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Uranium fate during crystallization of magnetite from ferrihydrite in conditions relevant to the disposal of radioactive waste

TIMOTHY A. MARSHALL¹, KATHERINE MORRIS¹, GARETH T.W. LAW², J. FREDERICK W. MOSSELMANS³, PIETER BOTS¹, HANNAH ROBERTS¹ AND SAMUEL SHAW^{1,*}

¹ Research Centre for Radwaste Disposal and Williamson Research Centre for Molecular Environmental Science, School of Earth, Atmospheric and Environmental Science, The University of Manchester, Manchester M13 9PL, United Kingdom

² Centre for Radiochemistry Research, School of Chemistry and Research Centre for Radwaste Disposal, School of Earth, Atmospheric and Environmental Science, The University of Manchester, Manchester M13 9PL, United Kingdom

³ Diamond Light Source Ltd, Diamond House, Harwell Science and Innovation Campus, Didcot, Oxfordshire OX11 0DE, United Kingdom

[Received 5 November 2014; Accepted 29 June 2015; Associate Editor: Nicholas Evans]

ABSTRACT

Uranium incorporation into magnetite and its behaviour during subsequent oxidation has been investigated at high pH to determine the uranium retention mechanism(s) on formation and oxidative perturbation of magnetite in systems relevant to radioactive waste disposal. Ferrihydrite was exposed to $U(VI)_{aq}$ containing cement leachates (pH 10.5–13.1) and crystallization of magnetite was induced via addition of $Fe(II)_{aq}$. A combination of XRD, chemical extraction and XAS techniques provided direct evidence that $U(VI)$ was reduced and incorporated into the magnetite structure, possibly as $U(V)$, with a significant fraction recalcitrant to oxidative remobilization. Immobilization of $U(VI)$ by reduction and incorporation into magnetite at high pH, and with significant stability upon reoxidation, has clear and important implications for limiting uranium migration in geological disposal of radioactive wastes.

KEYWORDS: uranium, magnetite, incorporation, EXAFS, radioactive waste, geodisposal.

Introduction

DUE to its long half-life and significance in the nuclear fuel cycle, uranium is the dominant radionuclide by mass in a geological disposal facility (GDF). In addition, uranium is a significant contaminant at many nuclear facilities. In terms of

environmental behaviour, under oxic conditions $U(VI)$ is dominant and is relatively soluble in aqueous solution and can form neutral and anionic carbonate complexes which make it environmentally mobile (Clark *et al.*, 1995); however, sparingly soluble uranyl silicates and uranate phases can also form in some natural/engineered environments (e.g. high-pH cementitious conditions, Wellman *et al.*, 2007). Under reducing conditions $U(IV)$ dominates, which is sparingly soluble and may precipitate as $U(IV)$ oxide phases

* E-mail: sam.shaw@manchester.ac.uk

DOI: 10.1180/minmag.2015.079.6.02



The publication of this research has been funded by the European Union's European Atomic Energy Community's (Euratom) Seventh Framework programme FP7 (2007–2013) under grant agreements n°249396, SeclGD, and n°323260, SeclGD2.

(Murphy and Shock, 1999; Lloyd and Renshaw, 2005). Additionally, formation of stable uranium colloids may also promote migration of uranium in cementitious environments (Bots *et al.*, 2014; Smith *et al.*, 2014).

Uranium adsorbs to a wide range of geological materials under oxic and reducing conditions (Payne *et al.*, 1994; Waite *et al.*, 1994; Shuibo *et al.*, 2009). In addition, both goethite (α -FeOOH) and hematite (α -Fe₂O₃) have been shown to be capable of incorporating U(VI) and U(V) into their structures, offering a pathway for long-term immobilization (Nico *et al.*, 2009; Boland *et al.*, 2011, 2014; Ilton *et al.*, 2012; Marshall *et al.*, 2014b). Magnetite (Fe₃O₄) is a mixed Fe(II)/Fe(III) oxide (Fleet, 1981; Cornell and Schwertmann, 2003) relevant in both natural and engineered environments and its interactions with uranium have been explored by several workers. Magnetite can remove U(VI) from solution through reduction to U(IV) coupled to oxidation of Fe(II), with precipitation of UO₂ on the magnetite surface (Dodge *et al.*, 2002; Missana *et al.*, 2003; Scott *et al.*, 2005; Skomurski *et al.*, 2011; Singer *et al.*, 2012a). Recent work has shown that at low uranium loadings, mononuclear U(IV) complexes form on the magnetite surface which then nucleate the precipitation of UO₂ as the surface loading increases (Latta *et al.*, 2014). Indeed, with time, it appears that adsorbed uranium may become incorporated into the magnetite structure with XPS suggesting U(VI) and U(V), and XANES suggesting U(IV) and U(V)/U(VI). EXAFS modelling was more challenging but the authors suggested that U was incorporated into the magnetite octahedral site (Huber *et al.*, 2012).

U(IV) associated with iron oxide minerals as UO₂ is easily remobilized to solution under oxidizing, and potentially even mildly reducing conditions (Senko *et al.*, 2002, 2007; Campbell *et al.*, 2011). By contrast, if U is incorporated within the mineral, it may be protected from oxidative release, as observed with goethite (Stewart *et al.*, 2009). Many geodisposal concepts utilize cementitious materials in the design of the disposal facility, e.g. an engineered barrier system. Leaching of the cement will cause hyperalkaline pH from >13 to 10 over geological time, creating a chemically disturbed zone in the host rock (Berner, 1992). In addition, radioactively contaminated land can be alkaline (Catalano *et al.*, 2004; Wallace *et al.*, 2012, 2013). In addition, magnetite will undoubtedly be important in any geological disposal facility as anaerobic corrosion of steel leads to the formation

of magnetite (Musić *et al.*, 1993). However, to date few studies have examined the structural environment of U incorporated into magnetite and the effect of oxidation of the host mineral under alkaline conditions. Understanding this is important to enable prediction of the long-term stability and environmental fate of U in geodisposal relevant situations.

Here, we examine incorporation of U into magnetite under alkaline conditions to explore the speciation of the incorporated U, the effects of changing pH, the amount of U sequestered, and the fate of U following air oxidation. Ferrihydrite has been used as the starting material and magnetite was crystallized by addition of aqueous Fe(II) (Hansel *et al.*, 2003). We have followed the partitioning of U during this process in two solutions at pH 13.1 and 10.5. After a period of ageing, the systems were then oxidized in air to test the recalcitrance of any iron-oxide-associated uranium. We have combined solution chemistry data with quantitative powder X-ray diffraction (qXRD), transmission electron microscopy (TEM), and X-ray absorption spectroscopy (XAS) to characterize the fate of uranium during magnetite crystallization and reoxidation at high pH.

Materials and methods

The Fe(II)-induced crystallization of ferrihydrite to magnetite was studied in U(VI)-amended synthetic cement leachates chosen to simulate early- (pH 13.1) and late- (pH 10.5) stage evolution of a GDF (Berner, 1992; Marshall *et al.*, 2014a; Moyce *et al.*, 2014). The young cement leachate (YCL, pH 13.1) is dominated by potassium and sodium hydroxide dissolution and the old cement leachate (OCL, pH 10.5) is dominated by calcium hydroxide. The YCL composition was: 5.2 g l⁻¹ KOH, 3.8 g l⁻¹ NaOH and 0.1 g l⁻¹ Ca(OH)₂; the OCL composition was: 15 mg l⁻¹ Ca(OH)₂. Experiments were maintained in an anaerobic atmosphere with <1 ppm O₂ and CO₂.

2-line ferrihydrite was synthesized (Cornell and Schwertmann, 2003) and the Fe(III) content was determined by dissolution in 1 M HCl and analysis (Viollier *et al.*, 2000). Ferrihydrite was equilibrated with the cement leachates at 4 g l⁻¹ for 1 hour at room temperature. A pH drop of 1–2 pH units in the OCL occurred over 2 hours and the system was manually readjusted to pH 10.5 with KOH. The pH was constant in the YCL. In the coprecipitation experiments, the ferrihydrite-equilibrated leachates were

spiked with U(VI) to 4 ppm (1.7×10^{-5} mol l⁻¹). Crystallization of magnetite was then induced by addition of Fe(II) as 0.1 M FeCl₂, to give stoichiometric magnetite [1 Fe(II) : 2 Fe(III)] and pH was immediately re-adjusted using KOH. The magnetite was aged for two days in an anaerobic CO₂-free environment and then moved to a CO₂-free air (Zero Grade Air) environment and oxidized for 14 days. Parallel unspiked experiments for full characterization of the solid products were set up and the data presented in Marshall *et al.* (2014a). Sorption ‘controls’ were also set up where U(VI) was spiked after 1 day of magnetite ageing and left for one further day before reoxidation in CO₂-free air for 14 days.

Partitioning of uranium between the solid and solution was determined by analysis of the solution U concentration by ICP-MS (Agilent 7500cx). The uranium distribution in the solid phase was determined by a dissolution time-series of the solid samples in 1 M HCl and the solution analysed for U and Fe by ICP-MS, after the method of Sidhu *et al.* (1978). Solids were characterized by X-ray Diffraction (XRD) using a Bruker D8 diffractometer. For TEM, solid samples were characterized using an FEI CM200 FEG-TEM. Uranium L_{III}-edge XAS spectra were collected at B18, Diamond Light Source, at room temperature in fluorescence mode using a 9-element Ge detector (Dent *et al.*, 2009). U(VI) and U(IV) standard spectra (schoepite [(UO₂)₈O₂(OH)₁₂·12(H₂O)] and uraninite (UO₂),

respectively) were collected in transmission mode. Data reduction and fitting of the EXAFS spectra were performed using Athena and Artemis (Ravel and Newville, 2005) with FEFF6.

Results and discussion

All uranium was removed from solution during the reduced stage in both systems (Fig. 1). After 14 days reoxidation, $6 \pm 1\%$ and $21 \pm 1\%$ of the uranium was remobilized to solution in the YCL and OCL, respectively (Fig. 1). Acid digestions on the reduced and oxidized magnetites, from both systems, were used to explore the U distribution profile in the solids (Fig. 2). In the YCL system, the gradient of 1 suggested that U was uniformly distributed through the particles and oxidation of the host mineral had no effect on the U distribution (Fig. 2). U distribution through the reduced magnetite particles in the OCL system also shows a uniform concentration throughout the particles. After 21 days oxidation in the OCL system the leaching data suggested that the distribution of U changed significantly with ~15% of the U released before any significant dissolution of Fe (<2%). This was then followed by a more uniform distribution of U within the particles. These data indicate that ~15% of the U was associated with the surface of the iron oxide particles (Fig. 2). This is consistent with the observation that after 21 days oxidation, 21% of the total U was present in solution so that ‘available’ solution-associated U could re-adsorb to the oxidized magnetite surface.

In the YCL system no pool of 1M HCl extractable adsorbed U was evident in the oxidized sample. The extent of magnetite oxidation in the YCL system after 21 days oxidation was limited, whereas in the OCL there was appreciable magnetite oxidation over the same period (Marshall *et al.*, 2014a). Different magnetite oxidation rates in the different leachates may explain the different behaviour of U; the extent of oxidation and hence oxidative release of U to solution is lower in the YCL, but additionally, any remobilized U(VI) may form colloidal U(VI) phases (Bots *et al.*, 2014) and limit re-adsorption behaviour with the magnetite.

Magnetite oxidation to maghemite proceeds via outward migration of Fe(II) (Sidhu *et al.*, 1977; Tang *et al.*, 2003; Gorski *et al.*, 2010). During the oxidation process, U is likely to be expelled from the particle core as the structure re-orders itself owing to the larger crystal radius of U versus Fe (Shannon, 1976). U(VI) and U(IV) have crystal radii (in octahedral coordination) of 0.87 Å and

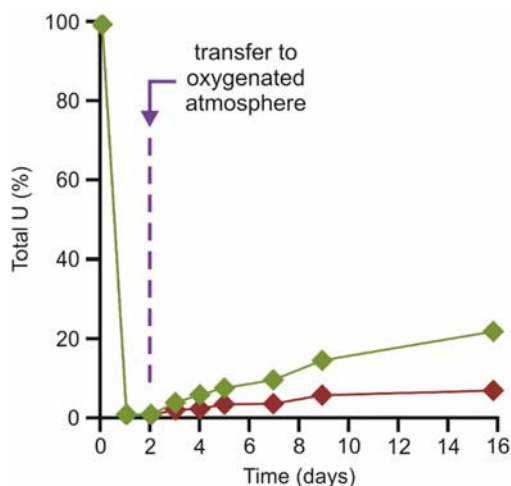


FIG. 1. Concentration of U in solution (%) during magnetite ageing (days 0–2), and subsequent air oxidation (days 2–16) in YCL (red) and OCL (green).

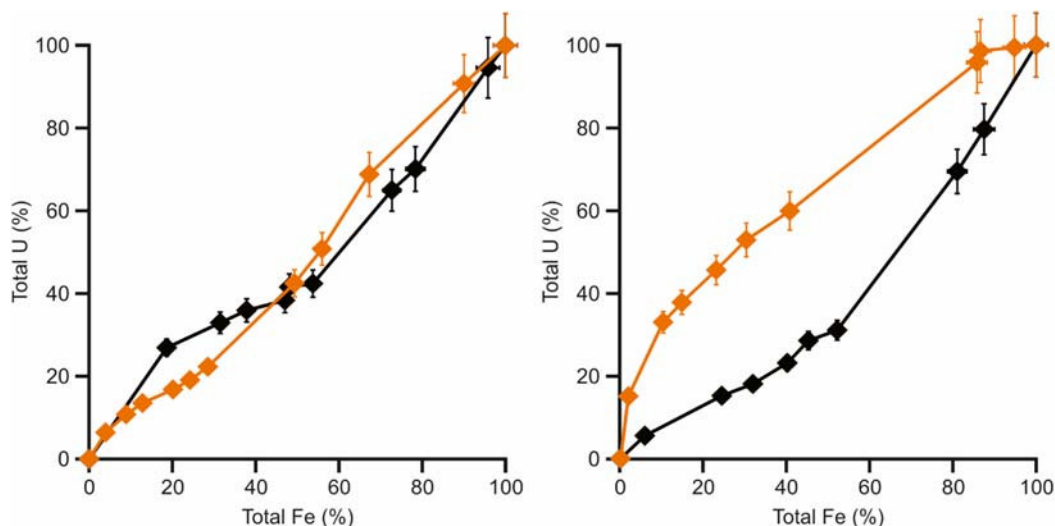


FIG. 2. Release of U and total Fe during dissolution of magnetite in 1M HCl in the YCL system (pH 13.1) (left) and OCL system (pH 10.5) (right). Black lines are reduced magnetite experiments, orange lines are oxidized magnetite experiments.

1.03 Å, respectively, whilst the crystal radii of Fe(III) and Fe(II) (in octahedral coordination, high spin) are 0.785 Å and 0.92 Å, respectively (Shannon, 1976). However, the uniform distribution of U in the oxidized samples, particularly in the OCL system, suggests that the majority of U is retained during oxidation and it is not ejected from the structure (Fig. 2). It has been suggested that the comparative ionic charge, radii and electronegativity all have an influence on the substitution behaviour of different elements within host minerals (Sidhu *et al.*, 1978). In magnetite, similar uniform distributions have been observed for Co, Ni and Zn, whereas Cu, Mn and Cd have been observed to have a dominant surface accumulation (Sidhu *et al.*, 1978). Recently, Tc was also observed to have a near surface distribution predominance (Marshall *et al.*, 2014a).

X-ray absorption spectroscopy

U L_{III} -edge XANES and EXAFS spectra were collected on reduced magnetite from the YCL and OCL coprecipitation and adsorption experiments, along with an oxidized magnetite from the OCL coprecipitation experiment. The XANES spectra are shown in Fig. 3 with the U(VI) and U(IV) reference spectra. It is clear that in all experiments U(VI) reduction occurred with an obvious shift in the energy

of the adsorption edge of all experimental spectra towards that of the U(IV) standard (Fig. 3). Interestingly, the edge position of the three coprecipitation experiments' spectra [(d)–(f)] were ~ 1 eV higher than the U(IV) standard and the parallel adsorption experiments [(a)–(c)]. This higher edge energy could potentially result from either mixed valence U in multiple sites within the sample [e.g. incorporated U(IV) and adsorbed U(VI)] or it could indicate U(V) stabilized through incorporation into the magnetite structure (Huber *et al.*, 2012), and which has also been reported for other iron oxides (Ilton *et al.*, 2010, 2012; Boland *et al.*, 2014). Additionally, the coprecipitation spectra displayed a broad white line with a shoulder on the rising edge (Fig. 3) that was absent in the sorption samples. This has been attributed to U(V) or U(VI) in octahedral uranate coordination (Van den Berghe *et al.*, 2002; Soldatov *et al.*, 2007). Therefore, both the partial shift to lower absorption edge energy, and the shoulder feature in the coprecipitation experiments suggest that U(V) may be present within the samples with uranate-like coordination. The YCL adsorption sample also displayed a broad white line, although there was no clear shoulder on the rising edge and the edge position was aligned to the U(IV) reference. In contrast to the coprecipitation experiments, this suggests that the dominant signal was from U(IV), likely to be UO_2 but the broadening of the white line may indicate a small

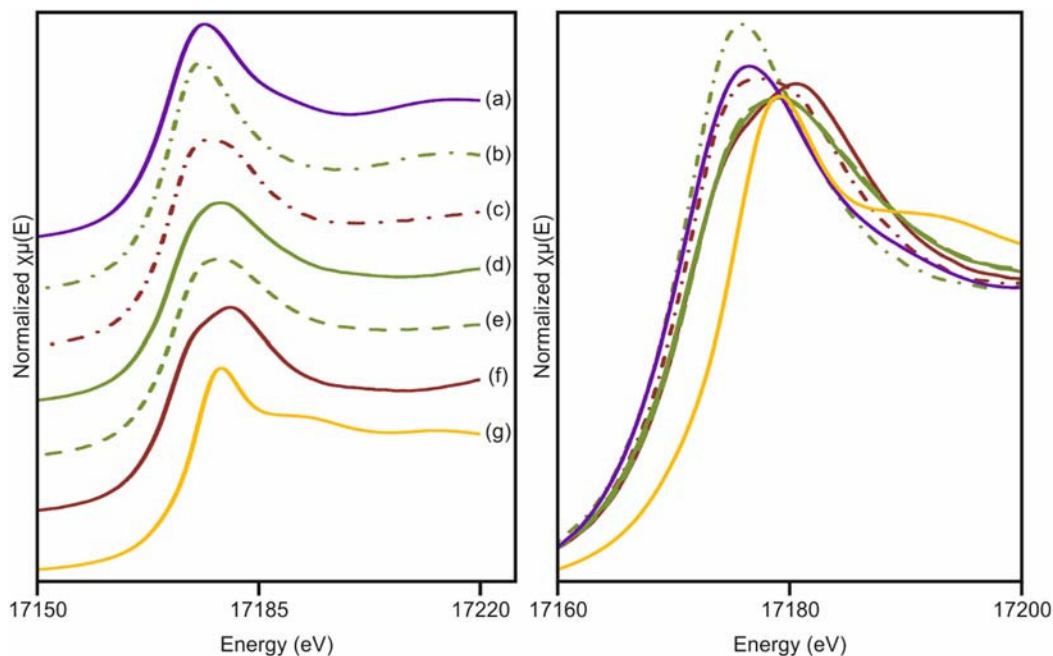


FIG. 3. U L_{III} -edge XANES spectra; (a) uraninite (U(IV) reference); (b) OCL adsorption; (c) YCL adsorption; (d) OCL coprecipitation reduced; (e) OCL coprecipitation oxidized; (f) YCL coprecipitation reduced; (g) schoepite (U(VI) reference).

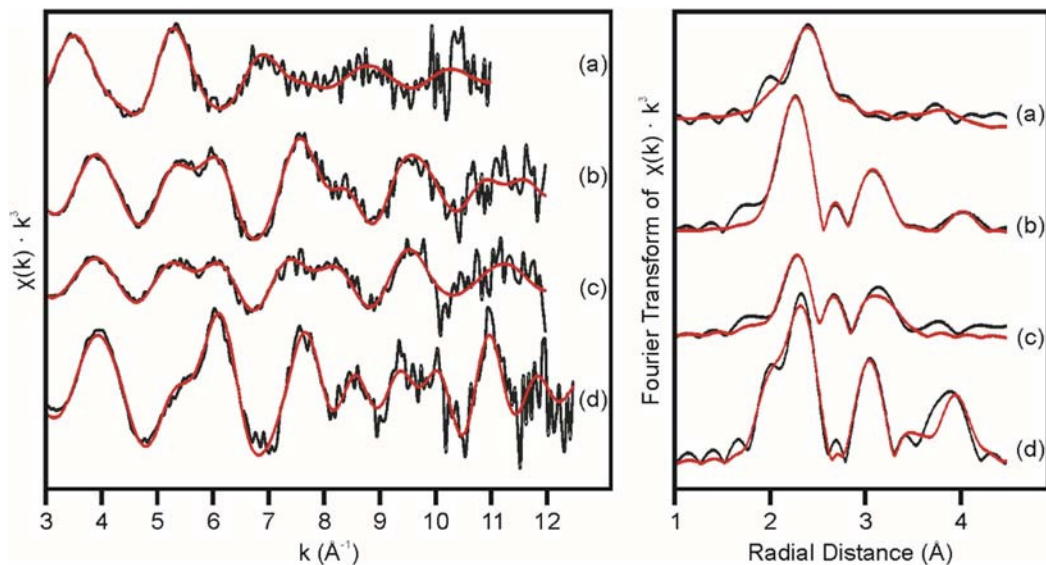


FIG. 4. U L_{III} -edge k^3 -weighted EXAFS spectra (left) and Fourier transforms plotted without a phase shift (right). Black lines are data and red lines are fits to the data; (a) OCL adsorption; (b) OCL coprecipitation reduced; (c) OCL coprecipitation oxidized; (d) YCL coprecipitation reduced.

proportion of the U becoming incorporated into the magnetite as a result of dynamic equilibrium or crystal ageing (e.g. Ostwald ripening; Kahlweit, 1975).

The EXAFS data and the respective fits for each sample are shown in Fig. 4, and the refined fit parameters are given in Table 1. The OCL adsorption sample is markedly different from the other spectra, lacking a peak in the EXAFS at $k \approx 6 \text{ \AA}^{-1}$, and also lacking a peak in the Fourier transform (FT) at 3 \AA . It was expected that UO_2 would reductively precipitate on the preformed magnetite surface (Singer *et al.*, 2012b; Latta *et al.*, 2014) and this sample was best fit using the uraninite (UO_2) structure with the addition of an Fe shell at 2.97 \AA (Table 1). Interestingly, the addition of an Fe shell suggests that some component of U(IV) was present as an adsorbed species (Latta *et al.*, 2014). Note that due to the lower quality of the data from the OCL adsorption sample it was not possible to achieve a reasonable fit while refining all parameters simultaneously; therefore the coordination numbers for each shell were changed systematically until the best fit to the data was achieved using the model of UO_2 (Singer *et al.*, 2012b) and U(IV) adsorbed to magnetite (Latta *et al.*, 2014) as the underlying structural models. Fitting of the reduced OCL coprecipitated sample assumed incorporation of U into the magnetite octahedral site (Fleet, 1981). The refined U–Fe distances for the two Fe shells (3.21 \AA and 3.46 \AA) in this fit corresponded well to the Fe–Fe distances in octahedral magnetite (2.97 \AA and 3.48 \AA). The longer U–Fe distance of the first shell (3.21 \AA versus 2.97 \AA) is presumably due to expansion of the octahedral site to accommodate U in place of Fe, as predicted by Kerisit *et al.* (2011). The crystal radius depends upon the coordination and the oxidation state of the element in question; in six-fold coordination, U(IV), U(V) and U(VI) have crystal radii of 1.03 \AA , 0.90 \AA and 0.87 \AA , respectively (Shannon, 1976), whilst the crystal radii of Fe(II) and Fe(III) in six-fold coordination (high spin) are 0.920 \AA and 0.785 \AA , respectively (Shannon, 1976). Both Fe(II) and Fe(III) occupy the octahedral sites in the magnetite inverse spinel structure, but substitution of U into the structure is likely to preferentially replace Fe(III) to aid charge balance. Indeed, the observed expansion of the site occurs in a similar way to other instances of metal incorporation (Dardenne *et al.*, 2002; Walter *et al.*, 2009). The presence of the shoulder on the rising edge in the XANES spectrum is not inconsistent with U(V) in the OCL reduced sample. However, the axial U–O distance ($1.81 \pm 0.01 \text{ \AA}$) is very short compared to a U(V) uranyl distance ($1.90\text{--}2.05 \text{ \AA}$; Burns *et al.*,

1997; Docrat *et al.*, 1999). This may indicate that U(VI) is also present (Singer *et al.*, 2012a). Additionally, a small contribution from a U–U shell was required at 3.95 \AA to fit the peak in the FT at around 4 \AA and attributed to a small component of U(IV) as uraninite, presumably formed at the magnetite surface. The coordination number of the second Fe shell (3.1 ± 1.5) is lower than expected (6), possibly due to the nanosize of the magnetite particles leading to a significant proportion of near-surface incorporated U (Yanina and Rosso, 2008), and if some of the U is present as uraninite, this will also decrease the effective coordination numbers. The same model was applied to the OCL oxidized sample. The U–O distance of the uranyl shell lengthened to 1.87 \AA , suggesting a diminished adsorbed U(VI) component and increased incorporated U(V) component. The Debye–Waller factor for this shell is relatively large which indicates a degree of static disorder from overlapping contributions at similar U–O distances. The other U–O shells remain relatively unchanged from the reduced model, as does the first Fe shell at $\sim 3.2 \text{ \AA}$. However, a second Fe shell was not resolved in the fit and there was no uraninite component (Fig. 4). This suggests that oxidative dissolution of precipitated UO_2 is the dominant mechanism for the elevated release of U back into solution in this system.

The EXAFS data for the YCL reduced coprecipitated sample was not similar to the OCL reduced and oxidized samples, despite the obvious similarities in the XANES spectra. The FT of the YCL reduced sample has three distinct peaks which roughly correlate with the three peaks observed in the FT of the OCL reduced sample, but the relative magnitude of the peaks $>R = 2.5 \text{ \AA}$ were greater in the YCL sample, particularly for the U–U shell (3rd peak, $\sim 4 \text{ \AA}$), indicating the presence of significant U–U interactions. The EXAFS was best fit using similar U–O and U–Fe shells to the OCL reduced and oxidized samples (at 2.17 \AA and 3.12 \AA , respectively) but also with contributions from a calcium–uranate phase comprising; one U–O distance at 1.69 \AA , two Ca at 3.41 \AA , and two U–U shells at between $3.8\text{--}4.2 \text{ \AA}$, similar to the structure of Ca-clerkeite (Bots *et al.*, 2014). Note that due to the complexity of this 2-site fit it was not possible to achieve a reasonable fit while refining all parameters simultaneously; therefore the coordination numbers for each shell were changed systematically until the best fit to the data was achieved. U(VI) in solution in the YCL (pH 13.1) has been shown to be present as colloidal calcium

TABLE 1. Details of EXAFS fit parameters of U adsorbed to, and coprecipitated with magnetite and subsequent air oxidation.

	Sample	Path	CN	R+ ΔR (Å)	σ^2 (Å ²)	ΔE_0 (eV)	S0 ²	χ^2_ν	R
(a)	OCL adsorption	U–O ₁	5 ^a	2.33 (1)	0.013 (1)	2.8 ± 0.8	1.00 (0)	5.9	0.013
		U–Fe ₁	1 ^a	2.97 (2)	0.014 (3)				
		U–U ₁	9 ^a	3.81 (3)	0.034 (4)				
(b)	OCL co-ppt reduced	U–O ₁	1 ^a	1.81 (1)	0.006 (1)	14.4 ± 1.3	1.00 (0)	18.8	0.006
		U–O ₂	3.7 (4)	2.21 (1)	0.008 (2) ^b				
		U–O ₃	2.0 (4)	2.44 (3)	0.008 (2) ^b				
		U–Fe ₁	6.1 (24)	3.21 (2)	0.018 (4) ^b				
		U–Fe ₂	3.1 (15)	3.46 (4)	0.018 (4) ^b				
		U–U ₁	1 ^a	3.95 (3)	0.007 (3)				
(c)	OCL co-ppt oxidized	U–O ₁	2 ^a	1.87 (2)	0.020 (3)	13.5 ± 1.6	1.00 (0)	16.4	0.023
		U–O ₂	3.3 (5)	2.20 (1)	0.008 (2) ^b				
		U–O ₃	2.6 (5)	2.45 (2)	0.008 (2) ^b				
		U–Fe ₁	2.5 (4)	3.23 (2)	0.015 ^a				
(d)	YCL co-ppt reduced	U–O ₁	1 ^a	1.69 (3)	0.009 (5)	11.2 ± 1.2	0.80 (4)	40.8	0.022
		U–O ₂	5 ^a	2.19 (1)	0.008 (1)				
		U–Fe ₁	6 ^a	3.14 (2)	0.012 (2)				
		U–Ca ₁	2 ^a	3.43 (2)	0.005 (3)				
		U–U ₁	3 ^a	3.90 (2)	0.006 (2) ^b				
		U–U ₂	1 ^a	4.20 (8)	0.006 (2) ^b				

CN denotes coordination number; R + ΔR denotes atomic distance; σ^2 denotes Debye-Waller factor; ΔE_0 denotes the shift in energy from the calculated Fermi level; S0² denotes the amplitude factor; χ^2_ν denotes the reduced Chi square value; R denotes the ‘goodness of fit’ factor. ^aParameter was fixed. ^bParameters were tied in a given fit. S0² was restrained to a value between 0.8–1.0. Numbers in parentheses are one standard deviation on the last decimal place.

uranate in systems with and without mineral phases present (Bots *et al.*, 2014; Smith *et al.*, 2014). This suggests that during coprecipitation the U(VI) interaction with the crystallizing magnetite was via a nanoparticulate colloidal phase and not via ionic solution species, leading to the possibility of nanoparticulates adsorbed to the magnetite surface, and even potentially entrapped Ca-uranate inclusions within the iron oxide particles.

Incorporation of uranium into magnetite has previously been modelled using atomistic (Kerisit *et al.*, 2011) and quantum-mechanical (Shuller-Nickles *et al.*, 2014) approaches. The atomistic approach modelled the structure using incorporated U(IV), U(V) and U(VI) atoms in both octahedral and tetrahedral sites (Kerisit *et al.*, 2011). In contrast, the quantum mechanical modelling used specific incorporation modes including: U(VI) incorporation into an octahedral site coupled with either octahedral or tetrahedral vacancies; and U(IV) incorporation into an octahedral site with either creation of an octahedral vacancy or Fe(III) reduction to Fe(II) to maintain charge balance (Shuller-Nickles *et al.*, 2014). In the latter study, incorporation of U(VI) in

an octahedral site in place of an Fe(III), with creation of an octahedral Fe(III) site vacancy was considered the most energetically favourable. Furthermore, they reported a U–Fe coordination of 4 rather than 6 for charge compensation purposes. With this arrangement, the U–Fe₁ atomic distances refined to 3.08 Å, whilst the equatorial U–O₂ distances were 2.13 Å. These are in close agreement with the refined interatomic distances and coordination numbers in the best fit model presented in Table 1.

However, in the study of Shuller-Nickles *et al.* (2014) the axial U–O₁ distance was 2.26 Å, longer than their equatorial U–O₂ distance, and significantly longer than the 1.8–1.9 Å observed here, and suggesting a component of adsorbed U(VI) in the current study. The equivalent model from atomistic simulations (Kerisit *et al.*, 2011) does not distinguish between axial and equatorial O, however, the U–O distance reported (2.13 Å) corresponds well with the average distance reported here (2.11–2.22 Å) and from quantum-mechanical modelling (2.17 Å) (Shuller-Nickles *et al.*, 2014). The Fe coordination observed here agrees well with quantum-mechanical modelling (Shuller-Nickles *et al.*,

2014), but was in excess of that observed experimentally elsewhere (Huber *et al.*, 2012). However, in contrast to the study of Huber *et al.* (2012), here U was present during magnetite formation, indicating that U incorporation occurs mainly in the rapid crystallization phase, but continues during ageing.

Conclusions

Chemical dissolution data suggest that uranium is incorporated into magnetite in both cement leachates (pH 10.5–13.1). Partial oxidation resulted in some release of U(VI) but a significant fraction of incorporated uranium was retained (Fig. 2). XANES data suggest that U was present as U(V), although contributions from both U(IV) and U(VI) could not be excluded. Fitting of the EXAFS data suggests that U, possibly as U(V), is incorporated into an expanded octahedral site within the magnetite structure in place of Fe. In the YCL coprecipitation sample, in addition to U(IV), some of the U appears to be present as a U(VI)-Ca-uranate phase. This is due to the formation of U(VI) colloids in the YCL as opposed to U(VI) being present as an aqueous species as predicted by PHREEQC modelling (Bots *et al.*, 2014). Despite the formation of U(VI) colloids at high pH, magnetite appears to have the ability to reductively immobilize U(VI) from solution through incorporation.

Magnetite is a common corrosion product of steel (Musić *et al.*, 1993; Dodge *et al.*, 2002), and may therefore potentially provide a pathway for significant incorporation and immobilization of uranium derived from radioactive waste contained within a geological disposal facility.

Acknowledgements

This work has been funded through the UK Natural Environment Research Council (NERC) BIGRAD consortium (NE/H007768/1) and Diamond Light Source beamtime grants SP7367, SP7593 and SP8070. The authors thank Steve Parry and Richard Doull for assistance at Diamond, and Paul Lythgoe, Mike Ward and John Waters for assistance with ICP-MS, TEM and XRD analyses, respectively.

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