

1 Modeling the fate of UV filters in subsurface: co-metabolic degradation 2 and the role of biomass in sorption processes

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7 Abstract

8 Ultraviolet filters (UVFs) are emerging organic compounds found in most water systems. They
9 are constituents of personal care products, as well as industrial ones. The concentration of UVFs
10 in the water bodies in space and time is mostly determined by degradation and sorption, both
11 processes being determinant of their bioavailability and toxicity to ecosystems and humans.
12 UVFs are a wide group of compounds, with different sorption behavior expected depending on
13 the individual chemical properties (pK_a , K_{oc} , K_{ow}). The goal of this work is framed in the context
14 of improving our understanding of the sorption processes of UVFs occurring in the aquifer; that
15 is, to evaluate the role of biomass growth, solid organic matter (SOM) and redox conditions in
16 the characterization of sorption of a set of UVFs. We constructed a conceptual and a numerical
17 model to evaluate the fate of selected UV filters, focused on both sorption and degradation. The
18 models were validated with published data by Liu et al. (2013), consisting in a suite of batch
19 experiments evaluating the fate of a cocktail of UVs under different redox conditions. The
20 compounds evaluated included ionic UV filters (Benzophenone-3; 2-(3-t-butyl-2-hydroxy-5-
21 methylphenyl)5-chloro-benzotriazole; 2-(2'-hydroxy-5'-octylphenyl)-benzotriazole) and neutral
22 ones (octyl 4-methoxycinnamate; and octocrylene).

23 Keywords

24 UV filters; benzophenone; sorption; numerical modeling; biomass; co-metabolic degradation

25 **1 Introduction**

26 Ultraviolet filters (UVFs) are emerging organic compounds (EOCs) commonly found in water
27 systems (Jurado et al., 2014). They are found worldwide as they are constituents of a large
28 number of widely used personal care products (e.g., sunscreens, perfumes, creams, body lotions
29 or shampoos), but also of a number of industrial applications (furniture varnishes, paints). UVFs
30 form a wide group of organic compounds with different chemical properties; the most common
31 ones found are Benzophenone-3 (BP-3), plus a number of its derivatives, 3-(4-
32 methylbenzylidene) camphor (4-MBC), octyl 4-methoxycinnamate (OMC), octocrylene (OC), 2-
33 (3-*t*-butyl-2-hydroxy-5-methylphenyl)5-chloro-benzotriazole (UV-326), and 2-(2'-hydroxy-5'-
34 octylphenyl)-benzotriazole (UV-329). Some of them (4-MBC and OMC) have been reported in
35 the literature as endocrine disruptors (Calafat et al., 2008; Fent et al., 2006; Schlumpf et al.,
36 2004). Human exposure to benzophenone derivatives has been associated with estrogen-
37 dependent diseases such as women endometriosis (Kunisue et al., 2012).

38 UVFs can enter the water systems either directly through recreational activities (swimming and
39 bathing), runoff from rainfall in contact with construction sites or garden furniture, but also
40 indirectly, by the supply of untreated (raw effluent), insufficiently treated (primary effluent), or
41 partially treated (secondary effluent) sewage waters. Consequently, they have been detected in
42 all types of water bodies: surface waters (Giokas et al., 2004; Kawaguchi et al., 2006), seawater
43 (Kawaguchi et al., 2006), wastewater (Giokas et al., 2004; Li et al., 2007), and groundwater
44 (Jurado et al., 2014). Furthermore, UVFs have also been found sorbed in a number of
45 environmental matrices: sewage sludge (Gago-Ferrero et al., 2011b; Nieto et al., 2009),
46 sediments (Baron et al., 2013; Gago-Ferrero et al., 2011a), and biota – fish and invertebrates
47 (Balmer et al., 2005; Fent et al., 2010; Gago-Ferrero et al., 2013), altogether suggesting that,
48 aside from degradation, sorption is arguably the most determinant factor governing the
49 concentration of UVFs in the water bodies, and thus their bioavailability and toxicity to
50 ecosystems and humans.

51 The extent of sorption partitioning -defined as the distribution of an organic compound between
52 the solid and the aqueous phase- of a given EOC is compound-dependent, and it is governed
53 either by its affinity for organic phases (hydrophobic partitioning) or by electrostatic and similar
54 interactions between ionized molecules and charged solid surfaces (non-hydrophobic
55 partitioning) (Franco et al., 2009; Polesel et al., 2015; Torresi et al., 2017). At equilibrium,
56 sorption and desorption rates are equal. Then, the ratio of concentrations of sorbed C_s^* -mass
57 per unit of mass of solids [MM^{-1}]- and dissolved C_w^* -mass per volume of water [ML^{-3}]- species is
58 characterized by the solid-liquid partitioning coefficient, $K_d = C_s^*/C_w^*$. K_d [L^3M^{-1}] is a lumped
59 sorption coefficient, being the sum of different species-specific partitioning coefficients (Franco
60 et al., 2009) use the term “apparent coefficient”).

61 As a general rule, neutral organic compounds have a quite limited water solubility, mainly
62 because water molecules change their overall H-bonding to their surroundings when forced to
63 interact with nonpolar compounds. On the other hand, organic matter minimizes the
64 hydrophobic surface area, because it exists in large part as organic chains coiled into globular
65 units, much like globular proteins, and occurs in somewhat isolated patches coating mineral
66 solids. Consequently, neutral organic compounds can physically penetrate between the chains
67 and find themselves “dissolved” in the non-aqueous medium (Schwarzenbach et al., 1993). This
68 type of sorption of neutral compounds is termed “absorption”, and it is mainly related to the
69 partitioning between the organic matter and water (K_{oc} , [L^3M^{-1}], tabulated in many chemical
70 databases), and to the amount of organic matter (f_{oc} , [-]). Then, sorption partitioning is given
71 as:

$$72 \quad K_d = K_{oc} \cdot f_{oc} \quad (1)$$

73 When an organic compound includes ionized structural components ($-\text{COO}^-$, $-\text{NH}_3^+$, $-\text{SO}_3^-$,
74 ...), a variety of processes become significant for sorption: (1) the electrostatic interactions with
75 charged sites on the sorbent, and (2) exchange reactions with ligands previously bound to the

76 solid. Note that the extent of solid association of ionic compounds also varies as a function of
77 external factors like the pH of the solution, since pH governs both the presence of charges on
78 mineral surfaces and the fraction of organic compound present in ionized form (through pK_a)
79 (Franco et al., 2009; Schaffer et al., 2012). Solution ionic strength and ionic composition also
80 affect the sorption of charged organic chemicals, especially if inorganic and organic ions
81 compete for the binding sites. The mineral surface composition of the sorbent is also key; for
82 example, oxides and hydroxides - like quartz or goethite mineral surfaces - present ionic radicals
83 in their surfaces. Besides this, the age of organic matter also plays an important role in sorption
84 properties, implying a distributed reactivity and increasing the heterogeneity of the
85 environment (Kleineidam et al., 2002; Weber et al., 1992).

86 Traditionally, in natural environments, sorption of organic compounds is characterized by only
87 regarding the properties of the compounds (K_{oc}), and a somewhat static sorption environment,
88 considering only the fraction of organic carbon (f_{oc}), leading to an individual K_d value if
89 equilibrium conditions are assumed. Then, when sorption is introduced in the advection-
90 dispersion equation, the storage term becomes multiplied by a constant retardation factor (R [-
91]) given by:

$$92 \quad R = 1 + \frac{\rho}{n} K_d, \quad (2)$$

93 where ρ [ML^{-3}] denotes the bulk density and n [-] the effective porosity of the soil. Although,
94 most of the reviewed literature considers this simple model (Burke et al., 2013; Henzler et al.,
95 2014; Schaffer et al., 2015), the assumption of constant R is only realistic when the environment
96 is static, meaning the simultaneous verification of three conditions: 1) no changes in
97 hydrochemistry (constant pH and ionic strength); 2) the organic compounds being always
98 neutral; and 3) the amount and the properties of sorbent being kept constant. As an example,
99 all three conditions are met in a sterile environment (e.g., Burke et al., 2013).

100 Therefore, considering a constant retardation factor in dynamic environments could lead to
101 making mistakes in predicting the fate of organic compounds, mainly the ionic ones. Knowing
102 the actual relevance of this simplification remains a challenge in most environments, but would
103 be definitely an issue in those with active redox reactions, such as biodegradation, for several
104 reasons. First, redox reactions in groundwater can vary pH, typically in the 6 to 8 range (Brun
105 and Engesgaard, 2002); this could be relevant in sorption of ionic compounds that have pK_a 's
106 values in that same range, such as a number of benzophenones (pK_a 's values being 7.07 for
107 Benzophenone-3; 7.09 for Benzophenone-1; 6.75 for Benzophenone-2, and 7.85 for 4-
108 Benzophenone (estimated by Chemaxon)). Second, in active redox systems, organic carbon
109 sediments could be oxidized, thus changing some sorbent properties; this includes its potential
110 complete disappearance by dissolution or mineralization, or the reappearance of other organic
111 sediment surfaces with different sorption properties. Last, in redox active zones, there is a
112 growth of microorganisms due to the oxidation of organic matter, implying the production of
113 solid biological compounds, like extracellular polymeric substances (EPS), and the formation of
114 biofilms. Both microorganisms and biological material can act as sorbents of organic
115 compounds. Although in wastewater treatment, sorption of organic compounds in active sludge
116 has been observed (Torresi et al., 2017), this assumption has not yet been addressed in the
117 groundwater literature.

118 All the three mentioned issues are quite relevant in Managed Aquifer Recharge (MAR) facilities,
119 hyporheic zones and bioremediation applications. In all cases, the biological processes are very
120 active and many redox reactions are occurring simultaneously (Greskowiak et al., 2006). This
121 phenomenon is enhanced in recharge applications, e.g., by the presence of a reactive layer of
122 organic matter (Beganskas et al., 2018; Grau-Martínez et al., 2017; Valhondo et al., 2014), that
123 could lead to biomass growth and bioclogging (Barba et al., 2019; Massmann et al., 2006; Perujo
124 et al., 2018). Besides this, recharge processes are a relevant source of EOCs (here including UVFs)
125 into aquifer systems (Kim et al., 2017; Laws et al., 2011; Maeng et al., 2011; Park and Lee, 2018).

126 As stated above, the chemical properties (pK_a , K_{oc}) of the different UVFs govern the individual
127 expected sorption behavior. This work aims at improving the understanding of sorption
128 processes occurring in the aquifer of a set of UVFs, with emphasis in the specific role of biomass
129 growth, organic sediment, redox conditions and pH. In section 2 we present a conceptual model
130 to evaluate the fate of selected UV filters, mainly focused on sorption processes but also on
131 degradation, plus the development of the corresponding numerical model. The models were
132 validated with data published in the literature, consisting in batch experiments evaluating the
133 fate of a cocktail of UVFs (both ionic and neutral) under different redox conditions. Section 3
134 provides the results of the modeling process followed by a discussion on the role and
135 significance of the different processes in the fate of UV filters. This is followed by a
136 concluding section.

137 **2 Model development**

138 **2.1 Experimental data description and conceptualization of the fate of UV filters**

139 To frame the conceptual and the numerical models, we start by introducing the experimental
140 data from Liu et al. 2013, involving a set of batch experiments targeting on the biodegradation
141 of five UVFs in aquifer materials. Both the groundwater and the sediments (mainly composed
142 by carbonates), were obtained from an aquifer at Bolivar (South Australia). The physicochemical
143 properties of the UVFs are shown in Table 1. The experiments were performed under different
144 redox conditions (aerobic and anaerobic-nitrate, sulfate and iron reducing conditions), all in
145 triplicate, plus sterile controls. Each batch test included 5 g of aquifer material and 5 mL of
146 groundwater, and spiked with 50 μ L of UVF mixture stock solution, prepared in methanol. The
147 experiments were incubated at 20°C and sampled at days 0, 7, 14, 21, 28, 35, 49, 63, and 77.
148 The sterile control was autoclaved and the metabolic activity was inhibited by adding sodium
149 azide (NaN_3). In the aerobic experiment, this condition was maintained by opening the caps
150 three times a day inside a laminar flow chamber. For the anaerobic ones, all the preparations

151 were carried out under an $N_{2(g)}$ atmosphere. The anaerobic microcosms were constructed using
152 Hungate tubes, flushed with nitrogen gas and sterilized, and then either un-amended with the
153 electron acceptors as the anaerobic control sulfate in the case of sulfate-reducing conditions.
154 Liu et al. (2013) analyzed both the aqueous and the solid concentrations of the five UVFs; from
155 those results and considering that sorption was in equilibrium, we calculated the K_d values.
156 Different authors concluded that considering sorption in equilibrium compared to degradation
157 is quite plausible, especially when dealing with recalcitrant organic compounds, like UVFs (Barret
158 et al., 2011; Joss et al., 2006; Rodriguez-Escales et al., 2013).

159 **2.2 Model Conceptualization**

160 Any conceptual model on the fate of UVFs should include all the relevant processes that
161 influence either degradation or sorption, sketched in Figure 1. Degradation was conceptualized
162 as co-metabolic (process 4 in Figure 1) occurring due to the oxidation of labile and dissolved
163 organic carbon (DOC) (process 3) by the relevant electron acceptor (oxygen, nitrate, or sulfate),
164 and mediated by microorganisms (biomass). Consequently, besides the production of inorganic
165 carbon (process 3.2), there is a growth of heterotrophic biomass (process 3.1), eventually
166 decaying and being oxidized to inorganic carbon. We hypothesized that biomass also acted as a
167 sorbent of UVF. At this point, biomass was treated as a pool of organic carbon from the oxidation
168 of DOC, thus, we did not distinguish where the sorption occurred (e.g. surface of wall cells, EPS's
169 of biofilm). In the experiments, the labile organic carbon was added externally, but it could be
170 also produced by the hydrolysis of sedimentary organic matter (SOM) (process 2).

171 Several factors, such as pH, temperature and oxidant concentrations, affect the reactivity of
172 SOM towards oxidants and environmental conditions (Hartog et al., 2004; van Bergen et al.,
173 1998). As the experiments were amended with high concentration of oxidants, it is also
174 reasonable to believe that SOM reacted to become labile organic carbon. At this point, we
175 introduced the possibility that SOM was not degraded directly into dissolved organic carbon and

176 we tested a second scenario considering the creation of a secondary SOM (SOM_{sec}) (see process
177 2.1 in Figure 1). Thus, we proposed two possibilities of conceptual model: 1) with two organic
178 surfaces as sorbents: SOM and biomass material (surface of microbes or EPS); and 2) with three
179 surfaces as sorbents (SOM, SOM_{sec} and biomass material).

180 **2.3 Numerical model**

181 The conceptual model was turned into a set of equations to reproduce all the processes
182 mentioned, then described in the following subsections. Similarly as Rodríguez-Escales and
183 Sanchez-Vila (2016), the kinetic reactions postulated were introduced into the geochemical code
184 PHREEQC (Parkhurst and Appelo, 1999), thus coupling the processes affecting the fate of UVFs
185 with the more common hydrochemical reactions. PHREEQC calculates the equilibrium chemistry
186 of aqueous solutions interacting with minerals, gases, solid solutions, exchangers, and sorption
187 surfaces by applying the law of mass action complemented by the thermodynamic parameters
188 present in a geochemical database that can be modified or extended by the modeler.

189 **2.3.1 Extension of the geochemical database: incorporation of the ionic UV-filters into the** 190 **geochemical speciation and interaction with pH**

191 We incorporated into the PHREEQC geochemical database the speciation of three ionic UVFs:
192 BP-3, UV-326 and UV-329. As complex organic molecules, the acidity and the basicity depend on
193 the functional groups that have either proton donor or proton acceptor properties, leading
194 consequently to different pK_a values. In Table 2, we list the reactions involving UVFs that were
195 incorporated into the database. In this case, the three compounds could behave as organic acids
196 due to the hydroxyl group bounded to aromatic rings, with a pK_a of 7.07 for BP-3, 10.08 for UV-
197 326 and 9.3 for UV-329. The logK introduced in the database was taken as equivalent to the
198 pK_a . Note that, in the case of BP-3, the potential formation of the ionic compound is relevant at
199 pH values typical of natural aquifers (on the range 6-8). In fact, the fraction of the neutral
200 compound (α_{UV_0}) is a function of pH (e.g., Schwarzenbach et al., 1993), by:

201
$$\alpha_{UV_0} = \frac{1}{1+K_a/[H^+]}, \quad (3a)$$

202 while, similarly, the fraction of the ionic compound (α_{UV_a}) is given by

203
$$\alpha_{UV_a} = \frac{K_a}{[H^+]+K_a}. \quad (3b)$$

204 Both the pK_a and the different ionic molecular forms were extracted from Chemicalize,
 205 <https://chemicalize.com/> developed by ChemAxon (<http://www.chemaxon.com>). UV-326 and
 206 UV-329 have also a basic pK_a equal to -0.96 and 0.84, respectively, in the nitrogen atoms of the
 207 benzotriazole rings. We did not include them, as significant formation of the ionic form is not
 208 expected for the pH values typical of natural aquifers.

209 **2.3.2 Sorption model between the active surfaces (SOM and biomass) and UV-filters**

210 Sorption and desorption were modeled as surface complexations reactions between the two
 211 active surfaces (SOM and biomass) and the different UVFs (both ionic and neutral forms). We
 212 assumed that sorption was always in equilibrium, thus characterized by K_d . In Table 3, there are
 213 the sorption reactions we defined in the PHREEQC database with the corresponding values of
 214 $\log K$ (from K_{oc} , see equation 1). As Liu et al. (2013) performed control and abiotic experiments,
 215 considering only sorption (no degradation), we determined the corresponding experimental K_{oc}
 216 and K_d values taking into account the amount of organic carbon (f_{oc} equal to 0.004). As there
 217 were different sorbents and some of the UV filters were ionizable, the total K_d , denoted
 218 $K_{d_{UV,TOT}}$, is defined as the sum of the K_d values in the different media and in the different forms
 219 (neutral and ionic):

220
$$K_{d_{UV,TOT}} = \sum_{j=1}^2 \sum_{i=1}^3 K_{d_{ij}} \quad (4)$$

221 where “i” is referred to the different sorbents (1=SOM, 2= SOM_{sec} and 3=biomass), and “j” to
 222 the different form of the UVFs (1=neutral and 2=ionic form).

223 SOM can sorb both to the neutral and the ionic forms of organic compounds (Schwarzenbach et
 224 al., 1993). Since the concept of SOM is very large and global – it covers multiples of molecules
 225 (cellulose, lignin, fulvic acids or humic acids) - there are several sorption mechanisms involved:
 226 H-bonding, ionic interactions, van der Waals forces, and absorption. If an ionic form of a UV
 227 molecule is present, Schwarzenbach et al. (1993) defines $K_{d_{UV,SOM}}$ as:

$$228 \quad K_{d_{UV,SOM}} = \frac{K_{oc,UV_0} + K_{oc,UV_a} K_a / [H^+]}{1 + K_a / [H^+]} f_{oc} \quad (5)$$

229 where K_{oc,UV_0} and K_{oc,UV_a} are referred to the organic carbon partition of the UVF at neutral and
 230 at ionic forms, respectively, which in this work were considered equal for simplicity and lack of
 231 additional data. This expression is equivalent to the K_{oc} definition of Franco et al. (2009) and
 232 Schaffer et al. (2012). The neutral form is relevant for low pH values, equivalent to $K_a / [H^+] \ll$
 233 1. In this case, expression (5) can be approximated as:

$$234 \quad K_{d_{UV,SOM}} = K_{oc,UV_0} f_{oc} \quad (6)$$

235 Contrarily, the ionic form is only quantitatively relevant when the pH is close to pK_a , i.e.,
 236 $1 \ll K_a / [H^+] \ll 1000$. If this is not occurring, expression (5) can be approximated as:

$$237 \quad K_{d_{UV,SOM}} \approx \frac{K_{oc,UV_0} f_{oc}}{1 + \frac{K_a}{[H^+]}} \quad (7)$$

238 This is the case of UV-326 and UV-329 with pK_a 's of 10.08 and 9.3, respectively. But, using only
 239 this statement for simplifying (5) to (7) was not valid for BP3, since the pK_a is 7.07, quite close
 240 to the working pH's. Nevertheless, we also simplified the expression. Our reason, also followed
 241 by Schwarzenbach et al. (1993), was that the formation of the conjugate acid of BP-3 was in the
 242 phenolic ring, forming phenolate. As $K_{oc,BP3_a}$ is completely unknown, we looked at the K_{ow} of
 243 the BP-3 fragments following Lyman et al. (1992). The K_{ow} of the phenolic ring was more than
 244 three orders of magnitude higher than that of the phenolate (Lyman et al., 1992; Schwarzenbach
 245 et al., 1993), indicating that phenolate was much more soluble with less interaction to the solid

246 surfaces. Consequently, we hypothesized than $K_{oc, BP3a} \ll K_{oc, BP3o}$, so that the simplification of
247 (7) could be used.

248 The sorption of organic compounds into biomass (X) has been only related to ionic compounds
249 (Flemming, 1995; Franco et al., 2009; Torresi et al., 2017). In porous media, biomass is organized
250 in biofilms, which contains living organisms and also other biological materials such as EPS. In
251 biofilms, sorption can act upon different regions (Flemming, 1995): i) EPS with cationic groups
252 in amino sugars and proteins, and anionic groups in uronic acids and proteins; ii) outer
253 membrane and lipopolysaccharides of gram-negative cells with their lipid membranes, the
254 lipoteichoic acids in gram positive cells; iii) cell walls consisting of N-acetylglucosamine and N-
255 acetylmuramic acid, offering more cationic and anionic sites. Although cation sorption in
256 biofilms has been widely reported in the literature (Franco et al., 2009; Torresi et al., 2017), e.g.,
257 a wide variety of metal ions bound to EPS (Flemming et al., 1996), anionic sorption is also
258 physically possible, due to the positive charges in amino groups in sugars, sugar acids and
259 proteins. The literature on sorption of organic compounds in biofilms is poor and contradictory;
260 for example, Torresi et al. (2017) only observed sorption into biofilm of 9 of 23 compounds,
261 being the cationic ones. On the other hand, Späth et al. (1998) observed that BTEX sorbed to
262 biomass, mainly to EPS. This was surprising, even for the authors, because EPS are mainly formed
263 by water (Brangarí et al., 2018) and sorption should be preferentially for polar compounds. As
264 literature is not conclusive, we followed the hypothesis that sorption into biomass only occurred
265 with polar compounds, thus, with ionic forms of UVs. Then, $K_{dUV,X}$ was defined as follows:

$$266 \quad K_{dUV,X} = \frac{K_{X,UVa} K_a / [H^+]}{1 + K_a / [H^+]} f_{sites} \quad (7)$$

267 where f_{sites} is related to the available sites for sorption of biomass and it is conceptually
268 equivalent to the f_{oc} defined in (5); $K_{X,UVa}$ is the partitioning coefficient of ionic UVFs into

269 biomass ($[\text{UV}_x^-]_{\text{eq}} / [\text{UV}_w]_{\text{eq}}$). This parameter is directly related to the logK values displayed in
 270 Table 2, fitted from experimental data.

271 **2.3.3 Organic matter and co-metabolic degradation model.**

272 UVFs degradation was driven by co-metabolism, with the oxidation of labile organic matter as
 273 the main process, and, consequently, the broken up of the UVF molecules. The oxidation was
 274 conducted by electron acceptors (oxygen and sulfate) in a separate way, and coupled with the
 275 biomass growth/decay and with the degradation of UVFs. The degradation of the labile dissolved
 276 organic carbon and the growth of biomass were modeled using equations (8) to (10), which were
 277 set up in previous works (Rodríguez-Escales et al., 2016; Rodríguez-Escales et al., 2014):

$$278 \quad r_{ED} = -k_{\max} \frac{[ED]}{[ED] + K_{S,ED}} \frac{[EA]}{[EA] + K_{S,EA}} [X] \quad (8)$$

$$279 \quad r_{EA} = Q r_{ED} - S b [X] \quad (9)$$

$$280 \quad r_X = -Y_h r_{ED} - b [X] \quad (10)$$

281 where [ED] is the concentration of the electron donor (organic carbon, methanol); [EA] that of
 282 the electron acceptor (oxygen, nitrate or sulfate depending on the experiment), and [X] the
 283 biomass concentration (aerobic or sulfate-reducing, respectively), all in $[\text{ML}^{-3}]$; $k_{\max} [\text{T}^{-1}]$ is the
 284 consumption rate of electron donor per unit value of biomass; $K_{S,ED}$ and $K_{S,EA} [\text{ML}^{-3}]$ the half
 285 saturation constants of electron donor and acceptor, respectively; $b [\text{T}^{-1}]$ a biomass decay
 286 constant; Y_h the microbial yield [C biomass / C organic matter], and Q [EA / ED] and S [EA / C
 287 endogenous] are stoichiometric coefficients. The values of Q and S are 2.92 and 0.4 for aerobic
 288 conditions and 3.62 and 0.1 for sulfate-reducing ones. We did not consider an inhibition process
 289 due to the co-existence of different electron acceptors (typical in some Monod kinetics models,
 290 (Rodríguez-Escales et al., 2017)), since the different experiments involved only one electron
 291 acceptor each. The calibration process is described in Section 2.3.4.

292 Note that in the experiments of Liu et al. (2013), as it is also the case in most aquifers, there was
 293 a little amount of sedimentary organic carbon. The literature reports that, if a certain activity of
 294 oxidants is occurring, SOM can be hydrolyzed, releasing dissolved organic carbon into the system
 295 (Hartog et al., 2004). As the experiments were amended with a high quantity of oxidants, we
 296 incorporated the release of DOC from SOM as:

$$297 \quad r_{DOC} = -k_{\text{prim}}[SOM] \quad (11)$$

298 Note that (11) has to be added to (8) to have the overall rate of DOC. As pointed out in the
 299 conceptual model, we also tested the possibility of the creation of a secondary SOM from the
 300 primary one, understood as an intermediate step until the formation of dissolved organic
 301 carbon. Then, the rate of SOM_{sec} was conceptualized as:

$$302 \quad r_{SOM_{\text{sec}}} = k_{\text{prim}}[SOM] - k_{\text{sec}}[SOM_{\text{sec}}] \quad (12)$$

303 On the other hand, the degradation rates of UVFs were assumed to depend on redox conditions
 304 as described by Liu et al. (2013), and simulated using the following first-order kinetic reaction:

$$305 \quad r_j = -C_j \sum_{i=1}^n k_{j,i} F_i \quad (13)$$

306 where j represents the actual UVF, i represents the redox condition, up to the n ones where
 307 UVFs degradation were studied, and $k_{j,i} [T^{-1}]$ the first-order degradation constant for each
 308 species j at a specific redox state i . Since different redox conditions occurred simultaneously in
 309 the aquifer material during the simulations, we incorporated in (13) the activation factor $F_i [-]$,
 310 which activated the degradation rate of UVF when oxidation of organic carbon occurred; it is
 311 defined as:

$$312 \quad F_i = \frac{p}{p+q_i} \quad (14)$$

313 where $p [ML^3T^{-1}]$ is the rate of organic carbon degradation and $q_i [ML^3T^{-1}]$ a constant which was
 314 related to the rate at degradation of UVF was activated. If $p \geq q_i$ degradation occurred without

315 any problem, but if $p < q_i$ degradation was limited. The $k_{j,i}$ values were calibrated by fitting
316 the model results with the experimental observations of UVFs in both aqueous and solid phases.

317 **2.3.4 Initial conditions of the model and calibration process**

318 Table 4 shows the initial concentration of the compounds considered in the biogeochemical
319 model. The calcium concentration of Table 4 was calculated assuming equilibrium with calcite,
320 acknowledging the sediment was composed of 10% calcite (Liu et al. 2013). On the other hand,
321 as the solution in Liu et al. (2013) was not chemically balanced, we added enough sodium into
322 the model to keep the electrostatic balance equal to 0. The organic carbon source was the
323 methanol (2.47×10^{-1} M) added with the UVF mixture stock solution. For the sulfate-reducing
324 condition, lactate (1×10^{-2} M) was considered the electron donor. The initial concentrations of
325 UVFs were approximately 1 $\mu\text{g/g}$ aquifer. The aerobic conditions were simulated by assuming
326 equilibrium with the atmospheric oxygen. The sulfate-reducing conditions were kept by adding
327 20mM of NaSO_4 and Na_2S as a reducing agent.

328 Regarding model parameters, the biogeochemical kinetic ones (k_{max} , $K_{\text{S,ED}}$, $K_{\text{S,EA}}$, b) for the
329 aerobic oxidation model and for the sulfate reduction model; the kinetic parameters of co-
330 metabolic degradation of UVF ($k_{i,j}$ and F_i); and the sorption constant for ionic UVFs (K_{X,UV_a})
331 were first calibrated by hand and then refined automatically using code PEST (Doherty, 2005).
332 PEST allowed computing the sensitivities, correlations, and confidence intervals for the
333 optimized model parameters using the Levenberg-Marquardt algorithm. The weights of each
334 chemical species associated to the measurement errors were applied using the inverse of the
335 standard deviation of the confidence interval of measurements (95%). For calibration, we used
336 the experimental information from Liu et al. 2013 (aerobic biomass, sulfate reducing biomass,
337 sulfate, mass of UVFs in the aqueous and in the solid phases, and Kd 's of ionic UVFs). Finally,
338 we evaluated the likelihood of the two conceptual models proposed regarding the formation of
339 a secondary SOM comparing to different Information Criteria values calculated by PEST.

340 **3 Results and discussion**

341 **3.1 Coupling degradation model of UVs with sorption**

342 Figure 2 shows the results of the co-metabolic degradation model of UVFs coupled to sorption
343 processes considering both SOM and biomass as active surfaces. In the Figure, the two tested
344 scenarios of conceptual models are shown: the first one, considering only one step between the
345 SOM hydrolysis and the DOC formation, and the second one, considering the formation of a
346 secondary SOM before the DOC formation (see Figure 1). Besides this, we also added to the plot
347 the results of the K_d without considering the role of the biomass as a sorbent. The model was
348 developed for aerobic conditions (subplots a-e), and for sulfate-reducing ones (subplots a'-e').
349 Table 5 shows the constant parameters for all conditions and for different UVs.

350 Note that whereas the first conceptual model of SOM (dashed lines) fitted quite acceptably the
351 experimental data of UVFs mass as a function of time in both dissolved (blue) and solid (red)
352 phases, the corresponding fittings of K_d values were not good enough (solid black lines of side
353 plots), especially for OC, UV-326 and UV-329. This indicated that the model did not consider all
354 the processes involved and the formation of a secondary SOM was a reasonable hypothesis;
355 actually, at around days 15-20 a peak in sorption was observed (large values of lumped K_d),
356 followed by a decay with time. As this transition was not coincident with the time evolution of
357 biomass and SOM (see Figure 3), it could indicate that a transitional appearance of a secondary
358 sorbing surface. The reasoning behind this assumption is that the degradation of SOM to DOC
359 does not take place in just one single metabolic step, but rather a complex degradation pathway
360 is always associated to the presence of recalcitrant organic matter (Hartog et al., 2004). The
361 addition of this secondary sorption process helped improving the fitting of the experimental
362 lumped K_d value (solid line in Figure 2). The plausibility of this hypothesis was verified using
363 several information criteria (Akaike's, AIC, corrected Akaike's, AICc, Bayesian, BIC, and Kashap's,
364 KIC). Regardless the criteria used, plausibility was larger for the second scenario (formation of a

365 secondary surface) rather than for the first one (aerobic and sulfate-reducing conditions), see
366 Table 1 of Supporting information.

367 The importance of biomass as a sorbent in ionic compounds is clearly shown in Figure 2 (grey
368 line in the insert plot), most particularly for BP-3, in agreement with the literature on the
369 sorption of BP-3 onto biological matrices (fish and crustaceans) (Mao et al., 2019). In the case of
370 the neutral compounds (OC and OMC), the best fit was obtained by assuming no sorption to the
371 biomass (results not shown). Both the experimental results and the model parameters show
372 that the most degraded UVF was OMC followed in descending order by UV-329, UV-326, BP-3,
373 and OC. The constants of activation factors for aerobic and anaerobic conditions (Equation 14)
374 were 1×10^{-4} mol/Ls for BP-3, OC and OMC, and 2×10^{-3} mol/Ls for UV-326 and UV-329. It is noted
375 that the results and the constant parameters did not vary considerable from different redox
376 conditions, as already reported by Liu et al. (2013). The extreme values were BP-3, showing the
377 largest amount of degradation in sulfate-reducing conditions, and OC, with the fastest overall
378 degradation observed for aerobic conditions. This could indicate that degradation pathways
379 were governed by the redox state. For BP-3, the first degradation step is the o-demethylation to
380 produce 2,4-hydroxybenzophenone (BP-1) in both aerobic and anaerobic conditions (Liu et al.,
381 2012). Nevertheless, under anoxic conditions, the relatively long half-life of BP-3 in nitrate
382 reducing conditions could indicate that these processes were inhibited by nitrate (Liu et al.,
383 2012; Milligan and Häggblom, 1999), justifying that redox conditions could affect the
384 disappearance rate of BP-3.

385 Interestingly, biomass sorption constants did not differ more than one order of magnitude
386 amongst the different redox conditions (see Table 3). Note that the fitted K_d values were
387 corrected considering the amount of biomass, quite different in the two conditions (see Figure
388 3). Note that K_x of BP-3 was comparable to $K_{d,SOM}$ (equal to 1.35, considering f_{oc} of 0.004), but
389 its relevance was lower in the cases of UV-326 and UV-329 ($K_{d,SOM}$ equal to 8.35 and 9.62,

390 respectively). We want to highlight that, to our knowledge, this is the first time these parameters
391 were determined, so it was not possible to compare them to the literature. In the case of K_{oc}
392 for BP-3, our results (339 L/kg, Table 3) were in the same order of magnitude of those reported
393 by Cao et al. (2014) for natural organic matter formed mainly by humic acids.

394 Finally, we remark that the values of K_{oc} 's fitted from the abiotic experiments of Liu et al. (2013)
395 (reported in Table 3) differ considerably from the theoretical values, determined by EPI-SUITE
396 (Table 1). On the other hand, empirical approximations considering logP provide values of K_{oc}
397 close to the theoretical ones. The best expression for BP-3 was the empirical expression of Sabljíć
398 et al. (1995): $K_{oc} = 10^{0.61\log P + 0.32}$, predicting a log value of 2.49 (the experimental one was
399 2.53). On the other hand, in the cases of UV-326 and UV-329, the best fit was obtained using the

400 expression of Franco et al. (2009): $K_{oc} = \frac{10^{0.54\log P_n + 0.11}}{1 + 10^{(\text{pH}_{\text{soil}} - 0.6 - \text{p}K_a)}} + \frac{10^{0.11\log P_n + 1.54}}{1 + 10^{(\text{p}K_a - \text{pH}_{\text{soil}} + 0.6)}}$, which
401 predicted $\log K_{oc}$ values of 2.98 and 3.32 (the experimental ones were 3.32 and 3.38,
402 respectively).

403 **3.2 Evaluating the effect of pH in sorption of ionic UV-filters**

404 Besides biomass, pH plays a significant role in sorption processes of ionic compounds (Schaffer
405 et al., 2012). Subsurface water pH is commonly governed by calcite equilibrium. Then, we tested
406 the effect of different scenarios of calcite equilibrium in the K_d values of UVFs (Figure 4). We
407 evaluated scenarios of calcite sub-saturation (saturation indexes, SI, from -2 to -1), and scenarios
408 of over-saturation (SI from +2 to +1). These scenarios are realistic in aquifers, representative of
409 different mineralogy. For example, SI = -2 was observed in granitic geologies, whereas SI = +2
410 can represent carbonate geologies (Folch et al., 2011). From the modeling effort, the UVF most
411 affected by these changes was BP-3, especially in aerobic conditions, that we associated to: 1)
412 $\text{p}K_a$ of BP-3 (7.07) was quite close to working pH values (plots d, and d' in Figure 4), and 2)
413 aerobic conditions implied a wider range of pHs (that itself controls the fate of BP-3) compared
414 to sulfate reducing conditions (Barry et al., 2002). Note that K_d spans two orders of magnitude

415 between the two extreme cases, with direct consequences in BP-3 mobility and, consequently,
416 in risk assessment. Besides, in a calcareous hyporheic zone or a recharge facility, the mobility of
417 BP-3 will be conditioned to carbonate equilibrium. It would be expected than in the first
418 centimeters, generally in aerobic conditions, the mobility of BP-3 will be larger than in depth
419 (and reduced conditions). For the two other ionic compounds, UV-326 and UV-329, the effects
420 of changes in calcite saturation was not that important as compared to BP-3.

421 **3.3 The role and significance of the different processes in the fate of UV-filters**

422 Figure 5 presents the results of a simulation for the relative importance of the different pools
423 on the fate of UVFs: liquid phase, organic carbon solid phase (SOM and SOM_{sec}), biomass, and
424 degradation as a function of time, for two different redox conditions and assuming a two-step
425 conceptual SOM.

426 Regarding the fate of ionic UVFs, BP-3 was mostly degraded in less than 20 days, with a
427 remaining fraction in biomass pool, more significant under aerobic conditions, probably due to
428 large production of aerobic biomass. The fate of UV-326 and UV-329 did not differ much
429 between the two conditions (similar sorption parameters, see Tables 1 and 5); these compounds
430 had more affinity to SOM than BP-3, explained by the values of $\log K_{oc}$ (3.32 for UV-326, 3.38
431 for UV-329, and 2.53 for BP-3). For UV-326 and UV-329, SOM_{sec} was the main pool in aerobic
432 conditions, indicating that these two compounds have more affinity to SOM than to biomass.
433 This is quite logical since their pK_a 's are 9.30 and 10.08 and pH oscillated from 7.9 to 5.3 in the
434 aerobic experiment and 7.9 to 6.4 in the anaerobic ones, indicating that the formation of ionic
435 compound was not quantitatively important.

436 Regarding the fate of neutral compounds (OC and OMC), they were fast degraded, especially
437 OMC with a half-life time of 5 days. In the case of OC, SOM was a more important pool compared
438 to OMC, again responding to the K_{oc} values, higher for OC (3.55) than for OMC (3.33).

439 Our results show that UVFs fate will depend on the environmental conditions, especially on the
440 activity of biogeochemical processes. If such processes are not occurring, e.g., in a groundwater

441 poor of labile organic carbon, co-metabolic degradation would not occur, and the fate of UVFs
442 will be conditioned by sorption and the presence of SOM. As typically the SOM content in
443 groundwater is low, most of the compounds will be found in dissolved (water) phase. On the
444 other hand, in active biogeochemical environments (e.g., in managed aquifer recharge surface
445 facilities), UVFs would be degraded because other biogeochemical processes are occurring and
446 also they would be sorbed into biomass. In these cases, biofilms will grow in warm seasons from
447 carbon fixation (Barba, 2018; Massmann et al., 2006); moreover, if the origin of the recharge
448 water is treated wastewater, it might contain a significant amount of UVFs (Biel-Maeso et al.,
449 2019). Thus, in these cases, biomass sorption and co-metabolic degradation should be
450 incorporated into the conceptual model in order to have a good picture of the processes
451 governing the fate of UVFs.

452 **4 Conclusions**

453 In this work, we explored the relevance of biomass to act as an important sorbent for ionic UVFs
454 (BP-3, UV-326 and UV-329) in porous media. To our knowledge, this is the first work indicating
455 that biomass can be an important sorbent capable of retaining UVFs in the subsurface.
456 Furthermore, our work improves the understanding of the UVs fate in the subsurface.

457 We have developed a geochemical model for UVF sorption, that considers co-metabolic
458 degradation and sorption to two different pools of sedimentary organic carbon, being the
459 sediment and the biomass formed during the process. This coupling was based in an extension
460 of a geochemical database, incorporating the speciation and sorption reactions of UVFs. In this
461 way, this work opens the door to use the capabilities for allowing editable and customizable
462 geochemical databases to incorporate a great extension of organic pollutants to allow evaluating
463 their fate in relation to all the hydrogeochemical reactions occurring in an aquifer.

464 We further conclude that redox conditions, by themselves, do not condition the fate of UV filters
465 (similar degradation constants). Nevertheless, redox conditions affect the biomass growth,

466 showing a much higher growth of aerobic biomass than anaerobic one, and also modifying the
467 pH. Thus, redox conditions modify the sorption properties of the media, which condition the
468 fate of the studied UVFs.

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665

666

667 **Table captions**

668 **Table 1.** Physicochemical properties of the UV-filters included in this study. (1) Estimated
669 by using EPIWEB (PCKOCWIN v1.66); (2) Estimated by ChemAxon; (*) Neutral
670 compound.

671 **Table 2.** Set of reactions added to the database concerning the formation of the ionic
672 compounds of UV-filters.

673 **Table 3.** Set of reactions added to the database in order to consider the sorption of UV-
674 filters into SOM and biomass (X). (1) $\log K$ was related to K_{oc} , determined by
675 $\log K_{d,PHREEQC} = \log K_{oc} - \log m_{org} = \log K_{oc} - \log(10^{100})$. The $\log K_{oc}$ values were determined from the
676 control experiments without degradation of Liu et al. (2013)), and not the theoretical
677 values (Table 1).

678 **Table 4.** Initial concentrations considered in the biogeochemical model. In the
679 sulfate-reducing scenario Sulfate concentration was 20 mM.

680 **Table 5** Model parameters for each redox condition. (1) It was determined
681 considering that the whole surface of biomass was suitable to sorb. (2) Sulfate
682 concentrations were as high that sulfate Monod term (Eq 8) was equivalent
683 to 1.

684

685

686 **Figure captions**

687 **Figure 1.** Processes affecting the fate of UV filters in aquifers and soils. DOC Dissolved
688 Organic Carbon; sed – sediment; lab – labile; SOM- Sedimentary Organic Matter; BM –
689 biomass; EA - Electron Acceptor; TP – Transformation Product.

690 **Figure 2.** Results of the model in terms of mass of UV filters as a function of time in both
691 dissolved (blue) and solid (red) phases. Left: aerobic conditions, Right: sulfate-reduction
692 conditions). Dashed lines correspond to first scenario model and the solid ones to the
693 second one, where a transitional appearance of a secondary surface for sorption, being
694 formed from SOM. Green solid line in the insert plot, showing the evolution of lumped
695 K_d as a function of time, was obtained using a model without considering the biomass
696 sorption.

697 **Figure 3.** Evolution of the different sorbents in the experiments. The circles in a) and c)
698 represent the experimental data of aerobic and sulfate reducer biomass, respectively.
699 b) and d) evolution of SOM; solid lines correspond to the one-step model; red dashed
700 lines model including a secondary sorption surface.

701 **Figure 4.** Effect in K_d of ionic UV-filters at different saturation index –SI- values of calcite
702 at different redox conditions, aerobic ones (a-d) and sulfate-reducing ones (a'-d'). Black
703 line represents calcite equilibrium, red lines represent different scenarios of calcite
704 saturation (SI = 2, solid; SI = 1, dotted), blue lines represent scenarios of calcite sub-
705 saturation (SI = -2, solid; SI = -1, dotted).

706 **Figure 5.** Temporal evolution of the relative importance of the processes involved in the
707 fate of UV filters under aerobic (left) and sulfate-reducing (right) redox conditions.