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3 Feedbacks on climate in the Earth system: introduction
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13 In the past century the Earth has undergone a very fast and unusual change in the radiative forcing
14 of its climate, resulting from human actions. This change in forcing has resulted mainly from rising
15 concentrations of greenhouse gases, particularly carbon dioxide (CO₂), but also from changes in the
16 nature of the land surface and changes in the concentrations of aerosol particles in the atmosphere
17 and of ozone-destroying chemicals in the stratosphere. If the changes in forcing continue on their
18 current trajectory, then very substantial changes in climate are predicted by the end of the century
19 [1].
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22 It is reasonably straightforward to calculate the equilibrium change in temperature at global scale
23 that would result from a particular change in forcing if the Earth behaved as a simple black body and
24 there were no additional effects [2]. For example it can be calculated that a doubling of CO₂
25 concentration results in a long-wave forcing of about 3.7 W m⁻², which by itself would cause an
26 equilibrium warming of about 1.2°C.
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29 However in practice the initial perturbation causes a range of other feedback effects, which may
30 weaken or strengthen the global temperature response. These effects are substantial, and it is the
31 net effect of such feedbacks that determines the overall sensitivity of the planet to forcing. Negative
32 feedbacks (such as that due to the change in atmospheric lapse rate, i.e. changes in the vertical
33 temperature gradient of the atmosphere) cause the climate to be less sensitive to changes of
34 forcing, while positive feedbacks like the ice-albedo feedback cause it to be more sensitive. The most
35 recent IPCC synthesis report concludes that the combined effects of all feedbacks is likely to be
36 significantly positive [1].
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40 Importantly, the nature and magnitude of these feedbacks is the principal cause of uncertainty in
41 the response of Earth's climate (over multidecadal and longer periods) to a particular emissions
42 scenario or greenhouse gas concentration pathway. Moreover, although feedbacks are often
43 discussed at a global scale, and in terms of single global numerical feedback parameters, feedbacks
44 also play a key role in inducing regionally variable responses to climate forcing, both in terms of
45 temperature and other key variables such as rainfall and the occurrence of extreme events. For
46 these reasons, improving understanding of feedback processes is an urgent priority.
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49 Because the emphasis in climate change research has often been on the next few decades, up to a
50 century, most discussions of climate feedbacks have focussed on the fast feedbacks, such as those
51 caused by changes in water vapour, clouds and sea ice extent, which are included in the traditional
52 definition of climate sensitivity [3]. However, the eventual response of the planet, and the one that
53 is observed in records of past climate changes, includes some feedbacks with a much longer
54 timescale. These include changes in ice sheet extent, and carbon cycle feedbacks. A final group of
55 effects, often discussed under the umbrella of abrupt events, but representing highly non-linear
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3 feedbacks, may also be considered: these would include hypothesised releases of high latitude
4 methane, or rapid changes in ocean circulation resulting from climate change.
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7 In designing this discussion meeting, we included work on a deliberately broad range of feedbacks,
8 taking into account feedbacks with timescales ranging from instantaneous up to thousands of years.
9 We split the topic broadly into three categories: feedbacks involving water in the atmosphere, as
10 water vapour, clouds and aerosol; feedbacks involving ice (on both land and the sea); and feedbacks
11 involving greenhouse gases and the carbon cycle. Presenters were asked to consider the processes,
12 modelling studies and the palaeo-records of past climate change. Process studies contribute the
13 detailed knowledge needed to improve models. Modelling studies indicate how the different
14 processes combine to produce observed large-scale behaviour. Palaeo-records contains numerous
15 examples of the response of the climate system to natural forcings such as volcanic eruptions,
16 changes in the distribution of solar energy due to changes in Earth's orbit and axial tilt, and natural
17 emissions of greenhouse gases. These can be investigated [4, 5] to challenge and support our
18 understanding of feedbacks in the present and the future.
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21 The papers in this volume represent contributions from the three broad topics, delving into the past,
22 present and future, and into process studies, modelling work and palaeo-climate analysis. The
23 contributions outline the current state of knowledge and highlight future directions of research.
24 Some of the papers take a more general look at the overall impact of feedbacks [3, 6], and the
25 variability of that impact over time [7] and space [8]. In the two days following the discussion
26 meeting, a satellite meeting considered some of the issues raised in more detail, basing discussion
27 around a set of questions, paraphrased here as:
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31 "Which feedbacks (and what aspect of them) provide the most uncertainty in response, (a) for
32 global response by 2100, (b) for regional variations of temperature, precipitation, extreme events
33 and (c) for long term outcomes and commitments beyond 2100?" The following paragraphs draw on
34 the conclusions of both meetings.
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38 The response from water vapour itself and from the lapse rate feedback are considered to be
39 reasonably well-understood, but that from clouds [9, 10] remains the major known cause of
40 uncertainty in both short term (decadal) and long-term (multi-centennial) climate prediction.
41 Improvements in the understanding of processes affecting a number of different cloud types have
42 been made in recent years. An approach involving multi-model comparisons, cloud process
43 modelling and ground-based observations offers a path to further reduce uncertainty, but it appears
44 to remain challenging to translate new knowledge into more skilful practicable representations of
45 the effects of clouds (parameterisations) in GCMs.
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49 For aerosols a first concern has been to determine whether the forcings assessed in the historical
50 period due to aerosols are accurate, as this affects the correct assessment of the forcing-response
51 relationship. However beyond that, there remains much work to be done on the role of aerosols in
52 a range of feedbacks (including those involving clouds). The meeting heard about new experimental
53 facilities, that are starting to have a major impact on our understanding of the complex chemistry
54 determining the rate of formation and growth of cloud condensation nuclei (particles on which cloud
55 condensation can occur).
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3 The presence of absence of sea ice causes a range of feedbacks, starting with the change in albedo
4 (reflectivity), but encompassing a wide range of physical and chemical effects. Along with changing
5 snow cover, this is a crucial component in producing the biggest regional contrast in the response to
6 climate change: Arctic amplification. The Antarctic differs from the Arctic in a number of respects:
7 not least, one is a continent surrounded by ocean, and the other an ocean surrounded by continent.
8 Amplification and sea ice retreat have not yet emerged around most of the sectors of Antarctica.
9 Overall, there is still some way to go in understanding how ice dynamics and thermodynamics, and
10 other factors involving atmospheric and ocean dynamics and atmospheric chemistry, interact to
11 determine the way in which feedbacks are manifested in the polar regions. The possible impact of
12 Arctic sea ice changes on mid-latitude weather patterns is a topical example of the processes that
13 are not yet properly explored and are still poorly understood.
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18 The ice sheets on land react on much longer timescales, and are not included in the traditional short-
19 term definition of climate sensitivity, but of course were the dominant feature of glacial-interglacial
20 changes and some other past climate changes. Their dynamic behaviour is one of the most
21 important uncertainties for the commitment beyond 2100, as there are good reasons to believe that
22 there is a threshold (whose value remains uncertain) beyond which the loss of much of the
23 Greenland ice sheet becomes inevitable. Additionally the processes that can destabilise a marine ice
24 sheet such as most of West Antarctica are now much better understood, but not to a level where an
25 accurate prognosis of their contribution to the future rates of sea level change can yet be given. The
26 role of the cryosphere was extensively discussed at the meeting, and is considered in some
27 contributions here, but is not represented by a dedicated paper in the volume.
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31 The feedbacks discussed so far describe the amplifiers and attenuators of climate change in
32 response to a given change in a forcing (such as CO₂). However there are further feedbacks within
33 the carbon cycle itself. More than half of anthropogenic carbon emissions are currently absorbed by
34 the ocean and the land biosphere and it is critical to climate projections to accurately understand
35 how this proportion may evolve in response to changes in CO₂ concentration and climate. The
36 palaeo-climate record suggests that large changes in concentrations of greenhouse gases are
37 possible, and were a major amplifying feedback in glacial cycles. Remote sensing and measurement
38 networks are transforming our ability to understand carbon fluxes but much more needs to be done.
39 Promising approaches include the study of “emergent constraints” – the evidence of recent
40 variability in the ocean and land sinks of CO₂, responding to climate variations such as the El Niños
41 [11], and detailed data-constrained models of major unknowns, such as the response of permafrost
42 soils to climate warming [12].
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47 Finally, discussion of the economic benefits of narrowing the uncertainties in the strength of
48 feedbacks [13] reminds us that a whole new set of feedbacks between the climate and the human
49 response to it will play a major role in defining the trajectory of emissions and climate change in the
50 next few decades.
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53 The papers in this volume review what is known about many of the major feedback processes that
54 have affected past, and will affect future, climate change. They also identify the issues that we know
55 still need to be addressed, and demonstrate how they may be approached. Even so, it remains
56 possible that there are additional feedback processes that have yet to be discovered and may prove
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3 to be significant; to avoid surprises, we must continue to use improved process understanding along
4 with the palaeo-record to identify missing feedbacks that should be included in models.
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6 References

- 7
8 [1] IPCC. 2013 *Climate Change 2013: The physical science basis. Contribution of Working Group I to*
9 *the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge and New
10 York, Cambridge University Press; 1535 p.
11 [2] Knutti, R. & Hegerl, G.C. 2008 The equilibrium sensitivity of the Earth's temperature to radiation
12 changes. *Nature Geoscience* **1**, 735-743. (doi:10.1038/ngeo337).
13 [3] Knutti, R. & Rugenstein, M.A.A. 2015 Feedbacks, climate sensitivity, and the limits of linear
14 models. *Philosophical Transactions of the Royal Society A*.
15 [4] Braconnot, P. & Kageyama, M. 2015 Shortwave forcing and feedbacks in Last Glacial Maximum
16 and Mid-Holocene PMIP3 simulations. *Philosophical Transactions of the Royal Society A*.
17 [5] Kennedy, A.T., Farnsworth, A., Lunt, D.J., Lear, C.H. & Markwick, P.J. 2015 Atmospheric and
18 oceanic impacts of Antarctic glaciation across the Eocene-Oligocene transition. *Philosophical*
19 *Transactions of the Royal Society A*.
20 [6] Gillett, N.P. 2015 Weighting climate model projections using observational constraints.
21 *Philosophical Transactions of the Royal Society A*.
22 [7] Gregory, J.M., Andrews, T. & Good, P. 2015 The inconstancy of the transient climate response
23 parameter under increasing CO₂. *Philosophical Transactions of the Royal Society A*.
24 [8] Sutton, R., Suckling, E. & Hawkins, E. 2015 What does global mean temperature tell us about local
25 climate? *Philosophical Transactions of the Royal Society A*.
26 [9] Bretherton, C.S. 2015 Insights into low-latitude cloud feedbacks from high-resolution models.
27 *Philosophical Transactions of the Royal Society A*.
28 [10] Webb, M.J., Lock, A.P., Bretherton, C.S., Bony, S., Cole, J.N.S., Idelkadi, A., Kang, S.M., Koshiro,
29 T., Kawai, H., Ogura, T., et al. 2015 The impact of parametrized convection on cloud feedback.
30 *Philosophical Transactions of the Royal Society A*.
31 [11] Friedlingstein, P. 2015 Carbon cycle feedbacks and future climate change. *Philosophical*
32 *Transactions of the Royal Society A*.
33 [12] Koven, C.D., Schuur, E.A.G., Schädel, C., Bohn, T.J., Burke, E.J., Chen, G., Chen, X., Ciais, P.,
34 Grosse, G., Harden, J.W., et al. 2015 A simplified, data-constrained approach to estimate the
35 permafrost carbon-climate feedback. *Philosophical Transactions of the Royal Society A*.
36 [13] Hope, C. 2015 The \$10 trillion value of better information about the transient climate response.
37 *Philosophical Transactions of the Royal Society A*.
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