

**₁ Cassini observations of ionospheric plasma in
₂ Saturn's magnetotail lobes**

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Abstract. Studies of Saturn's magnetosphere with the Cassini mission have established the importance of Enceladus as the dominant mass source for Saturn's magnetosphere. It is well known that the ionosphere is an important mass source at Earth during periods of intense geomagnetic activity, but lesser attention has been dedicated to study the ionospheric mass source at Saturn. In this paper we describe a case study of data from Saturn's magnetotail, when Cassini was located at $\simeq 2200$ hours Saturn local time at $36 R_S$ from Saturn. During several entries into the magnetotail lobe, tailward-flowing cold electrons and a cold ion beam were observed directly adjacent to the plasma sheet and extending deeper into the lobe. The electrons and ions appear to be dispersed, dropping to lower energies with time. The composition of both the plasma sheet and lobe ions show very low fluxes (sometimes zero within measurement error) of water group ions.

The magnetic field has a swept-forward configuration which is atypical for this region and the total magnetic field strength is larger than expected at this distance from the planet. Ultraviolet auroral observations show a dawn brightening and upstream heliospheric models suggest that the magnetosphere is being compressed by a region of high solar wind ram pressure. We interpret this event as the observation of ionospheric outflow in Saturn's magnetotail. We estimate a number flux between $(2.95 \pm 0.43) \times 10^9$ and $(1.43 \pm 0.21) \times 10^{10} \text{ cm}^{-2} \text{ s}^{-1}$, one or about two orders magnitude larger than suggested by steady state MHD models, with a mass source between 1.4×10^2

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25 and 1.1×10^3 kg/s. After considering several configurations for the active at-
26 mospheric regions, we consider as most probable the main auroral oval, with
27 associated mass source between 49.7 ± 13.4 and 239.8 ± 64.8 kg/s for an av-
28 erage auroral oval, and 10 ± 4 and 49 ± 23 kg/s for the specific auroral oval
29 morphology found during this event. It is not clear how much of this mass
30 is trapped within the magnetosphere and how much is lost to the solar wind.

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1. Introduction

31 Saturn's magnetosphere is a complex multi-component plasma system with several inter-
 32 nal plasma sources in addition to the solar wind. The largest internal plasma source is from
 33 photoionisation and electron-impact ionisation of neutral water and nitrogen molecules
 34 from the icy moon Enceladus. These ions are subsequently processed by photolytic and
 35 radiolytic processes to produce H^+ , and a variety of water group ions such as OH^+ and
 36 O^+ that are collectively referred to as W^+ . The other natural satellites, the rings, and
 37 Saturn's atmosphere are minor internal sources. The solar wind also plays a role as an
 38 external plasma source. A number of studies have focused on the moons, rings and solar
 39 wind as plasma sources, to constrain the extent to which they drive the system. In this
 40 paper we provide the first in situ constraints on the role that the ionosphere plays as a
 41 mass source for Saturn's magnetotail, via the first observation of ionospheric outflow at a
 42 giant planet.

1.1. Plasma sources and transport in Saturn's magnetosphere

43 *Shemansky et al.* [1993] presented Hubble Space Telescope (HST) observations of an
 44 OH torus extending from 3 to 8 R_S ($1 R_S = 60268 \text{ km}$). They identified Enceladus, and
 45 to a lesser extent the other icy moons, as H_2O sources for the magnetosphere. *Jurac et al.*
 46 [2002] and *Richardson and Jurac* [2004] estimated the amount of H_2O needed to maintain
 47 the OH cloud and found that a source rate of 3.75×10^{27} H_2O molecules/s (112 kg/s)
 48 was required to maintain this cloud, of which 93 kg/s must be coming from the orbit of
 49 Enceladus. This estimate sits within a range of estimated rates between 10^{26} and 10^{28}

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50 molecules/s ($\simeq 35\text{-}350$ kg/s) [*Tokar et al.*, 2006; *Waite et al.*, 2006; *Hansen et al.*, 2006].

51 The large variability in these figures may be a natural result of the time-variability of the
52 Enceladus source. Following processing of these neutrals by neutral-plasma chemistry, the
53 total plasma source rate is around 60-100 kg/s [*Fleshman et al.*, 2013].

54 Titan has also been studied as a source of mass for Saturn's magnetosphere. *Johnson*
55 *et al.* [2009] estimates a total ion loss rate from Titan of $1 - 5 \times 10^{26}$ amu/s (0.16-0.83
56 kg/s). *Coates et al.* [2012] estimated a loss rates of $(8.9, 1.6, 4.0) \times 10^{25}$ amu/s for three
57 crossings of Titan's tail, for an average loss rate of 0.8 kg/s.

58 Saturn's main rings have an O^+ and O_2^+ atmosphere which can be ionised and act as
59 a mass source for the magnetosphere [*Tokar et al.*, 2005; *Johnson et al.*, 2006a; *Johnson*
60 *et al.*, 2006b; *Bouhram et al.*, 2006; *Luhmann et al.*, 2005; *Martens et al.*, 2008; *Tseng*
61 *et al.*, 2010]. The ring atmosphere was predicted to vary seasonally as the incidence angle
62 of the solar radiation on the main rings varies seasonally [*Tseng et al.*, 2010]. Using a
63 photochemical model and Cassini plasma spectrometer (CAPS) data, *Elrod et al.* [2012]
64 demonstrated that observed changes in the ring plasma over time were due to seasonal
65 change in the production of neutrals from Saturn's ring atmosphere. We are not aware of
66 any published estimates of the mass loading rate due to the rings.

67 Plasma produced in the inner magnetosphere from these sources is transported to the
68 outer magnetosphere. This transport is regulated by the centrifugally driven interchange
69 instability [*Mauk et al.*, 2009, and references therein]. The most detectable signature of
70 this process is the injection of hot plasma into the inner magnetosphere accompanied by

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71 magnetic pressure enhancements or deficits [e.g. *Hill et al.*, 2005; *André et al.*, 2005, 2007;
72 *Thomsen*, 2013].

73 The solar wind and ionosphere are thought to be secondary sources but the source
74 rates have only been estimated and there are no observational constraints. To estimate
75 the magnitude of the solar wind source a common approach is to multiply the solar
76 wind mass flux $n_{SW}v_{SW}$ by the cross-sectional area of the magnetosphere to obtain an
77 upper limit for the source rate: $n_{SW}v_{SW}\pi R_0^2$. An efficiency factor $O(10^{-3})$ is included
78 to account for diversion of the the solar wind and magnetosheath plasma around the
79 magnetosphere and the ability of magnetosheath plasma adjacent to the magnetopause
80 to enter the magnetosphere [*Hill*, 1979; *Hill et al.*, 1983; *Vasyliuñas*, 2008; *Bagenal and*
81 *Delamere*, 2011]. Applying this logic with a solar wind number density between 0.002
82 and 0.4 cm^3 , and a solar wind speed between 400 and 600 km/s [*Cravry et al.*, 2005]
83 with a magnetopause of cross-sectional area $\pi(30 R_S)^2$ (using the terminator radius of
84 the magnetopause from *Kanani et al.* [2010]), gives an upper limit of between 8.21×10^{27}
85 and 2.46×10^{30} protons/s (hence between about 13 and 4119 kg/s). Combined with the
86 efficiency factor of 10^{-3} the solar wind is a minor source.

1.2. Ionospheric outflow from Saturn's atmosphere

87 The physical mechanisms which lead to the ionosphere outflowing into space were the-
88 orised before the ionospheric outflow was detected at Earth. *Dessler and Michel* [1966]
89 and *Bauer* [1962] argued that, since the magnetospheric tail has a lower pressure than the
90 ionosphere, there should be a continuous escape of thermal plasma from the ionosphere

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91 into the tail (referred to just H^+ and He^+ at Earth). By analogy with the solar wind,
92 *Axford* [1968] suggested that this flow should be supersonic and named it the polar wind.

93 The classical polar wind is an ambipolar outflow of thermal plasma from the high lati-
94 tude ionosphere. The faster upflowing electrons create a charge separation with the more
95 gravitationally-bound ions, generating an ambipolar electric field that accelerates the ions
96 to achieve charge neutrality. The plasma, travelling and then escaping the topside of the
97 ionosphere, undergoes four transitions: from chemical to diffusion dominance, from being
98 subsonic to supersonic, from a collision dominated to a collisionless regime, a transition
99 from heavy to light ions (at Earth O^+ and H^+) since the light ions are less gravitationally
100 bound.

101 A steady state polar wind outflow is highly improbable. Magnetospheric electric fields
102 make the ionosphere-polar wind system convect constantly across the polar region, polar
103 cap, nightside auroral oval, nighttime trough, and sunlit hemisphere. When the mag-
104 netic activity increases, plasma convection speeds and particle precipitations intensify.
105 Three-dimensional time-dependent simulations of the global ionosphere and polar wind
106 have shown that, when the geomagnetic activity changes, the temporal variations and
107 horizontal plasma convection affect the polar wind and its dynamics. Three-dimensional
108 models (a global ionosphere-polar wind model) studied how much a geomagnetic storm
109 (for different solar cycles conditions) would have influenced the atmospheric system [*Gan-*
110 *guli*, 1996; *Schunk and Nagy*, 2009, and references therein]. Polar wind outflow increases
111 with geomagnetic activity.

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112 Different mathematical approaches have been used over the years to model the complex-
113 ity of the polar wind, such as hydrodynamical and hydromagnetic modelling, generalized
114 transport, and kinetic models. Also, numerous studies have been conducted of the non-
115 classical polar wind, which may contain, for example, ion beams or hot electrons. A
116 wealth of processes might be acting in the polar wind and still understanding is needed
117 [*Ganguli*, 1996].

118 Observational evidence of the polar wind at Earth was presented by *Hoffman* [1970]
119 using data from Explorer 31 showing field-aligned aligned H^+ with speed $\simeq 10$ km/s, and
120 flux $\simeq 10^8$ $\text{cm}^{-2}\text{s}^{-1}$ above 2500 km altitude. Using ISIS 2 and OGO data, similar results
121 for H^+ were obtained, plus O^+ and He^+ observations were added to the picture by *Brinton*
122 *et al.* [1971]; *Taylor and Walsh* [1972]; *Hoffman et al.* [1974]; *Taylor Jr. and Cordier*
123 [1974]; *Hoffman and Dodson* [1980]. More recently, *Chandler et al.* [1991] measured ion
124 density, velocity and flux variations of polar wind outflows using DE 1 data.

125 Electron temperature anisotropies, the relationship between the plasma pressure gradi-
126 ent between the ionosphere and deep magnetosphere, and the process of ambipolar diffu-
127 sion along magnetic field lines was established using Akebono data [*Abe et al.*, 1993a, b;
128 *Yau et al.*, 1995]. Observations of ionospheric outflow in the magnetosphere are harder
129 to make given the temperature of the plasma and charging of the spacecraft. Using Clus-
130 ter data, *Engwall et al.* [2009a] and *Engwall et al.* [2009b] inferred a total outflow from
131 Earth's polar ionosphere of the order of 10^{26} ions/s, which confirmed previous simulation
132 results arguing for the continuous presence of a low-energy ion population in the lobes. In

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133 addition, they inferred that the solar wind dynamic pressure and interplanetary magnetic
134 field played a role in influencing these populations in the lobes.

135 The polar wind is an important source of plasma in Earth's magnetosphere during
136 periods of geomagnetic activity. The extent of the ionosphere as a plasma source at
137 Saturn has been investigated using numerical models [*Frey, 1997; Glocer et al., 2007*].
138 These models solve the field-aligned gyrotropic transport equation [*Gombosi and Nagy,*
139 1989] for ions and electrons and simulate multiple convecting field line solutions.

140 *Glocer et al.* [2007] applied this model to Saturn, adapting the chemistry for the compo-
141 sition of Saturn's thermosphere. The model considers the behavior of H^+ and H_3^+ . It as-
142 sumes a stationary neutral atmosphere and models a range in altitude from 1400 (chemical
143 and thermal equilibrium) to 61000 (lower pressure) km. The background neutral atmo-
144 sphere required as an input relies on analysis of the stellar occultation measurements (low
145 latitude) of the Voyager 2 Saturn flyby, presented by *Smith et al.* [1983]. The estimates
146 for temperature and density of the neutrals were made at low latitudes, therefore the
147 model counts for the uncertainty on these parameters with a wide array of temperatures
148 (420-1500 K) which takes into account the possible density and temperature variations
149 from low to high latitudes. From this model, *Glocer et al.* [2007] estimate the polar wind
150 number flux of 7.3×10^6 to $1.7 \times 10^8 \text{ cm}^{-2}\text{s}^{-1}$ at 10000 km, providing a total source rate
151 to the magnetosphere of 2.1×10^{26} to $7.5 \times 10^{27} \text{ s}^{-1}$, for a source rate between 0.35 kg/s
152 and 1.25 kg/s.

153 Unfortunately, there are no observational constraints with which to compare these model
154 results. In this paper we report the detection of cold plasma in Saturn's magnetotail

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lobes, consider the interpretation of polar wind outflow, and use these observations to
constrain the ionosphere as a source of plasma for Saturn's magnetosphere. In section 2
we describe the instrumentation used for this study, in section 3 we show an overview of the
observations, the spacecraft trajectory, and the inferred upstream solar wind conditions.
The detailed case study is presented in section 4 and various interpretations discussed in
section 5. The implications for the physics of Saturn's magnetosphere are presented in
section 5.

2. Instrumentation

We use data from the Cassini Dual Technique Magnetometer (MAG) [*Dougherty et al.*,
2004], the Cassini Plasma Spectrometer (CAPS) [*Young et al.*, 2004], the Magnetospheric
Imaging Instrument (MIMI) [*Krimigis et al.*, 2004], the Radio and Plasma Wave Science
instrument (RPWS) [*Gurnett et al.*, 2004], and the Ultraviolet Imaging Spectrometer
(UVIS) [*Esposito et al.*, 2004].

CAPS measures the energy per charge and arrival direction of electrons and ions. The
instrument consists of three sensors: the Electron Spectrometer (ELS) which measures
electrons from 0.7 eV/q to 29 keV/q, the Ion Beam Spectrometer (IBS) which measures
narrow ion beams from 1 eV/q to 50 keV/q, and the Ion Mass Spectrometer (IMS) that
measures ions from 1 eV/q to 50 keV/q, followed by a time of flight (TOF) analyzer for the
determination of mass per charge of incoming particles. A motor-driven actuator rotates
the sensor package to provide 208-degree scanning in the azimuth of the spacecraft, nearly

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174 2π sr of the sky can be swept across every 3 minutes; spacecraft rolls can occasionally
175 increase the field of view to 4π sr.

176 MAG measures the strength and direction of the magnetic field around Saturn via a
177 fluxgate magnetometer and a vector helium magnetometer mounted on an 11 m spacecraft
178 boom, with the FGM located in the middle of the boom and the VHM at the end. The
179 magnetometer boom distances the sensors from the stray magnetic field associated with
180 the spacecraft and its subsystems and, especially with spacecraft generated field variations,
181 spacing the sensors at different distances along the boom allows this field to be better
182 characterised and removed from the observations. This study uses data from the fluxgate
183 magnetometer.

184 MIMI consists of three detectors: Charged Energy Mass Spectrometer (CHEMS), the
185 Low Energy Magnetospheric Measurement System (LEMMS), and the Ion and Neutral
186 Camera (INCA). CHEMS measures charge and compositions of ions with energy range
187 between $\simeq 3$ to 220 keV/q, combining electrostatic deflection and TOF to measure the
188 energy and composition of the energetic particles. INCA operates in two different modes,
189 over the energy range between 7 keV/nuc and 3 MeV/nuc. In its ion mode INCA measures
190 directional distribution, energy spectra and composition of ions and, in its neutral mode, it
191 takes remote images of the global distribution of the energetic neutral atoms, determining
192 their composition and energy spectra for each image pixel. INCA has a field of view of
193 120° in latitude and 90° in azimuth, whereas when the spacecraft is rotating the camera
194 covers about 4π sr. LEMMS consists of two oppositely directed telescopes, a low-energy

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195 telescope designed to detect ions with energy ≥ 30 keV and electrons with energy between
196 15 keV and 1 MeV, and a high-energy telescope for ions with energy range between 1.5
197 and 160 MeV/nuc and electrons (0.1-5 MeV).

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198 RPWS measures radio emissions, plasma waves, thermal plasma and dust in the vicinity
199 of Saturn. Three nearly orthogonal electric field antennas detect electric fields over a
200 frequency range from 1 Hz to 16 MHz, and three orthogonal search coil antennas measure
201 magnetic fields between 1 Hz to 12 kHz. A Langmuir probe is used to measure the
202 electron density and temperature. Five receiver systems process signals from the electric
203 and magnetic antennas.

204 UVIS measures ultraviolet light between the wavelengths of 55.8 and 190 nm for imaging
205 spectroscopy and spectroscopic measurements of the structure and composition of the
206 atmospheres of Titan and Saturn, rings, and surfaces, through two telescopes. It comprises
207 two spectrographic channels: an extreme ultraviolet channel (EUV), that measures spectra
208 between 55.8 and 118 nm, and a far ultraviolet channel (FUV), which measures spectra
209 between 110 and 190 nm.

3. Overview and upstream conditions

210 Figure 1 shows Cassini's trajectory on our day of interest, 21 August 2006 (day of year
211 233). The spacecraft was located in the dusk flank, about $36 R_S$ from the planet, north
212 of the equator at $\simeq 13.3^\circ$ latitude, and in the pre-midnight sector at 22:13 Local Time.
213 Cassini was on the outbound leg of revolution (orbit) 27.

214 In Figure 2 we show time-energy electron and ion spectrograms, and magnetic field
215 components in the KRTP (Kronocentric Radial-Theta-Phi) coordinate system plus the
216 field magnitude, for the time interval from 20-23 August 2006 (day of year 232-235).
217 Both electron and ion spectrogram are represented in differential energy flux units (DEF)
218 [$\text{m}^{-2}\text{s}^{-1}\text{sr}^{-1}\text{eV eV}^{-1}$]. IMS measures Energy/q of incoming ions - hence ions with same
219 Energy/q are recorded in the same bin - but the instrument has different response functions

220 for different ion species. However, the best calibration available at the moment is the one
221 that considers all the ion population made of protons. In KRTP coordinate system the B_r ,
222 component of the magnetic field is positive pointing outward from Saturn, hence positive
223 when the spacecraft is northward of the center of the current sheet, B_θ is positive pointing
224 southward, B_ϕ is positive in the corotation direction.

225 The colored boxes indicate when the spacecraft was located in various regions as deter-
226 mined from the magnetic field and plasma data. For example the lobes are characterized
227 by a strong and steady magnetic field, almost entirely in the B_r and B_ϕ directions, lack of
228 both energetic particles and 100 eV plasma electrons. Centrifugal forces confine plasma
229 to the equatorial region in giant planet magnetospheres. Since the field lines in the tail
230 extend for long distances, the lack of thermal plasma on these tail field lines does not
231 necessarily mean that the field lines are open: it may simply mean that the spacecraft is
232 sufficiently far from the equatorially-confined plasma that it cannot be detected. Current
233 sheet crossings and encounters are identified with vertical dashed lines. The arrow in
234 Figure 2d indicates a dipolarization event studied by *Jackman et al.* [2015]. Apart from
235 the current sheet encounters and crossings, the radial component of the field is generally
236 positive, until 22 August when it tends to be more negative, suggesting that typically the
237 spacecraft was north of the mean current sheet location until 22 August. The azimuthal
238 field is close to zero but fluctuates, sometimes indicating a significantly swept-forward
239 field (B_r and B_ϕ having same sign), but sometimes swept-back as it is over the rest of the
240 Saturnian magnetosphere [*Vasyliuñas*, 1983]. B_θ is generally positive suggesting closed
241 field lines.

242 The first period in the lobe is preceded by the passage of a plasmoid at 1001 on
243 20 August 2006 and shortly after a data gap, from 1515 to $\simeq 1530$, is followed by a
244 dipolarization at 1610 UT suggesting an extended interval of tail driving and subsequent
245 relaxation [Jackman *et al.*, 2015]. Between 1530 and 1800 UT following the dipolarization,
246 the plasma sheet is disturbed with an electron energy about 600 eV and fast directional
247 planetward flow between 1 and 10 keV/q. Following this period the electrons and ions
248 slowly reduce in energy, and hence Cassini detects a cooler, more typical plasma sheet.
249 During this period the magnetic field is swept-forward.

250 The following four periods in the lobes are characterised by low energy ions and elec-
251 trons, where the electrons are found just above the population of trapped spacecraft
252 photoelectrons, sometimes almost indistinguishable from the spacecraft photoelectrons
253 (around 10 eV). In each case the surrounding plasma sheet has electron energies typically
254 found in the tail plasma sheet [Arridge *et al.*, 2009]. In the third lobe period during 1200-
255 1800 on 21 August the electrons reach very low energies and appeared to be dispersed in
256 time with lower electron and ion energies observed towards the end of the period in the
257 lobe. During each of these four lobe periods the magnetic field is either purely radial or
258 is significantly swept-forward. After 23 August 2006 the plasma sheet and lobe period
259 structure returns to that typically found in the magnetotail [Arridge *et al.*, 2009].

260 There is no upstream solar wind monitor at Saturn and so models and propagations from
261 1 AU are often used to infer upstream solar wind conditions [e.g. Zieger and Hansen, 2008;
262 Hsu *et al.*, 2013; Badman *et al.*, 2015; Baker *et al.*, 2009; Jasinski *et al.*, 2014]. Propagation
263 models, e.g., mSWiM [Zieger and Hansen, 2008], which propagate solar wind conditions
264 measured at 1 AU, cannot be used for this interval since Saturn is far from apparent

265 opposition during this period. In this work we use the ENLIL model, which is a time
266 dependent 3D MHD heliospheric model [*Odstreil et al.*, 2004] operated at the Community
267 Coordinated Modeling center at NASA Goddard Space Flight Center. This is the only
268 heliospheric model that simulates solar wind conditions beyond 5 AU. ENLIL simulates
269 supersonic, low β plasmas, and must have inner coronal boundary conditions provided
270 by either the Wang-Sheeley-Arge (WSA) [*Arge and Pizzo*, 2000] (inner boundary located
271 at 21.5 solar radii) or MHD-Around-a-Sphere (MAS) [*Riley et al.*, 2001] (inner boundary
272 located at 30 solar radii) models. The outer boundary can be chosen to extend up to 10
273 AU as appropriate for simulations for Saturn.

274 ENLIL was run using Carrington Rotation 2046 as appropriate for this interval. Figure
275 3 shows the global heliosphere simulation during this period. In Figure 3 we can see
276 density scaled with r^2 , where r is heliocentric distance, from the heliosphere model. This
277 shows a sequence of compression regions passing over Saturn. In Figure 4 we show a time
278 series of solar wind conditions extracted at Saturn. We can compare the times of high solar
279 wind density shown in Figure 3 with what we see in Figure 4, namely solar wind density,
280 speed, dynamic pressure, and total magnetic field strength in Radial Tangential Normal
281 (RTN) coordinates. These show that this event is included in a solar event compression
282 period. *Jian et al.* [2011] presented comparisons between ENLIL and Ulysses data at 5
283 AU and showed that the ENLIL predictions for the arrival of solar wind structures had
284 an error of approximately two days. Even if this error is doubled to four days at 10 AU,
285 Saturn is still immersed in a compression region during this event.

286 We also studied Cassini remote sensing data to check if there were increases in activity
287 of Saturn Kilometric Radiation (SKR) emissions and auroras that could support the

simulation results, which suggest that Saturn is immersed in a solar wind compression
region [Desch and Rucker, 1983; Kurth et al., 2005; Badman et al., 2008a; Clarke et al.,
2009; Kurth et al., 2013]. It has to be considered, anyway, that Stallard et al. [2012]
showed a delay of $\simeq 8$ h between the arrival of a solar wind compression and brightening
of the aurora. In figure 5 we show two auroral images taken by the UVIS instrument on
Cassini. The data are projected onto a latitude and local time grid at 1000 km altitude:
the figure shows the total FUV intensity, which is predominantly H and H₂ emissions.
Unfortunately, since Cassini is far from the planet and close to the equatorial plane, the
view of the polar region is only partial and at low spatial resolution. Figure 5a shows a
bright aurora, seen on the dawnside from 0030 to 0700 local time (with no viewing beyond
0700) of the northern hemisphere. The aurora reaches 30 kR between 0100 and 0700 local
time from 12° to about 16° colatitude. Figure 5b shows a more extended aurora: we have
two areas, one from midnight to 0600 local time, from about 4° to about 20°, with a
brightness between 10 and 30 kR and a second area from 1500 to 1900 local time, from 8°
to about 14°, with a brightness that reaches $\simeq 7$ kR. Clarke et al. [2009] reports similar
brightness for aurora during disturbed conditions. Since Cassini is far away from Saturn,
orbiting in the equatorial plane, the auroral emissions observed by UVIS are subject to
significant limb-brightening, whereby the emissions are viewed through a long column of
atmosphere near the poles, compared to lower latitudes. This was corrected using the
sine of the emission angle, but since each UVIS pixel covers a large area on the planet,
the images could still be partially affected by the limb-brightening, which, however, does
not affect the extension of the aurora, or the presence of aurora itself. These auroral
emissions, the fact that the aurora is extending poleward and is brighter on the dawn

311 side, suggest that there are some tail dynamics influencing the auroral region and which
312 has been shown to be a consequence of the passage of solar wind compression regions
313 [*Stallard et al.*, 2008; *Cowley et al.*, 2005].

314 Figure 6 shows electric field spectrogram, up to 2 MHz, from RPWS. In this time range
315 the emissions above 3 kHz are Saturn Kilometric Radiation (SKR) and extend until 0600
316 UT on 22 August with low frequency extensions [*Jackman et al.*, 2009]. After this time,
317 narrowband periodic emissions are observed near 5 kHz that are probably generated closer
318 to the planet, hence not likely to be associated with plasma detected near the spacecraft.
319 Narrow band emissions are also observed around 2 kHz, notably at 1530 on 21 August
320 and at 1430 on 23 August, possibly associated with electron plasma oscillations. These
321 can be used to infer the electron density from the frequency which suggests a density of
322 0.05 cm^{-3} , compatible with the CAPS/ELS electron moments. The spectrum below 1.5
323 kHz is noisy, mostly probably given by interference from the spacecraft reaction wheels.
324 However, below about 50 Hz, there are quite visible features (middle of 20 August and
325 just after 06:00 on 21 August) not generated by spacecraft interference. An examination
326 of the corresponding magnetic spectrogram (not shown) shows that these features do not
327 have a magnetic component.

328 More diffuse broadband emissions below 10 Hz might be associated with ionospheric
329 outflow and are seen to correlate with the observation of cold plasma in the tail lobes as
330 identified in figure 2. These emissions may be whistler mode emissions as detected in the
331 magnetotail of Uranus [*Kurth et al.*, 1989]. However, the corresponding magnetic field
332 spectrogram (not shown) does not show a magnetic component to this diffuse broadband
333 emission, suggesting an electrostatic mode.

334 The SKR observations in Figure 6 show evidence of a brightening in SKR near 1800
335 UT on 20 August as noted by *Jackman et al.* [2015] and which may be associated with
336 the dipolarization event at 1610 UT reported in that study. SKR emissions are active
337 throughout the rest of the 20 August and 21 August, appearing to switch off early on 22
338 August, clearly showing evidence of magnetospheric dynamics during this period [*Desch,*
339 *1982; Kurth et al., 2005; Badman et al., 2008b; Jackman et al., 2009*]. The SKR main spec-
340 trum, which typically ranges between 100-400 kHz, is generated by the cyclotron maser
341 instability. This is in contrast to the lower frequency narrowband emissions mentioned in
342 the previous paragraph which are likely caused by a different mechanism altogether.

4. Data analysis

343 Turning our attention to the specific time interval that we focus on in this case study.
344 Figure 7 is a zoom in of the case study interval, from 1200 to 2400 UT on 21 August,
345 from Figure 2, and shows the characteristics of this event from different instruments on
346 Cassini.

347 Looking at the electron distributions first (Figure 7a) we see that at 1330 on the 21st
348 August (when the spacecraft moves completely into the northern lobe) the population
349 below around 5 eV are trapped spacecraft photoelectrons and the upper edge of this
350 distribution shows that the spacecraft potential is around 5 V. Typically the potential
351 in the lobes is 30-50 V and so this is consistent with the presence of dense plasma in
352 the lobes. From 1330 UT the ambient electron energy drops from 10^2 eV to a few eV
353 and this ambient population is sometimes hard to distinguish from the trapped spacecraft
354 photoelectron distribution, especially towards the end of the interval. Further evidence of
355 the unique nature of this event is revealed by the low energy of these electrons since in the

356 quiet tail lobes, *Arridge et al.* [2009], finds an electron temperature of $\simeq 100$ eV. Moreover,
 357 Figure 8, where we compare two electron spectra, one from this event and one from the
 358 magnetosheath, shows how colder the electron population for this event is compared to
 359 another region of the magnetosphere.

360 Figure 7b show the ion populations with $\simeq 3$ keV/q ions in the plasma sheet and lower
 361 energies in the lobe. The measured ion fluxes are larger than in the plasma sheet and are
 362 also seen to slowly disperse in energy from $\simeq 500$ eV/q to $\simeq 100$ eV/q over a period of
 363 around four hours. During the period in the plasma sheet, the field of view of IMS does
 364 not cover the ideal corotation direction and so sees only weak fluxes from directions $>$
 365 30° from corotation. However, during the period in the lobes, IMS views flows coming
 366 from the direction of Saturn. Figure 9 shows measured ion fluxes as a function of the
 367 look direction around the spacecraft, in a polar projection, expressed in OAS coordinate.
 368 In this coordinates system \mathbf{S} is the axis along the Cassini-to-Saturn line, \mathbf{O} is defined by
 369 $\mathbf{S} \times (\boldsymbol{\Omega} \times \mathbf{S})$, where $\boldsymbol{\Omega}$ is the planet spin axis, and \mathbf{A} completes the right-hand system.
 370 We can represent a point around the spacecraft with two angles relative to the \mathbf{S} axis: θ
 371 (range from 0° to 180°) is the latitude angle, so it is the polar angle away from Saturn,
 372 and ϕ (range from 0° to 360°) is the azimuth around \mathbf{S} axis, referenced to 0° in the \mathbf{O}
 373 direction. Specifically in figure 9, $\theta = 90^\circ$ is represented by the inner circle and $\theta = 180^\circ$
 374 is the outer circle. Hence, these plots show the presence of a cold ion population with a
 375 width of $\approx 40^\circ$ flowing tailward.

376 The ion composition during this interval is also unusual and was determined by a fit
 377 of CAPS/IMS time-of-flight data to a forward model [e.g. *Thomsen et al.*, 2010]. In
 378 the plasma sheet between 1200 and 1320, where Cassini crosses the plasma sheet twice,

379 passing from the north lobe to the south lobe, and coming back to the north lobe again,
 380 H^+ counts are $\simeq 10^4$ and ($m/q = 2$) $\simeq 10^3$, and the ratio of water group ions to hydrogen,
 381 $[W^+]/[H^+]$, and $m/q=2$ to hydrogen, $[m/q=2]/[H^+]$, are $1.79 \pm 1.58 \%$ and 2.45 ± 0.15
 382 $\%$ respectively. Hence the plasma sheet appears to be devoid of water group ions. After
 383 1320, once the spacecraft is in the north lobe, H^+ counts are $\simeq 10^5$, one order of magnitude
 384 larger than the counts of when the spacecraft was crossing the plasma sheet and $m/q = 2$
 385 counts are five times larger. During this time period, the ratio between water group ions
 386 and hydrogen $[W^+]/[H^+]$ is zero within error and $[m/q=2]/[H^+] = 2.23 \pm 0.04 \%$. From
 387 2130 to 2400, the spacecraft returns to the plasma sheet, H^+ counts are $\simeq 10^4$, one order
 388 of magnitude lower than in the lobes and $m/q = 2$ counts diminish to 10^2 . During this
 389 time period the ratio between water group ions and hydrogen $[W^+]/[H^+]$ is again zero
 390 within error and $[m/q=2]/[H^+] = 3.22 \pm 0.29 \%$.

391 We checked previous and following spacecraft orbits at the same latitude and the same
 392 local time. For the previous orbit (28th July 2006), the lobes are empty of ions and
 393 when the spacecraft crosses the plasma sheet twice between 0400 and 1200, $[W^+]/[H^+] \simeq$
 394 $30.02 \pm 15.11 \%$ and $[m/q=2]/[H^+] = 24.79 \pm 0.26 \%$. On the following orbit (13th-14th
 395 September 2006) the spacecraft seems located always in the lobes, which are mostly empty
 396 of ions. Therefore, we consider this an atypical time interval.

397 Looking at MIMI/CHEMS data in Figure 7c and 7d, we find that between 1400 and
 398 2030, the lobes are populated by hot H^+ with foreground. Hot O^+ ions start to appear at
 399 around 18:45 and they seem slightly dispersed in energy. Higher intensities are observed
 400 between 100 and 300 keV, while INCA sees O^+ even for larger than 500 keV (see next

401 paragraph). The pitch angle distribution for this population is generally between 30 and
402 90 degrees, implying an outward flow.

403 Throughout this interval the MIMI/INCA camera is in ion mode and so provides ad-
404 ditional information on the energetic ions. When the spacecraft is in the lobes, we find
405 no O^+ during most of the interval. Figure 10 shows O^+ distributions observed by INCA
406 from 18:45 to 19:28 and from 20:29 to 21:02.

407 By looking at the INCA look direction and the pitch angle coverage in the columns
408 corresponding to these intervals, we know that, during this time, the spacecraft orientation
409 is steady. Afterwards, we see energy peaks periodically between 18:55 and 19:01, 19:15 and
410 19:28, 19:42 and 19:55 and 20:08 and 20:22, associated with a first order anisotropy, when
411 the spacecraft starts rolling: an entire rotation is enclosed by approximately four white
412 squares corresponding to the period in the intensity peaks [e.g. *Kane et al.*, 2008]. The
413 pitch angle distribution is peaked between 0° and 90° indicating ions flowing downtail,
414 with scattering accounting for the intensities that appear between 90° and 120° . The
415 highest flux is detected for energies between 89 keV to 589 keV and a very low flux
416 for lower energies, until 20:35, when the O^+ covers energies from 46 and 589 keV. The
417 gyroradius for 89 keV to 589 keV ions is $\simeq 0.6$ to $1.5 R_S$, hence we interpret the ions
418 before 20:35 as a remote detection of the plasma sheet whilst the spacecraft is the lobes.
419 After this point the spacecraft approaches the plasma sheet (around 20:55) and we see
420 O^+ of all energies in the detector; the flow has a broader pitch angle distribution, that is
421 more focused between 0° and 120° with increasing energy. Observing what happens to the
422 magnetic field at the same time, we notice that the magnetic field has lowered, showing
423 a step of about 0.5 nT before 2000 in B_r and B_{tot} trend, then a bigger drop in the field

424 of more than 1 nT. Finally, at 2122 UT, the flux seems to be isotropic and the magnetic
425 field reaches low field strengths, indicating the plasma sheet encounter.

426 INCA observations of H^+ are contaminated with O^+ due to an instrumental effect but
427 are consistent with isotropic H^+ in the lobes and an increasing flux of H^+ in the plasma
428 sheet towards the end of the interval.

429 The magnetic field is swept-forward during this all interval, namely B_r and B_ϕ compo-
430 nents of the field maintain the same sign for more than 8 hours. With a single spacecraft it
431 is difficult to separate spatial and temporal effects and it is possible that this swept-forward
432 field configuration was a characteristic of this local time in Saturn's magnetosphere. Many
433 of Cassini's orbits have nearly identical coverage in local-time and latitude so, to check
434 if the field is typically swept-forward at this radial distance and local time, we examined
435 the sweep-back angle during the orbits of Cassini before (28th July 2006) and after (13th
436 and 14th September 2006) the orbit during this case study. The spiral angle of the field
437 indicated that, although the field was generally almost meridional (not swept-forward
438 or swept-back) during these orbits, only this orbit had the swept-forward configuration
439 indicating an unusual configuration.

440 *Jackman and Arridge* [2011] studied the magnetic field strength in the lobes and es-
441 tablished the average field strength at various radial distances fitting this to a power-law
442 function of radial distance, r in units of R_S , such that $B_{lobe}(nT) = (251 \pm 22)r^{-1.2 \pm 0.03}$. At
443 $36 R_S$ this expression predicts a field strength of 3.4 ± 0.3 nT which is 2 nT smaller than
444 observed, indicating either a highly compressed magnetosphere or where the magnetotail
445 was loaded with open magnetic flux, or both.

5. Interpretation and discussion

446 In interpreting these observations we have considered several possibilities for the pres-
 447 ence of cold ion beams at large distances in Saturn's magnetotail. Magnetic reconnection
 448 would result in rapid ion flows in the range 144-1240 km/s ($\simeq 10$ keV) [*Hill et al.*, 2008;
 449 *Jackman et al.*, 2014] and energised electrons in the beam with planetward and tailward
 450 ion flows depending on location relative to the X-line. In this case no such energised
 451 electrons are observed, the observed ion energies are small, and no large B_θ deflections
 452 are observed.

453 In the Saturn system cold plasma usually originates from ionization in the inner mag-
 454 netosphere. These cold plasma observations could potentially be the result of rapid cold
 455 plasma transport from the inner magnetosphere. In this case, however, we would detect
 456 water group ions from the moons and we would expect the ions to be centrifugally con-
 457 fined. Furthermore, in our observations we find ($W_\perp/B \simeq 2$ eV/nT) and conservation of
 458 the first adiabatic invariant implies that we should find a similar ratio close to the source
 459 of this ion population. According to *Arridge et al.* [2011, and references therein], who
 460 synthesised the results of many studies, we see that at distance of $5 R_S$, $8.7 R_S$ and $20 R_S$
 461 we would find respectively a $W_\perp/B \simeq 0.005$, $W_\perp/B \simeq 0.3$ and $W_\perp/B \simeq 9$ eV/nT. Hence,
 462 we do not find a ratio $\simeq 2$ close to the planet. Furthermore, the composition is quite
 463 different to what is usually seen in the inner and middle magnetosphere. It is also difficult
 464 to envisage a physical mechanism for removing ions from the inner magnetosphere to the
 465 tail in the form of a narrow directional ion population.

466 Finally, an alternative interpretation is that we detect Saturn's plasma mantle: ions
 467 that have entered the dayside magnetosphere via dayside reconnection and have mirrored

468 and flowed out tailward into the magnetotail. We would expect a solar wind plasma
 469 composition of $[m/q=2]/[H^+]$, namely $\simeq 4\%$ which is similar to $2.23 \pm 0.04 \%$ from
 470 time-of-flight fits. We would also expect the particles to conserve the first adiabatic
 471 invariant between the magnetotail and the cusp. From *Jasinski et al.* [2014] the electron
 472 temperature in the cusp is $\simeq 40$ eV in a field strength of 8 nT, thus $W_{\perp}/B \simeq 5$ eV/nT, and
 473 so we would expect electron energies in the magnetotail to be 25 eV, much higher than
 474 observed. At Earth, the electrons in the mantle have same energy as the electrons in the
 475 magnetosheath [*Formisano, 1980*] and we would expect the same to happen at Saturn.
 476 Figure 8 shows that the electron population for this event is significantly colder than the
 477 electron population in Saturn magnetosheath.

478 Furthermore, we can also examine the convection timescale for newly opened flux tubes
 479 compared with the speed of the ions. Assuming a flux tube moves tailward at 40 km/s
 480 ($\simeq 10\%$ of the solar wind speed) it would take 16.7 hours to traverse the 40 R_S from the
 481 dayside to the magnetotail. H^+ with 1 keV energy would have covered a distance of 436
 482 R_S in the same time, and the 50 eV H^+ a distance of about 100 R_S . This suggests that by
 483 the time the flux tube travels from dayside to the spacecraft position, it would be already
 484 emptied of ions. If we considered the rotation time instead (about 5 hours) we would find
 485 that 1 keV and 50 eV H^+ would travel respectively to distances of 131 R_S and 29 R_S .
 486 This leads us to not consider valid a mantle provenance for the plasma in our event.

5.1. Ionospheric outflow

487 For ionospheric outflow we would expect cold electrons and ions flowing tailward from
 488 Saturn into the magnetotail via the magnetotail lobes, as observed with CAPS/IMS and
 489 CAPS/ELS. We would expect the ion composition to be consistent with Saturn's iono-

490 sphere, i.e., H^+ , H_2^+ and H_3^+ . In the CAPS data, the ions are dominated by H^+ with a
491 smaller contribution from a species with $m/q=2$ which cannot be separated into H_2^+ and
492 He^{++} . Unfortunately, H_3^+ has a time-of-flight in CAPS/IMS which lies near an instru-
493 mental artifact and therefore cannot be extracted at this time. Hence, we interpret this
494 event as ionospheric outflow via a polar wind.

495 It is not possible to determine the connectivity of field lines (open or closed) during
496 the period of ionospheric outflow. This is a period of intense magnetospheric activity
497 and we think that precipitating electrons producing auroral emissions could happen si-
498 multaneously on the same field line as ionospheric outflow, but still in an upward current
499 region. Hence, the auroral emission and source for ionospheric outflow could be collocated
500 in the same region of the ionosphere. *Bunce et al.* [2008], used Cassini and Hubble Space
501 Telescope data to show that the Southern auroral oval is located at the boundary between
502 open and closed field lines. However, *Jinks et al.* [2014], using Cassini data, found that
503 the poleward edge of the upward current region is displaced equatorward from the polar
504 cap boundary in both the northern and southern hemispheres. Thus, the closed field line
505 region can be present also beyond the upward current region poleward boundary. This
506 could imply that the spacecraft was located on closed field lines.

507 Whilst in CAPS/IMS we see cold dispersed ions, we argue that the ions detected in
508 MIMI/INCA and MIMI/CHEMS from 18:45 belong to the plasma sheet: these ions are
509 flowing downtail at speeds $\simeq 1000$ km/s. The fact that the plasma sheet is emptied from
510 W^+ in the range of CAPS/IMS might suggest that the plasma sheet has been emptied
511 through a reconnection. In this scenario, whilst in the lobes CAPS/IMS detects ions
512 flowing downtail coming from the ionosphere, MIMI/INCA is remotely sensing ions flowing

513 downtail accelerated by reconnection. Hence, we think that reconnection is happening at
514 the boundary between the lobe and the plasma sheet beneath the spacecraft, while the
515 spacecraft is on open field lines (see Figure 11).

516 According to magnetospheric magnetic field models [*Khurana et al.*, 2006; *Bunce et al.*,
517 2003], this region of the magnetosphere is only slightly swept-forward, but the sweep-
518 forward increases during periods of increased solar wind dynamic pressure. Since these
519 models only include azimuthal fields due to magnetopause currents, this then shows that
520 the swept-forward configuration is due to magnetopause currents. The reason why we do
521 not see a strongly swept forward configuration in the previous and following orbits, at
522 about same latitude and same local time, is due to the CIR that is passing the planet
523 during this specific time period.

524 *Glocer et al.* [2007] coupled a polar wind outflow model with an MHD model of Saturn's
525 magnetosphere to estimate the number flux of ions outflowing from Saturn's ionosphere
526 in a steady state. They found value between 7.3×10^6 and 1.7×10^8 $\text{cm}^{-2}\text{s}^{-1}$. To com-
527 pare our observations with the *Glocer et al.* [2007] simulations we estimated the number
528 flux, nv , where n is the number density and v is the speed of ions in the tail, and used
529 conservation of magnetic flux to scale these to their values closer to Saturn.

530 The generally low numbers of counts during this event make the ion moment calculations
531 challenging, so we assumed the ion number density was equal to the electron number
532 density. The ion speeds were estimated by fitting the ion spectra with Gaussian plus a
533 background. Fits were filtered using the χ^2 for each fit and a manual inspection of the fit.
534 The fits were performed on the IMS anodes where peak fluxes were observed. The peak
535 energy from this fit was taken as the ion bulk flow energy (actually an upper limit since

536 this assumes the ions are completely cold). The ion speed was found to be $\simeq 400 \text{ km s}^{-1}$
 537 at about 1340, with the speed slowly diminishing to get to $\simeq 200 \text{ km s}^{-1}$ at about 1700
 538 UT.

539 Figure 12 shows the number density, speed, and calculated tail number flux, $n_t v_t$. We
 540 assumed 10% uncertainty on the electron densities [Arridge *et al.*, 2009]; the speed uncer-
 541 tainties were obtained by propagating the uncertainties in the peak energies found from
 542 our non-linear fits. The number flux uncertainties were calculated by propagating the
 543 uncertainties on n_t and v_t .

544 *Glocer et al.* [2007] presented number fluxes at an altitude of 10000 km. To map our
 545 observed number fluxes to this altitude, we assume that the number of outflowing ions
 546 are conserved in a flux tube from the ionosphere to the tail, and so use $B_t A_t = B_i A_i$, and
 547 therefore scale the tail number flux to get the ionospheric number flux by $n_i v_i = n_t v_t B_i / B_t$.
 548 The ionospheric field strength was calculated from a dipole at an altitude of 10000 km at
 549 an auroral colatitude of 12° . Using the observed tail field strength shown in Figure 12d
 550 we then calculate the ionospheric number fluxes as shown in Figure 12e. We obtained
 551 ionospheric number fluxes between $(2.95 \pm 0.43) \times 10^9$ and $(1.43 \pm 0.21) \times 10^{10} \text{ cm}^{-2} \text{ s}^{-1}$.
 552 These estimates are one order of magnitude larger than the value obtained by *Glocer et al.*
 553 [2007].

554 One possible interpretation for this discrepancy is due to the fact that the model was
 555 run for a steady atmosphere and steady magnetosphere, hence classical polar wind as
 556 defined in *Schunk et al.* [2007]. We argued that this event occurred during a CIR (co-
 557 rotating interaction regions) compression and with substantial magnetospheric activity,
 558 which produced enhanced outflows, namely generalised polar wind, as defined in *Schunk*

559 *et al.* [2007]. Moreover, due to low counts it is very difficult to evaluate the velocities
560 with fits that also consider the temperatures. From a preliminary estimate, we think that
561 the speed we calculated overestimate the velocities of a factor of about two. This would
562 affect the number flux of a factor of two, which is a minor contamination for our estimate.
563 In addition, *O'Donoghue et al.* [2015, and references therein] find a neutral temperature
564 not higher than 650 K, which from *Glocer et al.* [2007] should lead us to higher values of
565 number flux.

566 We calculated the total particle source rate using the polar cap area range extracted
567 by *Glocer et al.* [2007], obtaining an estimate between 8.6×10^{28} and $6.3 \times 10^{29} \text{ s}^{-1}$, that
568 leads to a mass source between 1.4×10^2 and $1.1 \times 10^3 \text{ kg/s}$, if there were no loss in the
569 tail.

570 By contrast, using the mean position of the northern and southern aurora (respectively
571 $15.1 \pm 1.0^\circ$ and $15.9 \pm 1.9^\circ$, *Carbary* [2012]) we can recalculate the area of the polar cap
572 and we obtain instead a rate of $49.1 \pm 9.8 \times 10^{27}$ and $23.7 \pm 4.7 \times 10^{28} \text{ s}^{-1}$, and a mass
573 source between 82.0 ± 16.5 and $395.6 \pm 79.2 \text{ kg/s}$.

574 If we considered instead, more realistically, an active area covering the region of auro-
575 ral emission (northern and southern extending respectively between 13.4° and 16.8° and
576 between 12.9° and 18.9° , average values taken from *Carbary* [2012]), we would obtain
577 a source rate between $29.7 \pm 8.1 \times 10^{27}$ and $14.3 \pm 3.8 \times 10^{28} \text{ s}^{-1}$ (mass source between
578 49.7 ± 13.4 and $239.8 \pm 64.8 \text{ kg/s}$).

579 Specifically for our event, we can see, from UVIS data, the active area from auroral
580 emissions in the hours prior the event: it extends from 5° and 12° , from 0000 to 1900
581 LT, with a data gap from 0600 to 1430 LT. If we considered an outflow only from the

582 brightest area (we then consider the area of the ring which includes the brightest area
583 from dawn to dusk, divided by two to consider the area with no auroral emission and the
584 data gap) with an estimated error of 2.5° (half of the projected polar grid resolution), we
585 would obtain $6.1 \pm 2.9 \times 10^{27}$ and $2.9 \pm 1.4 \times 10^{28} \text{ s}^{-1}$ for a total mass source between of
586 10 ± 4 and $49 \pm 23 \text{ kg/s}$.

587 Evaluating the fact that the whole auroral oval is not completely active in our event
588 and the active area has a different brightness, so possibly the outflow is not coming out
589 uniformly, a useful quantity to be defined is the rate and the mass source per hour of local
590 time, spanning 5° in latitude (in our event mainly from 5° and 10°): $3.2 \pm 1.9 \times 10^{26}$ and
591 $1.5 \pm 0.9 \times 10^{27} \text{ s}^{-1}$ per hour of local time, hence between 0.5 ± 0.3 and $3 \pm 1.5 \text{ kg/s}$ per
592 hour of local time.

593 The mass source estimates for different active areas are here calculated only for the
594 northern cap, the northern auroral and the aurora in the north pole we remote sense with
595 Cassini/UVIS. Moreover, the source rates calculated are upper limits of the ionospheric
596 mass contribution to the magnetosphere: we do not have any estimate at the moment on
597 how much of this mass stays indeed in the magnetosphere.

598 The ion dispersion can be generated by three different processes, two of them consisting
599 in a temporal variation and one of them in a spatial variation. The first temporal effect
600 is due to the fact that the energies of the particles emitted from the same area in the
601 ionosphere, have a Maxwellian distribution. Consequently, there is a velocity filter effect
602 whereby faster ions reach the spacecraft before slower moving ions, producing a continuous
603 velocity dispersion in the data. The second temporal scenario can be caused by a time
604 variability of the source. Therefore the spacecraft, in this case, would be detecting ions

605 originated from the same source, but a source that was in an excited state at the beginning
606 of the time interval, and then relaxed afterwards, emitting lower energy ions towards the
607 end of the interval. Lastly, the spatial scenario could be caused by the fact that the
608 spacecraft moves across different field lines which have their feet in different latitudes
609 in the ionosphere. At the beginning of the interval, the spacecraft would have been
610 located on a field line connected to an intensely excited region of the ionosphere, and
611 then the spacecraft moved onto field lines connected to a calmer region of the ionosphere.
612 Unfortunately, from the data it is not possible to distinguish among these three different
613 scenarios. A combination of two or all of these processes may be involved in producing
614 the observed ion energy dispersion.

615 We know the presence of an electron dispersion in the data, maybe produced by the
616 same process that caused the ion dispersion.

617 If we interpret the ion dispersion as a velocity filter effect, whereby faster ions reach
618 the spacecraft before slower moving ions from a spatially-restricted source, then we can
619 obtain the distance to that source from a time-of-flight expression $t = t_0 + d/v$, where
620 t is the time of observation, t_0 is the time at which all the ions left the source, v is the
621 speed of an ion observed at time t , and d is the distance to the source. We used the
622 velocities and times from Figure 12, calculated the inverse velocity, and then fitted a
623 straight line to t as a function of $1/v$ to obtain the intercept (t_0) and the gradient (d).
624 Therefore d is an estimate of the field-aligned distance that the ions travelled and the time
625 at which they left their source t_0 . The ion counts were not considered to be sufficiently
626 far above background to perform this analysis using energy-time spectra and, hence, we
627 have performed the moment analysis as an alternative. The distance was found to be

628 $69 \pm 3 R_S$. Considering that our speeds are possibly over-estimated by a factor of $\simeq 2$.
 629 we then obtain a distance of $34 \pm 1 R_S$. This is consistent with the length of a field line
 630 from Cassini to Saturn's ionosphere calculated by tracing field lines in a magnetospheric
 631 field model *Khurana et al.* [2006], thus strengthening the interpretation of this event as
 632 ionospheric outflow.

6. Conclusions

633 We presented a case study of an event from Saturn's magnetotail from 21 August (day
 634 of year 233) 2006. The event is enclosed in a time period when the magnetosphere was
 635 compressed by a region of high solar wind dynamic pressure, as identified from a global
 636 heliosphere simulation (ENLIL) and auroral and SKR intensifications. Cold ions and
 637 electrons are detected in the lobes and the cold plasma ion composition does not show
 638 evidence for W^+ ions, neither when the spacecraft is located in the plasma sheet or when
 639 Cassini is in the lobes.

640 After considering different interpretations, we conclude that this event is an example
 641 of ionospheric outflow in Saturn's magnetotail. This is the first time that a low-energy
 642 ionospheric outflow event has been detected at planets other than Earth, helping un-
 643 derstand how ionospheric outflow contributes to the magnetosphere. We estimate an
 644 ionospheric escape number flux (at an altitude of 10000 km) between $(2.95 \pm 0.43) \times 10^9$
 645 and $(1.43 \pm 0.21) \times 10^{10} \text{ cm}^{-2}\text{s}^{-1}$ one or two orders of magnitude larger than the estimate
 646 obtained by *Glocer et al.* [2007], which represents a mass source between 1.4×10^2 and 1.1
 647 $\times 10^3 \text{ kg/s}$. Furthermore, we estimated the mass source provided by the ionosphere, for
 648 different configurations of the active region at the northern pole. Considering the most
 649 probable scenario as an active region that covers the main auroral oval, we obtain a mass

650 source between 49.7 ± 13.4 and 239.8 ± 64.8 kg/s, comparable with what was found for
651 Enceladus (60-100 kg/s [*Fleshman et al.*, 2013]). Specifically for the auroral morphology
652 found during this event, we obtained a mass source between of 10 ± 4 and 49 ± 23 kg/s.
653 However, we do not have any current estimate on how much of the mass provided by the
654 ionosphere, stays indeed in the magnetosphere and how much instead gets lost downtail.
655 As such, our estimates are an upper limit to the magnetospheric mass source.

656 Future work will include a survey to search for evidence of other ionospheric outflow
657 events at Saturn; besides modelling ionospheric outflow from Saturn's ionosphere with a
658 dynamic magnetosphere and atmosphere is needed to understand the relationship between
659 reconnection or magnetosphere-ionosphere coupling and the outflow from the atmosphere.
660 Moreover, other interesting steps could involve an estimate how much of the mass flowing
661 in the tail from the ionosphere remains in the magnetosphere and how much is lost down-
662 tail. A search for evidence of low energy H_3^+ in the magnetosphere will also provide further
663 evidence for ionospheric outflow. Cassini proximal orbits during its Grand Finale at the
664 end of mission will provide complementary measurements to place better constraints on
665 the ionosphere as a mass source at Saturn. These studies will also provide a valuable
666 context before Juno's arrival at Jupiter.

667 **Acknowledgments.** MF was supported in this work by Il Circolo, the European Space
668 Agency, the Royal Astronomical Society, Agenzia Spaziale Italiana, and the Science and
669 Technology Facilities Council. MF completed this work whilst a visiting student at Lan-
670 caster University and would like to thank the Department of Physics, staff and students,
671 for their kindness and support. MF would like to thank Rob Wilson, K.C. Hansen, Tamas
672 Gombosi, Elias Roussos, Doug Hamilton, Frank Crary, Krishan Khurana, Norbert Krupp

673 and Todd Smith for their comments. CSA was funded by a Royal Society Research Fellow-
674 ship. SVB is supported by a Royal Astronomical Society Fellowship. CMJ was funded by
675 a Science and Technology Facilities Council Ernest Rutherford Fellowship. The research
676 at the University of Iowa was supported by NASA through contract 1415150 with the Jet
677 Propulsion Laboratory.

678 Simulation results have been provided by the Community Coordinated Modeling Cen-
679 ter at Goddard Space Flight Center through their public Runs on Request system
680 (<http://ccmc.gsfc.nasa.gov>). The CCMC is a multi-agency partnership between NASA,
681 AFMC, AFOSR, AFRL, AFWA, NOAA, NSF and ONR. The ENLIL Model was devel-
682 oped by D. Odstrcil at the U. Colorado Boulder. All the data for this study can be found
683 at NASA's planetary data system (<https://pds.jpl.nasa.gov>).

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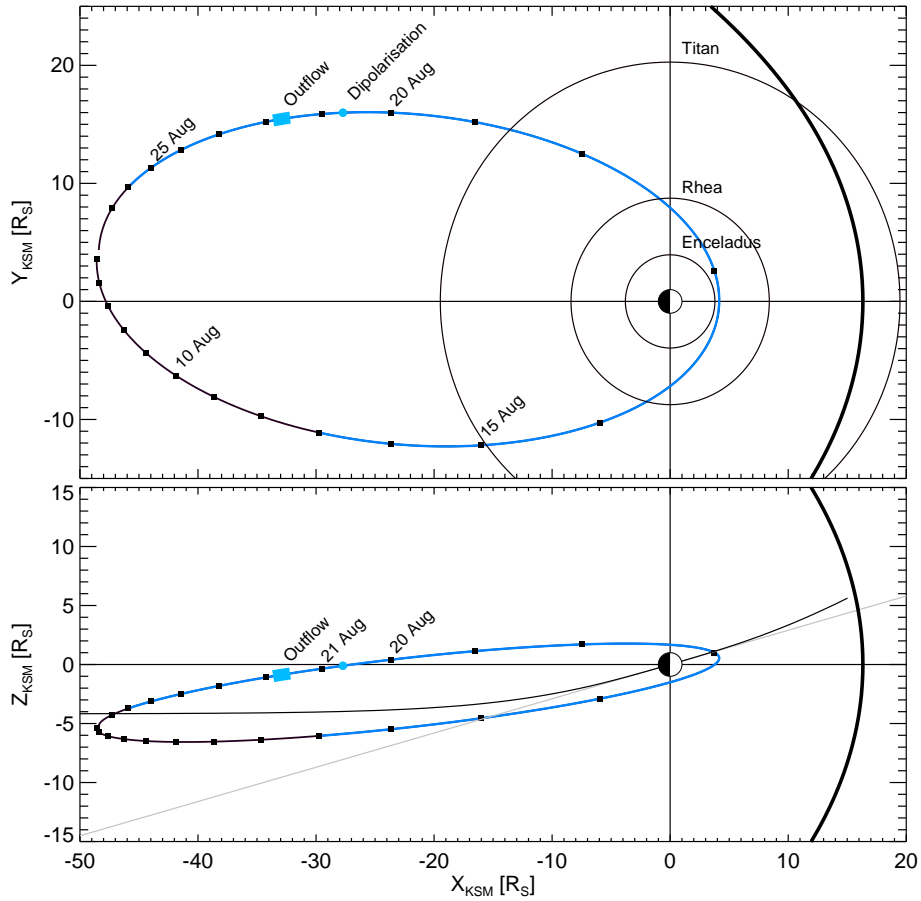


Figure 1. Cassini's trajectory in Kronocentric Solar Magnetospheric Coordinates (KSM) for Cassini's rev. 27 of Saturn from 04 August to 28 August 2006 projected in the X-Y plane (top) and noon-midnight meridional plane (bottom). The segment of the trajectory colored in thin blue lines indicates the time when Saturn was immersed in a solar wind compression region. The dot indicates the dipolarization observed on 20 August and the case study period is indicated by the thick blue line. The magnetopause is shown as the thick black curve in both panels. In the bottom panel the gray line indicates Saturn's equator and the curved black line the average location of Saturn's current sheet.

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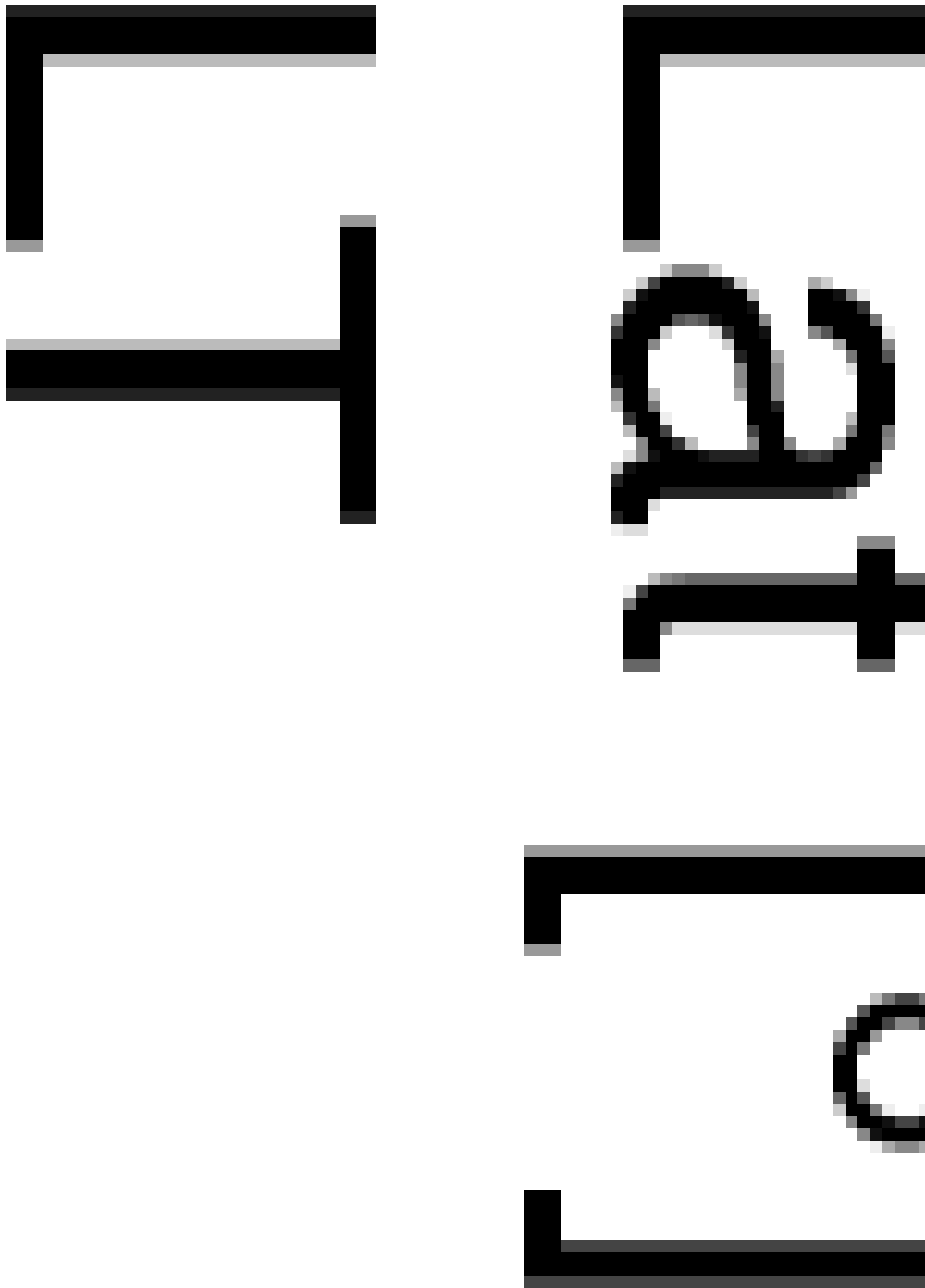


Figure 2. Overview of the interval studied in this paper, from top-to-bottom: panel a) ELS data from anode 5; b) IMS data

averaged over all anodes; **both panels a and b are in differential energy flux units (DEF)** [$\text{m}^{-2}\text{s}^{-1}\text{sr}^{-1}\text{eV eV}^{-1}$]; c-e)

magnetic field components in KRTP coordinates B_r , B_θ , B_ϕ , f) $|\mathbf{B}|$. Intervals when the spacecraft is in the lobes and plasma

sheet are highlighted at the top of the plot. Dotted lines indicate plasma sheet crossings. The arrow indicates the dipolarization

observed on 20 August and discussed in *Jackman et al.* [2015]. The disturbed plasma sheet following the dipolarization is

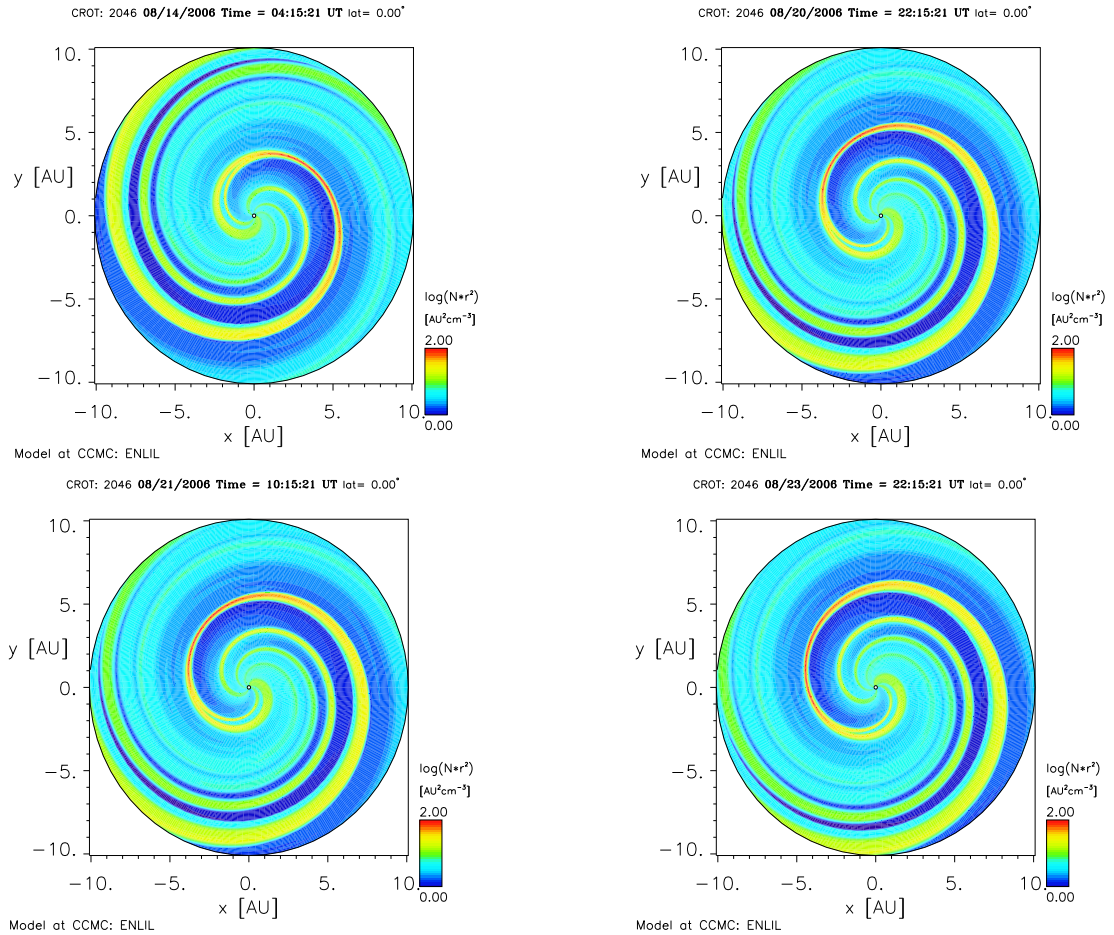


Figure 3. Simulated views of the plasma density (scaled with distance) in the ecliptic plane out to 10 AU. Saturn is near (-10,0) on the left-hand side of these figures. These compare with various times surrounding the compression identified in figure 4 and show the increases in solar wind dynamic pressure as the solar wind compression region rotates over Saturn.

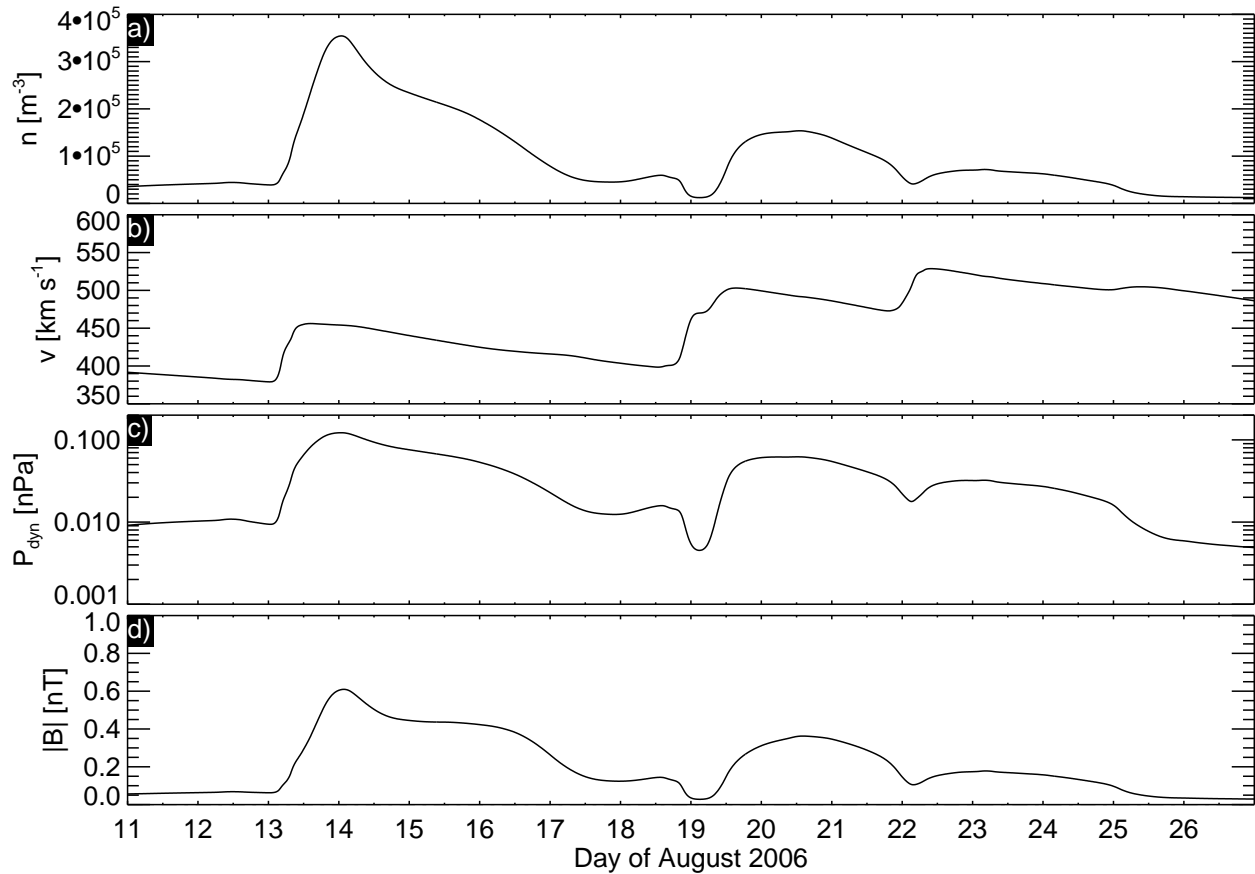


Figure 4. ENLIL simulation results for Carrington Rotation 2046 extracted at Saturn and plotted from 11 August to 27 August 2006. The panels show (a) solar wind plasma number density, (b) solar wind speed, (c) solar wind dynamic pressure, (d) the magnetic field strength in the solar wind.

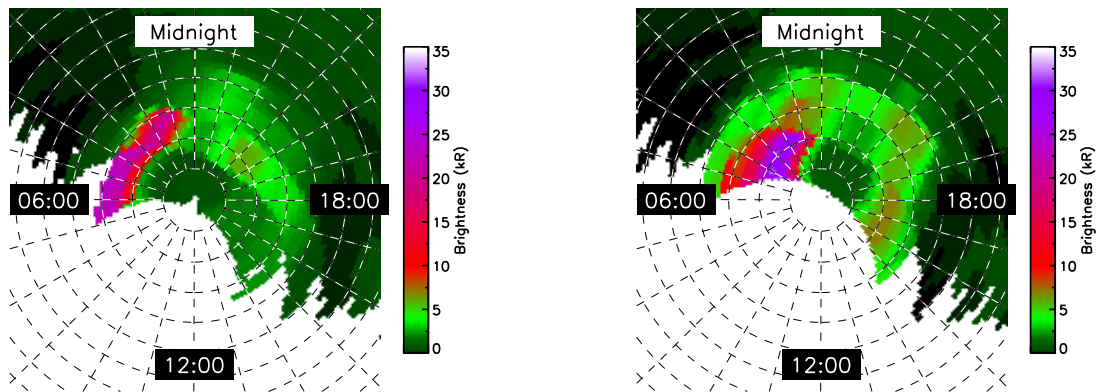


Figure 5. UVIS data for two periods before Cassini moves into the northern lobe from 05:40:16 and from 10:24:06 for about 4 hours of data. The UVIS images are projected onto a colatitude-local time grid where local noon is to the bottom of each figure and dusk to the right. The color scale indicates the intensity of the emission in kR.

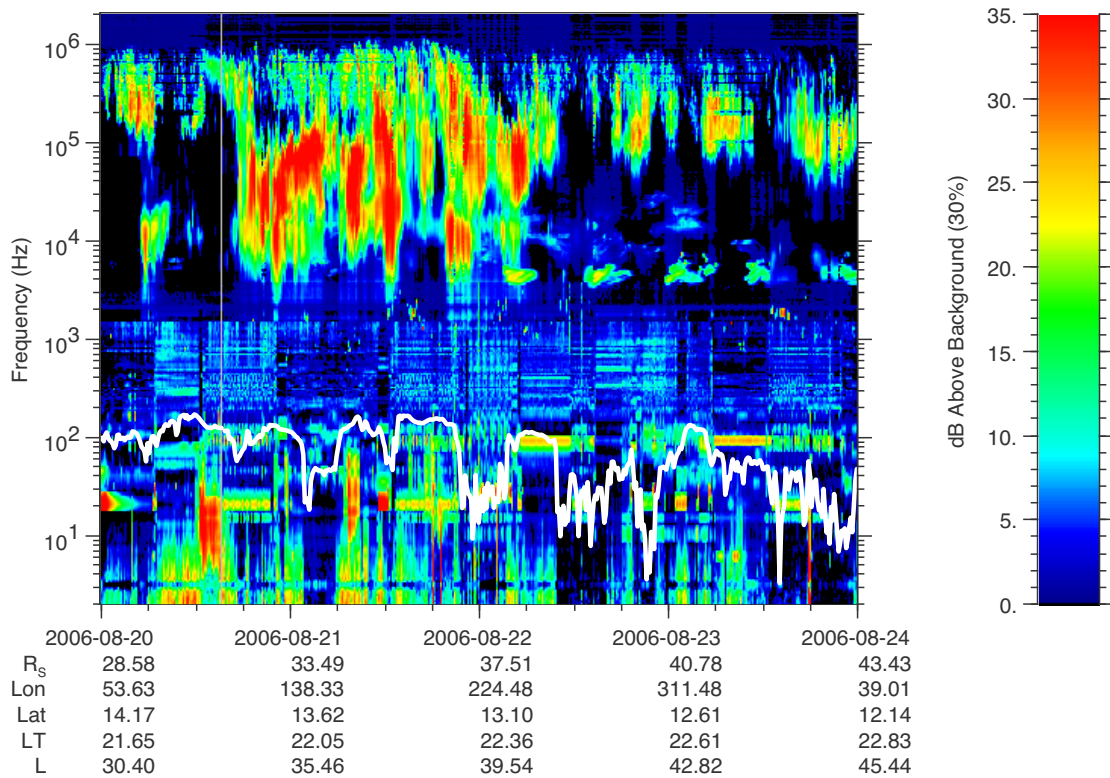


Figure 6. RPWS electric field data from 20 August to 23 August 2006 showing the received power relative to the background with the electron cyclotron frequency (calculated from the magnetic field strength) overlaid in white.

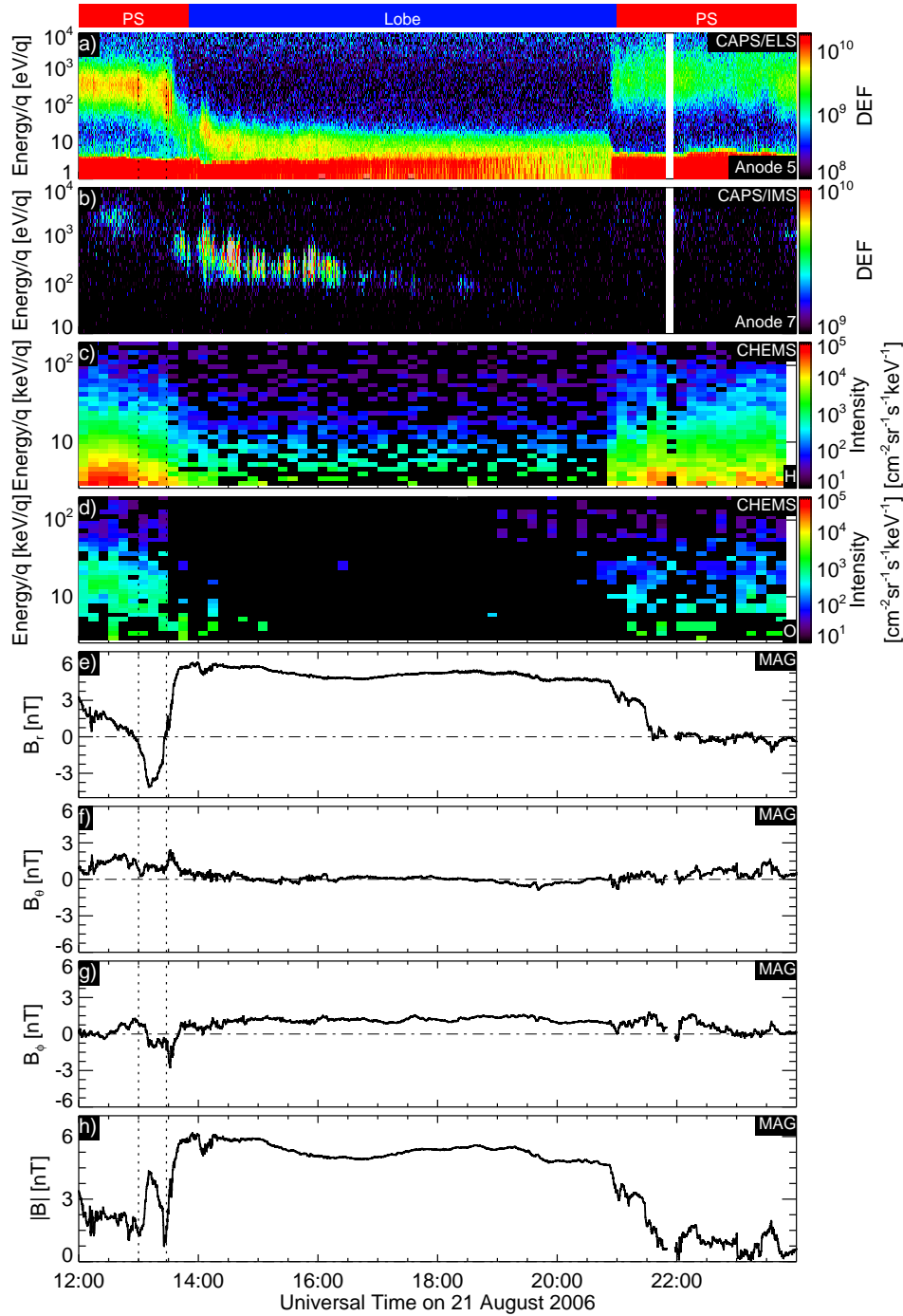


Figure 7. Interval studied in this paper, from top-to-bottom: panel a) ELS in anode 5, b) IMS data from anode 7. **Both panels a and b are in differential energy flux units (DEF) [$\text{m}^{-2}\text{s}^{-1}\text{sr}^{-1}\text{eV eV}^{-1}$];** c) CHEMS energetic H^+ spectrogram, d) CHEMS energetic W^+ spectrogram, e-g) magnetic field components in KRTP coordinates B_r , B_θ , B_ϕ , h) $|\mathbf{B}|$. The two dotted lines indicates the two plasma sheet crossing before the spacecraft moves in the northern

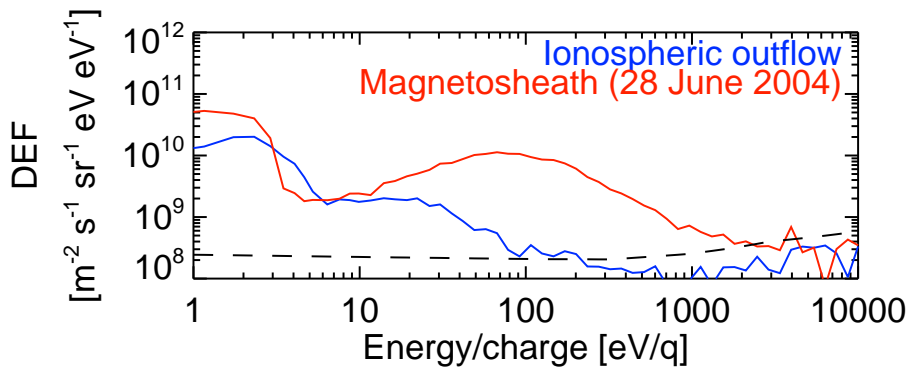


Figure 8. Comparison between two electron spectra: electron spectrum from this case study, 21 August 2006, 1400 UT (blue line), electron spectrum from 28th June 2004, 1400 UT, a typical magnetosheath spectrum (red line), with superimposed (dashed line) the one count level. Both spectra are from anode 5 in Cassini ELS. The magnetosheath population seen below $\simeq 3$ eV are photoelectrons, whereas in our case study we find photoelectrons below $\simeq 8$ eV.

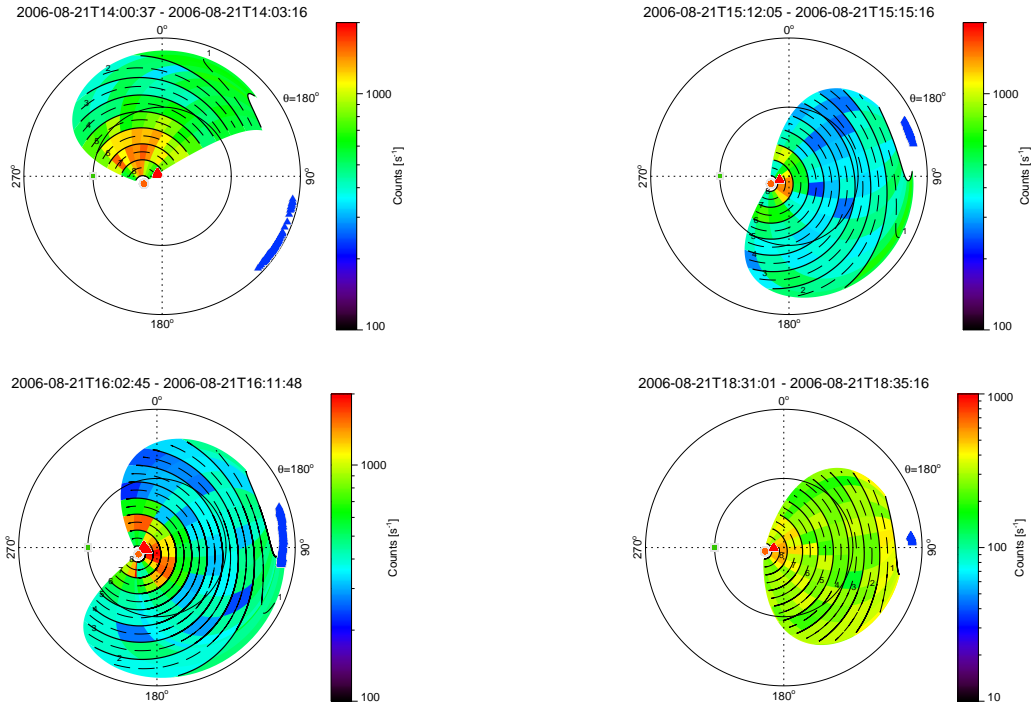
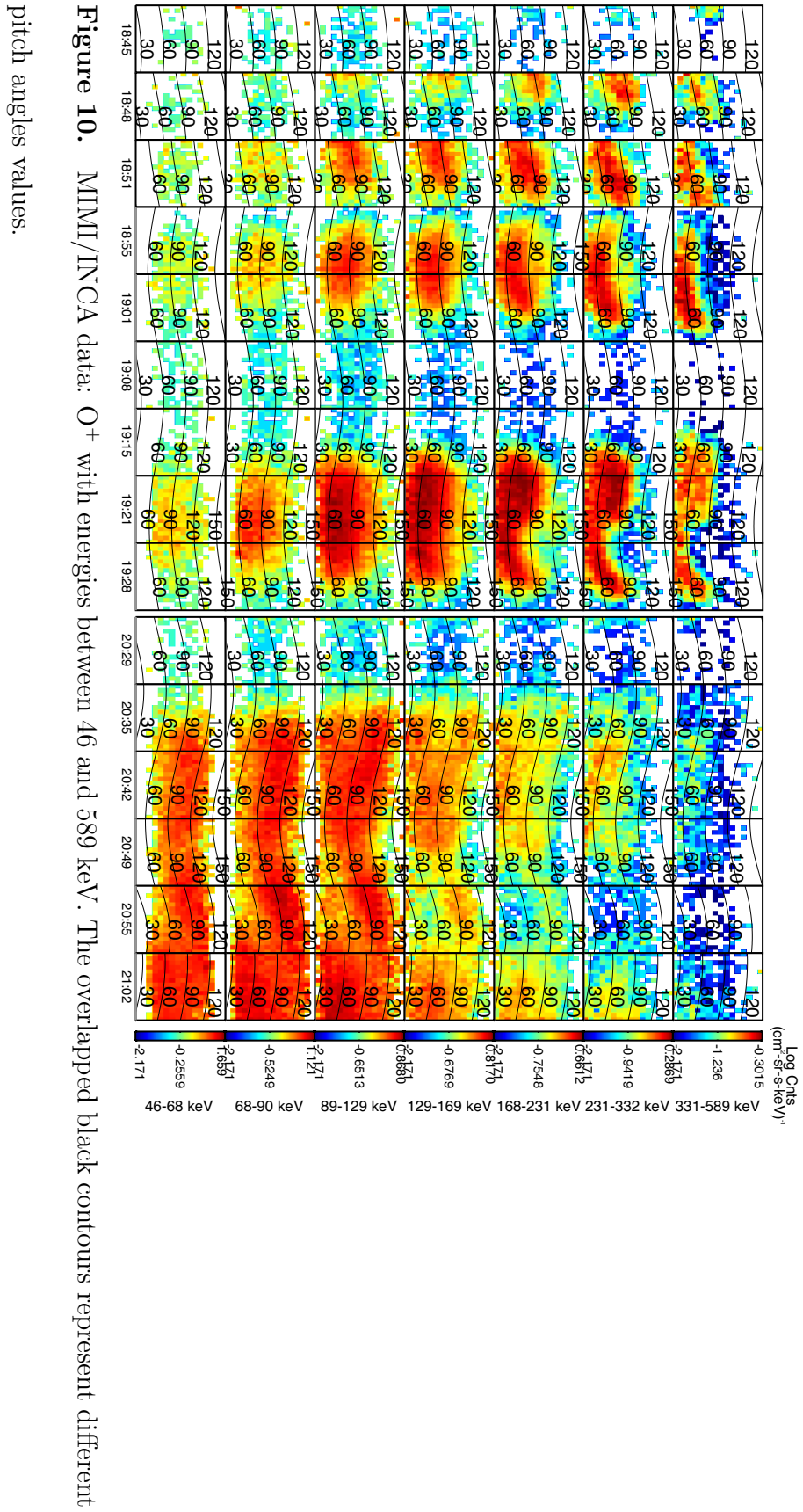


Figure 9. Measured ion counts as a function of look direction about the spacecraft for four different times in the northern lobe. In each case ions are observed flowing tailward from the direction near Saturn. The energy/ q ranges are respectively, from left to right, from top to bottom, 181.1-2048, 107.7-724.1, 107.7-724.1, 90.40-361.9 eV/ q .



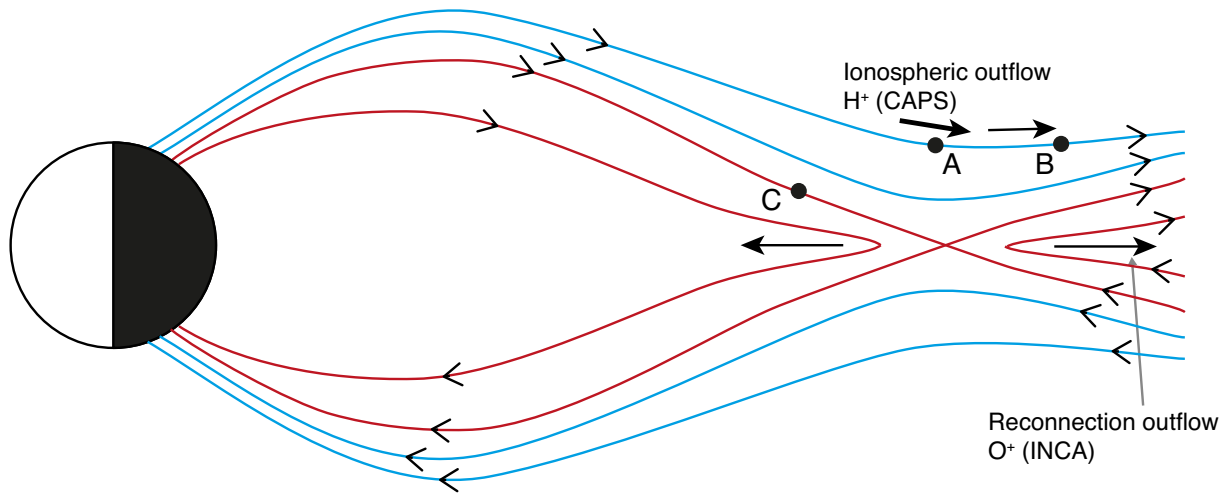


Figure 11. In the schematic we represent different field lines with different colors. In point A the spacecraft is in the north lobe, only detecting ionospheric ions in CAPS/IMS. The ionospheric ions are represented with arrows, that get thinner approaching point B, in order to represent the dispersion. When the spacecraft is located in B, from 1845, cold ions from ionosphere are still detected in CAPS/IMS but meantime MIMI/INCA remote senses hot O^+ from the plasma sheet. O^+ is flowing downtail accelerated by reconnection, and is represented in the plot through an arrow. After 2100 the spacecraft returns to the plasma sheet (point C).

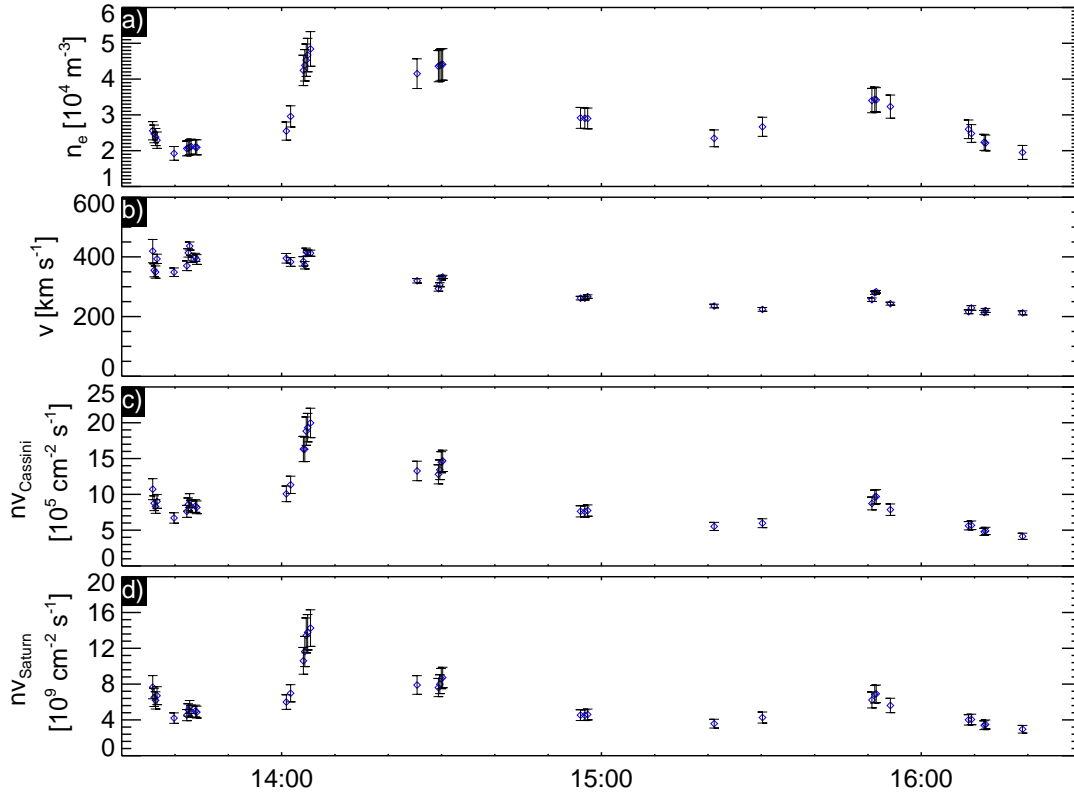
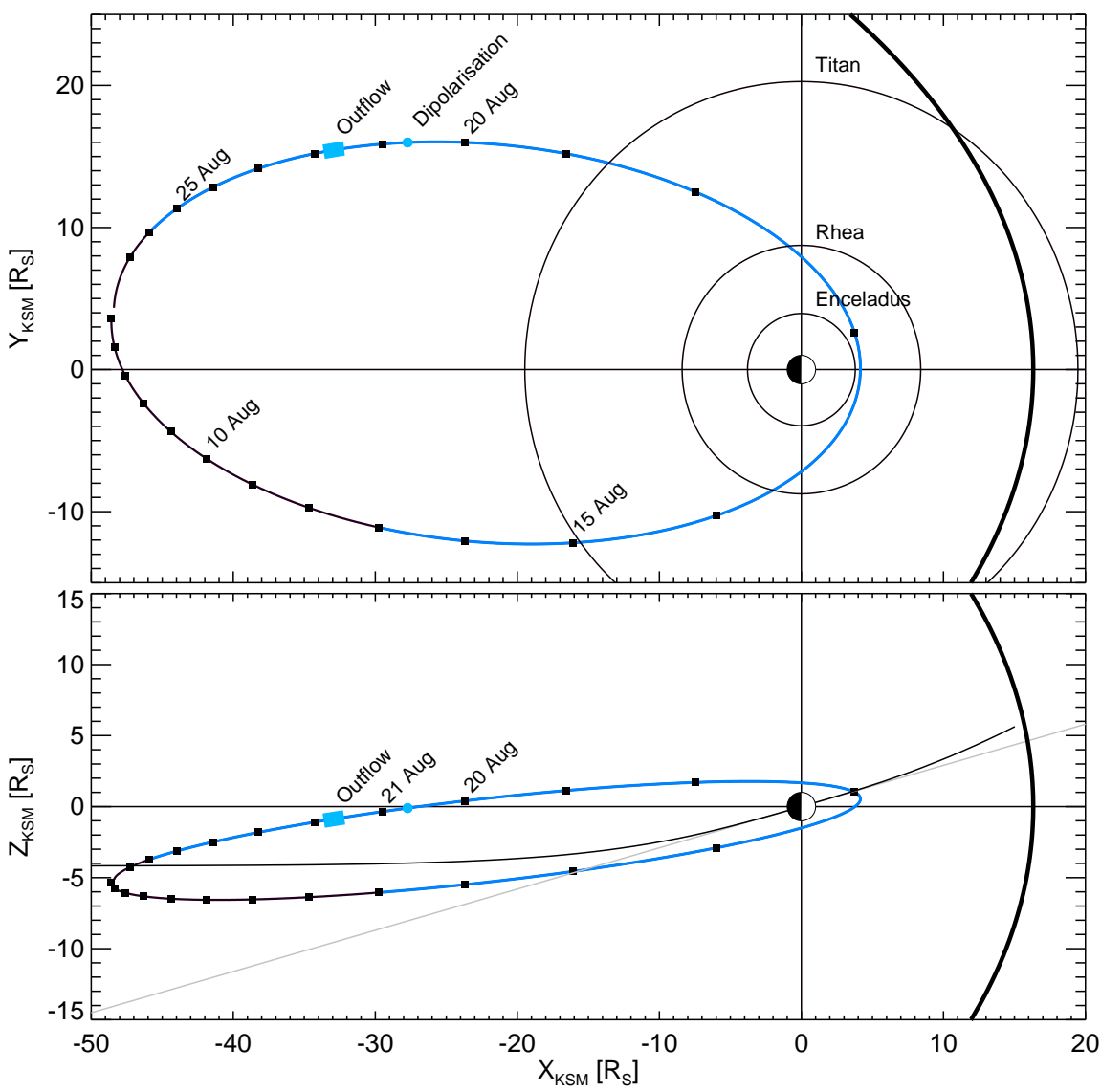
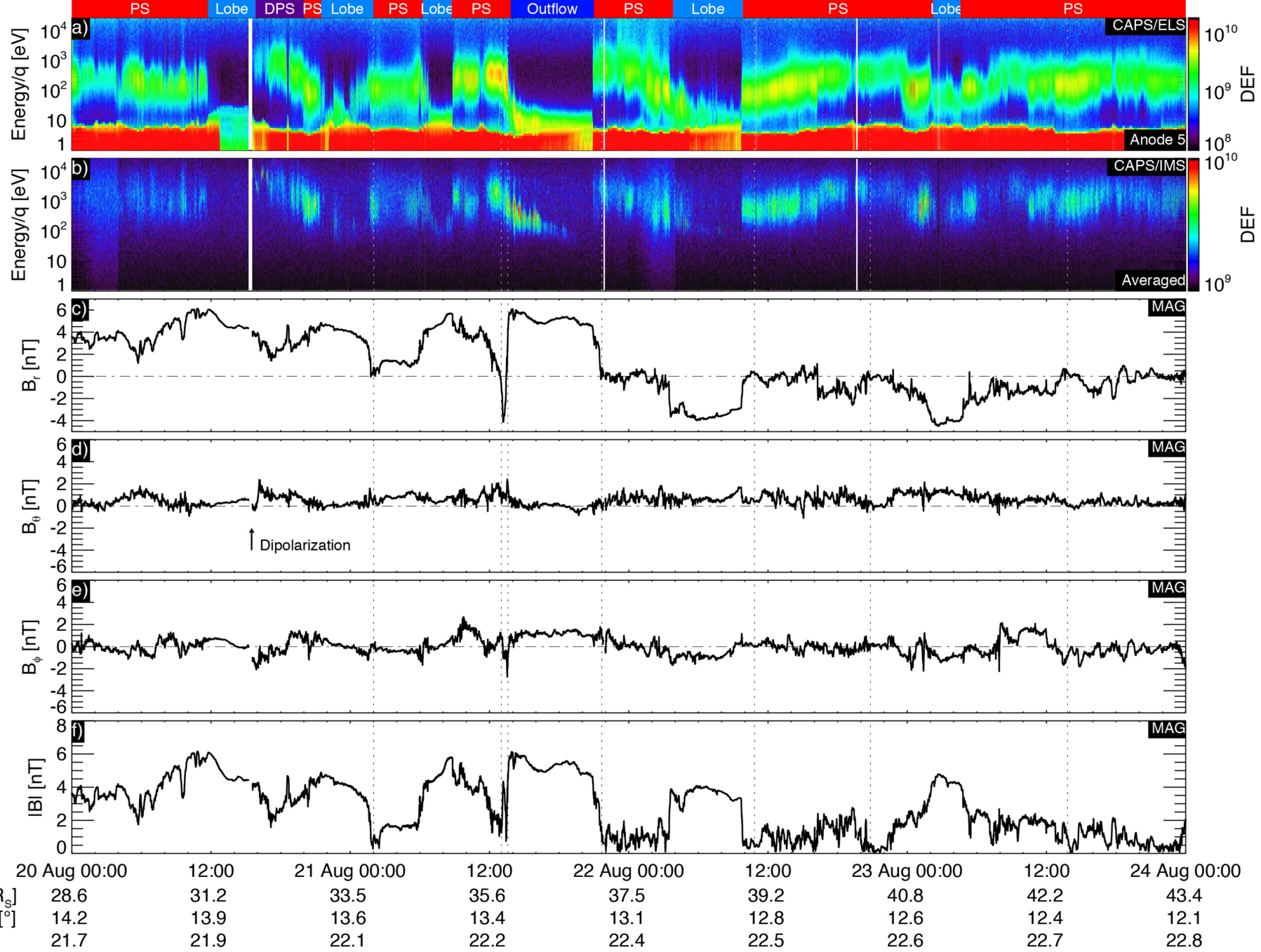
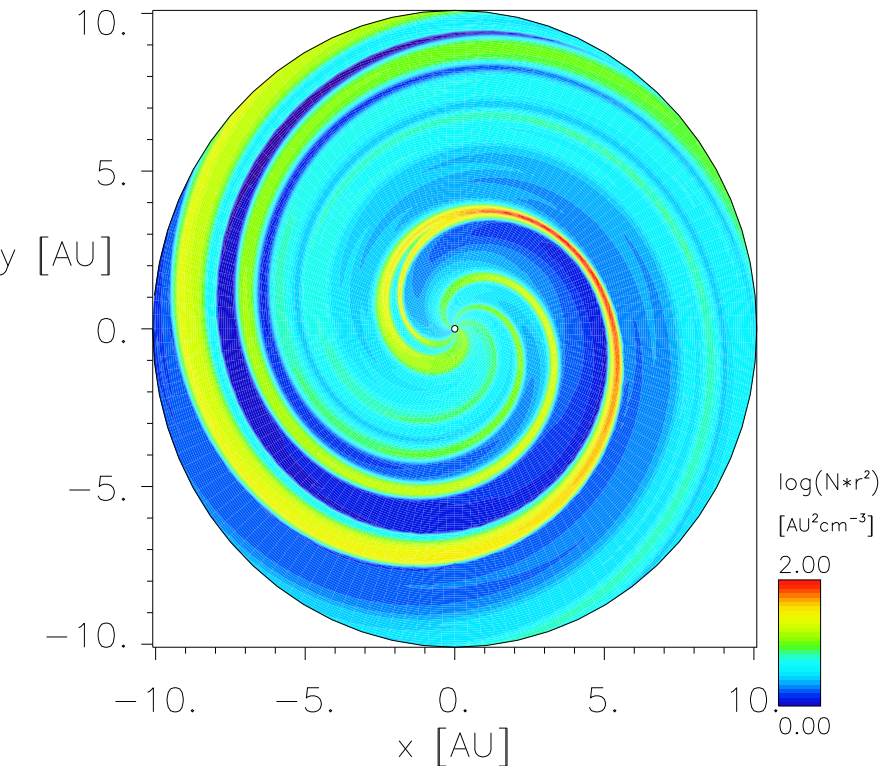


Figure 12. Estimates of the number flux associated with the polar wind. From top-to-bottom: a) the measured electron density at Cassini, b) the estimated ion speed at Cassini, c) the estimated number flux at Cassini, d) the number flux estimated at an altitude of 10000 km above Saturn at 78° latitude.



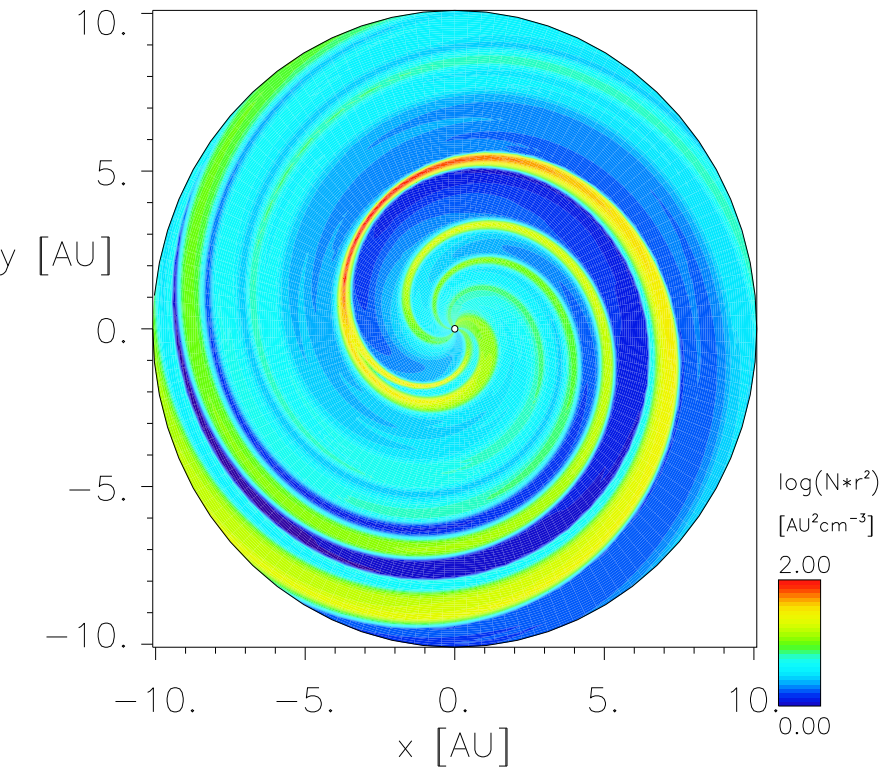


CROT: 2046 08/14/2006 Time = 04:15:21 UT lat= 0.00°



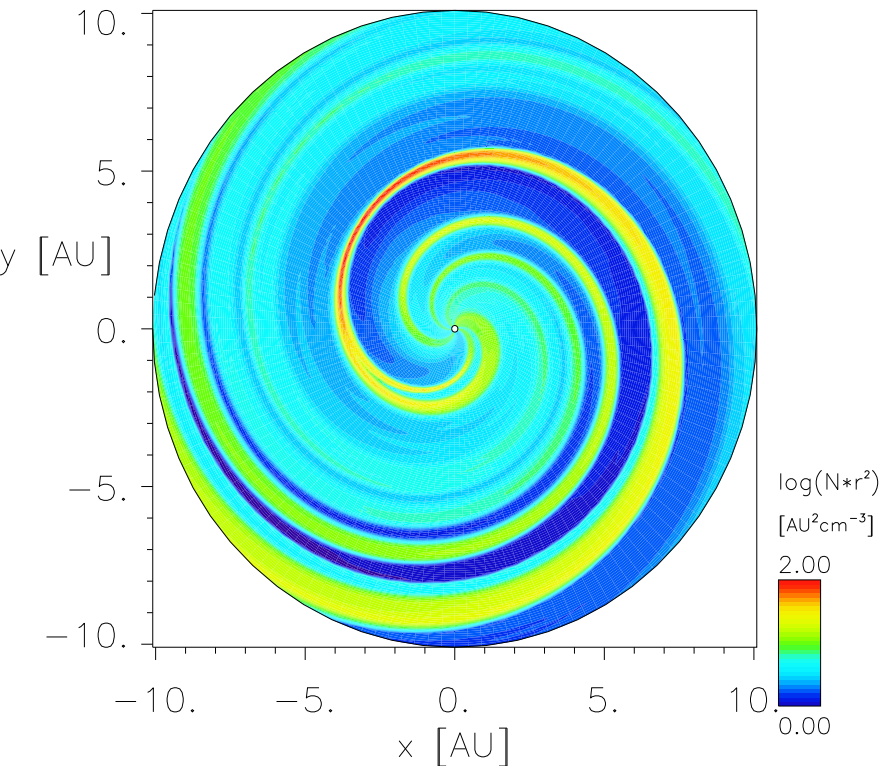
Model at CCMC: ENLIL

CROT: 2046 08/20/2006 Time = 22:15:21 UT lat= 0.00°



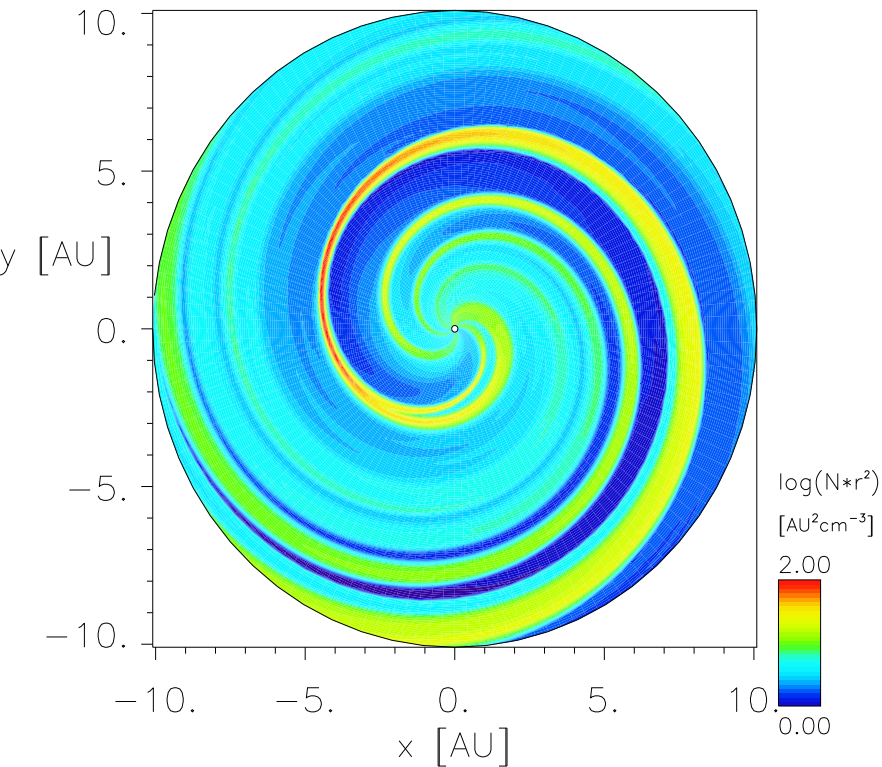
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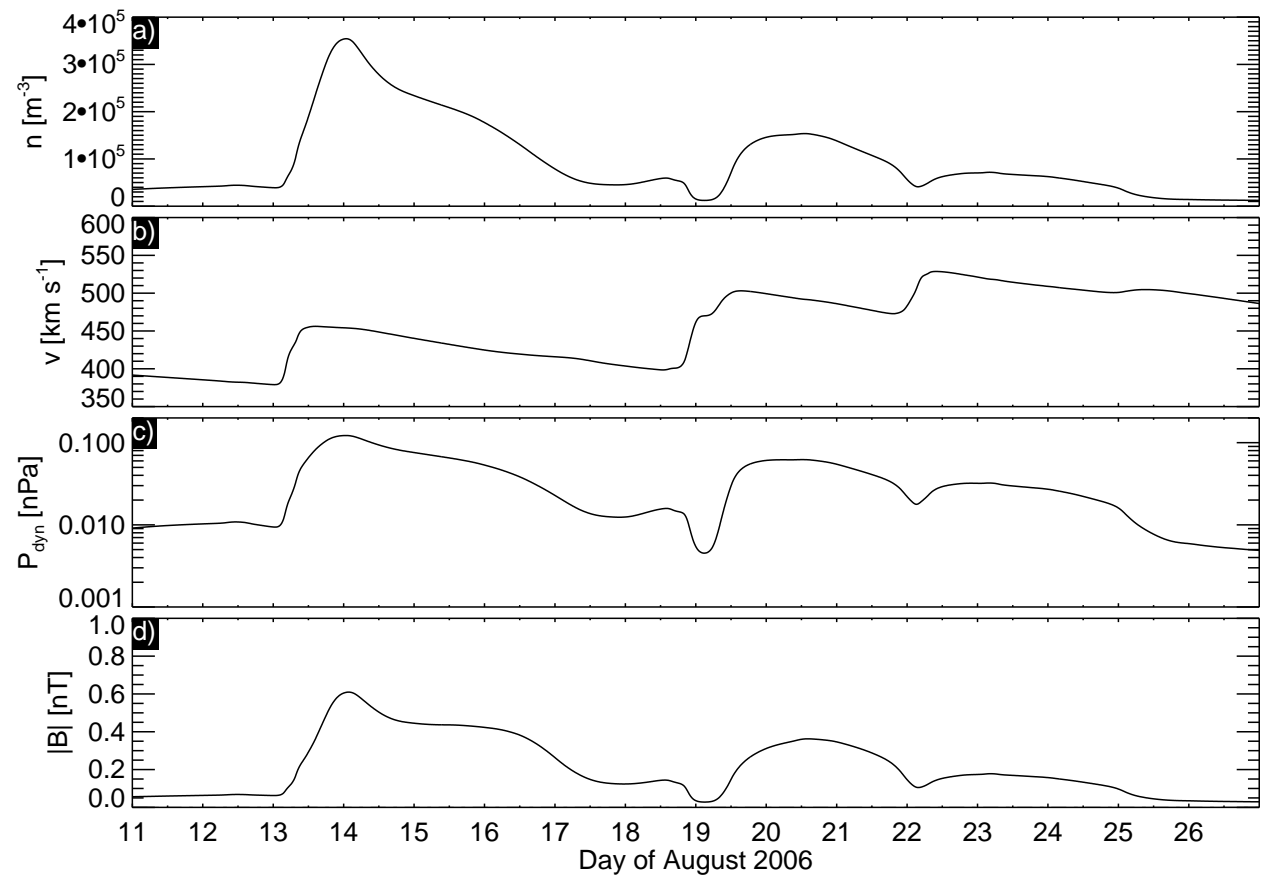


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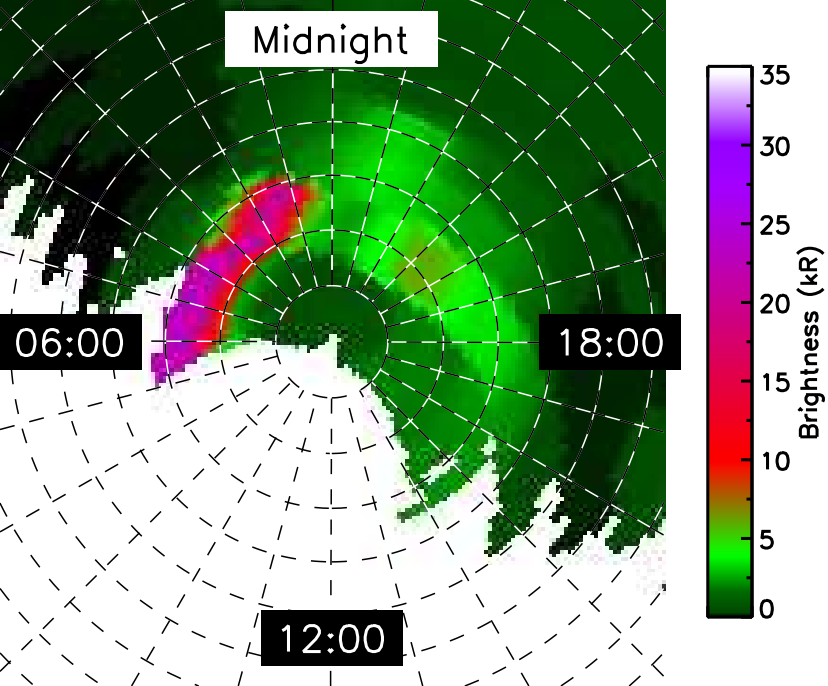
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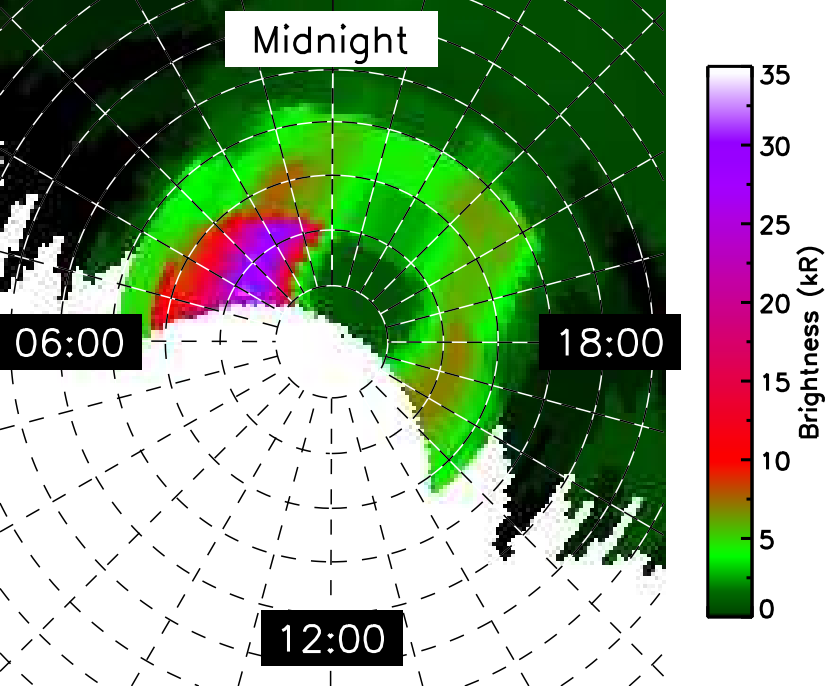
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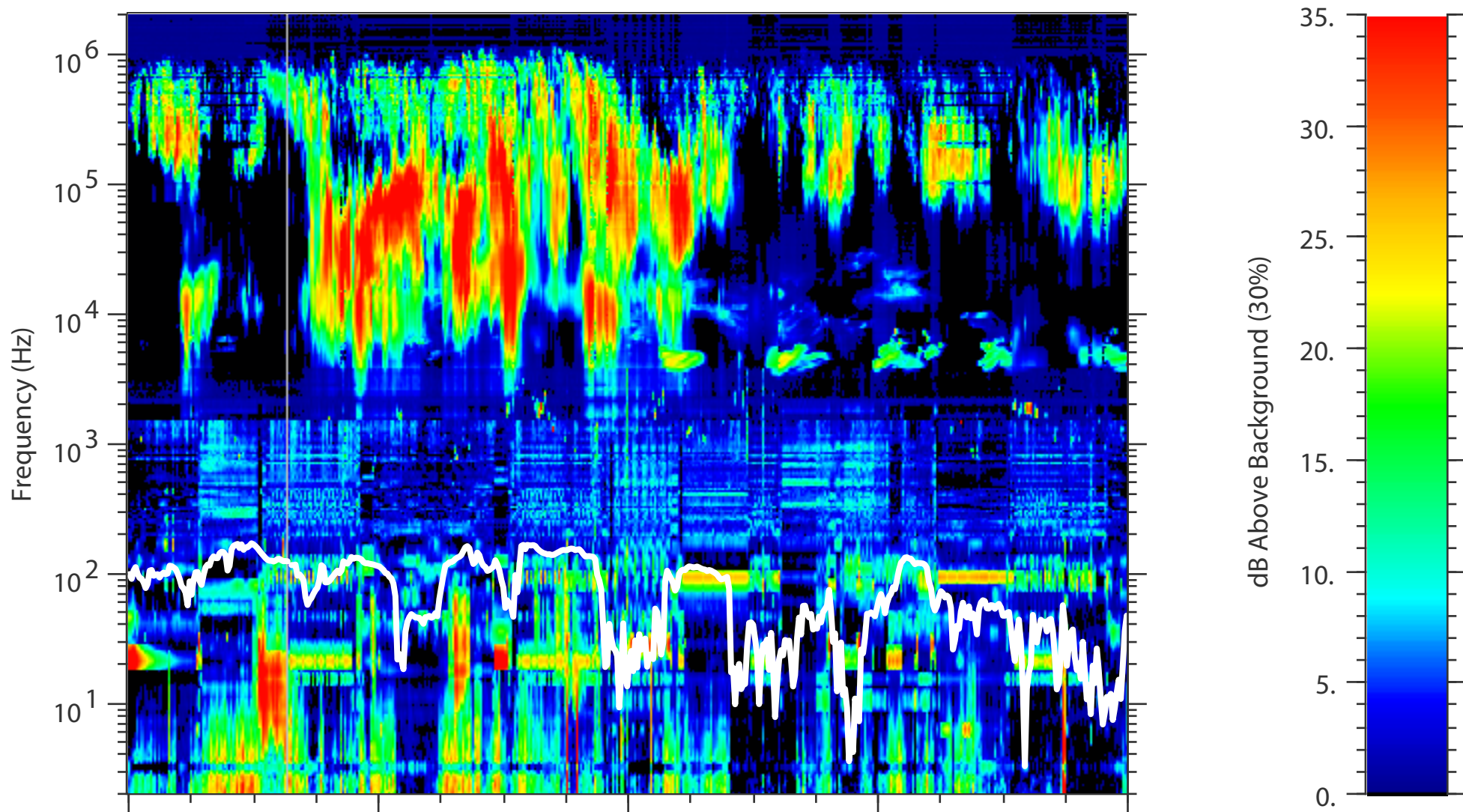


Midnight

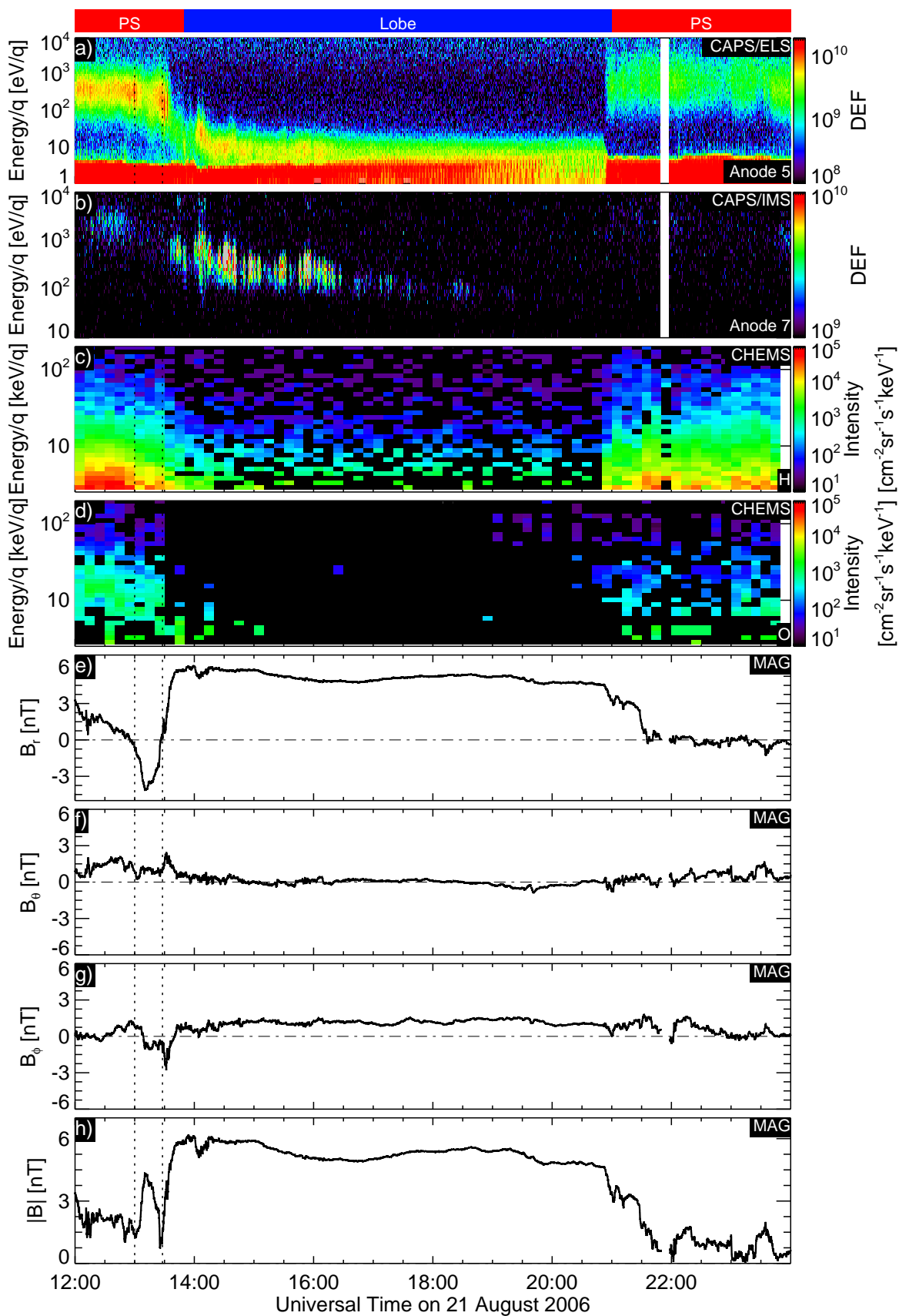


Midnight



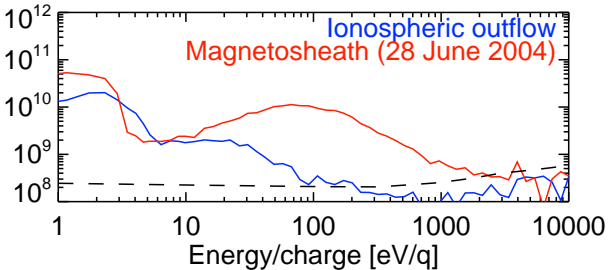


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R_s	28.58	33.49	37.51	40.78	43.43
Lon	53.63	138.33	224.48	311.48	39.01
Lat	14.17	13.62	13.10	12.61	12.14
LT	21.65	22.05	22.36	22.61	22.83
L	30.40	35.46	39.54	42.82	45.44

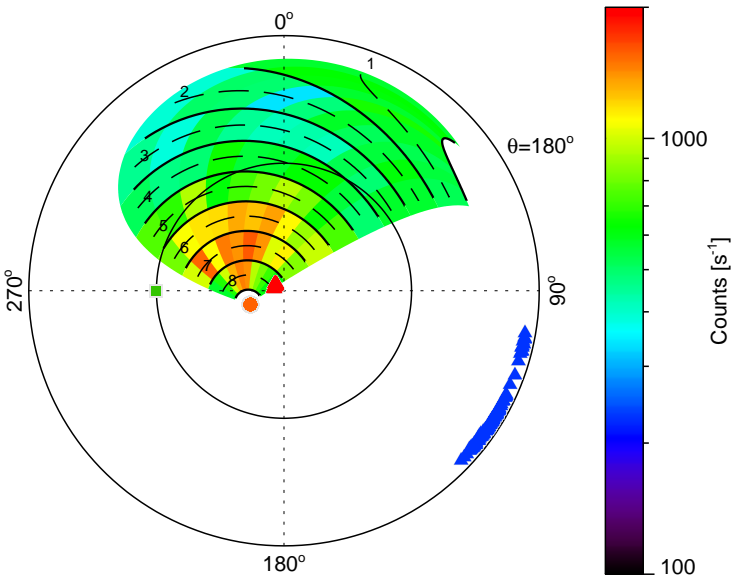


DEF

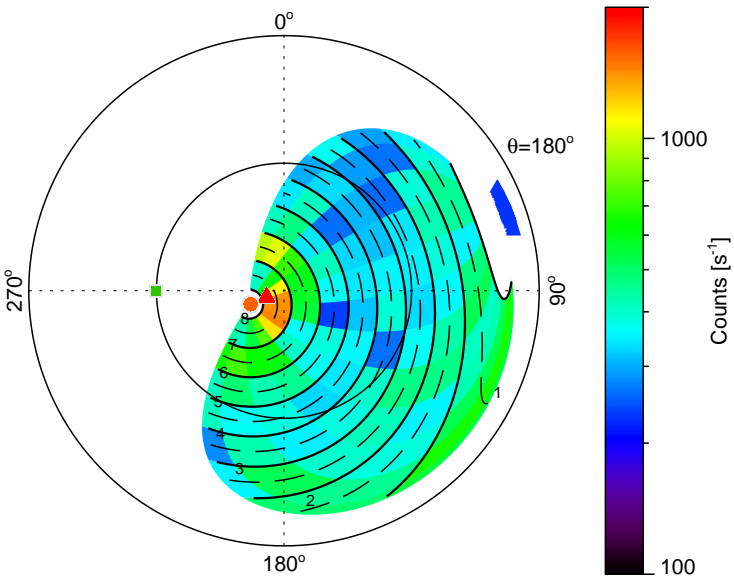
$[\text{m}^{-2} \text{s}^{-1} \text{sr}^{-1} \text{eV}^{-1}]$



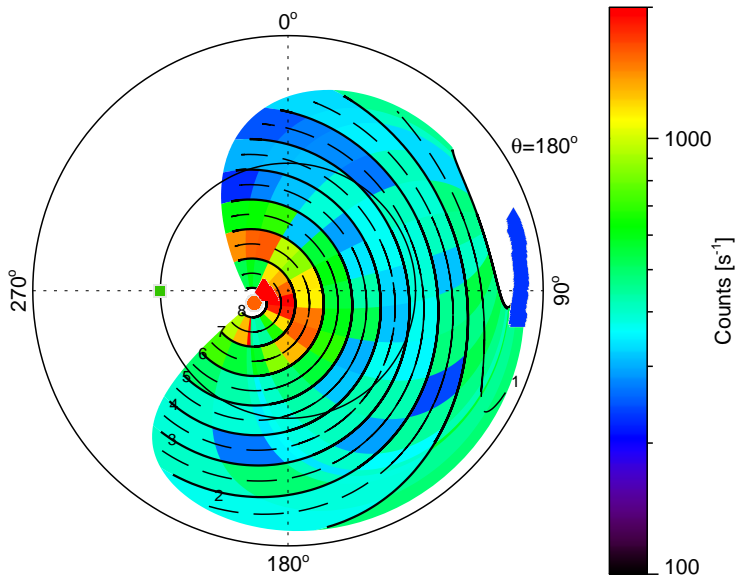
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2006-08-21T18:31:01 - 2006-08-21T18:35:16

