

Article

Evaluation of a Global Soil Moisture Product from Finer Spatial Resolution SAR Data and Ground Measurements at Irish Sites

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Abstract: In the framework of the European Space Agency Climate Change Initiative, a global, almost daily, soil moisture (SM) product is being developed from passive and active satellite microwave sensors, at a coarse spatial resolution. This study contributes to its validation by using finer spatial resolution ASAR Wide Swath and *in situ* soil moisture data taken over three sites in Ireland, from 2007 to 2009. This is the first time a comparison has been carried out between three sets of independent observations from different sensors at very different spatial resolutions for such a long time series. Furthermore, the SM spatial distribution has been investigated at the ASAR scale within each Essential Climate Variable (ECV) pixel, without adopting any particular model or using a densely distributed network of *in situ* stations. This approach facilitated an understanding of the extent to which geophysical factors, such as soil texture, terrain composition and altitude, affect the retrieved ECV SM product values in temperate grasslands. Temporal and spatial variability analysis provided high levels of correlation ($p < 0.025$) and low errors between the three datasets, leading to confidence in the new ECV SM global product, despite limitations in its ability to track the driest and wettest conditions.

Keywords: ESA Climate Change Initiative; Essential Climate Variable; soil moisture; ENVISAT ASAR WS; temporal variability; spatial variability

1. Introduction

Although only a small percentage (~1%, [1]) of the total freshwater budget is contained in the soil layer of the Earth's surface, it is a key parameter in the exchange of mass and energy at the soil-atmosphere boundary, as well as in eco-hydrological processes [2–5]. Soil moisture (SM) controls the partitioning of incoming radiant energy into latent and sensible heat fluxes through evaporation and transpiration [6,7]. Moreover, surface runoff and infiltration of precipitation are strongly related to the soil moisture content, which, as a result, has important control over the biogeochemical cycling of nutrients [8], the availability of transpirable water for plants and groundwater recharge. Hence, the knowledge of soil moisture behaviour is of vital importance for flood and drought forecasting, water resource management, irrigation, weather prediction and for global climate change research [9].

For these reasons, the Global Climate Observing System (GCOS) secretariat has identified soil moisture as an “Essential Climate Variable” (ECV), thereby requiring enhanced observation of its spatial and temporal variability in support of the United Nations Framework Convention on Climate Change (UNFCCC). It has been observed that particular meteorological conditions, geological characteristics, topography and land cover can affect the soil moisture variation in a small area as much as in a large region [10,11]. Moreover, considering the top layer of the soil, which is more subject to the influences of the atmosphere (than deeper soil layers), soil moisture can change significantly within a few hours [12]. Such high spatial variability and temporal dynamics make the monitoring of soil moisture over wide areas challenging.

Given the increasing interest of the scientific community in understanding the relationship between soil moisture and climate change, the *in situ* stations network is currently expanding worldwide [13]. Nevertheless, mapping local- and regional-scale variations in soil moisture requires a high spatial density of observations over time, making *in situ* measurements time consuming and costly. On the other hand, remote sensing represents a useful tool, as satellites can regularly provide information over large areas [14].

A first global soil moisture product meeting the requirements set by the GCOS was created within the framework of the European Space Agency (ESA) Water Cycle Multi-mission Observation Strategy (WACMOS) project [15], by merging soil moisture products derived from multi-frequency radiometer and C-band scatterometer observations into a single dataset covering the period from 1979 to 2010 [16–18]. Global soil moisture maps are provided on an almost daily basis and with a spatial resolution of 0.25 degrees.

Building on the WACMOS project, the ECV SM global time series is currently being extended and enhanced, by improving the retrieval and merging algorithms, in the context of the ESA-funded Climate Change Initiative (CCI) programme [19].

Despite the advantageous daily frequency of such products, there is still a relatively coarse spatial resolution. The assessment of soil moisture retrieved from satellite acquisitions is commonly carried

out through a comparison with *in situ* measurements. In [20], promising results have been found by validating three global soil moisture products, including the WACMOS time series, through a comparison with *in situ* measurements taken at 196 stations from five networks across the world. While the satellite measurements represent hundreds of km² of a more variable soil layer (0–5 cm), the *in situ* measurements describe the smaller areas (few dm²) of a deeper stratum.

However, the higher spatial resolution and the regular coverage provided by spaceborne Synthetic Aperture Radars (SARs) makes them a promising approach for measuring monthly, seasonal and long-term variations in surface soil moisture content [21,22]. In addition, the characterization of the errors inherent in the coarser spatial resolution ECV products can be studied. On the other hand, the retrieval of soil moisture by using SAR acquisitions is a non-trivial task. In fact, not only the soil moisture content, but also the speckle noise, terrain typology, soil roughness, vegetation cover and incidence angle affect the microwave backscattering. All of these factors must be taken into account to reduce the uncertainty and to improve the accuracy of the soil moisture estimate. In addition, the characterization of the errors inherent in the coarser spatial resolution ECV products can be studied. The comparison of soil moisture time series datasets acquired by different sensors and representing different spatial scales remains challenging due to the large-scale differences [23,24]. Nevertheless their inter-comparisons are still useful owing to the temporal stability of soil moisture patterns [25], where soil moisture values at smaller scales are representative of the mean soil moisture content over larger areas.

The opportunities to characterize large areas through the use of higher spatial resolution (hundreds of metres to several kilometres) image data acquired every few days make the ENVISAT Advanced Synthetic Aperture Radar (ASAR) sensor a suitable tool for improving the understanding of the ECV SM global products. In [26–35], the capability of ASAR backscattering to estimate the surface soil moisture has been demonstrated.

In this study, the ASAR Wide Swath (WS) mode of ASAR images (150-m spatial resolution) acquired at VV polarization have been used for the retrieval of soil moisture in Southern Ireland, where three *in situ* stations have been taken into account. The overall aim of the study was to determine how representative the ECV SM product is of the SM condition, both spatially and temporally. To achieve this, temporal and spatial analyses of the soil moisture are examined with the goals of:

- (1) Improving the understanding of the temporal variability of the difference between soil moisture values derived from the different sensors.
- (2) Improving the understanding of the main factors affecting the soil moisture spatial variability and ECV SM values.

Firstly, in order to analyse the soil moisture temporal behaviour, the derived ASAR pixel-based information has been averaged over the corresponding ECV pixel, which includes a single *in situ* station. Then, a triple comparison between ASAR WS, ECV and *in situ* soil moisture time series has been carried out. The spatial variability of the soil moisture within the ECV size pixel has been successively studied by exploiting the finer spatial resolution of the ASAR WS images. Ancillary data, such as a European Digital Elevation Model (EU-DEM) [36], the Irish soil map provided by the

Environmental Protection Agency (EPA) and the Corine Land Cover map (CLC2006), have been used to interpret the spatial and temporal soil moisture information retrieved.

2. Test Site Description

The study focused on the south of Ireland using three *in situ* sites, namely Kilworth (KW), Pallaskenry (PK) and Solohead (SH) (Figure 1). The region is characterized by a humid mild temperate climate, with a mean annual precipitation of $\sim 1200 \text{ mm}\cdot\text{y}^{-1}$. The *in situ* stations are installed in grassland areas, which represent almost 80% of the agricultural area of Ireland (4.4 million hectares) [37]. The region is typically low lying, with altitudes ranging between 15 m and 104 m above sea level, and relatively flat (slope lower than 6°). On the basis of the United States Department of Agriculture (USDA), the soil texture is classified as sandy loam in Kilworth and as loam in Pallaskenry and Solohead. In Table 1, the specific characteristics of each location are reported.

Figure 1. ASAR WS image of the area under study. The location of each *in situ* station used in this work is also shown within each Essential Climate Variable (ECV)-sized pixel ($^\circ$ Google Earth).

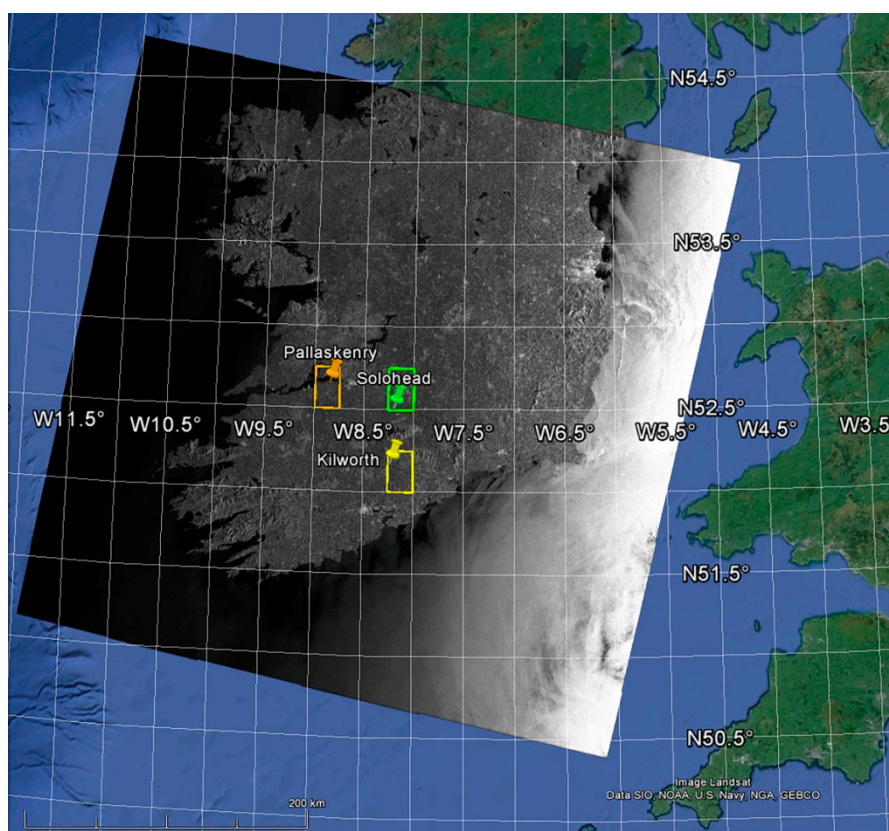


Table 1. Test site soil characteristics. KW, Kilworth; PK, Pallaskenry; SH, Solohead.

Site	Latitude	Longitude	Bulk Density (g/cm^3)	Sand%	Silt%	Clay%	Porosity%
KW	$52^\circ 10' \text{N}$	$-8^\circ 14' \text{E}$	1.08 ± 0.07	52.68 ± 1.38	25.50 ± 2.91	17.81 ± 1.52	59 ± 3.82
PK	$52^\circ 39' \text{N}$	$-8^\circ 51' \text{E}$	1.03 ± 0.10	49.50 ± 0.25	31.50 ± 1.61	19 ± 1.36	61 ± 5.92
SH	$52^\circ 30' \text{N}$	$-8^\circ 12' \text{E}$	0.99 ± 0.14	45.85 ± 1.32	31.90 ± 2.90	22.25 ± 1.57	63 ± 8.90

3. Material and Methods

3.1. In Situ SM Data

Soil moisture *in situ* measurements have been collected using time domain reflectometers (TDR), Campbell Scientific CS616, installed horizontally at each of the three study sites. The instruments have not been calibrated with the soil texture, but a standard manufacturer calibration closest to the soil type has been used. This indicates that while the absolute soil moisture may not be exact, the relative soil moisture and changes of soil moisture from day to day are correct. The instrument manufacturer's accuracy is $\pm 2\%$. Measurements have been recorded at 30-min intervals, together with precipitation and soil temperature, since 2007.

The *in situ* instruments provide soil moisture values to a depth of 5 cm from the surface and are expressed as the soil water-filled pore space (WFPS). The ECV SM data are in volumetric units ($\text{m}^3 \cdot \text{m}^{-3}$). The *in situ* values have been converted to the same units, by using the porosity f as in Equation (1):

$$\theta_v = \text{WFPS} \frac{f}{100} \quad (1)$$

3.2. Remotely Sensed Data

3.2.1. ECV SM Product

In the first version of the global ECV SM product, released in June 2012, by the Vienna University of Technology (TUW) and freely available at [19], soil moisture has been retrieved from several microwave multi-frequency radiometers (Scanning Multichannel Microwave Radiometer (SMMR), Special Sensor Microwave/Imager (SSM/I), Tropical Rainfall Measuring Mission's (TRMM) Microwave Imager (TMI) and Advanced Microwave Scanning Radiometer for Earth Observation System (ASMR-E), and C-band scatterometers European Remote Sensing (ERS), Advanced SCATterometer (ASCAT)). The Water Retrieval Package (WARP) developed at TUW has been used to derive the soil moisture from active system acquisitions, while all passive products were generated using the VUA-NASA (Vrije Universiteit Amsterdam and NASA) land Parameter Retrieval Model (LPRM) software package. Secondly, two homogenized products have been generated from all of the passive and active data [38]. The error characterization of the merged passive and merged active product has been carried out through triple collocation and error propagation modelling [39]. Finally, active-passive datasets have been merged in a unique soil moisture product [16–18,40]. The rationale behind this approach is that calibration differences and other structural biases between the products in the individual product groups can be removed, before merging two products based on completely different retrieval principles. The merging procedure involves firstly rescaling the different products into common reference climatology by using AMSR-E and ASCAT-based soil moisture data as a reference. Such systems provide the most reliable climatology for the individual product groups. To supply a global coverage, a supplementary dataset is needed to scale the merged active and passive dataset into a globally-consistent climatology. Hence, the GLDAS-Noah (Global Land Data Assimilation System) data assimilation system [41] is used. Scaling is performed using cumulative

distribution function (cdf) matching [16,42–44]. The drivers of the merging approach are data availability and the relative accuracy of the products.

Data availability (at a daily time step) and data sensitivity to vegetation are taken into account to combine the merged active and merged passive datasets. The average vegetation optical depth (VOD) over transitional regions (*i.e.*, regions between sparsely and moderately vegetated areas) is calculated and used as the threshold for separating sparsely- from moderately-vegetated regions outside of the transition zones. Active soil moisture data are used for regions with moderate vegetation density (a VOD value higher than the threshold), whereas the passive product is used for (semi-) arid regions (a VOD value lower than the threshold) [38]. When, at a given location in transition zones, the correlation coefficient (*R*) between the merged active and passive soil moisture products is greater than 0.65, both products are used [39]. This is done by simply averaging merged passive and merged active products for time steps where both products are available; if only one product type is available, that one is used [38]. In [16], it has been shown that this procedure increases the number of observations while minimally changing the accuracy of the merged soil moisture product. Moreover, it has been observed that when the active and passive merged datasets at a given location have an *R* lower than 0.65, using them both reduces the quality of the merged product relative to the single products.

The SM time series maps are provided on an almost daily basis, covering a period of more than thirty years, from 1979 to 2010. The spatial resolution of the ECV SM product is approximately 0.25 degrees. Data are provided in volumetric units ($\text{m}^3 \cdot \text{m}^{-3}$), together with quality flags indicating the possible presence of dense vegetation, snow or a temperature below 0 °C. The ECV data used in this study are ASCAT acquisitions exclusively.

3.2.2. ENVISAT ASAR WS SM Data

The ENVISAT Advanced Synthetic Aperture Radar (ASAR) data are used in this research. The satellite was launched in 2002 by the European Space Agency (ESA), and the mission ended in April, 2012. ASAR used a C-band (5.3 GHz) SAR, which acquired images in multiple modes, polarizations and at various incidence angles [45]. In particular, the ScanSAR modes included the Wide Swath (WS) and Global Monitoring (GM) modes, covering swaths of a 405-km width under varying incidence angles. The WS mode allowed a spatial resolution of 150 m, a radiometric accuracy of less than 0.6 dB and a maximum duty cycle (the measure of the fraction of the time the radar is transmitting) of 30%, whereas the GM mode provided data with a spatial resolution of 1 km, a radiometric accuracy of about 1.2 dB and a duty cycle of 100%. Although a number of soil moisture studies have been carried out by using ASAR GM images [29,46,47], the high level of noise affecting such data led us to focus on ASAR WS products. Given their higher spatial resolution and the possibility to collect theoretically up to 3–5 images a month, ASAR WS have the potential for better characterization and monitoring of soil moisture. As regards the best geometry of acquisition for the soil moisture retrieval, previous works demonstrated that more accurate results can be achieved by considering descending orbits [48,49], whereas others by using the ascending ones [50]. However, since the combination of ascending and descending orbits allows image acquisitions of a region up to 10-times a month, both orbits were considered here. Concerning the choice of polarization, the largest amount of VV acquisitions over the region under investigation (more than 300 archive images available) led us to discard the HH-polarized data (less than 25 archive images available).

Radar backscattering mostly depends on both sensor parameters (incidence angle, polarization) and land characteristics, such as soil roughness, vegetation cover and soil moisture [51]. In particular, in the C-band, the sensitivity of the backscattering to the soil wetness decreases in the presence of vegetation [52]. However, it has been found in various studies that sparse or low vegetation cover has little influence on the backscattered signal and can generally be neglected. For example, in [53], the authors found that a grass cover (average height of 40 cm) had little influence on ERS-1 (VV polarization) backscattering coefficients, attenuating the signal by less than 0.2 dB. Since all of the study sites are cultivated with relatively short grass, the influence of the vegetation can be considered insignificant on the ASAR WS dataset.

The retrieval of soil moisture from multi-temporal SAR images has been effectively carried out in many studies with a change detection approach [54]. In this work, the change detection algorithm developed by TUW originally for ERS scatterometer images has been used [55]. It is based on the assumption of the time-invariance of surface roughness and vegetation cover, which allows differences in the backscatter to be directly related to the soil moisture variations. Such a hypothesis is met in the grassland areas under investigation, which we can assume to have undergone minimal variation during the period of observation. The original method has been successively adapted for ENVISAT ASAR GM data (1-km spatial resolution). A complete explanation of the algorithm can be found in [29], where the effects on backscattering due to varying incidence angle are taken into account by adopting a standard approach, that is the pixel-wise multi-temporal incidence angle normalization. The estimation of the maximum soil moisture retrieval error, which could affect the ASAR GM SM product, has been evaluated in [29] as:

$$\Delta\vartheta_{max} \approx \sqrt{\left(\frac{1.2}{S}\right)^2 + \left(\frac{\beta}{S}\right)^2 + 0.01} \quad (2)$$

where S is the sensitivity of the backscatter coefficient to soil moisture variations, β the angular correction coefficient and $\Delta\sigma^\circ = 1.2$ dB is the noise of the ASAR GM backscatter measurements. Since ASAR WS acquisitions are characterized by a lower noise ($\Delta\sigma^\circ < 0.6$ dB), the maximum error of the soil moisture retrieval is expected to be smaller than for ASAR GM. Given the similarity of the ASAR GM and WS products and the promising results achieved by exploiting the former [31,56], the same technique has been used to handle the finer spatial resolution SAR data (150-m spatial resolution). ASAR WS scenes have been firstly geocoded, calibrated and resampled to a regular grid characterized by a sampling interval of 15 arcsec. Doing so, the spatial resolution of ASAR WS acquisitions has been degraded to that of the GM mode (1 km), with the advantage of achieving a much improved signal-to-noise ratio.

By using the Corine Land Cover 2006 Map, the retrieved soil moisture maps have been then masked in order to exclude from the subsequent analysis pixels representing classes where the soil moisture values are not reliable (*i.e.*, urban, evergreen broadleaf forest, water bodies, barren or sparsely vegetated areas, snow or ice). In the specific areas under investigation, only a few pixels representing water bodies in the region of Pallaskenry have been masked out. As a further indicator of the reliability of each single pixel, an additional masking has been carried out. In principle, the idea is to select only the points where the adaptation of the algorithm to the ASAR data works better. The temporal stability of soil moisture fields gives rise to an associated temporal stability in the backscatter

signal [57]. A strong correlation between local and regional backscatter is usually a good indicator of high sensitivity to soil moisture dynamics at the local scale. At locations with a weak correlation, either the backscatter response to soil moisture dynamics is dominated by noise and speckle or the backscatter characteristics are adversely influenced by factors, such as dense vegetation, complex topography or soil structure/texture characteristics, which inhibit the retrieval of reliable soil moisture estimates. Therefore, for each $1 \text{ km} \times 1 \text{ km}$ ASAR pixel, the correlation between the time series of the local σ^0 and the average of the backscattering over the $25 \text{ km} \times 25 \text{ km}$ area covering the ASAR one has been evaluated. There is no specific correlation value that is optimally applicable for the purpose of masking ASAR-based soil moisture products in general, as this value depends on backscattering mechanisms, environmental conditions and on the requirements of the intended application of the soil moisture data. In this study, a threshold of $R^2 = 0.3$ was set based on the mentioned assumptions and the experience of other researchers working with ASAR data [58]. After the masking process, it has been observed that in all of the sites, the classes cropland/grassland mosaic and pastures occur over 90% of the ECV pixel areas.

In order to compare the soil moisture expressed by the same unit ($\text{m}^3 \cdot \text{m}^{-3}$), the ASAR relative SM index values have been transformed by applying Equation (1).

3.3. Regional Scale Analysis of SM Temporal Variability

The ECV SM product was compared, on a regional scale basis, with the finer spatial resolution ASAR WS SM data and with the ground measurements. Bearing in mind the different temporal frequency of each dataset, only the ECV and *in situ* SM data corresponding to the ASAR WS acquisitions dates have been considered. The absolute soil moisture values derived from each sensor and their relative changes over time have been considered. This study mainly focused on the latter, aiming at an understanding of the capability of the ECV SM product in capturing soil moisture temporal variations.

The geographical position of the *in situ* stations has been plotted on a georeferenced grid, and the surrounding regions that exactly coincide with the ECV pixels have been selected. Because of the masking process, the number of ASAR data in each ECV pixel is lower than the maximum obtainable. Specifically, when the ECV pixels are completely covered by the ASAR acquisitions, the percentage of available ASAR pixels in each site is: 92.1% for Kilworth, 80.5% for Pallaskenry and 94.7% for Solohead. Moreover, because of the variability of the ASAR coverage over the areas of interest, the number of accessible soil moisture values can change significantly in each acquisition taken during the period of observation. In order to make the subsequent analysis statistically consistent, only the ASAR acquisitions that cover the ECV pixels for more than the 50% of the available pixels (after masking) were considered in the study. For each ASAR acquisition, the average of the soil moisture values retrieved within the corresponding ECV cell has been taken into account for the dataset comparison.

At first, ascending (A) and descending (D) ASAR acquisitions (soil moisture and backscattering) were compared separately with the ECV and *in situ* soil moisture values. The number of ASAR data temporally compatible with the other two datasets is reported in Table 2. Subsequently, two more studies have been carried out: in the first experiment, the whole set of ASAR WS data (AD) has been used together, whereas in the second one, ascending and descending SM values referring to the same day have been averaged (\overline{AD}).

Table 2. Number of ascending and descending ASAR WS images temporally compatible with the ECV and *in situ* soil moisture (SM) data.

Site	Temporal Interval	N. ASAR Ascending (~10:00 pm UTC)	N. ASAR Descending (~11:00 am UTC)	N
KW	6/27/2007–09/15/2009	26	88	114
PK	08/17/2007–09/15/2009	16	70	86
SH	05/23/2007–09/15/2009	24	72	96

The *in situ* SM values have been estimated by adopting two different approaches:

- A. Only the soil moisture measurements recorded at the time closest to the ASAR acquisitions have been used.
- B. The soil moisture values have been evaluated on a daily basis, by averaging the measurements recorded from 0:00 to 23:30 of the same day as the ASAR acquisition.

Summarizing, according to the different approaches, seven comparisons have been performed between the ECV SM product and the other two datasets. Specifically:

1. A_A: only ASAR ascending acquisitions, Approach A for the *in situ* values;
2. A_B: only ASAR ascending acquisitions, Approach B for the *in situ* values;
3. D_A: only ASAR descending acquisitions, Approach A for the *in situ* values;
4. D_B: only ASAR descending acquisitions, Approach B for the *in situ* values;
5. AD_A: both ASAR ascending and descending acquisitions, Approach A for the *in situ* values;
6. AD_B: both ASAR ascending and descending acquisitions, Approach B for the *in situ* values;
7. AD: both ASAR ascending and descending acquisitions, averaging the ASAR and the *in situ* soil moisture values on a daily basis.

In addition, a seasonal comparison has been also carried out between the three datasets, aiming at evaluating the performance of ASAR and ECV soil moisture products in capturing the annual cycle of surface SM and its short-term variability. Since it has been demonstrated that the seasonal vegetation C-band signal is much weaker than the soil moisture signal [59], its effects are generally neglected in the retrieval algorithms [49]. However, the vegetation cycle or intense precipitation periods may reduce the quality of satellite soil moisture products. Therefore, long time series have been split into four seasons (*i.e.*, winter (DJF), spring (MAM), summer (JJA) and autumn (SON)) and analysed in comparison with the *in situ* measurements, which are considered as a reference for the product quality assessment.

The Pearson correlation coefficient (R) has been used to characterize the temporal agreement between each dataset (ECV, ASAR, and *in situ*) at each test site. However, due to the limited number of collocated measurements (approximately 100 for the annual analysis and consequently only around 25 for the seasonal analysis), a certain chance exists that the obtained correlation levels are achieved by coincidence rather than a real statistical dependency. The null hypotheses of the non-significant correlation was tested using the Student's *t*-test, which relates the correlation value to the number of measurements used to calculate it, in order to estimate the probability (*p*-value) of the achieved correlation to be a coincidence, *i.e.*, not significant. The used threshold for accepting the null hypotheses was 0.025 (*i.e.*, 2.5% chance for a statistical coincidence). It should be emphasized

that statistical significance does not mean that the correlation is high, but that the correlation estimate is reliable.

3.4. Analysis of SM Spatial Variability

Heterogeneity of terrain, vegetation cover and topography can make the soil moisture extremely variable even over small distances [60], a variability that the coarse spatial resolution of the ECV SM product may mask. Previous studies [61–63] used soil moisture measurements recorded in several *in situ* stations aiming at the understanding of soil moisture variability over relatively small distances. However, contradictory results have been found, highlighting the complexity of the phenomenon and thus the difficulty of accurately predicting the dynamics of the soil moisture in different areas.

In this work, the SM spatial variability has been studied at the resampled ASAR scale (1-km spatial resolution after resampling). The aim is to investigate which geophysical factors (e.g., terrain composition, altitude, land cover) mainly affect the representativity of the ECV SM product.

Firstly, the average of the soil moisture retrieved from the ASAR WS acquisitions over each ECV-sized cell has been compared with the coefficient of variation (CV) expressed as:

$$CV_j = \frac{\sigma_j}{\bar{\vartheta}_j} = \frac{\sqrt{\frac{1}{N-1} \sum_{i=1}^N (\vartheta_{ij} - \bar{\vartheta}_j)^2}}{\bar{\vartheta}_j} \quad j = 1, \dots, M \quad (3)$$

where N is the number of ASAR pixels within the ECV cell, ϑ_{ij} is the soil moisture estimated in the i ASAR pixel and at time j and σ_j and $\bar{\vartheta}_j$ are the SM spatial standard deviation and mean, respectively.

Successively, the ASAR SM time series values recorded in each pixel have been compared with the ECV SM data. Finally, correlation maps have been provided also on a seasonal basis. The objective of this analysis was to investigate the correlation patterns within the ECV cell and their relationship with specific soil characteristics in terms of texture, land cover and altitude.

3.5. Error Characterization

A statistical approach has been used to quantify the error terms, by computing the root mean square difference (RMSD). This signifies the closeness of two independent datasets representing the same phenomenon. Because of the complexity of the soil moisture phenomenon and the different spatial scale representativity, none of the investigated datasets can be assumed to be the truth, although the *in situ* measurements are normally considered as the reference. As a consequence, the RMSD does not provide a measure of accuracy, but represents the relative error of the soil moisture dynamic range [64]. The RMSD varies with the variability within the distribution of error magnitudes and with the square root of the number of data, as well as with the average-error magnitude (MAE) [65]. Since the ASAR, and the *in situ* SM values are expressed in degrees of saturation, while the ECV SM product is given in volumetric soil moisture; and because of their different spatial resolution, the RMSD is generally affected by a bias. Therefore, all of the SM datasets have been expressed in the same unit ($\text{m}^3 \cdot \text{m}^{-3}$), and the bias has been removed, achieving a better and more reliable estimate of the error by evaluating the centred (unbiased) RMSD:

$$ubRMSD = \sqrt{\frac{1}{N} \sum_{n=1}^N \left(\left(\vartheta_s - \bar{\vartheta}_s \right) - \left(\vartheta_{situ} - \bar{\vartheta}_{situ} \right) \right)^2} \quad (4)$$

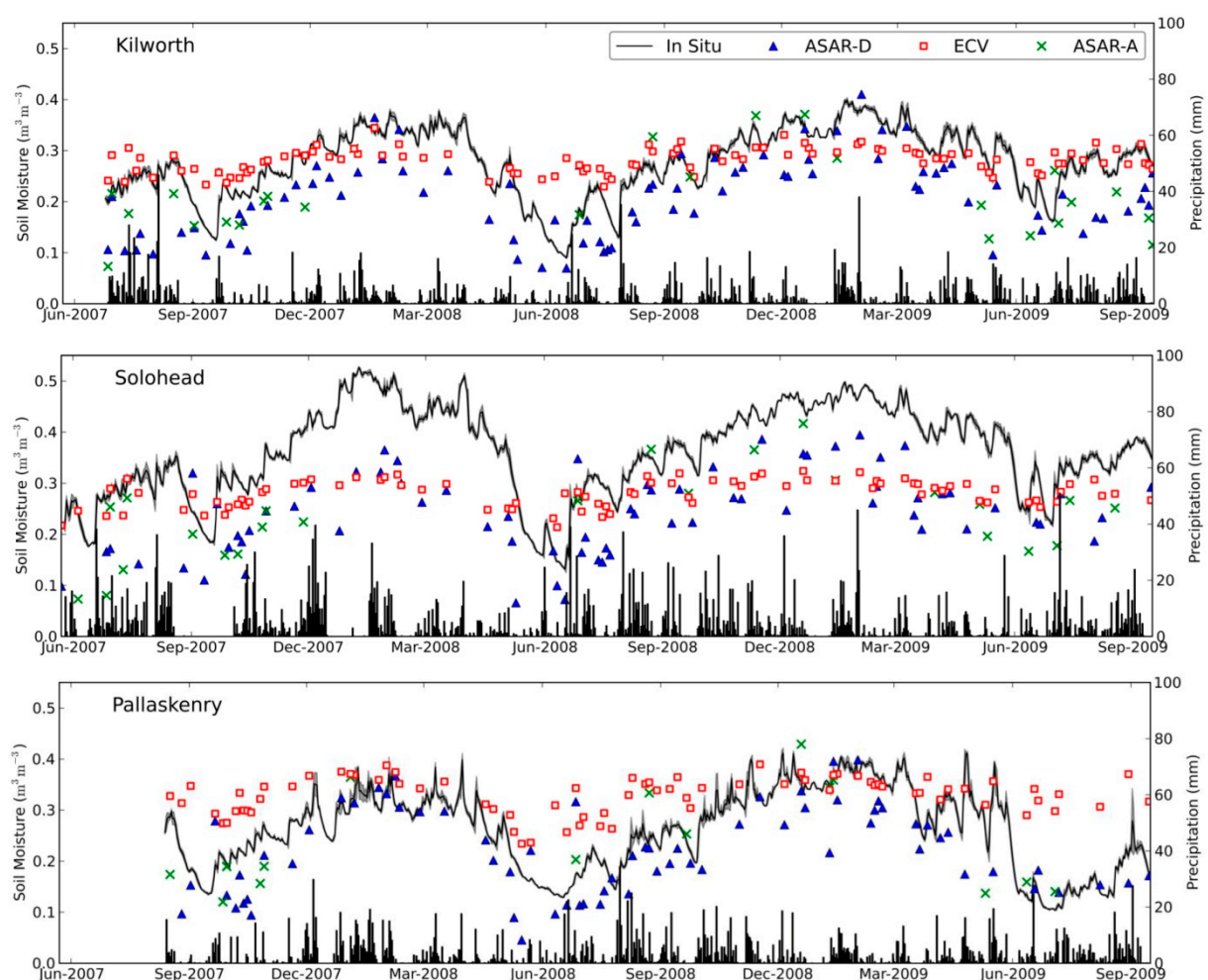
where N is the number of samples, ϑ_s and ϑ_{situ} are the soil moisture values retrieved from the satellite acquisitions and measured *in situ* and $\bar{\vartheta}_s$ and $\bar{\vartheta}_{situ}$ their correspondent means.

4. Results

4.1. Soil Moisture Temporal Variability

In Figure 2, the soil moisture time series retrieved from ascending and descending ASAR WS acquisitions are shown with the ECV values and the *in situ* continuous measurements. The accumulated daily precipitation values are displayed in Figure 2.

Figure 2. Temporal trend of the soil moisture in the sites under investigation. The *in situ* instruments provide measurements every 30-min. The soil moisture retrieved from ascending and descending ASAR acquisitions is compared with that provided by the ECV product. The total daily precipitation is also reported.



Comparable trends are seen for Kilworth and Pallaskenry, where the ECV SM values are about 25% and 31% greater than those retrieved from ASAR, respectively. For Pallaskenry, there is good

agreement with the temporal trend of the *in situ* measurements. While the ECV SM varies in a relatively small interval ($0.1 \text{ m}^3 \cdot \text{m}^{-3}$), a more significant temporal variation occurs in the ASAR time series. The same behaviour is much more evident in the *in situ* measurements, which vary in an interval of approximately $0.25 \text{ m}^3 \cdot \text{m}^{-3}$. Generally, the ECV SM values tend towards the highest *in situ* values. On the contrary, the driest ground conditions are not well represented by the ECV product. Such an observation confirms what has already been shown in other studies [66], where *in situ* measurements and the soil moisture derived from scatterometer acquisitions have been compared. The bias between ASAR and ECV soil moisture is smaller in Solohead, while neither dataset follows the temporal variation exhibited by the *in situ* time series. In this case, the ECV SM values vary in a small interval ($0.21\text{--}0.32 \text{ m}^3 \cdot \text{m}^{-3}$) below the average of the soil moisture ground measurements ($0.35 \text{ m}^3 \cdot \text{m}^{-3}$) recorded during the period under investigation. Contrary to the other two sites, here the best agreement between absolute SM values (ECV, ASAR and *in situ*) occurs in the driest conditions. Such a different behaviour of the soil moisture may be explained as a consequence of the low infiltration capacity of the clay, whose percentage is larger in the soil at Solohead than in the other regions studied. The comparison between the soil moisture retrieved from the remote sensing acquisitions and the *in situ* measurements show very similar results for both approaches, “A” and “B”. The daily standard deviation of the *in situ* values for all sites is 0.006 on average. This means that the variation of soil moisture during the day is negligible. Therefore, the value recorded at the time of the ASAR acquisition (Approach “A”), does not differ significantly from the daily mean (Approach “B”).

Table 3. Correlation values evaluated between each pair of soil moisture dataset, by adopting Approach “B”. No significant difference has been observed between the Approaches “A” and “B”. Because of the very small SM daily variation, the use of both ascending and descending ASAR SM data or their average leads to similar results. Correlation levels are statistically significant ($p < 0.025$).

R	ASAR vs. ECV			ASAR vs. <i>in situ</i>			ECV vs. <i>in situ</i>		
Orbit	KW	PK	SH	KW	PK	SH	KW	PK	SH
A	0.76	0.79	0.88	0.63	0.74	0.71	0.52	0.73	0.60
D	0.77	0.72	0.77	0.78	0.74	0.74	0.74	0.69	0.77
\overline{AD}	0.77	0.73	0.80	0.76	0.75	0.76	0.69	0.69	0.75

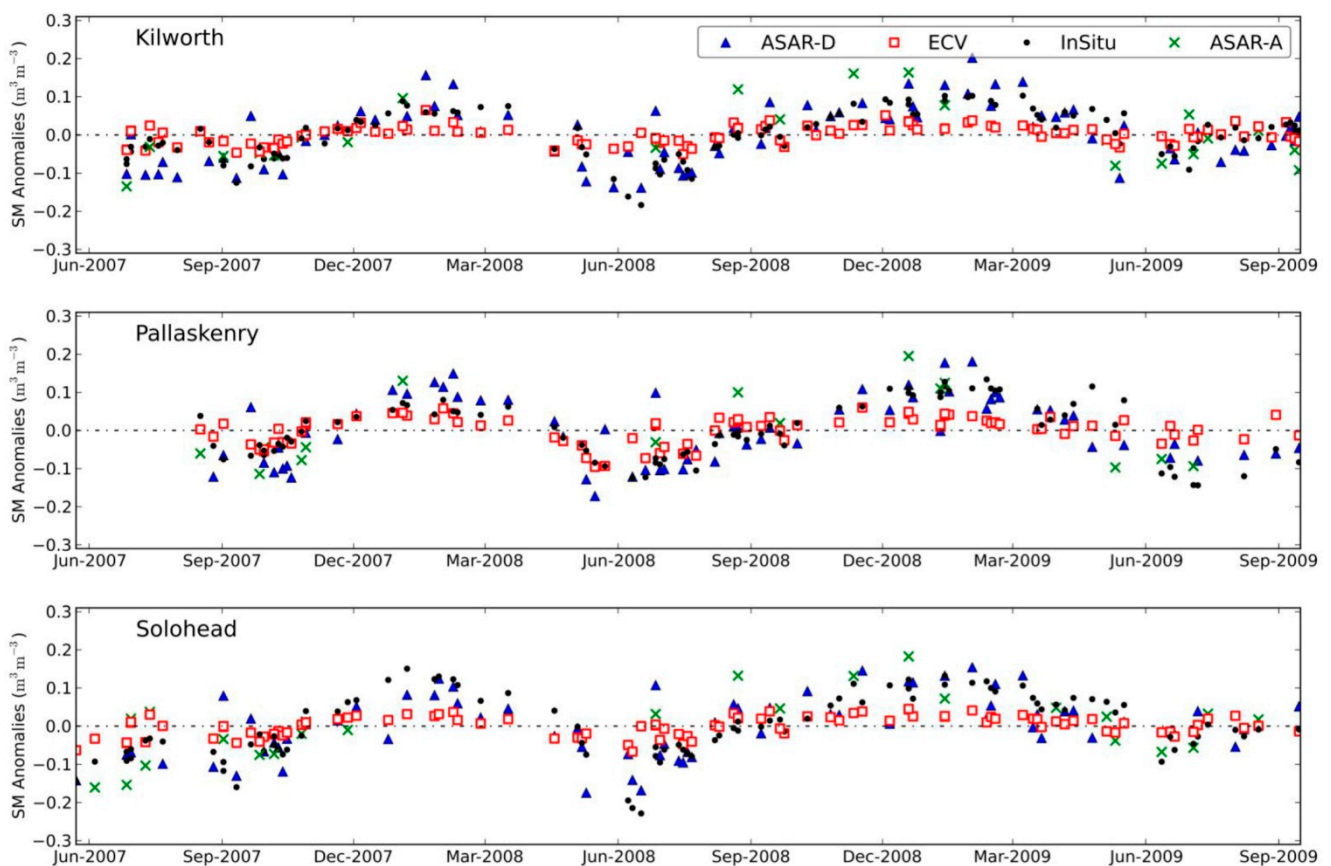
Table 4. Unbiased RMSD evaluated for each soil moisture dataset comparison. ubRMSD unbiased RMSD.

ubRMSD	ASAR vs. ECV			ASAR vs. <i>in situ</i>			ECV vs. <i>in situ</i>		
Orbit	KW	PK	SH	KW	PK	SH	KW	PK	SH
A	0.06	0.07	0.06	0.06	0.06	0.05	0.04	0.06	0.05
D	0.06	0.06	0.06	0.05	0.06	0.06	0.05	0.05	0.06
\overline{AD}	0.06	0.06	0.05	0.05	0.06	0.05	0.05	0.06	0.06

Moreover, in terms of correlation and error (Tables 3 and 4), the results are essentially independent of the choice of the ASAR dataset (ascending, descending or both of the acquisitions). It should be noted that the number of ascending acquisitions is much smaller than the number of descending ones.

In general, a lower degree of correlation occurs when only the ascending ASAR SM time series are used, with the exception of Pallaskenry. However, statistically significant correlation levels ($p < 0.025$) and low unbiased RMSD (ubRMSD) between the independent and multi-scale soil moisture time series have been noted in all of the sites. Specifically, the ASAR *vs.* ECV SM analysis demonstrate high correlations (KW: $R = 0.76$ – 0.77 , PK: $R = 0.72$ – 0.79 , SH: $R = 0.77$ – 0.88), associated with a quite low ubRMSD generally equal to 0.06. High agreement has been found between the ASAR and *in situ* SM time series, with a degree of correlation larger than 0.71 and ubRMSD ranging between 0.05 and 0.06, with the only anomaly being the ascending ASAR SM data at Kilworth. Probably due to the spatial scale difference, the worst outcome in terms of correlation has been estimated between ECV and *in situ* SM data. Nevertheless, R is still quite high (generally higher than 0.6), and the ubRMSD varies in the interval 0.04–0.06.

Figure 3. Temporal evolution of ASAR, ECV and *in situ* SM anomalies estimated in the 2007–2009 time interval.



Soil moisture anomalies have been estimated for all three datasets, during the time interval 2007–2009 (Figure 3). The seasonal cycles are well represented by the temporal evolution of the ASAR and *in situ* SM anomalies, which exhibit similar positive higher values in winter and negative lower values in summer. On the other hand, the poor variability of the ECV SM data leads to a very small range of variation of the associated anomalies. Despite the short period of observation (about two years), it is still possible to detect atypical moisture conditions through the analysis of ASAR and *in situ* SM anomalies. Specifically, in 2007, a drier autumn occurred at all sites compared to 2008.

Based on the *in situ* measurements, in Kilworth and Solohead, the summer of 2008 was drier than the previous and following year. While ASAR SM data are able to capture such a condition, the ECV SM product does not detect any significant anomaly. In Pallaskenry, similar dry soil conditions are revealed by the computation of *in situ*, ASAR and ECV SM anomalies. Successively, in 2009, a wetter summer occurred in Kilworth and Solohead, where the SM anomalies estimated by using all three datasets are very close to zero, meaning that the soil moisture was close to the average evaluated over the whole period of observation (2007–2009). In Pallaskenry, the *in situ* and ASAR anomalies are similar to those recorded in the summer of the previous year; however, the ECV SM anomalies hover about the null value, thereby suggesting wetter conditions compared to the previous summer.

4.1.1. Seasonal-Based Analysis

To investigate the ability of ASAR and ECV SM products to capture the annual cycle of surface SM and its short-term variability, a seasonal comparison has been carried out. For each dataset comparison, the correlation R , the unbiased RMSD and the normalized standard deviation (SDV) have been estimated and plotted in the Taylor diagrams in Figure 4. The normalized standard deviation is defined as the ratio between the standard deviations of soil moisture product and *in situ* measurements:

$$SDV = \frac{\sigma_{SM}}{\sigma_{in\ situ}} \quad (5)$$

Figure 4. Taylor diagrams illustrating the remote sensing soil moisture product statistics against the *in situ* observations. While the seasonal ECV SM values are less variable than the *in situ* records (left side of the dotted arc), the ASAR SM data exhibit a larger variability (right side of the dotted arc). By comparing both the satellite datasets with the ground measurements, it has been found that the highest correlation occurs in spring and autumn. On the contrary, in the extremely wet conditions in winter or relatively dry summer seasons, a very poor agreement has been found between both ASAR and ECV SM time series and ground measurements.

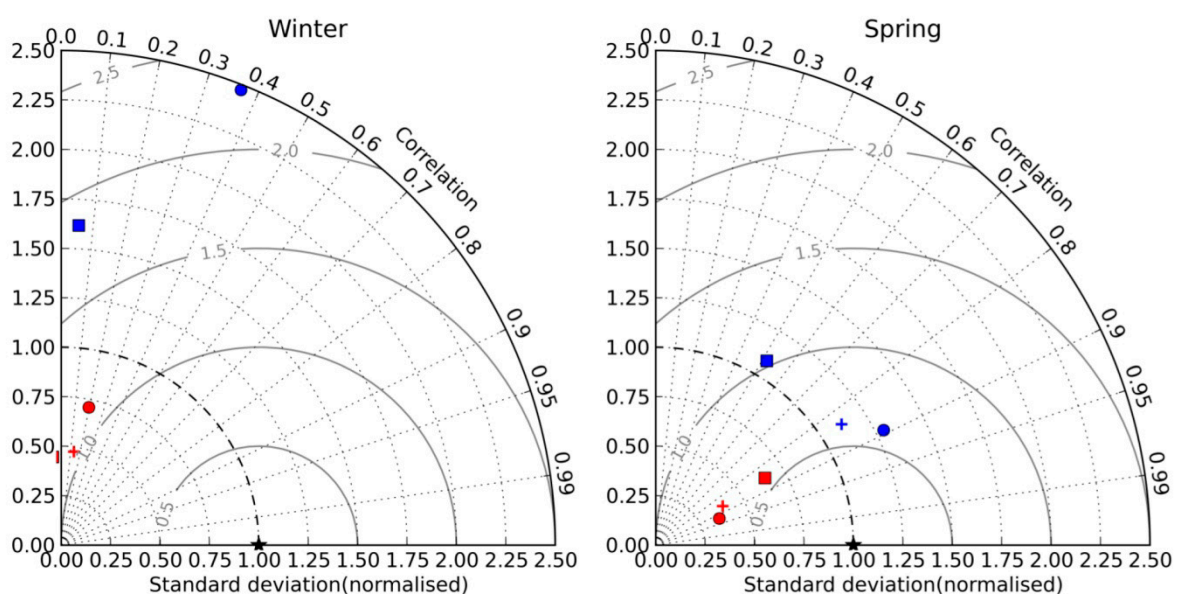
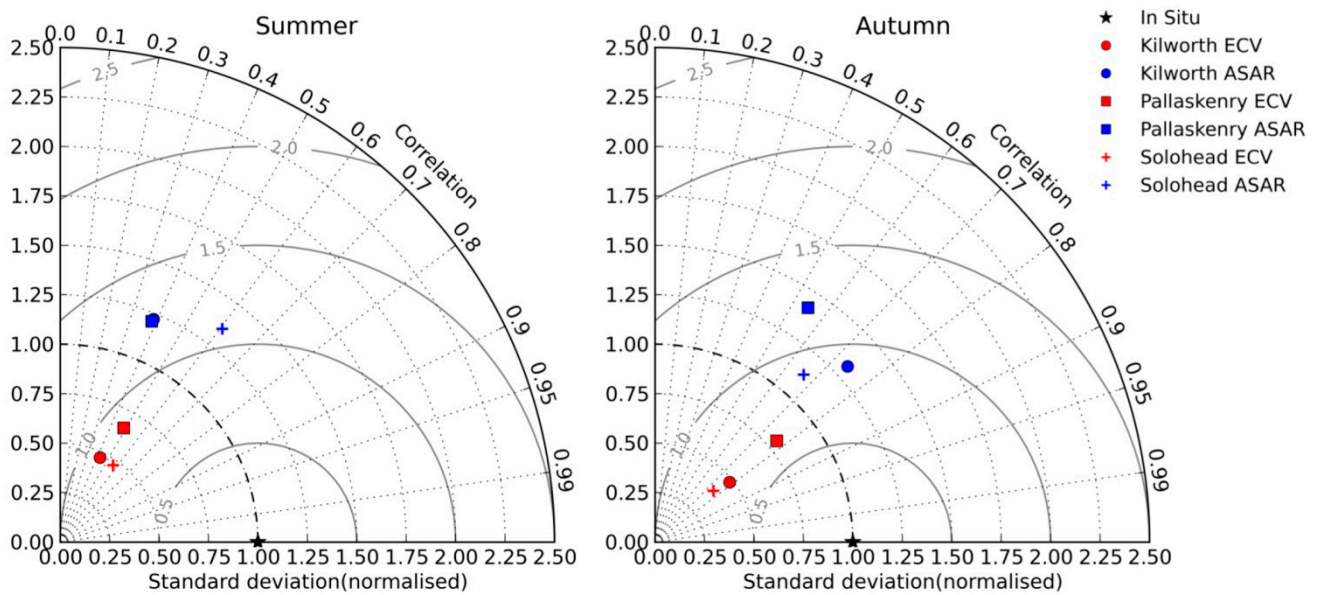


Figure 4. Cont.



In the Taylor representation, the SDV is displayed as a radial distance, with R as an angle in the polar plot. The point defined by $R = 1$ and $SDV = 1$ represents the *in situ* measurements. The distance to this point is evaluated as:

$$E^2 = SDV^2 + 1 - 2SDV \cdot R \quad (6)$$

which is also expressed as a function of the RMSD and bias:

$$E^2 = \frac{(RMSD^2 - Bias^2)}{\sigma_{in situ}^2} \quad (7)$$

The best agreement between the different remote sensing SM products and the *in situ* measurements occurs at a short distance to the $R = 1$, $SDV = 1$ point.

4.1.2. ASAR vs. *In Situ* SM

The seasonal Taylor plots in Figure 4 with the ASAR SM symbols, always located on the right side of the arc with a radius equal to one, highlight the larger variability of the ASAR SM mean with respect to the local *in situ* measurements. For both Solohead and Kilworth, the highest correlation value occurs in spring (MAM), with the poorest agreement in winter (DJF) and summer (JJA) (Figure 4). For Pallaskenry, similarly high correlations are seen in autumn (SON) and spring, while the lowest correlation occurs in winter (Table 5). However, in this case, the p -value is 0.4, making the analysis statistically unreliable. The seasonal-based analysis shows the poorest agreement generally occurring in Pallaskenry (Table 5).

Concerning the error estimation, no significant seasonal variations are seen across the sites. In fact, the ubRMSD ranges between 0.03 and 0.06. The only exception occurs in Solohead, where higher values (0.07–0.08) are observed in all the seasons, except in winter.

Table 5. Seasonal correlation (R) and unbiased RMSD (ubRMSD) between ASAR and *in situ* SM datasets (winter: DJF; spring: MAM; summer: JJA; autumn: SON). The number of available data (n) and the *p*-values (*p*) are also reported.

Seasons	Kilworth				Pallaskenry				Solohead			
	n	R	<i>p</i>	ubRMSD	n	R	<i>p</i>	ubRMSD	n	R	<i>p</i>	ubRMSD
DJF	23	0.37	0.04	0.05	23	0.06	0.	0.05	18	−0.05	0.4	0.04
MAM	18	0.89	<0.001	0.03	15	0.52	<0.025	0.06	17	0.84	<0.001	0.08
JJA	38	0.39	<0.025	0.05	25	0.38	0.03	0.06	34	0.60	<0.001	0.08
SON	35	0.74	<0.001	0.04	24	0.55	<0.025	0.05	27	0.66	<0.001	0.07

4.1.3. ECV vs. *In Situ* SM

As the ECV SM symbols are always on the left side of the arc with a radius equal to one (Figure 4), this indicates that the ECV values are less variable than the *in situ* measurements. The seasonal comparison between ECV and *in situ* SM data (Table 6) confirmed what has been generally found in the previous analysis: the best agreement is achieved in spring and the worst in winter at all sites. However, the results of the analysis carried out for the winter datasets are characterized by high *p*-values, indicating less statistical reliability. At the same time, low correlation values, with $p < 0.025$, have been observed also in summer. This suggests that the retrieval algorithm may be sub-optimal in very wet or in the driest conditions.

Table 6. Seasonal correlation (R) and unbiased RMSD (ubRMSD) between ECV and *in situ* SM datasets (winter: DJF; spring: MAM; summer: JJA; autumn: SON). The number of available data (n) and the *p*-values (*p*) are also reported.

Seasons	Kilworth				Pallaskenry				Solohead			
	n	R	<i>p</i>	ubRMSD	n	R	<i>p</i>	ubRMSD	n	R	<i>p</i>	ubRMSD
DJF	23	0.20	0.18	0.04	23	−0.09	0.34	0.04	18	0.13	0.3	0.04
MAM	18	0.92	<0.001	0.06	15	0.85	<0.001	0.05	17	0.86	<0.001	0.05
JJA	38	0.43	<0.025	0.04	25	0.48	<0.025	0.05	34	0.57	<0.001	0.05
SON	35	0.78	<0.001	0.05	24	0.77	<0.001	0.04	27	0.75	<0.001	0.05

The highest ubRMSD occurs in spring, varying between 0.05 and 0.06. However, no significant variations are noted in the other seasons and in every site. (Table 6)

4.1.4. ASAR vs. ECV SM

Regarding the ASAR-ECV SM comparison (Table 7), the highest correlation values are in spring in the Kilworth and Pallaskenry sites and in winter in Solohead. The lowest correlation values are in winter for Kilworth and Pallaskenry and in summer for Solohead. Different to the previous comparison, the correlation between ASAR and ECV SM datasets is less sensitive to the season. The highest variability in the seasonal R values can be observed in Kilworth ($R = 0.62$ – 0.84), whereas an even smaller range of variation has been evaluated for Pallaskenry ($R = 0.53$ – 0.66) and Solohead ($R = 0.70$ – 0.78). As the *p*-values are lower than 0.025, all of the results can be regarded as statistically reliable.

Table 7. Seasonal correlation (R) and unbiased RMSD (ubRMSD) between ASAR and ECV SM datasets (winter: DJF; spring: MAM; summer: JJA; autumn: SON). The number of available data (n) and the *p*-values (*p*) are also reported.

Seasons	Kilworth				Pallaskenry				Solohead			
	n	R	<i>p</i>	ubRMSD	n	R	<i>p</i>	ubRMSD	n	R	<i>p</i>	ubRMSD
DJF	23	0.62	<0.001	0.04	23	0.53	<0.025	0.04	18	0.78	<0.001	0.02
MAM	18	0.84	<0.001	0.06	15	0.66	<0.025	0.05	17	0.72	<0.001	0.04
JJA	38	0.66	<0.001	0.04	24	0.60	<0.001	0.05	34	0.70	<0.001	0.04
SON	35	0.71	<0.001	0.05	24	0.59	0.001	0.05	27	0.77	<0.001	0.04

Regarding the ubRMSD, the lowest errors occur in winter in all of the sites, although in Kilworth, the same value has been found in summer. In other seasons, there is minimal or no variability at all sites.

4.2. Soil Moisture Spatial Variability

As a measure of the spatial variability of the soil moisture, for each ASAR acquisition, the coefficient of variation (CV) has been plotted as a function of the mean of the SM (Figure 5). Furthermore, seasonal values are highlighted with different colours. In wetter conditions, the soil moisture is less variable for all sites. As a consequence, in winter, the soil moisture is more homogeneously distributed, whereas in summer, high spatial variability occurs. The trend of the CV approximates a decreasing power function, with a high coefficient of determination ($R^2 > 0.87$). Our results agree with the findings reported in [60]. However, it can be observed that, above values of $0.2 \text{ m}^3 \cdot \text{m}^{-3}$, the CV tends to vary linearly with the mean of the soil moisture, as reported also in [67,68].

Figure 5. Coefficient of variation of the ASAR SM mean, evaluated for each ASAR acquisition over the ECV-sized pixels, during the observation period 2007–2009. Seasonal-based values are highlighted by different colours. (Winter: DJF; Spring: MAM; Summer: JJA; Autumn: SON).

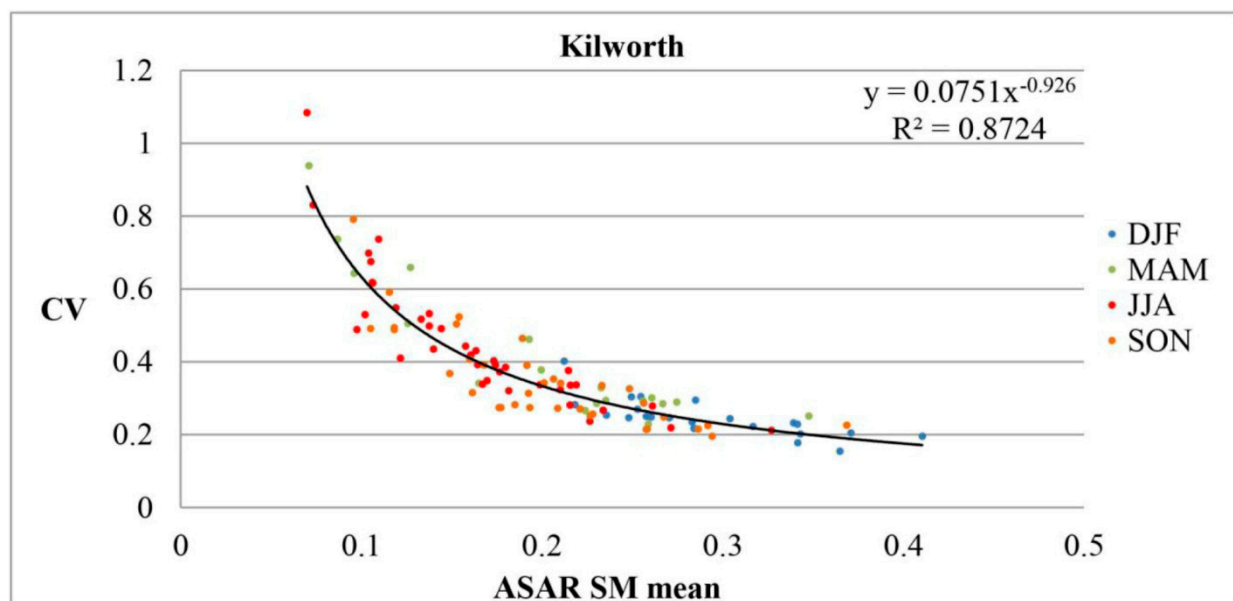


Figure 5. Cont.

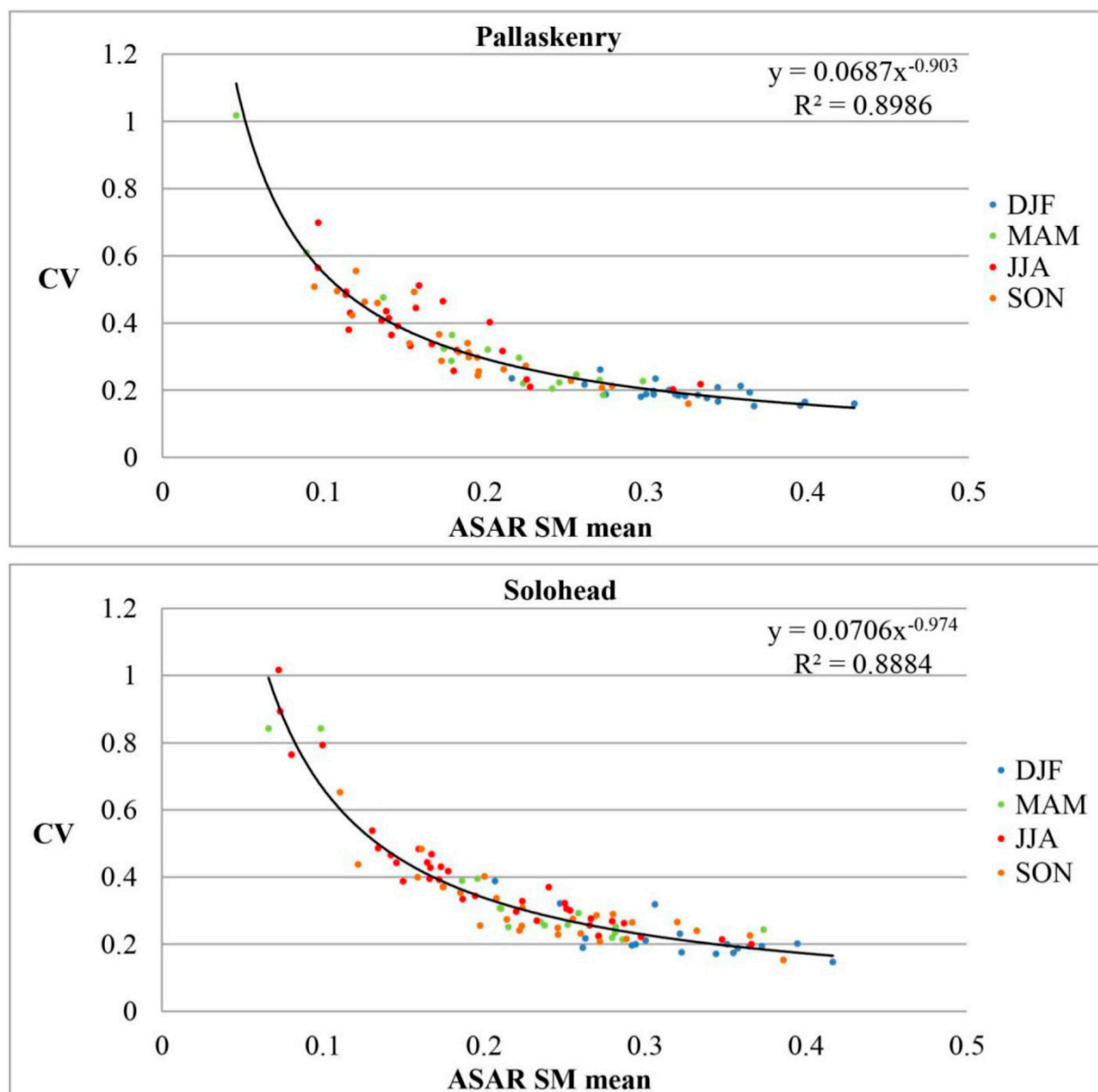


Figure 6. (Left) ASAR SM vs. ECV SM correlation maps evaluated for each ASAR pixel. **(Middle)** DEM maps. **(Right)** Soil maps.

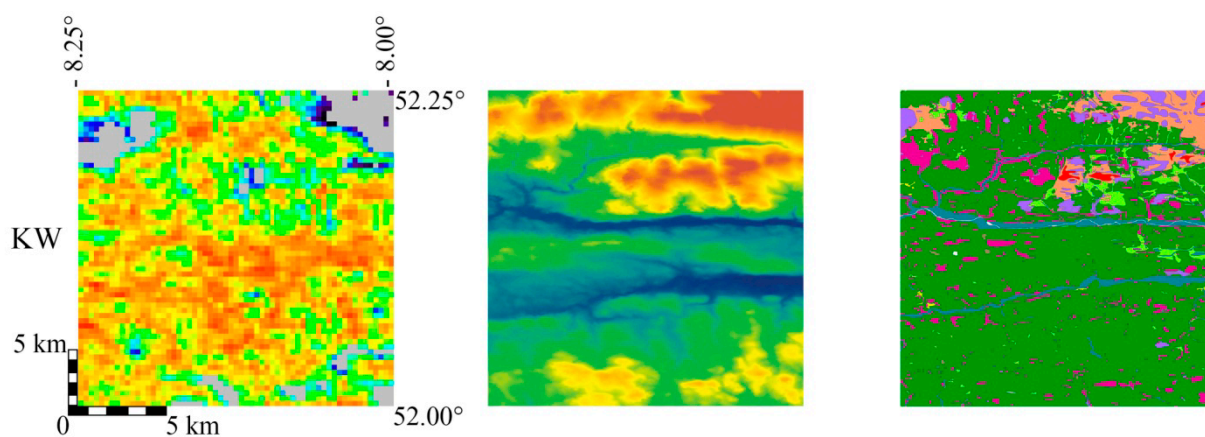
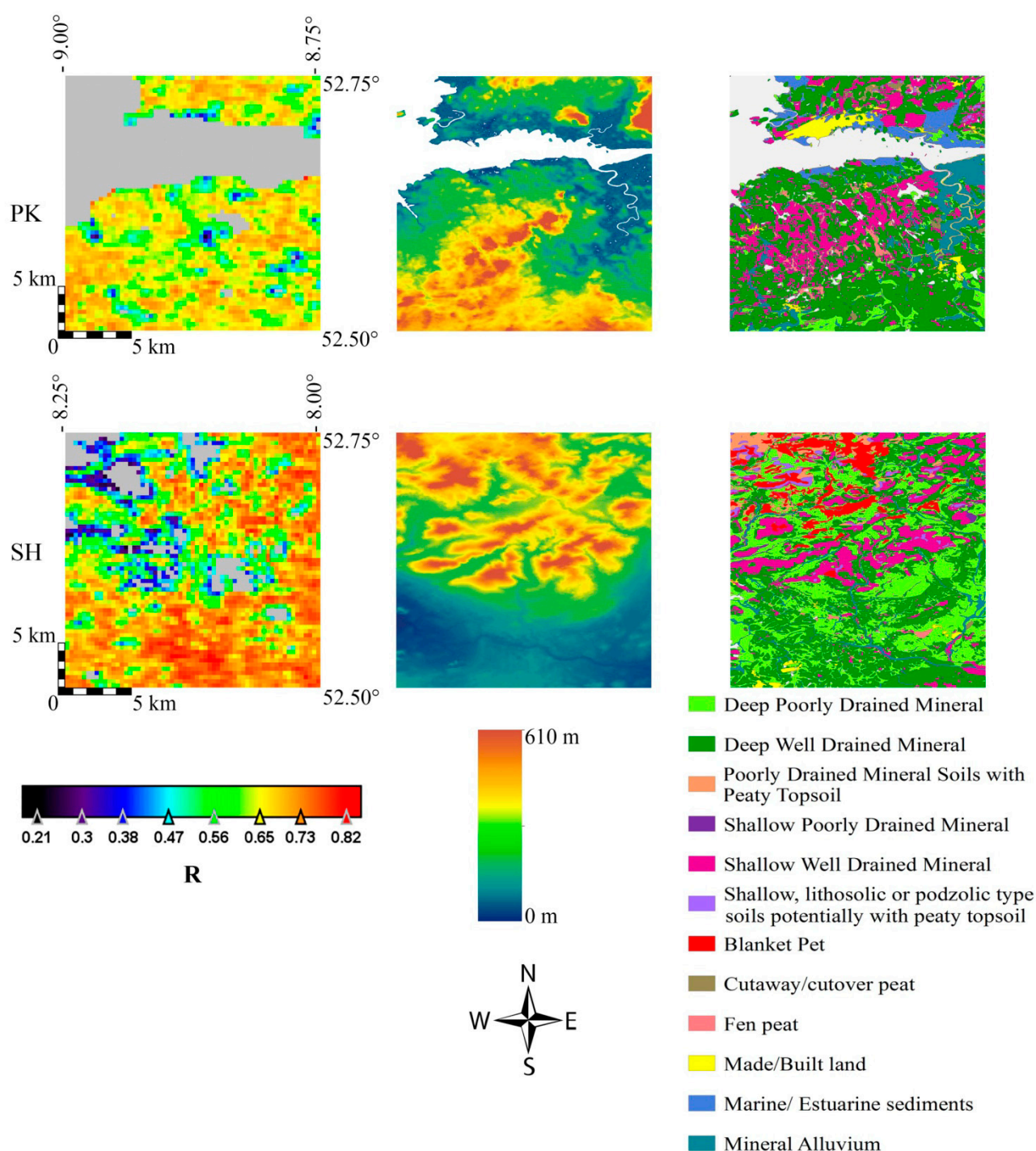
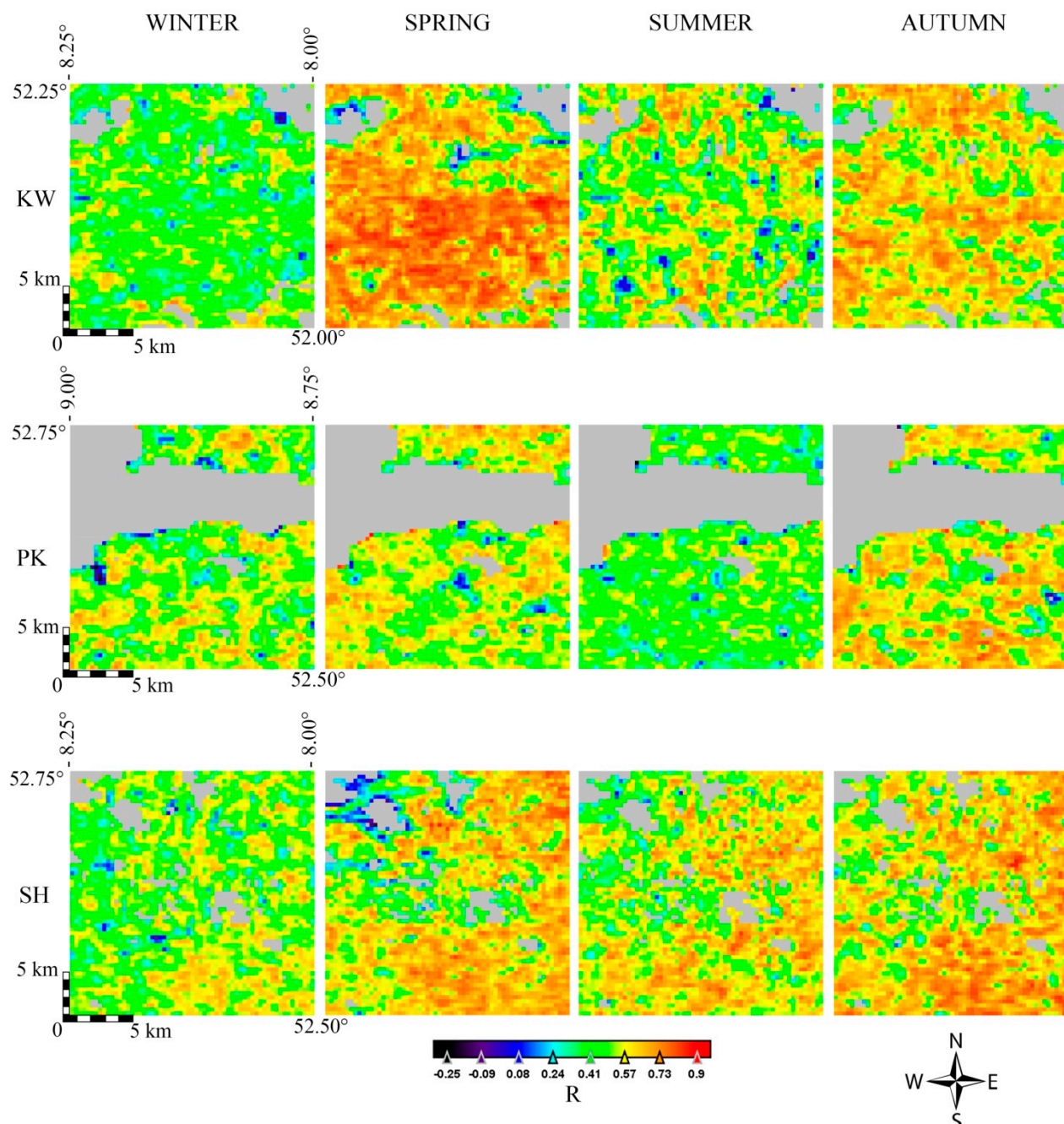


Figure 6. Cont.



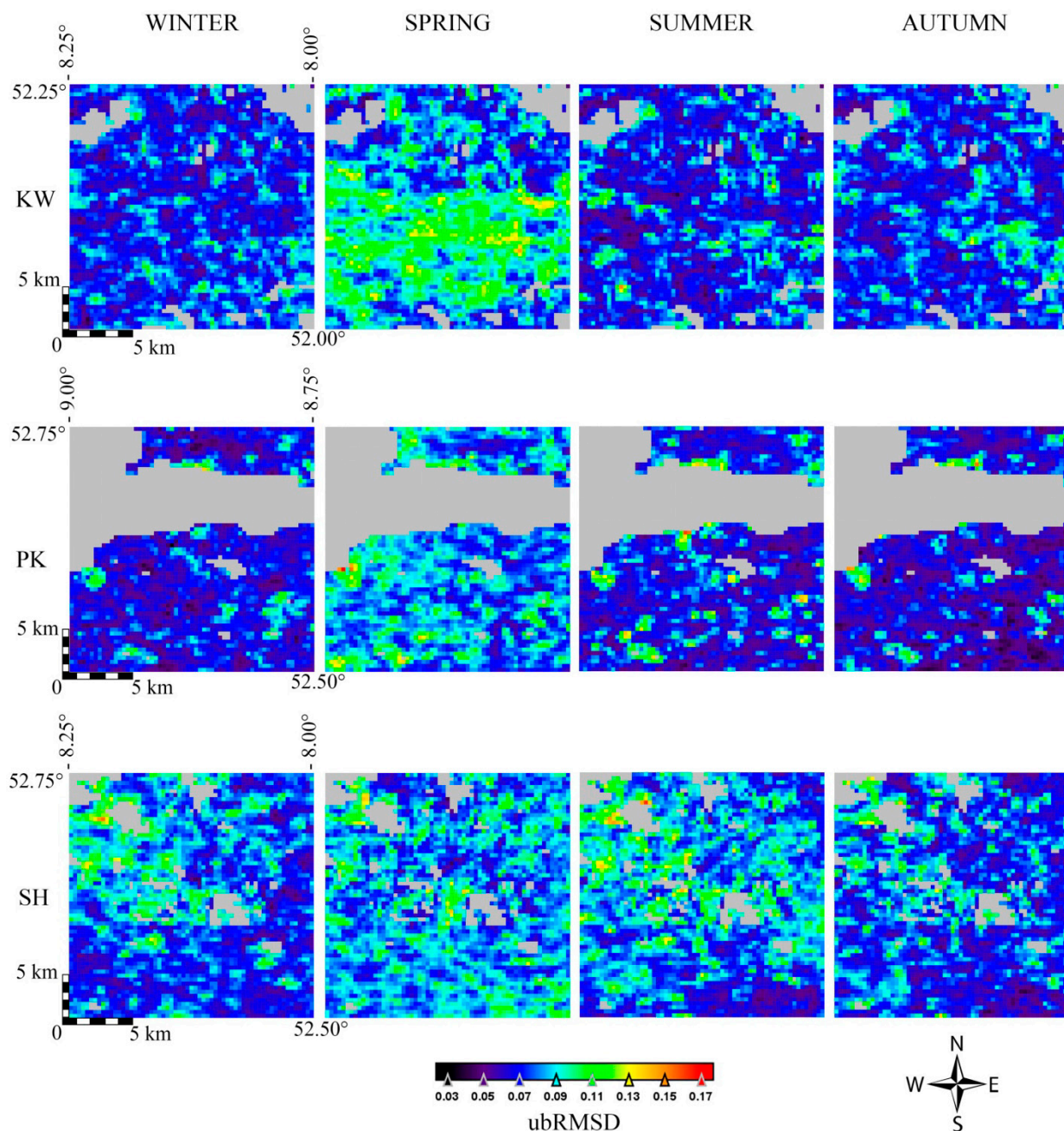
Each ASAR SM local time series value has been successively correlated with the ECV SM data (Figure 6, left). In general, there are quite high R values across all three sites. The highest correlations are seen in Kilworth and Solohead, although the top left corner of the Solohead ECV-sized pixel has quite a low correlation. The comparison with the DEM and soil maps (Figure 6, middle and right) highlights the influence of the combination of altitude and soil type on the soil moisture variations and, hence, on the representation of the ECV SM product. The soil moisture behaviour over those areas characterized by lower altitudes is better described by the ECV SM dataset (high correlation values). These regions mainly correspond to zones where the soil type is classified as deep, well-drained, mineral and mineral alluvium.

Figure 7. ASAR SM vs. ECV SM correlation maps evaluated for each ASAR pixel and in each season. (**Top**) Kilworth; (**Middle**) Pallaskenry; (**Bottom**) Solohead.



The soil moisture spatial analysis has been carried out also on a seasonal basis. The correlation and ubRMSD values between ASAR SM time series collected in each $1 \text{ km} \times 1 \text{ km}$ pixel and the ECV SM dataset are shown in Figures 7 and 8, respectively. A seasonal cycle of the correlation can be noted: in the spring and autumn seasons, the highest values are reached; in summer, R exhibits lower values, as well as in winter. Conversely, the seasonal ubRMSD shows a different behaviour, with higher values generally occurring in spring. In the other seasons, no significant variation of the ubRMSD can be observed. Such seasonal variation is slightly different in Solohead, where high and more homogeneously distributed ubRMSD values occur in spring and summer, while generally lower values can be observed in winter and autumn in most of the ASAR pixels.

Figure 8. ASAR SM vs. ECV SM ubRMSD maps evaluated for each ASAR pixel and in each season. **(Top)** Kilworth; **(Middle)** Pallaskenry; **(Bottom)** Solohead.



5. Discussion

As a contribution to the validation of the innovative CCI-ECV SM global product, in this work, a triple way comparison has been carried out for the first time for such a long time series of SM observations collected using different sensors and provided at very different spatial resolution. The results achieved by comparing the three soil moisture datasets temporally and spatially demonstrate the capability of the ECV SM product in representing the soil moisture variability, despite its coarser spatial resolution. The comparison of the ECV SM product with the ASAR WS and *in situ* time series has shown the capability of the former to capture the temporal dynamics of soil moisture in all three test sites. The ECV SM dataset is highly correlated with the finer spatial resolution ASAR SM product

and with the *in situ* measurements. Such outcomes are consistent with the results obtained in previous works. For instance, in [20], the authors related ECV SM data and *in situ* measurements taken at several stations worldwide, representing different climate zones and characterized by different topography, but mostly located over grassland or agricultural fields. They found correlation values ranging between 0.5 and 0.7 and ubRMSD equal to 0.05–0.06, which are in agreement with what has been observed in the present study. In [50,69], correlation values calculated between ASCAT SM and *in situ* values taken in humid regions are similar to those calculated here between ECV SM data (in our work, generated by using only ASCAT acquisitions) and *in situ* measurements. Our results also agree with those in [29], where the soil moisture retrieved from ASAR GM acquisitions has been compared with that derived from ERS scatterometer data and *in situ* measurements.

It should be noted that the lowest correlations have been estimated when only ascending (10:00 pm) ASAR acquisitions have been used, while improved outcomes have been achieved by using images taken during the descending orbit (10:00 am). Similar findings have been reported in other papers, where ASCAT- and/or AMSR-E-derived SM data have been analysed [50,69,70]. For instance, in [69], the authors suggest the use of morning measurements, because of the day-time decoupling that could happen in the afternoon due to the non-hydraulic equilibrium of the soil.

By analysing the satellite soil moisture values, it has been observed that those provided by the ECV product are generally higher than those derived from the ASAR WS acquisitions. The reason may be related to the change detection algorithm used to retrieve the soil moisture from the SAR images. It is based on parameters that have been shown to be sensitive to the season [71] and on assumptions derived from the use of SM time series by ERS scatterometer data [29]. Nevertheless, we have demonstrated here that ASAR WS data are suitable to assess the quality of the coarser spatial resolution ECV SM product. By considering the *in situ* soil moisture dataset as the reference, high correlation values and small bias have been estimated between ASAR SM time series and the ground measurements.

The comparison between the three soil moisture datasets shows that the temporal variability of the ECV SM product is lower than for the ASAR SM and *in situ* data, as noted previously in [20]. The different spatial scale of the SM products used may lead to such a behaviour [29]. In particular, the processing performed to generate the ECV SM data, which are scaled to the GLDAS-Noah model, yields a coarser spatial resolution with a limited ability to track the seasonal soil moisture variability. At the same time, the absolute SM values variations are reduced throughout the whole period of observation. In this regard, the analysis of the soil moisture anomalies during approximately two years of observations has shown a seasonal periodicity, which is particularly evident in the ASAR and *in situ* SM datasets, rather than in the ECV SM time series. Higher anomalies have been observed in winter and lower ones in summer. Specifically, the ASAR and *in situ* soil moisture products highlight the occurrence of anomalous dry and wet summers in Kilworth and Solohead in 2008 and 2009, respectively, whereas the ECV SM product fails to capture dry extremes. However, it should be noted that anomalies were calculated with respect to just over two years of data. For this specific study, only ASCAT data were available to generate the ECV product. As has been already observed in [66], in the driest conditions, the SM retrieved from only scatterometer acquisitions tends to overestimate the real moisture content as measured by *in situ* instruments. The very small anomalies of the ECV SM time series reflect this trend. The poor capability of the ECV SM product in capturing the driest and wettest soil conditions has been highlighted also in a seasonal-based analysis. It demonstrated that very low

correlations occur in winter and summer between both satellite products and the ground soil moisture dataset, whereas the best agreement has been achieved in spring and autumn. Winter is the wettest season in Southern Ireland. When heavy and/or continuous precipitation occur over a poorly-drained soil (e.g., Solohead), a water layer could persist on the surface, reducing the satellite microwave backscattering sensitivity to soil moisture and, hence, providing incorrect estimates of the moisture content. Summer is the driest season, during which the vegetation reaches the maximum growth, which may affect the quality of the retrieved soil moisture from SAR and scatterometer images. Furthermore, when the change detection algorithm is applied on humid sites by using a limited number of images, the sensitivity of the microwave signal to soil moisture (the difference between the driest and wettest signal) is likely to be underestimated [72,73]. This can explain the low correlation between ASAR and *in situ* SM data in winter and summer.

Concerning the soil moisture spatial variability study, while published works are typically carried out through geostatistical analysis by using hydrological models and *in situ* networks over rather wide regions, in this work, the spatial distribution of soil moisture has been investigated at the ASAR scale (1 km) within a smaller area (*i.e.*, the ECV size pixel). In doing so, actual observations covering the whole area have been used, without the need for adopting any particular model. This approach facilitated an understanding of the extent to which geophysical factors, such as soil texture, terrain composition and altitude, affect the retrieved ECV SM product values in temperate grasslands. The spatial variability of the soil moisture retrieved from the ASAR WS acquisitions over each ECV-sized pixel shows that wet conditions lead to more homogeneous moisture content across the area, while in the driest conditions, a significant variation in soil moisture values can occur. Therefore, in winter, the distribution of soil moisture is rather homogenous over the ASAR pixels within the ECV cells. On the contrary, a larger spatial variability has been observed in summer. This may be due to the frequent, but localized, heavy shower activity, which characterizes summer in Ireland. Under such conditions, the combination of soil texture type, terrain composition, evapotranspiration and percolation processes and water table levels is likely to play a significant role in the soil moisture dynamics. In the present work, we have been able to demonstrate how CV varies against the ASAR mean values for a full range of moisture conditions, and this highlights the seasonal variability. The trend of the CV approximates a decreasing power function, with a high coefficient of determination ($R^2 > 0.87$). Winter mean soil moisture values are associated with the lowest and most stable CV. On the contrary, summer mean soil moisture values correspond to high and more variable CV values. Our results agree with the findings reported in [60], where the authors described the SM spatial variability for different types of soil texture. In particular, the authors note that the CV of sandy soil increases as the soil dries and reaches its maximum near the residual moisture content. However, it can be observed that above values of $0.2 \text{ m}^3 \cdot \text{m}^{-3}$, the CV tends to vary linearly with the mean of the soil moisture. This result is in line with [67,68], where, focusing on soil moisture behaviour in humid grassland, the authors found the relationship between CV and the mean soil moisture to be a decreasing linear function. In [74], the authors observed that in wet catchments in New Zealand, the variability decreases with increasing moisture content, whereas it increases in the drier Australian catchments. They assert that these differences in behaviour are due to differences in the seasonal patterns of controlling processes associated with seasonal changes in spatial mean soil moisture, particularly the lateral flow processes,

which are related to climatic differences. Our results underline the ability of ASAR WS retrieved SM data to track this full spectrum of varying moisture content and seasonal behaviour.

The comparison between soil moisture time series in each ASAR pixel and in the corresponding ECV cells has shown high R values, larger than 0.56 on average at all sites. This highlights the effectiveness of the ECV SM product in representing the soil moisture conditions, despite the coarse spatial resolution. A further confirmation of the quality of the ECV SM product is given by the fact that the seasonal R values are quite homogeneously distributed over all sites, following the same periodic trend observed by carrying out the regional based analysis (higher R values in spring, lower in winter). By studying the spatial distribution of the correlation between ASAR and ECV SM, it has been observed that the soil moisture behaviour over those areas characterized by lower altitudes is better described by the ECV SM dataset (high R values). These regions mainly correspond to zones of deep, well-drained, mineral and mineral alluvium soils. On the contrary, areas located at a higher altitude, mainly characterized by poorly-drained soil, show lower correlation between the two satellite SM products. In addition, high slope variability contributes to the further loss of correlation. Indeed, infiltration, drainage and runoff depend on the slope angle. Steeper slopes generally cause lower infiltration rates, rapid subsurface drainage and higher surface runoff [75]. Since ASAR WS acquisitions are not necessarily taken at the same time as the ECV satellite data, different moisture content is likely to be detected by each product over these regions. The lower spatial correlation observed in Pallaskenry, which is characterized by a more complex topography and higher altitude, supports such a hypothesis.

6. Conclusions

The representativeness of the global soil moisture product provided in the framework of the ESA CCI programme has been investigated through the use of finer spatial resolution ASAR WS data and ground measurements across three grassland sites in Southern Ireland. In doing so, this is the first time a comparison has been carried out between three sets of independent observations from different sensors at very different spatial resolutions for such a long time series. Neither dense *in situ* station networks, nor hydrological models have been used, but the results obtained from the adopted approach are consistent with those reported in other papers using different sensors and classical methods. Therefore, the applicability and reliability of the approach presented has the potential to be an efficient and cost-effective validation method for low resolution SM products.

Both the regional-based and the ASAR pixel-based analysis demonstrated that there is a high correlation between the different soil moisture products. By carrying out the regional-based analysis, it has been observed that for all of the sites, the correlation between ASAR and ECV SM datasets varies between 0.72 and 0.8 (ubRMSD = 0.05–0.07). Slightly lower values have been found when the satellites products are compared to the *in situ* SM time series (ASAR vs. *in situ*: R = 0.63–0.78, ubRMSD = 0.05–0.06; ECV vs. *in situ*: R = 0.52–0.77, ubRMSD = 0.04–0.06). The ASAR pixel-based analysis highlighted the presence of high correlation patterns (R = 0.7–0.8) in each area under investigation. Moreover, despite the topography and soil type playing a role in the quality of the ECV SM product, we found limited differences in the correlation values across all of the studied areas, where R is generally higher than 0.6. Such results highlight the capability of the ECV SM product to describe the actual moisture dynamics in quite a large area (0.25×0.25 degrees), with a rather low

ubRMSD. However, the evaluation of the SM anomalies revealed a limitation in the ECV product, which is its poor capability in capturing the wettest and driest conditions. It is important to be aware of this when the product is used in climate change studies, as not capturing the extremes may lead to underestimation of evapotranspiration in summer and underestimation of surface runoff in winter. On the basis of these observations, we can conclude that further tuning of the algorithms is necessary in order to provide a higher quality product and, hence, gain more confident use of the ECV SM global dataset for climate change studies.

The recently improved and updated global ECV SM product will be released in mid-2014, after testing and validation. Quality assessment will be carried out by using *in situ* measurements, ASAR WS, as well as Sentinel-1 data, as soon as they become available. In order to verify any improvement of the new ECV SM dataset, the analysis will be performed by considering the same test sites. However, with the aim of comprehensively understanding the ECV global product, it will be necessary to test it with further studies in other regions worldwide. Therefore, different climate zones, various soil types and topography will be taken into account. Since the Irish sites are mainly located in grassland areas, in this work, it has not been possible to investigate the sensitivity of the soil moisture to the land cover. Yet, it is a major factor to take into account when soil moisture dynamics are studied. Therefore, future work will address such issues by selecting more heterogeneous regions, where it will be possible to establish how specific land covers affect the soil moisture provided by the ECV product.

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Author Contributions

Chiara Pratola carried out all of the analyses under the supervision of Edward Dwyer and Brian Barrett. Alexander Gruber performed the ASAR SM retrieval. The discussion of the findings has benefitted from the expertise of Gerard Kiely. All of the authors contributed to the preparation of the paper through their review, editing and comments. Brian Barrett provided some of the plots shown in the manuscript.

Conflicts of Interest

The authors declare no conflict of interest.

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