



Investigation of the impact of anthropogenic heat flux within an urban land surface model and PILPS-urban

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1 Investigation of the impact of anthropogenic heat flux within an Urban Land Surface model and PILPS-urban

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9

10 Abstract

11 Results from the first international urban model comparison experiment (PILPS-Urban) suggested that models which
12 neglected the anthropogenic heat flux within the surface energy balance performed at least as well as models that
13 include the source term, but this could not be explained. The analyses undertaken show that the results from PILPS-
14 Urban were masked by the signal from including vegetation, which was identified in PILPS-Urban as being important.
15 Including the anthropogenic heat flux does give improved performance, but the benefit is small for the site studied
16 given the relatively small magnitude of this flux relative to other terms in the surface energy balance. However, there is
17 no further benefit from including temporal variations in the flux at this site. The importance is expected to increase at
18 sites with a larger anthropogenic heat flux and greater temporal variations.

19

20 Keywords: Anthropogenic heat, JULES, Surface energy balance, Urban environment

21

22 1. Introduction

23

24 Numerical weather prediction (NWP) models have included surface processes for many years (e.g., Manabe, 1969), but
25 it is only within the last decade that representations of urban areas have been included (e.g., Best, 2005, Lemonsu et al.,
26 2009, Hamdi et al., 2014), even though urban energy balance models themselves have been developed over a much
27 longer period (e.g., models evaluated by Ross and Oke, 1988, Masson, 2000, Martilli et al., 2002). The increasing
28 resolution for NWP models has now reached the stage where urban areas can make up a large proportion of a grid-box,
29 or in some instances actually be resolved. This has led to additional interest from this community to include a
30 representation of urban areas within their modelling systems. Also, the move towards more integrated impacts for
31 climate change has seen a move away from the post processing of urban areas from climate change signals, to including
32 cities within the climate change simulations themselves (e.g., Oleson et al. 2008, McCarthy et al. 2010).

33

34 Unlike natural surfaces, the energy exchange within an urban environment includes additional source terms from the
35 human activities (e.g., Sailor, 2011). These include the energy that is released from the heating of buildings and the
36 emission of heat from vehicular transport. There is also a contribution from the metabolism of humans themselves,
37 although this is typically small compared to the other sources (Sailor, 2011). Calculations of the total magnitude of this
38 anthropogenic heat vary considerably between cities (e.g., Christen and Vogt, 2004, Offerle et al., 2005, Quah and
39 Roth, 2011, Kotthaus and Grimmond, 2014), and between different areas of any particular city (e.g., Ichinose et al.,
40 1999, Pigeon et al., 2007, Iamarino et al., 2012). In some locations the magnitude of the anthropogenic heat flux can be

41 a substantial source term, similar to the daily mean solar forcing (e.g., Ichinose et al., 1999, Hamilton et al. 2009,
42 Iamarino et al. 2012).
43
44 *A priori* it would be expected that such an additional source term would need to be accounted for in any urban model.
45 However, results from the first international urban model comparison project (PILPS-Urban) consistently suggested that
46 models which do not include an anthropogenic heat flux performed at least as well as models that did include this flux
47 (Grimmond et al., 2011, Best and Grimmond, 2013). These studies were not able to suggest the reasons for this and
48 indicated that additional investigation is required.

49
50 Further to the results of these studies, Figure 1 shows the seasonal errors derived from the results of PILPS-Urban for
51 the median model in each group, when the models are categorised by the complexity with which they represent the
52 anthropogenic heat flux. These have been calculated using the methodology presented in Best and Grimmond (2013).
53 The results show that the group of models that do not include the anthropogenic heat flux have the smallest root mean
54 square errors (RMSE) for all four of the surface fluxes at all times of the year, compared to any of the other model
55 groups that include this additional energy flux in various forms. In addition, the models without the anthropogenic heat
56 flux also have the smallest bias for the sensible and latent heat fluxes, although they are the only group of models that
57 have a negative bias in the sensible heat flux during the winter months.

58
59 Here we aim to understand the counter intuitive results from PILPS-Urban. To do this we use one of the models from
60 the comparison which did not include any anthropogenic heat flux. We have rerun the simulations that were done for
61 PILPS-Urban whilst introducing the additional source of energy to this model and analysed the subsequent impact on
62 the results without this term, as submitted to the comparison.

65 **2. Methodology**

67 2.1 Observations

68
69 The observational site chosen for PILPS-Urban was Preston, a northern suburb of Melbourne, Australia. Details of the
70 site are given in Coutts et al. (2007), and have also been summarised in the various analyses of PILPS-Urban results
71 (Grimmond et al., 2011, Best and Grimmond, 2013, 2014). The site is described as urban climate zone (UCZ) 5, with
72 moderately developed low density housing (Coutts et al., 2007). Two methods were used to determine the surface cover
73 fractions over a 500 m radius with the average giving 45% building area, 5% concrete, 13% roads, 15% grass, 23%
74 other vegetation and 1% other (Coutts et al., 2007), thus a total impervious surface of 62% and pervious of 38%.

75
76 Observations of the radiative, sensible and latent heat fluxes were all undertaken whilst the net storage heat flux was
77 determined as the residual of the surface energy balance. The net advective heat flux is difficult to determine and was
78 assumed to be negligible, and the anthropogenic heat flux was derived using an inventory approach. Details of the
79 instrumentation, sampling and averaging periods used by Coutts et al. (2007) are not discussed further here.

80

81 To determine the anthropogenic heat flux Coutts et al. (2007), following Sailor and Lu (2004), consider the heat
82 released from three different sources: vehicles, buildings (which is subsequently split further into electricity and natural
83 gas), and human metabolism. The vehicular contribution were based on surveys undertaken during November 2002 –
84 October 2003 (Coutts et al., 2007) and the 2001 census population to estimate a distance travelled per person per day.
85 To disaggregate to hourly values, the mean hourly traffic profile for U.S. cities from Hallenbeck et al. (1997) was
86 invoked.

87
88 For building electricity half-hourly demand data were used, but only the fraction used for direct heating was accounted
89 for. This is 43.1% of the total electricity usage, whilst the remaining 56.9% was used for refrigeration, lighting and
90 appliances, for which heat is only a small by-product (and is thus neglected). For natural gas the diurnal heating profile
91 was estimated using the diurnal variability in consumption which was modelled as a function of the daily range in
92 temperature, using mean maximum and minimum temperatures occurring at 1700 and 0700 local solar time (LST)
93 respectively, and linear interpolation.

94
95 The human metabolic rates, along with the day-time and night-time periods, were taken from Sailor and Lu (2004), i.e.,
96 175 W between 0700 – 2100 LST and 75 W between 2300 – 0500 LST, respectively. During the transition periods
97 (0600 and 2200 LST) a fixed value of 125 W was used, in contrast to Sailor and Lu (2004) who used linear
98 interpolation between the day and night-time values. These metabolic rates were then used along with the population
99 density to determine the contribution to the anthropogenic heat flux.

100
101 The contribution of the anthropogenic heat flux from human metabolism calculated by Coutts et al. (2007) is small
102 compared to the other source terms, as seen in other studies (e.g., Grimmond, 1992; Sailor and Lu, 2004). For the other
103 three components, the magnitude of their contribution is similar. The vehicle term has distinctive double peak during
104 morning and afternoon rush hours which, as documented by Sailor and Lu (2004), but this is to be expected because the
105 diurnal variations were determined from the same U.S. datasets. The natural gas term peaks in the morning at the time
106 of minimum temperature (0700 LST) and has its smallest value in the afternoon at the time of maximum temperature
107 (1700 LST), whereas the electricity term is fairly constant throughout the day. The resultant diurnal cycle for the
108 anthropogenic heat flux has two peaks, but with the morning peak being greater and the afternoon peak (Coutts et al.,
109 2007).

110

111

112 2.2 Urban land surface model

113

114 The model used in the current study was the Joint UK Land Environment Simulator (JULES, Best et al., 2011). JULES
115 can be run off-line (as used within this study) or coupled to provide the land surface component within the Unified
116 Model (UM, Cullen, 1993), which is used by the Met Office for weather forecasting and climate applications. Four sets
117 of JULES results, in two configurations, were contributed to PILPS-Urban. These represented the urban fraction as a
118 single bulk surface, or the roof surface and street canyons separately; namely the 1-tile (Best 2005) and 2-tile versions
119 (Best et al. 2006). These two configurations were run by two modelling centres and ensured that the physical set up of
120 the models was consistent, but the assumptions about the initial conditions, especially the soil moisture, were different.
121 One set of simulations had more initial soil moisture than the other, which was shown by Best and Grimmond (2014) to

122 have important implications for the model performance. None of the four sets of JULES simulations included an
123 anthropogenic heat flux.

124
125 Despite the differences in physical configuration and initial conditions between the four sets of JULES simulations,
126 results from PILPS-Urban show that all of them perform well compared to other models (Figure 2, adapted from
127 Grimmond et al., 2011), especially for the sensible and latent heat fluxes. Hence this is a good model to use to
128 investigate the impact of including the additional anthropogenic heat flux.

129
130 For the simulations presented here, both the 1-tile and 2-tile versions of JULES were used, but with the initial soil
131 moisture set to the same values as used for the dryer set of results contributed to PILPS-Urban.

132
133 To put the results for the impact of including the anthropogenic heat flux into context with other aspects of the physical
134 system represented in the urban models, an additional simulation using JULES with no representation of vegetation was
135 undertaken. Results from the previous studies of PILPS-Urban concluded that a representation of vegetation was critical
136 in order to obtain good performance from the urban models, especially for the sensible and latent heat fluxes
137 (Grimmond et al., 2010, 2011, Best and Grimmond, 2013, 2014). However, the method by which the vegetation is
138 represented, i.e., though an independent surface (tile scheme) or interacting with the urban surface (integrated) was
139 shown to be less important.

140
141 The JULES model included a tile scheme representation for vegetation in all four of the simulations submitted to
142 PILPS-Urban, and the same representation has been used in all of the simulations with the anthropogenic heat flux.
143 However, an additional simulation was completed with no anthropogenic heat flux and with the vegetation removed.
144 This was done by setting the fraction of the vegetation surface to zero and re-scaling the urban surface fractions to sum
145 to unity.

146
147

148 2.3 Anthropogenic heat flux

149
150 The JULES model was adapted to include the anthropogenic heat flux as an additional term in its surface energy
151 balance:

$$152 \quad C \frac{\partial T}{\partial t} = Q^* - Q_H - Q_E - \Delta Q_S + Q_F$$

153 where C is the areal heat capacity of the surface, T is the surface temperature, Q^* is the net all-wave radiation, Q_H is the
154 turbulent sensible heat flux, Q_E is the latent heat flux, ΔQ_S is the net storage heat flux and Q_F is the anthropogenic heat
155 flux.

156
157 The anthropogenic heat flux can hence be considered as an additional source term to the surface energy balance,
158 equivalent to additional radiative flux forcing. As such it will lead to an increase surface temperature with associated
159 larger values of the sensible and latent heat fluxes, the net storage heat flux and the outgoing longwave radiation. The
160 net impact on an atmospheric model to the inclusion of anthropogenic heat flux would be to increase the heat, moisture
161 and longwave radiative flux boundary conditions from the surface.

162

163 The anthropogenic heat flux was specified at every time-step of the model run, based upon the observed values of the
164 anthropogenic heat flux. By using the observed values, the analysis removes the impact of using a scheme to represent
165 the flux which would inevitably have its own inaccuracies. Hence we can identify the true impact of including the
166 anthropogenic heat flux.

167

168 For the 1-tile version, the anthropogenic heat flux was applied to just the urban surface energy balance. However, for
169 the 2-tile version there is more flexibility regarding the addition of the flux. It can be added to the canyon surface
170 energy balance only, the roof surface energy balance or to both the canyon and the roof surface energy balances. To
171 understand the full impact of the anthropogenic heat flux on the JULES simulations, the 2-tile version was run in all
172 three configurations. The anthropogenic heat flux applied to each surface was scaled to ensure that the total flux
173 integrated over all surfaces is equal to the observed values.

174

175 In addition to investigating the impact of including the anthropogenic heat flux, the impact of the temporal variation in
176 the flux is considered. Four methods were used (Fig. 3a-d), namely:

- 177 a) The average value of the period of the observations: constant (i.e., no temporal variation)
178 b) Monthly mean values, constant diurnal cycle, causing a step change between consecutive months
179 c) Average diurnal cycle over the entire observational period, constant variation between months
180 d) Monthly mean diurnal cycle, with variations between months (estimated Q_F data, Coutts et al., 2007)

181 The latter is the full temporal resolution of Q_F available. Note there are no differences in diurnal cycle between days
182 within the same month given the methodology used to determine the observed values; i.e., no response to
183 meteorological conditions or human behavior (e.g., days of the week).

184

185 The mean anthropogenic heat flux in the observational dataset is 11 W m^{-2} , which is assumed to occur only on the built
186 land cover in JULES (i.e., not applied to vegetation or bare soil surfaces), and so requires a value of 17 W m^{-2} to be
187 used in JULES. This corresponds to a diurnal maximum of 26 W m^{-2} and a minimum of 10 W m^{-2} . Figure 3e shows the
188 average diurnal cycle over the observational period, along with the diurnal cycles for the months with the maximum and
189 minimum values. This figure shows two important aspects for understanding the results presented here. Firstly, the
190 magnitude of the flux is small compared to the average of the other terms in the surface energy balance ($Q^* = 83 \text{ W m}^{-2}$,
191 $Q_H = 40 \text{ W m}^{-2}$, $Q_E = 34 \text{ W m}^{-2}$ and $\Delta Q_S = 20 \text{ W m}^{-2}$), and in particular to the mean incoming shortwave radiation ($K_{\downarrow} =$
192 168 W m^{-2}). Secondly, the minimum values in the diurnal cycle occur during the night-time hours, when the sensible
193 heat flux is small.

194

195

196 2.4 Analysis methods

197

198 The mean bias error (MBE) and root mean square error (RMSE) for Q^* , Q_H and Q_E for each of the simulations are
199 presented in Figure 4, for all data points and separately for the night-time values only (defined by $K_{\downarrow} = 0.0 \text{ W m}^{-2}$). The
200 statistics for the net storage heat flux are not shown as this is taken to be the residual of the energy balance in the
201 observations and as such aggregates the observational errors. To calculate these statistics, any time-steps with missing
202 observational data for any flux are omitted from the analysis. This is to ensure that the results are consistent between the

203 fluxes, and with the methodology that was adopted in PILPS-Urban by Grimmond et al. (2011) and Best and Grimmond
204 (2013, 2014).

205
206

207 **3. Results**

208

209 Figure 4 shows that the JULES simulations that include the anthropogenic heat flux have better performance (smaller
210 MBE and RMSE) for the sensible heat flux, compared to the simulations without the anthropogenic heat flux (as
211 submitted to PILPS-Urban). This holds for the analysis using all of the data and using the night-time data only. The
212 latent heat flux has a very small increase in the positive bias over all of the data and a decrease in the negative bias for
213 the night-time only data, but virtually no change to the RMSE. However, for Q^* there is an increase in the MBE for
214 both all data and the night-time data and a corresponding increase in the RMSE.

215

216 The increase in negative bias for the night-time results for Q^* are a result from the higher surface temperatures leading
217 to more outgoing longwave radiation. However, since Q_H still has a negative bias when Q_F is included, this suggests
218 that the relation between the radiative surface temperature and the thermodynamic temperature within JULES is not
219 optimal and could be improved.

220

221 The results show that for the anthropogenic heat flux with temporal variations there is little impact on all three of the
222 fluxes (Figure 4), with only small differences in either the MBE or the RMSE. Hence for these simulations, including
223 the average value of the anthropogenic heat flux is more important than having time varying values, either diurnally or
224 monthly. However, it should be noted that for this study only the mean monthly variations in Q_F are available and not
225 the true temporal variability that depends on the meteorological conditions and human behavior, such as weekdays
226 versus weekends.

227

228 All of the 2-tile versions perform better than the equivalent 1-tile version for Q^* and Q_H for both MBE and RMSE for
229 all data and night-time data only. The only exception is for the MBE in Q_H with the simulations that do not include
230 vegetation. There is no difference between the 1-tile and 2-tile versions for Q_E for either the MBE or the RMSE, except
231 for the night-time MBE for the 2-tile version with the anthropogenic heat flux only applied to the canyon surface, which
232 has a slightly larger negative bias.

233

234 There are also differences with the 2-tile version of JULES for Q^* and Q_H when considering the surfaces over which
235 the anthropogenic heat flux are implemented. The simulations with the anthropogenic heat flux applied to just the
236 canyon surface have noticeably smaller MBE than the other simulations for both all data and night-time only data, but
237 these improvements are not evident in the RMSE apart from the night-time data for Q_H which has a very small
238 improvement. Consistent with these results, we find that not applying the anthropogenic heat flux to the canyon tile
239 (i.e., applying it only to the roof tile) leads to larger MBE and RMSE, particularly for the sensible heat flux.

240

241 Despite the improvements that can be detected from the model simulations that include the anthropogenic heat flux,
242 these are far smaller than the improvements that are obtained from including a representation of vegetation, especially
243 for the RMSE for all of the data. These results are robust over all of the sets of simulations.

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The maximum values of the anthropogenic heat flux occur in the winter months (Fig. 3d), when the diurnally averaged sensible heat flux is at its lowest values. Hence we might expect to see a larger impact from the anthropogenic heat flux on Q_H at this time of the year. Seasonal variations of 60 day means in the results for the 2-tile version of JULES, with the anthropogenic heat flux applied only to the canyon energy balance, are shown in Figure 5 for the night-time. This figure shows the mean flux, MBE and the RMSE for all of the surface fluxes and is equivalent to the analysis presented in Best and Grimmond (2013) for all of the models in PILPS-Urban. For both the day-time (not shown) and night-time results (Fig. 5), the seasonal cycle of the anthropogenic heat flux generally makes no difference to the improvement of the model in terms of MBE for any of the fluxes. The results are consistent across all of the months, with the only exception being a slight improvement to the MBE of $\sim 1.5 \text{ W m}^{-2}$ for the night-time Q_H during June/July (JJ) compared to the summer months. However, there is a slightly larger improvement in the RMSE for Q_H ($\sim 3 \text{ W m}^{-2}$) in both the day-time (not shown) and night-time (Fig. 5) during the winter months (JJ) compared to not including the anthropogenic heat flux. For nocturnal data this improvement results from the RMSE not increasing by as much as for the results from the model without anthropogenic heat flux.

For the net storage heat flux, there is a slight improvement in both the MBE and RMSE during the day-time from including the anthropogenic heat flux (not shown), but a degradation in both statistics for the night-time (Fig. 5). These changes to both MBE and RMSE are consistent throughout the seasonal cycle.

As well as the anthropogenic heat flux having larger impacts at certain times within the seasonal cycle, we might expect the impact on the surface fluxes to vary during the diurnal cycle, especially when Q_H has its smallest values (i.e., during the night-time). However, an equivalent figure to Fig. 1 of Best and Grimmond (2013), for the average diurnal cycle for each 60 day period of the seasonal cycle, shows that the impact on the diurnal cycle for all of the surface fluxes is small (not shown). At the scale of the range of the diurnal cycle, the differences in the various model simulations that include the anthropogenic heat flux are not discernible from those without an anthropogenic heat flux.

Focusing on just the night-time part of the diurnal cycle (Fig. 6), results show that for this period there are some noticeable differences in the sensible and net storage heat fluxes between the model simulations. All of the JULES runs that include Q_F are almost indistinguishable from each other, but for the sensible heat flux they are clearly closer to the observed values than the JULES run that excludes Q_F (hence reducing the bias, Fig. 6). This reduced negative bias in the sensible heat flux from including Q_F ranges from a minimum of 4.8 W m^{-2} in the summer to a maximum of 7.3 W m^{-2} in the winter.

For the net storage heat flux, there is a reduction in the magnitude of the negative flux during the night-time from the runs with Q_F compared to the run without Q_F . The results for the JULES simulation without an anthropogenic heat flux have only a small bias in the net storage heat flux at night-time over most of the seasonal cycle. So subsequent changes to the night-time values in the net storage heat flux from implementing an anthropogenic heat flux into JULES introduces a more notable positive bias at all times of the year (results are further from the observations compared to the JULES run without Q_F , Fig. 6).

285 **4. Conclusions**

286

287 The JULES model, one of the better performing models in PILPS-Urban (especially for Q_H and Q_E) did not include any
288 representation of the anthropogenic heat flux. Hence it is a good tool to investigate if the inclusion of Q_F could improve
289 the model performance still further.

290

291 The results from the runs presented here to investigate this have shown that there is an improvement in both the MBE
292 and RMSE throughout the seasonal cycle from including the anthropogenic heat flux in JULES. Whilst the
293 improvement to the MBE is fairly constant throughout the year for the day-time results, the improvement in the night-
294 time MBE and the day-time MBE and RMSE is greater in the winter months, when Q_H has its smallest average diurnal
295 values, compared to the summer months.

296

297 A positive impact is also evident for Q_H during the night-time, with the consistent negative bias from the simulations
298 without Q_F reduced when the Q_F term is included in the surface energy balance of JULES. These changes also lead to a
299 greater reduction in the night-time RMSE in the winter months than in the summer, as might be expected since Q_H has
300 its lowest values.

301

302 The impact of including Q_F on Q^* is not so beneficial, leading to a slight degradation in both the MBE and RMSE.
303 However, this negative impact of Q_F within the JULES simulations is more indicative of issues with the radiation
304 balance rather than a direct influence of Q_F itself. It is possible that if the radiation issues within JULES were improved,
305 the impact of including Q_F might also give beneficial results for Q^* .

306

307 There are also negative impacts on the errors for the net storage heat flux, with both the MBE and RMSE being
308 degraded when Q_F is included in the simulations. The negative bias in ΔQ_S during the day-time throughout the year,
309 along with the positive night-time bias, suggests that insufficient energy is being stored during the day-time and
310 subsequently released during the night-time. This is consistent with the results presented by Best and Grimmond (2014)
311 who suggested that urban models have energy partitioning issues in general between Q_H and ΔQ_S .

312

313 Whilst the inclusion of the anthropogenic heat flux leads to some improvements within the simulations, the magnitude
314 of these improvements is small, even to the extent that it is difficult to identify the changes when looking at the full
315 diurnal cycle of the fluxes. The magnitude of the anthropogenic heat flux within the observational dataset has an
316 average value of 11 W m^{-2} , which is typical of suburban areas. This could suggest that Q_F is not significant in these
317 environments, especially since the minimum values of Q_F occur at similar times during the night to the smallest values
318 of the sensible heat flux. However, the temporal variations in Q_F used in this study were only the mean monthly
319 variations and do not take into account the meteorological conditions, or human behavior such as weekday and weekend
320 activities, hence the actual variations in Q_F might have a greater impact. In addition, for urban areas in colder climates,
321 Q_F could contribute a relatively larger fraction to the surface energy balance, due to the smaller size of radiative fluxes,
322 and hence be more significant. The flux is also known to be much greater in dense, urban centres (e.g., Ichinose et al.,
323 1999).

324

325 The seasonal variations in Q_F are also small, which explains why the increased positive impact from including the flux
326 during the winter months is also small. The small magnitude and diurnal cycle of the anthropogenic heat flux at this site
327 could well be responsible for this result. It is likely that at sites with large variations in both the diurnal and seasonal
328 cycles in Q_F there will be additional benefits from resolving the temporal behavior of the anthropogenic heat flux.

329
330 The impact of including Q_F is much smaller on both the MBE and RMSE for all of the data and night-time data only,
331 than including a representation of vegetation for the site. Again this result is influenced by the relatively small
332 magnitude of the anthropogenic heat flux. However, nearly all of the models that neglected Q_F in PILPS-Urban did
333 include a representation of vegetation, whilst other categories that included Q_F contained a greater proportion of models
334 that neglected vegetation. Hence the counter intuitive results presented in the PILPS-Urban, suggesting that the group of
335 models that did not include the anthropogenic heat flux performed at least as well as the models that did include this
336 flux, were being influenced by the treatment of vegetation within these models.

337
338 From the results presented in this study we can conclude that a representation of the anthropogenic heat flux is
339 important for urban models and can lead to improved results. Moreover, the influence of the anthropogenic heat flux is
340 likely to be greater at sites with a larger flux, increasing the need for the urban models to include this term in their
341 surface energy balance.

342
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348 simulations to the comparison.

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421

422 **Figure captions**

423

424 Fig. 1: Median of the mean modelled flux, mean bias error (MBE), and root mean square error (RMSE) for the surface
425 fluxes from all models within PILPS-Urban, determined for two month periods, for the models classified by their
426 representation of the anthropogenic heat flux (None - Q_F neglected, Prescribed - Q_F prescribed during model simulation,
427 Internal Temperature - models use an internal building temperature to calculate Q_F , Prescribed + Internal T -
428 combination of prescribed Q_F (e.g. vehicular component) and internal building temperature to calculate additional Q_F ,
429 Modelled - all aspects of Q_F modelled). Note the scales are different for each flux. For details of PILPS-Urban see
430 section 2, and for method of analysis see Best and Grimmond (2013).

431

432 Fig. 2: Ranked model performance for stage 4 of PILPS-Urban, for each of the surface fluxes. The four sets of results
433 from the JULES model are highlighted. These were generated from the 1- and 2-tile configurations, each run with low
434 (dry) and high (wet) initial soil moisture. Adapted from Grimmond et al. (2011).

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437 Fig. 3: Temporal variations in the anthropogenic heat flux applied to the urban surface fraction for the model
438 simulations: (a) mean for observational period, (b) mean monthly value, (c) mean diurnal cycle for observational
439 period, (d) monthly mean diurnal cycle, (e) mean diurnal cycle for observational period applied to the urban surface
440 fraction (dashed line), with the average diurnal cycles for the months with maximum and minimum values (solid lines).

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442 Fig. 4: MBE (a-c) and RMSE (d-f) for net all-wave radiation (a,d), sensible heat flux (b,e) and latent heat flux (c,f), for
443 both day- (filled) and night-time (hollow), for all implementations of temporal variation in the anthropogenic heat flux,
444 and simulations without vegetation. Results are presented for the 1 and 2 tile configurations with the anthropogenic heat
445 flux applied to only the canyon, only the roof, or both canyon and roof. Note the scales are different for each flux.

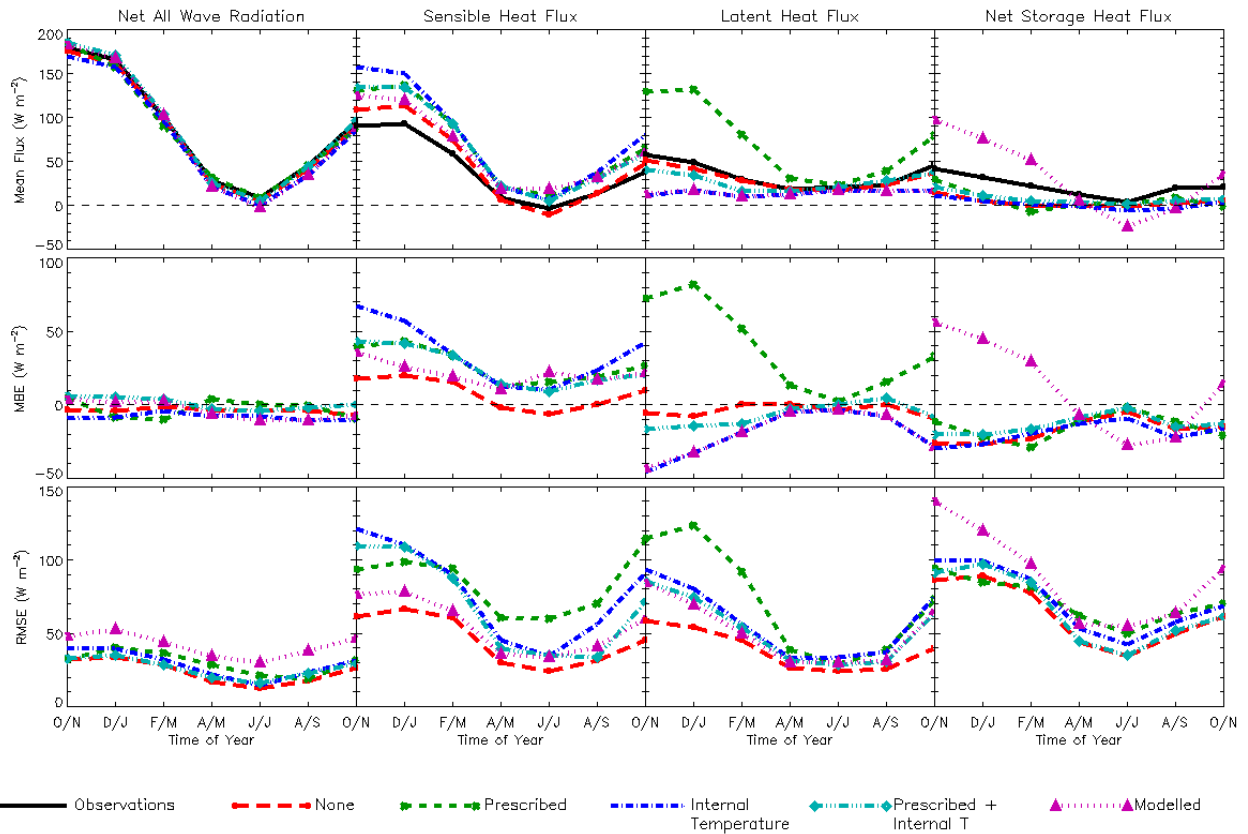
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447 Fig. 5: Nocturnal mean modelled flux (row 1), MBE (row 2), and RMSE (row 3) for the surface fluxes determined for
448 two month periods.

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450 Fig. 6: Mean diurnal cycle for each 60-day period throughout the seasonal cycle, scaled to focus on the night-time
451 results, for the sensible and net storage heat fluxes. Note the scales are different for both fluxes.

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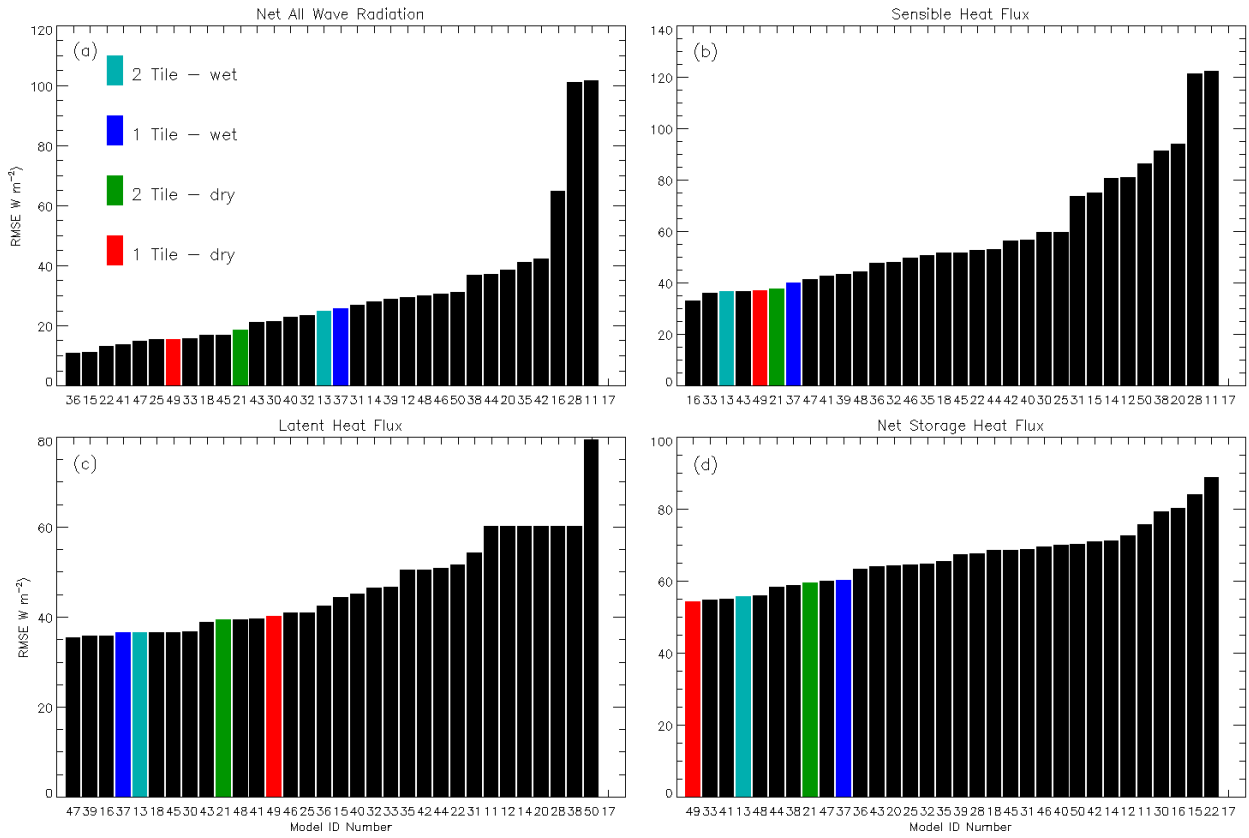
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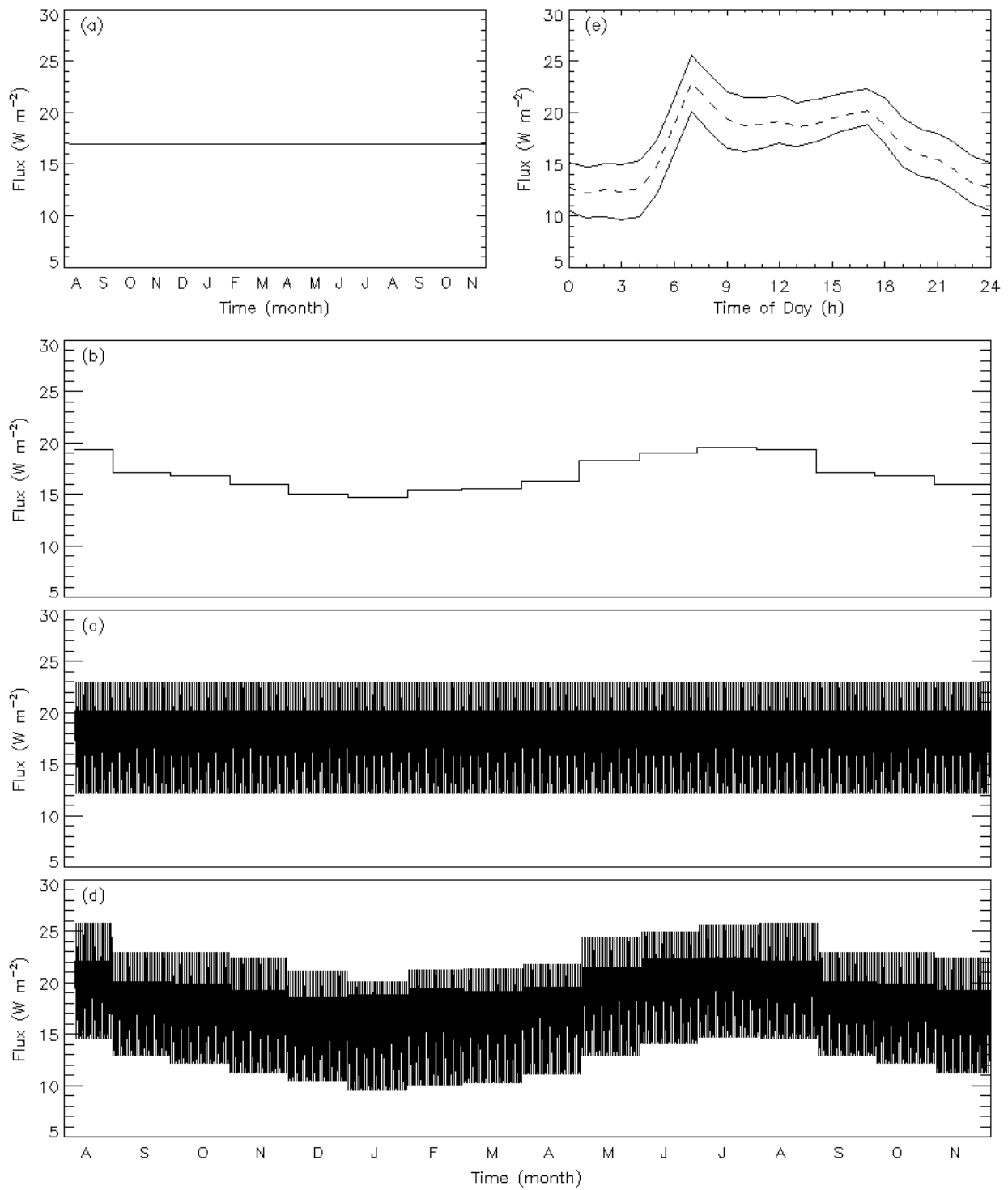
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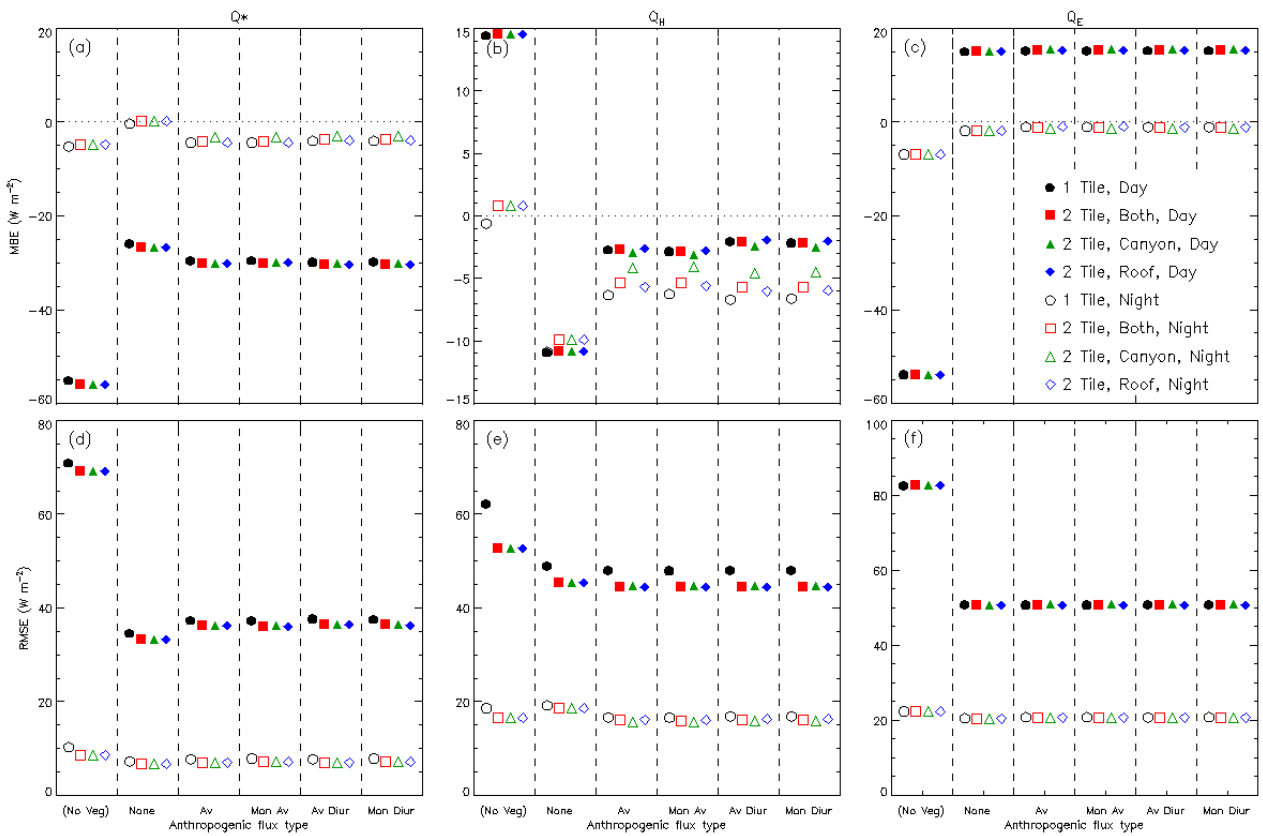


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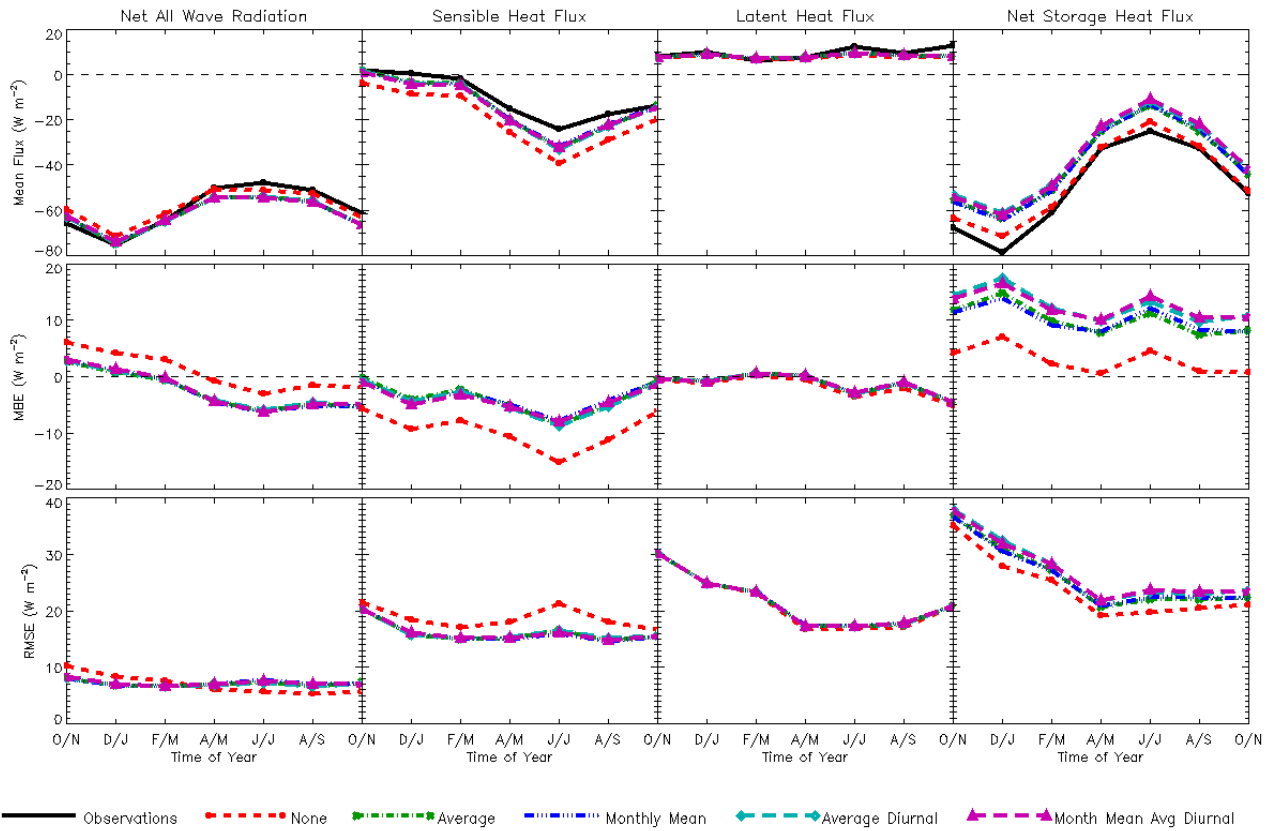
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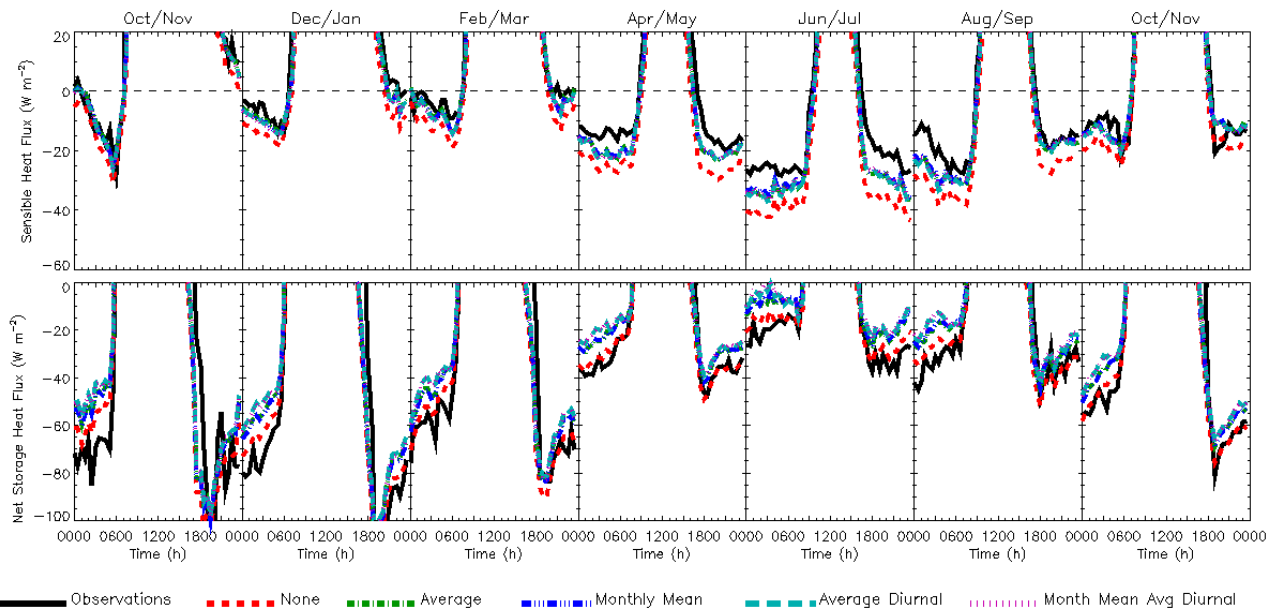


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