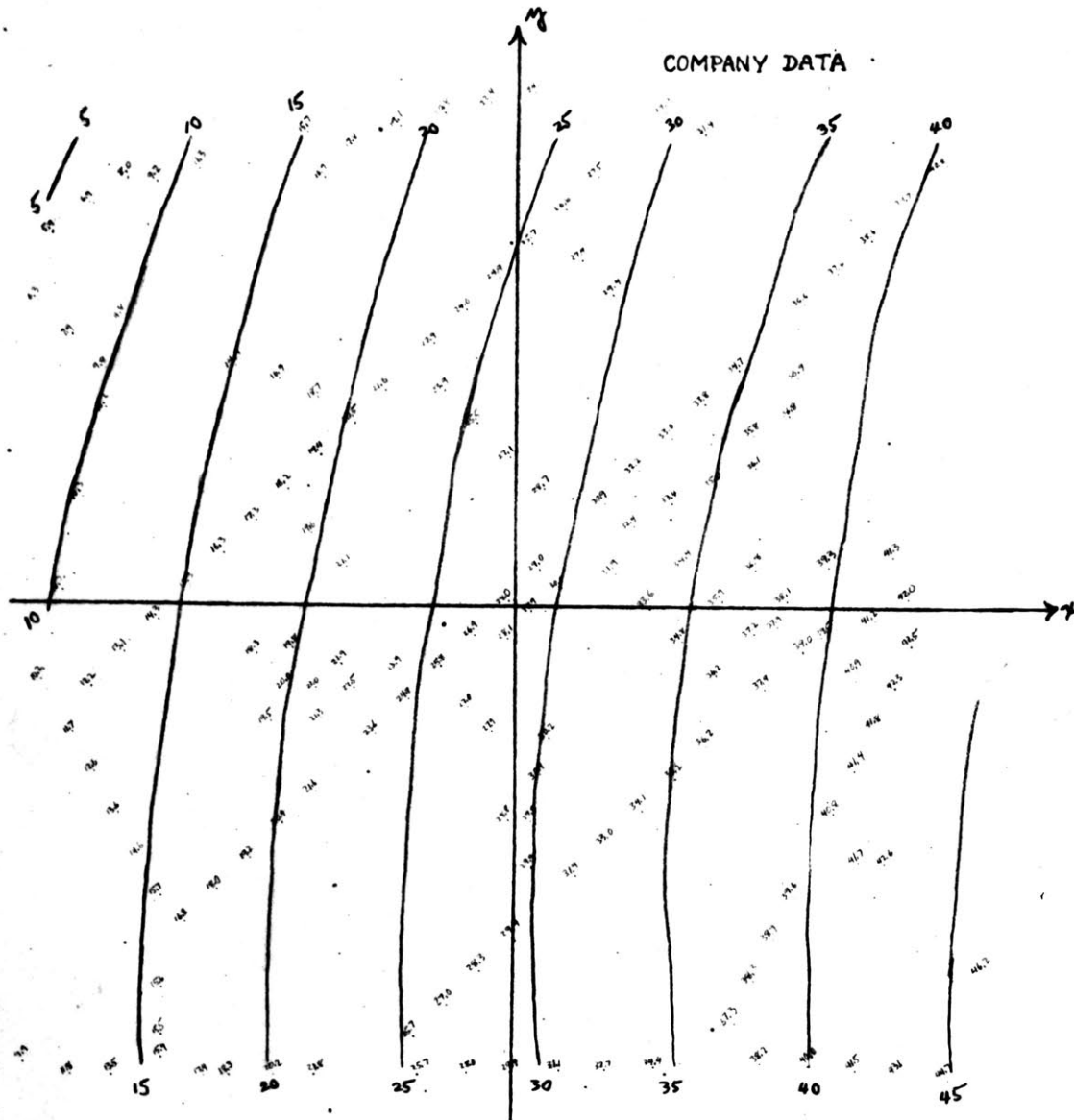


Fig. 1.11



### Setting Up the Normal Equations

We are concerned here with the problem of setting up the normal equations for various grids and polynomial degrees. If we limit ourselves to polynomials of degree 4 or less, we are then interested in finding the following quantities:

$$\begin{aligned} & \Sigma x^2 \quad \Sigma x^4 \quad \Sigma x^6 \quad \Sigma x^8 \quad \Sigma y^4 \quad \Sigma y^6 \quad \Sigma y^8 \\ & \Sigma x^2 y^2 \quad \Sigma x^2 y^4 \quad \Sigma x^2 y^6 \quad \Sigma x^4 y^2 \quad \Sigma x^4 y^4 \quad \Sigma x^6 y^2 \quad \Sigma x^0 y^0 \end{aligned} \quad 1.3$$

where the summations are to be taken over the particular grid we are dealing with.

If the grid has the dimensions  $2N$  by  $2M$  as shown in Fig. 1.2, we may set up a fairly simple procedure for finding these summations.

First we note that  $\Sigma x^k$  over the grid is equal to the  $\Sigma x^k$  on a single horizontal line, times the number of lines. Thus

$$\Sigma x^k = (2M + 1) \sum_{i=-N}^N i^k$$

but since in our case  $k$  is always even

$$\Sigma x^k = 2(2M + 1) \left( \sum_{i=1}^N i^k \right) \quad 1.4$$

Likewise

$$\Sigma y^k = 2(2N + 1) \sum_{i=1}^M i^k$$

For the cross terms  $\Sigma x^k y^l$  we have

$$\begin{aligned} \Sigma_{\text{grid}} x^k y^l &= \Sigma_{j=-M}^M \Sigma_{i=-N}^N i^k j^l \\ &= \begin{bmatrix} \Sigma i^k \\ -N \end{bmatrix} \begin{bmatrix} \Sigma j^l \\ -M \end{bmatrix} \\ \Sigma_{\text{grid}} x^k y^l &= 4 \left( \Sigma_{i=1}^N i^k \right) \left( \Sigma_{j=1}^M j^l \right) \end{aligned}$$

Hence the sums 1.3 are easily derivable from equations 1.4 and 1.5 if we tabulate the quantities  $\Sigma_{i=1}^L i^k$ . Table I gives values of  $i^k$  from which the sums  $\Sigma_{i=1}^L i^k$  are derived, and Table II tabulates these sums for L up to 25 and  $k = 2, 4, 6, 8$ . The latter Table allows us to compute the sums 1.3 for any grids measuring up to 50 by 50. a grid this size would encompass 2601 gravity readings which seems ample for most applications.

Table III contains the sums 1.3 computed for six representative grids 10 by 10, 10 by 20, 20 by 20, 30 by 30, 40 by 40, and 50 by 50.

TABLE I  $1^k$ for  $k = 2, 4, 6, 8,$   $1 = 1, 2, 3 \dots 25$ 

$1$	$1^2$	$1^4$	$1^6$	$1^8$
1	1	1	1	1
2	4	16	64	256
3	9	81	729	6561
4	16	256	4096	65536
5	25	625	15625	390625
6	36	1296	46656	1679616
7	49	2401	117649	5764801
8	64	4096	262144	16777216
9	81	6561	531441	43046721
10	100	10000	1000000	100000000
11	121	14641	1771561	214358881
12	144	20736	2985984	429981696
13	169	28561	4826809	815730721
14	196	38416	7529536	1475789056
15	225	50625	11390625	2562890625
16	256	65536	16777216	4294967296
17	289	83521	24137569	6975757441
18	324	104976	34012224	11019960576
19	361	130321	47045881	16983563041
20	400	160000	64000000	25600000000
21	441	194481	85766121	37822859361
22	484	234256	113379904	54875873536
23	529	279841	148035889	78310985281
24	576	331776	191102976	110075314176
25	625	390625	244140625	152587890625

TABLE II  $\sum_{i=1}^L i^k$

for  $k = 2, 4, 6, 8$   $L = 1, 2, 3, \dots, 25$

L	k=2	k=4	k=6	k=8
1	1	1	1	1
2	5	17	65	257
3	14	98	794	6818
4	30	354	4890	72354
5	55	979	20515	462979
6	91	2275	67171	2142595
7	140	4676	184820	7907396
8	204	8772	446964	24684612
9	285	15333	978405	67731333
10	385	25333	1978405	167731333
11	506	39974	3749966	382090214
12	650	60710	6735950	812071910
13	819	89271	11562759	1627802631
14	1015	127687	19092295	3103591687
15	1240	178312	30482920	5666482312
16	1496	243848	47260136	9961449608
17	1785	327369	71397705	16937207049
18	2109	432345	105409929	27957167625
19	2470	562666	152455810	44940730666
20	2870	722666	216455810	70540730666
21	3311	917147	302221931	108363590027
22	3795	1151403	415601835	163239463563
23	4324	1431244	563637724	241550448844
24	4900	1763020	754740700	351625763020
25	5525	2153645	998881325	504213653645

TABLE III  
 $\Sigma x^k y^l$

Over various grids measuring  $2N \times 2M$

for  $k = 0, 2, 4, 6, 8$  and  $l = 0, 2, \dots, (8-k)$

	N=5 M=5	N=5 M=10	N=10 M=10	N=15 M=15	N=20 M=20	N=25 M=25
$\Sigma x^0 y^0 =$	121	231	441	961	1681	2601
$\Sigma x^2 =$	1210	2310	16170	76880	235340	563550
$\Sigma x^4 =$	21538	41118	1063986	11055344	59258612	219671790
$\Sigma x^6 =$	451330	861630	83093010	1889941040	17749376420	101885895150
$\Sigma x^8 =$	10185538	19445118	7044715986	351321903344	5784339914612	51429792671790
$\Sigma y^2 =$	1210	8470	16170	76880	235340	563550
$\Sigma y^4 =$	21538	557326	1063986	11055344	59258612	219671790
$\Sigma y^6 =$	451330	43524910	83093010	1889941040	17749376420	101885895150
$\Sigma y^8 =$	10185538	3690089326	7044715986	351321903344	5784339914612	51429792671790
$\Sigma x^2 y^2 =$	12100	84700	592900	6150400	32947600	122102500
$\Sigma x^2 y^4 =$	215380	5573260	39012820	884427520	8296205680	47595554500
$\Sigma x^2 y^6 =$	4513300	435249100	3046743700	151195283200	2484912698800	22075277282500
$\Sigma x^4 y^2 =$	215380	1507660	39012820	884427520	8296205680	47595554500
$\Sigma x^6 y^2 =$	4513300	31593100	3046743700	151195283200	2484912698800	22075277282500
$\Sigma x^4 y^4 =$	3833764	99204028	2567043556	127180677376	2088984590224	18552747144100

PART II

## SEISMIC RECORD ANALYSIS BY LINEAR OPERATORS

### Introduction

In the study of reflection seismic records, taken in the exploration for oil, it is becoming increasingly difficult to pick reflection times by the standard procedures. The reason for this is that, as the simpler geologic areas are being fully exploited, exploration is being forced into the more complicated areas. Seismic records taken in these structurally complex areas contain much in the way of unwanted information and much not-understood information. Energy reflected from the strata of interest is largely masked by this "noise". At least two different approaches to unscrambling these records are being developed at present.

The first of these approaches is largely instrumental. Its principle is: take more and more information (more traces on each record, etc.), filter it in different ways and mix it up in a variety of combinations to see if a procedure for averaging out the unwanted information can be arrived at. This approach has led oil companies to the use of 24-trace records, each trace representing the responses from up to thirty geophones. The success of these methods is not publicly available, but the oil industry is expressing great interest in the approach described below, so probably they are not completely satisfactory.

The second of these approaches is basically analytic. Rather than taking more information, we attempt to sharpen up the interpretive procedure on the information we have.

The search for such procedure has been largely carried out at MIT in the Mathematics Department, and subsequently in the Mathematics and Geology Departments. The tools in this analysis have been the statistics of time series.

### Statistical Methods

After some experimentation it was found at MIT that the use of the "linear operator" seemed most promising in the determination of reflection times. The exact methods used are described in Refs. 5 and 6. The linear operator permits a measure of the change in dynamics as we proceed down a seismic record. As these dynamics are amplitude, frequency, and phase relationships, it was hoped that the dynamical change at a reflection could be discriminated even when the changes due to amplitude were small. This was hoped for since the usual interpretive procedures depend heavily on "amplitude reflections". The results were very encouraging and stimulated increased research.

One direction this study has taken is the empirical one. We know the linear operator gives us added information. But, since there is considerable freedom in the choice of the exact mathematical form of the operator we use, we try many different forms and see which ones give us the most information. This is a trial and error procedure and involves an immense amount of computation. For this reason a program was written for the WWI Digital Computer which would compute automatically the measure of dynamic change, for a great



variety of forms of linear operators, at very high speed. A copy of this program and a description of its functions is contained in Appendix B. Case studies designed to test the effects of individual parameters of the linear operator are being run with this program, but the results are as yet incomplete.

Along with this empirical approach an attempt is being made to study the linear operator from theoretical grounds. Although the form of the operator which is being tested at present is relatively complicated, it is instructive to consider a simpler form, the so-called "cosine operator". This operator is a mathematical expression which generates a pure cosine wave of given frequency. We can determine quite simply the effects of this type of operator on various time series including those found on seismic records. We hope to gain insight into the physical function of such operators as well as correspondence between them and simple filters.

We shall also consider two other more practical problems connected with the statistical analysis of seismograms by the use of linear operators. One concerns certain iterative methods for approaching the values of the linear operator coefficients for least squares fitting. The other is a related problem, the necessity for accuracy in finding these values.

## Single Frequency Cosine Operators

A cosine operator is a prediction mechanism which exactly predicts equally spaced points on a cosine wave. It has the general form †

$$\hat{x}_{1+2} = c + ax_{1+1} + bx_1 \quad 2.1$$

$$\text{where } a = 2 \cos 2\pi hf \equiv 2u$$

$$b = -1$$

$$h = \text{time between observation}$$

$$c = (1-a-b)\bar{x} = 2(1-u)\bar{x}$$

$$\bar{x} = \text{mean of series}$$

$$f = \text{frequency of cosine wave}$$

$$\hat{x}_{1+2} = \text{predicted value of } x_{1+2}$$

Suppose we use this operator to predict an arbitrary series  $x$ . Then the error of prediction  $x_{1+2} - \hat{x}_{1+2}$  will be

$$\begin{aligned} E_{1+2} &= x_{1+2} - [2(1-u)\bar{x} + 2ux_{1+1} - x_1] \\ &= x_{1+2} - 2(1-u)\bar{x} - 2ux_{1+1} + x_1 \\ &= (x_{1+2} - \bar{x}) + (x_1 - \bar{x}) - 2u(x_{1+1} - \bar{x}) \end{aligned} \quad 2.2$$

For simplicity let us deal with a series  $X_1$  measured around its equilibrium mean  $\bar{x}$ , i.e.  $X_1 = x_1 - \bar{x}$  then 2.2 becomes

$$E_{1+2} = X_{1+2} + X_1 - 2uX_{1+1} \quad 2.3$$

Now if we sum the squares of these errors over an interval of the series we get

† Ref. 6

$$\sum_1 E_{1+2}^2 = \sum (X_{1+2}^2 + X_1^2 + 4u^2 X_{1+1}^2 + 2X_{1+2}X_1 - 4uX_{1+2}X_{1+1} - 4uX_1X_{1+1}) \quad 2.4$$

If the series is stationary and the interval sufficiently great we may write this in terms of the auto-correlations.

Let the series be normalized so that  $\sum_1 X_1^2 = 1$  then

$$\sum_1 E_{1+2}^2 = R_0 + R_0 + 4u^2 R_0 + 2R_2 - 8uR_1 \quad 2.5$$

where  $R_1 = 1^{\text{th}}$  lag auto-correlation and  $R_0 = 1$

or

$$\sum E_{1+2}^2 = 2(1 + R_2 - 4uR_1 + 2u^2) \quad 2.6$$

This expression has a minimum value when

$$u = \frac{-(-4R_1)}{2 \times 2} = R_1 \quad 2.7$$

or

$$\cos 2\pi hf = R_1$$

$$f = \frac{1}{2\pi h} [\cos^{-1} R_1 \pm 2n\pi] \quad 2.8$$

Hence we have the least squares fit for a cosine operator predicting an arbitrary stationary series. We find that  $f$  is determined only by the first lag of the auto-correlation function of the series, and that  $f$  is only determined modulo  $1/h$ . This last fact is apparent if we refer back to equation 2.1, where we see that cosine operators have identical forms for angular frequencies differing by  $1/h$ .

Hence

$$\text{Min } \sum_1 E_{1+2}^2 = 2(1 + R_2 - 2R_1^2) \quad 2.9$$

If we want a perfect least squares fit we have

$$R_1 = \pm \sqrt{\frac{1}{2} (1 + R_2)} \quad 2.10$$

with the restriction  $\cosine \ 2\pi hf = R_1$

One way to meet this condition is to let the interval  $h$  shrink toward zero so that  $R_1 \rightarrow 1$  and  $R_2 \rightarrow 1$ . This is equivalent to saying that any small segment of the original series approaches a straight line, in the case where the function is continuous and its first derivative exists.

### The Geometry of Cosine Operators

#### A. Error Sum as Function of $W$

Consider the sum of squared errors as a function of  $u$ . We have

$$\sum_1 E_{1+2}^2 = 2(1 + R_2 - 4uR_1 + 2u^2) \quad 2.6$$

This is a parabola in  $u$  as shown in Fig. 2.1.

$u = \cos \ 2\pi hf$  must lie in the range  $-1 \leq u \leq 1$ .

We have shown that  $u = R_1$  is the condition for a minimum fit, and since  $-1 < R_1 < +1$ ,  $\sum E^2$  will always have its minimum in this range.

This means that, for any series, we can always get a minimum fit with some cosine operator of frequency  $f$ , where  $f$  must be in the range

Fig. 2.1

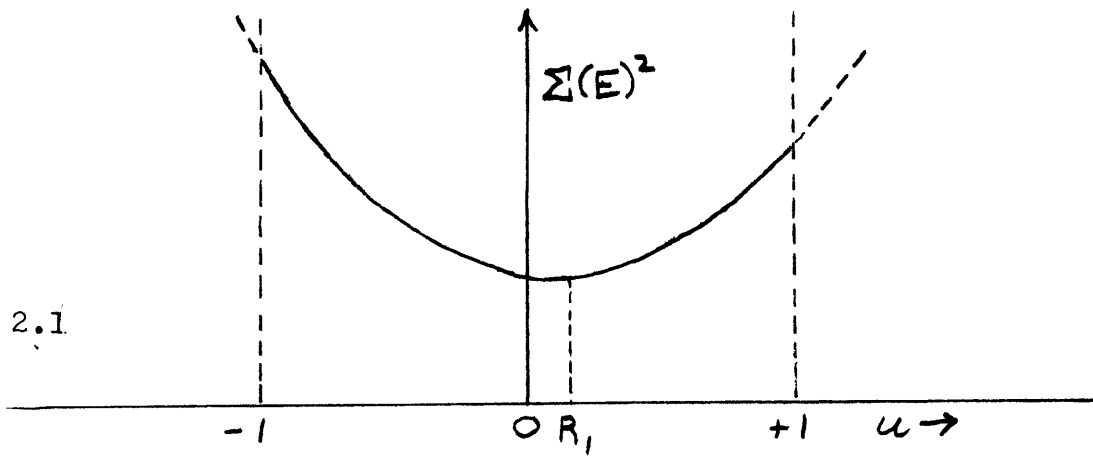


Fig. 2.2

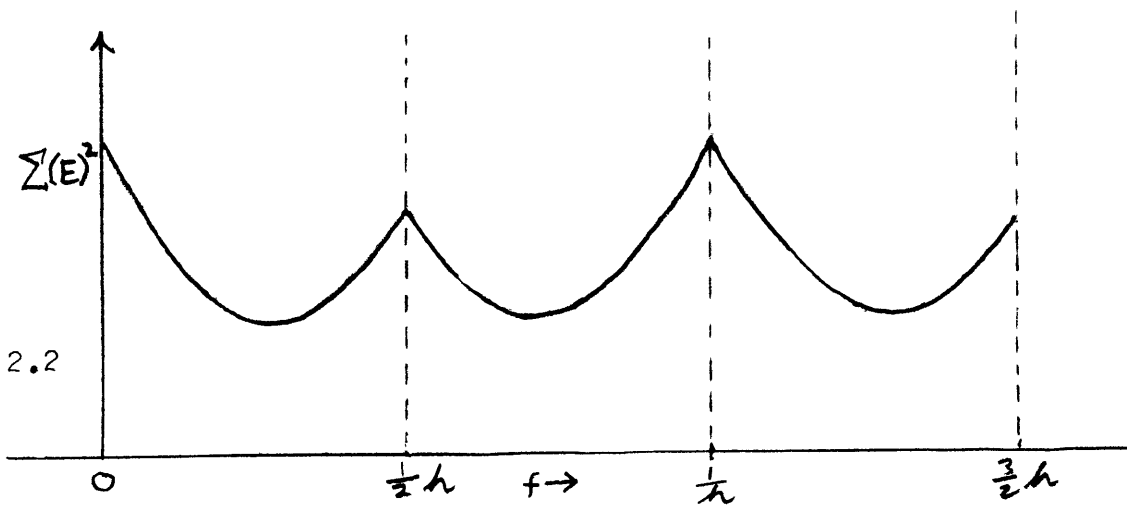
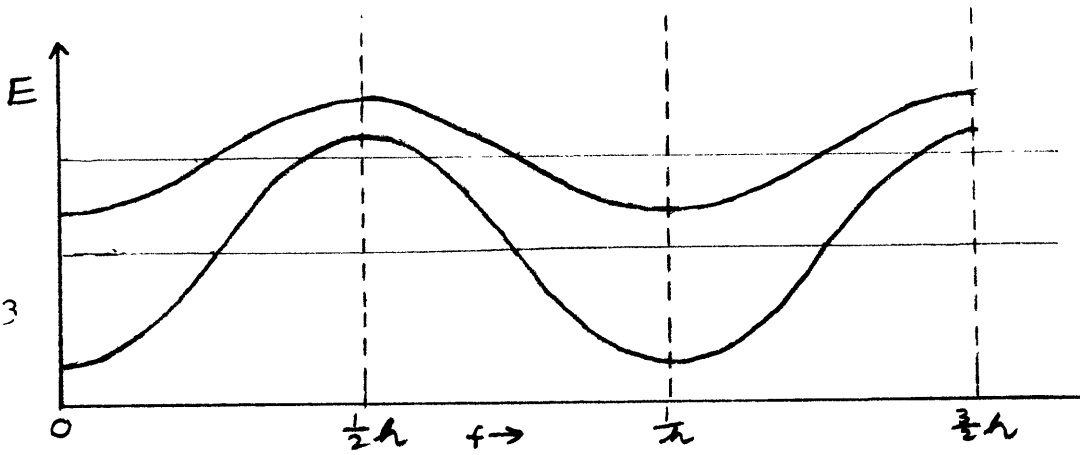


Fig. 2.3



$$\begin{array}{c}
 0 < f < \frac{1}{2h} \\
 \uparrow \qquad \qquad \searrow \\
 ( u = 1 \qquad ) \quad ( u = -1 \qquad ) \\
 ( 2\pi hf = 0 \quad ) \quad ( 2\pi hf = \pi \quad )
 \end{array}
 \qquad 2.11$$

We require  $\Sigma E^2$  to be non-negative. This means the discriminant of 2.9 must be  $\leq 0$  or

$$16 R_1^2 - 8(1+R_2) \leq 0 \qquad 2.12$$

Therefore the curve cannot cross the u axis, but can be tangent to it at one point, when the equality sign holds above. This is the condition for a perfect fit.

#### B. Error Sum as Function of f

The error sum as a function of frequency f is not truly parabolic but has the general shape of a parabola. It is periodic in f with a period 1/h. It appears as shown in Fig. 2.2 .

#### C. Individual Errors as Functions of u and f

Equation 2.3 gives us an expression for the individual errors

$$E_{i+2} = X_{i+2} + X_i - 2uX_{i+1} \qquad 2.3$$

If we fix attention on a single individual error (i constant) and let u vary we see that  $E_{i+2}$  is sinusoidal since  $u = \cos 2\pi hf$  . Thus  $E_{i+2}$  varies sinusoidally about a mean given by the sum of the  $i^{\text{th}}$  and the  $(i-2)^{\text{nd}}$  value of the series, with an amplitude of twice the  $(i-1)^{\text{rst}}$  value of the series. The period is 1/h. There is no phase shift

between these curves for different  $i$  values. Thus all individual errors must increase or decrease simultaneously with  $f$ .

Fig. 2.3 shows individual errors as functions of  $f$ . This figure explains why the error sum of Fig. 2.2 is reflected across the line  $f = 1/2h$ . This line in Fig. 2.3 is the axis of symmetry for the individual errors, so that it must also be the symmetry axis for the error sum.

### Conclusions

From the above, we can draw certain conclusions.

1. If one limits himself to the general class of cosine operators, there is a maximum error obtainable for the particular data, using any frequency whatever. That is, there is such a thing as a worst fit for cosine operators.
2. Since  $\Sigma E^2$  is parabolic, determining 3 values of  $\Sigma E^2$  is sufficient to determine the complete shape of the error curve for all other frequencies.
3. Moreover, since the individual errors are sinusoidal in  $f$ , determining the individual errors for 3 values of  $f$  determines the errors for all  $f$ .

Looking at the problem another way, much of the information obtainable from any data series by a study of

this type is contained in the first and second lags of the auto-correlation function of the series, for these two quantities determine the shape and position of the curve  $\Sigma E^2$ .

### Example

The prediction program described in Appendix B provided a means for testing the conclusions reached about cosine operators. Individual errors and sums of squared errors were computed for cosine operators of frequencies 25, 30, .... 75 cps. The data for which these were computed were readings taken from a typical seismic trace at intervals of 2 ms.

The sums of squared errors are plotted in Fig. 2.4 over two intervals of 240 readings each. Both curves exhibit very good parabolic shapes. The average minimum for the two curves occurs for  $u = .85$ . This should equal the first lag auto-correlation over the two intervals, which was computed by the correlation program (Appendix D) to be .853.

Fig. 2.4 shows several individual errors plotted as functions of the frequency of the cosine operator used. They appear to be sections of sinusoids as expected.

These curves, computed on an arbitrary time series, seem to be in remarkable agreement with the theory.



Fig. 2.4

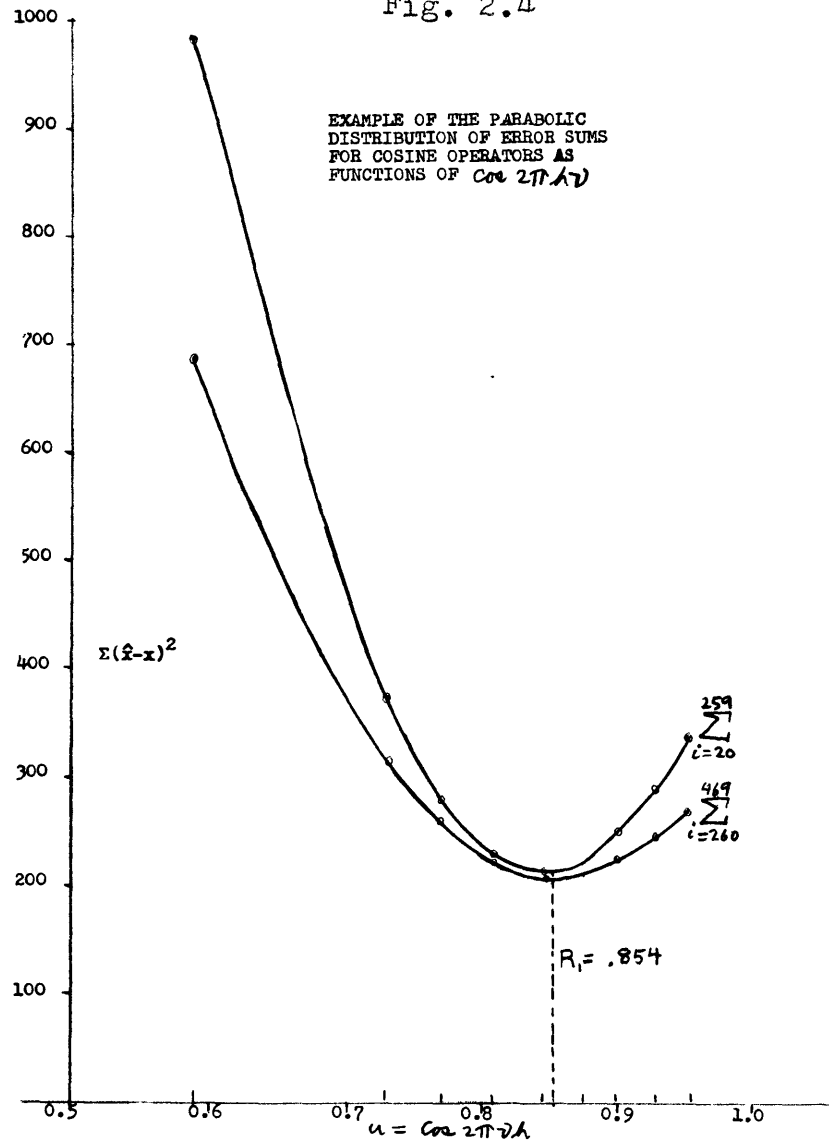
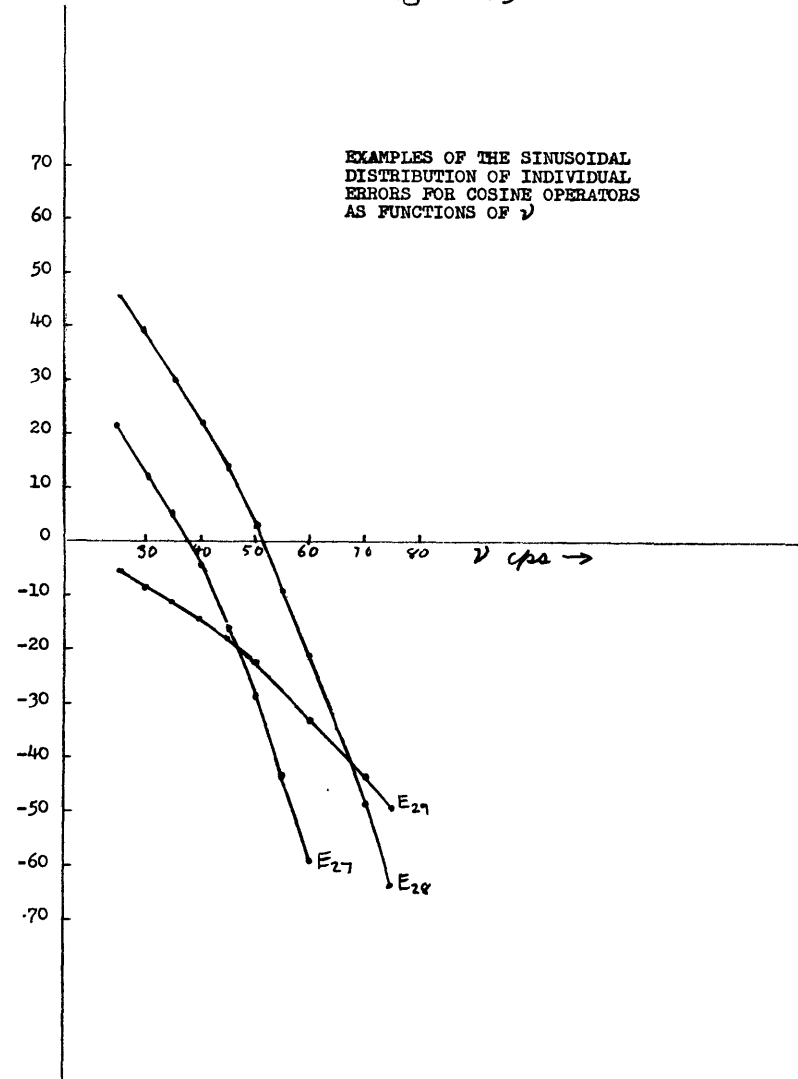


Fig. 2.5



## The Cyclical Nature of Cosine Operators

We have mentioned that cosine operators differing in frequency by  $n/h$  have identical forms. It is interesting to see what this means physically.

Suppose we are trying to represent a cosine wave of 1 cycle/second with a spacing of  $h = 1/4$  second. The points we would plot might appear as in Fig. 2.6 .

Now consider a cosine wave of frequency  $1 + 1/h = 5$  cycles/sec. If we try to plot this frequency with a spacing of  $1/4$  sec. we find that it can be exactly represented by the points we plotted for the one cycle wave. This is illustrated in Fig. 2.7 . We would find the same would be true for frequencies of  $1 + n/h = 1, 5, 9, 13, 17 \dots$ . Thus, it is the fact that we cannot uniquely represent frequencies differing by  $n/h$  that explains the identity of form for cosine operators whose frequencies differ by this amount.

This is also the explanation for the so-called "condensed" power spectra met with in computational procedures.† The computed power at a frequency  $f$  must represent the sum of the powers at frequencies  $f, f, + 1/h, f + 2/h \dots$ . Therefore power spectra can only have the range 0 to  $1/h$  cycles. In practice  $h$  must be chosen so that  $1/h$  is greater than the greatest frequency from which significant contribution is expected.

† Ref. 3

Fig. 2.6

COSINE WAVE 1 CYCLE/SEC.

PLOTTING INTERVAL  $h = \frac{1}{4}$  SEC.

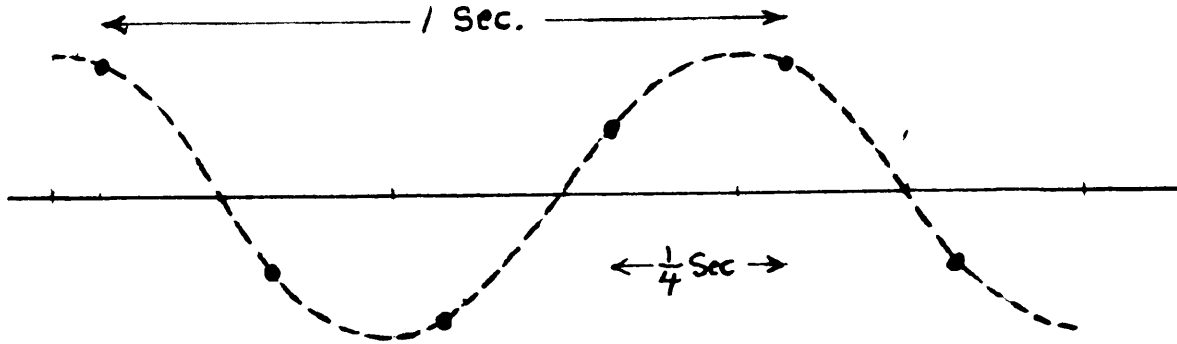
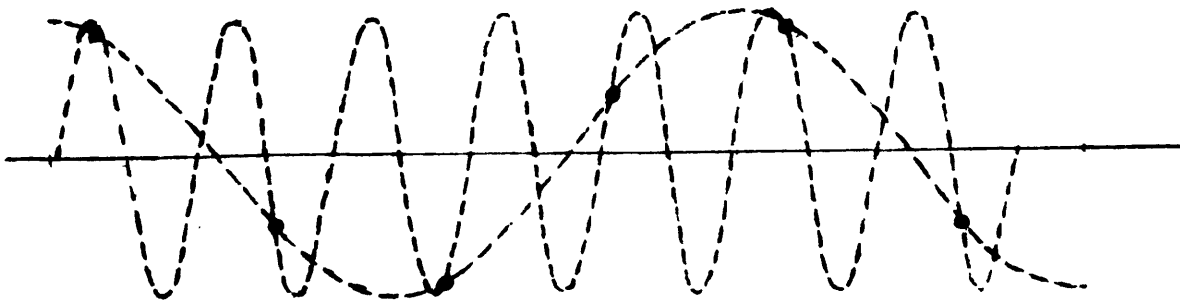


Fig. 2.7

COSINE WAVES 1 AND 5 CYCLES/SEC.



### Cosine Operator Predicting an Autoregressive Series

In order to examine the filter characteristics of cosine operators it is convenient to consider their effect on autoregressive-type series. The autoregressive series has a known Cauchy-type distribution of its spectrum.† It will be interesting to examine what the spectrum of the error function will be when we predict such a series with a cosine operator.

Referring back to equation 2.3 we have the error function for cosine operators.

$$E_{1+2} = X_{1+2} + X_1 - 2uX_{1+1} \quad 2.3$$

To get the spectrum of this function we first find the auto-correlation  $R_\tau$ .

$$R_\tau = \Sigma (X_{1+2} + X_1 - 2uX_{1+1})(X_{1+2-\tau} + X_{1-\tau} - 2uX_{1+1-\tau})$$

$$\begin{aligned} R_\tau = & \Sigma X_{1+2}X_{1+2-\tau} + \Sigma X_{1+2}X_{1-\tau} + \Sigma X_1X_{1+2-\tau} \\ & + \Sigma X_1X_{1-\tau} - 2u(\Sigma X_{1+1}X_{1+2-\tau} + \Sigma X_{1+1}X_{1-\tau} \\ & + \Sigma X_{1+1-\tau}X_{1+2} + \Sigma X_{1+1-\tau}X_1) + 4u^2\Sigma X_{1+1}X_{1+1-\tau} \end{aligned}$$

If the X series is properly normalized we may write this in terms of the correlations  $r_\tau$  of the X series.

$$\begin{aligned} R_\tau = & r_\tau + r_{\tau+2} + r_{\tau-2} + 4u^2r_\tau \\ & + r_\tau - 2u(r_{\tau-1} + r_{\tau+1} + r_{\tau+1} + r_{\tau-1}) \end{aligned}$$

† Ref. 7

$$\begin{aligned}
 R_{\tau} = & r_{\tau-2} + r_{\tau-1}(-4u) + r_{\tau}(2+4u^2) \\
 & + r_{\tau+1}(-4u) + r_{\tau+2}
 \end{aligned}
 \tag{2.13}$$

Since the series is taken to be autoregressive

$$r_{\tau} = \cos 2\pi f_0 \tau h e^{-a\tau h} \equiv \cos a\tau e^{-b\tau} \quad \ddagger \quad \tau \geq 0
 \tag{2.14}$$

Substituting equation 2.14 into 2.13 we have

$$\begin{aligned}
 R_{\tau} = & \cos[a(\tau-2)]e^{-b(\tau-2)} \\
 & - 4u\cos[a(\tau-1)]e^{-b(\tau-1)} \\
 & + (2+4u^2)\cos a\tau e^{-b\tau} \\
 & - 4u\cos[a(\tau+1)]e^{-b(\tau+1)} \\
 & + \cos[a(\tau+2)]e^{-b(\tau+2)}
 \end{aligned}$$

Using trigonometric identities

$$\begin{aligned}
 R_{\tau} = & e^{2b}[\cos a\tau \cos 2a + \sin a\tau \sin 2a]e^{-b\tau} \\
 & - 4ue^b[\cos a\tau \cos a + \sin a\tau \sin a]e^{-b\tau} \\
 & + (2+4u^2)(\cos a\tau)e^{-b\tau} \\
 & - 4ue^{-b}[\cos a\tau \cos a - \sin a\tau \sin a]e^{-b\tau} \\
 & + e^{-2b}[\cos a\tau \cos 2a - \sin a\tau \sin 2a]e^{-b\tau}
 \end{aligned}$$

or

$$\begin{aligned}
 R_{\tau} = & \cos a\tau e^{-b\tau} [e^{2b} \cos 2a - 4ue^b \cos a + 2 + 4u^2 \\
 & - 4ue^{-b} \cos a + e^{-2b} \cos 2a]
 \end{aligned}$$

‡ Ref: 5

$$+ \sin a \tau e^{-b\tau} [e^{2b} \sin 2a - 4ue^b \sin a + 4ue^{-b} \sin a - e^{-2b} \sin 2a]$$

Hence

$$\begin{aligned} R_{\tau} &= A \cos a \tau e^{-b\tau} + B \sin a \tau e^{-b\tau} \\ &= e^{-b\tau} (A \cos a \tau + B \sin a \tau) \end{aligned} \quad 2.15$$

where

$$\begin{aligned} A &= \cos 2a (e^{2b} + e^{-2b}) - 4u \cos a (e^b + e^{-b}) + 2 + 4u^2 \\ B &= \sin 2a (e^{2b} - e^{-2b}) - 4u \sin a (e^b - e^{-b}) \end{aligned} \quad 2.16$$

$$\begin{aligned} R_{\tau} &= e^{-b\tau} (A^2 + B^2)^{1/2} \left[ \frac{A}{(A^2 + B^2)^{1/2}} \cos a \tau + \frac{B}{(A^2 + B^2)^{1/2}} \sin a \tau \right] \\ &= e^{-b\tau} (A^2 + B^2)^{1/2} (\cos a \tau \cos \beta + \sin a \tau \sin \beta) \\ &\quad \text{where } \beta = \tan^{-1} \frac{B}{A} \end{aligned}$$

Thus

$$R_{\tau} = e^{-b\tau} (A^2 + B^2)^{1/2} \cos(a\tau + \beta) \quad 2.17$$

$R_{\tau}$  is now in a form similar to the  $r_{\tau}$  for the original series. The spectrum of this type is known to be a Cauchy distribution. The specific shape will be controlled by values of  $b$ ,  $A$ ,  $B$ , and  $\beta$ .

Rather than continuing with this example we shall proceed to another type of series. The autoregressive series is somewhat non-typical. Its spectrum, the Cauchy distribution, is very broad, in fact there is no mean value of frequency for this spectrum.

Cosine Operators on Series with Gaussian Spectrum Distribution

A much more stringent series than the autoregressive type is a series with a power spectrum composed of two Gaussian curves. The spectrum has the form †

$$\phi(\omega) = \frac{1}{2\sigma\sqrt{2\pi}} \left[ e^{-\frac{(\omega-a)^2}{\sigma^2}} + e^{-\frac{(\omega+a)^2}{\sigma^2}} \right] \quad 2.18$$

where +a and -a are the respective means of the two Gaussian curves, and  $\sigma$  is their standard deviation in radians. With such a series the normalized auto-correlation function may be written as †

$$r_\tau = e^{-\frac{\sigma^2 \tau^2}{2}} \cos a$$

If we predict such a series with a cosine operator, we generate an error series whose auto-correlation function is, as before

$$R_\tau = r_{\tau-2} + r_{\tau-1}(-4u) + r_\tau(2+4u^2) + r_{\tau+1}(-4u) + r_{\tau+2} \quad 2.13$$

or substituting

$$R_\tau = e^{-\frac{\sigma^2(\tau-2)^2}{2}} \cos[a(\tau-2)] + e^{-\frac{\sigma^2(\tau-1)^2}{2}} \cos[a(\tau-1)](-4u) + e^{-\frac{\sigma^2(\tau)^2}{2}} \cos[a\tau](2+4u^2) + e^{-\frac{\sigma^2(\tau+1)^2}{2}} \cos[a(\tau+1)](-4u) + e^{-\frac{\sigma^2(\tau+2)^2}{2}} \cos[a(\tau+2)] \quad 2.20$$

† Ref. 5

This can be reduced, as before, to the form

$$R_{\tau} = e^{-\frac{\sigma^2 \tau^2}{2}} [A(\tau) \cos a\tau + B(\tau) \sin a\tau] \quad 2.21$$

where

$$\begin{aligned} A(\tau) = & \cos 2a \left[ e^{-\frac{\sigma^2(-4\tau+4)}{2}} + e^{-\frac{\sigma^2(4\tau+4)}{2}} \right] \\ & - 4u \cos a \left[ e^{-\frac{\sigma^2(-2\tau+1)}{2}} + e^{-\frac{\sigma^2(2\tau+1)}{2}} \right] \\ & + (2 + 4u^2) \end{aligned} \quad 2.22$$

$$\begin{aligned} B(\tau) = & \sin 2a \left[ e^{-\frac{\sigma^2(-4\tau+4)}{2}} + e^{-\frac{\sigma^2(4\tau+4)}{2}} \right] \\ & - 4u \sin a \left[ e^{-\frac{\sigma^2(-2\tau+1)}{2}} + e^{-\frac{\sigma^2(2\tau+1)}{2}} \right] \end{aligned}$$

If we are interested in the power spectrum of this series we want

$$\phi(\omega) = 2 \int_0^{\infty} R(\tau) \cos \omega\tau \, d\tau \quad 2.23$$

Probably this integral cannot be expressed in closed form, and we shall have to resort to a computed example.

#### Computational Example

Here we illustrate the filter characteristics of cosine operators in a particular case. We choose a series with a Gaussian spectrum peaked at 50 cycles and with a standard deviation of 22.36 cycles. The power spectrum of such a series is shown in Fig. 2.8, and was computed from equation 2.18. In general shape this is not unlike power spectra dealt with on seismic traces. Fig. 2.9 shows



the normalized auto-correlation function for this type of series, as derived from equation 2.19 . The series has essentially no correlation for lags greater than about .03 sec.

To examine the effectiveness of cosine operators as frequency filtering mechanisms, a cosine operator of frequency 50 cps was taken. The spacing interval was chosen to be 2.5 ms. The auto-correlation function of the error series generated by this operator is shown in Fig. 2.9 , and is computed from equation 2.13 . In this case the function is unnormalized so that the zeroth lag auto-correlation is proportional to the total power contained in the power spectrum of the error series. Thus we see that less than 20 per cent of the power contained in the original Gaussian series remains in the error series. More than 80 per cent has been "filtered" out. However, since some of this is due merely to curve continuity, the shape of the spectrum of errors is more important than the total power.

The unnormalized spectrum of the error series is shown in Fig. 2.8 , and, as might be expected, is definitely bimodal. This curve clearly indicates that the operator is acting as a filter peaked at 50 cps, at which frequency all power has been removed. Lower frequencies are also well reduced but the higher ones are not so much affected. In fact, the power at 100 cycles is slightly greater than

in the original series. This is not a computational error. As discussed below it seems to be a necessary characteristic.

A more convenient way of showing the filter characteristics is to plot the quantity

$$\frac{\text{Power removed at } \omega}{\text{Initial power at } \omega}$$

This graph is shown in Fig. 2.10 . It shows how frequencies lower than 50 cycles are much preferred to those greater. It is possible that this curve would not represent the filter characteristics of a 50 cycle cosine operator used on another type of series. There is some reason, however, to suspect that it does, and that, in fact, the curve of Fig. 2.10 continues downward considerably below the axis (thus representing amplification rather than filtration). If we were to use a series containing mostly frequencies between 100 and 200 cycles, the 50 cycle operator would yield very high errors of prediction. The sum of squared errors would be far from the minimum of Fig. 2.1 . Hence the power in the error series would probably be greater than that in the original series. This could only come about by an amplification of certain frequencies, which would naturally occur for frequencies greatly different from 50 cycles. In this example 200 cycles is chosen as an upper limit, because with a spacing of 2.5ms unique curves only exist from 0 cycles to  $1/2h$  or 200 cycles.

Fig. 2.8

POWER SPECTRA

- For series with Gaussian spectrum with  $\sigma = 22.4$  cps  $a = 50$  cps (normalised).
- - - For error series generated by a cosine operator of frequency 50 cps on above series (unnormalized).

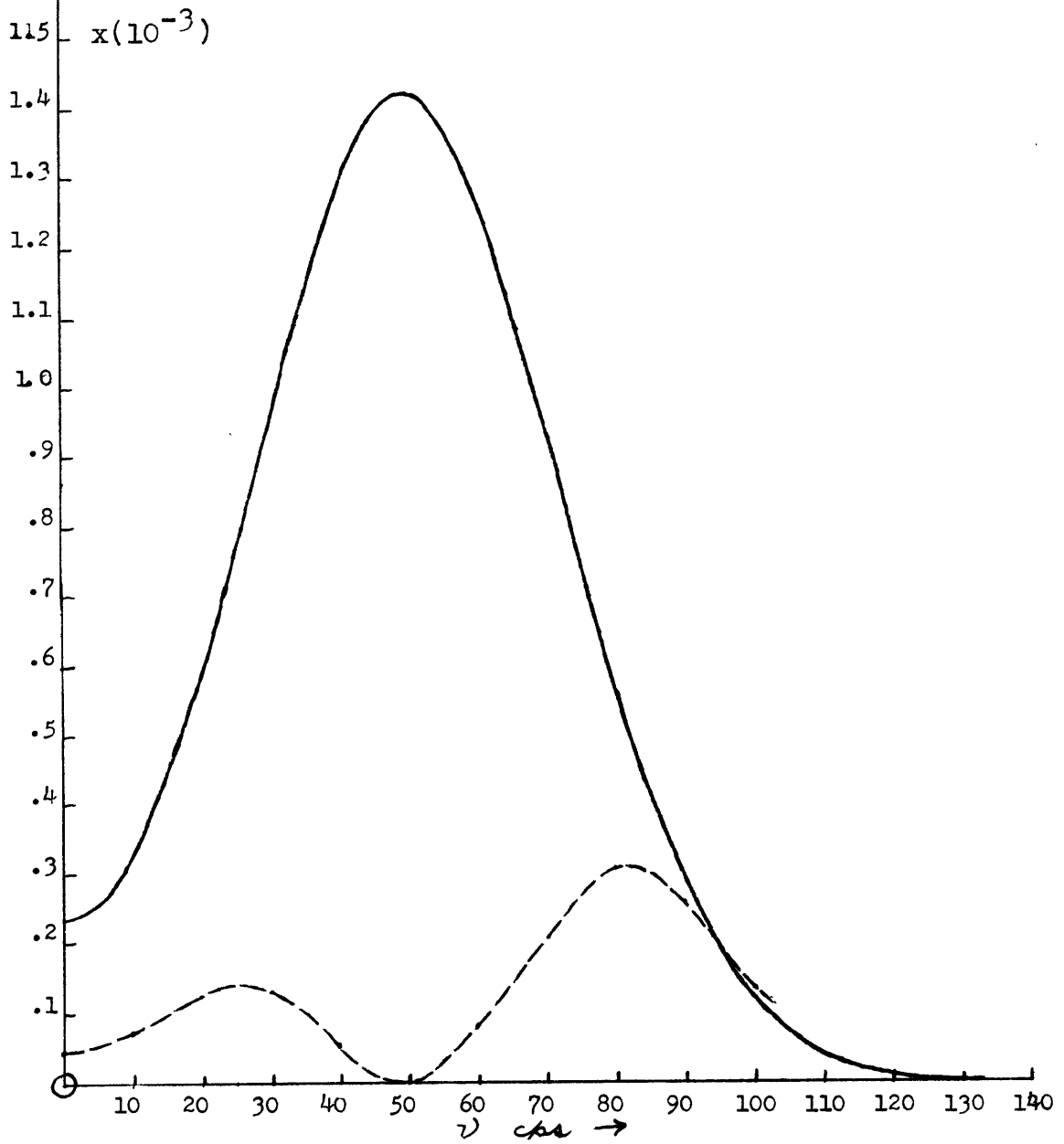


Fig. 2.9

AUTOCORRELATIONS  $R(ih)$

— For series with Gaussian spectrum  
with  $\sigma = 224$  cps  $a = 50$  cps (normalized).

----- For error series generated by a  
cosine operator of frequency 50 cps  
on above series (unnormalized).

$$h = 2.5 \cdot 10^{-3} \text{ sec.}$$

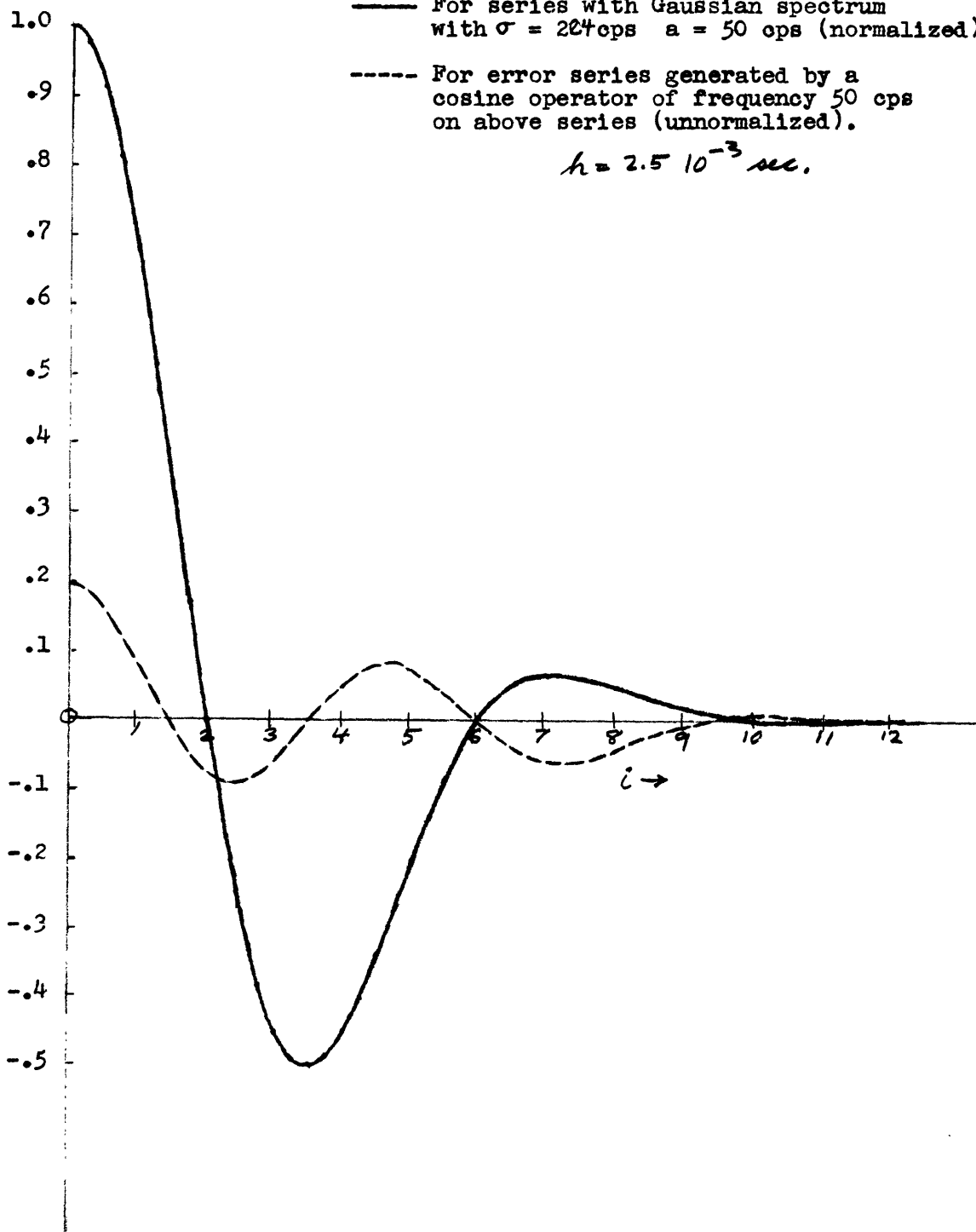
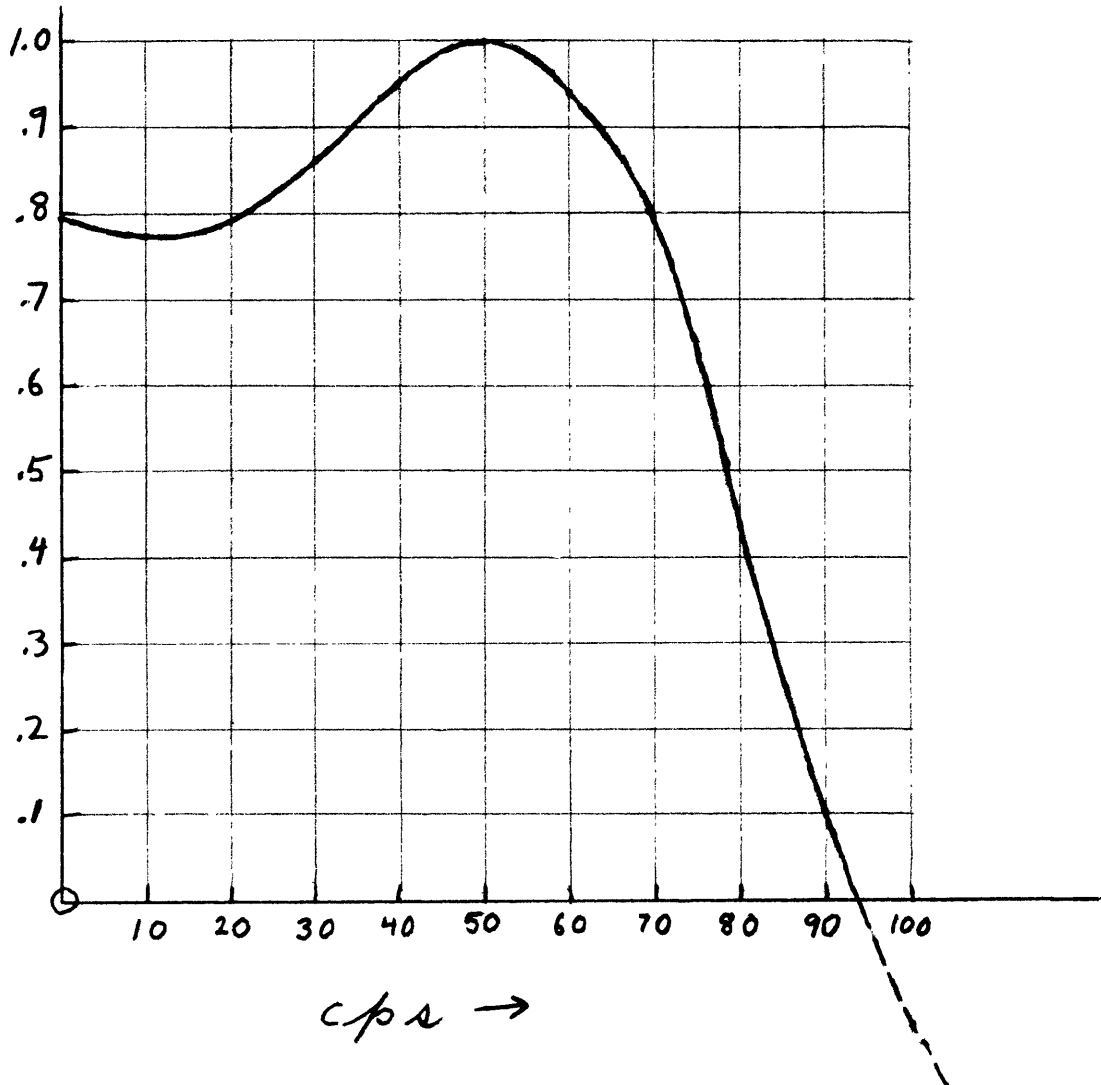


Fig. 2.10

FILTER CHARACTERISTICS  
FOR 50 CYCLE COSINE OPERATOR  
WITH  $h = 2.5$  ms.

% Power removed / 100



A Method of Finding Linear Operators for Least Squares Fitting Procedures

Described in this section is an iterative method of approaching the values of coefficients for a least squares fitting of linear operators to multiple time series. The problem arose in connection with the determination of linear operators to use in picking reflections from seismograms. † The method is extremely inefficient and is really only possible with the aid of very high speed computing machines, but it gives interesting insight into the behaviour of matrices, which helps in constructing other techniques.

We are trying to fit a linear operator of the form

$$\begin{aligned} \hat{x}_{1+k} = & c + a_0 x_1 + a_1 x_{1-1} \dots + a_M x_{1-M} \\ & + b_0 y_1 \dots + b_M y_{1-M} \\ & + c_0 z_1 \dots + c_M z_{1-M} \\ & + d_0 u_1 \dots + d_M u_{1-M} \end{aligned} \tag{2.24}$$

to an interval of the sequences  $x_1, y_1, z_1,$  and  $u_1$  so that

$$I = \sum_1 (x_{1+k} - \hat{x}_{1+k})^2 \text{ is a minimum.} \tag{2.25}$$

The plan is to guess initial values of the constants  $a, a_s, b_s, c_s,$  and  $d_s$  and compute 2.24 . Then adjust the constants so that  $I$  is continually reduced. The initial

† Ref. 5

values chosen are  $a = \bar{x}$  (mean of  $x_1$  series)  $a_s = b_s = c_s = d_s = 0$ . These values are the values which the constants would assume under least squares fitting procedures if the  $x_1$  series were truly random and had no predictability. With these values of the constants  $I = I(\bar{x}, 0, 0, \dots)$  becomes the sample variance about the sample mean.

The computational procedure is:

1. Find  $I(\bar{x}, 0, 0, \dots)$
2. Find  $I(\bar{x} + \Delta a, 0, 0, \dots)$
3. If  $2 < 1$ , continue adding  $\Delta a$  until  $I(\bar{x} + n\Delta a, 0, 0, \dots) > I(\bar{x} + (n-1)\Delta a, 0, 0, \dots)$ .  
If  $2 > 1$ , subtract  $\Delta a$  and continue to subtract until  $I(\bar{x} - n\Delta a, 0, 0, \dots) > I(\bar{x} - (n-1)\Delta a, 0, 0, \dots)$
4. Using  $\bar{x} \pm (n-1)\Delta a, 0, 0, \dots$ , as the new starting point, find  $I(\bar{x} \pm (n-1)\Delta a, \Delta a_0, 0, 0, \dots)$  and repeat the steps under 3.
5. Work successively in this fashion with each of the variables  $a, a_0, a_1 \dots d_M$ .
6. Start the process over again with the variable  $a$ .
7. Continue recycling until the desired accuracy is reached.

It is interesting to consider the geometry of this process. If we substitute equation 2.24 into 2.25,

we find that  $I$  is parabolic in each of the coefficients  $a, a_s, b_s, c_s, d_s$ . For simplicity consider the case where we have only two coefficients  $a$  and  $b$ . Then  $I$  is a two dimensional paraboloid in  $a$  and  $b$  whose minimum we wish to find.  $I$  is positive or zero for all  $a, b$  and has one minimum. Contours of  $I = c$  are ellipses in the  $a b$  plane of constant major to minor axis ratios, and are centered at the minimum. Figs. 2.11, 2.12, and 2.13 illustrate three situations that might arise. In Fig. 2.11 the contours are circular which is the case when the matrix of the normal equations associated with the minimum fit is well-behaved. Fig. 2.12 is the more usual situation where the contours are definitely elliptical. Fig. 2.13 shows a very badly-behaved situation corresponding to near singularity of the associated matrix.

The solid line shows how the iterative method described above would converge toward the minimum point in the three situations. The dashed curve shows how another iterative method, the steepest descent method, would converge in these situations. The steepest descent method runs into trouble in the near singular case because with finite increments it cannot land on the long axis of the ellipse. It is forced to wobble back and forth, much as a small ball would wobble rolling in such a trough. The method described above would also encounter trouble if the increment were not fine enough, for if it got near the



Fig. 2.11

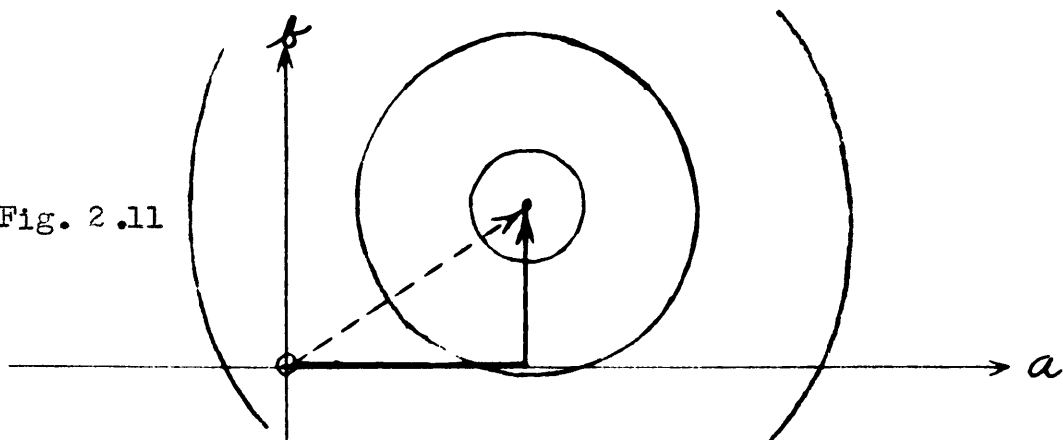


Fig. 2.12

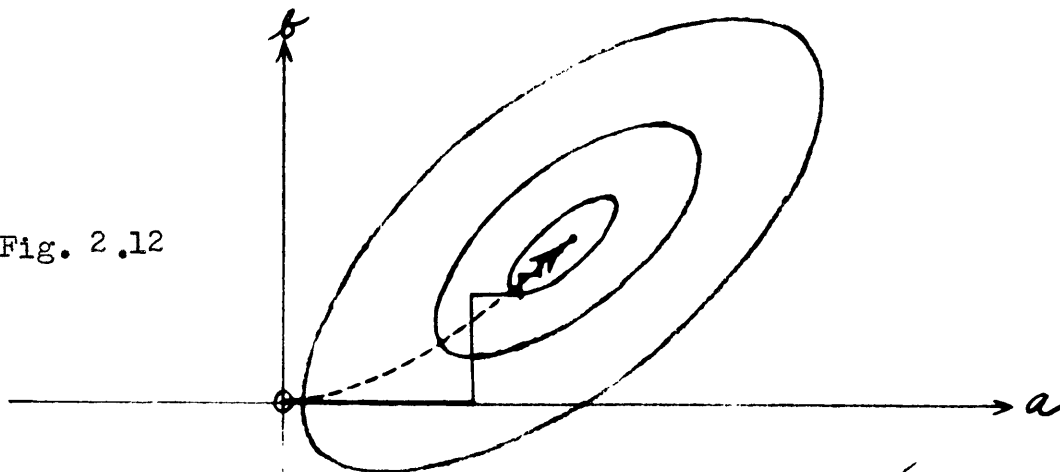
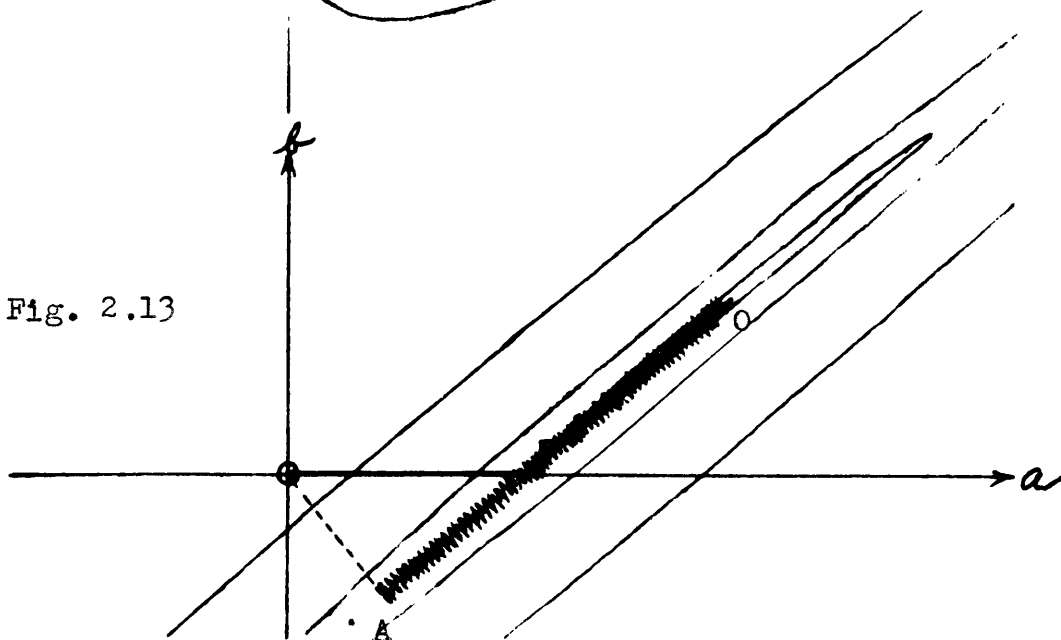


Fig. 2.13



trough the next increment would carry it across the trough to a greater value of  $I$ .

These figures illustrate a fundamental problem met with in iterative methods. The fine increments necessary in treating the near singular case are very inefficient when used on well-behaved data, whereas the larger increments applicable in Fig. 2.11 could never find the minimum of Fig. 2.13 .

A program was written for the WWI Digital Computer which would do this one-variable-at-a-time type of iteration. It is described in Appendix C. The computations it carries out take fifteen or twenty minutes of machine time, but they represent nearly a year of hand computation. The program can print out each successive value of  $I$  as it is computed. Fig. 2.14 shows a plot of these values as the program converges towards the solution of a particular problem. This diagram shows how  $I$  is parabolic in each coefficient. We also note that all the parabolas have approximately the same shape. This indicates that if there is a predominant long ellipse axis as in Fig. 2.13 , it cannot be close to parallel to any of the axes  $a$ ,  $a_s$ ,  $d_s$ , for if it were, the parabolic section in the corresponding direction would be quite flat. One surprising feature of this diagram is the failure of the parabolas to tend to flatten as  $I$  is diminished.

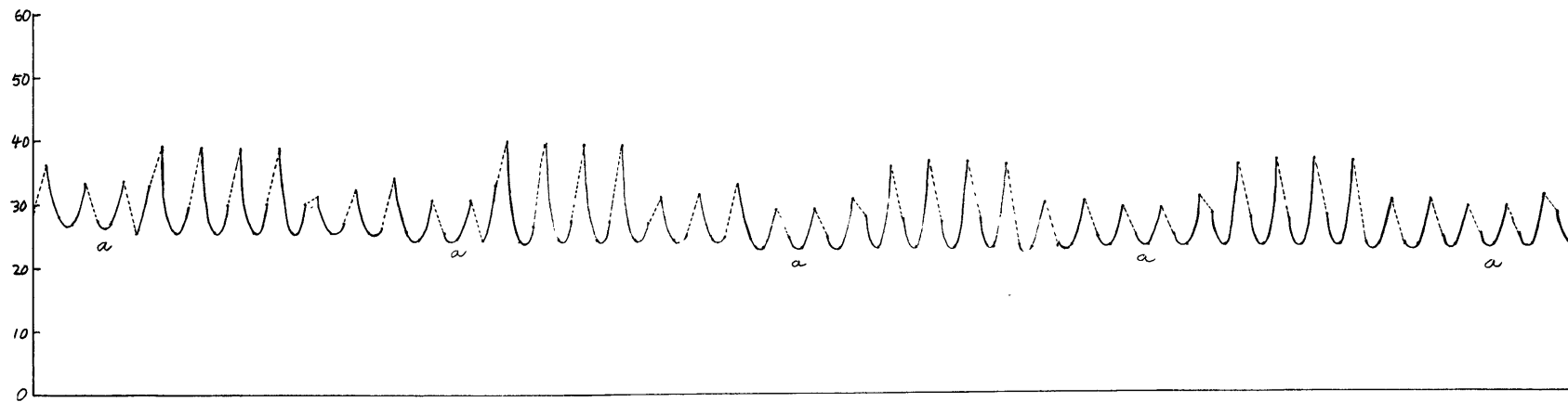
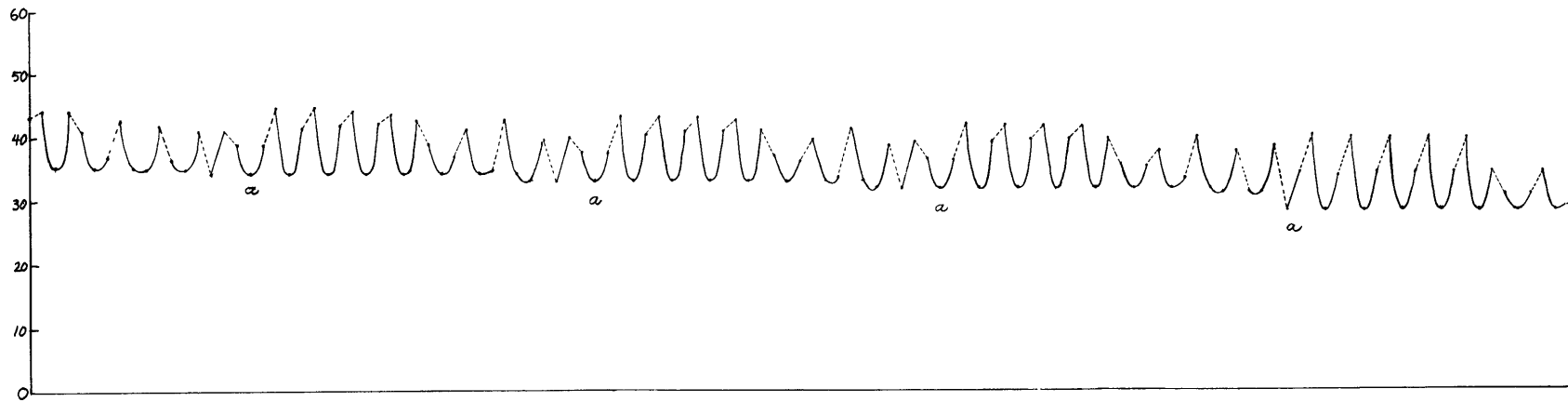
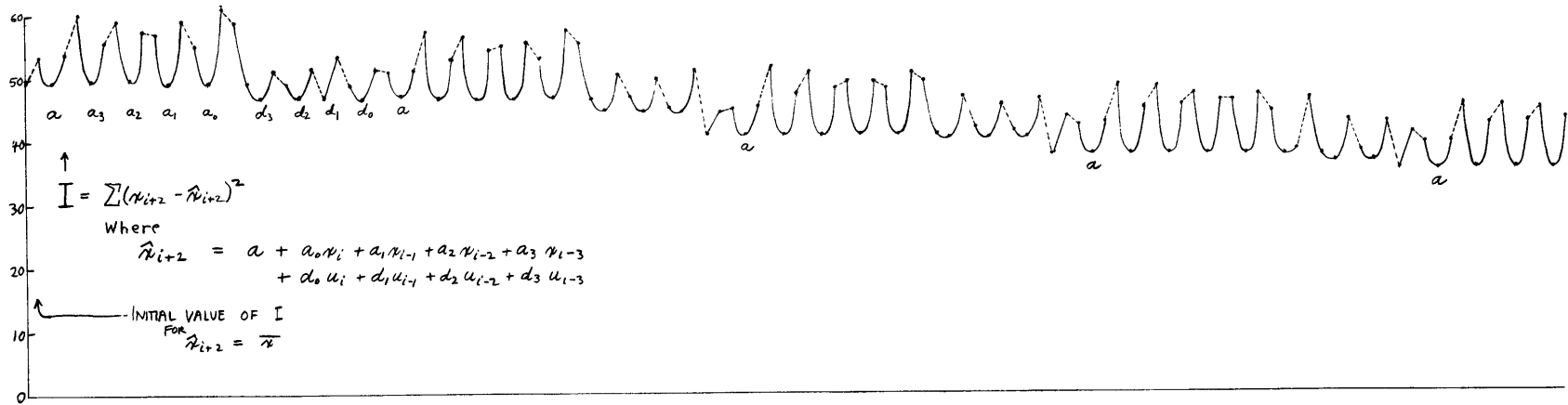
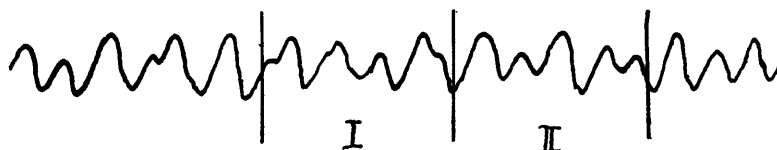


FIG. 2.14

## Accuracy

This is a convenient point to consider the problem of the importance of obtaining the exact solution. If we look at Fig. 2.13, we see that values of  $a$  and  $b$  at the point  $A$  will reduce  $I$  almost as well as values at the true minimum  $O$ . Individual errors  $(x_1 - \hat{x}_1)$  will likewise be practically identical. The effect of the displacement  $OA$  will not be felt until the values at  $A$  are used to predict outside the interval where the minimum fit is taken.

Suppose the series is



and the minimum fit is taken in the interval  $I$  of this series. What happens when we predict the interval  $II$  with coefficients chosen in  $I$ ?

Consider Fig. 2.15. The dark solid line represents the long axis of the ellipses for interval  $I$  and the light solid lines, the contours for this interval. The true minimum of these contours is at  $O$ . Likewise, we can draw a similar contour picture for the interval  $II$ . If we assume the dynamics are but slightly different in the two intervals, the second contours will be slightly rotated with respect to the first, and, there will be a small displacement of the minimum. The heavy and light dashed lines in Fig. 2.15

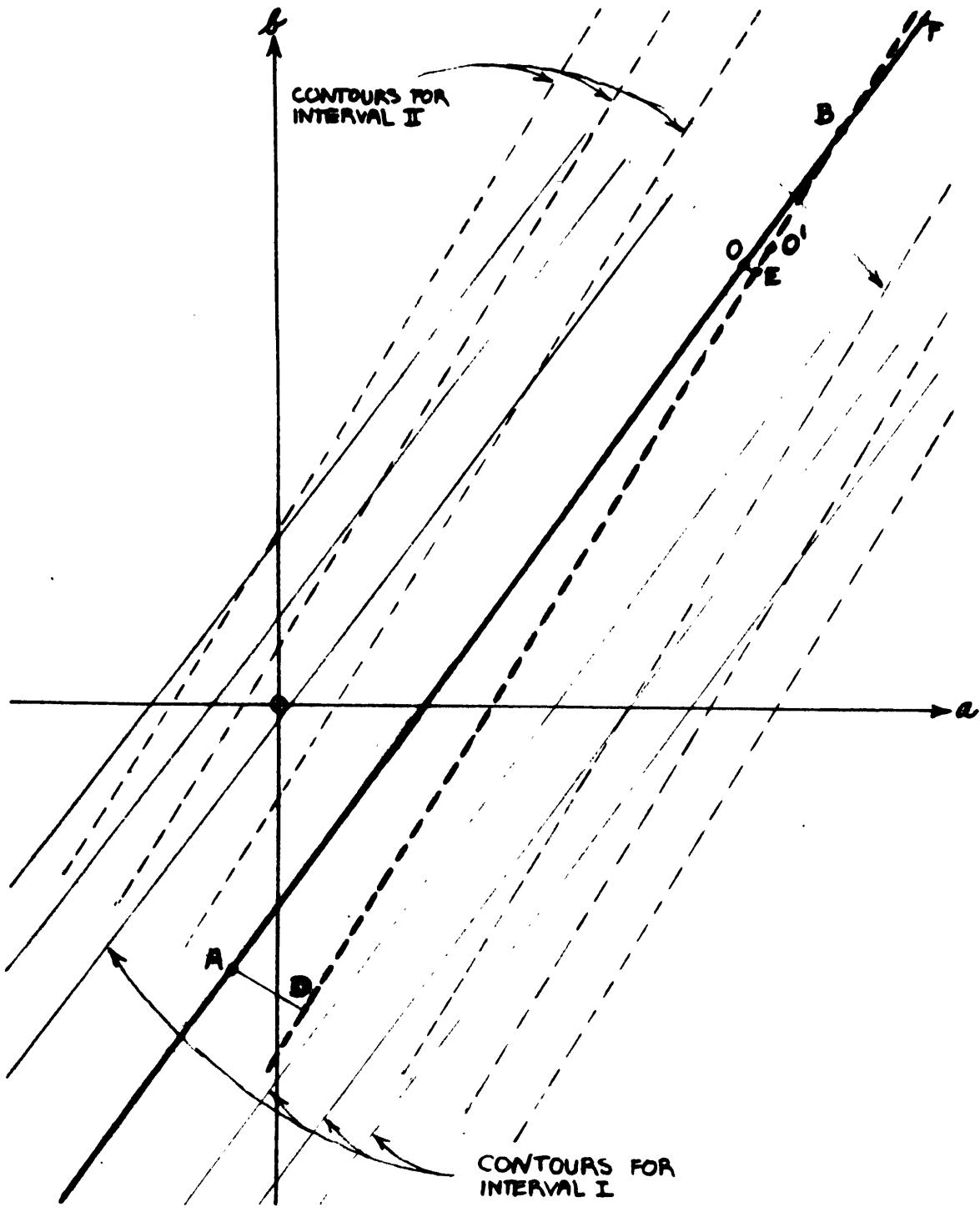


Fig. 2.15

represent these contours for interval II. Since we are considering the near singular case, the deviation of the sum of squared errors for the second interval from its value on the heavy dashed line will vary as the square of the distance from a point in the ab plane to the heavy dashed line.

Now suppose in finding our minimum point for interval I we had landed at the point A which satisfies the least squares criterion almost as well as the true point O. The deviation of the sum of squared errors when point A is used to predict interval II will be proportional to  $(AD)^2$  which would be about sixteen times greater than if the point O were used, since  $OE = 1/4 AD$ . On the other hand, if we had landed at the point B for the first interval we would get a sum of squared errors smaller than if the point O were chosen. Again, if the point F were taken, the sum of squared errors for interval II would not be appreciably different than for the point O.

These effects have been noted in computed data. The indication is that the true minimum point O must be chosen if we are to take the sum of squared errors as a valid comparison of the changing dynamics in various intervals of a series by this method.

PART III

## SOME INTERPRETIVE PROCEDURES

### Introduction

In this part we present several ideas which may be applicable to answering certain questions involving seismogram analysis. Two of the ideas have had some testing, the others none. With one exception these ideas relate specifically to reflection seismic records, and various possibilities in picking reflections therefrom. The questions are:

1. In a two velocity system, e.g., shear and compressional waves, can we set up a method for separating these velocities and can we apply it to reflection determination?
2. In the use of linear operators for seismogram analysis, is there another measure of prediction error, other than the "error curve", which will show reflections?
3. Can we obtain information on the step-out times of reflections, by the use of linear operators and the concept of ensemble averages?
4. Can a special seismometer set-up be used in conjunction with correlation analysis to pick reflections?



## Velocity Separation

The determination of velocities for compressional waves in the earth at shallow depths is relatively simple due to (1) the ease in generating such waves, and (2) the fact that the first arrivals are the compressional waves. Shear waves are more difficult to generate with sufficient amplitude to separate from the earlier arriving types. Although with the proper equipment this can be done by visual inspection of the seismogram, † it seemed of interest to consider if a statistical test could be devised to help in this problem.

The approach was to set up a simple model approximating the physical situation.

Assume we have two wave forms A and B traveling horizontally at velocities  $V_A$  and  $V_B$ , where  $V_A > V_B$ , past three geophones F, G, and H, equally spaced with separation  $d$ . The wave shapes do not change with time. Traces F, G, and H then represent composites of A and B with different time lags. Assuming  $V_A$  is known, the problem is to find  $V_B$  and, if possible, the wave forms A and B.

Divide the time scale into units such that the no. of units per sec. is  $L$ . Since  $V_A$  and  $d$  are known we may line up F, G, and H so that very nearly

† Ref. 8

$$F_N = A_N + B_N \quad 3.1$$

$$G_N = A_N + B_{N-j} \quad 3.2$$

$$H_N = A_N + B_{N-2j} \quad 3.3$$

where the time lag between traces is approximated by  $j$  units

so that 
$$\frac{j}{L} = \frac{-d}{V_A} + \frac{d}{V_B}$$

or

$$V_B = \frac{LdVA}{jV_A + Ld} \quad 3.4$$

This is illustrated in Fig. 3.1 .

From equations 3.1, 3.2, and 3.3 we can get

$$A_N - A_{N-j} = G_N - F_{N-j} \quad 3.5$$

$$A_N - A_{N-2j} = H_N - F_{N-2j} \quad 3.6$$

3.5 and 3.6 are recursion formulas giving  $A_{N-kj}$  and

$A_{N-2kj}$  respectively ( $k = 1, 2, \dots$ ) once  $A_N \dots A_{N-j+1}$

are known. Now if  $j$  has its correct value then it is easy

to show that regardless of how we choose the initial A's

both formulas give the same value for  $A_{N-2kj}$  . If  $j$  is

slightly wrong then the two series will differ slightly.

The difference will increase as  $j$  strays further from its

true value. We may now set up a procedure for finding this

value. Assume values for  $j$  and for  $A_N, A_{N-1}, \dots A_{N-j+1}$ ,

use equations 3.5 and 3.6 to calculate the two series (to

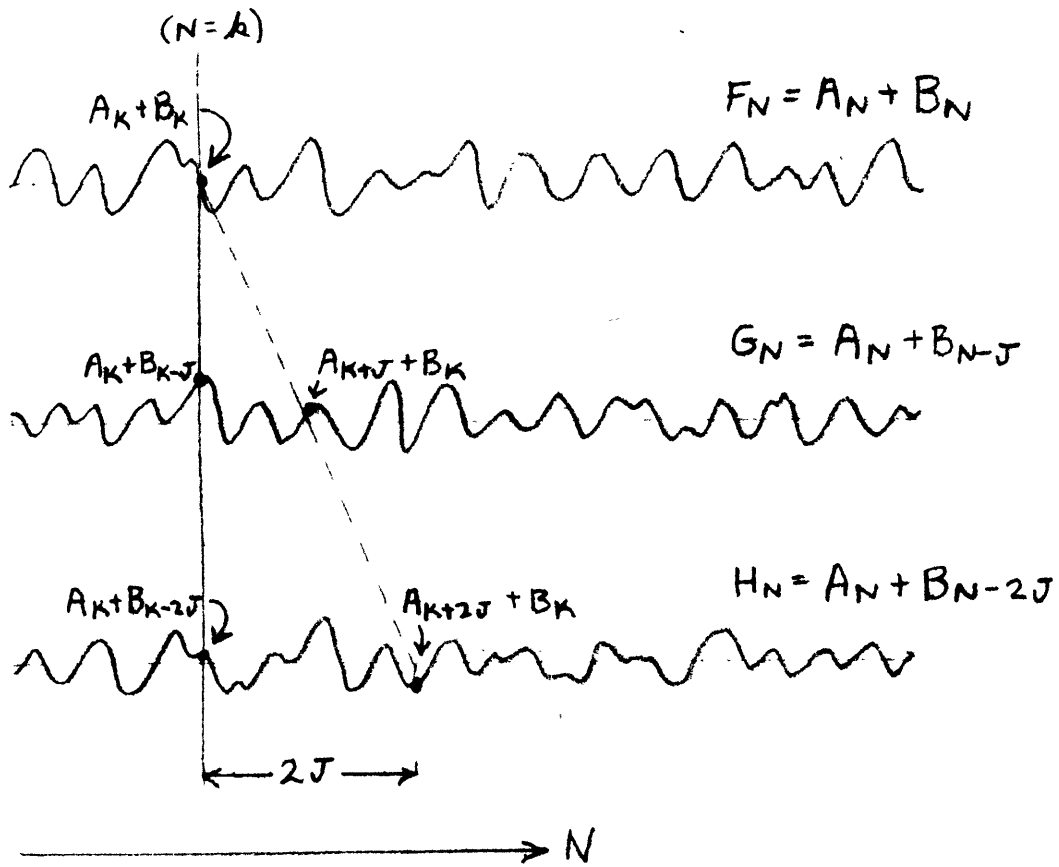


Fig. 3.1

ASSUMED REFLECTION PATTERN

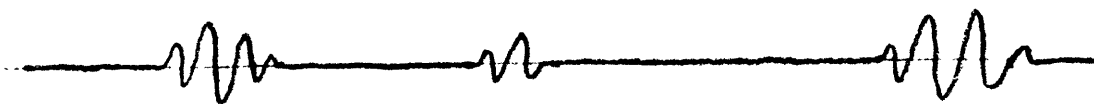


Fig 3.2

a certain length), find the mean square difference between the series, and plot this difference as a function of  $j$ . In the ideal case, this function will go to zero for the correct value of  $j$ . In practice we can expect this difference function to have a minimum at the correct value.

Thus in theory at least,  $j$  is determinable. Equation 3.4 may be used to find  $V_B$ . Although the exact wave shapes are indeterminate in the general case, there may be obtained some information about them. Assume the first  $j$  values of  $A_N$  are taken to be zero. If  $j$  is correct, then the series from equation 3.5 represents the true  $A_N$  with the first  $j$  values subtracted successively. Thus the series from equation 3.5 might be expected to have the same frequency characteristics as the true  $A_N$ .

In certain cases the assumption that  $A_N \dots A_{N-j+1} = 0$  will be fairly accurate. In these cases the wave forms should be determinable. Examples would arise in the separation of shear and compressional waves where it is known that the shear waves arrive late, and in reflection picking.

Another possibility in this problem would be the use of pure cross-correlation between two traces. We should expect to get a peak in the correlation at a lag corresponding to the velocity  $V_B$  and the particular geophone separation. However, if the wave form B were of small

amplitude, the shape of the cross-correlation curve would effectively be dominated by that of the auto-correlation of wave form A, and the selection of the peak would be somewhat arbitrary. On the other hand, the mean square difference between equations 3.5 and 3.6 should still show a true minimum at the correct lag.

We can adapt this idea to the selection of reflections on seismic records. Here we make the simplified assumptions that the reflection consists of a wave train with zero amplitude between reflections as in Fig. 3.2 . This is assumed to occur on two traces in the same form and at the same time (i.e., there is no step out time of the reflection which is assumed to be coming in vertically). In this case equation 3.5 alone is applicable and we need only two traces.

$$A_N - A_{N-j} = G_N - F_{N-j} \quad 3.5$$

$j$  is taken from the step-out time of the initial breaks on the seismogram. We then select some interval  $j$  units in length in which  $A_N$  is zero (a non-reflection interval), and use equation 3.5 to predict the remainder of the reflected wave. For interpretation it is convenient to plot the running variance of the predicted reflection.

Now the assumptions will certainly not be upheld exactly on any real seismic record. A certain amount of random energy will be in phase between any two traces and would be picked out by this method as part of the predicted

reflection. To alleviate this situation we can use three traces and predict the reflected wave from the three possible pairings of these traces. Adding the three predicted waves would tend to accentuate components in phase between all three, and to minimize the random, in-phase components between any two traces. For three traces  $F_N$ ,  $G_N$ , and  $H_N$  we can express this sum as

$$3A_{N+K} = A_N + A_{N+j} + A_{N+K-j} + G_{N+K} + 2H_{N+K} - F_N - F_{N+K-j} - G_{N+j} \quad 3.7$$

where  $j$  corresponds to the step-out between  $F_N$  and  $G_N$ , and  $K$  the step-out between  $F_N$  and  $H_N$ .

### Tests of the Method

#### 1. Selection of Shear Velocity

An initial test was constructed which showed that, when the assumptions were exactly upheld, the minimum of the plot of the squared differences between equations 3.5 and 3.6 was quite sharp.

On this basis three adjacent traces of a seismogram were converted to numerical form and the method applied to these real series. The seismogram was taken at Revere Beach, Mass., in unconsolidated sediments, by Peter Southwick.† Special generating apparatus was used so that the shear arrivals were quite prominent. This record is now lost, but

† Ref. 8

Fig. 3.3 shows a very similar seismogram taken with the same apparatus. The first line of check marks on this seismogram indicates the first arrivals, and the second line of check marks was picked as the arrivals of the shear waves. This second line permitted a direct computation of the shear velocity.

The readings for the three traces were lined up in accordance the first line of time breaks, and equations 3.5 and 3.6 were computed for a variety of values of  $j$ . In each case the first  $j$  values of  $A_N$  were assumed to be zero. The sum of squared differences between these two series were computed for each  $j$ , and normalized by the number of terms in the series for each  $j$ . A plot of these quantities appears in Fig. 3.4 .

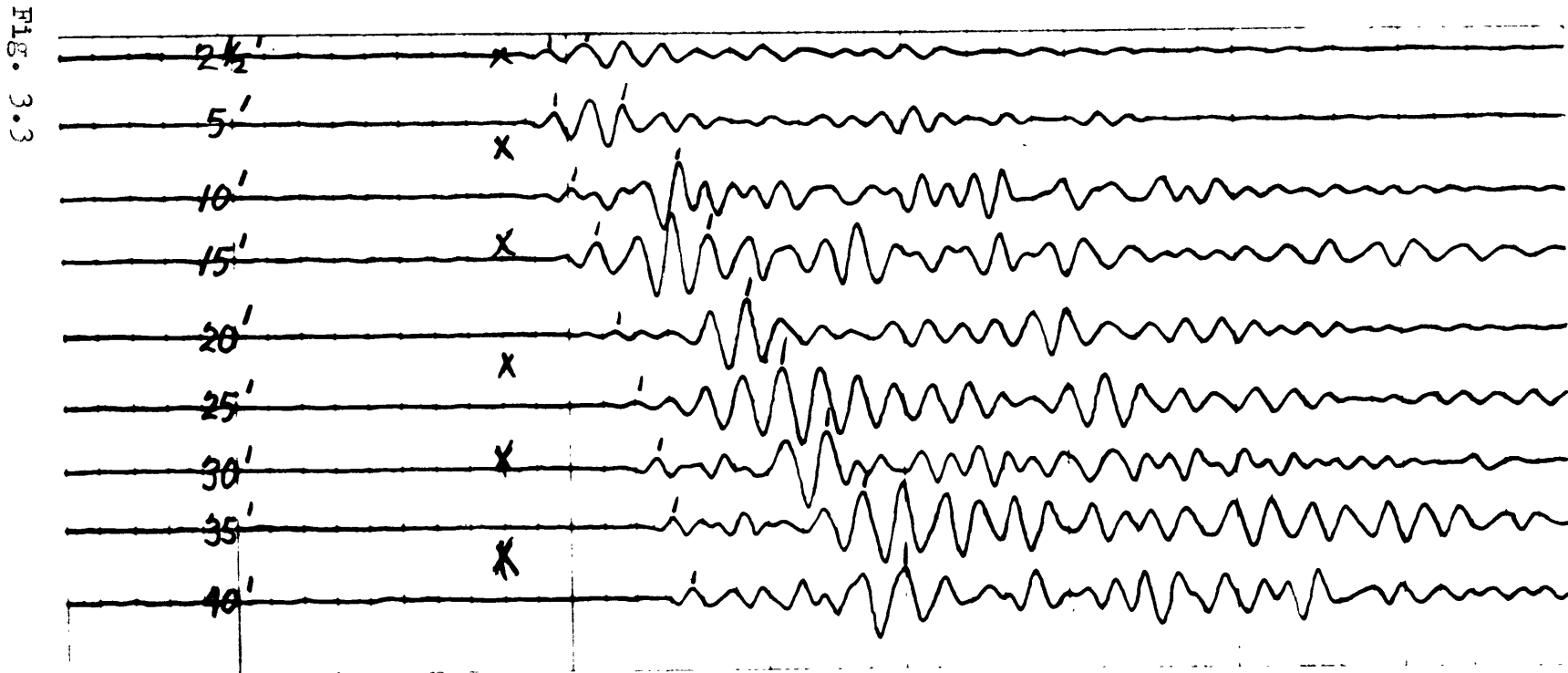
This figure shows two distinct minima (at  $j = 13.3$  and  $j = 16.0$ ) rather than just one. Upon examination it turned out that the value  $j = 13.3$  corresponded to a shear velocity which would have been computed by direct interpretation of the first two traces chosen. The second minimum corresponded to a velocity which would have been determined directly from the second and third traces chosen. The value of velocity computed by the entire second line of check marks of Fig. 3.3 lay between these two values.

Fig. 3.5 shows a running average of the points in Fig. 3.4 (by overlapping groups of three) which exhibits a flat minimum between  $j = 13.3$  and  $j = 16.0$  . The

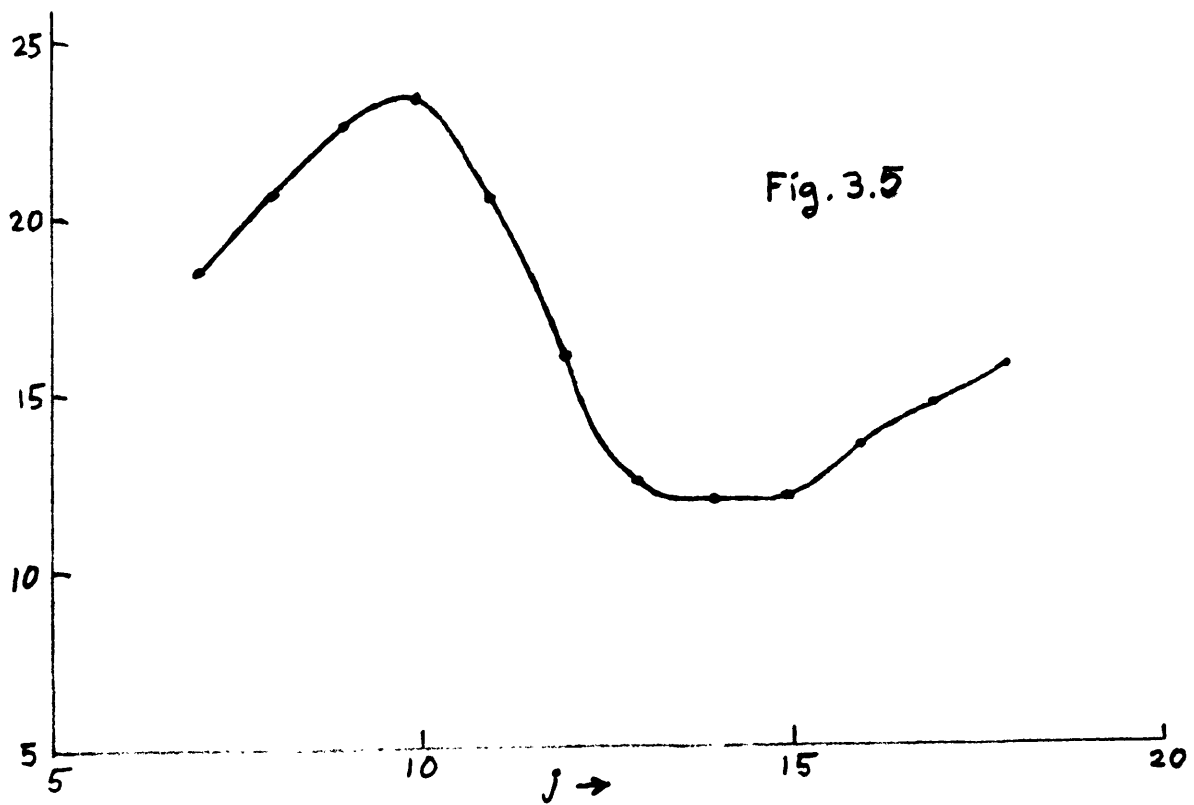
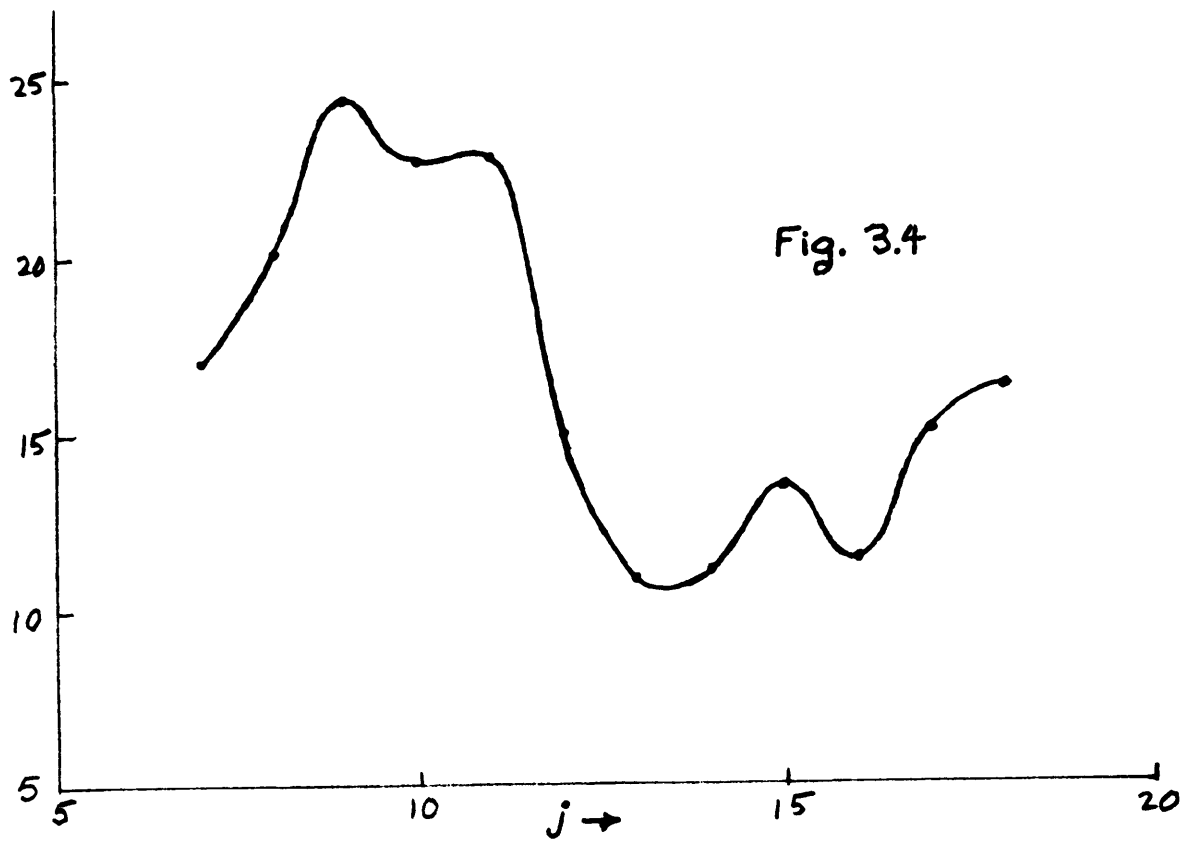
# LEXINGTON S+9

Shear test - med. sand

Geophones in line







corresponding velocity was quite close to that computed from the second line of check marks.

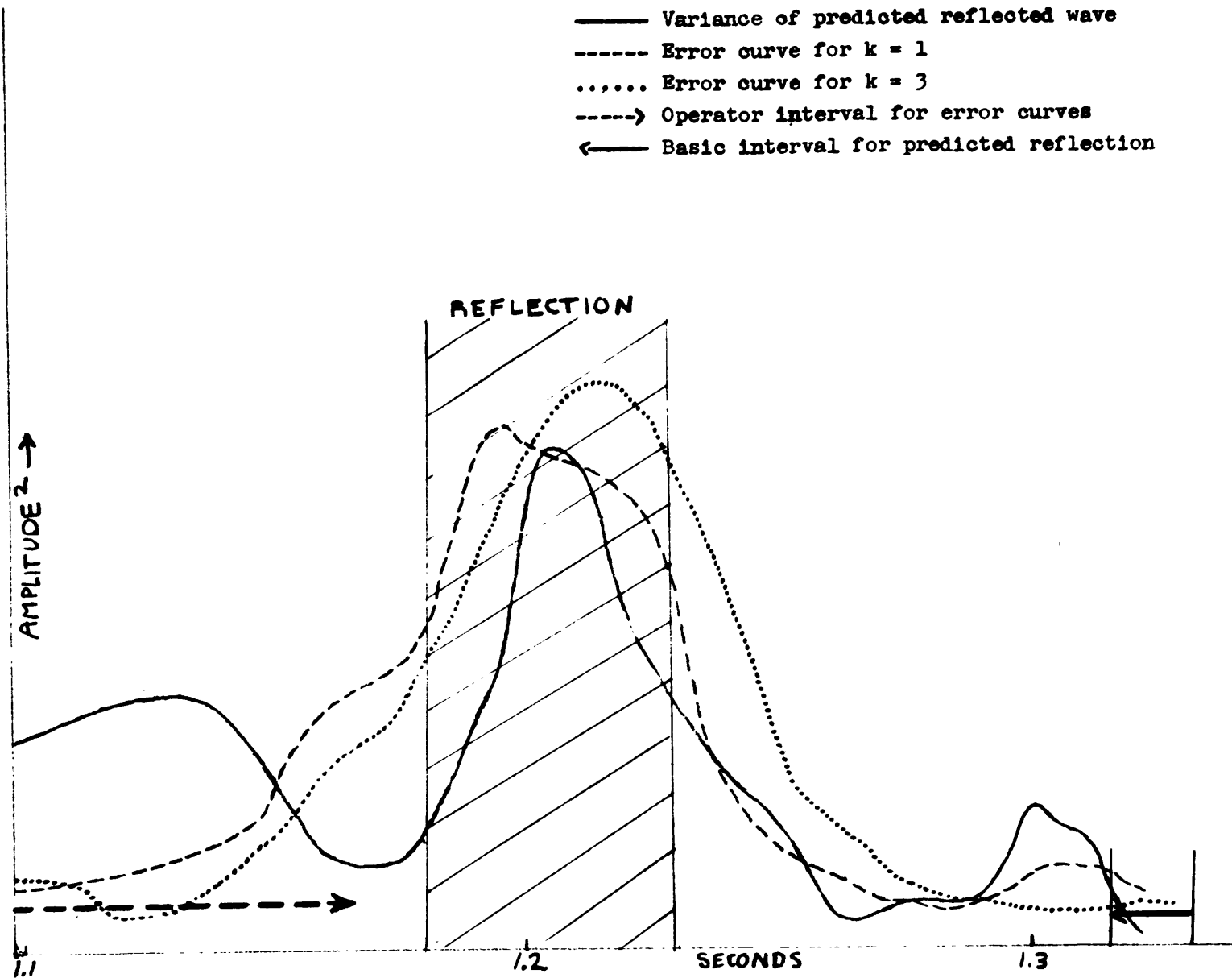
## 2. Predicting a Reflected Wave

To test whether or not a reflection could be predicted by these methods, a seismogram showing a prominent reflection was chosen (MIT Record No. 1 †). In this record linear operators had been computed and error curves derived. These error curves showed marked peaks at the reflection so the curves were taken as a basis of comparison.

From two traces on this record equation 3.5 was computed.  $j$  was selected from the initial step-out between the traces and the non-reflection interval chosen to occur after the reflection. The variance of the predicted wave (in overlapping groups of ten) is plotted in Fig. 3.6 . The dotted and dashed curves of this figure show error curves for linear operators with different prediction distances  $k$ . The variance curve does not reach a peak in the reflection as rapidly as do the error curves, but it does compare favorably with them in general shape during and after the reflection. Before the reflection the discrepancy is more noticeable. This may very well be attributable to the fact that operator interval was chosen just before the reflection. In the operator interval, the least squares fitting procedure forces the error curves to be as low as possible.

† Ref. 4

Fig. 3.6



## Conclusions

The two tests discussed show that the expected effects are noted. However, the data used are reasonably ideal, in the sense that ordinary methods of interpretation are adequate. Whether or not the statistical techniques are better can only be determined by many further trials. Situations difficult to treat by the ordinary methods will also fail to uphold the simple assumptions of the theory presented here. On the other hand, only the simplest forms of the theory were used in the examples. Refinements, such as the use of three or more traces for reflection picking may give more valid results.

### Phase Test

The "error curve", as used by the Geophysical Analysis Group for picking reflections, is a running average of the squared differences between a predicted and an actual seismic trace. Fig. 3.7 shows an actual trace (the solid curves), and three predictions of this trace, from linear operators with different values of prediction distance. From this diagram we see that the error curve is a running measure of the vertical differences between the predicted and actual traces.

At the reflection (shaded) these differences are seen to become large, and hence the error curve rises to a peak in this interval. The reason the differences become large is not because there is a big discrepancy between the average amplitudes of the predicted and actual traces. From the diagram it appears that the reason is that there is a horizontal displacement of the oscillations of one trace with respect to the other. In other words, there is a phase shift between the predicted and actual traces during the reflection, which disappears shortly after the reflection.

It seems then that a test of phase relationships might well show the reflections as well as the error curve does. A fairly rigorous way of testing this phase shift would be the following:

M.L.T. RECORD NO. 1 SHEET B

MAGNOLIA PETROLEUM CO  
 PROSPECT : PACE  
 PROFILE : FB-N  
 SPREAD : N250 - N750  
 RECORD : T  
 CHARGE : 5  
 DEPTH : 90'  
 DATE : MAR 15, 1950  
 A.V.C. TRACE, C.V. PHONES

SEISMOGRAM TRACE N750 ———  
 PREDICTED TRACE N750 (FROM N750 & N250) - - - -  
 OPERATOR ———→  
 COMPANY MARKED REFLECTIONS // //  
 ABSCISSA - TIME IN SECONDS FROM SHOT  
 ORDINATES - UNITS OF TRACE AMPLITUDE  
 SUPERVISOR ROBINSON  
 DATE AUGUST 1951

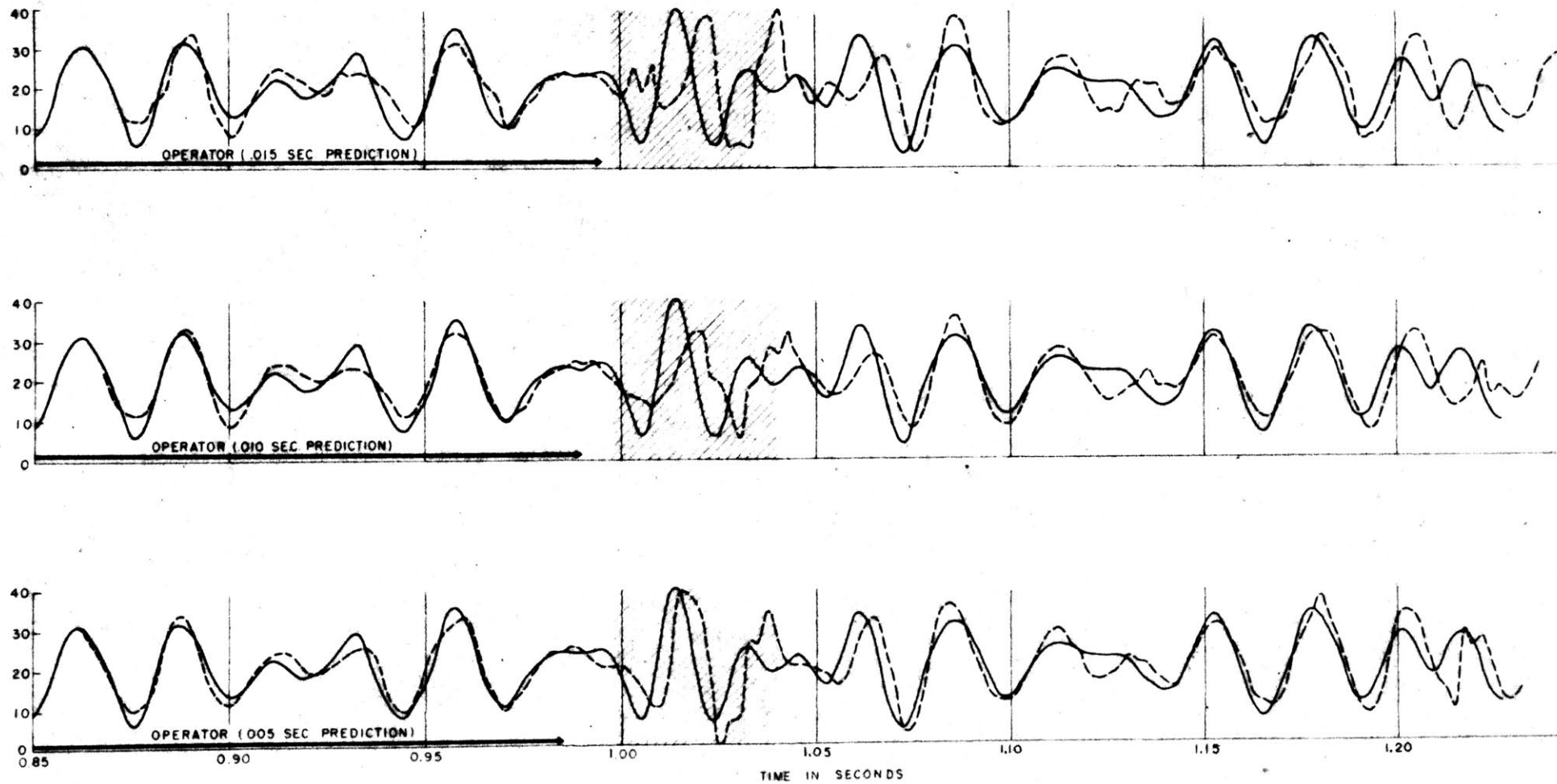


Fig. 3.7

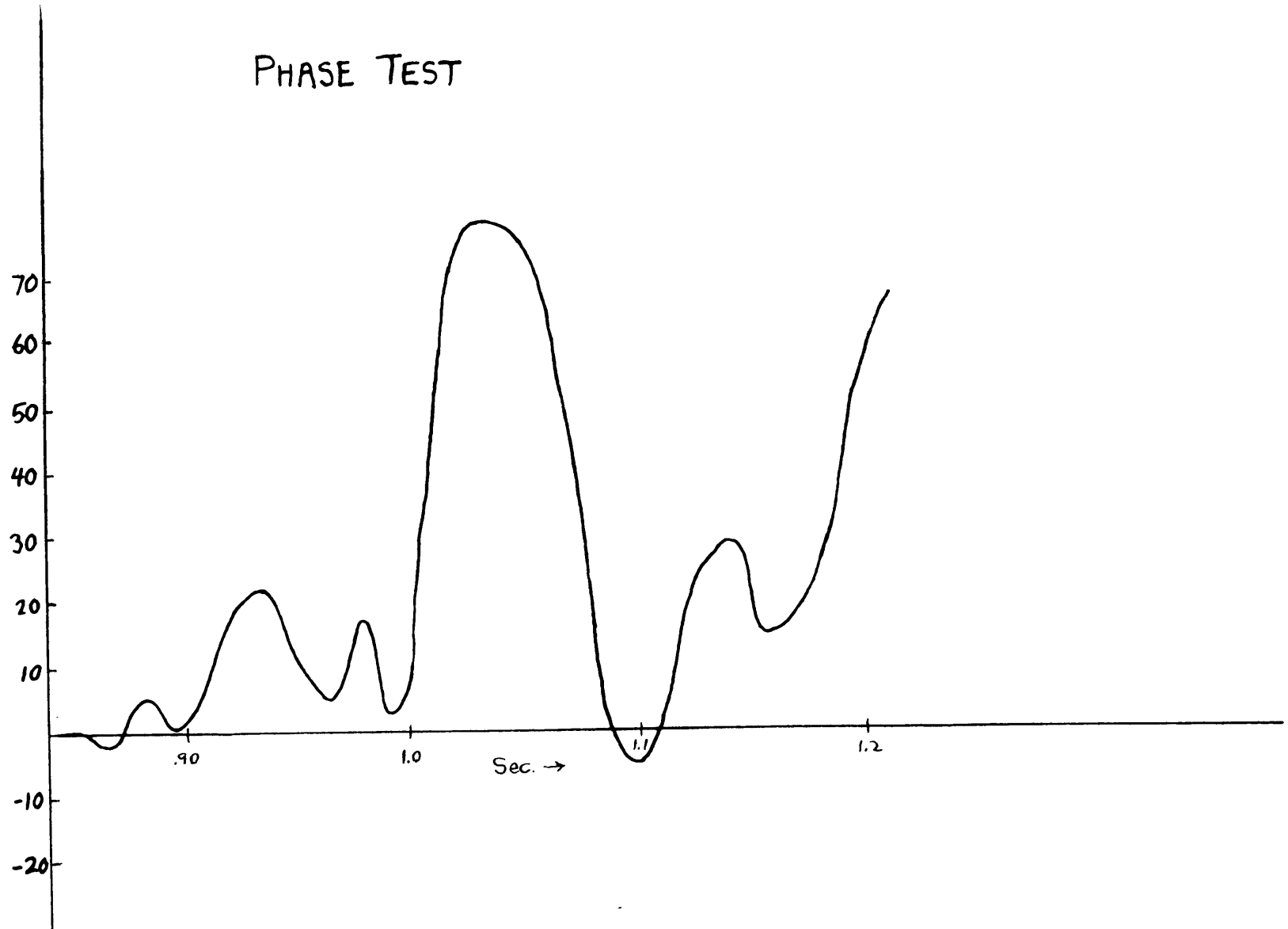
1. Select highly overlapping intervals of the record.
2. Compute the cross-spectrum between the predicted and actual traces in each interval, thus obtaining the phase relationships.
3. Plot the phase angle of the dominant frequency as a function of the interval chosen.

Practically, this is an involved procedure. We can use a simple but crude method to get approximately the same results. Since phase shift is expressed by horizontal displacement we can measure this displacement directly from graphs such as in Fig. 3.7 . This requires that we be able to follow corresponding waves in the two traces, which is subject to personal interpretation.

The displacement was measured for the upper set of curves in Fig. 3.7 . From equally spaced points (in time) on the solid curve the horizontal displacements to the dashed curve were measured. Displacements to the right were considered positive, those to the left negative. Where such measurements could not logically be made (for example on the peak occurring at about .96 sec.) values were taken midway between the last value that could logically be made and the next such value. Once this series of displacements was determined, its individual terms were summed in groups of twenty overlapping by ten, in order to smooth the data. These sums are plotted in Fig. 3.8 .

# PHASE TEST

FIG. 3.8





This curve indicates a rapid change of phase occurs at the reflection, the phase rising to a peak in the middle of the reflection, and falling off more gradually thereafter. It seems surprising that the curve is almost entirely positive. If this effect is characteristic, perhaps we should consider as significant only those portions of the curve above a certain mean (about 25 or 30 units in Fig. 3.7). From the original record it appears that there may be another reflection at about 1.23 seconds, which could conceivably cause the rise at the end of the curve.

This is a purely empirical curve. Perhaps it only holds for the particular case treated. One would suspect that the arrival of reflected energy would be accompanied by a rapid change in phase relationships. However, it does not seem reasonable that these changes should be in one direction since the times of arrival of reflected energy are random. Possibly we should deal with the original series only, and compute the rate of change of phase angle (between two overlapping intervals) as a function of interval.

### Ensemble Average

The step-out time of a reflection is a property of several seismic traces rather than just a single trace. The error curve for linear operators, as defined elsewhere in this paper, is a property of a single trace - a time average of a single error time series. To get information on the step-out time we must consider operators chosen for different traces. In this connection it is convenient to use an "ensemble" average. This is an average across the "ensemble" of error time series generated by the various operators chosen.

Let us suppose that we have taken a series of operators on a record which consists of traces from equally spaced seismometers. Suppose there are  $T$  traces, and on the  $l^{\text{th}}$  trace ( $l = 1, 2, \dots, T$ ) we have chosen  $N_l$  operators. For the  $k^{\text{th}}$  operator on this trace ( $k = 1, \dots, N_l$ ) there is an associated error time series which we define as  $e_1^{(kl)}$ . Then, for example, we may construct a single error time series  $\epsilon_1^l$  to be associated with the  $l^{\text{th}}$  trace by the expression

$$\epsilon_1^l = \sum_{k=1}^{N_l} [e_1^{(kl)}]^2 \quad 3.7$$

We may then average these error time series over the various traces. Between traces we observe the effect of step-out. Hence we construct the error time series  $\delta_1^{(\alpha)}$  with an arbitrary lag or lead  $\alpha$

$$S_1^{(\alpha)} = \sum_{l=1}^T \epsilon_{1-\alpha l}^{(l)} \quad \alpha = 0, \pm 1, \pm 2, \dots 3.8$$

with the expectation that a peak on this error time series, corresponding to a certain reflection, should be highest and narrowest for that value of  $\alpha$  most closely corresponding to the true step-out of the given reflection.

No attempt has been made yet to compute 3.8 . It would be a fairly simple task to program this equation for the WVI Digital Computer as a follow-up of the Prediction XV program described in Appendix B.

### Travelling Correlations

As an approach to the problem of using a special geophone layout for reflection picking, consider the following arrangement. Two geophones  $G_1$  and  $G_2$  are placed in the ground, one vertically under the other at a distance  $d$ . Assuming the ground homogeneous and non-dispersive around the geophones, the responses of  $G_1$  and  $G_2$  may be considered to be due to superpositions of many plane waves travelling with a velocity  $V$  from many different directions. In the absence of big reflections, the major contribution to the responses will come from waves having directions not far from the horizontal.

Now consider the cross-correlation of the two responses at  $G_1$  and  $G_2$ . In particular consider the value of the function for a time lag equal to  $d/V$ . It appears that this value will be strongly influenced by the amount of vertical wave contribution present in the responses, since  $d/V$  is the time of direct travel from  $G_2$  to  $G_1$ . The cross-correlation at the lag  $d/V$  should rise rapidly at a reflection and drop off afterward.

In practice we would have to compute this correlation over highly overlapping time intervals of the response functions in order to obtain the correlation as a function of time. The correlation program described in Appendix D is adaptable to this type of analysis. So far however, no seismograms with the above geophone arrangement have been available.

## CONCLUSIONS AND SUGGESTIONS FOR FUTURE WORK

It is difficult to make evaluations of validity on methods which have undergone little testing. Nevertheless we may draw certain conclusions from the work presented here.

Polynomial gravity approximation, as presented here, seems of sufficient simplicity and validity to justify a considerable amount of further study. If further trials show more promise it would be well worth-while to find the inverses of the matrices of Table III. In any event, polynomial approximations of this type have applications in many other fields, and the simplifications brought forward here may be of real value in these other applications.

The properties of cosine operators are of mathematical interest, but it is hoped that studies of this sort will lead to more practical results. In particular, further pursuit of the filter characteristics of linear operators will lead to a better understanding of the extent of realizability of equivalent electronic filters, and or to simplification in the determination of such operators.

The author is more hesitant about recommending the various procedures discussed in Part III. Seismograms exhibit extreme variability in their characteristics and, whereas the examples given here are encouraging, the procedures may fail on other types of records. However, the problems they attempt to settle are of great practical concern and all promising techniques should be either proved

or disproved. Phase is a crucial variable in these problems, and probably considerable effort should be spent studying this parameter.

As for the Appendixes, the author feels that the programs described therein have genuine value. Anyone concerned with research depending largely on computation appreciates the fact that obtaining errorless results is a major problem. Programs such as these effectively eliminate this type of problem, and are available for the use of persons interested in the sort of computations they perform.

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APPENDIX A





III. It prints out these residuals in the same network fashion that the grid was chosen.

Use of Polynomial I

There are certain conventions which must be observed in the use of this program. The constants defining the nature of the polynomial and grid must appear as follows:

Register	440	+n	order of polynomial (less than 7)
(Octal)	441	+N	greatest value of x
	442	+M	greatest value of y

The coefficients  $C_{jk}$  of the polynomial must be scale factored in a special way because they decrease in magnitude rapidly as  $j+k$  increases. The scale factor is  $10^{(j+k)-2}$ , which in most instances will guarantee that all are less than unity in absolute value, but not greatly so. They must appear in the machine as follows:

Register	443	$C_{00}$	x	$10^{-2}$	461	$C_{12}$	x	10
(Octal)	444	$C_{10}$	x	$10^{-2}$	462	$C_{22}$	x	$10^2$
	445	$C_{20}$	x	1	463	$C_{32}$	x	$10^3$
	446	$C_{30}$	x	10	464	$C_{42}$	x	$10^4$
	447	$C_{40}$	x	$10^2$	465	$C_{03}$	x	10
	450	$C_{50}$	x	$10^3$	466	$C_{13}$	x	$10^2$
	451	$C_{60}$	x	$10^4$	467	$C_{23}$	x	$10^3$
	452	$C_{01}$	x	$10^{-1}$	470	$C_{33}$	x	$10^4$
	453	$C_{11}$	x	1	471	$C_{04}$	x	$10^2$
	454	$C_{21}$	x	10	472	$C_{14}$	x	$10^3$

455	$C_{31}$	$\times 10^2$	473	$C_{24}$	$\times 10^4$
456	$C_{41}$	$\times 10^3$	474	$C_{05}$	$\times 10^3$
457	$C_{51}$	$\times 10^4$	475	$C_{15}$	$\times 10^4$
460	$C_{02}$	$\times 1$	476	$C_{06}$	$\times 10^4$

The data  $g(xy)$  which is presumed to be taken over the gridwork, is scale factored by  $10^{-2}$  and appears in the machine as follows:

Register	540	$g(-N, M)$	$\times 10^{-2}$
	541	$g(-N+1, M)$	$\times 10^{-2}$
	542	$g(-N+2, M)$	$\times 10^{-2}$
	.	.	
	.	.	
	.	$g(N, M)$	$\times 10^{-2}$
	.	$g(-N, M-1)$	$\times 10^{-2}$
	.	$g(-N+1, M-1)$	$\times 10^{-2}$
	.	.	
	.	.	
	.	$g(N, M-1)$	$\times 10^{-2}$
	.	$g(-N, M-2)$	$\times 10^{-2}$
	.	.	
	.	.	
	.	.	
	.	$g(-N, -M)$	$\times 10^{-2}$
	.	$g(-N+1, -M)$	$\times 10^{-2}$
	.	.	
	.	.	

$$\begin{array}{r}
\cdot \quad g(N, -M) \quad \times \quad 10^{-2} \\
\cdot \quad g(-N+1, -M) \quad \times \quad 10^{-2} \\
\cdot \quad \cdot \\
\cdot \quad \cdot \\
\cdot \quad g(N, -M) \quad \times \quad 10^{-2}
\end{array}$$

Now suppose the information  $n$ ,  $N$ ,  $M$ , and the  $C_{jk}$ 's are prepared on a tape with the tape number  $X$ , and the data  $g(xy)$  is prepared on a tape with the tape number  $Y$ . Then the instructions for the operation of this program would be

```

Erase storage
Read in 2492 m l
Read in X
Read in Y
Start at 127 (Octal)

```

The residuals are printed out by the direct printer in about three minutes or so depending on the size of the grid. They appear as four-digit numbers where the decimal point is understood to occur after the second digit.

As an example of the output we include a sample of three sets of residuals. These were derived for the three sets of coefficients used elsewhere in this paper. The sample illustrates the convenience of this form of answer for contouring purposes. In fact, with only slight modification (inserting two extra carriage returns between

lines) these numbers would appear on a grid with square unit cell, and could be contoured directly, on the result sheet.

### A Technical Feature in Polynomial I

We describe here a technical feature in this program which might be of use to other programmers. The problem is that we are multiplying numbers rapidly decreasing in magnitude with  $l+k$  (the  $C_{lk}$ 's) by numbers rapidly increasing in magnitude with  $l+k$  ( $x^l y^k$ ) while the product is of a relatively constant order of magnitude, which must be in a form which we can add to other such products.

What we want is the product  $C_{lk} x^l y^k$  to be scale factored finally by  $10^{-2}$ . To preserve accuracy during the computation of the product we do the following:

1. Form  $x^l 2^{-15}$  and  $y^k 2^{-15}$  and then scale factor to  $x^l 2^{-15+\alpha}$  and  $y^k 2^{-15+\beta}$  by use of the scale factor order.

$$2. \text{ Form } C_{lk} 10^{(l+k)-2} x^l 2^{-15+\alpha} y^k 2^{-15+\beta} \\ = C_{lk} x^l y^k 2^{-30+\alpha+\beta} 10^{(l+k)-2} \quad (1)$$

To get this product to the form  $C_{lk} x^l y^k 10^{-2}$  we must multiply by  $10^{-(l+k)} 2^{30-(\alpha+\beta)}$

It appears that we merely need to store the negative powers of 10, multiply the expression (1) by  $10^{-(l+k)}$  and then shift left  $30-(\alpha+\beta)$ . However the negative powers of 10 cannot be stored with any accuracy for high  $l+k$  so we write  $10^{-(l+k)} 2^{30-(\alpha+\beta)}$  in the form

$$\begin{aligned}
& 2^{-(l+k)\log_2 10 + 30 - (\alpha + \beta)} \\
= & 2^{-3.32193(l+k) + 30 - (\alpha + \beta)} \\
= & \frac{1}{2^{3.32193(l+k)}} [2^{-3(l+k) + 30 - (\alpha + \beta)}] \\
= & \frac{1}{(2^{3.32193})^{(l+k)}} [2^{-3(l+k) + 30 - (\alpha + \beta)}] \\
= & (.800)^{l+k} [2^{-3(l+k) + 30 - (\alpha + \beta)}]
\end{aligned}$$

We can store the powers of (.800) with ample accuracy. Thus we multiply by the appropriate power of (.800) and follow this by a shift left or right according to the exponent of 2. (The zeroth power of (.800) is put in as +.9999.)

$m = 4$

+0385	-0302	-0008	-0147	-0094	-0136	+0082	+0153	+0209	+0173	-0146
+0163	+0004	-0325	+0062	-0049	-0193	-0007	-0045	+0031	+0001	-0120
+0086	-0132	-0082	+0183	+0553	+0301	-0056	-0116	+0063	+0069	-0078
+0039	-0083	-0097	+0250	+0199	-0115	-0176	-0122	-0003	+0017	-0030
-0180	-0195	+0032	+0311	-0058	-0043	-0014	-0158	-0018	+0049	-0021
<del>-0203</del>	<del>-0151</del>	<del>+0025</del>	<del>+0141</del>	<del>+0050</del>	<del>-0007</del>	<del>-0247</del>	<del>-0250</del>	<del>+0145</del>	<del>+0141</del>	<del>+0074</del>
-0145	-0164	-0079	+0376	+0563	+0179	-0186	-0208	-0220	+0084	+0149
+0099	-0028	+0070	+0071	+0138	+0175	-0222	-0124	-0158	+0037	+0165
+0249	+0225	+0098	-0149	-0151	-0045	-0231	-0123	-0041	+0030	+0098
+0095	-0060	-0004	-0097	-0264	-0141	-0070	-0010	+0173	+0104	+0044
-0157	+0417	-0482	-0112	+0106	+0097	+0220	+0224	+0350	+0125	-0583

$m = 3$

+0372	-0293	-0001	-0172	-0181	-0300	-0154	-0115	-0007	+0146	+0223
+0193	+0085	-0211	+0179	+0035	-0166	-0051	-0140	-0058	+0021	+0175
+0067	-0087	+0020	+0318	+0684	+0393	-0027	-0150	-0000	+0050	+0078
-0057	-0114	-0057	+0343	+0313	-0021	-0132	-0145	-0081	-0065	-0021
-0333	-0299	+0003	+0350	+0019	+0036	+0027	-0179	-0108	-0082	-0124
-0363	-0292	-0050	+0138	+0100	+0063	-0196	-0253	+0070	+0001	-0084
-0249	-0296	-0170	+0355	+0605	+0259	-0108	-0169	-0248	-0018	-0002
+0107	-0112	-0008	+0047	+0181	+0270	-0108	-0030	-0117	+0056	+0070
+0404	+0205	+0033	-0184	-0122	+0046	-0100	+0012	+0062	+0075	+0079
+0390	-0041	-0095	-0193	-0309	-0115	+0014	+0106	+0288	+0186	+0076
+0212	+0390	-0700	-0382	-0131	-0067	+0133	+0200	+0358	+0134	-0595

$m = 2$

+0105	-0402	-0035	-0189	-0217	-0368	-0242	-0189	-0009	+0297	+0632
+0044	+0079	-0158	+0231	+0048	-0203	-0138	-0240	-0115	+0085	+0464
-0000	-0019	+0136	+0421	+0731	+0369	-0115	-0273	-0104	+0041	+0264
-0075	-0001	+0099	+0478	+0383	-0033	-0224	-0288	-0224	-0134	+0082
-0333	-0164	+0181	+0504	+0104	+0030	-0065	-0334	-0277	-0192	-0080
-0372	-0158	+0130	+0298	+0192	+0063	-0288	-0413	-0111	-0131	-0075
-0293	-0185	-0001	+0510	+0699	+0264	-0192	-0323	-0426	-0152	-0002
+0003	-0043	+0134	+0190	+0273	+0283	-0179	-0165	-0274	-0056	+0088
+0218	+0215	+0138	-0061	-0033	+0070	-0148	-0039	-0054	+0007	+0147
+0102	-0106	-0039	-0093	-0222	-0073	+0002	+0054	+0234	+0191	+0225
-0195	+0239	-0698	-0309	-0042	-0000	+0169	+0217	+0391	+0243	-0328

00			100	0.46000	
01			101	0.66000	
02			102	0.56000	
03			103	0.06000	
04			104	1.54000	
05			105	0.20000	
06			106	1.22000	
07			107	+3	
10			110	ai 0	
11			111	ai 0	
12			112	ai 0	
13			113	ai 0	
14			114	ca 123	
15			115	ca 121	
16			116	ca 116	
17			117	ca 107	
20			120	ca 45	
21			121	0.92000	
22			122	0.72000	
23			123	ca 122	
24			124	ca 116	
25			125	ai 0	
26			126	ai 0	
27			127	ca 441	
30			130	ca 382	
31			131	ca 441	
32			132	ca 317	
33			133	ca 112	
34			134	ca 442	
35			135	ca 442	
36			136	ca 321	
37			137	ca 442	
40	ca 63		140	ca 325	
41	ca 70		141	ca 341	
42	ca 225		142	ca 300	
43	ca 113		143	ca 325	
44	ca 114		144	ca 326	
45	ca 110		145	ca 317	
46	ca 113		146	ca 320	
47	ca 71		147	ca 326	
50	ca 72		150	ca 312	
51	ca 54		151	ca 326	
52	ca 17		152	ca 322	
53	ca 113		153	ca 312	
54	ca		154	ca 323	
55	ca 55		155	ca 323	
56	ca 110		156	ca 312	
57	ca 46		157	ca 324	
60	ca 105		160	ca 327	
61	ca 61		161	ca 331	
62	ca 112		162	ca 230	
63	ca		163	ca 440	
64	ca 106		164	ca 313	
65	ca 65		165	ca 314	
66	ca 111		166	ca 315	
67	ca 112		167	ca 316	
70	ca		170	ca 340	
71	+10		171	ca 264	
72	0.00073		172	ca 364	
73	1.74000		173	ca 233	
74	0.52000		174	ca 0	
75	0.36000		175	ca 343	
76	0.16000		176	ca 345	
77	0.26000		177	ca 376	

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104	1.54000		104	1.54000	
105	0.20000		105	0.20000	
106	1.22000		106	1.22000	
107	+3		107	+3	
110	ai 0		110	ai 0	
111	ai 0		111	ai 0	
112	ai 0		112	ai 0	
113	ai 0		113	ai 0	
114	ca 123		114	ca 123	
115	ca 121		115	ca 121	
116	ca 116		116	ca 116	
117	ca 107		117	ca 107	
120	ca 45		120	ca 45	
121	0.92000		121	0.92000	
122	0.72000		122	0.72000	
123	ca 122		123	ca 122	
124	ca 116		124	ca 116	
125	ai 0		125	ai 0	
126	ai 0		126	ai 0	
127	ca 441		127	ca 441	
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162	ca 230		162	ca 230	
163	ca 440		163	ca 440	
164	ca 313		164	ca 313	
165	ca 314		165	ca 314	
166	ca 315		166	ca 315	
167	ca 316		167	ca 316	
170	ca 340		170	ca 340	
171	ca 264		171	ca 264	
172	ca 364		172	ca 364	
173	ca 233		173	ca 233	
174	ca 0		174	ca 0	
175	ca 343		175	ca 343	
176	ca 345		176	ca 345	
177	ca 376		177	ca 376	

REG 440 + n	452	C <sub>01</sub>	REG 540	g (-N, M) x 10 <sup>2</sup>
441 + N	453	C <sub>11</sub>	541	g (-N+1, M)
442 + M		476 C <sub>02</sub>		
443 C <sub>02</sub>	760	C <sub>02</sub>		j (-N, M-1)
444 C <sub>02</sub>		are +10 (k+1)-2		g (-N+1, M-1)

200	ca 253		200	ca 0	
201	ca 324		201	ca 40	
202	ca 341		202	ca 0	
203	ca 350		203	ca 330	
204	ca 327		204	ca 310	
205	ca 347		205	ca 320	
206	ca 351		206	ca 155	
207	ca 343		207	ca 321	
210	ca 343		210	ca 145	
211	ca 353		211	ca 0	
212	ca 17		212	+1	
213	ca 352		213	ai 0	
214	ca 346		214	ai 0	
215	ca 347		215	ai 0	
216	ca 223		216	ai 0	
217	ca 373		217	ai 0	
220	ca 373		220	ai 0	
221	ca 234		221	ai 0	
222	ca 230		222	ai 0	
223	ca 375		223	ai 0	
224	ca 17		224	ai 0	
225	ca 374		225	ai 0	
226	ca 374		226	ai 0	
227	ca 234		227	ai 0	
230	ca 0		230	ai 0	
231	ca 443		231	ca 443	
232	ca 452		232	ca 452	
233	ca 460		233	ca 460	
234	ca 465		234	ca 465	
235	ca 471		235	ca 471	
236	ca 474		236	ca 474	
237	ca 476		237	ca 476	
240	ca 332		240	ca 332	
241	ca 17		241	ca 540	
242	ca 324		242	ai 0	
243	ca 230		243	ai 0	
244	ca 233		244	ai 0	
245	ca 343		245	ai 0	
246	ca 314		246	ai 0	
247	ca 201		247	ai 0	
250	ca 345		250	ai 0	
251	ca 0		251	ai 0	
252	ca 343		252	+30	
253	ca 0		253	+3	
254	ca 256		254	ai 0	
255	ca 0		255	+9999	
256	ca 233		256	+8000	
257	ca 253		257	+6400	
260	ca 327		260	+5120	
261	ca 326		261	+4096	
262	ca 17		262	+3276	
263	ca 327		263	+2621	
264	ca 312		264	ca 355	
265	ca 324		265	ca 356	
266	ca 0		266	ca 357	
267	ca 271		267	ca 360	
270	ca 0		270	ca 361	
271	ca 230		271	ca 362	
272	ca 264		272	ca 363	
273	ca 313		273	ca 0	
274	ca 314		274	ca 0	
275	ca 316		275	-1	
276	ca 201		276	ca 365	
277	ca 230		277		

200	ca 0		200	ca 0	
201	ca 40		201	ca 40	
202	ca 0		202	ca 0	
203	ca 330		203	ca 330	
204	ca 310		204	ca 310	
205	ca 320		205	ca 320	
206	ca 155		206	ca 155	
207	ca 321		207	ca 321	
210	ca 145		210	ca 145	
211	ca 0		211	ca 0	
212	+1		212	+1	
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214	ai 0		214	ai 0	
215	ai 0		215	ai 0	
216	ai 0		216	ai 0	
217	ai 0		217	ai 0	
220	ai 0		220	ai 0	
221	ai 0		221	ai 0	
222	ai 0		222	ai 0	
223	ai 0		223	ai 0	
224	ai 0		224	ai 0	
225	ai 0		225	ai 0	
226	ai 0		226	ai 0	
227	ai 0		227	ai 0	
230	ai 0		230	ai 0	
231	ca 443		231	ca 443	
232	ca 452		232	ca 452	
233	ca 460		233	ca 460	
234	ca 465		234	ca 465	
235	ca 471		235	ca 471	
236	ca 474		236	ca 474	
237	ca 476		237	ca 476	
240	ca 332		240	ca 332	
241	ca 540		241	ca 540	
242	ai 0		242	ai 0	
243	ai 0		243	ai 0	
244	ai 0		244	ai 0	
245	ai 0		245	ai 0	
246	ai 0		246	ai 0	
247	ai 0		247	ai 0	
250	ai 0		250	ai 0	
251	ai 0		251	ai 0	
252	+30		252	+30	
253	+3		253	+3	
254	ai 0		254	ai 0	
255	+9999		255	+9999	
256	+8000		256	+8000	
257	+6400		257	+6400	
260	+5120		260	+5120	
261	+4096		261	+4096	
262	+3276		262	+3276	
263	+2621		263	+2621	
264	ca 355		264	ca 355	
265	ca 356		265	ca 356	
266	ca 357		266	ca 357	
267	ca 360		267	ca 360	
270	ca 361		270	ca 361	
271	ca 362		271	ca 362	
272	ca 363		272	ca 363	
273	ca 0		273	ca 0	
274	ca 0		274	ca 0	
275	-1		275	-1	
276	ca 365		276	ca 365	
277			277</		



APPENDIX B

APPENDIX B

PREDICTION XV (2539 m 2) - Description and Use Of

This program was written to provide computational facility for predicting a series  $x_i$  ( $y_i, z_i$ , or  $u_i$ ) with a linear operator of the general form

$$\begin{aligned} x_{i+k} (y_{i+k}, z_{i+k} \text{ or } u_{i+k}) &= a + a_0 x_i + a_1 x_{i-1} \dots + a_M x_{i-M} \\ &+ b_0 y_i \dots \dots \dots + b_M y_{i-M} \\ &+ c_0 z_i \dots \dots \dots + c_M z_{i-M} \\ &+ d_0 u_i \dots \dots \dots + d_M u_{i-M} \end{aligned}$$

where the prediction distance  $k$  and the number of lags  $M$  are arbitrary but have the restrictions that  $M \geq 7$  and  $M+k \leq 19$ . The four series which this program handles contain 500 members each so that  $i$  ranges from 0 through 499. This first prediction computed is for  $i+k = 20$  and the last one for  $i+k = 499$ . After doing this computation the program forms the running average of squared errors between the predicted and actual series

$$\sum_{i=j-5}^{j+4} (x_i - \hat{x}_i)^2 \quad j = 25, 35, 45, 55, \dots, 495$$

which is called the "Error Curve".

There is considerable choice of output. The alternatives are any combination of or none of the following:

1. Print-out of the errors and sums of squares
2. Print-out of just the sums of squares of errors.

### 3. Photographs of oscilloscope displays of the sums of errors squares

An additional choice is the use of magnetic tape delayed output for 1 and 2 above, which is about fifteen times faster than direct print-out.

This program handles up to eight operators at a time in the above fashion. When the magnetic tape output is used, the error curves can be removed from the machine at the rate of one every ten seconds whereas the individual errors and error curves require fifty seconds for each operator. Each curve would represent about a week of hand computation. Once the computations are completed the individual errors  $(x_1 - x_1)$  for all operators are left in magnetic drum storage so other programs can use them for different types of averaging processes rather than just the Error Curve as described above.

On the next three pages are illustrated the various output forms. The first page is a reproduction of the individual errors and error curves for two operators. The results for each operator appear as a block of numbers 10 by 48 and a right-hand column of 48 numbers. The block represents the 480 individual errors whereas the right-hand column is the Error Curve, each member representing the sums of the squares of the 10 individual errors in the corresponding row to the left. The number appearing over the upper left corner of each block is a number assigned to the particular operator for identification purposes, and is printed by the

program. The number printed over the center of each block was inserted later.

The next page shows the output form for four operators when just the Error Curve is desired. The +0000 identification number indicates that the operator was chosen as the variance operator which has the form  $x_1 = \bar{x}$  (means of series). The Error Curve for this type operator becomes the sample variance curve and provides a basis for testing the statistical significance of other operators predicting the  $x_1$  series.

The third page is a photograph taken automatically by the program of an oscilloscope display of one-half of an Error Curve. A vertical and horizontal axis are also displayed.



+0000

~~+0454 +0375 +0437 +0366 +0605 +0253 +0518 +0435 +0273 +0648 +0530 +0342~~  
~~+0339 +0569 +0354 +0465 +0418 +0611 +0357 +0348 +0422 +0288 +0674 +0390~~  
~~+0396 +0383 +0422 +0384 +0343 +0468 +0429 +0512 +0332 +0355 +0775 +0256~~  
~~+0483 +0423 +0313 +0593 +0216 +0395 +0399 +0399 +0399 +0399 +0399 +0399~~

+0000

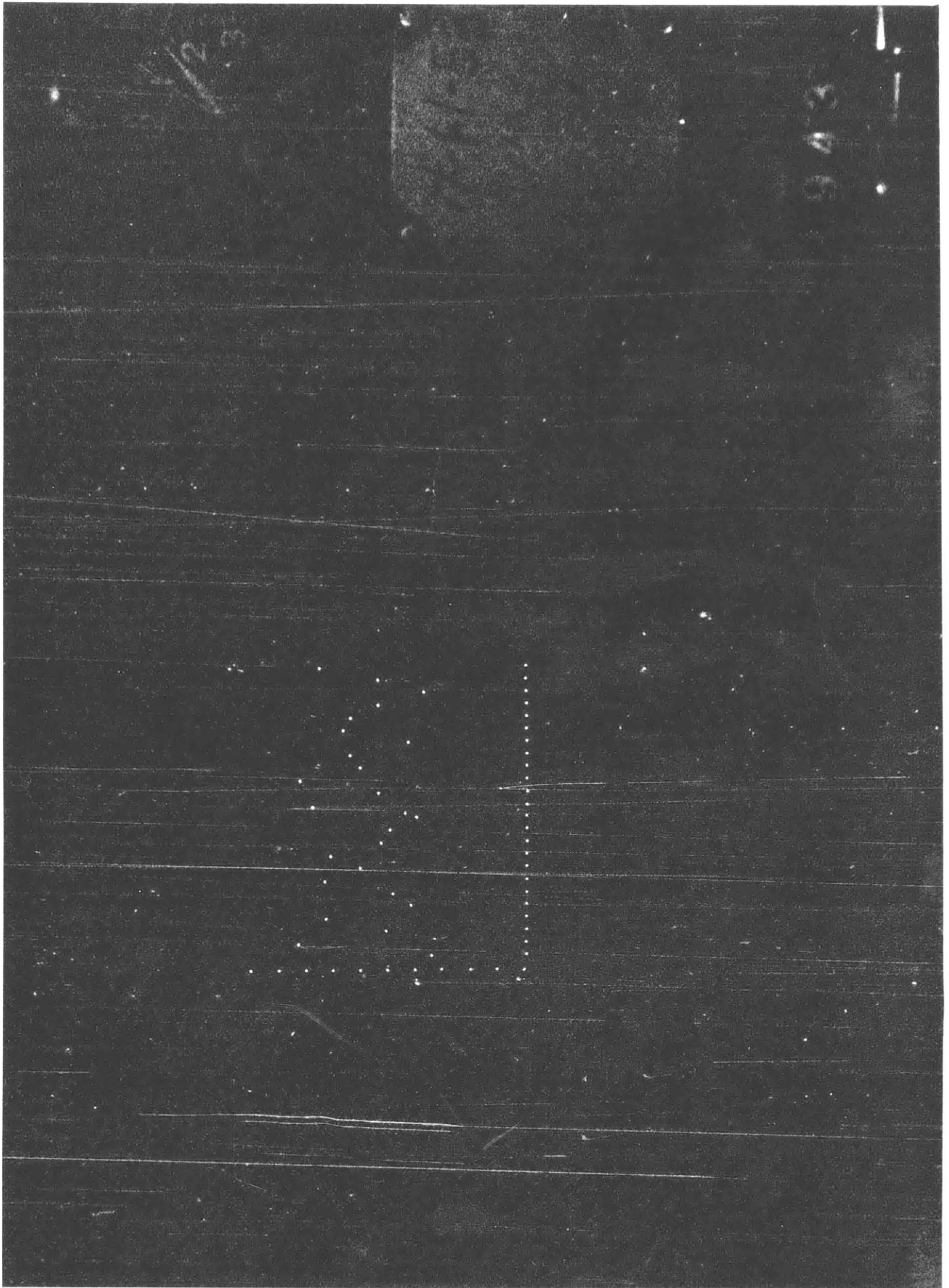
~~+0816 +1241 +0509 +1182 +0781 +0725 +0995 +0784 +0681 +0864 +0902 +0783~~  
~~+0875 +0728 +0834 +0761 +1013 +0594 +1030 +0871 +0437 +1340 +0646 +0771~~  
~~+1129 +0536 +0977 +0687 +0949 +0520 +0853 +1589 +0155 +1276 +0878 +0532~~  
~~+1091 +0550 +0848 +1040 +0384 +0900 +0900 +0900 +0900 +0900 +0900 +0900~~

+0000

~~+0373 +0337 +0589 +0437 +0484 +0318 +0519 +0374 +0431 +0507 +0351 +0485~~  
~~+0391 +0394 +0468 +0431 +0386 +0442 +0423 +0337 +0442 +0395 +0539 +0328~~  
~~+0359 +0406 +0510 +0307 +0372 +0408 +0413 +0437 +0326 +0395 +0427 +0326~~  
~~+0388 +0506 +0279 +0451 +0424 +0362 +0399 +0399 +0399 +0399 +0399 +0399~~

+0000

~~+0800 +0713 +0818 +0886 +0927 +0833 +0680 +0949 +0806 +0843 +0703 +0846~~  
~~+0691 +0887 +0893 +0804 +0659 +0938 +0780 +0797 +0695 +0823 +1077 +0800~~  
~~+0653 +0698 +0877 +0663 +0758 +0905 +0747 +0802 +0603 +0802 +0952 +0696~~  
~~+0770 +0613 +0856 +0755 +0855 +0677 +0784 +0784 +0784 +0784 +0784 +0784~~



Use of Prediction XV

It is necessary to prepare a tape containing the operators and a tape containing the traces  $x_1, y_1, z_1, u_1$ . These are prepared as described in the following two pages. Assume these are given tape numbers X and Y respectively. Then the operating instruction would be:

- Erase storage, put S11 switch off
- Read in 2539 m 2
- Read in 2539 P — (Control Tape)
- Read in X
- Place Y in Photoelectric Reader
- Start at 145

The control tapes control the output and serve the following functions:

- |            |  |                                  |
|------------|--|----------------------------------|
| 2539 - P0  | Print errors and sums of squares and scope display sums of squares |                                  |
| 2539 - P1  | Print sums of squares and scope display sums of squares            |                                  |
| 2539 - P2  | Scope display sums of squares                                      | DIRECT PRINTOUT                  |
| 2539 - P3  | Print errors and sums of squares                                   |                                  |
| 2539 - P4  | Print sums of squares  |                                  |
| 2539 - P5  | Print nothing, display nothing                                     |                                  |
| 2539 - P10 | Print errors and sums of squares and scope display sums of squares |                                  |
| 2539 - P11 | Print sums of squares and scope display sums of squares            | DELAYED PRINTOUT (MAGNETIC TAPE) |
| 2539 - P12 | Print errors and sums of squares                                   |                                  |
| 2539 - P13 | Print sums of squares  |                                  |



If one of the operators on X were badly prepared, it might happen that machine overflow would occur causing the machine to stop while computing for that operator. If this does happen, starting the machine over at 166 will have the effect of ignoring the bad operator and proceeding to the remaining ones.

Preparation of Data Parameter

Each set of data  $x_1, y_1, z_1,$  and  $u_1$  is prepared as a separate parameter and then the four parameters are combined into one long one. The form of each is identical.

Octal Address	$\bar{x}$	Octal Address	$\bar{y}$	Octal Address	$\bar{z}$	Octal Address	$\bar{u}$
1054	$x_1$	1054	$y_1$	1054	$z_1$	1054	$u_1$
1055	$x_2$	1055	$y_2$	1055	$z_2$	1054	$u_2$
1056		1056		1056		1056	
↓	↓	↓	↓	↓	↓	↓	↓
2037	$x_{499}$	2037	$y_{499}$	2037	$z_{499}$	2037	$u_{499}$
Start at 1033		Start at 1033		Start at 1033		Start at 1033	

Notes:

It is not necessary that the series contain 499 members. However, there must be four traces. If less than four are to be used, short dummy traces must be inserted. For consistency with the operator tape, the order of combination of the separate parameters must be  $x_1, y_1, z_1, u_1$ .

The data must appear as integers in the range -99 through +99.





400	ca 524		500	tl 532	
401	tl 410		501	ca 434	
402	ca 225		502	tl 435	
403	tl 423		503	sp 774	
404	ca 434		504	tl 440	
405	ca 435		505	ca 542	
406	ca 432		506	ca 543	
407	ca 433		507	ca 437	
410	ca		510	mp 547	
411	sp 40		511	ca 12	
412	ca 410		512	ca 600	
413	ca 433		513	ca 543	
414	ca 410		514	ca 514	
415	ca 107		515	ca 543	
416	ca 416		516	ca 544	
417	sp 135		517	ca 543	
420	ca 107		520	ca 520	
421	ca 421		521	ca 545	
422	sp 135		522	ca 515	
423	ca		523	ca 440	
424	sp 40		524	ca 440	
425	ca 423		525	mp 547	
426	ca 435		526	ca 12	
427	ca 406		527	ca 600	
430	ca 630		530	ca 542	
431	ca		531	ca 531	
432	ca		532	ca	
433	ca 0		533	ca 1	
434	ca 23		534	ca 534	
435	ca 0		535	ca 532	
436	ca 10		536	ca 435	
437	ca 77747		537	ca 523	
440	ca 77747		540	ca 500	
441	ca 415		541	ca	
442	ca 225		542	ca 15777	
443	ca 130		543	ca 15777	
444	ca 110		544	ca 06315	
445	ca 445		545	ca 06000	
446	sp 135		546	ca 02000	
447	ca 1007		547	ca 66667	
450	sp 40		550	ca 1053	
451	ca 110		551	ca 170	
452	ca 452		552	ca 170	
453	sp 135		553	ca 1047	
454	sp 460		554	ca 583	
455	ca 475		555	ca 606	
456	ca 225		556	ca 357	
457	sp 130		557	ca 373	
460	ca 434		560	ca 442	
461	ca 113		561	ca 456	
462	ca 114		562	ca 111	
463	ca 225		563	ca 1050	
464	ca 467		564	ca 574	
465	ca 434		565	ca 583	
466	ca 435		566	ca 574	
467	ca		567	ca 575	
470	sp 40		570	ca 571	
471	ca 467		571	ca 605	
472	ca 435		572	ca 583	
473	ca 467		573	ca 146	
474	ca 630		574	ca 107	
475	ca		575	ca 154	
476	ca 541		576	ca 64	
477	ca 225		577	ca 162	

NOTES

DL-481

600	ca 565		700	ca 711	
601	ca 155		701	ca 1004	
602	ca 156		702	ca	
603	ca 157		703	ca 770	
604	ca 6		704	ca	
605	ca -2		705	ca	
606	ca 136		706	ca 702	
607	ca 170		707	ca 704	
610	ca 151		710	ca 705	
611	ca 0		711	ca 1004	
612	ca 616		712	ca 702	
613	ca 613		713	ca 233	
614	ca 620		714	ca 1003	
615	ca 0		715	ca 711	
616	ca 2000		716	ca 1004	
617	ca 223		717	ca 1012	
620	ca 146		720	ca 1005	
621	ca 746		721	ca	
622	ca 1001		722	ca	
623	ca 746		723	ca	
624	ca 625		724	ca	
625	ca 1010		725	ca 721	
626	ca 1003		726	ca 722	
627	ca 760		727	ca 1005	
630	ca 670		730	ca 721	
631	ca 761		731	ca 721	
632	ca 702		732	ca 772	
633	ca 1011		733	ca 721	
634	ca 1012		734	ca 722	
635	ca 763		735	ca 773	
636	ca 762		736	ca 722	
637	ca 704		737	ca 722	
640	ca 705		740	ca 724	
641	ca 754		741	ca 1004	
642	ca 1002		742	ca 717	
643	ca 764		743	ca 752	
644	ca 755		744	ca 624	
645	ca 1012		745	ca 1002	
646	ca 765		746	ca	
647	ca 756		747	ca 753	
650	ca 1012		750	ca 624	
651	ca 766		751	ca 771	
652	ca 757		752	ca 666	
653	ca 1012		753	ca 625	
654	ca 767		754	ca 1023	
655	ca 244		755	ca 1032	
656	ca 0		756	ca 1043	
657	ca 1012		757	ca 1054	
660	ca 772		760	ca 764	
661	ca 1012		761	ca 1100	
662	ca 773		762	ca 1460	
663	ca 773		763	ca 0	
664	ca 330		764	ca 0	
665	ca 1002		765	ca 0	
666	ca 763		766	ca 0	
667	ca 771		767	ca 0	
670	ca		770	ca 1000	
671	ca 722		771	ca -239	
672	ca 762		772	ca 0	
673	ca 723		773	ca 0	
674	ca 424		774	ca 500	
675	ca 670		775	ca 437	
676	ca 1003		776	ca 504	
677	ca 715		777	ca 0	

MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE PREDICTION INDEX  
 AUTHOR SIMPSON DATE  
 TAPE NUMBER 2539 m 2

100	ai' 0			1100	
101	+2			1101	
102	ca 1001	A1		1102	
103	ai 707			1103	
104	ca 1010			1104	
105	ca 1053			1105	
106	ap 147			1106	
107	+4876			1107	
108	+297			1108	
109	ca 1025	A2		1109	
110	ai 707			1110	
111	ca 1074			1111	
112	ca 1054			1112	
113	ca 1025			1113	
114	ca 1024			1114	
115	ca 1025			1115	
116	ca 1026			1116	
117	ca 1023			1117	
118	ca 1023			1118	
119	ca 1023			1119	
120	ca 1023			1120	
121	ca 1023			1121	
122	ca 1023			1122	
123	ca 1023			1123	
124	ca 1023			1124	
125	ca 1023			1125	
126	ca 1023			1126	
127	ca 1023			1127	
128	ca 1023			1128	
129	ca 1023			1129	
130	ca 1023			1130	
131	ca 1023			1131	
132	ca 1023			1132	
133	ca 1023			1133	
134	ca 1023			1134	
135	ca 1023			1135	
136	ca 1023			1136	
137	ca 1023			1137	
138	ca 1023			1138	
139	ca 1023			1139	
140	ca 1023			1140	
141	ca 1023			1141	
142	ca 1023			1142	
143	ca 1023			1143	
144	ca 1023			1144	
145	ca 1023			1145	
146	ca 1023			1146	
147	ca 1023			1147	
148	ca 1023			1148	
149	ca 1023			1149	
150	ca 1023			1150	
151	ca 1023			1151	
152	ca 1023			1152	
153	ca 1023			1153	
154	ca 1023			1154	
155	ca 1023			1155	
156	ca 1023			1156	
157	ca 1023			1157	
158	ca 1023			1158	
159	ca 1023			1159	
160	ca 1023			1160	
161	ca 1023			1161	
162	ca 1023			1162	
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166	ca 1023			1166	
167	ca 1023			1167	
168	ca 1023			1168	
169	ca 1023			1169	
170	ca 1023			1170	
171	ca 1023			1171	
172	ca 1023			1172	
173	ca 1023			1173	
174	ca 1023			1174	
175	ca 1023			1175	
176	ca 1023			1176	
177	ca 1023			1177	

OPERATOR ONE

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OPERATOR FOUR

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OPERATOR SEVEN

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OPERATOR ELEVEN

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OPERATOR TEN THOUSAND SIX

OPERATOR TEN THOUSAND SEVEN

OPERATOR TEN THOUSAND EIGHT

OPERATOR TEN THOUSAND NINE

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NOTES

MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE PREDICTION XV INDEX \_\_\_\_\_  
 AUTHOR SIMPSON DATE \_\_\_\_\_  
 TAPE NUMBER 2539 m 2

APPENDIX C

## APPENDIX C

### ITERATION I (2615 m 2) - Description and Use Of

This program was written with the purpose of obtaining least squares fits for linear operators as described in Part IV. It computes essentially as described, but has provisions for changing its increment after cycling for a certain prescribed number of times.

The program was designed to be run in conjunction with the Prediction XV program described in Appendix B, and to illustrate the conveniences which programs can include. The data to which the linear operator is to be fitted is prepared in the same fashion as in Prediction XV. The information about the operators to be found (Iteration I solves up to eight operators one after the other) is prepared as a single tape. The operator coefficients once formed are printed out, and also are left in the machine in a form to be used directly with the prediction program.

The output of Iteration I was designed to eliminate identification problems. In addition to printing out the coefficients identified, it prints out the operator number, the operator parameters including which set of data is predicted, and the variance and minimum sums of square errors. These last two numbers

allow a rapid computation of the percent reduction

$R = 1 - I_{\min}/I_{\text{var}}$ , which is a measure of the goodness of the least squares fit. A sample of the output appears below.

```
Variance sum = 004953
Minimum sum = 001840
Operator No. -1010
N = 066
n = 050
k = 002
M = 003
T4 predicted
a/1000 = +0292

a3/10 = -0000
a2/10 = +0565
a1/10 = -0000
a0/10 = +0005

d3/10 = -0341
d2/10 = -0117
d1/10 = -0410
d0/10 = +0708
```

The program may be used to print out all the values of I as they are computed. A plot of these values for a particular operator appears in Part IV.

One other feature in this program is a "roll back" procedure. This permits us to avoid having to start from scratch if the machine fails in the middle of the long computations. Every fifteen seconds during the



computation, all of electrostatic storage is transferred to the magnetic drums which are very reliable. Then if electrostatic storage is destroyed, we can call back the program from the magnetic drums and start over where we left off not more than fifteen seconds ago.

#### Use of Iteration I

If the operators are prepared as described on the next page with a tape number X then the instructions for the operation of this program would be:

Erase storage, S11 switch down  
Read in 2615 m 2  
Read in 2615 P (control tape)  
Read in Y (data tape)  
Place X in photoelectric reader  
Start over at 145

The control tapes control the output and serve the following functions:

2615 P 0 Print out operator and identification (direct printer)  
2615 P 1 Print out operator, identification, and all values of I (delayed printer).

## Preparation of Operator Parameters

The information for each operator is prepared as a short separate tape and the tapes are then combined in any order. The form of each operator is identical.

Octal Address	Contents	Explanation
1001	+N	First member op. interval
1002	+n	Length of op. interval
1003	+0, or -1	-1 if $x_1$ not used
1004	+0, or -1	-1 if $y_1$ " "
1005	+0, or -1	-1 if $z_1$ " "
1006	+0, or -1	-1 if $u_1$ " "
1007	+.XXXX	First operator no.
1010	-0, -1, -2, or -3	-0 if $x_1$ , -1 if $y_1$
1111	+k	-2 if $z_1$ , -3 if $u_1$
1112	+M	
1113	+.XXXX	= mean of pred. series $\times 10^{-3}$

Start at 147

The first of the separate tapes must have one additional register, register 1000, which contains + no. of operators on the combined tape.

Register 2123 contains +.0010 which is the first increment to be used for the a term. Register 2223 contains +.0100 which is the first increment to be used with the remaining constants. Register 3421 is the counter for the cycles at these increments. The second set of increments is  $1/10$  the first set, and appears in registers 2124 and 2224. The counter for this set is register 3433. These registers may be changed to adapt to the particular problem.

The "roll back" procedure in case of electrostatic storage failure is

Erase storage

Read in 2615 P 13

Start over at 145

00		100	0.6000
01		101	0.26000
02		102	0.46000
03		103	0.66000
04		104	0.86000
05		105	0.06000
06		106	1.58000
07		107	0.20000
10		110	1.22000
11		111	+3
12		112	AL O
13		113	AL O
14		114	AL O
15		115	AL O
16		116	CP 126
17		117	CA 124
20		120	AC 120
21		121	AP 135
22		122	CA 111
23		123	AP 44
24		124	0.32000
25		125	0.72000
26		126	CA 125
27		127	AP 120
30		130	AL 134
31		131	CA 142
32		132	AL 144
33		133	CA 132
34		134	AP
35		135	AL 141
36		136	CA 143
37		137	AL 144
40	CA 64	140	CA 137
41	CA 72	141	AP
42	CA 115	142	-140
43	CA 116	143	128
44	CA 112	144	+1
45	CA 115	145	AP 146
46	CA 73	146	AP 25
47	CA 74	147	AP 3232
50	CA 53	150	AP 261
51	CA 17	151	AP 617
52	CA 115	152	AP 2614
53	CA	153	AP 2100
54	CA 58	154	AP 2160
55	AP 135	155	AP 2204
56	AP 112	156	AP 2227
57	AP 95	157	AP 2252
60	CA 107	160	AP 3428
61	CA 61	161	AP 3425
62	AP 135	162	AP 163
63	AP 114	163	AP 2670
64	CA	164	AP 2737
65	CA 110	165	AP 2316
66	CA 66	166	AP 2550
67	AP 135	167	AP 2560
70	CA 113	170	AP 2570
71	CA 114	171	AP 2600
72	AP	172	AP 225
73	+10	173	AP 600
74	0.00075	174	AP 217
75	1.74000	175	AL O
76	0.52000	176	AL O
77	0.36000	177	AL O

NOTES

SHEET 1

200	CA 215	300	AL O
201	AL 703	301	AL O
202	AL 202	302	AL O
203	AL 214	303	-3
204	AL 214	304	-3
205	AP 207	305	+500
206	AL O	306	AL O
207	CA 215	307	AL O
210	AL 707	310	CA 255
211	CA 214	311	AL 320
212	AL 212	312	AL 325
213	AL 173	313	CA 344
214	AL O	314	AL 323
215	+4076	315	AL O
216	AL O	316	CA 361
217	AL O	317	AL 342
220	CA 224	320	AL
221	AL 221	321	CA 345
222	AL 630	322	AL 17
223	AL O	323	AL
224	142000	324	CA 345
225	AL 630	325	AL 17
226	CA 237	326	AL 342
227	AL 707	327	AL 342
230	CA 240	330	CA 341
231	AL 1007	331	AL 341
232	AL 237	332	AL 320
233	AL 240	333	AL 323
234	AL 241	334	AL 343
235	AP 200	335	CA 320
236	AL O	336	AL 451
237	+4077	337	AP
240	+37	340	AL O
241	AL O	341	AL O
242	AL O	342	AL O
243	CA 256	343	AL O
244	AL 257	344	AL O
245	CA 255	345	+10
246	AL 250	346	AL O
247	CA 1015	347	AL 424
250	AL	350	CA 74
251	AP 250	351	AL 473
252	AP 257	352	AL 416
253	CA 247	353	CA 404
254	AP 657	354	AL 365
255	CA 1460	355	CA 402
256	AL O	356	AL 403
257	AL O	357	CA 401
260	AL O	360	AL 370
261	AL 77	361	CA 341
262	CA 301	362	AL 407
263	AL 703	363	AP 407
264	CA 300	364	CA 407
265	AL 1054	365	AL
266	CA 301	366	CA 372
267	AL 305	367	AL 407
270	AL 301	370	AL
271	AL 303	371	AL 364
272	CA 297	372	AL 365
273	AL 302	373	CA 370
274	AL 301	374	CA 370
275	CA 304	375	AL 370
276	AL 303	376	AL 403
277	AL	377	CA 364

MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE ITERATION I INDEX \_\_\_\_\_  
 AUTHOR SIMPSON DATE \_\_\_\_\_  
 TAPE NUMBER 2.615-REV 0

SHEET 2.

400	ac 411	400	ac 0
401	ca 413	401	ac 0
402	-	402	ac 0
403	ac 0	403	ac 0
404	ca 405	404	ac 0
405	+10	405	ca 573
406	+1	406	ac 709
407	ac 0	407	ca 514
410	ac 0	410	ca 40
411	ca 225	411	ac 516
412	ap 130	412	ac 0
413	ca	413	+6145
414	ca 414	414	11857
415	ap 135	415	ac 0
416	ca	416	ca 1047
417	ca 417	417	ca 535
420	ap 135	420	ca 522
421	ca 342	421	ca 136
422	ap 40	422	ca 521
423	ac 130	423	ca 217
424	ap	424	ca 411
425	ac 0	425	ca 2361
426	ac 0	426	ca 2376
427	ca 467	427	ca 2614
430	ca 434	430	ca 2670
431	ca 474	431	ca 535
432	ca 342	432	ac 0
433	ca 475	433	ac 0
434	ca 476	434	ac 0
435	ca 473	435	ca 1050
437	ca 473	437	ca 646
440	ca 474	440	ca 562
441	ca 457	441	ca 567
442	ca 466	442	ca 2114
443	ca 444	443	ca 570
444	ca 465	444	ca 2137
445	ca 477	445	ca 571
446	ca 475	446	ca 2147
447	ca 457	447	ca 572
448	ca 465	448	ca 1150
450	ac 0	450	ca 573
451	ca 341	451	ca 2214
452	ca 474	452	ca 574
453	ca 342	453	ca 2237
454	ca 475	454	ca 575
455	ca 337	455	ca 2262
456	ac 0	456	ca 576
457	ca 462	457	ca 2314
460	ca 500	460	ca 577
461	ca 467	461	ca 2324
462	+0	462	ca 604
463	ac 0	463	ca 707
464	ac 0	464	ca 0
465	ca 490	465	ca 585
466	ca 500	466	ac 0
467	ap	467	ca 2115
470	ca 1	470	ca 2144
471	+2	471	ca 2150
472	ac 0	472	ca 2171
473	ac 0	473	ca 2115
474	ac 0	474	ca 240
475	ac 0	475	ca 2263
476	ac 0	476	ca 2315
477	ac 0	477	ca 2344

NOTES

600	ca 177	700	ca 791
601	ca 146	701	ca 614
602	ca 174	702	ca
603	ac 0	703	ca 770
604	+4096	704	ca
605	ac 0	705	ca
606	ac 0	706	ca 702
607	+10	707	ca 704
610	-3	710	ca 705
611	+2	711	ca 614
612	ac 0	712	ca 702
613	ac 0	713	ca 607
614	ac 0	714	ca 613
615	ac 0	715	ca 771
616	ac 0	716	ca 614
617	ca 337	717	ca 1012
620	ca 746	720	ca 615
621	ca 746	721	ca
622	ca 611	722	ca
623	ca 746	723	ca
624	ca 625	724	ca
625	ca 1010	725	ca 721
626	ca 613	726	ca 722
627	ca 760	727	ca 615
630	ca 670	730	ca 721
631	ca 761	731	ca 721
632	ca 702	732	ca 772
633	ca 1011	733	ca 721
634	ca 1012	734	ca 702
635	ca 763	735	ca 773
636	ca 762	736	ca 722
637	ca 764	737	ca 723
640	ca 705	740	ca 724
641	ca 754	741	ca 614
642	ca 812	742	ca 717
643	ca 764	743	ca 752
644	ca 755	744	ca 624
645	ca 1012	745	ca 612
646	ca 765	746	ca
647	ca 756	747	ca 753
650	ca 1012	750	ca 624
651	ca 766	751	ca 310
652	ca 757	752	ca 625
653	ca 1012	753	ca 625
654	ca 767	754	ca 1023
655	ca 243	755	ca 1033
656	ac 0	756	ca 1043
657	ca 1012	757	ca 1053
660	ca 712	760	ca 914
661	ca 1012	761	ca 1106
662	ca 773	762	ca 1160
663	ca 773	763	ac 0
664	ca 610	764	ac 0
665	ca 612	765	ac 0
666	ca 613	766	ac 0
667	ca 721	767	ac 0
670	ca	770	ca 1000
671	ca 722	771	ac 0
672	ca 762	772	ac 0
673	ca 723	773	ac 0
674	ca 724	774	ac 0
675	ca 640	775	ac 0
676	ca 613	776	ac 0
677	ca 715	777	ac 0

MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE ITERATION 1 INDEX \_\_\_\_\_  
 AUTHOR SIMPSON DATE \_\_\_\_\_  
 TAPE NUMBER 2615-220

1000	AI 0	3500	AI 225
1001	AI 0	3501	CA 2756
1002	AI 0	3502	TC 3502
1003	AI 0	3503	AP 135
1004	AI 0	3504	AP 3577
1005	AI 0	3505	CP 3501
1006	AI 0	3506	CA 3513
1007	AI 0	3507	TA 3512
1010	AI 0	3510	AP 150
1011	CA 1035	3511	AI 0
1012	AI 707	3512	AP 3772
1013	CA 1024	3513	AP 3772
1014	TA 1054	3514	AI 0
1015	CA 1025	3515	AI 0
1016	TA 1024	3516	AI 0
1017	TA 1025	3517	AI 0
1020	CA 1026	3520	CA 513
1021	CA 1023	3521	AI 707
1022	AI 0	3522	CA 514
1023	AP 1027	3523	TA 40
1024	+500	3524	AP 2100
1025	+2049	3525	AI 0
1026	-3	3526	AI 0
1027	AP 25	3527	AI 0
1030	+0100	3530	CA 1013
1031	TA 2037	3531	AI 2123
1032	CA 1034	3532	TA 1013
1033	CA 1032	3533	CA 2126
1034	TA 1036	3534	CA 3531
1035	TA 1031	3535	CA 476
1036	CA	3536	TA 474
1037	TA 1030	3537	CA 477
1040	TA 17	3540	TA 475
1041	TA	3541	AP 154
1042	AI 1036	3542	AI 0
1043	AI 1041	3543	AI 0
1044	AI 1031	3544	AI 0
1045	CA 1036	3545	AI 0
1046	AP 1011	3546	AI 0
1047	AI 0	3547	AI 0
1050	AI 0	3550	CA 2355
1051	AI 0	3551	TA 2333
1052	AI 0	3552	CA 0
1053	AI 0	3553	AI 2322
1054		3554	AI 0
1055		3555	CA 2326
1056		3556	TA 3553
1057		3557	CA 476
1060		3560	TA 474
1061		3561	CA 477
1062		3562	TA 475
1063		3563	AP 2347
1064		3564	AI 0
1065		3565	TA 3571
1066		3566	TA 2354
1067		3567	TA 3552
1070		3570	TA 3554
1071		3571	AP 0
1072		3572	CP 3551
1073		3573	AP 3553
1074		3574	AP 161
1075		3575	AI 0
1076		3576	AI 0
1077		3577	TA 2155

NOTES

3600	TA 3531	00	
3601	AP 3357	01	
3602	AI 0	02	
3603	TA 2355	03	
3604	TA 3533	04	
3605	AP 3353	05	
3606	AI 0	06	
3607	AI 0	07	
3610	CA 2355	10	
3611	TA 3533	11	
3612	AP 2300	12	
3613	AI 0	13	
3614	AI 0	14	
3615	TA 3621	15	
3616	AP 3552	16	
3617	AP 3554	17	
3620	AP 2277	20	
3621	AP 0	21	
3622	AI 0	22	
3623	AI 0	23	
24		24	
25		25	
26		26	
27		27	
30		30	
31		31	
32		32	
33		33	
34		34	
35		35	
36		36	
37		37	
40		2040	AP 3774
41		2041	CA 3452
42		2042	TA 2042
43		2043	AP 135
44		2044	AI 2041
45		2045	AI 2040
46		2046	CP 2041
47		2047	CA 3452
50		2050	TA 2080
51		2051	CA 2076
52		2052	TA 2041
53		2053	CA 1010
54		2054	TA 3451
55		2055	AI 3451
56		2056	AI 74
57		2057	TA 2060
60		2060	CA 0
61		2061	TA 2061
62		2062	AP 135
63		2063	AI 2060
64		2064	TA 2064
65		2065	AP 135
66		2066	AI 2063
67		2067	AI 3475
70		2070	AP 2063
71		2071	CA 3476
72		2072	TA 3475
73		2073	CA 2077
74		2074	TA 2063
75		2075	AI 165
76		2076	CA 3452
77		2077	CA 3460

MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE ITERATION 1 INDEX \_\_\_\_\_  
 AUTHOR SIMPSON DATE \_\_\_\_\_  
 TAPE NUMBER 2615-000

2100	ca 474	2200	ap 3615
2101	ca 476	2201	ap 2170
2102	ca 475	2202	ac 0
2103	ca 477	2203	ac 0
2104	ap 2105	2204	ca 1012
2105	ca 1013	2205	ca 2202
2106	ap 2123	2206	ca 2226
2107	ca 1013	2207	ca 2305
2110	ca 2127	2210	ca 2307
2111	ca 2104	2211	ca 2332
2112	ap 261	2212	ap 3565
2113	ca 617	2213	ap 2277
2114	ap 347	2214	ap 2260
2115	ca 427	2215	ca 2202
2116	ca 500	2216	ca 2220
2117	ca 2100	2217	ca 156
2120	ca 2126	2220	ca 2307
2121	ca 2133	2221	ca 2332
2122	ap 2151	2222	ca 2334
2123	ca 0010	2223	ca 2305
2124	ca 0001	2224	ca 3615
2125	ca 4	2225	ca 2214
2126	ca 2123	2226	ac 0
2127	ap 2132	2227	ca 1012
2130	ac 0	2230	ca 2202
2131	ac 0	2231	ca 2251
2132	ca 1013	2232	ca 2305
2133	ap 2123	2233	ca 2307
2134	ca 1013	2234	ca 2332
2135	ap 261	2235	ap 3565
2136	ap 617	2236	ca 2277
2137	ap 347	2237	ap 2360
2140	ap 427	2240	ca 2202
2141	ca 500	2241	ca 2243
2142	ca 2100	2242	ap 157
2143	ca 2158	2243	ca 2307
2144	ca 2104	2244	ca 2332
2145	ca 2155	2245	ca 2334
2146	ca 2133	2246	ca 2305
2147	ap 2360	2247	ap 3615
2150	ap 3536	2250	ap 2237
2151	ca 2155	2251	ac 0
2152	ca 3531	2252	ca 1012
2153	ap 2100	2253	ca 2202
2154	ap 2105	2254	ca 2274
2155	ap 2123	2255	ca 2305
2156	ac 0	2256	ca 2307
2157	ac 0	2257	ca 2332
2160	ca 1012	2260	ap 3565
2161	ca 2202	2261	ap 2277
2162	ca 2203	2262	ap 2360
2163	ca 2305	2263	ca 2202
2164	ca 2307	2264	ca 2266
2165	ca 2582	2265	ap 160
2166	ap 3565	2266	ca 2307
2167	ap 2277	2267	ca 2332
2170	ap 2360	2270	ca 2334
2171	ap 2202	2271	ca 2305
2172	ap 2174	2272	ap 3615
2173	ap 155	2273	ap 2266
2174	ca 2307	2274	ac 0
2175	ca 2332	2275	ac 0
2176	ca 2334	2276	ac 0
2177	ca 2301	2277	ca 347
NOTES			

DL-481

SHEET 4

2300	ca 474	2400	ca 2421
2301	ca 476	2401	ca 2401
2302	ca 475	2402	ap 135
2303	ca 477	2403	ca 2400
2304	ap 2305	2404	ca 2437
2305	ca 0	2405	ca 2400
2306	ap 2327	2406	ca 2440
2307	ca 0	2407	ca 2437
2310	ca 2327	2410	ca 2447
2311	ca 2304	2411	ca 2400
2312	ap 261	2412	ca 2442
2313	ap 617	2413	ca 113
2314	ap 347	2414	ca 114
2315	ap 427	2415	ca 1013
2316	ca 500	2416	ap 40
2317	ca 2300	2417	ca 2423
2320	ca 2326	2420	ca 2420
2321	ca 2333	2421	ap 135
2322	ap 3610	2422	ap 166
2323	ca 0100	2423	ca 2400
2324	ca 0010	2424	ca 14000
2325	ca 0001	2425	ca 162000
2326	ca 2323	2426	ca 52000
2327	ap 2332	2427	ca 172000
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2334	ca 0	2434	ca 20000
2335	ap 261	2435	ca 22000
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2337	ap 347	2437	-11
2340	ap 427	2440	-11
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2355	ca 2223	2455	ca 2437
2356	ac 0	2456	ca 2444
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2360	ca 2323	2460	ca 0
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2363	ca 363	2463	ca 2463
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2366	ca 2360	2466	ac 0
2367	ap 135	2467	ca 0
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2371	ca 114	2471	ca 22000
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2373	ca 0	2473	ca 52000
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2377	ca 130	2477	ca 24000
MIT DIGITAL COMPUTER LABORATORY OCTAL PROGRAM FORM			
TITLE ITERATION INDEX _____			
AUTHOR SIMPSON DATE _____			
TAPE NUMBER 1615 0000			

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2512	ca 1012	2512	ca 225
2513	ca 2535	2513	ca 2444
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2515	ca 2467	2515	ca 155
2516	ca 2445	2516	ca 2615
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2524	ca 2516	2524	ca 1644
2525	ca 14000	2525	ca 2721
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2527	ca 70000	2527	ca 262
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MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE ITERATION I INDEX  
 AUTHOR SIMPSON DATE  
 TAPE NUMBER 2615-110



SHEET 6

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3103	Ca 3011	3203	Ca 20075
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3105	Ca 3116	3205	Ca 2
3106	Ca 0	3206	Ca 0
3107	Ca 2200	3207	Ca 2210
3110	Ca 24000	3210	Ca 100
3111	Ca 20000	3211	Ca 10
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3113	Ca 20000	3213	Ca 0
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3115	Ca 0	3215	Ca 0
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3121	Ca 3116	3221	Ca 3221
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3333	Ca 157	3433	Ca 0
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MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE ITERATION I INDEX \_\_\_\_\_  
 AUTHOR SIMPSON DATE \_\_\_\_\_  
 TAPE NUMBER 2615-am.c

APPENDIX D

APPENDIX D

AUTO CROSS-CORRELATION I 2559 mo, ml) - Description and Use Of

This program was written for the WWI Digital Computer to compute the unnormalized sample correlations

$$\sum_{i=1}^{N+n-1} x_{i-j} y_i \quad j=0, 1, 2 \dots m$$

The conventions for preparation of the data  $x_1$  and  $y_1$  are identical with those described for Prediction XV in Appendix B, with the exception that the data tapes need not be combined after preparation. A short tape is prepared containing the information  $N$ ,  $n$ , and  $m$  as follows

Register	1047	+ N	(First data point in block)
(Octal)	1050	+ n	(No. data points in block)
	1051	+ m	(No. lags)

2559 mo handles individual data tapes and is used as follows

Erase storage, put Sil down  
Read in 2559 mo  
Read in Z (tape for  $N$ ,  $n$ ,  $m$ )  
Read in X ( $x_1$  data tape)  
Read in Y ( $y_1$  data tape)  
Start at 770

If  $X$  and  $Y$  are not identical we get half of the cross-correlation curve (for  $j = 0$ ). To get the other half, we repeat the instructions interchanging the order of read-in

for X and Y. If X and Y are the same tape, we get the entire auto-correlation curve, since since auto-correlations are symmetric about the zeroth lag.

The correlations are printed out by the direct printer as seven-place numbers, ten per line, the 0<sup>th</sup> lag being the first no. on the first line, the 1<sup>rst</sup> lag being the second no. on the first line, etc.

2559 ml performs the same functions as 2559 mo, but is adapted for handling the combined tapes used with Prediction XV. It assumes there are 3 real data sets plus a dummy set and forms the nine correlations representing the permutations of the 3 real sets. The correlations are over 380 values of the data, and are taken to 100 lags. The output is the delayed printer, and requires one minute for each 100 lag block. At this rate the program can perform 8 or 10 million multiplications in 4 hours of machine time. A sample of the output is shown on the next page.

RECORD 12.4 T1, T3, T5

CORRELATIONS = n' Q' T--T--(p)

N' = 120 m = 100

n' = 370 p = 0, -1, ... -m

T1T1

0179270	0175549	0167943	0159926	0154392	0152791	0154375	0157500	0160762	0163416
0165321	0166495	0166739	0165643	0163303	0160803	0159622	0160656	0163452	0166588
0168334	0167537	0164736	0161344	0158992	0158675	0159932	0161713	0162974	0163369
0163530	0164083	0164976	0165711	0165652	0164657	0163094	0161890	0161770	0162461
0163614	0163930	0162907	0161306	0160078	0160104	0161483	0163441	0164974	0165438
0165304	0165441	0166164	0166961	0166724	0164631	0161027	0157575	0156120	0157333
0160723	0164614	0167174	0167842	0166975	0165494	0164158	0163148	0162496	0161983
0161894	0162752	0164481	0166457	0167456	0166564	0163950	0160785	0158852	0158994
0161202	0164385	0166921	0167969	0167684	0166672	0165477	0164313	0163157	0161878
0160869	0160636	0161380	0162796	0164121	0164902	0164818	0164335	0164288	0165130
0166570									

T1T3

0212315	0205051	0197189	0192194	0191950	0195871	0201416	0206027	0208399	0208603
0207664	0206130	0204026	0201293	0198671	0197364	0198592	0202105	0206145	0208921
0209045	0206355	0202538	0199672	0198995	0200283	0202215	0203689	0204039	0203696
0203808	0204576	0205619	0206097	0205229	0203193	0200919	0199747	0200230	0201889
0203855	0204841	0204594	0203746	0203076	0203147	0203796	0204298	0204135	0203438
0203065	0203750	0205317	0206341	0205592	0202834	0199241	0196915	0197372	0200464
0204846	0208429	0209584	0208257	0205568	0203089	0201738	0201436	0201732	0202350
0203410	0205086	0206897	0207946	0207196	0204530	0201016	0198505	0198493	0201004
0205076	0208558	0209854	0208729	0206316	0204030	0202591	0202043	0202011	0201972
0202154	0202876	0204135	0205599	0206456	0206337	0205253	0204088	0203818	0204721
0206194									

T1T5

0209619	0202418	0194992	0190346	0189852	0192921	0197627	0202017	0204816	0205892
0205566	0203982	0201230	0197769	0194739	0193592	0195427	0199779	0204563	0207811
0208031	0205093	0200592	0196696	0195019	0195593	0197376	0199289	0200417	0200991
0201573	0202362	0202995	0202941	0202052	0200418	0198843	0198474	0199452	0200979
0202027	0201708	0200246	0198710	0198077	0198724	0200240	0201682	0202178	0201819
0201532	0202012	0202953	0203148	0201688	0198684	0195532	0194158	0195609	0199468
0203951	0206958	0207149	0204766	0201330	0198496	0197102	0197084	0197926	0199340
0201398	0204089	0206618	0207553	0205910	0201907	0197105	0193901	0193894	0197030
0201804	0205869	0207635	0206750	0204291	0201551	0199462	0198337	0198049	0198426
0199550	0201315	0203328	0204797	0204898	0203555	0201258	0199341	0198886	0200234
0202612									

T3T1

0212315	0214230	0211085	0204997	0198972	0195694	0195736	0198067	0201243	0204156
0206571	0208142	0208394	0206896	0203669	0199515	0196011	0195022	0197272	0201639
0206213	0208676	0208284	0205975	0203138	0201366	0201268	0202300	0203388	0203614
0203158	0202595	0202514	0202965	0203421	0203311	0202552	0201699	0201672	0202735
0204740	0206443	0206737	0205565	0203606	0202126	0201797	0202280	0202770	0202386
0201275	0200503	0201133	0203442	0206422	0208297	0207773	0205107	0202023	0200134
0200408	0202296	0204504	0205787	0205749	0204874	0203941	0203295	0202763	0201748
0200511	0200046	0201303	0204326	0207785	0209951	0209440	0206467	0202834	0200394
0200202	0202080	0204417	0205863	0205834	0204829	0203754	0203166	0203132	0203255
0203320	0203525	0204280	0205588	0206847	0207156	0205938	0203269	0200563	0199272
0200207									

T3T3

0269909	0265706	0256802	0247574	0241891	0241348	0244756	0249524	0253665	0256073
0257163	0257240	0256225	0253851	0250495	0247405	0246542	0248938	0253606	0258193
0260357	0258809	0254777	0250492	0247962	0248053	0249871	0251853	0252758	0252534
0252496	0253346	0255092	0256991	0257643	0256624	0254305	0251877	0250637	0250896
0252183	0253218	0253426	0253018	0252728	0253066	0253749	0254090	0253453	0251917
0250751	0251120	0253476	0256785	0259049	0258778	0255780	0251673	0248717	0248300
0250421	0253577	0255923	0256450	0255513	0254166	0253326	0253018	0252810	0252240
0251881	0252629	0254885	0257828	0259654	0259015	0255674	0251246	0248299	0248297
0251198	0255194	0257873	0258140	0256306	0254031	0252778	0252825	0253670	0254324
0254635	0254841	0255504	0256390	0256789	0255943	0253700	0250774	0248926	0249346
0252154									

T3T5

0261579	0258137	0251129	0244258	0240413	0240513	0243623	0247978	0252005	0254770
0256106	0257702	0253348	0249351	0244530	0240709	0239840	0242870	0248441	0254093
0257400	0256909	0253511	0249110	0245885	0245181	0246616	0248910	0250438	0250744
0250528	0250280	0250387	0250502	0250142	0249157	0247968	0247496	0248343	0250252
0252304	0253160	0252283	0250366	0248668	0248153	0248641	0249393	0249538	0248787
0247998	0248275	0250144	0252699	0254218	0253440	0250483	0247015	0252271	0246195
0249339	0252849	0254797	0254377	0252184	0249797	0248130	0247402	0247085	0246813
0247203	0249075	0252563	0256210	0257999	0256650	0252239	0247033	0243746	0244026
0247474	0252038	0255261	0255942	0254281	0251630	0249399	0248193	0247942	0248265
0248955	0250277	0252326	0254405	0255303	0254251	0251263	0247568	0245103	0245277
0248127									

T5T6

0209619	0212038	0209335	0203211	0196693	0192461	0191581	0193231	0196093	0199052
0201607	0203523	0204436	0204000	0201893	0198661	0195687	0194688	0196313	0199712
0203166	0204746	0203810	0201116	0198049	0196123	0196040	0197551	0198966	0199887
0200298	0200580	0201113	0201808	0202130	0201613	0200460	0199281	0198848	0199346
0200502	0201601	0201728	0200788	0199293	0198161	0197832	0198239	0198926	0199360
0199702	0200359	0201613	0203350	0204737	0204639	0202457	0198811	0195347	0193572
0194388	0197391	0201058	0203921	0205094	0204666	0203335	0201642	0199916	0198151
0196876	0196723	0198201	0201054	0204070	0205587	0204639	0201459	0197718	0195364
0195317	0197459	0200387	0202653	0203103	0202116	0203503	0201188	0200492	0199758
0199000	0198644	0199082	0200357	0201726	0202234	0201376	0199169	0196832	0195679
0196311									

T5T3

0261579	0258503	0251130	0243226	0238177	0237759	0241323	0246520	0251191	0253999
0255065	0254873	0253629	0251155	0247622	0243998	0241989	0242889	0246389	0250736
0253821	0254032	0251731	0248511	0246101	0245658	0246846	0248696	0249412	0249165
0248910	0249316	0250478	0251667	0251882	0250681	0248649	0246845	0246399	0247486
0249367	0250990	0251434	0250840	0249792	0249027	0248655	0248462	0248069	0247487
0247406	0248440	0250518	0252857	0253862	0252575	0249266	0245444	0243258	0243786
0246985	0251304	0254515	0255452	0254075	0251484	0249071	0247744	0246661	0246353
0246782	0248200	0250730	0253666	0255346	0254600	0251279	0246782	0243491	0243226
0246072	0250554	0253998	0254889	0253248	0250477	0248352	0247568	0247855	0248337
0248540	0248758	0249401	0250568	0251588	0251501	0249919	0247283	0245280	0245305
0247641									

T5T5

0260952	0257123	0248990	0240358	0234742	0233724	0236575	0241565	0246583	0250383
0252531	0252789	0251067	0247427	0242673	0238511	0236926	0239032	0243935	0249400
0253017	0253106	0249976	0254577	0241455	0239781	0240557	0243023	0245084	0246490
0247433	0248105	0248531	0248305	0247163	0245313	0243479	0242682	0243616	0245862
0248378	0249779	0249186	0247196	0244753	0243055	0242447	0242861	0243792	0244802
0246006	0247554	0249418	0250780	0250442	0247878	0243718	0243863	0238360	0240083
0244495	0249677	0253109	0253525	0251174	0247524	0244207	0242064	0241157	0241370
0242851	0245681	0249543	0253041	0254352	0252348	0247416	0241608	0237715	0237649
0241224	0246595	0250978	0252717	0251670	0248923	0246000	0243937	0242980	0242918
0243585	02450								

### Use of 2559 ml

The instructions for use are  
Erase storage, S11 switch down  
Read in 2559 ml  
Read in W (combined data)  
Start at 677

If the combined tape has 4 real data sets, and we want the 16 permutations of correlation, then an additional tape is used and the instructions are

Erase storage, S11 switch down  
Read in 2559 ml  
Read in 2559 p4  
Read in W  
Start at 677

2559 ml is equipped with the same "roll back" procedure that Iteration I is (Appendix C). In case of machine failure

Erase storage  
Read in 2559 pl3  
Start at 677

### Traveling Correlations

With the aid of tape 2559 pl0 we can use 2559 mo to obtain correlations from highly overlapping blocks of the data. The correlations are over blocks 50 in length and the number of lags is taken to be 20. The first reading in each block has an index (N) equal to  $k \times 10$  where  $k = 3, 4, \dots$   
44. The procedure for using 2559 mo in this way is

Erase storage, put Sil down

Read in 2559 mo

Read in X

Read in Y

Read in 2559 pl0 leave in P.E.T.R.

Start at 770 (21 lags are printed for N = 30)

Read in

Start at 770 (21 lags are printed for N = 40)

Read in

Start at 770 (21 lags are printed for N = 50)

etc.

Start at 770 (21 lags are printed for N = 440)

00	ca 1051		100	0.46000	
1001	ca 1023		101	0.66000	
1002	ca 1034		102	0.56000	
1003	si 0		103	0.06000	
1004	si 0		104	1.54000	
1005	si 0		105	0.24000	
1006	si 0		106	1.22000	
1007	si 0		107	+	
1010	si 0		110	si 0	
1011	si 0		111	-1	
1012	+ 570		112	-9	
1013	+		113	si 0	
1014	si 0		114	14	
1015	si 0		115	15	
1016	si 0		116	16	
1017	si 0		117	17	
1020	si 0		20		
1021	ca 2030		21		
1022	ca 1054		22		
1023	si 0		23		
1024	si 0		24		
1025	si 0		25		
1026	si 0		26		
1027	si 0		27		
30			1030	+ 0.100	
31			1031	ca 2021	
32			1032	ca 1054	
33			1033	ca 1032	
34			1034	ca 1036	
35			1035	ca 1041	
36			1036	ca	
37			1037	ca 1030	
40	ca 63		1040	ca 1030	
41	ca 70		1041	ca 1030	
42	ca 225		1042	ca 1021	
43	ca 103		1043	ca 1041	
44	ca 107		1044	ca 1031	
45	ca 110		1045	ca 1036	
46	ca 113		1046	ca 560	
47	ca 117		1047		
50	ca 122		1050		
51	ca 54		1051		
52	ca 117		52		
53	ca 113		53		
54	ca		54		
55	ca 55		55		
56	ca 110		56		
57	ca 46		57		
60	ca 105		560	ca 1032	
61	ca 61		561	ca 566	
62	ca 112		562	ca 603	
63	ca		563	ca 604	
64	ca 106		564	ca 566	
65	ca 65		565	ca 567	
66	ca 111		566	ca	
67	ca 112		567	ca 3014	
70	ca		570	ca 566	
71	+ 10		571	ca 604	
72	ca 00073		572	ca 565	
73	1.74000		573	ca 567	
74	1.52000		574	ca 605	
75	0.36000		575	ca 605	
76	0.16000		576	ca 606	
77	0.26000		577	ca 567	

NOTES

SHEET 1

600	si 0		700	ca 1032	
601	si 0		701	ca 111	
602	si 0		702	ca 9	
603	-499		703	ca 722	
604	si 0		704	ca 723	
605	si 0		705	ca 1050	
606	ca 0.06027		706	ca 724	
607	si 0		707	ca 724	
610	si 0		710	ca 723	
611	si 0		711	ca	
612	si 0		712	ca 723	
613	si 0		713	ca 722	
614	si 0		714	ca 722	
615	si 0		715	ca 711	
616	si 0		716	ca 724	
617	ca 630		717	ca 718	
620	ca 643		720	ca 617	
621	ca 645		721	si 0	
622	ca 667		722	si 0	
623	ca 654		723	si 0	
624	ca 634		724	si 0	
625	ca 632		725	si 0	
626	ca 650		726	ca 1015	
627	ca 651		727	ca 1013	
630	ca 637		730	ca	
631	ca 722		731	ca 1015	
632	ca 660		732	ca 710	
633	ca 660		733	si 0	
634	ca		734	si 0	
635	ca 641		735	ca 1050	
636	ca 660		736	ca 1024	
637	ca		737	ca 1024	
640	ca 633		740	ca 753	
641	ca 637		741	ca 1016	
642	ca 637		742	ca 752	
643	ca 637		743	ca 1015	
644	ca 633		744	ca 751	
645	ca 633		745	ca 751	
646	ca 633		746	ca 751	
647	ca 633		747	si 0	
650	ca 10013		750	si 0	
651	ca 663		751	ca	
652	ca 2		752	ca	
653	si 0		753	ca	
654	ca 655		754	ca 751	
655	+ 100		755	ca 1013	
656	+ 10		756	ca 751	
657	+		757	ca 752	
660	si 0		760	ca 1013	
661	si 0		761	ca 752	
662	ca 225		762	ca 753	
663	ca		763	ca 1024	
664	ca 664		764	ca 751	
665	ca		765	ca 726	
666	ca 666		766	si 0	
667	ca		767	si 0	
670	ca 670		770	ca 1021	
671	ca 723		771	ca 1047	
672	ca 40		772	ca 1014	
673	ca 1023		773	ca 1054	
674	ca 734		774	ca 1013	
675	si 0	STOP	775	ca 1015	
676	si 0		776	ca 1012	
677	si 0		777	ca 1016	

STAIR

MIT DIGITAL COMPUTER LABORATORY  
 OCTAL PROGRAM FORM  
 TITLE AUTO CROSS-CORRELATION INDEX  
 AUTHOR SIMPSON DATE  
 TAPE NUMBER 2559-mc



300	ca 313	500	ca 507
301	ca 707	501	ca 320
302	ca 312	502	ca 176
303	ca 1654	503	ca 510
304	ca 313	504	ca 504
305	ca 312	505	ca 504
306	ca 313	506	ca 0
307	ca 314	507	-2
310	ca 177	510	142000
311	ca 320	11	
312	ca 500	12	
313	+2049	13	
314	-3	14	
15		15	
16		16	
17		17	
320	ca 334	20	
321	ca 703	21	
322	ca 312	22	
323	ca 2030	23	
324	ca 325	24	
325	ca 340	25	
326	ca 334	26	
327	ca 312	27	
330	ca 334	30	
331	ca 336	31	
332	ca 335	32	
333	ca 340	33	
334	+2049	34	
335	-2	35	
336	-2	36	
37		37	
340	ca 356	540	ca 111
341	ca 703	541	ca 112
342	ca 312	542	ca 176
343	ca 3014	543	ca 104
344	ca 356	544	ca 544
345	ca 312	545	ca 554
346	ca 356	546	ca 543
347	ca 360	547	ca 553
350	ca 770	550	ca 554
351	ca 336	551	ca 134
352	ca 360	552	ca 580
353	ca 357	553	-4
354	ca 366	554	-4
355	ca 770	555	
356	+2049	556	
357	+2049	557	
360	-2	60	
61		61	
62		62	ca 176
63		63	
64		64	
65		65	
66		66	
67		67	
370	ca 376	70	
371	ca 707	71	
372	ca 375	72	
373	ca 40	673	ca 134
374	ca 320	674	ca 1023
375	ca 01054	675	ca 134
376	+4097	676	ca 540
77		677	ca 32
NOTES:			

01-481

SHEET 2

00		00	
01		01	
02		02	
03		03	
04		04	
05		05	
06		06	
07		07	
10		10	
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32		32	
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36		36	
37		37	
40		40	
41		41	
42	ca 43	42	
43		43	
44		44	
45		45	
46	ca 300	46	
47	+120	47	
48		48	
49	+370	49	
50	+100	50	
51		51	
52		52	
53		53	
54		54	
55		55	
56		56	
57		57	
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74		74	
75		75	
76		76	
77		77	
MIT DIGITAL COMPUTER LABORATORY OCTAL PROGRAM FORM TITLE AND CROSS-CORRELATION INDEX AUTHOR SIMPSON DATE TAPE NUMBER 657 ca 1			

APPENDIX E

APPENDIX E

BIOGRAPHICAL NOTE    Stephen Milton Simpson, Jr.

Attended Yale University September 1946 - June 1950, receiving B.S. in Physics. Entered the Massachusetts Institute of Technology in the Department of Geology in September 1950. Member of Phi Beta Kappa and Sigma Xi.

Presently under appointment as Instructor in the Department of Geology and Geophysics at the Massachusetts Institute of Technology.