

LATE HOLOCENE HYDROLOGIC AND VEGETATION CHANGES AT TURNEFFE ATOLL, BELIZE, COMPARED WITH RECORDS FROM MAINLAND CENTRAL AMERICA AND MEXICO

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ABSTRACT

We present pollen and stable isotope ($\delta^{13}\text{C}$, $\delta^{18}\text{O}$, $\delta^{15}\text{N}$) data from a ~4 m core (TNF-1) of primarily mangrove peat taken from Turneffe Atoll, Belize. Radiocarbon (accelerator mass spectrometry) dates show that the record represents ca. 5000 years of sediment accumulation. Vegetation composition varied between dominant mangroves (primarily *Rhizophora mangle*) and Chenopodiaceae-Amaranthaceae, most likely *Salicornia bigelovii*. The pollen data, along with inferences from stable isotope analyses of bulk peat and fossil leaf fragments, indicate that marked environmental changes occurred at this location over the past ca. 5000 years. There was a transition between ca. 4100 and 2900 cal yr BP, from vegetation dominated by relatively tall mangroves (*R. mangle*) to one dominated by Chenopodiaceae-Amaranthaceae and then *Myrica*, most likely wax myrtle (*M. cerifera*). These changes bracket a period centered at ca. 3500 calibrated years before present, where there is a peak in the $\delta^{18}\text{O}$ of mangrove leaf fragments. This timing corresponds with other paleoenvironmental records of climate drying in Central America and increases the geographic and habitat scope (i.e., mangrove habitat) of records documenting these changes. Interpretations of shifts in mangrove habitat, however, require consideration of additional environmental influences, including changes in groundwater hydrology and relative inputs of seawater and freshwater (i.e., precipitation) during the Holocene.

INTRODUCTION

The primary goal of this research was to examine the vegetation history of a mangrove site on the Turneffe Islands, Belize. Previous studies of the paleoecology of mangrove ecosystems in Belize have largely been restricted to cays inside the Belize Barrier Reef (Wooller et al., 2004, 2007). Research at the Turneffe Islands (Gischler and Hudson, 1998; Gischler and Lomando, 1999; Gischler, 2003; Gischler et al., 2003) identified mangrove peat and reported that mangrove ecosystems developed at the Turneffe Atoll beginning ca. 6000 cal yr BP (calibrated years before present) (Gischler and Hudson, 1998; Gischler, 2003). This body of research, however, did not document the paleoecological changes that took place in the vegetation that developed on the atoll. We present paleoenvironmental data from a dated (basal age ca. 6100 cal yr BP) ~4 m peat core (core TNF-1) from a mangrove habitat on the Turneffe Atoll. Accelerator mass spectrometry (AMS) radiocarbon measurements of material preserved in the core were used to determine the timing of observed changes throughout the sequence and to allow comparison with other paleoecological studies in Central America. Pollen analyses of samples from the core illustrate vegetation dynamics. Elemental (carbon and nitrogen) composition was combined with stable isotope analyses of

total organic matter ($\delta^{15}\text{N}$, $\delta^{13}\text{C}$) and mangrove leaf fragments ($\delta^{15}\text{N}$, $\delta^{13}\text{C}$, $\delta^{18}\text{O}$) to infer past environmental conditions at the site. Relatively high $\delta^{15}\text{N}$ values of sediment and leaves in mangrove ecosystems can indicate the input of an organic nitrogen source, such as bird guano (Wooller et al., 2003a). Lower $\delta^{15}\text{N}$ values can be used to infer the stand structure of mangroves, with very negative values ($< -3\text{‰}$) indicating the uptake of atmospheric ammonia by mangroves that are nutrient limited (e.g., phosphorus) (Fogel et al., 2008). The $\delta^{13}\text{C}$ values of mangrove leaves can be used to infer mangrove stand structure (Wooller et al., 2003b, 2004, 2007), with relatively low $\delta^{13}\text{C}$ values ($\approx -27\text{‰}$) typical of tall mangroves (McKee et al., 2002; Wooller et al., 2003b). In addition to indicating low stand structure, higher $\delta^{13}\text{C}$ values in mangrove leaves can result from increased photosynthetic water-use efficiency (PWUE), as, for example, in more saline conditions (Lin and Sternberg, 1992a, 1992b, 1992c). In previous paleoenvironmental studies of mangrove peat, these types of analyses have usually been restricted to mangrove leaves preserved in peat (Wooller et al., 2004, 2007; Monacci et al., 2009). Here we investigate whether long-term trends in the $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ of mangrove leaves are also reflected in the total organic matter (TOM) of mangrove peat. We have used $\delta^{18}\text{O}$ values from analyses of mangrove leaves to infer the source water available to mangroves (Sternberg and Swart, 1987; Lin and Sternberg, 1994; Wooller et al., 2007; Monacci et al., 2009), and to this end, we analyzed $\delta^{18}\text{O}$ in preserved mangrove leaves from core TNF-1.

STUDY SITE

The Turneffe Islands are surrounded by deep water and have a breakwater reef encircling a main lagoon (Fig. 1). Mangrove species that encircle the central lagoon, which is ~8 m deep (Gischler, 2003), include *Rhizophora mangle* L. (red mangrove), *Laguncularia racemosa* (L.) Gaertn. f. (white mangrove), and *Avicennia germinans* (L.) Stearn (black mangrove). Also present on the northern end of Turneffe are other tree species, including *Myrica cerifera* L. along with other littoral woodland species, palms, and *Conocarpus erecta* L. (Murray et al., 2003). At our coring site, the vegetation consisted exclusively of dwarf *R. mangle*. Other plants present included epiphytic ferns and orchids. The geological setting of the Turneffe Islands has been described previously (Gischler and Hudson, 1998; Gischler and Lomando, 1999; Gischler, 2003; Gischler et al., 2003). In brief, the general facies succession begins with a basal soil overlying Pleistocene limestone, and is followed by mangrove peat and carbonate sediments (Gischler, 2003). One of the objectives of this study was to provide the longest paleoecological record from a mangrove habitat in the Turneffe Islands. We therefore used the sedimentary geological analyses of Gischler (2003) to guide selection of our coring location. We selected a location near the site of a previous core that yielded continuous mangrove peat (Gischler, 2003, core location T17) and

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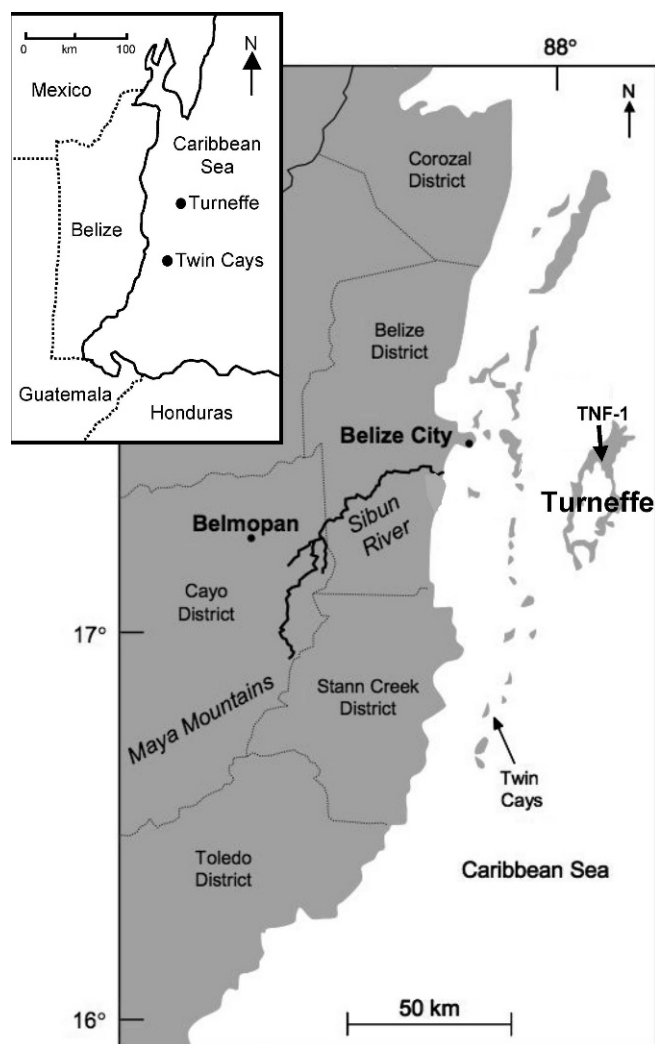


FIGURE 1—Site map of coring location (core TNF-1) at Turneffe, Belize.

seemed to have sustained mangroves for the longest period at Turneffe.

METHODS

An Eijkelpkamp Russian peat corer was used to obtain a 440-cm-long, continuous core (TNF-1) that was composed primarily of peat. Core sections, up to 50 cm, were wrapped in plastic, and kept in coolers at the Spanish Bay Conservation and Research Center on Spanish Lookout Cay. Sections were transported to the Water and Environmental Research Center (WERC) at the University of Alaska Fairbanks (UAF), where they are archived at $\sim 4^{\circ}\text{C}$.

AMS radiocarbon measurements (^{14}C) were made on wood pieces, macroscopic mangrove roots, and bulk peat at the National Ocean Sciences Accelerator Mass Spectrometry Facility (NOSAMS) at Woods Hole Oceanographic Institute (WHOI). All ^{14}C dates were calibrated using CALIB 5.0.2 and rounded to the nearest decade.

Pollen Analysis

All pollen preparation and analytical procedures were performed in the Department of Palynology and Climate Dynamics, University of Göttingen. At ≤ 20 cm intervals, a 1 cm^3 sample was prepared using previously described methods (Wooller et al., 2004, 2007). One tablet of exotic *Lycopodium* spores ($18,583 \pm 762$ spores/tablet) was added prior

to processing to calculate pollen concentrations ($\text{grains}/\text{cm}^3$). Samples were mounted in glycerin gelatin medium. At least 300 pollen grains were identified for each pollen sample, with a few exceptions that were counted to >100 due to low pollen concentrations. Pollen grains were identified using previously published procedures (Wooller et al., 2004, 2007). Pollen data are reported as percentages of the total pollen sum, which excludes aquatic taxa, fern and fungal spores, and microforaminifera. Pollen diagrams were created and zoned using TILIA, TILIAGRAPH, and CONISS (Grimm, 1987).

Elemental (C and N) and Stable (C, N, O) Isotope Analyses

Bulk peat samples were freeze dried, ground, and acidified with 2 N HCl to remove carbonate in preparation for elemental and stable isotopic analyses of TOM. Mangrove leaf fragments were also picked from bulk peat at points throughout the core for stable isotope analyses. Stable isotope analyses were performed at the Alaska Stable Isotope Facility (ASIF) at UAF and reported in standard delta (δ) notation (‰). Due to the low organic content of the basal portion of core TNF-1, larger sample masses (2–3 mg) were used to yield reliable stable isotope data. Elemental composition (%C and %N, wt. %) and stable isotope ratios ($\delta^{13}\text{C}$ and $\delta^{15}\text{N}$, ‰) were measured using a Costech Instruments Elemental Analyzer (EA) coupled with a ConFlo III and a Thermo Electron Delta V Plus Isotope Ratio Mass Spectrometer (IRMS). The $\delta^{13}\text{C}$ values are reported relative to Vienna Pee Dee Belemnite (VPDB), and the $\delta^{15}\text{N}$ values are reported relative to atmospheric nitrogen. Peptone was used as a reference standard to determine analytical precision (defined here as 1 standard deviation on 15 replicate analyses of peptone throughout the analytical run): 0.9‰ (wt. %N), 2.8‰ (wt. %C), 0.5‰ ($\delta^{15}\text{N}$), and 0.3‰ ($\delta^{13}\text{C}$). The $\delta^{18}\text{O}$ of leaf fragments from core TNF-1 was analyzed using a previously published protocol and instrumentation (Wooller et al., 2007). The $\delta^{18}\text{O}$ data are expressed relative to Vienna Standard Mean Ocean Water (VSMOW). Benzoic acid ($\text{C}_7\text{H}_6\text{O}_2$) was used to determine analytical precision (defined here as 1 standard deviation resulting from $n = 10$ analyses of benzoic acid) associated with $\delta^{18}\text{O}$ and ‰O analyses for each run, which were 0.4‰ and 2‰, respectively.

RESULTS

All data from core TNF-1 are presented relative to four pollen zones: TNF-I (375–260 cm), TNF-II (260–210 cm), TNF-III (210–135 cm), and TNF-IV (135–0 cm). These zones were established through a Constrained Incremental Sum of Squares (CONISS) analysis performed in TILIA using the pollen data.

Chronology and Sediment Characteristics

Table 1 presents the AMS radiocarbon data (^{14}C) and their corresponding calibrated dates and ranges. Ages for all depths of core TNF-1 and sedimentation rates were interpolated using four linear equations between calibrated AMS radiocarbon dates. The age at the top of the core (0–1 cm) is assumed to be modern: -55 cal yr BP. The lithology of core TNF-1 is divided into two main parts, with the bottom ~ 60 cm composed of grey mud with a few fine plant macrofossils and the upper portion composed of reddish-brown to brown peat. Basal rock was encountered at the bottom of the sequence. The thickness of the peat is consistent with previous data from Turneffe (Gischler, 2003). The boundaries between the pollen zones (TNF-I to TNF-IV) are at 4100, 3900, and 3300 cal yr BP, respectively (Fig. 2).

Pollen, Elemental, and Stable Isotope Analyses

Pollen Zone TNF-I (6100–4100 cal yr BP).—Poor pollen preservation in the basal portion (440–380 cm) of the core, which was composed

TABLE 1—AMS radiocarbon dates, analyzed at National Ocean Sciences Accelerator Mass Spectrometry Facility (NOSAMS) for core TNF-1 from the Turneffe Atoll, Belize.

NOSAMS accession number	Material dated	Depth (cm)	¹⁴ C yr BP	Age error	Calibrated age yr BP	Calibrated age range yr BP (2 sigma)
OS-65273	Peat	73.5	2480	30	2540	2366–2718
OS-65271	Wood	224.5	3720	30	4070	3990–4150
OS-66026	Fine roots	295.5	3750	45	4120	3994–4242
OS-65272	Peat	439.5	5390	30	6160	6026–6284

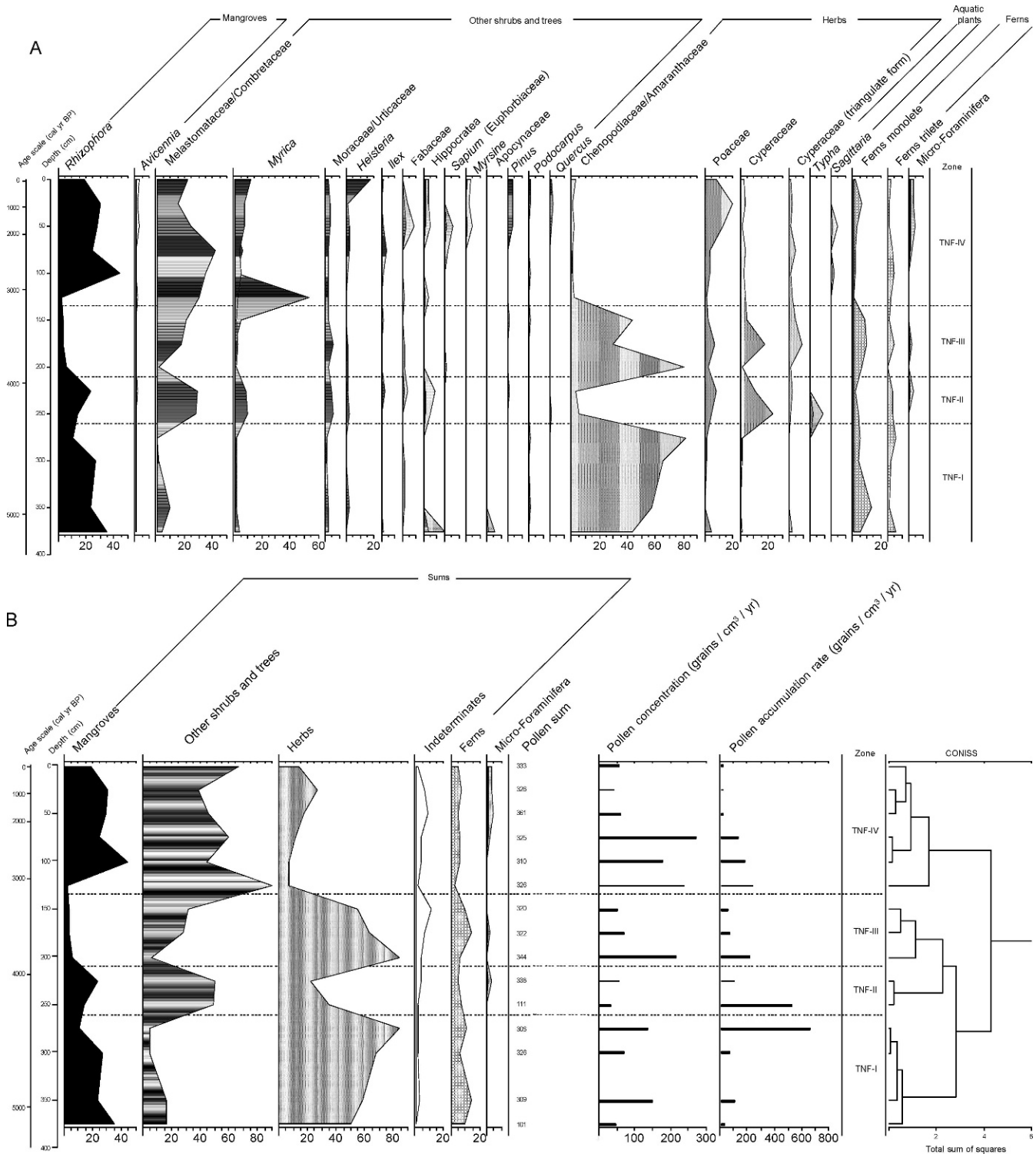


FIGURE 2—Summary (A) and percentage (B) pollen diagrams for core TNF-1 from Turneffe (cal yr BP = calibrated years before present).

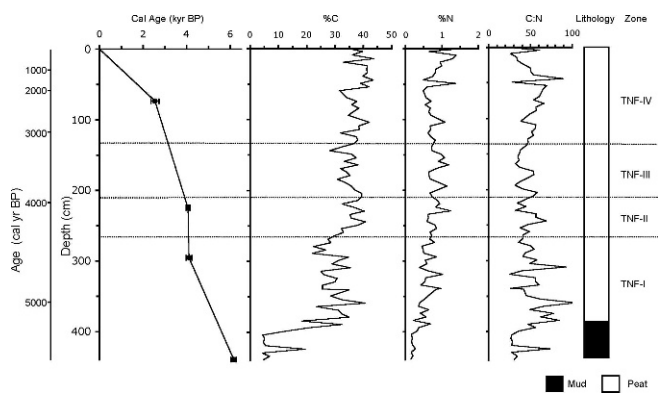


FIGURE 3—The chronology, C:N ratio, %C, %N, and lithology of core TNF-I from Turneffe (cal yr BP = calibrated years before present).

of mud with low organic carbon content (<10% C) (Fig. 3), prevented acquisition of reliable pollen data. The pollen profile in Figure 2 begins at ca. 5300 cal yr BP. The zone begins with a number of pollen types including Hippocratea, Ilex, Apocynaceae, and Poaceae, along with fern spores. The percentage composition of pollen in zone TNF-I consists primarily of Rhizophora pollen, which declined from ~35% near the start of the zone to ~10% at the end of the zone, along with ChenAm (Chenopodiaceae-Amaranthaceae pollen, which increased from ~45% at the start of the zone to ~80% at the end of the zone. Microforaminifera were present in very low numbers, and bulk sedimentation during this pollen zone was relatively low (0.07 cm/yr). Mud at the base of the core had low total nitrogen (Fig. 3) that had a relatively high $\delta^{15}\text{N}$ value, up to 6‰ (Fig. 4). The $\delta^{13}\text{C}$ composition of the TOM was relatively constant throughout the zone, with a mean of -27‰ (sd = 0.2‰) (Fig. 4). The $\delta^{13}\text{C}$ data from TOM tracked the $\delta^{13}\text{C}$ data derived from mangrove leaf fragments throughout the entire core (Fig. 5), yielding a strong positive linear relationship ($R^2 = 0.72$). No significant relationship was evident for $\delta^{15}\text{N}$ analyses of TOM and

leaf fragments ($R^2 = 0.07$), although the range of $\delta^{15}\text{N}$ variation in the entire core was considerably less than the range of $\delta^{13}\text{C}$ values. The $\delta^{13}\text{C}$ findings have significance for future isotope studies of mangrove peat. In the past, $\delta^{13}\text{C}$ analyses of mangrove peat have been restricted to preserved leaf fragments (Wooller et al., 2004, 2007), but it is time-consuming to remove and prepare leaf fragments. These findings show that $\delta^{13}\text{C}$ analyses of TOM in mangrove peat cores provides the same paleoecological information as data derived from fossil mangrove leaf fragments.

Pollen Zone TNF-II (4,100–3,900 cal yr BP).—Pollen zone TNF-II was marked by a slight increase in the percentage of Rhizophora pollen and a marked increase in the percentage of Melastomataceae pollen (Fig. 2). The percentage of ChenAm pollen declined markedly at the start of the zone, at which point the percentage of Myrica type, Moraceae-Urticaceae, and Cyperaceae pollen also increased. Additional pollen types were evident during this zone, including Typha, Sagittaria, Poaceae, Hippocratea, Ilex, Pinus, and Fabaceae. These changes in pollen occurred during a period when the sedimentation rate increased (≥ 0.5 cm/yr) (Fig. 3).

Pollen Zone TNF-III (3,900–3,300 cal yr BP).—The peat accumulation rate was slower (0.11 cm/yr) in this zone relative to the previous one. This change in sedimentation rate corresponds with marked changes in both pollen and stable isotopic composition of both TOM and mangrove leaf fragments in the core. Decreases in the percentages of Rhizophora and Melastomataceae pollen were matched by a marked increase in the percentage of ChenAm pollen. Low percentages of mangrove pollen (Fig. 2A) indicate that these plants were not likely present at the site (Ellison, 2008). Other features in the pollen data included a peak in the percentage of Cyperaceae pollen and a rise in the number of fern spores. Myrica-type pollen was evident towards the end of the zone. Although variations in the elemental composition characteristics (i.e., C%, N%, and C:N) of the peat are minor (Fig. 3), the stable isotopic signatures of the leaf fragments ($\delta^{13}\text{C}$ and $\delta^{18}\text{O}$) and peat ($\delta^{13}\text{C}$) shifted to some of the highest values ($\delta^{13}\text{C}$ up to

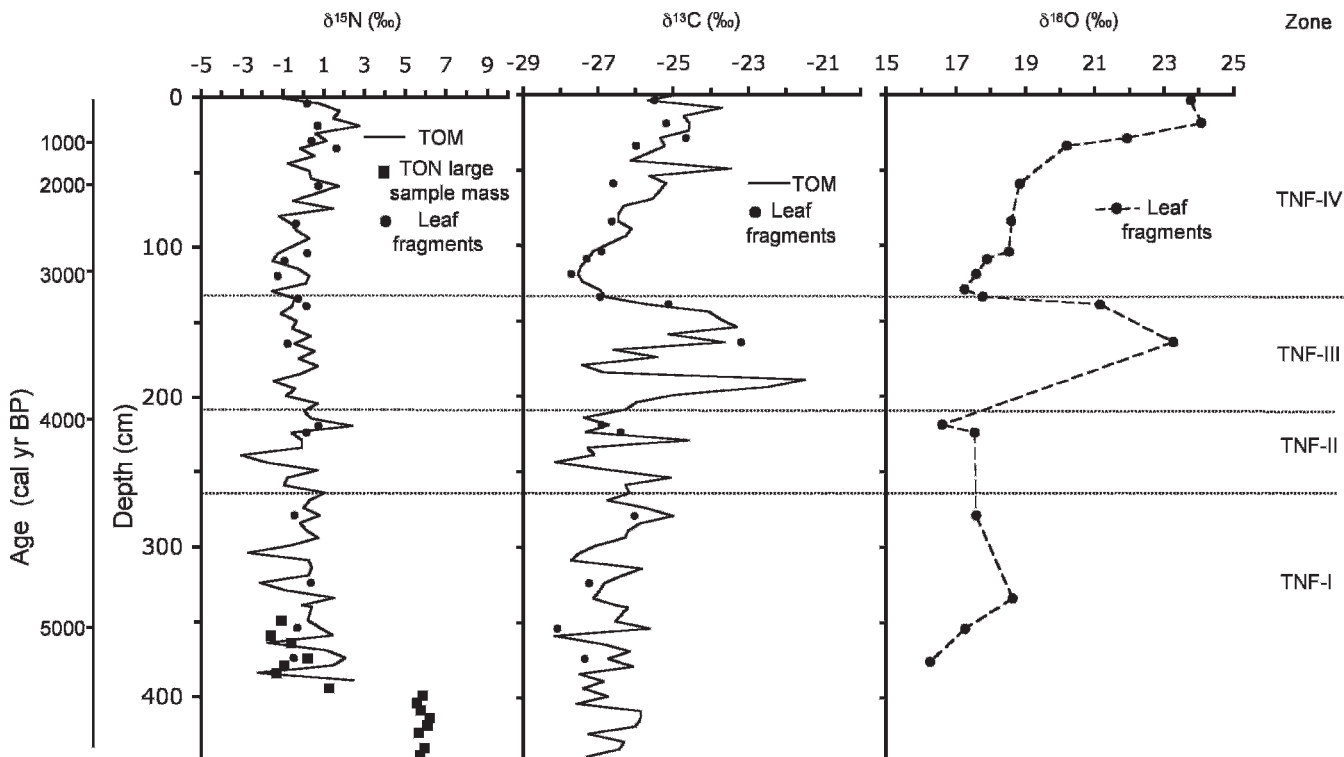


FIGURE 4—Stable nitrogen, carbon, and oxygen isotope data from analyses of organic components (leaves and total organic matter) from core TNF-I from Turneffe; TOM = total organic matter; TON = total organic nitrogen.

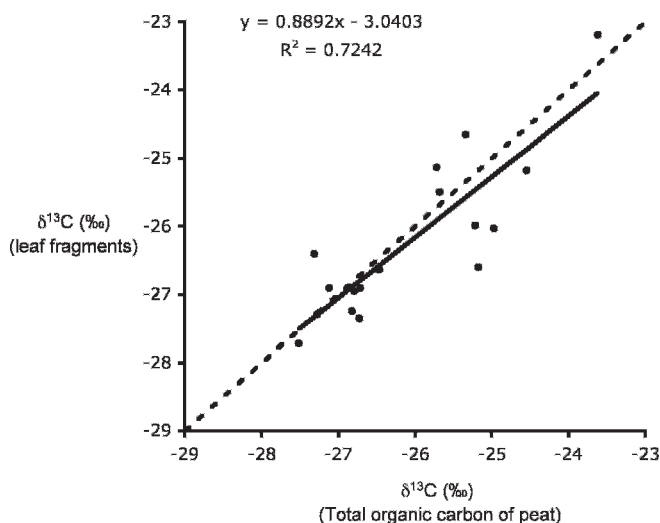


FIGURE 5—The relationship between the $\delta^{13}\text{C}$ of leaf fragments and total organic matter from peat samples taken from core TNF-1 from Turneffe (dotted line = 1:1 line).

$\sim -21.5\text{‰}$ and $\delta^{18}\text{O}$ up to 23‰ recorded (Fig. 4). A minor change in the $\delta^{15}\text{N}$ of peat or leaf fragments was recorded during this zone (mean = 0‰).

Pollen Zone TNF-IV (3300 cal yr BP to present).—The lowest peat accumulation rate (0.03 cm/yr) was recorded in the top half of pollen zone TNF-IV. The zone began with a marked increase in *Myrica*-type pollen, from $\sim 5\%$ to $\sim 55\%$. This was followed by a sharp increase in *Rhizophora* pollen, from $<5\%$ to $\sim 45\%$, with percentages $\geq 20\%$ sustained through the rest of the zone. Other notable changes included an increase in *Poaceae* pollen approximately halfway through the zone and the presence of microforaminifera. Changes in the elemental compositions (C and N) of peat during this zone remained small ($\text{C}\% \pm 1$, $\text{sd} = <4\%$; $\text{N}\% \pm 1$, $\text{sd} = <0.3\%$). This contrasted with $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ data derived from this zone, which showed an increase from $\sim -28\text{‰}$ to $\sim -24.5\text{‰}$ in the $\delta^{13}\text{C}$ of peat and leaf fragments and an increase from $\sim 17\text{‰}$ to $\sim 24\text{‰}$ in the $\delta^{18}\text{O}$ of leaf fragments (Fig. 4).

INTERPRETATION AND DISCUSSION

The sediment sequence at core site TNF-1 on Turneffe consists of limestone bedrock overlain by ~ 60 cm of mud that transitions to mangrove peat. In every respect, this sequence is consistent with that previously described by Gischler (2003). Although no micromorphological features were examined in the mud, high $\delta^{15}\text{N}$ values relative to the rest of the core ($\sim 6\text{‰}$) imply the sediment was receiving organic nitrogen, likely marine-derived and possibly from bird guano. This chemical characteristic has previously been documented from the basal mud that transitioned to mangrove peat during the Holocene at Twin Cays, Belize (Wooller et al., 2004, 2007) as well as in a modern mangrove habitat in Belize (Man-O-War Cay) (Wooller et al., 2003a). Mangroves were established at or near the coring site by at least 5,200 cal yr BP, and the stand seems to have been essentially monospecific, composed primarily of *R. mangle*. The habitat was also composed of a vegetation type containing ChenoAm. There are a number of Amaranthaceae species in the extant Belize flora (Balick et al., 2000); however, there are only two species of Chenopodiaceae (*Chenopodium ambrosioides* L. and *Salicornia bigelovii* Torr.) (Balick et al., 2000). Given the lithology (mangrove peat) associated with the presence of ChenoAm pollen, it is probable that this pollen represents *Salicornia*, which coexists with *R. mangle* in some mangrove habitats. If the high percentage of this pollen type is accepted as indicating the presence of *Salicornia*, then this pollen would tend to indicate a particular

hydrological setting (Alexander and Dunton, 2002; Forbes and Dunton, 2006). For instance, Geographic Information System analyses of the Nuecas estuary in Texas have shown that considerable vegetation change can occur following sporadic freshwater inundation of coastal habitats, including increased cover of the annual succulent *S. bigelovii* (Alexander and Dunton, 2002). The duration of freshwater inundation also seems to be critical to the growth of this plant, with brief periods occurring at specific times in the year, allowing seed germination as the stress of hypersaline conditions is removed (Alexander and Dunton, 2002). The ecology of *S. bigelovii* suggests that the site at Turneffe may have experienced brief freshwater events (precipitation input) in the past that punctuated predominantly saline conditions. This low salinity may have come about because of decreased seawater inundation and/or sporadic input of precipitation at the site. Absence of microforaminifera supports the scenario that the site may not have been regularly inundated with seawater until ca. 4100 cal yr BP (pollen zone boundary). Sedimentation rate also increased to the highest recorded from the core until 3,900 cal yr BP. This period of increased sedimentation rate accompanied some marked ecosystem changes. For instance, percentage of *Rhizophora* pollen increased slightly, and the percentage of ChenoAm pollen decreased sharply. The low $\delta^{13}\text{C}$ values of mangrove leaf fragments during the period from ca. 4100–3900 cal yr BP indicate the site was inhabited by relatively tall mangroves (McKee et al., 2002; Wooller et al., 2003b). This period likely represents a time when the site was being inundated with seawater more frequently.

After 3,900 cal yr BP and until 3,300 cal yr BP, the vegetation composition at the site changed significantly again, with the percentage of mangrove pollen dropping to its lowest values for the entire core, most likely indicating the absence of *R. mangle* from the site and the general vicinity (as reviewed by Ellison, 2008). The decrease in *R. mangle* at the site, which plays an important role in peat dynamics (McKee et al., 2007), may account for the drop in sedimentation rate after 3,900 cal yr BP. This brief and dramatic change in the vegetation composition at this site is consistent with the view that mangroves can come and go rapidly (Alongi, 2008). The high percentage of ChenoAm pollen during this period probably indicates that *Salicornia* was present at the site and that the site was not inundated with seawater as frequently (Alexander and Dunton, 2002). This evidence for community change is accompanied by indications that the hydrology also changed. One indication of this is a shift towards higher $\delta^{18}\text{O}$ values (17‰ to $\sim 24\text{‰}$) from leaf fragments in the core. The $\delta^{18}\text{O}$ of mangrove leaves is primarily dictated by the $\delta^{18}\text{O}$ of the source water used by mangrove trees (Wooller et al., 2007). Wooller et al. (2007) previously suggested that shifts to higher $\delta^{18}\text{O}$ in fossil mangrove leaves probably were driven by an increased proportion of seawater relative to rainwater (Sternberg and Swart, 1987; Lin and Sternberg, 1994). In the record from core TNF-1, the $\delta^{18}\text{O}$ record may represent a similar shift, which may be overprinted by an evaporative alteration of the source water available to mangroves. Evaporation increases the $\delta^{18}\text{O}$ of water, and an evaporative influence is suspected because the $\delta^{18}\text{O}$ leaf values between ca. 3600 and 3300 cal yr BP and also from ca. 1000 cal yr BP (Fig. 4) are higher by $\sim 4\text{‰}$ than fossil leaf data from Twin Cays (Wooller et al., 2007) or Spanish Lookout Cay (Monacci et al., 2009). Moreover, in contrast to the record from Twin Cays, the $\delta^{18}\text{O}$ record from core TNF-1 is accompanied by a vegetation shift, possibly towards a vegetation composed of *Salicornia*. As discussed above, *Salicornia* tends to inhabit sites that experience sporadic and brief episodes of freshwater inundation (Alexander and Dunton, 2002). Unlike the $\delta^{18}\text{O}$ record from Twin Cays, the TNF-1 core $\delta^{18}\text{O}$ data is positively correlated with a shift towards higher $\delta^{13}\text{C}$ values in mangrove leaves (Fig. 4). Higher $\delta^{13}\text{C}$ values in mangroves can result from a change in stand structure (McKee et al., 2002; Wooller et al., 2003b), indicating a shift towards a dwarf mangrove stand. These values can also indicate a shift towards greater photosynthetic water-

use efficiency (PWUE) in mangroves (Lin and Sternberg, 1992a, 1992b, 1992c). The lack of shift towards more negative $\delta^{15}\text{N}$ values (Fig. 4) during this time, which can relate to nutrient (P) limitation (McKee et al., 2002) and result from an increased reliance on atmospheric ammonia as a nitrogen source (Fogel et al., 2008), implicates factors other than nutrient-driven changes in the mangrove stand structure at the Turneffe site between 3900 and 3300 cal yr BP. Given the other lines of evidence (pollen and $\delta^{18}\text{O}$ data), the most likely explanation for the change in $\delta^{13}\text{C}$ values is a shift towards predominantly more saline or drier conditions, resulting from either sporadic input of precipitation, decreased frequency of seawater inundation, and/or greater evaporation during that time. The timing of this drying or more saline conditions at Turneffe (3900–3300 cal yr BP) correlates with a peak of climatic drying at 3500 cal yr BP in the lowlands of Guatemala (Mueller et al., 2009). Our peak in $\delta^{18}\text{O}$ from analyses of mangrove leaf fragments at ca. 3500 cal yr BP also correlates with a peak in climatic drying (Mueller et al., 2009) as well as a peak in $\delta^{18}\text{O}$ values in sediments from Haiti (Hodell et al., 1991) and northern Venezuela (Curtis et al., 1999), which have been interpreted as documenting late Holocene climatic drying (Mueller et al., 2009). The period of climatic drying described by Mueller et al. (2009) between 4000 and 3500 cal yrs BP in the lowlands of Guatemala corresponds with a period of marked vegetation and hydrological changes in our record from Turneffe Atoll. Our record broadens the geographic range of these environmental changes in Central America to include mangrove ecosystems, which were also relied upon by Mayan populations (McKillop, 2005). Nevertheless, we must consider the additional environmental influences, which could include changes in groundwater hydrology and the degree of inundation with seawater relative to freshwater (i.e., precipitation) during the late Holocene.

After 3300 cal yr BP, the site seems to shift towards a habitat composed of *Myrica*, which is a species typical of a more littoral forest (Russell, 1997). Wax myrtle (*Myrica cerifera*) is an early successional tree-shrub species and one of the first woody plants to invade secondary dunes and beach meadows, naturally reseeding in disturbed sites from adjacent forests. Wax myrtle occupies diverse habitats including fresh-to-slightly-brackish banks and shores, flats, interdune swales, pine and palmetto flatwoods and savannas, cypress-gum ponds and swamps, wet and dry prairies, pitcher-plant bogs, upland mixed woodlands, old fields, and fence and hedge rows. Such sites include tree islands, cypress heads, and wet and dry hummocks (Russell, 1997). In the sediment record this vegetation shift is accompanied by a marked shift towards lower $\delta^{18}\text{O}$ values from ca. 3300 to ca. 2000 cal yr BP, lower $\delta^{13}\text{C}$ values, and a decrease in the percentage of chenopod pollen. All of these shifts can be explained if the site was receiving an increasing proportion of precipitation relative to seawater. Changing salinity at the site could have been a consequence of increased precipitation, decreased inundation with seawater, and decreased evaporative losses, either singly or in combination. Wet conditions were recorded for the period ca. 3500 cal yr BP to ca. 1700 cal yr BP on the Yucatan Peninsula (Hodell et al., 1995; Curtis et al., 1996) and the Caribbean island of Hispaniola (Hodell et al., 1991). *Rhizophora mangle* seems to reestablish at the Turneffe site after ca. 2600 cal yr BP. Both the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ increase begin at ca. 2600 cal yr BP, and *Avicennia germinans* pollen is evident. Since black mangrove is a low pollen producer, this relatively minor change could indicate that it was growing close to the coring site (Behling et al., 2001; Behling and Costa, 2001). All of these sediment variables indicate an increasingly saline habitat after ca. 2000 cal yr BP.

CONCLUSIONS

The record from core TNF-1 shows that the mangrove habitat that exists at the study site on Turneffe Atoll today experienced hydrologic and vegetation changes since it became established ca. 5000 cal yr BP.

In some instances, these changes occurred over short (submillennial) time scales. Hydrological changes, inferred from $\delta^{18}\text{O}$ changes, correlate with some of the marked vegetation changes. The most marked changes in vegetation composition, mangrove physiology, and site hydrology seem to have occurred in the TNF-1 record ca. 4000–2500 cal yr BP, a period when other mangrove ecosystems also demonstrated marked ecosystem and environmental changes in Belize (Wooller et al., 2004, 2007; Monacci et al., 2009). Further paleoenvironmental evidence is needed to examine the factors that influenced the coastal region during the Holocene (Leyden, 2002). Our findings add to the growing suite of records demonstrating how mangrove ecosystems changed during the Holocene in Belize. These types of records will ultimately lead to a synthetic picture of how mangroves responded to Holocene environmental changes, including changes in sea level.

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