

Bowley, Hannah E. and Young, Scott D. and Ander, E. Louise and Crout, Neil J.M. and Watts, Michael J. and Bailey, Elizabeth H. (2016) Iodine binding to humic acid. Chemosphere, 157 . pp. 208-214. ISSN 1879-1298

Access from the University of Nottingham repository:

http://eprints.nottingham.ac.uk/35195/1/Bowley%20et%20al%202016.pdf

Copyright and reuse:

The Nottingham ePrints service makes this work by researchers of the University of Nottingham available open access under the following conditions.

This article is made available under the Creative Commons Attribution Non-commercial No Derivatives licence and may be reused according to the conditions of the licence. For more details see: http://creativecommons.org/licenses/by-nc-nd/2.5/

A note on versions:

The version presented here may differ from the published version or from the version of record. If you wish to cite this item you are advised to consult the publisher's version. Please see the repository url above for details on accessing the published version and note that access may require a subscription.

For more information, please contact eprints@nottingham.ac.uk

Iodine binding to Humic Acid

Bowley H.E.^{1,2}, Young S.D.^{1,*}, Ander E.L.², Crout N.M.J.¹, Watts M.J.² & Bailey E.H.¹

¹ University of Nottingham, School of Biosciences, Sutton Bonington Campus, Loughborough, Leics,

LE12 5RD UK.

² Inorganic Geochemistry, Centre for Environmental Geochemistry, British Geological Survey,

Keyworth, Nottingham, NG12 5GG, UK.

*Corresponding Author

1 1 Abstract

2 The rate of reactions between humic acid (HA) and iodide (I) and iodate (IO_3) have been investigated in suspensions spiked with ¹²⁹I at concentrations of 22, 44 and 88 μ g L⁻¹ and stored at 3 10°C. Changes in the speciation of ${}^{129}I^{-}$, ${}^{129}IO_{3}^{-}$ and mixed (${}^{129}I^{-}+{}^{129}IO_{3}^{-}$) spikes were monitored over 77 4 5 days using liquid chromatography inductively coupled plasma mass spectrometry (LC-ICP-MS). In 6 suspensions spiked with ¹²⁹I⁻ 25% of the added I⁻ was transformed into organic iodine (Org-¹²⁹I) within 77 days and there was no evidence of ${}^{129}IO_3^-$ formation. By contrast, rapid loss of ${}^{129}IO_3^-$ and 7 increase in both ¹²⁹I⁻ and Org-¹²⁹I was observed in ¹²⁹IO₃-spiked suspensions. However, the rate of 8 9 $Org^{-129}I$ production was greater in mixed systems compared to $^{129}IO_3^{-}$ -spiked suspensions with the same total ¹²⁹I concentration, possibly indicating $IO_3 - I$ redox coupling. Size exclusion 10 chromatography (SEC) demonstrated that Org-¹²⁹I was present in both high and low molecular 11 12 weight fractions of the HA although a slight preference to bond with the lower molecular weight 13 fractions was observed indicating that, after 77 days, the spiked isotope had not fully mixed with the native ¹²⁷I pool. Iodine transformations were modelled using first order rate equations and fitted 14 15 rate coefficients determined. However, extrapolation of the model to 250 days indicated that a pseudo-steady state would be attained after ~ 200 days but that the proportion of 129 I incorporated 16 into HA was less than that of ¹²⁷I indicating the presence of a recalcitrant pool of ¹²⁷I that was 17 18 unavailable for isotopic mixing.

19 Keywords: Humic acid; iodine; kinetics; speciation; iodine-129; soil.

20 2 Introduction

21 Iodine is an essential micro-nutrient for all mammals, with a recommended daily dietary intake of

100-150 µg d⁻¹ for humans (Johnson, 2003). Iodine deficiency diseases (IDDs) are a global health problem, estimated to affect up to one third of the world's population (WHO, 2004). They are a significant social and economic stress on developing countries and an area of concern for developed countries (WHO 2008). Mechanisms that control the transfer of iodine from the terrestrial environment to the food chain are a poorly understood component of global iodine cycles (Johnson 2003a).

Soil iodine concentrations in the UK are typically within the range $0.5 - 100 \text{ mg kg}^{-1}$ (Whitehead, 28 29 1984, Johnson, 2003b). Seawater is considered to be the largest source of iodine to the terrestrial 30 biosphere (Muramatsu and Wedepohl, 1998, Fuge and Johnson, 1986) therefore higher soil 31 concentrations tend to occur in coastal areas. High soil iodine concentrations are also often 32 associated with high organic matter content, with humus the primary reservoir of soil iodine (Dai et 33 al., 2009, Smyth and Johnson, 2011, Xu et al., 2011b). In general, a larger soil iodine concentration 34 generally results in greater iodine uptake by plants (Weng et al., 2008, Dai et al., 2006), however 35 iodine associated with organic matter appears to be relatively unavailable for plant uptake (Keppler 36 et al., 2003, Xu et al., 2011b, Xu et al., 2012).

37 Humic acid (HA) is the colloidal fraction of humus, containing both aliphatic and aromatic moieties, 38 depending on the degree of humification and the original vegetation source of the organic matter. 39 Its large surface area and significant presence in soil organic matter mean that it is highly influential 40 in determining soil iodine dynamics (Xu et al., 2011b, Yamada et al., 2002, Francois, 1987, Hansen et 41 al., 2011, Allard, 2006). Although the composition of HA varies between soils, the functional groups 42 present are similar and therefore understanding iodine interactions with HA contributes significantly 43 to understanding its dynamics in soil (Schlegel et al., 2006, Warner et al., 2000, Saunders et al., 44 2012). A few authors have investigated reactions between HA and iodine (e.g. Christiansen and 45 Carlsen, 1991, Reiller et al., 2006, Choung et al., 2013) but the majority of studies involve other 46 reaction components, often in complex media such as aerosols; the rate of reaction between I and 47 HA has generally not been explicitly considered. Iodination of HA is thought to occur mainly via 48 reduction of IO_3^- to produce reactive intermediate species, such as I_2 or HOI, followed by 49 electrophilic substitution reactions with electron donor groups on the HA (Francois 1987). 50 Whitehead (1974) identified weakly acidic phenolic and amino acid groups as the most likely 51 reaction sites. Christiansen and Carlsen (1991) concluded that HA reacted with a transient iodine 52 species resulting from reaction between I, peroxidase enzymes and hydrogen peroxide over 20 53 minutes, although they were unable to determine whether the transient species was an enzyme-54 iodine complex or HOI/I_2 produced by action of the peroxidase on I^{-} . They suggested three possible 55 types of sites in the HA were available for reaction: primary sites where I was weakly bonded, 56 secondary sites susceptible to nucleophilic iodide-iodide substitution and tertiary electrophilic sites. 57 They also observed that iodine appeared to be uniformly distributed across all molecular size 58 Schlegel et al. (2006) used Extended X-ray Absorption Fine Structure (EXAFS) to fractions. 59 investigate naturally iodinated marine-derived humic substances with 10% iodine by weight and 60 concluded that iodine was covalently bonded, primarily to aromatic rings, but probably present as 61 more than one type of organic species. Preferential bonding of iodine to polycyclic aromatic moities 62 in black carbon was also demonstrated by Choung et al. (2013).

The objectives of this work were to (i) measure and model the dynamics of the reaction between IO₃⁻ and I⁻ and HA (ii) establish whether HA contains a recalcitrant (unreactive) pool of iodine and (iii) investigate whether iodine exhibits preferential association with high or low molecular weight moieties in HA. This directly addresses our currently poor understanding of the interaction of iodine with soil organic matter.

68 **3** Materials and methods

69 **3.1** Preparation and characterisation of humic acid solutions

70 Humic acid (HA) was extracted from a coniferous plantation soil, Leicestershire, UK (52° 42' N, 1° 14' 71 W; 195 m) by shaking for twelve hours with 0.1 M NaOH, followed by centrifugation (10,000 g, 72 15 min) and acidification to pH 2 of the supernatant, using concentrated HCl. Humic and fulvic 73 fractions were separated by centrifugation and the HA was purified using dialysis against 1 %v/v HCl 74 and HF then deionised water; the resultant HA was then freeze dried and finely ground (Marshall et 75 al., 1995). A portion of dried, ground HA was dissolved in 0.016 M NaOH and adjusted to pH 7.0 using NaOH to give a final concentration of 7.18 mg HA mL⁻¹. A neutral pH value was chosen to be 76 77 close to typical arable soil conditions and avoid volatilization of I₂. Dissolved organic carbon in the HA 78 solution was determined using a Shimadzu TOC-VCPH analyser. Each sample was acidified to pH 2-3 79 using HCl to remove inorganic carbon, before the remaining (organic) carbon was detected as CO₂ by 80 non-dispersive infrared detection after heating the sample to 720 °C with a platinum-coated alumina catalyst. Samples were quantified against standards of 2.125 g L⁻¹ potassium hydrogen phthalate (C 81 concentration = 1000 mg L⁻¹), diluted to appropriate concentrations with Milli-Q water (18.2 M Ω 82 83 cm). Total iodine concentration in the HA suspension was measured by ICP-MS (Thermo-Fisher Scientific X-Series^{II}) using internal detector cross-calibration with Rh and Re (10 μ g L⁻¹) as internal 84 standards. Stock standards for 127 I were prepared at iodine concentrations of 1000 mg L⁻¹ from 85 86 oven-dried KI and KIO₃, and stored at 4 °C in 1 % tetra methyl ammonium hydroxide (TMAH).

Iodine-129 was obtained as a solution of sodium iodide (¹²⁹I⁻) from the American National Institute of 87 Standards (NIST, Gaithersburg, Maryland, USA; CRM 4949C, 0.004 mol L⁻¹ Na¹²⁹I, 3451 Bq mL⁻¹). 88 lodate-129 (¹²⁹IO₃) was prepared by oxidation of ¹²⁹I with sodium chlorite using a method adapted 89 90 from Yntema and Fleming (1939). Successful oxidation to IO₃ was confirmed by ICP-MS with in-line 91 chromatographic separation using HPLC (Dionex, ICS-3000) operated in isocratic mode with an anion exchange column (Hamilton PRP-X100; 250 x 4.6 mm, 5 μm particle size). The mobile phase (1.3 mL 92 93 min⁻¹) was 60 mmol L⁻¹ NH₄NO₃, 1 x 10⁻⁵ mmol L⁻¹ Na₂-EDTA, 2% methanol; pH was adjusted to 9.5 94 with TMAH.

Triplicate aliquots of HA solution (pH 7) were spiked with ¹²⁹I⁻ and/or ¹²⁹IO₃⁻ on 8 occasions to give incubation times of 26, 79, 155, 328, 596, 992, 1404 and 1855 hr. Samples were stored in the dark at 10 °C, the average annual soil temperature measured at Armagh Observatory, Northern Ireland (Garcia-Suarez and Butler, 2006). Final concentrations of spiked iodine were 22.1, 44.1 and 88.2 μ g L⁻¹ (Table 1). The CRM 4949C contained ¹²⁷I equivalent to 12 % of the ¹²⁹I concentration; this was accounted for as described in Electronic Annex A so that, for simplicity, all iodine added in the spike (¹²⁹I + ¹²⁷I) is referred to as ¹²⁹I and native iodine is described as ¹²⁷I.

102 **3.2** Iodine speciation

103 lodine species were separated using HPLC (Dionex, ICS-3000), with a Superose 12 10/300 GL size 104 exclusion chromatography (SEC) column (GE Healthcare), linked to ICP-MS. A sample injection 105 volume of 25 μ l and isocratic elution with 0.1 M Tris (Tris(hydroxymethyl)aminomethane) adjusted 106 to pH 8.8 using 50 % Trace Analysis Grade (TAG) HNO₃ at a flowrate of 1 mL min⁻¹ was used. The 107 HPLC was controlled using Chromeleon software (Dionex, version 6.80SR12) and sample processing was undertaken using Plasmalab software. Working standards of ¹²⁷I⁻ and ¹²⁷IO₃⁻, at iodine 108 concentrations of $0 - 100 \,\mu g \, L^{-1}$, were used. Species-specific quantification was carried out with 109 standards of ¹²⁷I⁻, ¹²⁷IO₃⁻, ¹²⁹I⁻ and ¹²⁹IO₃⁻ prepared in Milli-Q water and mean, isotope-specific, 110 111 sensitivity values (signal (integrated CPS) per ppb) calculated from iodide and iodate standards were 112 used to quantify HA-bonded iodine (HA-I); drift correction was applied by analysis of repeated 113 standards through the experimental run. Limits of detection were 0.047 μ g L⁻¹ for ¹²⁷I and 0.014 μ g L⁻ 114 ¹ for ¹²⁹I. Humic acid controls, spiked with Milli-Q water, were analysed alongside ¹²⁹I-spiked samples to determine the equilibrium speciation of native ¹²⁷I. 115

A correction factor for ¹²⁹Xe was determined for each run and applied to all ¹²⁹I results, according to
 Equation 1:

118
$${}^{129}S = {}^{129}S_{\text{meas}} - (k \times {}^{131}S_{\text{meas}})$$
 (1)

119 where ¹²⁹S = corrected signal (CPS) for ¹²⁹I; ¹²⁹S_{meas} = measured signal at m/z = 129; k = a factor 120 determined by iteration for each run (typically 1.08) to give an average ¹²⁹S baseline of zero; ¹³¹S_{meas} 121 = measured signal for ¹³¹Xe. Matrix matching and standard addition were used to calculate mean 122 sensitivity for all samples in each run, from which concentrations of ¹²⁹I and ¹²⁷I in each peak were 123 quantified.

124 **3.3 Modelling**

125 Iodine-129 transformations were represented as simultaneous ordinary differential equations (Fig.126 1):

127
$$\frac{d[IO_3]}{dt} = -(k_1 + k_4)[IO_3] + k_5[Org-I]$$
(2)

128
$$\frac{d[I^{-}]}{dt} = -k_3[I^{-}] + k_1[IO_3^{-}] + k_2[Org-I]$$
(3)

129
$$\frac{d[Org-I]}{dt} = -(k_2+k_5)[Org-I] + k_4[IO_3] + k_3[I^-]$$
(4)

130 Where k_1 , k_2 , k_3 , k_4 and k_5 are unknown rate coefficients (hr¹) estimated by fitting the model to the 131 observed concentrations of ¹²⁹IO₃⁻, ¹²⁹I⁻ and Org-¹²⁹I.

The differential equations were solved using 4th order Runge-Kutta and fitting was performed using a Marquardt procedure (Press et al, 1986) implemented in <u>OpenModel</u> (Tarsitano et al 2011, http://openmodel.info/) to minimize the residual sum of squares (RSS) between modelled and observed values over all the time points and different spike combinations. Alternative model structures were considered, the arrangement described (and shown schematically in Fig. 1) provided the best fit to the data.

138 4 Results and discussion

139 **4.1** Iodine dynamics

140 Measured characteristics of the HA are given in Table 2. Concentrations of ¹²⁷I species represent 141 iodine in equilibrium with HA suspensions with median values of 98.0 μ g L⁻¹ of Org-¹²⁷I and 15.1 μ g L⁻¹

142 1^{127} , iodate $(1^{27}$ IO₃) was not detected. The C:I mole ratio in the HA was 396,000.

Figure 2 shows the change in speciation of spiked ¹²⁹I over time; model fits to the data are discussed in a later section. In suspensions spiked with Γ only (Table 1: I_{22} , I_{44} and I_{88}) a gradual decrease in ¹²⁹I concentrations was observed over time with a concomitant increase in Org-¹²⁹I. By the final time

point (1855 hr) ~25% of added ¹²⁹I⁻ had been transformed to Org-¹²⁹I irrespective of initial spike

concentration. No generation of $^{129}IO_3^-$ was observed. Suspensions spiked with $^{129}IO_3^-$ (Table 1: IO_{22} , 147 IO_{44} and IO_{88}) showed rapid initial loss of ¹²⁹ IO_3^- with associated increases in both ¹²⁹ I^- and Org-¹²⁹I. 148 Concentrations of ¹²⁹I⁻ were always greater than Org-¹²⁹I. By the end of the experiment (1855 hr) 149 concentrations of ${}^{129}IO_3^{-129}$ in each system were < 5% of the initial concentrations and $Org^{-129}I$ 150 151 accounted for ~40% of the added iodine, again irrespective of initial spike concentration. In mixed systems (Table 1; Mix₂₂, Mix₄₄ and Mix₈₈), where equal concentrations of ¹²⁹I⁻ and ¹²⁹IO₃⁻ were added, 152 results were similar to those observed when ¹²⁹IO₃⁻ alone was added (Figure 2). Concentrations of 153 ¹²⁹IO₃⁻ decreased rapidly with both ¹²⁹I⁻ and Org-¹²⁹I increasing over time. Again no concentration-154 dependence was observed, with similar proportions of ¹²⁹I⁻ and Org-¹²⁹I observed regardless of spike 155 156 concentration.

157 Humic acid has been reported to both reduce IO_3^- and oxidise I⁻ in soils (Yamaguchi et al., 2010), 158 however oxidation of I by organic matter to form IO_3^- is expected to be much slower than reduction of IO_3^- to I⁻ (Schlegel et al., 2006) unless Fe^{III} and Mn^{IV} oxides are present which may catalyse the 159 oxidation of I⁻ (Gallard et al., 2009, Allard et al., 2009, Fox et al., 2009). This difference in rate 160 (between I and IO_3), in the absence of metal oxides, was confirmed by the absence of ¹²⁹ IO_3 in ¹²⁹I 161 spiked systems, in contrast to the rapid production of $^{129}I^{-}$ in HA solutions spiked with $^{129}IO_{3}^{-}$. 162 163 Shetaya et al. (2012), selected soils with a range of pH values and concentrations of Fe/Mn oxides and organic matter for incubation with $^{129}I^{-}$ or $^{129}IO_3^{-}$ at 10°C and 20°C. In contrast to the current 164 165 study, and that of Choung et al. (2013), they found that conversion of added ¹²⁹I⁻ to (humic-bound) Org-¹²⁹I ($t_{\frac{1}{2}}$ = 0.38 – 45 hr) was faster than conversion of ¹²⁹IO₃⁻ to Org-¹²⁹I ($t_{\frac{1}{2}}$ = 9.0 -412 hr). 166 167 Comparison of these half-life values for whole soil with the data comparing iodide and iodate in Fig. 168 2 may indicate the importance of metal oxides in soils in rapidly adsorbing IO₃, and thereby slowing 169 its transformation into Org-I, or possibly enabling oxidation of I⁻ to an intermediate species capable 170 of reacting with humus. Additionally the large overall differences in reaction rates seen between 171 soils and HA may reflect enzyme activity in soils which is likely to be absent from isolated HA 172 fractions.

The rate of $Org^{-129}I$ production was greater in the systems spiked with $^{129}IO_3^-$ than those spiked with $^{129}I^-$ at the same concentration (Fig. 2). However, in mixed systems the *initial* rate of $Org^{-129}I$ production was greater than in solutions spiked with $^{129}IO_3^-$ alone, although a slightly lower concentration of $Org^{-129}I$ was eventually present after 500 hr incubation (e.g. IO_{44} and Mix_{44}). However comparison of systems with the same initial concentration of $^{129}IO_3^-$ (e.g. IO_{22} and Mix_{44}) showed that $Org^{-129}I$ concentrations were substantially greater in the mixed systems where $^{129}I^-$ was

also present compared to the equivalent ${}^{129}IO_3^-$ -spiked system. This may suggest that redox coupling between ${}^{129}I^-$ (or ${}^{127}I^-$) and ${}^{129}IO_3^-$ had occurred in the mixed spike solutions (Eq. 2):

181

 $IO_3^- + 5I^- + 6H^+ \rightleftharpoons 3I_2 + 3H_2O$ (2)

182 When only ¹²⁹I⁻ was added, its transformation to $Org^{-129}I$ did not depend upon concentration (Fig. 3), 183 suggesting its oxidation mechanism was independent of the presence of another species. However, 184 considering the presence of native iodide (¹²⁷I⁻) in the HA solutions, the *relative* rate of ¹²⁹I⁻ 185 production, from ¹²⁹IO₃⁻, in the mixed spike system was greater when lower concentrations of ¹²⁹IO₃⁻ 186 were added (Fig. 3); this result is consistent with faster reduction of IO₃⁻ (to I⁻) at higher 187 iodide/iodate ratios. For the mixed-spike system the ratio of measured iodide to added iodide 188 progressively exceeded 1.0, reflecting the production of I⁻ from reduction of added IO₃⁻.

The *relative* rates of ¹²⁹IO₃⁻ loss in both ¹²⁹IO₃⁻ and mixed-spike systems were slightly greater at lower IO₃⁻ concentrations with similar rates observed for systems with the same total ¹²⁹I concentration (Fig. 4). Again, this suggests that the rate of iodate reduction is limited, either by the concentration of iodide (Eq. 2) or some other reductive mechanism associated with the HA.

193 Figure 5 compares the concentrations of species measured in systems where a mix of species was 194 added with the sum of species from the appropriate ${}^{129}I^{-}$ and ${}^{129}IO_{3}^{-}$ treatments spiked with a single species (e.g. Mix_{44} compared with $I_{22}+IO_{22}$). Although it could be argued that this comparison is not 195 completely valid because the ¹²⁹I:HA ratio differed, the rate of ¹²⁹I⁻ transformation to Org-¹²⁹I (Fig. 3, 196 197 open symbols) indicated that HA concentration is not a limiting factor. Figure 5 suggests that 198 initially there was less ¹²⁹I, and consequently greater concentrations of ¹²⁹IO₃ and Org-¹²⁹I, in the 199 mixed spike systems (circles fall below the 1:1 line). However, the difference in species composition 200 between the mixed and single-species systems generally declined with time - eg highest values for 201 iodide (open and closed circles).

- 202 Comparison of the concentrations of the species present in single species spiked systems with ¹²⁹I-
- 203 spiked systems at an equivalent total iodine concentration (e.g. $IO_{44} + I_{44}$ compared with I_{88}) also
- show that ¹²⁹I^{\circ} was only transformed to Org-¹²⁹I and that this happened faster when ¹²⁹IO₃^{\circ} was
- 205 present, perhaps reflecting the 5:1 ratio in Eq. 2.

206 **4.2** Iodine distribution within the HA

The distribution of iodine within the HA was examined using SEC-ICP-MS. Lower molecular weight (LMW) HA may be expected to react more easily with iodine due to its greater surface area which may provide greater accessibility to reactive sites. Conversely, the greater negative charge density 210 expected on LMW fractions may act to exclude or delay I^{-} or IO_{3}^{-} ions from interaction. Xu et al. 211 (2011a), in a study where IO_3^{-1} was added to HA at pH 3, observed that the LMW fractions (3 – 50 212 kDa) sorbed more iodine than the > 50 kDa fractions after 72 hr. By contrast, Christiansen and 213 Carlsen (1991) observed no dependence on MW size fraction following a rapid (20 min) reaction of 214 Aldrich HA with I⁻ in the presence of peroxidase enzymes. Figure 6 shows the cumulative integrated 215 counts per second (ICPS) for the organic portion of the SEC chromatogram for added ¹²⁹I and pre-216 existing ¹²⁷I in the HA after a 77 day incubation of the Mix₈₈ solution. Higher molecular weight 217 (HMW) molecules elute first due to exclusion from the column matrix - shown by the prominent 218 peak around 400 s. The lines representing the cumulative integrated signals for ¹²⁷I and ¹²⁹I in Figure 219 6 suggest that a greater proportion of ¹²⁷I than ¹²⁹I was present in larger HA molecules. Approximately 50% of the cumulative integrated signal for ¹²⁷I had been detected by 530 s, 220 221 compared to 600 s for ¹²⁹I. This indicates that ¹²⁹I had not fully mixed with the pre-existing HA-bound 222 ¹²⁷I and had reacted preferentially with the LMW fractions of HA. It also indicates the presence of 223 recalcitrant iodine (¹²⁷I), unavailable to mix with the added ¹²⁹I, and with a greater presence in the HMW fractions of HA. Comparison of the Mix₈₈ 77 day chromatogram for ¹²⁷I and ¹²⁹I (Figure 6) also 224 shows a smaller excluded peak maxima compared to the main organic peak for ¹²⁹I and a shift in the 225 226 ¹²⁹I peak maxima towards LMW fractions, again suggesting preferential binding of freshly added 227 iodine to LMW fractions of HA in contrast to (native) ¹²⁷I.

228 **4.3** Modelling

229 Fitted model (Fig. 1) predictions are compared to the observations in Fig. 2 and estimated rate coefficients are given in Table 3. Overall the model fit was good (for all data: $r^2 = 0.96$, p < 0.001), 230 231 supporting the model structure. The values of the (apparent) rate constants broadly reflect the 232 dynamics of the system. lodate is rapidly converted to humic-bound forms ($k_4 = 2.62 \text{ hr}^{-1}$) but the 233 optimal model fit required a reverse reaction ($k_5 = 0.157 \text{ hr}^{-1}$) suggesting an approach to an 234 equilibrium position rather than a zero sink for iodate. The further reduction of iodate to iodide was slower ($k_1 = 4.11 \times 10^{-3} hr^{-1}$) and both the 'direct' conversion of iodide to humic-bound iodine and the 235 re-mineralisation of iodide were comparatively very slow reactions ($k_3 = 3.16 \times 10^{-4}$ and $k_2 = 4.67 \times 10^{-4}$ 236 237 10^{-4} hr^{-1}).

In order to investigate the availability of ¹²⁷I for mixing with ¹²⁹I, modelling was extended to 250 d. At this point a pseudo-steady-state was apparent, with no significant changes in species concentrations from 200 d after spiking. After 250 d contact between ¹²⁹I and HA the model predicted that ¹²⁹I⁻ /Org-¹²⁹I = 0.24, whereas the measure ratio ¹²⁷I⁻ /Org-¹²⁷I was lower, at 0.17. Therefore although the model prediction suggested that ¹²⁹I was at equilibrium, a greater proportion of ¹²⁷I than ¹²⁹I existed as Org-I, suggesting the presence of a recalcitrant pool of ¹²⁷I, unavailable for isotopic mixing. From the ¹²⁹I⁻ /Org-¹²⁹I ratio and measured ¹²⁷I⁻ in solution, the estimated labile Org-¹²⁷I was 62.9 μ g L⁻¹, suggesting that 64% of the HA bound iodine was labile and 36% was incapable of isotopic mixing within 250 d.

247 Keppler et al. (2003) and Xu et al. (2011b) suggested that iodination of HA occurs early in the 248 humification of fresh plant material with fewer binding sites available as humification progresses. 249 Steric hindrance by aliphatic chains may also make some aromatic binding sites less accessible to 250 freshly added iodine, while effectively 'fixing' native iodine (Xu et al., 2012). Schwehr et al (2009) 251 also observed, in an experiment where natural sediments were spiked with I⁻, that recently added I⁻ 252 was less strongly sorbed than native iodine, and that greater added concentrations resulted in a 253 smaller proportion bound. Therefore although organic-rich soils may well contain higher total 254 concentrations of iodine much of that iodine may not be readily accessible to plants.

255 **5** Conclusions

Inorganic iodine (both I^{-} and IO_{3}^{-}) reacted with HA to produce Org-I. The reaction was slower with I^{-} compared to systems spiked with IO_{3}^{-} or a mixture of IO_{3}^{-} and I^{-} . In mixed and IO_{3}^{-} -spiked systems I^{-} concentrations increased rapidly and remained higher than Org-I concentrations throughout the duration of the experiment.

Native iodine in the HA solutions was present only as Γ and Org-I; no IO_3^- was observed. There was evidence that the presence of native ¹²⁷ Γ allowed more rapid reduction of spiked ¹²⁹ IO_3^- , potentially by redox coupling with Γ . This was supported by the relative rates of reaction of the two species, with a faster reaction observed when a mix of inorganic species (Γ and IO_3^-) was added, rather than a single species (Γ or IO_3^- alone). In soils, by contrast, Γ has been reported to transform to Org-I more quickly than IO_3^- , possibly due to the presence of metal oxides or enzymatic oxidation. In a 'purified' HA solution the same reaction mechanisms are not available and Γ oxidation was relatively slow.

Size exclusion chromatography showed that ¹²⁹I became associated with both high and low molecular weight HA although a slight preference for lower molecular weight fractions was suggested. The native and spiked isotopes were not fully mixed after ~ 2 months of reaction indicating the presence of a recalcitrant pool of ¹²⁷I; this was also strongly indicated by modelling over longer timescales.

272 6 Acknowledgements

- 273 Funding for H. E. Bowley was provided by UoN and the BGS University Funding Initiative (BUFI-S178).
- 274 This work is published with the permission of the Executive Director, British Geological Survey.

275

276 **7** References

- Allard, B. (2006) A comparative study on the chemical composition of humic acids from forest soil,
 agricultural soil and lignite deposite. Bound lipid, carbohydrate and amino acid distributions.
 Geoderma, 130, 77 96.
- Allard, S., von Gunten, U., Sahli, E., Nicolau, R. & Gallard, H. (2009) Oxidation of iodide and iodine on
 birnessite (d-MnO₂) in the pH range 4-8. *Water Research*, 43, 3417-3426.
- Christiansen, J. V. & Carlsen, L. (1991) Enzymatically controlled iodination reactions in the terrestrial
 environment. *Radiochimica Acta*, 52-3, 327-333.
- Choung, S., Um, W., Kim, M. & Kim, M-G. 2013. Uptake mechanism for iodine species to black
 carbon. Environmental Science and Technology 47, 10349 10355.
- Dai, J. L., Zhang, M., Hu, Q. H., Huang, Y. Z., Wang, R. Q. & Zhu, Y. G. (2009) Adsorption and
 desorption of iodine by various Chinese soils: II. Iodide and iodate. *Geoderma*, **153**, 130-135.
- Dai, J. L., Zhu, Y. G., Huang, Y. Z., Zhang, M. & Song, J. L. (2006) Availability of iodide and iodate to
 spinach (Spinacia oleracea L.) in relation to total iodine in soil solution. *Plant and Soil*, 289,
 301-308.
- Fox, P. M., Davis, J. A. & Luther, G. W. (2009) The kinetics of iodide oxidation by the manganese
 oxide mineral birnessite. *Geochimica et Cosmochimica Acta*, **73**, 2850-2861.
- Francois, R. (1987) The influence of humic substances on the geochemistry of iodine in nearshore
 and hemipelagic marine sediments. *Geochimica et Cosmochimica Acta*, **51**, 2417 2427.
- Fuge, R. & Johnson, C. C. (1986) The Geochemistry of Iodine a Review. *Environmental Geochemistry and Health*, 8, 31-54.
- Gallard, H., Allard, S., Nicolau, R., von Gunten, U. & Croue, J. P. (2009) Formation of iodinated
 organic compounds by oxidation of iodide-containing waters with manganese dioxide.
 Environmental Science & Technology, 43, 7003-7009.
- Garcia-Suarez, A. M. & Butler, C. J. (2006) Soil temperatures at Armagh Observatory, Northern
 Ireland, from 1904 to 2002. *International Journal of Climatology*, 26, 1075-1089.
- Hansen, V., Roos, P., Aldahan, A., Hou, X. & Possnert, G. (2011) Partition of iodine (¹²⁹I and ¹²⁷I)
 isotopes in soils and marine sediments. *Journal of Environmental Radioactivity*, **102**, 1096 1104.
- Johnson C. C. (2003a) The geochemistry of iodine and its application to environmental strategies for
 reducing the risks from iodine deficiency disorders (idd). Brit. Geol. Surv., DFID kar project
 R7411, Report CR/03/057N.
- Johnson, C. C. (2003b) Database of the iodine content of soils populated with data from published
 literature. British Geological Survey Commissioned Report.

- Keppler, F., Biester, H., Putschew, A., Silk, P. J., Scholer, H. F. & Muller, G. (2003) Organoiodine
 formation during humification in peatlands. *Environmental Chemistry Letters*, 1, 219-223.
- Marshall, S.J. (1992) The complexation of aluminium by humic acids in fresh waters. PhD Thesis.
 Agricultural and Environmnetal Sciences, University of Nottingham.
- Marshall, S. J., Young, S. D. & Gregson, K. (1995) Humic acid-proton equilibria: A comparison of two
 models and assessment of titration error. *European Journal of Soil Science*, 46, 471-480.
- Muramatsu, Y. & Wedepohl, K. H. (1998) The distribution of iodine in the Earth's crust. *Chemical Geology*, 147, 201-216.
- Press, W.H., Teukolsky S.A., Vetterling, W.T., Flannery, B.P. (2007) Numerical Recipes 3rd Edition:
 The Art of Scientific Computing, Cambridge University Press, New York, NY
- 320Reiller, P., Mercier-Blon, F., Gimenez, N., Barre, N. & Miserque, F. (2006) Iodination of humic acid321samples from different origins. *Radiochimica Acta*, **94**, 739-745.
- Saunders, R.W., Kumar, R., MacDonald, S.M. and Plane J.M.C. (2012). Insights into the
 Photochemical Transformation of Iodine in Aqueous Systems: Humic Acid Photosensitized
 Reduction of Iodate. Environ. Sci. Technol., 46, 11854–11861.
- Schlegel, M. L., Reiller, P., Mercier-Bion, F., Barre, N. & Moulin, V. (2006) Molecular environment of
 iodine in naturally iodinated humic substances: Insight from X-ray absorption spectroscopy.
 Geochimica et Cosmochimica Acta, **70**, 5536 5551.
- Schwehr, K. A., Santschi, P. H., Kaplan, D. I., Yeager, C. M. & Brinkmeyer, R. (2009) Organo-iodine
 formation in soils and aquifer sediments at ambient concentrations. *Environmental Science* & *Technology*, 43, 7258-7264.
- Shetaya, W. H., Young, S. D., Watts, M. J., Ander, E. L. & Bailey, E. H. (2012) Iodine dynamics in soils.
 Geochimica et Cosmochimica Acta, **77**, 457-473.
- Smyth, D. & Johnson, C. C. (2011) Distribution of iodine in soils of Northern Ireland. *Geochemistry- Exploration Environment Analysis*, **11**, 25-39.
- D Tarsitano, D, Young, S.D., Crout, N.M.J. (2011) Evaluating and reducing a model of radiocaesium
 soil-plant uptake. Journal of Environmental Radioactivity, **102**, 262-269.
- Warner, J. A., Casey, W. H. & Dahlgren, R. A. (2000) Interaction kinetics of I_{2(aq)} with substituted
 phenols and humic substances. *Environmental Science & Technology*, **34**, 3180-3185.
- Weng, H. X., Weng, J. K., Yan, A. L., Hong, C. L., Yong, W. B. & Qin, Y. C. (2008) Increment of iodine
 content in vegetable plants by applying iodized fertilizer and the residual characteristics of
 iodine in soil. *Biological Trace Element Research*, **123**, 218-228.
- Whitehead, D. C. (1974) The infl uence of organic matter, chalk and sequisoxides on the solubility of
 iodide, elemental iodine and iodate inclubated with soil. *Journal of Soil Science*, 25, 461-470.

- Whitehead, D. C. (1984) The Distribution and Transformations of Iodine in the Environment.
 Environment International, **10**, 321-339.
- WHO (2004) In Iodine Status Worldwide (eds. B. Benoist, et al.). Department of Nutrition for Health
 and Development, World Health Organization, Geneva.
- WHO (2008) WHO expert consultation on salt as a vehicle for fortification, Luxembourg, 21-22
 March 2007. ISBN 978 92 4 159678 7.
- Xu, C., Miller, E. J., Zhang, S., Li, H.-P., Ho, Y.-F., Schwehr, K. A., Kaplan, D. I., Otosaka, S., Roberts, K.
 A., Brinkmeyer, R., Yeager, C. M. & Santschi, P. H. (2011a) Sequestration and remobilization
 of radioiodine (¹²⁹I) by soil organic matter and possible consequences of the remedial action
 at Savannah River site. *Environmental Science & Technology*, **45**, 9975-9983.
- Xu, C., Zhang, S. J., Ho, Y. F., Miller, E. J., Roberts, K. A., Li, H. P., Schwehr, K. A., Otosaka, S., Kaplan,
 D. I., Brinkmeyer, R., Yeager, C. M. & Santschi, P. H. (2011b) Is soil natural organic matter a
 sink or source for mobile radioiodine (¹²⁹I) at the Savannah River Site? *Geochimica et Cosmochimica Acta*, **75**, 5716-5735.
- Xu, C., Zhong, J., Hatcher, P. G., Zhang, S., Li, H.-P., Ho, Y.-F., Schwehr, K. A., Kaplan, D. I., Roberts, K.
 A., Brinkmeyer, R., Yeager, C. M. & Santschi, P. H. (2012) Molecular environment of stable
 iodine and radioiodine (I-129) in natural organic matter: Evidence inferred from NMR and
 binding experiments at environmentally relevant concentrations. *Geochimica et Cosmochimica Acta*, 97, 166-182.
- Yamada, H., Hisamorl, I. & Yonebayashi, K. (2002) Identification of organically bound iodine in soil
 humic substances by size exclusion chromatography/inductively coupled plasma mass
 spectrometry (SEC/ICP-MS). *Soil Science and Plant Nutrition*, 48, 379-385.
- Yamaguchl, N., Nakano, M., Takamatsu, R. & Tanida, H. (2010) Inorganic iodine incorporation into
 soil organic matter: evidence from iodine K-edge X-ray absorption near-edge structure.
 Journal of Environmental Radioactivity, **101**, 451-457.
- 369 Yntema, L. F. & Fleming, T. (1939) Volumetric oxidation of iodide to iodate by sodium chlorite.
 370 *Industrial and Engineering Chemistry-Analytical Edition*, **11**, 0375-0377.

371

Tables

- Table 1:
 Summary of added iodine (¹²⁹I) species in incubated HA solutions.
- Table 2:Measured characteristics of a humic acid (HA) isolated from the Ah horizon of a coniferous plantation soil
in Leicestershire (UK); (52° 42' N, 1° 14' W; 195 m).
- Table 3:Optimised parameter values (first-order rate constants) describing HA-iodine dynamics in the model
shown in Figure 1.

Solution	¹²⁹ Ι ⁻ added (μg ¹²⁹ Ι L ⁻¹)	¹²⁹ IO ₃ added (μg ¹²⁹ I L ⁻¹)	
I ₂₂	22.1	0	
I_{44}	44.1	0	
I ₈₈	88.2	0	
IO ₂₂	0	22.1	
IO ₄₄	0	44.1	
10 ₈₈	0	88.2	
Mix ₂₂	11.0	11.0	
Mix ₄₄	22.1	22.1	
Mix ₈₈	44.1	44.1	

 Table 1: Summary of added iodine (¹²⁹I) species in incubated HA solutions.

Table 2: Measured characteristics of a humic acid (HA) isolated from the Ah horizon of a coniferous plantation soil in Leicestershire (UK); (52° 42' N, 1° 14' W; 195 m).

Characteristic	Value	Units	Source
Ash content	0.66*	%	Marshall (1992)
Total acidity	6.49*	mol _c kg⁻¹	Marshall (1992)
Dissolved organic carbon	3.67	mg ml⁻¹	DOC analysis
¹²⁷ lodide [#]	15.1	µg I L ⁻¹	SEC analysis
Org- ¹²⁷ l [#]	98.0	µg I L ⁻¹	SEC analysis

*Ash content and total acidity are the mean of two measurements quoted by Marshall (1992). *Measured in HA suspensions

 Table 3: Optimised parameter values (first-order rate constants) describing HA-iodine dynamics in the model shown in Figure 1.

Rate constant	Mean (hr⁻¹)	S. D.
k1	0.00411	0.00010
k2	0.000467	0.00004
k3	0.000316	0.00002
k4	2.62	0.00000
k5	0.157	0.00003

Figures

Figure 1: Conceptual model describing transformations of spiked ¹²⁹I in the presence of HA. Rate constants $k_1 - k_5$ describe first-order rate equations.

Figure 2: Modelled (lines) and measured (symbols) change in ¹²⁹I concentrations with time following spiking with ¹²⁹I at a range of concentrations and species compositions. Species measured included ¹²⁹I⁻ (black circles \bullet , black lines), ¹²⁹IO₃⁻ (open squares \Box , dashed lines) and Org-¹²⁹I (grey triangles \blacktriangle , grey lines). Error bars show standard error of triplicate measurements.

Figure 3: Change in the ratio of measured iodide to added iodide over time, following addition of iodide (*open* symbols) and mixed iodide/iodate ¹²⁹I spikes (*closed* symbols). Total concentrations of ¹²⁹I added were: 22.1 μ g L⁻¹ (circles), 44.1 μ g L⁻¹ (squares) and 88.2 μ g L⁻¹ (triangles).

Figure 4: Change in ratio of measured iodate to added iodate over time, following addition of iodate (*open* symbols) and mixed iodide/iodate ¹²⁹I spikes (*closed* symbols). Total concentrations of ¹²⁹I added were: 22.1 μ g L⁻¹ (circles), 44.1 μ g L⁻¹ (squares) and 88.2 μ g L⁻¹ (triangles).

Figure 5: Comparison of total concentrations of iodine-129 species in solution in mixed and single spiked systems at total iodine concentrations of 44 μ g L⁻¹ (*open* symbols) and 88 μ g L⁻¹ (*closed* symbols). Dashed line = 1:1, square symbols = iodate, circles = iodide, triangles = Org-I.

Figure 6: Cumulative ICPS (integrated counts per second) as a percentage of total counts, from size exclusion chromatograph of ¹²⁹I (thick black line) and ¹²⁷I (thick grey line); 25%, 50% and 75 % of cumulated ICPS are indicated by circles, squares and diamonds, respectively. The Org-I section of ¹²⁷I (thin black line) and ¹²⁹I (thin grey line) SEC chromatograms (2° Y-axis) are also shown. All data are for the Mix₈₈ solution (Table 1) after 77 days incubation.



Figure 1: Conceptual model describing transformations of spiked ¹²⁹I in the presence of HA. Rate constants k_1 - k_5 describe first-order rate equations.



Figure 2: Modelled (lines) and measured (symbols) change in ¹²⁹I concentrations with time following spiking with ¹²⁹I at a range of concentrations and species compositions. Species measured included ¹²⁹I⁻ (black circles \bullet , black lines), ¹²⁹IO₃⁻ (open squares \Box , dashed lines) and Org-¹²⁹I (grey triangles \blacktriangle , grey lines). Error bars show standard error of triplicate measurements.



Figure 3: Change in the ratio of measured iodide to added iodide over time, following addition of iodide (*open* symbols) and mixed iodide/iodate ¹²⁹I spikes (*closed* symbols). Total concentrations of ¹²⁹I added were: 22.1 μ g L⁻¹ (circles), 44.1 μ g L⁻¹ (squares) and 88.2 μ g L⁻¹ (triangles).



Figure 4: Change in ratio of measured iodate to added iodate over time, following addition of iodate (*open* symbols) and mixed iodide/iodate ¹²⁹I spikes (*closed* symbols). Total concentrations of ¹²⁹I added were: 22.1 μ g L⁻¹ (circles), 44.1 μ g L⁻¹ (squares) and 88.2 μ g L⁻¹ (triangles).



Figure 5: Comparison of total concentrations of iodine-129 species in solution in mixed and single spiked systems at total iodine concentrations of 44 μ g L⁻¹ (*open* symbols) and 88 μ g L⁻¹ (*closed* symbols). Dashed line = 1:1 relationship, square symbols = iodate, circles = iodide, triangles = Org-I.



Figure 6: Cumulative ICPS (integrated counts per second) as a percentage of total integrated counts, from size exclusion chromatograph of ¹²⁹I (thick black line) and ¹²⁷I (thick grey line); 25%, 50% and 75 % of cumulated ICPS are indicated by circles, squares and diamonds, respectively. The Org-I section of ¹²⁷I (thin black line) and ¹²⁹I (thin grey line) SEC chromatograms (2° Y-axis) are also shown. All data are for the Mix₈₈ solution (Table 1) after 77 days incubation.

Electronic Annex A: Correction applied for the presence of ¹²⁷I in the ¹²⁹I SRM 4949C:

All measured ¹²⁹I concentrations had a correction applied to nominally ascribe *all* iodine added in ¹²⁹I spikes (¹²⁹I + ¹²⁷I) to ¹²⁹I, for ease of description:

$$^{129}I = {}^{129}I_{meas} \times 1.12 \times (127/129)$$
 (A1)

Where ¹²⁹I = 'corrected' concentration of I from the spike in solution (μ g L⁻¹) i.e. (¹²⁹I + ¹²⁷I present in the spike), ¹²⁹I_{meas} = measured concentration of ¹²⁹I (μ g L⁻¹), 1.12 corrects for the presence of 12 % ¹²⁷I in the ¹²⁹I and (127/129) is a gravimetric correction. The corresponding correction was also applied to measurements of ¹²⁷I, according to Eqn. A2:

$${}^{127}I = {}^{127}I_{meas} - (0.12 \times {}^{129}I_{m})$$
(A2)

Where ¹²⁷I = 'corrected' concentration of ¹²⁷I in solution (μ g L⁻¹) i.e. ¹²⁷I minus the ¹²⁷I present as a result of the spiking, and ¹²⁷I_{meas} = measured concentration of ¹²⁷I (μ g L⁻¹).