

# Probing geomagnetic storm-driven magnetosphere-ionosphere dynamics in D-region via propagation characteristics of very low frequency radio signals

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## Abstract

The amplitude and phase of VLF/LF radio signals are sensitive to changes in electrical conductivity of the lower ionosphere which imprints its signature on the Earth-ionosphere waveguide. This characteristics makes it useful in studying sudden ionospheric disturbances, especially those related to prompt X-ray flux output from solar flares and gamma ray bursts (GRBs). However, strong geomagnetic disturbance and storm conditions are known to produce large and global ionospheric disturbances, which can significantly affect VLF radio propagation in the D region of the ionosphere. In this paper, using the data of three propagation paths at mid-latitudes ( $40^\circ - 54^\circ$ ), we analyze the trend of aspects of VLF diurnal signal under varying solar and geomagnetic space environmental conditions in order to identify possible geomagnetic footprints on the D region characteristics. We found that the trend of variations generally reflect the prevailing space weather conditions in various time scales. In particular, the ‘dipping’ of mid-day signal amplitude (MDP) of VLF always occurs after geomagnetic perturbed or storm conditions in the time scale of 1-2 days. The mean signal before sunrise (MBSR) and mean signal after sunset (MASS) also exhibit storm-induced dipping, but they appear to be influenced by event’s exact occurrence time and highly variable conditions of dusk-to-dawn ionosphere. We observed fewer cases of the signals rise (e.g., MDP, MBSR or MASS) following a significant geomagnetic event, though this effect may be related to storms associated phenomena or

effects arising from sources other than solar origin. The magnitude of induced dipping (or rise) significantly depends on the intensity and duration of event(s), as well as the propagation path of the signal. The post-storm day signal (following a main event, with lesser or significantly reduced geomagnetic activity), exhibited a tendency of recovery to pre-storm day level. In the present analysis, We do not see a well defined trend of the variations of the post-storm sunrise terminator (SRT) and sunset terminator (SST). The SRT and SST signals show more post-storm dipping in GQD-A118 propagation path but generally an increase along DHO-A118 propagation path. Thus the result could be propagation path dependent and detailed modeling is required to understand these phenomena.

*Keywords:* D-region ionosphere, Geomagnetic storm, Ionospheric response, magnetosphere-ionosphere dynamics, VLF radio signals

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## 1. Introduction

Although separated by thousands of kilometers, the magnetosphere and ionosphere are known to be physically connected through the Earth's magnetic field into one global system. The ionosphere responds to (a) prompt changes in solar energetic events, mainly the solar flare associated bursts in EUV, X-ray and relativistic particles (Mitra, 1974; Bounsanto, 1999; Alfonsi et al., 2008), (b) delayed changes mainly due to geomagnetic storm conditions with time scale from several hours to 1-3 days (Lastovika, 1996; Bounsanto, 1999; Kutiev, 2013), and (c) periodic changes with time scales of several days to months, and those of several solar cycles (Alfonsi, 2008; Kutiev, 2013). The ionosphere also exhibits diurnal (day/night) and seasonal (summer/winter) variations (Miller and Brace, 1969; Zhang et al., 1999). Solar and geomagnetic induced phenomena drive changes in magnetosphere conditions, whose coupling effects modify ionospheric signatures including atmospheric density distribution, total electron content (TEC), ionospheric current system, ionisation rates, and crucial D-region parameters such as conductivity gradient and reference height (Wait, 1959; Wait and Spies, 1964; Mitra, 1974; Buonsanto, 1999; Burke, 2000; Simoes et al., 2012; Nwankwo and Chakrabarti, 2014b). The dynamics of ionospheric response to changes in solar and geomagnetic conditions, involve the exchange of particles and electromagnetic energy (absorbed, reprocessed and deposited in the ionosphere by the magnetosphere) between magnetically connected regions (Burke, 2000;

23 Streltsov and Lotko, 2004; Goldstein et al., 2006; Russell et al., 2010; Russell  
24 and Wright, 2012 Leonard et al., 2012; Kutiev et al., 2013).

### 25 *1.1. The ionosphere at a glance*

26 The ionosphere is composed of three distinct space regions [D (50 km to  
27 90 km), E (90 km to 120 km), and the F (from 120 km up to 500 km), which  
28 often split into two layers, namely, F1 and F2]. Its existence is primarily  
29 due to ionisation by solar ultraviolet (UV) radiation and X-ray wavelength  
30 (Kelley, 1989; Prolss, 2004; McRae and Thomson, 2004; Raulin et al., 2006;  
31 Heikkila, 2011) and isotropic cosmic rays. Recombination also occurs when  
32 free electrons are captured by positive ions. Ionisation and recombination  
33 efficiency controls the overall electron density at every instant of time. The  
34 D region ionosphere highly active during the day (roughly between the local  
35 sunrise and sunset) due to high rate of ionisation, but its density fall signif-  
36 icantly at night largely due to rapid recombination at the altitude. The E  
37 region also maintains the same dynamics (night/day fluctuations) as the D  
38 region but ionisation state persists longer due to slower rate of recombination  
39 at lower density. Thus, the reflection of signals mainly occurs at the bottom  
40 of the nighttime E region (Han and Cummer, 2010a and references therein).  
41 The F region is present both day and night; air density and recombination  
42 rate is very low in the region. Therefore, ionisation persists in the nighttime  
43 (also see Mimno, 1937; Poole, 1999; Prolss, 2004). In general, these layers  
44 are severely disturbed by phenomena of solar and geomagnetic origin, as well  
45 as planetary and tidal waves, thermospheric tides and stratospheric warming  
46 (Pancheva et al., 2008; Leonard et al., 2012; Chen et al., 2013; Goncharenko  
47 et al., 2012; Polyakova et al., 2014). However, effects at different heights, lo-  
48 cations or latitudes vary in development, depending on time and intensity (of  
49 driving force). Ionospheric signature variations reflect different mechanisms  
50 and aspects of solar and other induced phenomena.

### 51 *1.2. VLF propagation in the Earth-ionosphere waveguide*

52 The velocity, direction and amplitude of most electromagnetic waves are  
53 distinctly affected when propagating through the ionosphere. This character-  
54 istics makes Radio waves one of the ideal tools for ionospheric study (Prolss,  
55 2004). Very low frequency (VLF) radio waves in the 3-30 kHz are effective  
56 in the investigation of solar induced variable conditions in the ionosphere  
57 (especially the D region) because their amplitude and phase are sensitive to  
58 changes in electrical conductivity of the lower ionosphere (Wait and Spies,

59 1964; Mitra, 1974; Alfonsi et al., 2008). VLF radio signals are reflected  
60 alternately by the D region and the Earth's surface due to high conductiv-  
61 ity (Mimno, 1937; Poole, 1999). The transmitted wave is thus guided be-  
62 tween the Earth and the ionosphere enabling the signal to propagate globally  
63 through the Earth-Ionosphere waveguide. The signal is then received at var-  
64 ious receivers across the world. Variations in daytime VLF signal amplitude  
65 and phase appear to be well correlated with solar X-ray output, with almost  
66 prompt responses. Hence, it has been used by many researchers to study  
67 sudden ionospheric disturbances and changes in the atmosphere (e.g., Araki,  
68 1974; Hayakawa et al., 1996; Molchanov et al., 1998; Kleimenova et al., 2004;  
69 McRae and Thomson, 2004; Thomas et al., 2004; Chakrabarti et al., 2005;  
70 Grubor et al., 2005; Peter et al., 2006; Sasmal et al., 2009; Chakrabarti et  
71 al., 2010; Clilverd et al., 2010; Basak et al., 2011; Pal et al., 2012; Palit et  
72 al., 2013; Ray et al., 2013; Raulin et al., 2013; Nwankwo and Chakrabarti,  
73 2014b). Other methods used for ionospheric studies include observational and  
74 experimental techniques and tools such as Global Navigation Satellite system  
75 (GNSS) receivers, vertical and oblique sounding, Riometers, incoherent scat-  
76 ter radars (e.g., EISCAT), coherent scatter radars (e.g., Goose Bay radar,  
77 SuperDARN), magnetometers, etc. (Greenwald et al., 1995, 1996; Honary  
78 et al., 1995; Lastovicka, 1996; Wild et al., 2003; Burke, 2000; Danilov and  
79 Lastovicka, 2001; Goldstein et al., 2005; Ruohoniemi and Greenwald, 2005;  
80 Alfonsi et al., 2008).

### 81 *1.3. VLF signal detection mechanism of sudden ionospheric disturbances*

82 The D region ionosphere is maintained by Lyman- $\alpha$  radiation at a wave-  
83 length of about 121.5nm, which ionises neutral nitric oxide (NO). With high  
84 solar activity, hard X-ray ( $\lambda < 1nm$ ) may ionise  $N_2$  and  $O_2$ . Galactic cosmic  
85 rays are also responsible for the ionisation of the lowest part of the lower  
86 ionosphere and the low-lying atmosphere down to the troposphere (also, see  
87 Mitra, 1974; Lastovicka, 1996). A huge amount of energy is released during  
88 solar flare in the form of highly energetic ultraviolet radiation, mainly X-ray  
89 flux enhancement. The radiation penetrates the D region where it increases  
90 ionisation rate (of dominant neutral NO molecules), and enhances electron  
91 density. These processes enhance the 'thickness' of the D region, thereby  
92 decreasing the reflection height ( $h$ ) in the waveguide. This is normally de-  
93 tected as a sudden change (usually an increase) in the amplitude and phase  
94 enhancement of a VLF signal. VLF dusk-to-dawn signal exhibit high vari-  
95 ability (or, fluctuation) during the night due to a significant fall in density

96 of the D region. The signal is also sensitive to phenomena other than those  
97 originating from the Sun. Day time VLF signal is primarily controlled by  
98 the Sun.

#### 99 *1.4. Geomagnetic induced variations of the ionosphere and effects*

100 Geomagnetic disturbances and storms are also known to produce signifi-  
101 cant global disturbances in the ionosphere, including the middle atmosphere  
102 and troposphere (Lastovika, 1996; Danilov and Lastovika 2001). Geomag-  
103 netic storms are the products of highly variable solar wind speeds and density  
104 and associated shock waves (Lastovika, 1986; Baker, 1996, 2000; Borovsky  
105 and Denton, 2006; Tsurutani et al., 2006; Kozyra et al., 2006). The ef-  
106 fects of geomagnetic storms on the ionosphere manifest mainly through en-  
107 ergetic particles precipitation, which lose their energy by impact and X-ray  
108 bremsstrahlung production (Lastovika, 1996). There is also a consequent and  
109 significant enhancement of electron density (Chenette et al., 1993; Stoker  
110 1993; Lastovika, 1996), causing significant increase in radio wave absorp-  
111 tion and subsequent disappearance of radio signals in MF/HF values (Las-  
112 tovika, 1996). Galactic cosmic ray flux (which are modulated by geomagnetic  
113 storms) and global electric circuit and atmosphere electricity (affected by lo-  
114 cal changes of conductivity and ionosphere/magnetosphere electric fields and  
115 currents), are assumed to be the processes for ionospheric effects of geomag-  
116 netic storms (Danilov and Lastovika, 2001). VLF signals can be significantly  
117 affected by geomagnetic disturbances and storms induced ionosphere per-  
118 turbations (Kikuchi and Evans, 1983). Nevertheless, a few researchers have  
119 used it to study these perturbations with insightful findings (e.g., Araki,  
120 1974; Kleimenova et al., 2004; Peter et al., 2006; Clilverd et al., 2010; Ku-  
121 mar and Kumar, 2014; Tatsuta et al., 2015).

122  
123 Apart from X-ray flux induced enhancement of amplitude and phase,  
124 anomalies in diurnal VLF signature may convey other important informa-  
125 tion, especially those related to geomagnetic disturbance or storm-induced  
126 ionospheric variations. If substantiated, such information could be instruc-  
127 tive and resourceful to the study and understanding of the complex dynamics  
128 of Earth's ionosphere. Thus, in addition to well correlated VLF signal am-  
129 plitude variation and phase enhancement with X-ray flux induced sudden  
130 ionospheric disturbances (SID), this work seeks to understand possible ge-  
131 omagnetic activity footprints in the D region of the ionosphere and their  
132 dependence on the propagation path of VLF radio waves. First, the analysis

133 concentrates on four selected periods of significant solar and geomagnetic  
134 activities in order of increasing magnitude, followed by a detailed statistical  
135 analysis of up to 16 storm conditions.

## 136 2. Data and method of analysis

137 In this work, analysed data mainly include diurnal VLF signal ampli-  
138 tude (of up to three propagation paths) monitored at A118 SID monitor-  
139 ing station in Southern France (<http://sidstation.loudet.org/data-en.xhtml>),  
140 GOES solar X-ray flux, average z-components ( $B_z$ ) and total magnetic field  
141 ( $H_T$ ) (<http://satdat.ngdc.noaa.gov/sem/goes/data/>), global geomagnetic  $A_p$   
142 (NOAA) and disturbance storm time (Dst) index (from World Data Centre  
143 for Geomagnetism (WDCG)), solar wind speed ( $V_{sw}$ ) and particle density  
144 ( $PD$ ) (<ftp://sohoftp.nascom.nasa.gov/sdb/goes/ace/>). Analysis was con-  
145 ducted over four different 6-day periods with different geomagnetic condi-  
146 tions of varying disturbance. The space condition during 14th-19th February  
147 2011 is recognised as moderately disturbed, the condition during 26th-31st  
148 May 2011 is recognised as a moderate storm, and condition during 24th-29th  
149 September and 23rd-28th October 2011 are recognised as relatively intense  
150 storm conditions. The choice of a six days time frame is to give us a rea-  
151 sonable time interval for analysis of data before, during and after the main  
152 event(s). The three propagation paths are shown in Figure 1 and include  
153 GQD-A118, ICV-A118, and DHO-A118; GQD (22.1 kHz GQD, lat N54.73°  
154 long W002.88°), ICV (20.27 kHz, lat N40.92° long E009.73°), DHO (23.4  
155 kHz, lat N53.08° long W007.61°).

### 156 2.1. Data description

157 A solar flare is ranked based on its X-ray output, and classified according  
158 to the order of magnitude of the peak burst intensity (I), measured at the  
159 Earth in 0.1 to 0.8 nm band,  $B = I < 10^{-6}W/m^2$ ,  $C = 10^{-6}I < 10^{-5}W/m^2$ ,  
160  $M = 10^{-5}I < 10^{-4}W/m^2$ ,  $X = 10^{-4}IW/m^2$ . We investigate solar wind speed  
161 conditions because the velocity, density, strength and direction of the solar  
162 wind plasma, and strength and direction of its associated magnetic field,  
163 influence the structure of the surrounding magnetic field of the Earth and  
164 controls the processes by which mass, momentum and energy are transferred  
165 from the solar wind to the Earth's magnetosphere-ionosphere system (Las-  
166 tovika, 1989; Singer et al., 1996). The  $B_z$  component significantly contributes  
167 to energy transfer from the solar wind sector to the magnetosphere (Prolss,

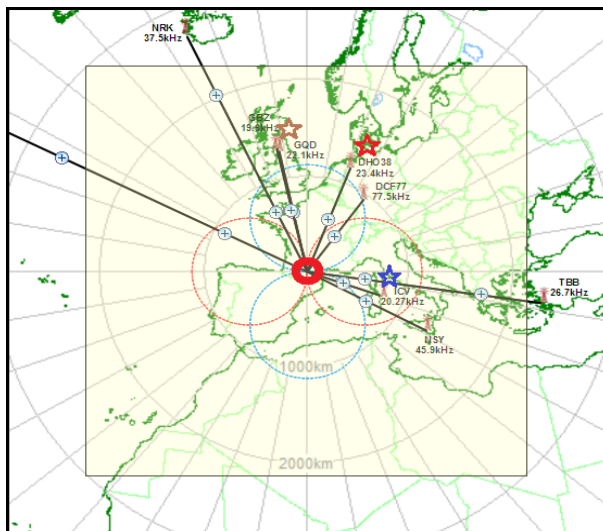


Figure 1: VLF signal propagation paths (PP) used in the study: A118 receiver (thick red circle), DHO transmitter (red star), GQD (brown star), ICV (blue star) [adopted from A118 SID station Web page]

168 2004).  $H_T$  data can be used to deduce and check solar wind influence on  
 169 the magnetosphere. Substorms advance and intensify current systems in the  
 170 magnetosphere and ionosphere, which can also be detected via  $H_T$  compo-  
 171 nent.  $A_p$  (or,  $K_p$ ) are planetary indices and are the indicators of geomag-  
 172 netic activity. The Dst is used to assess or measure the severity of magnetic  
 173 storms. The strength of the surface magnetic field is inversely proportional to  
 174 the energy content of the ring current, which increases during geomagnetic  
 175 storms (Hamilton et al., 1988). The solar wind condition and the men-  
 176 tioned geomagnetic parameters are important for studying and understand-  
 177 ing magnetosphere-ionosphere coupling and effects (Borovsky and Denton,  
 178 2006; Tsurutani et al., 2006; Kozyra et al., 2006; Weigel 2010; Nwankwo et  
 179 al., 2014, 2015). However, having provided a precise background of the pa-  
 180 rameters, we will concentrate mainly on how various aspects of diurnal VLF  
 181 signal varies in response to geomagnetic activity and storm footprints in the  
 182 D region ionosphere via these parameters, especially the Dst index. Details  
 183 of geomagnetic indices variation in response to solar wind conditions and  
 184 sources can be found in some literatures e.g Lastovika (1989), Tsurutani et  
 185 al. (1972, 1988, 1995, 1997, 2006, 2011), Baker (1996), Kozyra et al. (2006),

186 Weigel (2010) and references therein.

187

188 We analyse 2- to 4-hour Mean VLF signal amplitude before ‘local’ sun-  
189 rise and after sunset (hereafter respectively denoted as MBSR and MASS),  
190 and mid-day signal amplitude peak (MDP). We also identified variations in  
191 the so-called sunrise and sunset terminators (hereafter, denoted as SRT and  
192 SST). The aspects of a typical VLF signal (MBSR, MDP, MASS, SRT and  
193 SST) that were analysed are shown in Fig. 2 (a-d). In addition, daily so-  
194 lar flare count (for flares  $\geq C$ ) and the standard deviation or fluctuation of  
195 daily Dst were calculated. The main goal of the analysis is to investigate  
196 the trend in variations of these components under given solar and geomag-  
197 netic induced space environmental conditions, for possible identification of  
198 geomagnetic footprint in D-region ionosphere via the propagation character-  
199 istics of VLF signal, in addition to known X-ray flux induced prompt response  
200 of VLF amplitude and phase. Data were analysed for two signal propagation  
201 paths (PP) in each case. To begin with, we perform a detailed study of four  
202 particular cases, and then investigate the statistical significance of our results  
203 with more cases (up to 16).

### 204 3. Results and Discussion

205 Figure 3(a-h) shows diurnal VLF amplitude for GQD-A118 and ICV-  
206 A118 propagation paths, X-ray flux output, solar wind speed ( $V_{sw}$ ), particle  
207 density (PD),  $B_z$  magnetic field component,  $H_T$  magnetic field, daily Dst  
208 standard deviation and  $A_p$  variation during 14th-19th February 2011. The  
209 period is associated with high flare activity (up to 79 flares; C=69, M=9,  
210 X=1) and Dst variations of  $>-50$  (also see, Table 1). High flare events were  
211 observed on 14th, 16th and 18th (Fig. 3c), as well as significant geomag-  
212 netic activity on the 14th and 18th February (Fig. 3e-g). Highly variable  
213 solar wind speed ( $V_{sw}$ ) and associated magnetospheric impact (via  $B_z$  and  
214  $H_T$ ) were also observed from 06:00 pm, 14th - 12:00 noon, 15th and during  
215 most part of 18th February (Fig. 3d-f). The extent and severity of induced  
216 magnetospheric perturbations is highlighted by the Dst during late 14th and  
217 the considerable part of 18th (Fig. 3g). High  $A_p$  index of 18th February is  
218 therefore not surprising (Fig. 2h). VLF signal amplitude of the two propa-  
219 gation paths responded in a manner consistent with high flare events during  
220 the period. However, the flare-induced perturbations are distinct in VLF sig-  
221 nals (during local daytime), and appear to overshadow those of geomagnetic



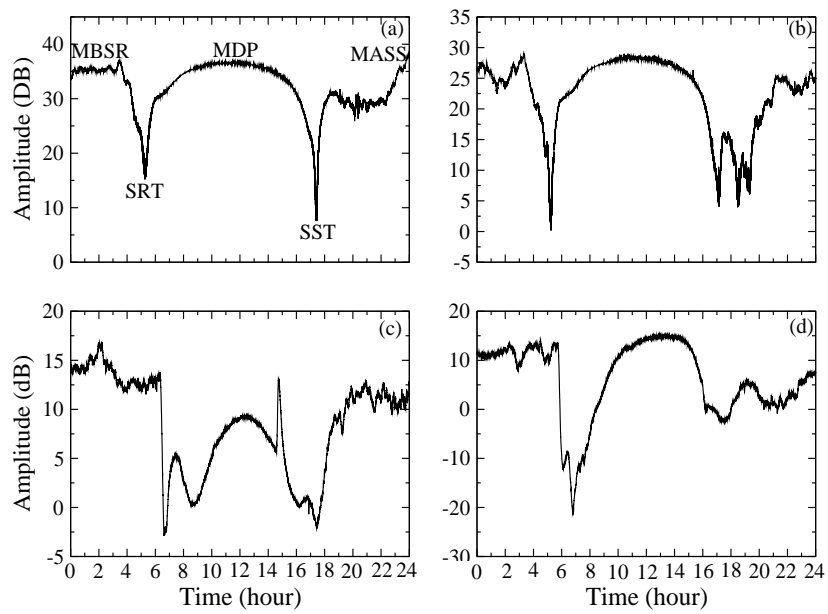


Figure 2: Diurnal signature of VLF signals from propagation paths showing various aspects as identified in (a).

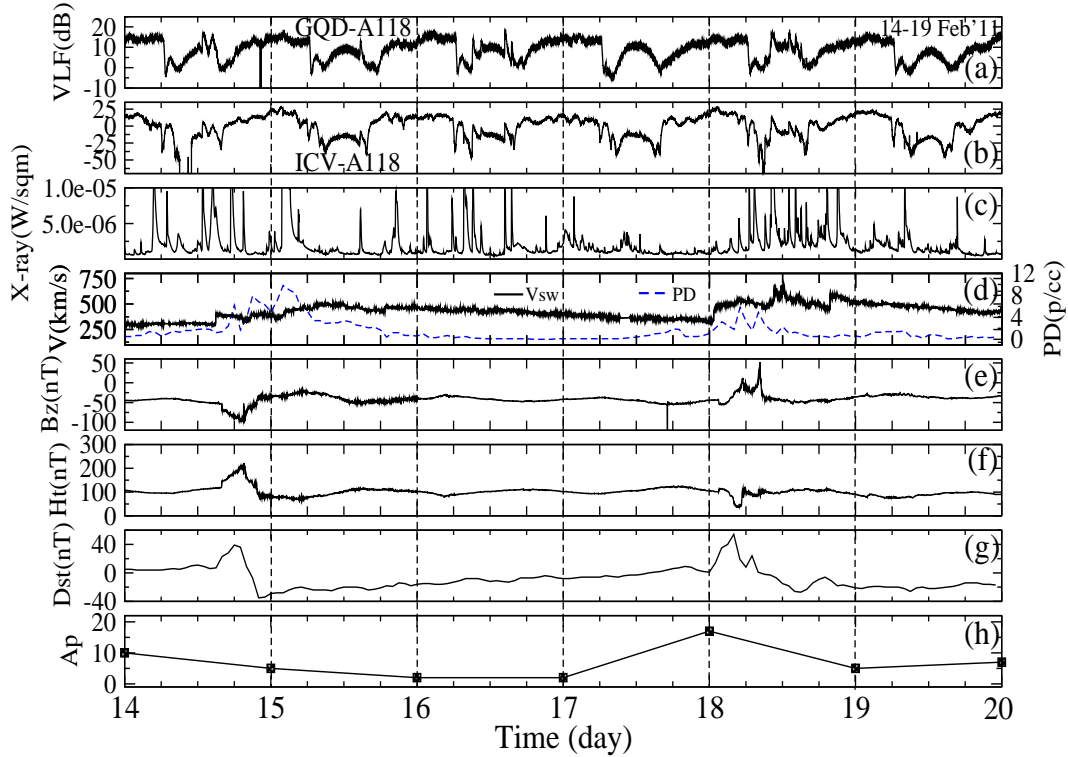


Figure 3: (a) Diurnal VLF amplitude for GQD-A118 PP; (b) VLF amplitude for ICV-A118 PP; (c) X-ray flux output; (d) solar wind speed ( $V_{sw}$ ) and particle density ( $PD$ ); (e)  $B_z$  magnetic field component; (f)  $H_T$  magnetic field; (g)  $Dst$  and (h)  $A_p$  variations during 14-19th February 2011.

222 activity origin. We therefore looked for the trend in the signal diurnal varia-  
 223 tions such as MBSR, MDP, MASS, SST and SRT, for possible separation of  
 224 distinct signatures of geomagnetic disturbance induced variations.

225

226 Figure 4 shows daily  $Dst$  standard deviation, 4-hour mean signal ampli-  
 227 tude before local sunrise (MBSR), mid-day signal peak (MDP), 4-hour mean  
 228 signal amplitude after sunset (MASS), variation in sunrise terminator (SRT)  
 229 and in sunset terminator (SST) for (a) GQD-A118 and (b) ICV-A118 prop-  
 230 agation paths during 14-19th February 2011. A summary of relative trend  
 231 in variations of the parameters over the period is provided in Table 1. Two  
 232 main geomagnetic disturbed days are the 14th (day 1) and the 18th (day 5)

233 presumably due to increase or spikes in solar wind speed ( $V_{sw}$ ) and parti-  
 234 cle density ( $PD$ ) (see, Fig. 3d). Proper analysis of a trend on a particular  
 235 day requires a comparison with the trend of the previous day and the day  
 236 after the event, because of the varying timescale of ionospheric response to  
 237 different aspects of solar forcing and mechanisms. Therefore, we consider  
 238 the trend of pre-event day in order to determine that of the event (s) day,  
 239 and also consider the post-event(s) day for extended effect. We observed  
 240 an increase in MBSR and SRT, but ‘dipping’ of MDP, MASS and SST on  
 241 15th (day 2) (Fig. 4a). Note the onset of perturbations on the 14th (day  
 242 1) - during and after sunset. The influence of the induced perturbations  
 243 are therefore expected to extend into a considerable part of 15th (day 2).  
 244 There was a quiet geomagnetic condition on the 16th (day 3), and almost all  
 245 the parameters increased. Of interest is the more (and longer) geomagnetic  
 246 disturbed condition on the 18th (day 5). Only the SST increased (during  
 247 which a decline in the initial induced perturbation was expected), while al-  
 248 most all other parameters (MBSR, MDP, MASS and SRT) experienced a  
 249 ‘dipping’. The observed trend is replicated in ICV-A118 propagation path  
 250 around 15th (day 2) but quite inconsistent on 18th (day 5) - mainly increase  
 251 of MBSR, MDP and MASS, but dipping of SRT and SST (Fig 4b). However,  
 252 the increase in MDP appeared to be related to flare induced signal amplitude  
 253 variation on the signal as well as high fluctuation in ICV-A118 propagation  
 254 path signal level, before and after sunset (see, Fig 3b).

255  
 256 Figure 5 shows the diurnal VLF signal amplitude variations for GQD-  
 257 A118 and ICV-A118 propagation paths, X-ray flux,  $V_{sw}$ ,  $PD$ ,  $B_z$ ,  $H_T$ , daily  
 258 Dst standard deviation and  $A_p$  variations during 26th-31st May 2011. Blue  
 259 and red lines in the Figure indicate the storm commencement and peak time,  
 260 respectively. The period is associated with moderate flare activity (up to 43;  
 261 C=41, M=2, X=0), as well as a moderate storm condition (Dst < -50 (up  
 262 to -91). The most disturbed days in this case are the 28th and the 29th  
 263 May, following a geomagnetic storm on the 28th (Fig. 5(c-h)). The geo-  
 264 magnetic storm of 28th February appears to be related to the sudden (and  
 265 significant) rise in  $V_{sw}$  and  $PD$ , possibly of coronal origin. Up to three CMEs  
 266 with the speed exceeding 1000 km/s occurred between 27th and 29th ([http :  
 267 //cdaw.gsfc.nasa.gov/CME.list/UNIVERSAL/2011\\_05/univ2011\\_05.html](http://cdaw.gsfc.nasa.gov/CME.list/UNIVERSAL/2011_05/univ2011_05.html)).  
 268 Solar wind density influences the capability of a given value of the solar wind  
 269 electric field (SWEF) to create a  $Dst$  disturbance or geo-efficiency (Weigel,  
 270 2010; Tsurutani et al., 2011; Nwankwo et al., 2016). Also, solar flares and

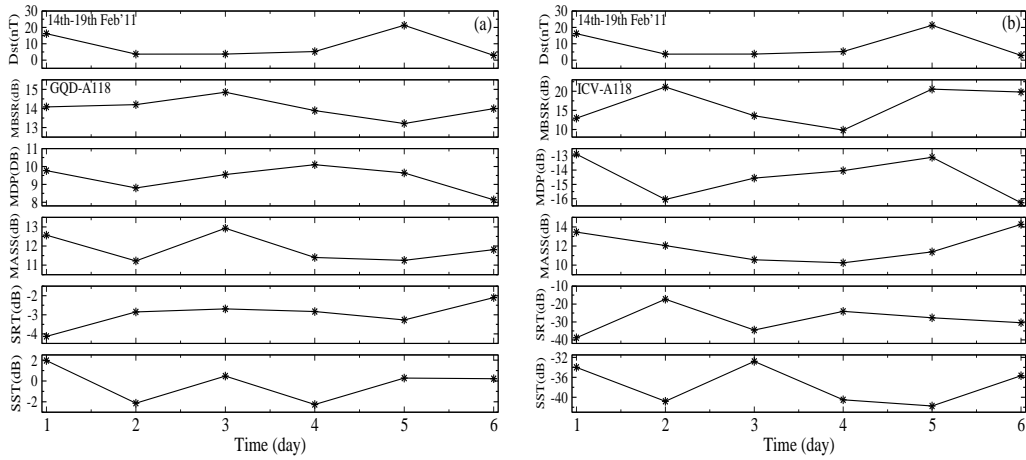


Figure 4: Daily Dst standard deviation, 4-hour mean signal amplitude before sunrise (MBSR), mid-day signal peak (MDP), 4-hour mean signal amplitude after sunset (MASS), sunrise terminator (SRT) and sunset terminator (SST) variations for (a) GQD-A118 and (b) ICV-A118 propagation path during 14-19th February 2011.

271 prominence eruptions are known independent and sporadic events, but they  
 272 do also occur in association with coronal mass ejections (CMEs). However,  
 273 we do not strictly attribute the solar wind and magnetosphere conditions  
 274 during this period to CMEs because of limited scope of our analysis in this  
 275 regard. In Fig. 5(a-c), we observed that with relatively high flare activity  
 276 around 28th-29th May, the known diurnal (daytime) signal amplitude-spike  
 277 in response to solar X-ray output in both propagation paths tend to be di-  
 278 minished under geomagnetic storm condition when compared with 14th-19th  
 279 February scenario (Fig. 5a-b). This situation is replicated in the other three  
 280 storm conditions investigated alongside.

281

282 Figure 6 shows daily Dst standard deviation, 2-hour mean MBSR, MDP,  
 283 2-hour mean MASS, SRT and SST variations for (a) GQD-A118 and (b)  
 284 ICV-A118 propagation paths during 26th-31st May 2011. A summary of  
 285 trend in variation of the parameters over the period is provided in Table 2.  
 286 Our main focus here is on 28th (day 3), being the most disturbed, as well as  
 287 the storm day. We observed an increase in MBSR, MDP and MASS, but a  
 288 dipping of SRT and SST in GQD-A118 propagation path (Fig. 6a). Notwith-  
 289 standing, dipping of the MBSR and MDP occurred on the day following the

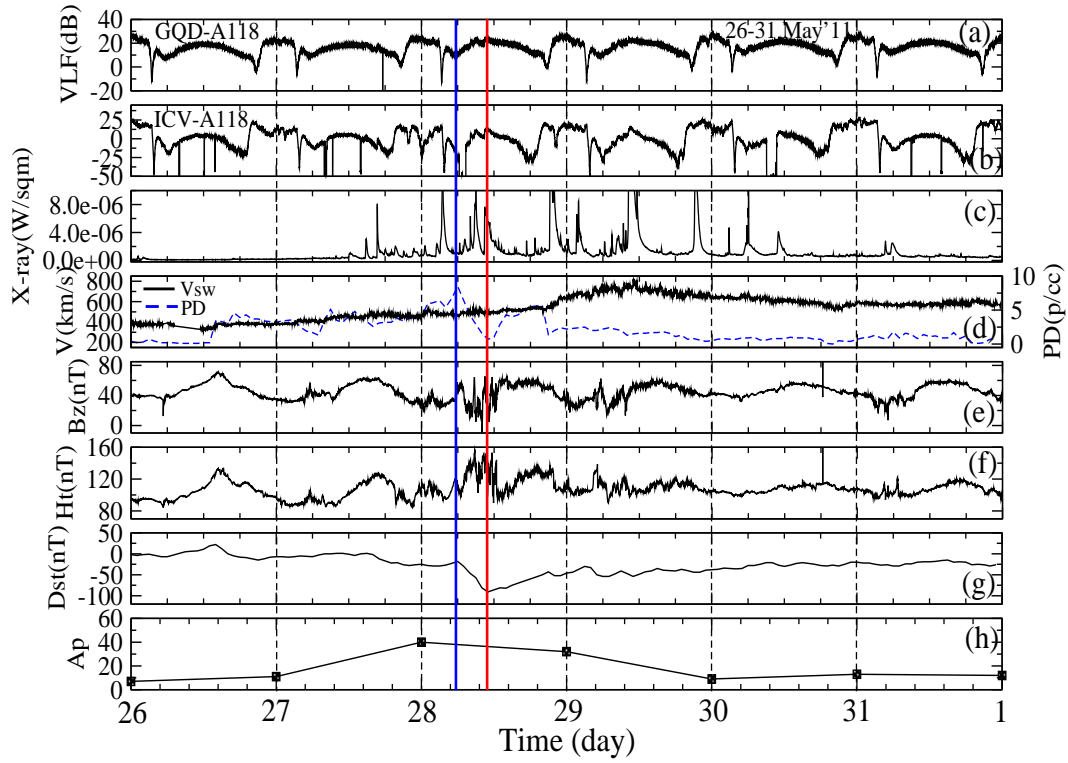


Figure 5: (a) Diurnal VLF amplitude for GQD-A118 PP; (b) VLF amplitude for ICV-A118 PP; (c) X-ray flux output; (d) solar wind speed ( $V_{sw}$ ) and particle density ( $PD$ ); (e)  $B_z$  magnetic field component; (f)  $H_T$  magnetic field; (g)  $Dst$  and (h)  $A_p$  variations during 26th-31st May 2011 (Blue and red lines in the Figure indicate storm commencement and peak time respectively)

Table 1: Trend of time variation of VLF amplitude, Dst and flare count during 15-18th February 2011 for GQD-A118 and ICV-A118 propagation path

GQD-A118 propagation path								
Date	Mean Signal peak (dB)			Signal dip (dB)		Dst (nT)	Flare count	
	MBSR	MDP	MASS	SRT	SST	$\sigma_{Dst}$	$\geq C$	C M X
14/2/11	14.08±0.78	9.77	12.57±2.18	-4.13	1.96	±16.19	12	11 1 0
15/2/11	14.20±1.15	8.80	11.22±0.72	-2.85	-2.13	±3.67	8	7 0 1
16/2/11	14.85±1.07	9.55	12.93±0.95	-2.69	0.47	±3.71	15	12 3 0
17/2/11	13.89±1.14	10.10	11.40±0.82	-2.83	-2.26	±5.27	12	12 0 0
18/2/11	13.21±0.90	9.64	11.25±1.09	-3.27	0.28	±21.29	20	15 5 0
19/2/11	13.99±1.10	8.14	11.81±2.23	-2.10	0.22	±2.90	12	12 0 0
ICV-A118 propagation path								
14/2/11	12.95±3.82	-12.89	13.46±3.40	-38.82	-33.99	±16.19	12	11 1 0
15/2/11	21.11±3.11	-16.05	12.05±4.17	-17.30	-40.80	±3.67	8	7 0 1
16/2/11	13.60±2.38	-14.56	10.56±3.49	-34.52	-32.80	±3.71	15	12 3 0
17/2/11	9.83±3.81	-14.04	10.24±2.57	-24.08	-40.50	±5.27	12	12 0 0
18/2/11	20.56±3.24	-13.11	11.39±3.95	-27.65	-41.75	±21.29	20	15 5 0
19/2/11	19.81±1.25	-16.28	14.26±3.88	-30.42	-35.67	±2.90	12	12 0 0

290 storm day (moderate but significantly disturbed 29th (day 2)). In ICV-A118  
 291 propagation path, the MASS increased slightly while MBSR, MDP, SRT and  
 292 SST dipped with high Dst (Fig. 6b). It is important to note that we had  
 293 to take a two hour mean due to increase in day length. Also note the spike  
 294 in MDP due to the possible influence of the flare particularly in GQD-A118  
 295 propagation path on 28th (dipping need to be large or significant to nullify  
 296 flare-induced influence). Understandably, geomagnetic effects are also not  
 297 expected on any portion of the signal (e.g., MBSR, MDP, MASS, SRT, SST)  
 298 before significant geomagnetic perturbations. The increase (MDP) could also  
 299 be due to the propagation characteristics of ICV-A118 propagation path, be-  
 300 cause mode interference significantly depends on ionospheric conditions at  
 301 the time, propagation paths and energetic electron precipitation level on the  
 302 ionosphere due to the magnetic storm, which depends on geomagnetic lati-  
 303 tude (Tatsuta et al., 2015).

304  
 305 Figure 7 shows the diurnal VLF amplitude variations for GQD-A118 and  
 306 DHO-A118 propagation paths, X-ray flux,  $V_{sw}$ ,  $PD$ ,  $B_z$ ,  $H_T$ , daily Dst stan-  
 307 dard deviation and  $A_p$  variations during 24th-29th September 2011. The

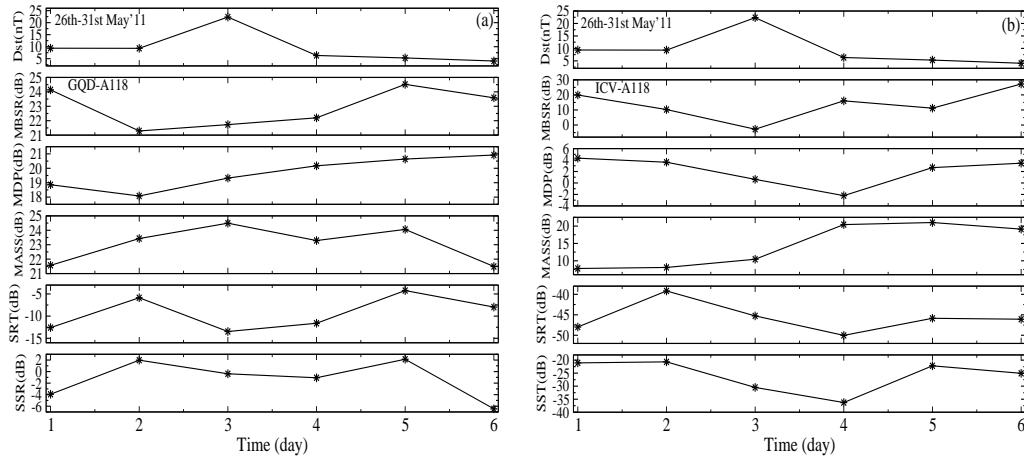


Figure 6: Daily Dst standard deviation, two-hour mean signal amplitude before sunrise (MBSR), mid-day signal peak (MDP), two-hour mean signal amplitude after sunset (MASS), sunrise terminator (SRT) and sunset terminator (SST) variations for (a) GQD-A118 and (b) ICV-A118 propagation path during 26th-31st May 2011.

308 period is associated with relatively high flare events (up to 51; C=33, M=17,  
 309 X=1) and intense storm conditions with  $Dst \leq -100$ . The unique feature of  
 310 the period is the associated sub-storm of late 26th (red line) following the  
 311 storm condition that commenced before noon with peak (broken red line),  
 312 which also marked the sub-storm commencement (Fig. 7e-g). Milder storm  
 313 conditions also occurred on 28th and 29th. The storm-driving high variable  
 314 solar wind (and  $PD$  spike) is clearly observed in Fig. 6d. Dipping of DHO-  
 315 A118 propagation path daytime (and MDP) signal on 26th is clearly visible  
 316 in Fig. 7b, with the post storm day signal (with lesser geomagnetic indices  
 317 and/or disturbances) on 27th exhibiting a tendency of recovery (or return)  
 318 to pre-storm level. The trend of variations of MBSR, MDP, MASS, SRT and  
 319 SST have also shown similar tendency.

320

321 Figure 8 shows daily Dst standard deviation, 4-hour mean MBSR, MDP,  
 322 4-hour mean MASS, SRT and SST variations for (a) GQD-A118 and (b)  
 323 DHO-A118 propagation paths during 24th-29th September 2011. Summary  
 324 of the trend in variation of the parameters over the period is provided in  
 325 Table 3. In GQD-A118 propagation path signal, dipping of MDP, SRT and  
 326 SST were observed on 26th (day 3), while MBSR and MASS increased (Fig.

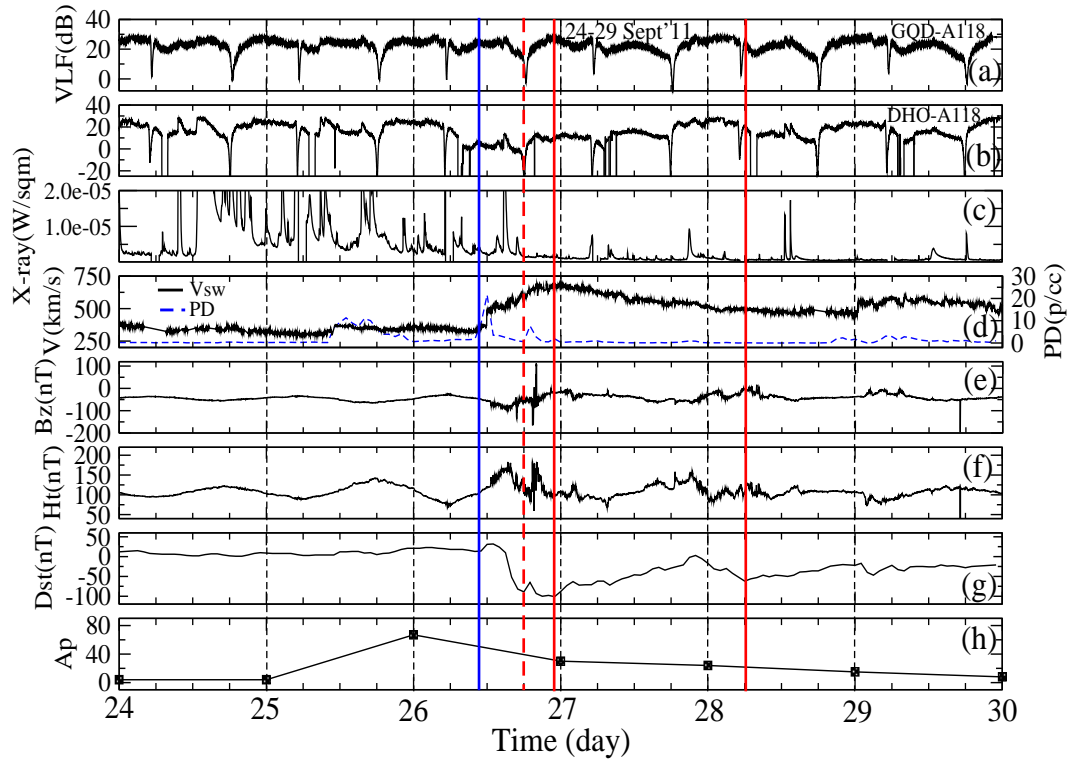


Figure 7: (a) Diurnal VLF amplitude for GQD-A118 PP (b) Diurnal VLF amplitude for DHO-A118 PP (c) X-ray flux output (d) solar wind speed ( $V_{sw}$ ) and particle density ( $PD$ ) (d)  $B_z$  magnetic field component (e)  $H_T$  magnetic field (f) Dst and (g)  $A_p$  variations during 24th-29th September 2011 (Blue and red lines in the Figure indicate storm commencement and peak time respectively)



Table 2: Trend of time variation of VLF amplitude, Dst standard deviation and flare count during 26-31st May 2011 for GQD-A118 and ICV-A118 propagation path.

GQD-A118 propagation path								
Date	Mean Signal peak (dB)			Signal dip (dB)		Dst (nT)	Flare count	
	BSR	Mid-day	ASS	SRT	SST	$\sigma_{Dst}$	$\geq C$	C M X
26/5/11	24.14±1.24	18.86	21.57±1.01	-12.59	-3.93	±9.37	0	0 0 0
27/5/11	21.29±1.05	18.08	23.43±0.65	-5.86	1.98	±9.31	5	5 0 0
28/5/11	21.73±1.00	19.32	24.49±1.22	-13.47	-0.38	±22.33	19	18 1 0
29/5/11	22.20±1.42	20.17	23.29±1.63	-11.60	-1.07	±6.35	13	12 1 0
30/5/11	24.52±1.74	20.64	24.06±1.07	-4.24	2.14	±5.31	4	4 0 0
31/5/11	23.59±2.14	20.92	19.11±4.10	-7.75	-6.46	±4.04	2	2 0 0
ICV-A118 propagation path								
26/5/11	19.92±4.32	4.33	7.79±2.62	-47.18	-21.05	±9.37	0	0 0 0
27/5/11	10.26±4.32	3.62	8.08±8.74	-39.18	-20.66	±9.31	5	5 0 0
28/5/11	-2.74±8.39	0.63	10.44±9.05	-45.27	-30.47	±22.33	19	18 1 0
29/5/11	16.07±2.28	-2.21	20.42±3.17	-50.02	-36.28	±6.35	13	12 1 0
30/5/11	11.19±2.94	2.68	21.02±3.28	-45.85	-22.17	±5.31	4	4 0 0
31/5/11	22.21±3.83	3.45	19.11±4.10	-46.08	-25.07	±4.04	2	2 0 0

327 8a). It is important to note that the peak of the geomagnetic storms-induced  
328 perturbations on the ionosphere, which commenced during the later part of  
329 26th are expected into greater part of 27th. As could be seen in Fig. 7g,  
330 the Dst recovery during 27th is associated with momentary perturbations,  
331 followed by the sub-storm commencement at 06:00 pm on that day. Further  
332 dippings of MBSR, MDP, MASS and SST were also observed on 27th (day 4;  
333 see Fig 8a). Thereafter, the MBSR, MDP and MASS increased with reduced  
334 Dst on the 28th. Notwithstanding, storm conditions were also recorded on  
335 the 28th and 29th, the perturbations are not comparable to those of 26th-  
336 27th. In DHO-A118 propagation path, dipping of the MDP, MASS and SST  
337 were observed on the 26th (day 3) and 28th (day 4; see Fig 8b). On the  
338 other hand, there is a relative increase in MBSR and SRT on the days (3 and  
339 4). While the trends in the two propagation paths appear to significantly re-  
340 flect the space weather conditions, the dipping or increase of the signal varied.

341  
342 Figure 9 shows the diurnal VLF amplitude variations for GQD-A118 and  
343 DHO-A118 propagation paths, X-ray flux,  $V_{sw}$ ,  $PD$ ,  $B_z$ ,  $H_T$ , daily Dst stan-  
344 dard deviation and  $A_p$  variations during 23rd-28th October 2011. This period

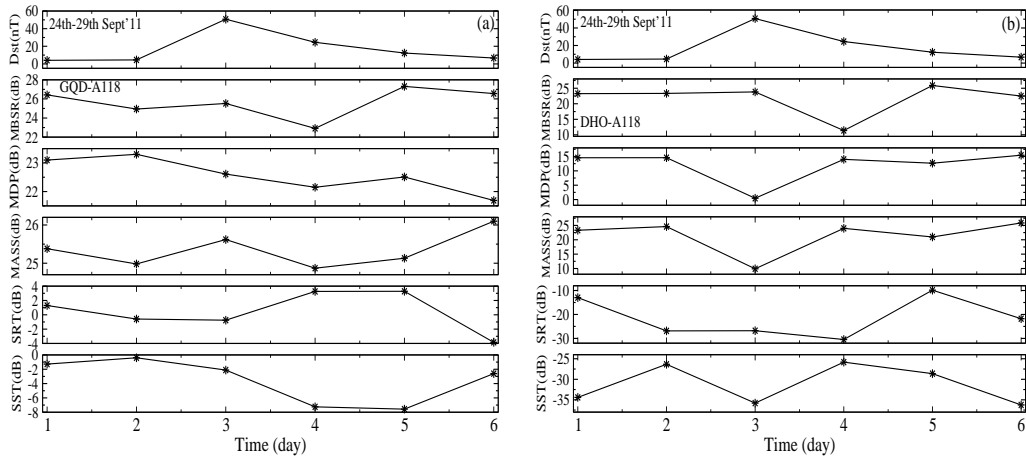


Figure 8: Daily Dst standard deviation, 4-hour mean signal amplitude before sunrise (MBSR), mid-day signal peak (MDP), 4-hour mean signal amplitude after sunset (MASS), sunrise terminator (SRT) and sunset terminator (SST) variations for (a) GQD-A118 and (b) DHO-A118 propagation path during 24th-29th September 2011.

345 is associated with relatively low flare activity (only 11 C class flares), but  
 346 with an intense storm condition of higher magnitude (Dst < -100 (down to -  
 347 132)). The storm occurred in the early hours of 25th, which commenced late  
 348 24th (around 06:00 pm), presumably due to high speed solar wind (HSS)  
 349 and *PD* condition of 24th October (Fig 9d-h). VLF signal data for GQD-  
 350 A118 propagation path during 12:00 noon, 25th - 06:00 pm, 26th October  
 351 (Fig. 9a) are not available. It is worth mentioning that only DHO-A118  
 352 propagation path (at A118 SID receiving station) recorded data during this  
 353 time interval. Data of about 6 other propagation paths (e.g., GBZ-A118,  
 354 ICV-A118, NAA-A118, TBB-A118) in the series are also not available (see,  
 355 Fig. 1 for PP identification). As this time interval probably corresponds  
 356 to the peak period of induced ionosphere perturbations, it will be interest-  
 357 ing to further investigate possible cause of the scenario (beyond the scope  
 358 of this work), with respect to the prevailing geomagnetic condition. Again,  
 359 dipping of DHO-A118 propagation path daytime and MDP signal on 25th  
 360 (most disturbed day) is clearly visible (Fig. 9b), with the post storm day  
 361 signal exhibiting a drop or recovery to pre-storm level.

362  
 363

Figure 10 shows daily Dst standard deviation, 4-hour mean MBSR, MDP,

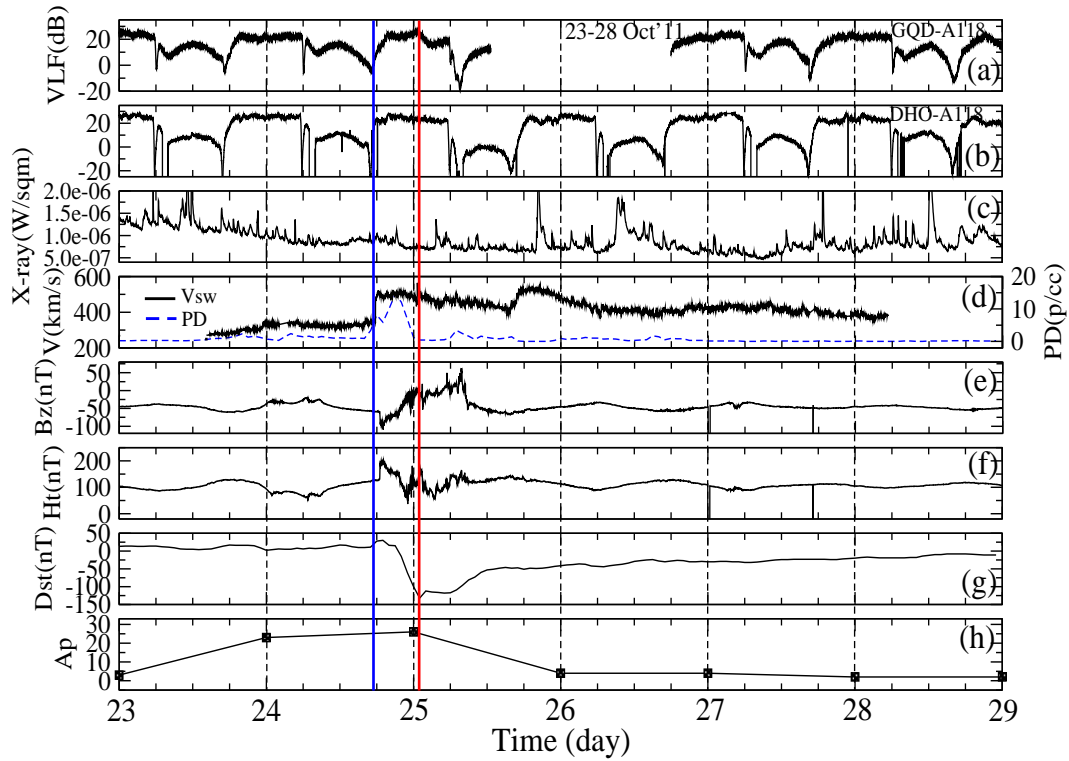


Figure 9: (a) Diurnal VLF amplitude for GQD-A118 PP (b) Diurnal VLF amplitude for DHO-A118 PP (c) X-ray flux output (d) solar wind speed (d)  $B_z$  magnetic field component (e)  $H_T$  magnetic field (f) Dst and (g)  $A_p$  variations during 23rd-28th October 2011

Table 3: Trend of time variation of VLF amplitude, Dst and flare count during 25th-28th September 2011 for GQD-A118 and DHO-A118 propagation path.

GQD-A118 propagation path								
Date	Mean Signal peak (dB)			Signal dip (dB)		Dst (nT)	Flare count	
	BSR	Mid-day	ASS	SRT	SST	$\sigma_{Dst}$	$\geq C$	C M X
24/9/11	26.42±1.02	23.10	25.38±2.10	1.30	-1.28	±4.08	13	4 8 1
25/9/11	24.94±1.16	23.30	24.98±0.96	-0.59	-0.40	±4.56	10	4 6 0
26/9/11	25.52±1.14	22.61	25.62±1.59	-0.75	-2.11	±50.73	11	9 2 0
27/9/11	22.91±1.35	22.15	24.87±1.63	-3.26	-7.25	±24.54	8	8 0 0
28/9/11	27.31±0.77	22.51	25.13±1.38	3.28	-7.57	±12.37	4	3 1 0
29/9/11	26.56±1.29	21.69	26.10±2.32	-3.85	-2.61	±6.73	3	3 0 0
DHO-A118 propagation path								
24/9/11	23.26±2.04	14.55	23.32±1.00	-12.96	-34.41	±4.08	13	4 8 1
25/9/11	23.33±1.29	14.57	24.60±0.99	-26.86	-26.34	±4.56	10	4 6 0
26/9/11	23.81±1.05	0.45	9.90±1.48	-26.79	-35.80	±50.73	11	9 2 0
27/9/11	11.38±1.05	14.00	23.68±1.90	-30.47	-25.82	±24.54	8	8 0 0
28/9/11	25.90±1.74	12.66	20.98±2.09	-9.85	-28.62	±12.37	4	3 1 0
29/9/11	22.49±2.04	15.43	25.87±3.31	-21.78	-36.25	±6.73	3	3 0 0

364 4-hour mean MASS, SRT and SST variations for (a) GQD-A118 and (b)  
365 DHO-A118 propagation paths during 23rd-28th October 2011. Summary of  
366 trend in variation of the parameters over the period is provided in Table 4.  
367 GQD-A118 propagation path data during 25th and 26th is inadequate for  
368 the present analysis (Fig. 10a). The DHO-A118 propagation path signal  
369 showed dipping of the MBSR, MDP and MASS on 25th (day 3), correspond-  
370 ing to the storm's peak day, but an increase in SRT and SST (Fig 10a). The  
371 prevailing space weather conditions (with peak) of 25th (day 3) commenced  
372 at around 06:00 pm on 24th (day 2). Interestingly, dipping of the MDP and  
373 MASS also commenced on 24th (day 2). There is a post-storm day increase  
374 of MBSR, MDP and MASS with significant Dst low on 26th, a scenario that  
375 is characteristic of most post-storm day signals. We, therefore viewed such  
376 scenario as post-storm day signal recovery tendency.

377

378 We now identify the most disturbed day in each of the four periods, and  
379 analyse the trend in the signal metrics variation on the day, namely, event 1  
380 ( $E_1$ ) on 18th February, 2011; event 2 ( $E_2$ ) on 28th May, 2011; event 3 ( $E_3$ )  
381 on 26-27 September, 2011; and event 4 ( $E_4$ ) on 25th October 2011. Due to

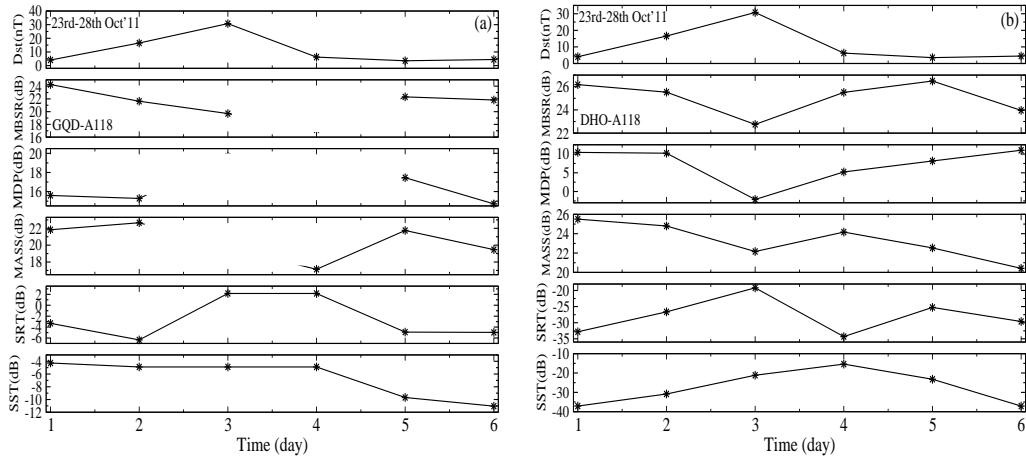


Figure 10: Daily Dst standard deviation, 4-hour mean signal amplitude before sunrise (MBSR), mid-day signal peak (MDP), 4-hour mean signal amplitude after sunset (MASS), sunrise terminator (SRT) and sunset terminator (SST) variations for (a) GQD-A118 and (b) DHO-A118 propagation path during 23rd-28th October 2011.

382 the peculiarity of the events during 26th-27th September, 2011 (recurrent  
383 substorm), two days have been allowed for the analysis. In general, two  
384 of three events ( $E_{1-3}$ ) showed dipping of MDP in GQD-A118 propagation  
385 path (VLF data during  $E_4$  is not available). Three of the four events ( $E_{1-4}$ )  
386 showed dipping of MDP in ICV/DHO-A118 propagation paths. We note  
387 that solar flare occurred around mid-day in the days when MDP showed no  
388 dipping. This suggests possible flare induced increase of signal amplitude  
389 on the MDP or resulting from other atmospheric phenomena. Two of four  
390 events ( $E_{1-4}$ ) showed dipping of MBSR in GQD-A118 propagation path, and  
391 dipping in all the four events in ICV/DHO-A118 propagation paths. Two  
392 of three events ( $E_{1-3}$ ) showed dipping of MASS in GQD-A118 propagation  
393 path (VLF data during  $E_4$  is not available), and two of the four events in  
394 ICV/DHO-A118 propagation path. Three of the four events showed dipping  
395 of SRT in GQD-A118 propagation path, and two of the four in ICV/DHO-  
396 A118 propagation paths. Two of the four events showed dipping of SST in  
397 GQD-A118 propagation path, and three of the four in ICV/DHO-A118 prop-  
398 agation paths. We have also observed that within the local day time interval  
399 (24 hours), the events occurred well before or after four of five MBSR and  
400 MASS, and five of six SRT and SST that showed no dipping (or, maintained

Table 4: Trend of time variation of VLF amplitude, Dst and flare count during 23rd-28th October 2011 for GQD-A118 and DHO-A118 propagation path

GQD-A118 propagation path								
Date	Mean Signal peak (dB)			Signal dip (dB)		Dst (nT)	Flare count	
	BSR	Mid-day	ASS	SRT	SST	$\sigma_{Dst}$	$\geq C$	C M X
23/10/11	24.35±0.88	16.59	21.83±0.87	-3.31	-4.27	±4.08	3	3 0 0
24/10/11	21.63±1.02	15.28	22.66±0.93	-6.35	-4.89	±16.35	0	0 0 0
25/10/11	19.70±3.77	-	-	2.16	-	±30.76	1	0 0 0
26/10/11	17.14±2.59	-	-	-	-	±6.25	1	1 0 0
27/10/11	22.32±1.43	17.45	21.74±1.33	-4.92	-9.69	±3.53	1	1 0 0
28/10/11	21.83±0.86	19.35	19.47±2.52	-4.97	-11.98	±4.48	5	5 0 0
DHO-A118 propagation path								
23/10/11	26.18±1.05	10.45	25.51±0.82	-32.81	-37.10	±4.08	3	3 0 0
24/10/11	25.53±0.92	10.23	24.80±1.33	-26.64	-30.84	±16.35	0	0 0 0
25/10/11	22.75±0.99	-2.12	22.16±1.68	-19.19	-21.17	±30.76	1	1 0 0
26/10/11	25.51±1.22	5.23	24.17±1.18	-34.30	-15.40	±6.25	1	1 0 0
27/10/11	26.49±1.72	8.16	22.53±4.45	-25.25	-23.23	±3.53	1	1 0 0
28/10/11	23.96±1.68	11.02	20.42±1.32	-29.63	-37.10	±4.48	5	5 0 0

401 amplitude) in accordance with the events. Among other possible inferences,  
402 this trend suggest that geomagnetic effects are not expected on any aspect of  
403 the signal (e.g., MBSR, MDP, MASS, SRT, SST) before significant geomag-  
404 netic perturbations, and if the event occurs well before the component, the  
405 induced ionospheric perturbations is expected to have significantly reduced at  
406 the time interval. Of the three propagation paths, the signal of DHO-A118  
407 appears to be the most sensitive to geomagnetic induced magnetosphere-  
408 ionospheric dynamics. However, given the few number of the cases analysed  
409 so far, drawing a firm conclusion would be difficult at this stage. Therefore,  
410 we include more cases in the next analysis (see Table 4), and combine differ-  
411 ent signal aspects on a single graph for a better view of the trends.

412

413 We analyse and study the trend in variations of combined signal aspects  
414 for 16 storm cases (Dst=-50 to -132) between February 2011 and June 2012  
415 for two propagation paths (GQD-A118 and DHO-A118). Details of the storm  
416 events are provided in Table 4. Analysis include taking (a) signal metrics  
417 (MBSR, MDP, MASS, SRT and SST) 1-day before an event (BE), during  
418 an event (DE) and after an event (AE), and (b) a 2-day mean signal metric

Table 5: Summary of analysed geomagnetic storm conditions

No.	Date	Max Dst (nT)	$\sigma_{Dst}$	Flare count( $\geq C$ )
				C M X
1	05022011	-51	$\pm 8.99$	0 0 0
2	01032011	-81	$\pm 36.28$	7 0 0
3	06042011	-65	$\pm 24.31$	3 0 0
4	12042011	-51	$\pm 22.11$	3 0 0
5	26092011	-101	$\pm 50.73$	9 2 0
6	25102011	-132	$\pm 30.76$	1 0 0
7	22012012	-67	$\pm 37.00$	4 0 0
8	15022012	-58	$\pm 9.63$	0 0 0
9	19022012	-54	$\pm 12.8$	1 0 0
10	07032012	-74	$\pm 25.41$	1 0 0
11	15032012	-74	$\pm 20.75$	1 0 0
12	28032012	-55	$\pm 12.09$	1 0 0
13	05042012	-54	$\pm 13.82$	3 0 0
14	23042012	-95	$\pm 32.23$	3 0 0
15	12062012	-51	$\pm 12.47$	13 0 0
16	16062012	95	$\pm 20.24$	4 0 0
17*	17062012	80	$\pm 46.75$	7 0 0

419 BE, DE and AE. An event is selected based on factors such as availability  
420 and quality of VLF signal data on the day, and relatively quiet BE and AE,  
421 particularly for the 2-day mean analysis. Although BE and AE data were  
422 carefully chosen to be consistent with relative geomagnetic quiet condition,  
423 a few choices on significantly perturbed days were unavoidable due to inter-  
424 vals of extended geomagnetic active condition and recurrent storms. This  
425 scenario can cause high variability of VLF radio signal. Other than solar  
426 induced fluctuations, the ionosphere and VLF radio signal also response to  
427 effects originating from a number of other sources (see Section 1.1). Some  
428 of the effects are interconnected (with possible interference), leading to a  
429 high variability of signal strength. Therefore, a ‘perfect’ consistency in trend  
430 across all the cases are not expected. Figure 11 shows Dst deviation (fluctu-  
431 ation) and trend in variation of signals MDP, MBSR, MASS, SRT and SST  
432 one day before and after (successive) each of the 16 selected storm conditions  
433 for (a) GQD-A118 and (b) DHO-A118 propagation paths. Detail of the data  
434 is provided in appendix I.

435

436 For GQD-A118 propagation path, 10 of 14 MDP, 10 of 15 MBSR, 7 of

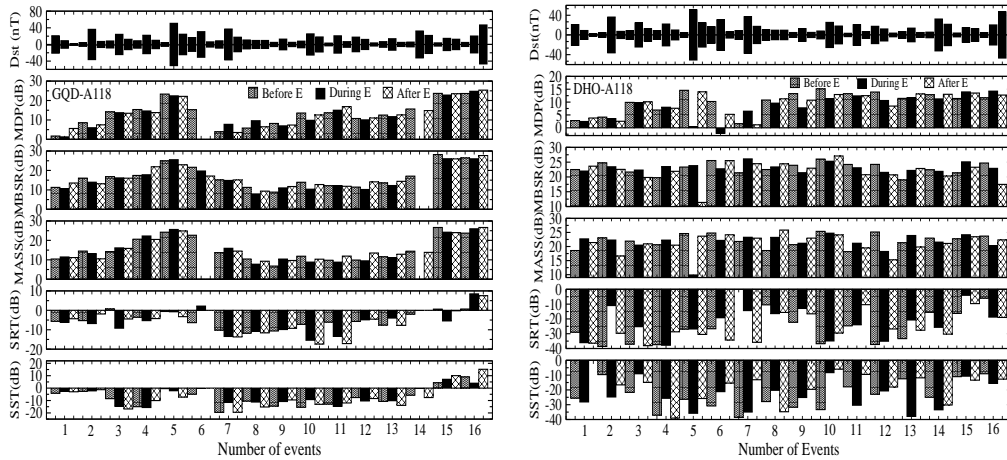


Figure 11: Daily Dst deviation and trend in variation of MDP, MBSR, MASS, SRT and SST signals one day before and after each of the 16 selected storm conditions for (a) GQD-A118 and (b) DHO-A118 propagation paths. A '0' indicate absence of data.

437 14 MASS, 9 of 14 SRT and 7 of 14 SST have shown a dipping of the signals.  
 438 These correspond respectively to 71.4%, 66.7%, 50%, 64.3% and 50.0% of  
 439 the combined cases. In DHO-A118 propagation path 13 of 16 MDP, 9 of  
 440 16 MBSR, 8 of 16 MASS, 5 of 14 SRT and 7 of 16 SST showed dipping  
 441 of the signals. These correspond respectively to 81.3%, 56.3%, 50%, 35.7%  
 442 and 43.8% of the combined cases. Note that dipping of any of DE and AE  
 443 signal metric in cases 15 and 16 is taken as a response to the event because  
 444 storm condition or the event commenced during late DE and peaked in AE.  
 445 Also, recurrent storms occurred on the day after case 16. Whereas majority  
 446 of MDP in both the propagation paths have shown a notable evidence of  
 447 dipping, few number of PP-mismatched incidences of MDP signal rise (or,  
 448 increase) on some events day have been observed (e.g., events 8, 11 and 16  
 449 in GQD and 4 and 13 in DHO). The increase may be related to flare induced  
 450 signal amplitude spike on the signal or phenomena arising from sources other  
 451 than storm events. We also observed a notable matched-increase of the diur-  
 452 nal signal level (including MDP, MBSR and MASS) on DE 7 (22 Jan 2012) in  
 453 both propagation paths. While further investigation is vital to accurate in-  
 454 terpretation, a closer look at the available data showed occurrence of storm  
 455 associated M-class flare with corresponding peaks, suggesting an enhance-  
 456 ment of not only the instantaneous but also background X-ray flux output.



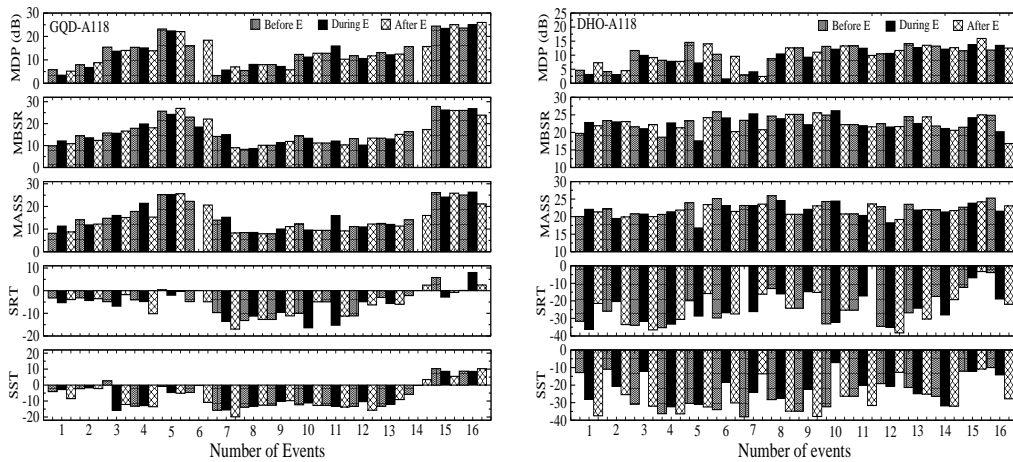


Figure 12: Daily Dst deviation (fluctuation) and trend in variation of 2-day mean MDP, MBSR, MASS, SRT and SST before, during and after an event for (a) GQD-A118 and (b) DHO-A118 propagation paths. A '0' indicate absence of data

457 Figure 12 shows Dst deviation (fluctuation) and trend in variation of 2-day  
 458 mean MDP, MBSR, MASS, SRT and SST signals before, during and after  
 459 each event for (a) GQD-A118 and (b) DHO-A118 propagation paths. Details  
 460 of the data is provided in Appendix II. Using a different criterion for data  
 461 selection, the analysis presented in Fig. 12 is a follow up on the one pre-  
 462 sented in Fig. 11, and expected to provide resourceful clue towards a better  
 463 conclusion of the results. Whereas BE, DE and AE represent data of three  
 464 consecutive days with reference to the event's day (DE) in the former anal-  
 465 ysis (Fig 11), each acronym (BE, DE or AE) represent a 2-day mean (VLF)  
 466 with respect to DE (but not necessarily in succession to DE). Besides data  
 467 availability and quality, an important data selection criterion is a relative  
 468 geomagnetic quiet BE- and AE-day with respect to DE - hence, a one or  
 469 more days gap before or after DE (in some cases).

470  
 471 For GQD-A118 propagation path, 10 of 14 MDP, 9 of 15 MBSR, 7 of 14  
 472 MASS, 11 of 16 SRT and 5 of 14 SST showed dipping of the signals. These  
 473 correspond respectively to 71.4%, 60.0%, 50.0%, 68.8% and 35.7% of the  
 474 combined cases. For DHO-A118 propagation path, 11 of 16 MDP, 11 of 16  
 475 MBSR, 10 of 16 MASS, 6 of 14 SRT and 7 of 16 SST showed dipping of the  
 476 signals, corresponding respectively to 68.8%, 68.8%, 62%, 42.9% and 43.8%

477 of the combined cases. In general, MDP signal has shown a high probability  
478 of a dipping scenario following significant geomagnetic disturbance or storm  
479 condition. The MBSR and MASS signals have also shown good probability  
480 of exhibiting such storm-induced dipping, but appear to be influenced by  
481 event's occurrence time and the highly variable conditions of dusk-to-dawn  
482 ionosphere. However, a few cases have shown a rise or increase of the com-  
483 ponents instead (e.g., MDP, MBSR, MASS) following a significant geomag-  
484 netic event. We speculate that such a scenario (signal rise) may be related to  
485 storm associated phenomena or of sources other than solar origin rather than  
486 being a case against the 'favoured' dipping - this need be studied further. In  
487 contrast, the SRT and SST signals have shown significant post-storm dipping  
488 in GQD-A118 propagation path but mostly increase in DHO-A118 propaga-  
489 tion path. Does the trend in post-storm SRT and SST variation depend on  
490 signal propagation path? This important question may not be conclusively  
491 answered based on this present analysis. Thus, a clear dependence of SRT  
492 and SST on geomagnetic disturbance or storm conditions seems inconclusive.

493  
494 We consider it to be important to highlight the constraints associated  
495 with this analysis that may have also influenced our results and findings.  
496 Besides the flare and X-ray flux induced amplitude variation (see, Fig 2c),  
497 the daytime diurnal signal between SRT and SST of VLF radio waves are  
498 generally quite stable. No doubt, their stability has contributed to the con-  
499 sistency of MDP trend in the overall pattern of the results - the combined  
500 analysis showed about 73% dipping of the MDP. On the other hand, high  
501 variability or fluctuation of dusk-to-dawn signal (see, Fig. 2a-d) remain a  
502 major drawback to analysis relating to MBSR and MASS - the combined  
503 analysis showed 63% and 53% dipping of the MBSR and MASS, respectively.  
504 Similarly, the pseudo-SRT and SST (occurrence of double or multiple-tipped  
505 sunrise and/or sunset terminator) exhibited by diurnal VLF signal also ham-  
506 pers proper analysis of the signals - the combined analysis showed 52% and  
507 43% dipping of the SRT and SST, respectively. Deciding which of the tips  
508 to measure (in case of a pseudo-SRT/SST) would be more important but  
509 challenging. Nevertheless, a proper study which probes the cause of such  
510 fluctuations and occurrence of pseudo-terminators in VLF signature will be  
511 highly valuable. Such a study in addition to further investigating the ob-  
512 served interesting propagation paths (matched and mismatched) signal-rise  
513 during some cases of geomagnetic storm conditions have been initiated. This  
514 is beyond the scope of the present work and will be published elsewhere in

515 due course.

#### 516 4. Summary and Conclusion

517 The characteristic response of diurnal VLF signal to space weather in-  
518 duced ionospheric disturbances vary from one propagation path to another,  
519 and also depend on location of the transmitters and receivers, ionisation and  
520 chemistry of the D region over the propagation path, and the intensity of in-  
521 duced perturbations. Other influencing factors include signal frequency and  
522 nature of Earth's surface (also see, Mimno, 1937; Poole, 1999; Melia, 2010).  
523 In principle, known strong perturbations from solar flares and gamma-ray  
524 bursts of VLF signals can be reproduced from ab-initio calculations (Palit  
525 et al. 2013). In this paper, we used various aspect of diurnal VLF signal  
526 (such as MBSR, MDP, MASS, SRT and SST) to investigate the footprint of  
527 geomagnetic activity in D layer ionosphere at mid-latitude ( $40^{\circ}$ - $54^{\circ}$ ) region,  
528 under varying degree of sixteen storm conditions (and consequent distur-  
529 bances). Although the strength of diurnal signals significantly varied from  
530 one propagation path to another, the trend of variations of the characteristic  
531 signal appear to reflect the prevailing space weather conditions of various time  
532 scales. We found a significant dipping of the mid-day amplitude peak (MDP)  
533 of the signal within 1-2 days of significant geomagnetic disturbance or storm  
534 conditions. The MBSR and MASS signals have also generally shown such  
535 storm-induced dipping. However, they appear to be influenced by events'  
536 occurrence time and highly variable condition of dusk-to-dawn ionosphere.  
537 We observed a fewer cases of rise of the signals (e.g., MDP, MBSR or MASS)  
538 following a significant geomagnetic event. However, this may be related to  
539 storm-associated events or due to effects arising from sources other than so-  
540 lar origin. The extent of the induced dipping (or, rise) significantly depends  
541 on the intensity and duration of event(s), as well as the propagation path of  
542 the signal. The post-storm day signal (following a main event, with lesser or  
543 significantly reduced geomagnetic activity), exhibited a tendency of recovery  
544 to pre-storm day level. In the present analysis, the post-storm SRT and SST  
545 variations do not appear to have a well defined trend - the SRT and SST  
546 signals have shown more post-storm dipping in GQD-A118 propagation path  
547 but mostly increase in DHO-A118 propagation path.

548

549 Many researchers have investigated and reported ionospheric and VLF  
550 signal anomalies before seismic events (e.g., Hayakawa et al., 2010; Ray

551 and Chakrabarti, 2013; Sasmal et al., 2014). Such anomalies were often  
552 attributed to seismicity and therefore viewed as pre-cursors. However, in  
553 order to ensure that such VLF anomalies are indeed due to seismic events, it  
554 is imperative that other possible and potential drivers of ionospheric anoma-  
555 lies around intervening period are investigated, identified and separated. In  
556 future, we will investigate possible solar and geomagnetic-induced perturba-  
557 tions of the ionosphere within the time frame in which ionospheric precursor  
558 (using VLF signal) were reported. This must be taken into consideration  
559 before marking anomalies as pre-cursors. For this two prong approach is  
560 necessary: (i) to reproduce propagation path dependent effects on VLF sig-  
561 nals due to number of specific types of solar induced perturbations as in Palit  
562 et al. (2013) and (ii) to find statistical correlations among various quanti-  
563 ties using data for longer duration. The work is in progress and would be  
564 published elsewhere.

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## 573 **References**

- 574 Alfonsi, L., Kavanagh, A. J., Amata, E., Cilliers, P., Correia, E., Free-  
575 man, M., Kauristie, K., Liu, R. Y., Luntama, J. P., Mitchell, C. N. and  
576 Zhrebtsov, G. A., 2008. Probing the high latitude ionosphere from ground-  
577 based observations: The state of current knowledge and capabilities during  
578 IPY (2007-2009). *Journal of Atmospheric and Solar-Terrestrial Physics*, 70  
579 (18), pp. 2293-2308.
- 580 Araki T., 1974. Anomalous Phase Changes of Trans equatorial VLF Radio  
581 Waves during Geomagnetic Storms. *Journal of Geophysical Research*, 79,  
582 4811-4813.

- 583 Baker D. N., 1996. Solar wind-magnetosphere drivers of space weather. *J.*  
584 *Atm. Solar-Terres. Phys.*, 58, 15OW526.
- 585 Baker D. N., 2000. Effects of the Sun on the Earth's environment. *J. Atm.*  
586 *Solar-Terres. Phys.*, 62, 1669-1681.
- 587 Basak T., S. Pal and S. K. Chakrabarti, 2011. VLF study of Ionospheric  
588 properties during solar flares of varied intensity for a fixed propagation  
589 path. General Assembly and Scientific Symposium, 2011 XXXth URSI,  
590 13-20 Aug. 2011, doi: 10.1109/URSIGASS.2011.6051004.
- 591 Borovsky J. E. and Denton M. H., 2006. Differences between CME-driven  
592 storms and CIR-driven storms'. *Journal of Geophysical Research*, 111.
- 593 Burke W. J., 2000. Magnetosphere-ionosphere coupling: selected topics.  
594 *Journal of Atmospheric and Solar-Terrestrial Physics* 62, 817-824.
- 595 Bucha, V., Bucha Jr. V., 1998. Geomagnetic forcing of changes in climate and  
596 in the atmospheric circulation. *J. Atmos. Sol-Terr. Phys.* 60 (2), 145-169.
- 597 Buonsanto M. J., 1999. Ionospheric Storms - A Review, *Space Science Re-*  
598 *views*, 88, 563-601, 1999
- 599 Chakrabarti S. K., M. Saha, R. Khan, Mandal, S., Acharyya, K., and Saha,  
600 R., 2005. Possible Detection of Ionospheric Disturbances during Sumatra-  
601 Andaman Islands. *IJRS Phys.*, 34, 314-317.
- 602 Chakrabarti, S. K., Sasmal, S., and Chakrabarti, S., 2010. Ionospheric  
603 anomaly due to seismic activities - Part 2: Evidence from D-layer prepa-  
604 ration and disappearance times. *Nat. Haz. Earth Syst. Sci.*, 10, 1751-1757.
- 605 Chen C. H., C. H. Lin, L. C. Chang, J. D. Huba, J. T. Lin, A. Saito, and J.  
606 Y. Liu, 2013. Thermospheric tidal effects on the ionospheric midlatitude  
607 summer nighttime anomaly using SAMI3 and TIEGCM, *J. Geophys. Res.*  
608 *Space Physics*, 118, 38363845, doi:10.1002/jgra.50340.
- 609 Chenette D.L., D. W. Datlowe, R. M. Robinson, T. L. Schumaker, R. R. Von-  
610 drak, and J. D. Winningham, 1993. Atmospheric energy input and ioniza-  
611 tion by energetic electrons during the geomagnetic storm of 8-9 November  
612 1991, *Geophys. Res. Lett.*, 20, 1323.

- 613 Clilverd M. A., Rodger C. J., Gamble R. J., Ulich T., Raita T., Seppala  
614 A., Green J. C., Thomson N. R., Sauvaud J. A., and Parrot M., 2010.  
615 Ground-based estimates of outer radiation belt energetic electron precip-  
616 itation fluxes into the atmosphere. *Journal of Geophysical Research*, 115,  
617 A12304.
- 618 Cranmer S. R., 2009. Coronal Holes. *Living Rev. Solar Phys.*, 6, 3.
- 619 Danilov A. D. and Lastovicka J., 2001. Effects of Geomagnetic Storms on the  
620 Ionosphere and Atmosphere, *International Journal of Geomagnetism and*  
621 *Aeronomy*, 2, 209-224.
- 622 Goldstein J., Burch J. L., Sandel B. L., Mende S. B., P. C:son Brandt  
623 and Hairston M. R., 2005. Coupled response of the inner magnetosphere  
624 and ionosphere on 17 April 2002, *Journal of Geophysical Research*, 110,  
625 A03205, DOI:10.1029/2004JA010712
- 626 Goncharenko, L. P., A. J. Coster, R. A. Plumb, and D. I. V. Domeisen, 2012.  
627 The potential role of stratospheric ozone in the stratosphere-ionosphere  
628 coupling during stratospheric warmings, *Geophys. Res. Lett.*, 39, L08101,  
629 doi:10.1029/2012GL051261
- 630 Greenwald, R. A., K. B. Baker, J. R. Dudeney, M. Pinnock, T. B. Jones,  
631 E. C. Thomas, J.-P. Villain, J.-C. Cerisier, C. Senior, C. Hanuise, R. D.  
632 Hunsucker, G. Sofko, J. Koehler, E. Nielsen, R. Pellinen, A. D. M. Walker,  
633 N. Sato, and H. Yamagishi, 1995. Darn/Superdarn: A Global View of  
634 the Dynamics of High-Latitude Convection, *Space Sci. Rev.*, 71, 761796.  
635 doi:10.1007/BF00751350
- 636 Greenwald R. A., J. M. Ruohoniemi, W. A. Bristow, G. J. Sofko, J.-P. Vil-  
637 lain, A. Huuskonen, S. Kokubun, and, L. A. Frank, 1996. Mesoscale day-  
638 side convection vortices and their relation to substorm phase, *Journal of*  
639 *Geophysical Research*, 101, 21,697-21,713.
- 640 Grubor D., Sulic D., and Zigman V., 2005. Influence of solar X-ray flares on  
641 the Earth-ionosphere waveguide. *Serbian Astronomical J*, 171, 29-35.
- 642 Hamilton D. C., Gloeckler G., Ipavich F. M., Stdemann W., Wilken B. and  
643 Kremser G., 1988. Ring current development during the great geomagnetic  
644 storm of February 1986'. *Journal of Geophysical Research: Space Physics*,  
645 93, pp 14343-14355.

- 646 Han F. and S. A. Cummer, 2010. Midlatitude nighttime D region ionosphere  
647 variability on hourly to monthly time scales. *J. Geophys. Res.*, 115, A09323,  
648 doi:10.1029/2010JA015437.
- 649 Hayakawa M, O. A. Molchanov, Ondoh T. and Kawai E., 1996. The precur-  
650 sory signature effect of Kobe earthquake on VLF subionospheric signals.  
651 *Phys. Earth Planet. Inter.*, 57, pp 64-67.
- 652 Heikkila W., 2011. *Earth's Magnetosphere*. Elsevier Kidlington, Oxford, UK.
- 653 Hejda P. and Bochncek J., 2005. Geomagnetically induced pipe-to-soil volt-  
654 ages in the Czech oil pipelines during October/November 2003. *Annales*  
655 *Geophysicae*, 23, 3089-3093.
- 656 Honary F, Stocker A. J., Robinson T. R. and Jones T. B., 1995. Ionospheric  
657 plasma response to HF radio waves operating at frequencies close to the  
658 third harmonic of electron gyrofrequency. *J. Phys. Res.*, 100, 21489-21501.
- 659 Kelley M. C., 1989. *The Earth's Ionosphere*. Academic Press Inc. San Diego,  
660 California.
- 661 Kikuchi T. and Evans D.S., 1983. Quantitative study of substorm-associated  
662 VLF phase anomalies and precipitating energetic electrons on November  
663 13, 1979. *J. Geophys. Res., Space Phys.*, 88, 871-880.
- 664 Kleimenova N. G., Kozyreva O. V., Rozhnoy A. A. and Soloveva M. S., 2004.  
665 Variations in the VLF signal parameters on the Australia-Kamchatka radio  
666 path during magnetic storms. *Geomagn Aeron* 44, 385-393.
- 667 Kozyra J. U, G. Crowley, B. A. Emery, X. Fang, G. Maris, M. G. Mlynczak,  
668 R. J. Niciejewski, S. E. Palo, L. J. Paxton, C. E. Randall, P.-P. Rong, J.  
669 M. Russell III, W. Skinner, S. C. Solomon, E. R. Talaat, Q. Wu and J.-H.  
670 Yee, 2006. Response of the Upper/Middle Atmosphere to Coronal Holes  
671 and Powerful High-Speed Solar Wind Streams in 2003. Recurrent Magnetic  
672 Storms: Corotating Solar Wind Streams. Geophysical Monograph 167.  
673 Edited by Bruce Tsurutani, Robert McPherron, Walter Gonzalez, Gang  
674 Lu, Jos H. A. Sobral and Natchimuthukonar Gopalswamy. ISBN-13: 978-  
675 0-87590-432-0. AGU Books Board, AGU, Washington, DC USA, 319.
- 676 Kumar A. and Kumar S., 2014. Space weather effects on the low latitude  
677 D-region ionosphere during solar minimum. *Earth, Planets and Space*, 66.

- 678 Kutiev I., Ioanna Tsagouri, Loredana Perrone, Dora Pancheva, Plamen  
679 Mukhtarov, Andrei Mikhailov, Jan Lastovicka, Norbert Jakowski, Dalia  
680 Buresova, Estefania Blanch, Borislav Andonov, David Altadill, Sergio  
681 Magdaleno, Mario Parisi and Joan Miquel Torta, 2013. Solar activity im-  
682 pact on the Earths upper atmosphere. *Journal of Space Weather and Space*  
683 *Climate*, 3.
- 684 Lastovicka J., 1989. Solar wind and high energy particle effects in the middle  
685 atmosphere, *Handb. MAP*, 29, 119.
- 686 Lastovicka J., 1996. Effects of geomagnetic storms in the lower ionosphere,  
687 middle atmosphere and troposphere, *Journal of Atmospheric and Terres-*  
688 *trial Physics*, 58, 831-843.
- 689 Leonard J. M., Forbes J., and Born G. H., 2012. Impact of tidal density  
690 variability on orbital and reentry predictions, *J. Geophys. Res.: Space*  
691 *Weather*, 10, S12003, doi:10.1029/2012SW000842
- 692 McRae W. M. and Thomson N. R., 2004. Solar flare induced iono-  
693 spheric D-region enhancements from VLF phase and amplitude observa-  
694 tions, *Journal of Atmospheric and Solar-Terrestrial Physics*, 66, 77-87.  
695 DOI:10.1016/j.jastp.2003.09.009
- 696 Melia A., 2010. Flare Detection using VLF Radio Signal, G3NYK.
- 697 Miller N. J. and Brace L. H., 1969. Some winter characteristics of the northern  
698 high-latitude ionosphere, *Journal of Geophysical Research*, 74, 5752-5762.
- 699 Mimno H. R., 1937. The Physics of the Ionosphere, *Reviews of Modern*  
700 *Physics*, 9, 1-45.
- 701 Mitra W. B., 1974. Ionospheric effects of solar flares. D. Reidel Publishing  
702 Company, Dordrecht, Holland.
- 703 Molchanov O. A. and Hayakawa M., 1998. On the generation mechanism of  
704 ULF seimogenic electromagnetic emissions, *Phys. Earth Planet. Int.*, 105,  
705 201-220.
- 706 Nwankwo V. U. J and Chakrabarti K. S., 2014. Theoretical Modeling of  
707 Drag Force Impact on a Model International Space Station (ISS) during  
708 Variation of Solar Activity. *Trans. JSASS Aerospace Technology Japan* ,  
709 12, pp 47-53



- 710 Nwankwo V. U. J. and Chakrabarti S. K., 2014. A Probe of magnetosphere-  
711 ionosphere coupling using Very Low Frequency (VLF) Radio Signal from  
712 North-West Cape (Australia) to Kolkata (India). 40th COSPAR Scientific  
713 Assembly. Held 2-10 August 2014, in Moscow, Russia, C0.4-21-14.
- 714 Nwankwo V. U. J., Sandip K. Chakrabarti and Weigel R. S., 2015. Effects of  
715 Plasma Drag on Low Earth Orbiting Satellites due to solar forcing induced  
716 perturbations and Heating. *Adv. Space Res.*, 56, 47-56.
- 717 Nwankwo V. U. J., Chakrabarti S. K., Sasmal S. and Ray S., 2016.  
718 Possible influence of solar activity on some seismically induced pre-  
719 cursors through Magnetosphere-ionosphere coupling. *J. Geophys. Res.*,  
720 2016JA022468 (submitted)
- 721 Pal S., Surya K. Maji and Sandip K. Chakrabarti, 2012. First ever VLF  
722 monitoring of the lunar occultation of a solar flare during the 2010 an-  
723 nular solar eclipse and its effects on the D-region electron density profile,  
724 *Planetary and Space Science*, 73, 310-317.
- 725 Palit S., Basak T., S. K. Mondal, S. Pal, and S. K. Chakrabarti, 2013.  
726 Modeling of very low frequency (VLF) radio wave signal profile due to solar  
727 flares using the GEANT4 Monte Carlo simulation coupled with ionospheric  
728 chemistry, *Atmos. Chem. Phys.*, 13, 9159-9168.
- 729 Pancheva D. V., P.J. Mukhtarov, N.J. Mitchell, D.C. Fritts, D.M. Riggin,  
730 H. Takahashi, P.P. Batista, B.R. Clemesha, S. Gurubaran, G. Ramku-  
731 mar, 2008. Planetary wave coupling (56-day waves) in the low-latitude  
732 atmosphere-ionosphere system. *J. Atm. Solar-Terr. Phys.*, 70, 101-122.
- 733 Peter W. B., Chevalier M. W., and Inan U. S., 2006. Perturbations of mid-  
734 latitude sub-ionospheric VLF signals associated with lower ionospheric dis-  
735 turbances during major geomagnetic storms. *Journal of Geophysical Re-*  
736 *search*, 111, AO3301.
- 737 Polyakov A. S., M.A. Chernigovskay, N.P. Perevalov, 2014. Ionospheric effects  
738 of sudden stratosphere warmings in Eastern Siberian region. *J. Atm. Solar-*  
739 *Terr. Phys.*, 120, 15-23.
- 740 Poole I., 1999. *Radio Waves and the Ionosphere*. American Radio Relay  
741 League, G3YWX.

- 742 Prolss G. W., 2004. Physics of the Earth's space environment. Springer Berlin  
743 Heidelberg, Germany.
- 744 Raulin, J.-P., Pacini, A.A., Kaufmann, P., Correia, E., Martinez, M.A.G.,  
745 2010. On the detectability of solar X-ray flares using very low frequency  
746 sudden phase anomalies. Journal of Atmospheric and Solar-Terrestrial  
747 Physics 68, 1029-1035.
- 748 Raulin J-P., Fernando C. P. Bertoni, Hernan R. Gaviln, Walter Guevara-Day,  
749 Rodolfo Rodriguez, Germn Fernandez, Emilia Correia, Pierre Kaufmann,  
750 Alessandra Pacini, Tardelli R. C. Stekel, Washington L. C. Lima, Nelson J.  
751 Schuch, Paulo R. Fagundes and Rubens Hadano, 2010. Solar flare detection  
752 sensitivity using the South America VLF Network (SAVNET). Journal of  
753 Geophysical Research, 115, A07301. DOI: 10.1029/2009JA015154
- 754 Raulin J-P, Grard Trottet, Matthieu Kretzschmar, Edith L. Macotela,  
755 Alessandra Pacini, Fernando C. P. Bertoni and Ingolf E. Dammasch,  
756 2013. Response of the low ionosphere to X-ray and Lyman- solar  
757 flare emissions. Journal of Geophysical Research, 118, 570-575. DOI:  
758 10.1029/2012JA017916
- 759 Ray S. and Chakrabarti S. K., 2012. A study of the behavior of the terminator  
760 time shifts using multiple VLF propagation paths during the Pakistan  
761 earthquake (M=7.2) of 18 January 2011. Nat. Hazards Earth Syst. Sci.,  
762 13, pp 1501-1506.
- 763 Ruohoniemi J.M., and Greenwald R.A., 2005. Dependencies of high-latitude  
764 plasma convection: consideration of interplanetary magnetic field, sea-  
765 sonal, and universal time factors in statistical patterns. Journal of Geo-  
766 physical Research, 110, A09204, doi:10.1029/2004JA010815
- 767 Russell, A. J. B., A. N. Wright, and A. W. Hood, 2010. Self-consistent  
768 ionospheric plasma density modifications by field-aligned currents: Steady  
769 state solutions. J. Geophys. Res., 115, A04216, doi:10.1029/2009JA014836.
- 770 Russell, A. J. B., and A. N. Wright, 2012. Magnetosphere-ionosphere waves.  
771 J. Geophys. Res., 117, A01202, 2012. doi:10.1029/2011JA016950
- 772 Sasmal S. and S. K. Chakrabarti, 2009. Ionospheric anomaly due to seismic  
773 activities Part 1: Calibration of the VLF signal of VTX 18.2 KHz station

- 774 from Kolkata and deviation during seismic events. *Nat. Hazards Earth*  
775 *Syst. Sci.*, 9, 1403-1408.
- 776 Simoes F., Pfaff R., Berthelier J. and Klenzing, 2012. A Review of Low  
777 Frequency Electromagnetic Wave Phenomena Related to Tropospheric-  
778 Ionospheric Coupling Mechanisms. *Space Science Reviews*, 168, 551-593.
- 779 Singer H. J., Matheson L., Grubb R., Newman A. and Bouwer S. D., 1996.  
780 Monitoring space weather with the GOES magnetometers. NOAA Space  
781 Environment Center.
- 782 Stoker P. H., 1993. Energetic Electron Power Flux Deposition at Sanae  
783 (L=4.0) from Riometer Recording. *Journal of Geophysical Research*, 98,  
784 19111-19116.
- 785 Streltsov A. V. and W. Lotko, 2004. Multiscale electrodynamics of  
786 the ionosphere-magnetosphere system. *J. Geophys. Res.*, 109, A09214.  
787 doi:10.1029/2004JA010457.
- 788 Tatsuta K., Hobara Y., Pal S. and Balikhin M., 2015. Sub-ionospheric VLF  
789 signal anomaly due to geomagnetic storms: a statistical study. *Ann. Geo-*  
790 *phys.*, 33, 1457-1467.
- 791 Thomson N. R., Rodger C. J. and Dowden R. L., 2004. Ionosphere gives size  
792 of greatest solar flares. *Geophys Res Lett* 31, L06803.
- 793 Tsurutani B. T. and Meng C. I., 1972. Interplanetary magnetic field varia-  
794 tions and substorms. *J. Geophys. Res.*, 77, 2964-2970.
- 795 Tsurutani B. T., Gonzalez W. D., Tang F., Akasofu S. I., and Smith E.  
796 J., 1988. Origin of interplanetary southward magnetic fields responsible  
797 for major magnetic storms near solar maximum (1978-1979). *J. Geophys.*  
798 *Res.*, 93, 8519-8531.
- 799 Tsurutani B. T., Gonzalez W. D., Gonzalez A. L. C., Tang F., Arballo J. K.,  
800 and Okada M., 1995. Interplanetary origin of geomagnetic activity in the  
801 declining phase of the solar cycle. *J. Geophys. Res.*, 100, 21717-21733.
- 802 Tsurutani B. T. and Gonzalez W. D., 1997. The interplanetary causes of  
803 magnetic storms: A review, in: *Magnetic Storms*. Edited by: Tsurutani,  
804 B. T., Gonzalez, W. D., Kamide, Y., and Arballo, J. K., Amer. Geophys.  
805 Un. Press, Wash. D.C., 98, 77-89.

- 806 Tsurutani B. T., Gonzalez W. D., Gonzalez A. L. C., Guarnieri F. L., Gopal-  
807 swamy N., Grande M., Kamide Y., Kasahara Y., Lu G., Mann I., McPher-  
808 ron R. L., Soraas F., and Vasyliunas V. M., 2006. Corotating solar wind  
809 streams and recurrent geomagnetic activity: A review. *J. Geophys. Res.*,  
810 111, A07S01, doi:10.1029/2005JA011273
- 811 Tsurutani B. T., Echer E., Guarnieri F. L. and Gonzalez W. D., 2011. The  
812 properties of two solar wind high speed streams and related geomagnetic  
813 activity during the declining phase of solar cycle 23. *J. Atmos. Solar-Terr.*  
814 *Phys.*, 73, 164, doi:10.1016/j.jastp.2010.04.003
- 815 Tsurutani B. T., Echer E., and Gonzalez W. D., 2011. The solar and inter-  
816 planetary causes of the recent minimum in geomagnetic activity (MGA23):  
817 a combination of midlatitude small coronal holes, low IMF BZ variances,  
818 low solar wind speeds and low solar magnetic fields. *Ann. Geophys.*, 29,  
819 839-849.
- 820 Wait J. R., 1959. Diurnal change of ionospheric heights deduced from phase  
821 velocity measurements at VLF. *Proc. IRE*, 47, 998.
- 822 Wait J. R. and Spies K. P., 1964. Characteristics of the Earth-ionosphere  
823 wave-guide for VLF radio waves. NBS Tech. Note 300.
- 824 Weigel R. S., 2010. Solar wind density influence on geomagnetic storm inten-  
825 sity. *J. Geophys. Res.*, 115, A09201, doi:10.1029/2009JA015062.
- 826 Wild J.A., Milan S.E., Cowley S.W.H., Dunlop M.W., Owen C.J., Bosqued  
827 J.M., Taylor M.G.G.T., Davies J.A., Lester M., Sato N., Yukimatu A.S.,  
828 Fazakerley A.N., Balogh A., Re'ime H., 2003. Coordinated interhemispheric  
829 SuperDARN radar observations of the ionospheric response to flux trans-  
830 fer events observed by the Cluster spacecraft at the high-latitude magne-  
831 topause. *Ann Geophys*, 21, 1807-1826.
- 832 Zhang S., Fukao S., Oliver W. L., and Otsuka Y., 1999. The height  
833 of the maximum ionospheric electron density over the MU radar,  
834 *Journal of Atmospheric and Solar-terrestrial Physics*, 61, 1367-1383,  
835 DOI:10.1016/S1364-6826(99)00088-7

836 **Figure Captions**

837 Figure 1: VLF signal propagation paths used in the study

838

839 Figure 2: Diurnal signature of VLF signals showing the aspects of the anal-  
840 ysed signal

841

842 Figure 3: (a) Diurnal VLF amplitude for GQD-A118 PP (b) Diurnal VLF  
843 amplitude for ICV-A118 PP (c) X-ray flux output (d) solar wind speed ( $V_{sw}$ )  
844 (d)  $B_z$  magnetic field component (e)  $H_T$  magnetic field (f) Dst and (g)  $A_p$   
845 variations during 14-19th February 2011

846

847 Figure 4: Daily Dst standard deviation, 4-hour mean signal amplitude before  
848 sunrise (MBSR), mid-day signal peak (MDP), 4-hour mean signal amplitude  
849 after sunset (MASS), sunrise terminator (SRT) and sunset terminator (SST)  
850 variations for (a) GQD-A118 and (b) ICV-A118 propagation path during  
851 14-19th February 2011

852

853 Figure 5: (a) Diurnal VLF amplitude for GQD-A118 PP (b) Diurnal VLF  
854 amplitude for ICV-A118 PP (c) X-ray flux (d)  $V_{sw}$  (d)  $B_z$  (e)  $H_T$  (f) Dst and  
855 (g)  $A_p$  variations during 26th-31st May 2011

856

857 Figure 6: Daily Dst standard deviation, 2-hour MBSR, MDP, 2-hour MASS,  
858 SRT and SST variations for (a) GQD-A118 and (b) ICV-A118 propagation  
859 path during 26th-31st May 2011

860

861 Figure 7: (a) Diurnal VLF amplitude for GQD-A118 PP (b) Diurnal VLF  
862 amplitude for DHO-A118 PP (c) X-ray flux (d)  $V_{sw}$  (d)  $B_z$  (e)  $H_T$  (f) Dst  
863 and (g)  $A_p$  variations during 24th-29th September 2011

864

865 Figure 8: Daily Dst standard deviation, 4-hour MBSR, MDP, 4-hour MASS,  
866 SRT and SST variations for (a) GQD-A118 and (b) DHO-A118 propagation  
867 path during 24th-29th September 2011

868

869 Figure 9: (a) Diurnal VLF amplitude for GQD-A118 PP (b) Diurnal VLF  
870 amplitude for DHO-A118 PP (c) X-ray flux (d)  $V_{sw}$  (d)  $B_z$  (e)  $H_T$  (f) Dst  
871 and (g)  $A_p$  variations during 23rd-28th October 2011

872

873 Figure 10: Daily Dst standard deviation, 4-hour MBSR, MDP, 4-hour MASS,  
874 SRT and SST variations for (a) GQD-A118 and (b) DHO-A118 propagation  
875 path during 23rd-28th October 2011

876

877 Figure 11: Daily Dst deviation (fluctuation) and trend in variation of signals  
878 MDP, MBSR, MASS, SRT and SST one day before and after each of the 16  
879 selected storm conditions for (a) GQD-A118 and (b) DHO-A118 propagation  
880 paths. A '0' indicate absence of data

881

882 Figure 12: Daily Dst deviation (fluctuation) and trend in variation of 2-day  
883 mean MDP, MBSR, MASS, SRT and SST before, during and after an event  
884 for (a) GQD-A118 and (b) DHO-A118 propagation paths. A '0' indicate  
885 absence of data