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## Thermodynamics shapes the biogeography of propionate-oxidizing syntrophs in paddy field soils

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1 **Thermodynamics Shapes the Biogeography of Propionate-Oxidizing Syntrophs in**  
2 **Paddy Field Soils**

3 Running title: Biogeography of Propionate-Oxidizing Syntrophs

4

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## 25 **Summary**

26 Soil biogeochemical processes are not only gauged by the dominant taxa in the  
27 microbiome but also depend on the critical functions of its “rare biosphere” members.  
28 Here we evaluated the biogeographical pattern of “rare biosphere” propionate-oxidizing  
29 syntrophs in 113 paddy soil samples collected across China. The relative abundance,  
30 activity and growth potential of propionate-oxidizing syntrophs were analyzed through  
31 sequencing bacterial 16S rRNA genes and anaerobic incubations to provide a panoramic  
32 view of syntroph biogeographical distribution at the continental scale. The relative  
33 abundances of four syntroph genera, *Syntrophobacter*, *Pelotomaculum*, *Smithella* and  
34 *Syntrophomonas* were significantly greater at the warm low latitudes than at the cool high  
35 latitudes. Correspondingly, propionate degradation was faster in the low latitude soils  
36 compared with the high latitude soils. The **low** rate of propionate degradation in the high  
37 latitude soils resulted in a greater increase of the total syntroph relative abundance,  
38 probably due to their initial low relative abundances and the longer incubation time for  
39 propionate consumption. The mean annual temperature (MAT) is the most important  
40 factor shaping the biogeographical pattern of propionate-oxidizing syntrophs, with the  
41 next factor being the soil’s total sulfur content (TS). We suggest that the effect of MAT is  
42 related to the thermodynamic conditions, in which the endergonic constraint of  
43 propionate oxidation is leveraged with the increase of MAT. The TS effect is likely due  
44 to the ability of some propionate syntrophs to facultatively perform sulfate respiration.

45

46 **Keywords:** Rare biosphere, Propionate syntrophs, Biogeography, Methanogenesis,  
47 Paddy soil.

48

## 49 **Introduction**

50       The soil microbiome plays a vital role in regulating global biogeochemistry. It has  
51 been recently documented that in spite of the immense biodiversity of the soil  
52 microbiome, only a few hundreds of phylotypes are prevalent across global soils and that  
53 these dominant phylotypes display distinct biogeographical distribution according to their  
54 habitat preferences (Delgado-Baquerizo et al., 2018). The factors shaping the  
55 biogeographic patterns of the global soil microbiome include climatic factors, edaphic  
56 properties and biological interactions (Garbeva et al., 2004; Schlatter et al., 2015; Bahram  
57 et al., 2018; Rillig et al., 2019; Steidinger et al., 2019). Identification of soil dominant  
58 taxa, their geographical distributions and the controlling factors provides a “priority” list  
59 for cultivation and genomic scrutiny that shall help in decoding biogeochemical  
60 mechanisms and pave a way toward developing global Earth system models that are soil  
61 context-dependent (Delgado-Baquerizo et al., 2018; Crowther et al., 2019; van den  
62 Hoogen et al., 2019). The dominants-focused approaches, however, will inevitably  
63 overlook the function of “rare biosphere” specialists (Lynch and Neufeld, 2015). It has  
64 been understood that biogeochemical processes are not only gauged by dominant taxa  
65 that dictate biomass turnover but also depend on critical functions of specialist members  
66 (Crowther et al., 2019). A typical example is microbial syntrophs, which play a critical  
67 role in anaerobic decomposition of organic matter and methanogenesis but have only a  
68 low relative abundance in natural environments like paddy field soils (Lueders et al.,  
69 2004; Gan et al., 2012).

70       Complex organic matter in anoxic environments is degraded by an anaerobic food  
71 chain comprising primary and secondary fermenters, homoacetogens and methanogens  
72 (Glissmann et al., 2001; Stams and Plugge, 2009). Polymers are initially converted to

73 oligomers and monomers and subsequently fermented to short-chain fatty acids (SCFAs),  
74 alcohols, acetate, and so on. A series of SCFAs and alcohols are transiently accumulated  
75 as intermediate products during the process (Glissmann and Conrad, 2000; Glissmann et  
76 al., 2001; Rui et al., 2009; Noll et al., 2010). Therefore during the anaerobic  
77 decomposition of complex organic matter, a quite portion of the carbon flows through the  
78 SCFAs, which are subsequently degraded to acetate and CO<sub>2</sub> by secondary fermenters (i.e.  
79 the syntrophs) (Schink, 1997; Schink and Stams, 2006; McInerney et al., 2009; Stams and  
80 Plugge, 2009). Propionic acid, one of the major SCFAs, accounts for up to 30% of the  
81 total CH<sub>4</sub> formation in paddy soils (Krylova et al., 1997; Glissmann and Conrad, 2000).  
82 The process of propionate degradation is thermodynamically more constrained than that  
83 of ethanol and butyrate, making propionate degraders being slower in growth and more  
84 sensitive to environmental conditions (Schink, 1997; Boe et al., 2010; Mueller et al., 2010;  
85 Hidalgo-Ahumada et al., 2018). Identifying propionate-oxidizing key taxa and delineating  
86 their biogeographical distribution are therefore important for better understanding and  
87 predicting organic matter decomposition and carbon cycling in anoxic environments  
88 including natural wetlands and paddy field soils.

89 All known propionate-oxidizing syntrophs had in pure cultures so far fall into four  
90 bacterial genera: *Syntrophobacter*, *Pelotomaculum*, *Smithella* and *Desulfotomaculum*  
91 (Sieber et al., 2012; Hidalgo-Ahumada et al., 2018; Dykema and Gallert, 2019). Their  
92 importance in anoxic environments, such as paddy field soils and natural wetlands, has  
93 been demonstrated by using various approaches such as enrichment cultivation,  
94 DNA/RNA stable isotope probing and meta-omics technology (Lueders et al., 2004; Li et  
95 al., 2015; Tveit et al., 2015; Peng et al., 2018; Xia et al., 2019). Recently, a few candidate  
96 taxa (e.g. *Ca. Syntrophosphaera thermopropionivorans* from the phylum *Cloacimonetes*,

97 *Ca. Syntrophopropionicum ammoniitolerans* from the family *Peptococcaceae* and  
98 member from the family *Desulfobulbaceae*) were proposed to be capable of performing  
99 syntrophic propionate metabolism (Dyksma and Gallert, 2019; Singh et al., 2021). The  
100 ecophysiology and environmental significance of these organisms however have  
101 remained unclear.

102 Paddy field soils are human-managed ecosystems and are important contributors to  
103 global CH<sub>4</sub> emissions (Thauer, 1998; Conrad, 2009). About 90% of paddy fields are  
104 located in Asia with 20% of that in China (Haeefele et al., 2014). Rice production in China  
105 is dominated by irrigated systems, which are distributed mainly in the lowland of eastern  
106 China, extending from the warm sub-tropics at 18°N latitude to the cool temperate  
107 regions at 50°N (Deng et al., 2019). The long history of rice cultivation makes the region  
108 an ideal site for microbial biogeography investigations that are independent of large scale  
109 variations in vegetation and soil type. In the present study, we collected 113 paddy field  
110 soils from different sites across eastern China covering a latitudinal distance of 3689  
111 kilometers. The potential for propionate degradation in each soil was evaluated through  
112 anaerobic incubations in the laboratory. The composition and relative abundance of four  
113 known genera of propionate oxidizing syntrophs in original soils and soil samples after  
114 anaerobic incubations were investigated using high throughput sequencing of bacterial  
115 16S rRNA genes. We show here that temperature and the total sulfur content in soil are  
116 the most important factors shaping the biogeographic distribution and functional potential  
117 of propionate-oxidizing syntrophs in paddy field soils at the continental scale.

118

## 119 **Results and Discussion**

## 120 **Biogeographical distribution of propionate syntrophs**

121 A total of 6204541 high quality sequences of V3-V4 regions of the bacterial 16S  
122 rRNA genes were obtained from 113 paddy soil samples, which were clustered into  
123 27411 OTUs. The most dominant OTUs were affiliated to *Proteobacteria*, *Firmicutes*,  
124 *Actinobacteria*, *Chloroflexi*, *Acidobacteria*, *Nitrospirae* and *Gemmatimonadetes* (Fig. S1).  
125 For the present study we focused on four genera of propionate-oxidizing syntrophs (i.e.  
126 *Syntrophobacter*, *Pelotomaculum*, *Smithella* and *Desulfotomaculum*) and  
127 *Syntrophomonas* (Fig. S2). A total of 30 OTUs belonging to these five syntroph genera  
128 were obtained. Members of the genus *Syntrophomonas* are not known to oxidize  
129 propionate but butyrate and fatty acids up to C<sub>10</sub> (Zhang et al., 2004). The inclusion of  
130 *Syntrophomonas* in our analysis was due to the fact that *Smithella*, which utilize the C<sub>6</sub>  
131 dismutation pathway [ $2\text{CH}_3\text{CH}_2\text{COO}^- \rightarrow \text{CH}_3\text{COO}^- + \text{CH}_3\text{CH}_2\text{CH}_2\text{COO}^-$ ], can release  
132 butyrate as an intermediate product, which is then metabolized by *Syntrophomonas* (Xia  
133 et al., 2019). *Desulfotomaculum* was included, even though their relative abundances  
134 were very low (Fig. S2, 3), because some *Desulfotomaculum* species such as *D.*  
135 *thermobenzoicum* and *D. thermocisternum* have been known to utilize propionate in  
136 coculture with methanogens (Nilsen et al., 1996; Plugge et al., 2002) and because  
137 *Pelotomaculum* species are phylogenetically a branch of *Desulfotomaculum* (Imachi et al.,  
138 2006; Imachi et al., 2007).

139 The total relative abundance of the five syntroph genera in combination, referred to  
140 as synTotal, ranged from 0.01 to 0.34% across 113 soil samples. This low relative  
141 abundance certificates that syntrophs in paddy soils belong to a subset of the “rare  
142 biosphere” community (defined as individual species or OTUs accounting for <0.1% of  
143 the total relative abundance). Albeit “rare” in relative abundance, the synTotal displayed



144 a distinct geographical distribution, showing the highest relative abundance in the warm  
145 low latitude soils and gradually decreasing towards the high latitude regions (Fig. 1A).  
146 Conventionally the geographical division separating China into the south (low latitude)  
147 and the north (high latitude) regions follows to the Qinling Mountains-Huaihe River Line  
148 (latitude  $\approx 32^\circ$ ), a critical geographical boundary for climate, landform and soil  
149 conditions (Qi et al., 2016). We summarized our soil samples accordingly into a low  
150 latitude and the a high latitude group, respectively (bottom-left of Fig. 1A). The mean  
151 relative abundance of the low latitude group (South) was significantly greater (by 1.84-  
152 fold) than that of the high latitude group (North). In line with this, the  $\alpha$ -diversity of  
153 syntrophs, estimated based on OTU richness, significantly decreased with increasing  
154 latitude (Fig. 1B). Having found this, we then identified the key environmental factors  
155 related to the synTotal distribution through Spearman correlation analysis. Mean annual  
156 temperature (MAT) and total soil sulfur (TS) were identified as the two most important  
157 factors, followed by ammonium-nitrogen ( $\text{NH}_4^+\text{-N}$ ), soil organic matter (OM), microbial  
158 biomass carbon (MBC) and total nitrogen (TN) (Fig. 1C). Soil cation exchange capacity  
159 (CEC) and pH showed negative correlations. The importance index estimated based on  
160 the Boruta algorithm reiterated that MAT and TS were the two most important factors  
161 linked to the biogeographical distribution of synTotal (Fig. 1C).

162 To compare the distribution of different syntrophs, the relative abundances of five  
163 syntroph genera were individually analyzed (Fig. S3). The distributions of  
164 *Syntrophobacter*, *Pelotomaculum*, *Smithella* and *Syntrophomonas* were consistent with  
165 the distribution of the synTotal (Fig. S3A-D). *Desulfotomaculum* was an exception,  
166 showing the highest abundance in the middle latitude regions (Fig. S3E). Division of the  
167 low and high latitude groups revealed that the mean relative abundances of

168 *Syntrophobacter*, *Syntrophomonas* and *Pelotomaculum* were significantly greater in the  
169 low latitude regions than in the high latitude regions (Fig. S3F), in consistence with the  
170 synTotal. *Syntrophobacter* had the highest mean relative abundance followed by  
171 *Syntrophomonas* and *Smithella*, while *Pelotomaculum* and *Desulfotomaculum* showed  
172 very low relative abundances (Fig. S3F).

173

### 174 **Biogeographical pattern of syntroph functioning potentials**

175 Next we evaluated the biogeographical pattern of syntroph functioning potentials.  
176 For this purpose, fresh soil samples were incubated under anaerobic conditions at 30°C  
177 with addition of 10 mM propionate. We tracked CH<sub>4</sub> production until >90% of the added  
178 propionate had been consumed. Three distinct methane production patterns were  
179 identified according to the time “lapse” required for CH<sub>4</sub> production from propionate  
180 oxidation (Fig. 2). Pattern I comprised 45 soil samples with a time lapse of 13 d to 27 d,  
181 representing the **high** rate group of syntrophic metabolisms (Fig. 2A). The majority of  
182 soil samples in this group was located at the low latitudes. Pattern II, comprising 51  
183 samples and representing the median rate of syntrophic metabolisms (28 d to 43 d) did  
184 not show a distinct latitudinal tendency (Fig. 2B). The pattern III group comprising 17  
185 soil samples required 44 d to 82 d for CH<sub>4</sub> production from propionate oxidation and  
186 hence represented the **low** rate group. The soils of this group were distributed mainly in  
187 the cool high latitude regions (Fig. 2C). The SCFAs analyses further underpinned and  
188 justified the separation of soil samples into three groups (Fig. 3A). Acetate was the major  
189 intermediate detected in all samples (Fig. 3B). Butyrate was detected in 28.3% of soil  
190 samples (detection limit of 0.05 mM) (Fig. 3C), most likely as an intermediate product

191 from the C<sub>6</sub> dismutation pathway of *Smithella*. The highest butyrate concentrations  
192 occurred in the pattern III soils (Fig. 3C). Taking all soil samples together, the time lapse  
193 for methanogenesis from propionate degradation displayed an explicit biogeographical  
194 pattern, being significantly slower in the high latitude soils than in the low latitude soils  
195 (Fig. 4A). Linear least squares regressions revealed that the time lapse for  
196 methanogenesis was significantly negatively correlated with the relative abundance of  
197 synTotal in original soils, MAT, TS, and other edaphic factors including soil OM, MBC,  
198 available Fe, Cu and Mn (Fig. 4C). The functional potential of propionate degradation  
199 can also be inferred from the maximum rate of methanogenesis (Fig. 4B), which supports  
200 the biogeographical tendency revealed by the time lapse.

201

### 202 **Shifts in microbial community during anaerobic incubation with propionate**

203 At the end of the anaerobic incubations, the structure of the microbial communities in  
204 the soil samples was revisited. The community composition at the phylum level did not  
205 show significant changes between the original soil samples and those after anaerobic  
206 incubations. The relative abundances of a few phyla, however, had changed markedly  
207 (Fig. S4A). Specifically, the relative abundances of *Proteobacteria* and *Acidobacteria*  
208 decreased during the incubations while those of *Actinobacteria*, *Firmicutes* and  
209 *Chloroflexi* were increased. As expected, the  $\alpha$ -diversity of the bacterial communities at  
210 the OTU level showed a substantial decline at the end of the incubation compared with  
211 the original soils (Fig. S4B). The  $\beta$ -diversity at the genera level showed the separation of  
212 soil samples into two clusters for the low latitude and the high latitude soils, respectively  
213 (Fig. S4C). The co-occurrence network analysis of the top 10% OTUs confirmed that the

214 OTUs for the low and high latitude regions tended to group separately (Fig. S4D). These  
215 results indicate that the bacterial community shifted significantly during anaerobic  
216 incubation, while the separation of metacommunities into the low latitude and the high  
217 latitude groups remained robust.

218 The relative abundance of syntrophs except *Desulfotomaculum* had increased  
219 markedly after the incubation with the maximum relative abundance of synTotal reaching  
220 12.5%, indicating significant growth. We estimated the increase of individual syntrophs  
221 by calculating the logarithmic ( $\log_2$ ) fold change (R) of the relative abundance before and  
222 after the incubation. The relative abundance of *Desulfotomaculum*, being low right at the  
223 beginning, did not change over the incubation (data not shown). The other four genera  
224 showed significant increases, but to different extents (Fig. 5). *Pelotomaculum* showed the  
225 greatest increase, followed by *Smithella* and *Syntrophomonas*, while *Syntrophobacter*  
226 exhibited the lowest increase (Fig. 5F). Strikingly, the relative increases of syntrophs  
227 showed an opposite geographic tendency compared with their relative abundances in the  
228 original soils (Fig. 2). The  $\log_2$  R values for *Syntrophobacter*, *Syntrophomonas* and  
229 *Smithella* increased with the increase of latitude (Fig. 5A, B, D). The values for  
230 *Pelotomaculum* also increased with latitude, though showing the highest value at the  
231 middle latitude (Fig. 5C). These results indicate that the relative abundance of propionate  
232 syntrophs after anaerobic incubation increased to a greater extent in the high latitude  
233 soils compared to the low latitude soils.

234 Given their low relative abundances and their critical roles in oxidizing intermediate  
235 products during anaerobic decomposition of organic matter, propionate-oxidizing  
236 syntrophs represent a typical example of rare but important taxa in paddy soil  
237 microbiome. Here we show that the relative abundance of propionate syntrophs is

238 significantly correlated with MAT (Fig. 1). Based on the time lapse for propionate-fueled  
239 methanogenesis, we further reveal that the biogeographical pattern of syntroph  
240 functioning potentials (Figs. 2, 3) is in accordance with their relative abundances in soils.  
241 It has been documented that within the physiological range, the microbial diversity,  
242 metabolic activity and population growth rates in terrestrial ecosystems increase  
243 exponentially with temperature (Zhou et al., 2016). Hence, climate warming is expected  
244 to accelerate temporal turnover and divergent succession of microbial communities in  
245 soils (Guo et al., 2018; Guo et al., 2019). In response to the increased temperature,  
246 microbiota can also modulate metabolic and trophic interactions through shifting  
247 functional guilds (Tveit et al., 2015). Syntrophs are known to have a specific lifestyle by  
248 living at thermodynamic limit. Syntrophic propionate degradation has been revealed to be  
249 temperature sensitive (Tveit et al., 2015). In order to obtain a deeper insight into the  
250 temperature effect, we calculated the Gibbs free-energy changes ( $\Delta G'$ ) for the reaction of  
251 propionate oxidation [ $\text{CH}_3\text{CH}_2\text{COO}^-_{(\text{aq})} + 2\text{H}_2\text{O}_{(\text{l})} \rightarrow \text{CH}_3\text{COO}^-_{(\text{aq})} + 3\text{H}_{2(\text{g})} + \text{CO}_{2(\text{g})}$ ] (Fig.  
252 6A).  $\Delta G^0_f$  and  $\Delta H^0_f$  values for acetate and propionate were obtained from the database  
253 established by Shock and Helgeson (Shock and Helgeson, 1990). We put temperature as  
254 the only variable with other variables set to standard conditions and calculated the  $\Delta G'$  as  
255 described by Hanselmann (Hanselmann, 1991). The calculated  $\Delta G'$  decreased linearly  
256 with temperature at a rate of  $-0.44 \text{ kJ mol}^{-1} \text{ }^\circ\text{C}^{-1}$  (Fig. S5). Consequently, the cool high  
257 latitude regions showed significantly higher positive values of  $\Delta G'$  than the warm low  
258 latitude regions (Fig. 6A). The geographical distribution of the Gibbs free-energy changes  
259 is in corroboration with the linear positive correlations of synTotal and the maximum rate  
260 of  $\text{CH}_4$  production with temperature (Fig. 6B,D) and with the negative correlation for the  
261 time lapse of methanogenesis from propionate degradation (Fig. 6C). These correlations

262 suggest that the decrease of Gibbs free-energy changes in the southern regions of China  
263 alleviates the thermodynamic tension, hence supporting the greater abundance and  
264 functional potential of propionate-oxidizing syntrophs compared with those in the  
265 northern regions. It has to be noted that the values of  $\Delta G'$  depicted in Fig. 6A are positive  
266 due to the use of standard conditions except temperature. In reality, negative  $\Delta G'$  values  
267 are to be expected due to low *in situ* concentrations of H<sub>2</sub> and acetate (Schutz et al., 1988;  
268 Kramer and Conrad, 1993). Fluctuations in concentrations of substrates and products  
269 actually can exert a significant impact on  $\Delta G'$ , especially the fluctuations in H<sub>2</sub> as three  
270 moles of H<sub>2</sub> are produced per mole of propionate oxidized in the reaction mentioned  
271 above. Indeed, due to significant seasonal and spatial variations, the explanatory power of  
272 thermodynamics hinging on *in situ* concentrations of substrates and products has to be  
273 applied with care to infer geographical patterns at a large distance scale. Only very few  
274 studies have determined methanogenic substrates and products under quasi-steady  
275 conditions in paddy soil, which showed that H<sub>2</sub> concentrations were mostly lower than a  
276 few Pa (Schutz et al., 1988; Kramer and Conrad, 1993). Accumulation of such  
277 observations from different locations shall help shape a better understanding about  
278 geographical distributions of syntrophic organisms.

279 We found that the total sulfur content of the soil was the second most important  
280 factor influencing propionate syntroph biogeography. This was possibly due to the fact  
281 that many of propionate syntrophs, such as *Syntrophobacter fumaroxidans* and  
282 *Pelotomaculum thermopropionicum*, are facultative sulfate reducers (Harmsen et al.,  
283 1998; Imachi et al., 2002; Sedano-Nunez et al., 2018). In the presence of sulfate these  
284 organisms tend to perform anaerobic sulfate respiration instead of syntrophy with  
285 methanogens, maximizing energy conservation (Van Kuijk and Stams, 1995; Rebac et al.,

286 1996; Worm et al., 2014; Li et al., 2018; Sedano-Nunez et al., 2018). Notably, while the  
287 relative abundances of *Syntrophobacter*, *Smithella* and *Syntrophomonas* in paddy soils  
288 were significantly correlated with TS, that of *Pelotomaculum* was not (Fig. 1C).  
289 Therefore, individual syntrophs have different responses to TS. Besides TS, the relative  
290 abundances of syntrophs were also positively correlated with the contents of soil OM,  
291 total N, MBC; additionally the functioning potential was correlated with the contents of  
292 trace elements Fe, Cu and Mn (Fig. 4C). Since a high soil OM content is usually  
293 associated with high TN, MBC and trace element contents and even TS, these edaphic  
294 factors are probably co-variables, yet the positive correlations likely indicate the  
295 favorable living conditions for most microbes including syntrophs.

296 Anaerobic incubation with propionate caused a marked shift in the soil bacterial  
297 community (Fig. S4). Specifically, the relative abundances of syntrophic populations  
298 increased by up to 29-fold over the incubation period. The increase in relative abundances,  
299 however, differed among the five syntroph genera. *Pelotomaculum* showed the greatest  
300 growth, followed by *Smithella* and *Syntrophomonas* while *Syntrophobacter* and  
301 *Desulfotomaculum* grew the least (Fig. 5). These results indicate that *Pelotomaculum*  
302 species have the most potential to be active in paddy soils across a wide range of  
303 geographical locations when favorable conditions become available. *Smithella* in  
304 combination with *Syntrophomonas* are the next group with high metabolic potentials,  
305 while *Desulfotomaculum* showed virtually no response. Compared to the other three  
306 syntroph genera (*Pelotomaculum*, *Smithella* and *Syntrophomonas*), *Syntrophobacter* had  
307 the highest mean relative abundance in original soils (Fig. S3) but showed the least  
308 increase after the incubation (Fig. 5). Probably, under a warmer 30°C condition during the  
309 incubation (compared with the mean summer temperature of 17-29°C across sampling

310 sites), *Syntrophobacter* were less competitive in utilizing propionate than the other  
311 syntrophs (Gan et al., 2012; Chen et al., 2020). Alternately, it is also likely that  
312 *Pelotomaculum* and *Smithella* reacted faster than *Syntrophobacter* when an adequate  
313 concentration of propionate (10 mM in this study) was supplied. *Pelotomaculum* and  
314 *Smithella* are known to use the methylmalonyl-CoA (MMC) pathway and the C<sub>6</sub>  
315 dismutation pathway, respectively (Liu et al., 1999; de Bok et al., 2001; Mueller et al.,  
316 2010; Xia et al., 2019). Due to the necessity of only two electrons released from the C<sub>6</sub>  
317 pathway compared with six electrons from the MMC pathway per propionate oxidized,  
318 *Smithella* are considered to have a larger window of opportunity in environments than the  
319 MMC-utilizing syntrophs (Dolfing, 2013). The calculation of  $\Delta G'$  values indicates that  
320 the energetics of the C<sub>6</sub> pathway is less sensitive to temperature than the MMC pathway  
321 (Fig. S6). Consistent with this, we found that the maximum increase of relative  
322 abundance occurred at the middle latitudes for *Pelotomaculum* but farther north for  
323 *Smithella* and *Syntrophomonas* (Fig. 5A vs B and D). Together with the detection of  
324 butyrate as a significant intermediate (Fig. 3C), our results suggest that *Smithella* and  
325 *Syntrophomonas* likely form syntrophic interaction during propionate degradation and  
326 they are potentially more active in the cool high latitude soils.

327 The increase in relative abundances, expressed as  $\log_2 R$ , tended to be greater in the  
328 high latitude soils relative to the low latitude soils (Fig. 5), in contrast to the geographical  
329 distribution of relative abundances in original soils (Fig. 1). There are three plausible  
330 explanations for this result. Firstly, in a mixed community, not only syntrophs are present  
331 (Fig. S4), and a relatively slow response of those other organisms would result in a  
332 relative increase of syntrophs. Secondly, for an initially low biomass (which appears to be  
333 the case in the present experiment), consumption of an identical quantity of substrate is



334 expected to produce a greater fold change of biomass compared with a group of  
335 organisms having an initial high biomass. Thirdly, the anaerobic incubations were  
336 stopped when >90% of propionate were consumed, which resulting in incubations of  
337 between 13 d and 82 d for different soils (Fig. 2). A longer duration may have resulted in  
338 a greater biomass yield per unit substrate consumption by preventing energetic loss from  
339 overflow metabolism (e.g. futile cycles and protein synthesis) (Thauer et al., 2008;  
340 Beardmore et al., 2011; Lipson, 2015; Seel et al., 2016). Whether this rate-yield tradeoff  
341 relationship really exists in paddy soil syntrophs requires further investigation.

342 In summary, we demonstrate in the present study that both the relative abundance  
343 and functioning potential of propionate syntrophs in paddy field soils exhibit a distinct  
344 biogeographical pattern. MAT and TS are identified as the two most important factors  
345 related to the biogeographical distribution. The temperature effect is very likely related to  
346 thermodynamic conditions, which improve when the temperature increases. The effect of  
347 soil sulfur content is possibly due to the fact that some facultative propionate syntrophs  
348 perform sulfate respiration in the presence of sulfate. Though originating from areas with  
349 a MAT lower by only 6.7°C, the mean time lapse for methanogenesis from propionate  
350 degradation was approximately 30 d longer in the high latitude soils relative to the low  
351 latitude soils. This result, on the one hand, indicates that the rates of organic matter  
352 decomposition and C cycling are potentially lower in the cool high latitude regions. On  
353 the other hand, it suggests that the global warming, which is occurring to a greater extent  
354 at the high latitudes, can cause a significant acceleration of organic matter decomposition  
355 and C cycling in the high latitude soils.

356

357 **Data availability**

358 The datasets supporting the conclusions of this article are available in the NCBI  
359 Sequence Read Archive under BioProject PRJNA544819 and PRJNA601098 that are  
360 publicly accessible at <https://www.ncbi.nlm.nih.gov>. R codes on the statistical analyses  
361 are available at [https://github.com/jinyidan/Rare-Biosphere-Propionate-Syntrophs-in-](https://github.com/jinyidan/Rare-Biosphere-Propionate-Syntrophs-in-Paddy-Field-Soils)  
362 [Paddy-Field-Soils](https://github.com/jinyidan/Rare-Biosphere-Propionate-Syntrophs-in-Paddy-Field-Soils).

363

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370 analysis with constructive discussions. Y.L. and Y.J. wrote the original manuscript. All  
371 authors read, revised and approved the final manuscript.

372 The authors declare that they have no competing interests.

373

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542

543 **Figure Legends**

544

545 **Figure 1:** Biogeographic distribution of propionate syntrophs in paddy field soils across  
546 eastern China. (A) The map shows the spatial distribution of total relative abundance (Rel.  
547 abun.) of five syntroph genera (referred to synTotal) comprising *Syntrophobacter*,  
548 *Pelotomaculum*, *Smithella*, *Desulfotomaculum* and *Syntrophomonas* together. The map  
549 was built using Kriging interpolation method with cross-validation (cv) based on Pearson

550 correlation. The inset in lower-left corner shows the difference in mean relative  
 551 abundance of synTotal between the low (South) and the high (North) latitude soils. The  
 552 red asterisks indicate the significant difference at  $P \leq 0.001$ . (B) The  $\alpha$ -diversity of  
 553 propionate syntrophs expressed as OTU richness, which decreases with the increase of  
 554 latitude. The solid line denotes the least-squares linear regression, which was significant  
 555 at  $P < 0.0001$ . (C) Correlations between the relative abundance of propionate syntrophs  
 556 (either in total or individually) and environmental factors. The color and size of circles  
 557 indicate the Spearman correlation coefficients. The black asterisks indicate the significant  
 558 correlations: \*  $0.01 < P \leq 0.05$ , \*\*  $0.001 < P \leq 0.01$ , \*\*\*  $P \leq 0.001$ . Bars in the right show  
 559 the importance of environmental factors to five syntroph genera estimated based on  
 560 Boruta algorithm. The environmental factors include the mean annual temperature (MAT)  
 561 and fifteen edaphic factors: the content of soil organic matter (OM), dissolved organic  
 562 carbon (DOC), pH, cation exchange capacity (CEC), microbial biomass carbon (MBC),  
 563 the total contents of N, P, K, S, Fe, Mn, Cu, Zn (TN, TP, TK, TS, TFe, TMn, TCu, TZn),  
 564 ammonium ( $\text{NH}_4^+$ ) and nitrate ( $\text{NO}_3^-$ ).

565

566 **Figure 2:** Geographic features of  $\text{CH}_4$  production from propionate degradation in paddy  
 567 field soils. The time till  $\text{CH}_4$  production reached the quantity corresponding to >90%  
 568 degradation of propionate added was defined as the time lapse. Three patterns can be  
 569 distinguished from 113 paddy field soils analyzed. (A) Pattern I, representing the high  
 570 rate group, comprises 45 soil samples (red dots), having time lapses of 13 d to 27 d. (B)  
 571 Pattern II, representing the medium rate group, includes 51 soil samples (green dots),  
 572 with time lapses of 28 d to 43 d. (C) Pattern III, the low rate group, contains 17 soil

573 samples (blue dots), having the time lapse of 44 d to 82 d. The CH<sub>4</sub> production curves of  
574 each soil in the three groups are illustrated in the lower-left corner of each panel. The  
575 CH<sub>4</sub> production curves within each group are colored randomly; the two vertical grey dot  
576 lines denote the ranges of time lapse. The error bars indicate the standard deviation of  
577 three replicates. All soil samples are displayed in maps with the pattern-specific samples  
578 highlighted by the increased sizes of the colored dots (red for Pattern I, green for Pattern  
579 II and blue for Pattern III).

580

581 **Figure 3:** The dynamics of short-chain fatty acids in paddy soils during anaerobic  
582 incubation. Propionate was added to a final concentration of 10 mM to each soil samples,  
583 which were incubated anaerobically until at least >90% of propionate had been oxidized.  
584 (A) The consumption of propionate. (B and C) The transient accumulations of acetate and  
585 butyrate, respectively. Soil samples are separated into three groups in accordance with  
586 three patterns revealed by methanogenesis, circles for Pattern I, triangles for Pattern II  
587 and diamonds for Pattern III. Soil samples within each group are colored randomly.

588

589 **Figure 4:** Geographical distribution of syntrophic potential and the correlative factors. (A)  
590 The map showing the functioning potential expressed as the time lapse for propionate  
591 degradation. The shorter the time lapse, the higher the functioning potential. (B) The map  
592 showing the functioning potential expressed as the maximum rate of CH<sub>4</sub> production,  
593 which was estimated during the period when CH<sub>4</sub> concentration was at the linear increase.  
594 The maps in (A) and (B) were build using Kriging interpolation method. The insets in the  
595 lower-left corner of each panel display the actual values of the time lapse (blue) and the

596 maximum rate (red) of CH<sub>4</sub> production, respectively. (C) The correlations of the time  
 597 lapse with the relative abundance of total syntrophs (Rel. abund. of synTotal), the mean  
 598 annual temperature (MAT), the contents of total soil sulfur (TS), organic matter (OM),  
 599 microbial biomass carbon (MBC), and the bioavailable contents of Fe, Cu and Mn (AF<sub>e</sub>,  
 600 ACu, AMn). The solid lines denote the least-squares linear regressions. Individual soil  
 601 samples are colored from blue to red in correspondence with their locations from the low  
 602 latitudes to the high latitudes.

603

604 **Figure 5:** The relative growth of propionate syntrophs during anaerobic incubation. The  
 605 relative growth was expressed as the logarithmic (log<sub>2</sub>) fold change (R) in relative  
 606 abundance of individual syntroph genera over the anaerobic incubation. (A-D) The  
 607 geographic distributions of relative growth of *Syntrophobacter*, *Syntrophomonas*,  
 608 *Pelotomaculum* and *Smithella*, respectively. (E) The mean relative abundances of four  
 609 syntroph genera at the end of the anaerobic incubation. (F) The mean log<sub>2</sub> fold change (R)  
 610 in relative abundance of four syntroph genera. Soil samples in (E) and (F) were clustered  
 611 into two geographic groups for the low latitude regions (South) and the high latitude  
 612 regions (North), respectively. The red asterisks denote the significant differences between  
 613 the “South” and “North” groups: \* 0.01 < P ≤ 0.05, \*\* P ≤ 0.01.

614

615 **Figure 6:** The geographical features of Gibbs free-energy changes for propionate  
 616 degradation. (A) The geographic distribution of Gibbs free-energy changes (ΔG') for  
 617 propionate oxidation corrected by temperature. The mean summer temperature (from  
 618 June to September) at sampling sites was used to correct the calculation of Gibbs free-  
 619 energy changes for the reaction: CH<sub>3</sub>CH<sub>2</sub>COO<sup>-</sup><sub>(aq)</sub> + 2H<sub>2</sub>O<sub>(l)</sub> → CH<sub>3</sub>COO<sup>-</sup><sub>(aq)</sub> + 3H<sub>2(g)</sub> +

620 CO<sub>2(g)</sub>. The values of standard Gibbs free-energy changes ( $\Delta G^0$ ) and enthalpy changes  
621 ( $\Delta H^0$ ) at 25°C (i.e. 298.15 K, 100 kPa, 1 M) equal to +73.5 kJ and +205.1 kJ per mol of  
622 propionate, respectively (for calculation details, see the Supplemental Material). The inset  
623 in the lower-left corner shows the difference in mean  $\Delta G'$  between two geographic groups  
624 for the low (South) and the high (North) latitude soils, respectively. The red asterisks  
625 indicate the significant difference at  $P \leq 0.001$ . (B) The correlation of synTotal relative  
626 abundance with temperature. (C) The correlation of time lapse for propionate degradation  
627 with temperature. (D) The correlation of maximum rate of CH<sub>4</sub> production with  
628 temperature. The solid lines in (B-D) denote the least-squares linear regressions.  
629 Individual soil samples are colored from blue to red in correspondence with their  
630 locations from the low latitudes to the high latitudes.

631

### 632 **Supporting Information**

633 **Figure S1.** Phylogenetic relationship of 528 dominant OTUs (top 2.5%) retrieved from  
634 113 paddy field soils. **Figure S2.** Neighbor-joining tree of 16S rRNA gene sequences  
635 related to five syntroph genera including representative sequences retrieved from this  
636 study and the reference sequences. **Figure S3.** Biogeographical distributions of five  
637 syntroph genera in paddy field soils across eastern China. **Figure S4.** Bacterial  
638 community in original soils and in soil samples after anaerobic incubations. **Figure S5.**  
639 The effect of temperature on the Gibbs free energy change of propionate oxidation via the  
640 methylmalonyl-CoA pathway. **Figure S6.** Temperature sensitivity of the Gibbs free  
641 energy changes of propionate oxidation through the methylmalonyl-CoA pathway vs the  
642 C<sub>6</sub> dismutation pathway. (PDF 2511 kb)

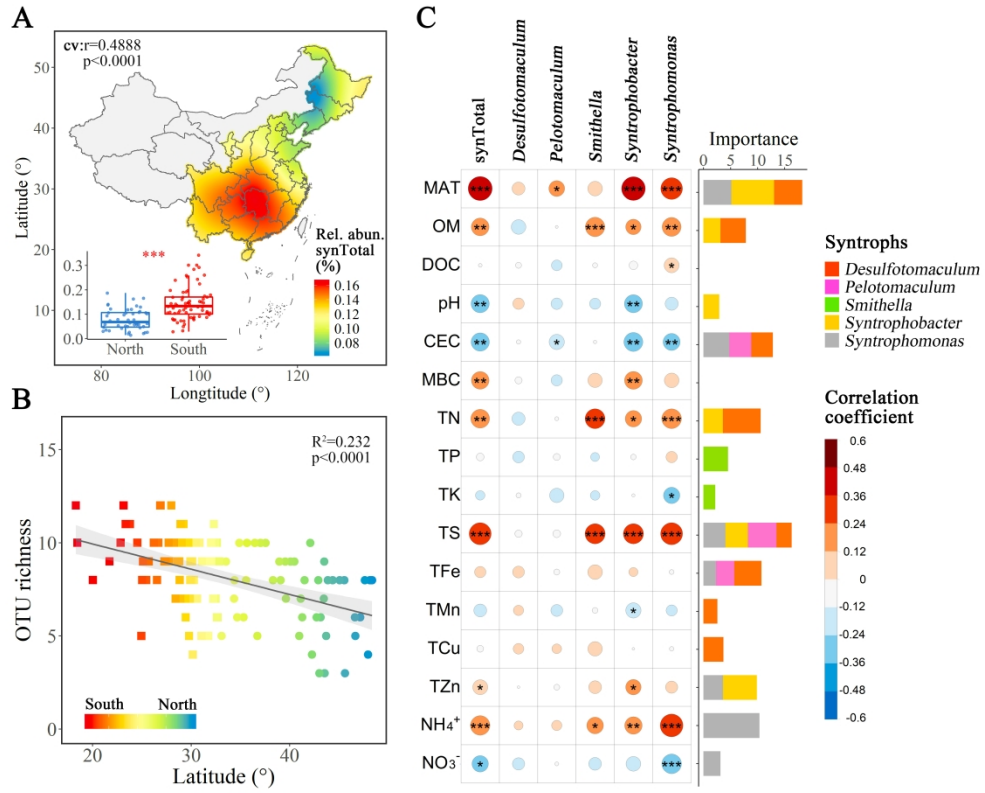


Figure 1: Biogeographic distribution of propionate syntrophs in paddy field soils across eastern China.

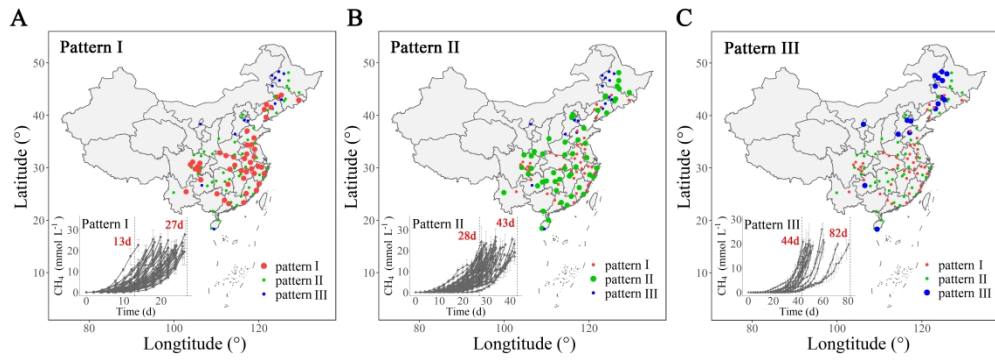


Figure 2: Geographic features of CH<sub>4</sub> production from propionate degradation in paddy field soils.

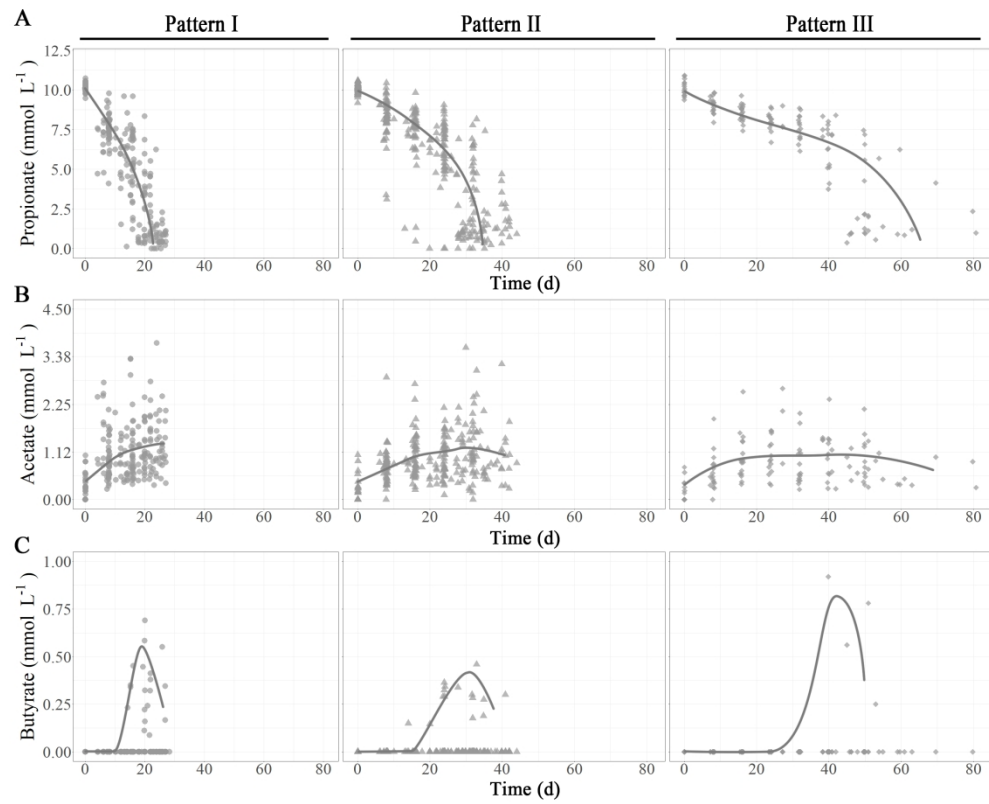


Figure 3: The dynamics of short-chain fatty acids in paddy soils during anaerobic incubation.



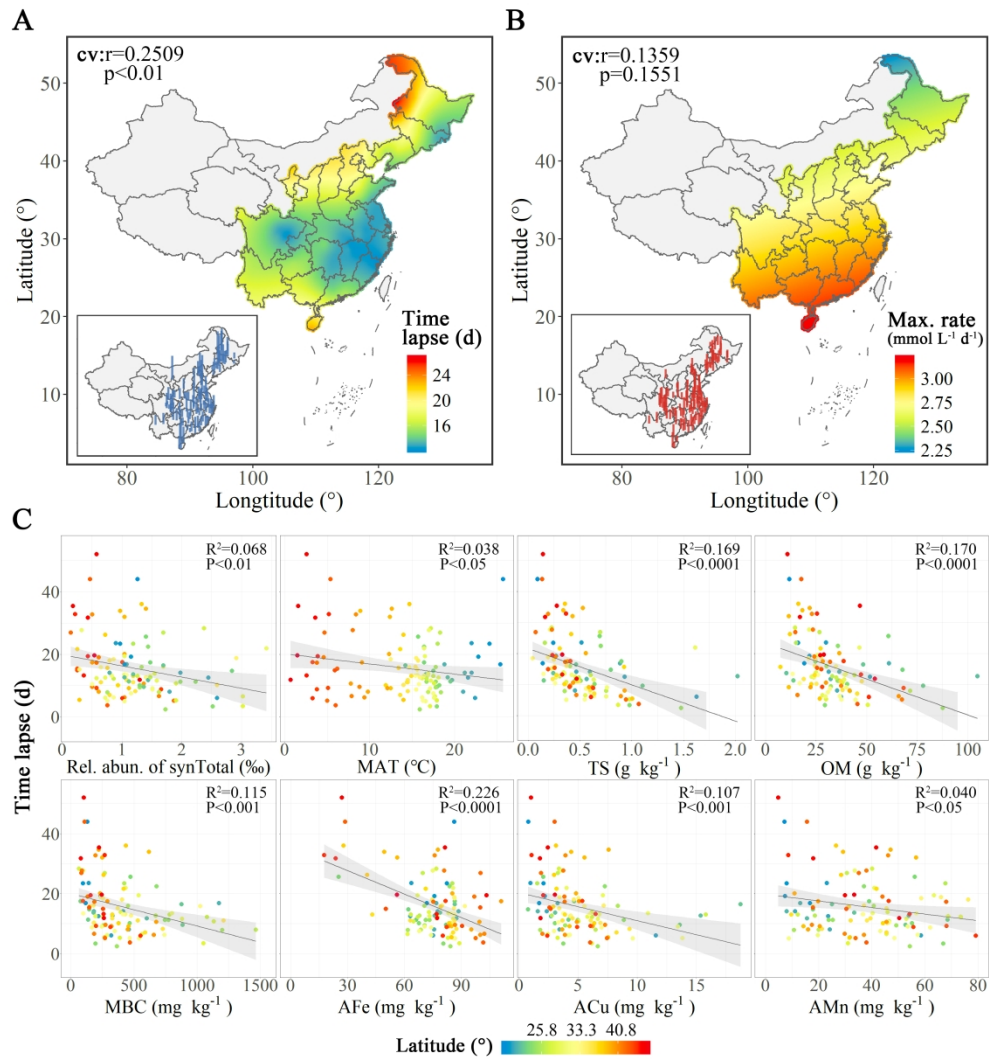


Figure 4: Geographical distribution of syntrophic potential and the correlative factors.

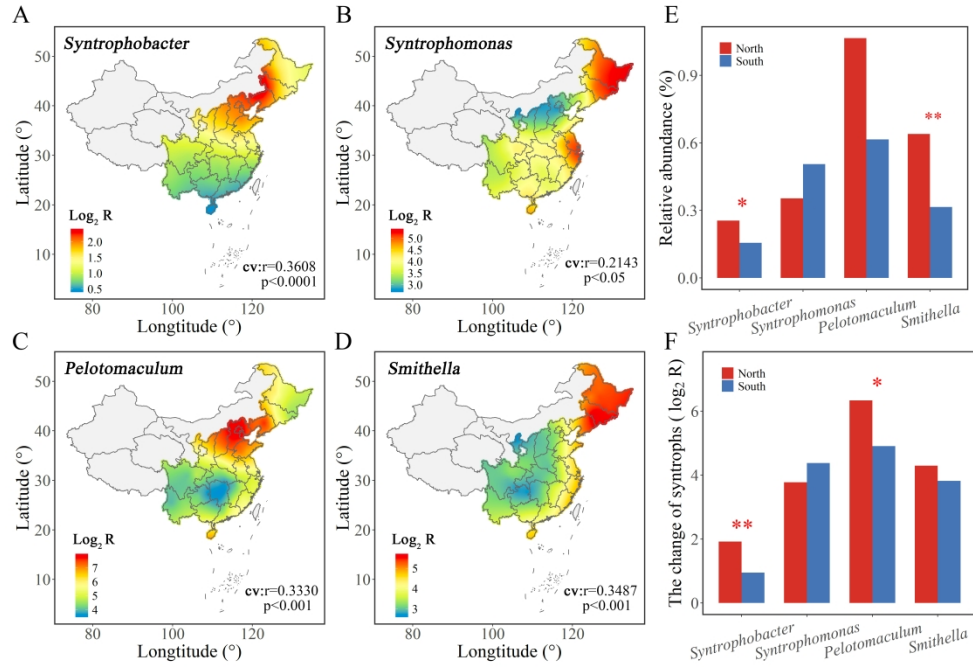


Figure 5: The relative growth of propionate syntrophs during anaerobic incubation.

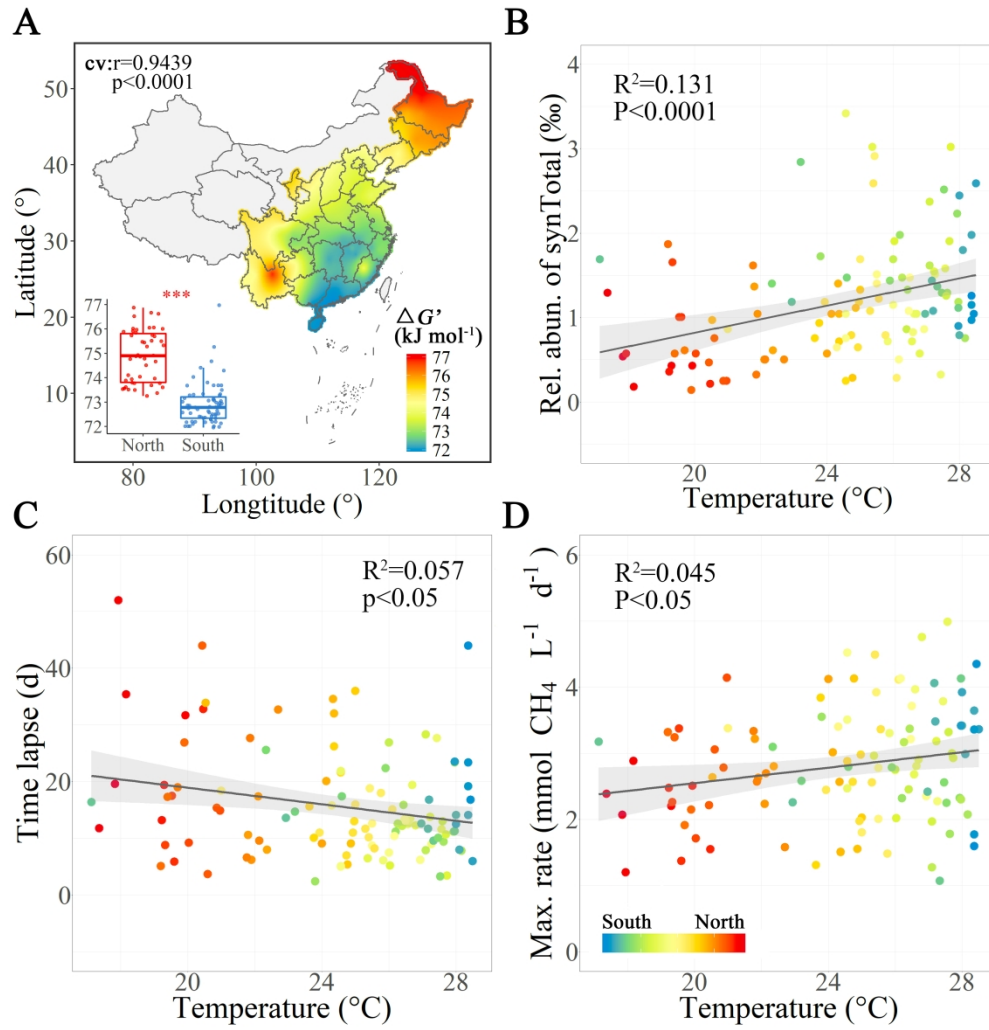


Figure 6: The geographical features of Gibbs free-energy changes for propionate degradation.