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1 Title: Alpha and Gammaproteobacterial Methanotrophs Co-Dominate the Active

2 Methane Oxidizing Communities in an Acidic Boreal Peat Bog

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16 Running Title: Co-Dominant Active Methanotrophs in a Boreal Peat Bog

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24 **Abstract**

25 The objective of this study was to characterize metabolically active, aerobic  
26 methanotrophs in an ombrotrophic peatland in the Marcell Experimental Forest,  
27 Minnesota, USA. Methanotrophs were investigated in the field and in laboratory  
28 incubations using DNA-stable isotope probing, expression studies on particulate methane  
29 monooxygenase (*pmoA*) genes, and amplicon sequencing of 16S rRNA genes. Potential  
30 rates of oxidation ranged from 14-17  $\mu\text{mol CH}_4 \text{ g dry wt soil}^{-1} \text{ d}^{-1}$ . Within DNA-SIP  
31 incubations, the relative abundance of methanotrophs increased from 4% *in situ* to 25-  
32 36% after 8 -14 days. Phylogenetic analysis of the  $^{13}\text{C}$ -enriched DNA fractions revealed  
33 the active methanotrophs were dominated by the genera *Methylocystis* (Type II;  
34 Alphaproteobacteria), *Methylomonas*, and *Methylovulum* (Type I;  
35 Gammaproteobacteria). In field samples, a transcript-to-gene ratio of 1 to 2 was observed  
36 for *pmoA* in surface peat layers which attenuated rapidly with depth, indicating the  
37 highest methane consumption was associated with the 0-10 cm depth interval.  
38 Metagenomes and sequencing of cDNA *pmoA* amplicons from field samples confirmed  
39 the dominant active methanotrophs were *Methylocystis* and *Methylomonas*. Although  
40 Type II methanotrophs have long been shown to mediate methane consumption in  
41 peatlands, our results indicate members of the genera *Methylomonas* and *Methylovulum*  
42 (Type I) can significantly contribute to aerobic methane oxidation in these ecosystems.

43

44 **Introduction**

45 Methane is the third most important greenhouse gas and has 28 times the potential  
46 of carbon dioxide to trap heat radiation on a molecular basis over a 100 year time scale  
47 (1, 2, 3). Wetlands, such as peatlands, represent the largest natural source of methane to  
48 the atmosphere (4). Aerobic methanotrophic bacteria live at the oxic-anoxic interface of  
49 wetland soils and it has been shown that they consume as much as 90 % of the methane  
50 produced belowground before it reaches the atmosphere, thus serving as a biofilter  
51 regulating emissions (3, 5, 6, 7). The response of methane dynamics in wetlands to  
52 global climate change is uncertain, and climate models would be improved through  
53 quantification of the response of microbially-mediated mechanisms of methanotrophy to  
54 temperature and moisture variation.

55 Aerobic methanotrophs are phylogenetically located in two phyla: the  
56 Proteobacteria and Verrucomicrobia (8). The majority of characterized methane-  
57 oxidizing organisms have been separated into Type I methanotrophs of the  
58 Gammaproteobacteria and Type II methanotrophs of the Alphaproteobacteria (9, 10, 11).  
59 The prevailing view has been that most methanotrophs only grow on methane or  
60 methanol as a source of carbon and energy (4). However, more recently, a number of  
61 Type II methanotrophs (*Methylocella*, *Methylocapsa*, and *Methylocystis*) have been  
62 characterized as facultative methanotrophs capable of conserving energy for growth on  
63 multi-carbon compounds such as acetate, pyruvate, succinate, malate, and ethanol (12).  
64 Although members of the phylum Verrucomicrobia have been widely detected in  
65 peatlands, none have been definitively linked to methanotrophy and thus more research is

66 needed to ascertain the role of Verrucomicrobia in the carbon cycle of peatlands (8, 13,  
67 14, 15, 16).

68 Aerobic methane oxidation in Proteobacterial methanotrophs is catalyzed by the  
69 enzyme methane monooxygenase (MMO), either particulate MMO (pMMO) or a soluble  
70 MMO (sMMO). The genes *pmoA* (encoding the 27kDa subunit of pMMO) and *mmoX*  
71 (encoding the alpha-subunit of the hydroxylase of sMMO) as well as 16S rRNA genes  
72 have been used most often as molecular markers to characterize methanotrophs in  
73 peatlands and other environments (6, 17, 18, 19, 20, 21, 22, 23). Previous studies point to  
74 a co-dominance of Alpha- and Gammaproteobacteria, and in particular, the genera  
75 *Methylocystis* and *Methylomonas*; however, the majority of past work in the field was  
76 conducted at the DNA level, and less information is available on which microbial  
77 populations are actively involved in methane oxidation *in situ*. Using a combination of  
78 stable isotope probing (SIP) and a functional gene (*pmoA*) array, Chen *et al.* (24)  
79 determined that *Methylocystis* populations predominated the active methanotrophs in a  
80 range of peatlands in the UK. Phospholipid fatty acid stable isotope analysis (PLFA-SIP)  
81 was also utilized in conjunction with mRNA analyses to probe the active methanotrophic  
82 communities in peatlands in the UK, again finding a community dominated by  
83 *Methylocystis* (6). Gupta *et al.* (7) also detected a predominance of Type II  
84 methanotrophs (*Methylocystis*, *Methylosinus*, *Methylocapsa*, and *Methylocella*) in a  
85 peatland in New York, USA, using SIP. As reviewed by Dedysh (25), a number of  
86 acidophilic and acidotolerant Type II methanotrophs have been cultivated from peatlands.  
87 The first acid-tolerant Type I methanotroph was only recently isolated and described by  
88 Danilova *et al.* (26); however, cultivation of both types might suggest their involvement

89 in active methane oxidation.

90 The objective of this study was to identify the microorganisms actively involved  
91 in methane oxidation in climatically sensitive boreal peatlands using multiple,  
92 independent molecular approaches in the field and laboratory. Based on previous studies,  
93 it was hypothesized that the Alphaproteobacterial methanotrophs were most active in  
94 methane oxidation with only a minor contribution from the Gammaproteobacterial  
95 methanotrophs. This study was conducted at the Marcell Experimental Forest (MEF) in  
96 northern Minnesota, USA, where the U.S. Department of Energy (DOE) Oak Ridge  
97 National Laboratory and the USDA Forest Service are conducting a large-scale field  
98 climate manipulation known as Spruce and Peatland Response Under Climatic and  
99 Environmental Change (SPRUCE).

100

## 101 **Materials and Methods**

### 102 *Site Description and Sample Collection*

103 Peat samples were collected at the S1 Bog located in the Marcell Experimental  
104 Forest (MEF; N 47°30.476'; W 93°27.162') north of Grand Rapids, MN (27). This site  
105 has been described in detail in other publications (14, 15). The S1 bog is acidic with an  
106 average pH of 3.5 – 4.0 and is oxygen limited with oxygen levels decreasing to below  
107 detection (limit of approximately 20 ppb) within the top five centimeters of the bog (14).

108 For use in DNA-SIP incubations, a 10 x 10 x 10 cm block of peat, approximately  
109 1 liter in volume, was sampled using a sterilized bread knife in hollows from the S1 bog,  
110 transect 3, in July 2012. The collected peat was homogenized by hand in a sterile bag and  
111 stored at 4°C until use in experiments. Samples for nucleic acid extraction were collected

112 in triplicate with a Russian peat corer as described by Lin *et al.* (14, 15). Each core was  
113 sub-sectioned into 10 cm intervals and immediately placed on dry ice. Samples were then  
114 stored at -80°C in a portable freezer until nucleic acid extractions were performed. DNA  
115 and RNA extractions were performed with MoBio PowerSoil DNA and Total RNA  
116 Extraction kits, respectively, according to the manufacturer's instructions.

117 *Microcosm incubations with <sup>13</sup>C-labelled methane*

118 Ten grams of homogenized peat from the 0 – 10 cm depth interval of S1 peat bog  
119 midway along the third transect (S1T3M) were added to 150 mL serum bottles in  
120 duplicate for each treatment. This site was chosen for consistency with other field  
121 samples obtained from the S1 peat bog. Bottles were sealed with blue-butyl rubber  
122 stoppers and crimped with aluminum crimp seals. Samples were stored in the dark at  
123 room temperature (appx. 24°C). Treatments included those for which the headspace was  
124 amended with either 1% (vol/vol) 99.9% <sup>12</sup>C-CH<sub>4</sub> (Sigma) or 1% (vol/vol) 99.9% <sup>13</sup>C-  
125 CH<sub>4</sub> (Sigma). This concentration is higher than *in situ* levels of methane to obtain enough  
126 labeled DNA for subsequent analyses. Headspace concentrations were monitored with a  
127 gas chromatograph – flame ionization detector (GC-FID) equipped with a methanizer  
128 over two weeks of incubation. Analysis of headspace gas (150 µL) was performed on a  
129 Shimadzu GC-2014 with a Supelco custom packed column (Packing 80/100 Hayesep Q).  
130 The flow rate was 30 mL/min with the injector and detector at 100°C, the column at  
131 40°C, and the methanizer at 380°C. Samples were taken on the day of preparation (day 0)  
132 and subsequently after 3, 8, 11, and 14 days. The samples were not replaced due to the  
133 relatively short incubation and minimal headspace sampling. In parallel, <sup>12</sup>C and <sup>13</sup>C-CH<sub>4</sub>  
134 samples were sacrificed at the initiation of the experiment (T0), after eight days (T1), and

135 after fourteen days (T2). A subsample of 5 grams was removed from each sample and  
136 frozen at -80°C until DNA was extracted for further analysis.

137 Wet to dry weight was determined by weighing out ~5 grams of peat from the  
138 incubation. Samples were then dried in a drying oven at 60°C until a stable mass was  
139 obtained (appx. 7 days).

#### 140 *Stable Isotope Probing – ultracentrifugation and gradient fractionation*

141 DNA was extracted from frozen peat samples with the Mo Bio Powersoil DNA  
142 kit according to the manufacturer's protocol and stored at -20°C until further analysis.

143 Stable isotope probing was conducted as described previously (28, 29). In brief, extracted  
144 DNA was added to a cesium chloride solution and centrifuged by ultracentrifugation at  
145 177,000 x g. After 40 hours, samples were removed from the ultracentrifuge and  
146 fractionated by needle fractionation into twelve or thirteen fractions and the density of  
147 each fraction was determined with a digital refractometer (Reichert AR200). <sup>13</sup>C-enriched  
148 DNA was expected within the "heavy" fractions (five to eight). DNA was precipitated  
149 from all fractions with polyethylene glycol and glycogen as a carrier (28, 29).

150 Precipitated DNA was stored at -20°C until further analysis.

#### 151 *Microbial community characterization in SIP incubations.*

152 All fractionated DNA samples from 8 and 14 day time points were fingerprinted  
153 with Automated Ribosomal Intergenic Spacer Analysis (ARISA). First, ARISA PCR was  
154 run on each sample with the S-D-Bact-1522-b-S-20 and L-D-Bact-132-a-A-18 primers  
155 (30). PCR reactions were performed with an initial denaturation step at 95°C for 5  
156 minutes followed by 30 cycles of 30 seconds at 94°C, 1 minute at 52°C, and 1 minute at  
157 72°C followed by a 10 minute final elongation step at 72°C. PCR products were run on a

158 1.5% w/v agarose gel with 1xTris/Borate/EDTA (TBE) buffer and successful reaction  
159 products were cleaned with the Mo Bio PCR cleanup kit following the manufacturer's  
160 instructions. ARISA PCR products were then separated and analyzed using an Agilent  
161 model 2100 Bioanalyzer and unique bands in heavy fractions were noted from  $^{13}\text{C}\text{-CH}_4$   
162 samples to determine the success of the SIP incubation.

163 Quantitative PCR (qPCR) was used to determine the abundance of *pmoA* genes in  
164 different DNA fractions retrieved from the  $^{13}\text{C}\text{-CH}_4$  incubated soil samples (after 14 days  
165 of incubation). To minimize effects of inhibitors, DNA from SIP fractions was diluted to  
166 1/20 of original concentrations. All fractions were analyzed with A189f/Mb661r PCR  
167 primers as described by Kolb *et al.* (31) to target the abundance of *pmoA* genes in 20 $\mu\text{L}$   
168 reactions with 2  $\mu\text{L}$  of template DNA (2.3 – 8.8 ng/ $\mu\text{L}$ ) added to a master mix of 10  $\mu\text{L}$  of  
169 Sybr green to a final concentration of 1x, 1.6  $\mu\text{L}$  each of forward and reverse primers to a  
170 final concentration of 0.8  $\mu\text{M}$ , and 4.8  $\mu\text{L}$  of PCR grade water. Samples were run against  
171 a standard curve in a StepOnePlus instrument with 96 wells with the following  
172 parameters: an initial denaturation step of 5 minutes at 95°C and 40 cycles of  
173 denaturation at 95°C for 15s, annealing at 64°C for 45s, extension at 72°C for 45s, and  
174 data acquisition at 86.5°C for 16s. Quantity of *pmoA* genes was normalized to the  
175 abundance of 16S rRNA genes, which were analyzed as described in Lin *et al.* (15) with  
176 1/20 dilution of DNA.

177 DNA fractions from one  $^{13}\text{C}\text{-CH}_4$  sample from T1 (8 day incubation) and one  $^{13}\text{C}\text{-}$   
178  $\text{CH}_4$  sample from T2 (14 day incubation) were sequenced on an Illumina MiSeq platform  
179 at the Michigan State Sequencing Facility with the 515F/806R primer set (32). Sequences  
180 were analyzed in QIIME 1.8 (33) as follows: overlapping reads were merged with fastq-



181 join (34) and quality filtering was performed with USEARCH 7.0, rejecting reads with an  
182 expected error greater than 0.5 (35). Read length was limited to approximately 250 base  
183 pairs after primer removal with the inclusion of only completely assembled reads.  
184 Subsequently, initial OTUs from the 16S rRNA gene sequence reads were picked *de novo*  
185 based on 97% similarity and representative sequences were picked from each OTU.  
186 Representative sequences were filtered by comparison to the Greengenes database  
187 (<http://greengenes.lbl.gov/cgi-bin/nph-index.cgi>) at 60% similarity. An OTU table was  
188 compiled and filtered by removing phylotypes comprising less than 0.05% of the library.  
189 Taxonomy was assigned to the parsed OTU table with the Greengenes database. 16S  
190 rRNA gene sequences assigned to the phylum Proteobacteria were screened for  
191 methanotrophs with a maximum-likelihood phylogenetic tree. Sequences were aligned  
192 and identity determined at 95% similarity to nearest neighboring sequence in SILVA  
193 (36). The methanotrophs represented in DNA-SIP samples were identified with a nucleic  
194 acid maximum-likelihood tree with bootstrap analysis (1,000 replications). Analysis of  
195 variance and regression analysis of the shift in methanotroph community composition in  
196 heavy and light fractions was conducted in R to test for significant differences in  
197 methanotroph populations (37).

#### 198 *Metagenomic analysis of field samples*

199 Libraries for metagenomic sequencing were generated from field DNA extracts  
200 using the Nextera DNA sample preparation kit (Illumina, Inc. San Diego, CA) as in Lin  
201 et al. (15). Libraries were size-selected using E-Gels (Life Technologies, Inc.) for an  
202 insert size range of 400-800 bp. Libraries were then quantified and quality checked using  
203 the Invitrogen Qubit and Agilent Bioanalyzer. FASTQ files from the metagenomic

204 sequencing were loaded into the MG-RAST server for quality filtering and downstream  
205 analysis (38). The paired-end reads from each library were joined and then filtered with  
206 the default parameters. Protein and pathway search was performed with SEED  
207 annotations (e-value of  $10^{-5}$ ) in MG-RAST. The amino acid sequences derived from the  
208 gene calling were downloaded for searching the *pmoA* gene. A Hidden Markov Model  
209 (HMM) for the *pmoA* gene was created by using HMMER v3.0 tools  
210 (<http://hmmer.janelia.org/>), based on HMM training sequences downloaded from the  
211 functional gene pipeline and repository (<http://fungene.cme.msu.edu/>). All HMM search  
212 hits with e-values below a threshold of  $10^{-5}$  were counted and retrieved. For the  
213 taxonomic assignment of gene sequences, the corresponding BLASTP search outputs  
214 were uploaded for analysis through the software Metagenome Analyzer (MEGAN) (39).

#### 215 *Quantification of pmoA genes and transcripts in field samples*

216 Quantification of *pmoA* genes and transcripts was conducted with DNA and  
217 cDNA extracted from field samples, respectively, according to methods described above  
218 for analysis of SIP fractions. cDNA standards and cDNA of environmental RNA samples  
219 were synthesized using GoScript Reverse Transcription system following the  
220 manufacturer's protocol (Promega). The *pmoA* gene fragment used for constructing  
221 plasmid standards of qPCR was amplified from genomic DNA of *Methylococcus*  
222 *capsulatus* Bath. The plasmid standard was prepared according to Lin *et al.* (40). To  
223 prepare cDNA standards, plasmid DNA with a positive *pmoA* insert was linearized with  
224 *NcoI* restriction enzyme following the manufacturer's protocol (Promega), and purified  
225 by MinElute PCR purification kit (Qiagen Inc., CA). RNA was synthesized from the  
226 linearized plasmid DNA using the Riboprobe *in vitro* transcription system according to

227 the manufacturer's protocol (Promega).

228 Power SYBR® Green PCR Master Mix was used for all qPCR assays. Plasmid  
229 DNA or cDNA standards with inserts of specific gene fragments were used to establish  
230 standard curves that were included in each run. The standard contains different quantities  
231 of cloned gene fragments, spanning 7 orders of magnitude from  $10^1$  to  $10^7$  gene copies  
232 per PCR well. To minimize the effects of inhibitors in assays, peat DNA was diluted to  
233 1/40 of original concentrations, and duplicate 20ul reactions each containing 2  $\mu$ l of  
234 diluted DNA were run for each sample. The *pmoA* amplicons from the synthesized  
235 environmental cDNA were sent to the University of Illinois at Chicago (UIC) for DNA  
236 sequencing using a 454 platform.

237 Raw *pmoA* sequences were demultiplexed, trimmed, and quality filtered in  
238 CLCbio.

#### 239 *Sequence Accession Numbers*

240 Gene sequences from the analysis of SIP fractions have been submitted to the  
241 GenBank database under accession number PRJNA286313. Metagenomes have been  
242 submitted to MG-RAST under identification numbers 4538779.3, 4538778.3, and  
243 4538997.3. Amplicon sequences for the *pmoA* gene are in the process of submission and  
244 are available upon request.

245

## 246 **Results**

### 247 *Abundance, activity and community composition of methanotrophs in the field.*

248 The abundance of *pmoA* genes and transcripts decreased rapidly with depth in the  
249 peat column, decreasing by two orders of magnitude from 0 to 100 cm depth (Figure 1).

250 The transcript-to-gene ratio, a proxy for *pmoA* expression or the activity of  
251 methanotrophs, decreased to background levels below 40 cm depth.

252 Multiple lines of evidence allowed us to determine the dominant methanotrophs  
253 in the S1 bog soils (Figure 2). In all of the soils sampled, sequences affiliated with the  
254 genus *Methylocystis* comprised over 75 % of the *pmoA* sequences retrieved from  
255 metagenomes and sequenced amplicons derived from cDNA (Figure 2). Overall, the  
256 remaining *pmoA* sequences were mainly affiliated with the genus *Methylomonas*. At  
257 mid-depth (approximately 30 cm), *Methylosinus*-like sequences showed a higher relative  
258 abundance compared to *Methylomonas*-like sequences in the metagenomes, while  
259 *Methylomonas*-like sequences were second in relative abundance to *Methylocystis* in  
260 cDNA amplicons. However, this could be due to the slightly different depths sampled or  
261 temporal variability, since the metagenomes and cDNA amplicons were sampled in  
262 successive years. However, microbial community composition was shown to be  
263 temporally stable at the DNA level in extensive field studies of the S1 bog (14, 15).

264

265 *Stable isotope probing incubations.*

266 Within microcosm incubations, the most rapid methane oxidation rates were  
267 observed within the first three days of incubation at room temperature (approximately  
268 24°C). Rates ranged from 13.8 to 17.3  $\mu\text{mol CH}_4 \text{ g dwt}^{-1} \text{ d}^{-1}$ . Samples amended with  $^{13}\text{C}$ -  
269  $\text{CH}_4$  and  $^{12}\text{C}$ - $\text{CH}_4$  demonstrated potential consumption rates of  $15.1 \pm 2.3 \mu\text{mol CH}_4 \text{ g}$   
270  $\text{dwt}^{-1} \text{ d}^{-1}$  and  $15.9 \pm 1.6 \mu\text{mol CH}_4 \text{ g dwt}^{-1} \text{ d}^{-1}$ , respectively. Rates of methane consumption  
271 were calculated in Excel utilizing the linest function from three point linear regions in  
272 methane depletion with time. After two weeks, nearly all of the methane in the headspace

273 had been consumed (Figure 3). Samples were sacrificed for DNA-SIP analysis after eight  
274 days for time point one, after peak methane consumption rates were observed, and after  
275 fourteen days for time point two, when nearly all of the methane had been consumed.

276           Fingerprinting of the fractionated DNA was conducted with automated ribosomal  
277 intergenic spacer analysis (ARISA) which showed a clear shift in microbial communities,  
278 indicating incorporation of  $^{13}\text{C}$  into DNA within “heavy” fractions 7 and 8  
279 (Supplementary Figures 1 & 2). Potential enrichment of active methanotrophs within  
280 these fractions was supported by a relative enrichment in *pmoA* gene abundance as  
281 determined by qPCR (Supplementary Figure 3). Within  $^{13}\text{C}$ -enriched fractions from  
282 DNA-SIP incubations, an enrichment of Proteobacteria was observed relative to the  $^{12}\text{C}$ -  
283 enriched fractions. The relative abundance of methanotrophs in the overall community  
284 increased from approximately 4 % in the field samples (data not shown) to 36 % in the  
285  $^{13}\text{C}$ -enriched fractions after 8 days of incubation (Figure 4). The shift in the abundance of  
286 the methanotrophic community between  $^{13}\text{C}$ -enriched fractions and  $^{12}\text{C}$ -fractions was  
287 shown to be significant with ANOVA analysis (F-value = 7.144, df = 3, p = 0.0439).  
288 Phylogenetic analysis of 16S rRNA gene sequences showed a co-dominance of  
289 Alphaproteobacterial and Gammaproteobacterial methanotrophs, and phylotypes were  
290 most closely related to the rRNA genes of the genera *Methylocystis*, *Methylomonas*, and  
291 *Methylovulum* (Figure 4, 5). None of the phylotypes were closely related to cultivated  
292 members of each genus. Environmental sequences obtained in other peat bogs were  
293 similar to the sequences enriched in  $^{13}\text{C}$ ; however, phylotypes most closely related to  
294 *Methylomonas* and *Methylovulum* in particular remained phylogenetically distinct while  
295 phylotypes related to *Methylocystis* were phylogenetically similar to sequences obtained

296 in other acidic forest and peat soils (Figure 5). Methanotrophic genera *Methylocella* and  
297 *Methyloferula*, possessing only the soluble methane monooxygenase, were not detected.

298

## 299 **Discussion**

300 The community composition and activity of methanotrophic bacteria was  
301 interrogated in field samples from the S1 bog at Marcell Experimental Forest in  
302 Minnesota using several cultivation-independent approaches. Expression of a key gene  
303 in the methane oxidation pathway (*pmoA*, encoding a subunit of particulate methane  
304 monooxygenase) was used as a proxy for the activity of methanotrophs in the bog.  
305 Although *pmoA* genes could be detected throughout the peat column, indicating the  
306 presence of methanotrophs, transcript abundance decreased with depth and no *pmoA*  
307 transcripts were detected below the 30-40 cm depth interval. Biogeochemical  
308 characteristics of the S1 peat bog have been described in detail by Tfaily *et al.* (27)  
309 including distinct layers within the peat column encompassing the acrotelm (0-30 cm),  
310 mesotelm (30-75 cm), and catotelm (75 cm and deeper). Oxygen diffusion is limited  
311 within the bog due to the height of the water table but may extend lower in the acrotelm  
312 due to zones of aeration within the rhizosphere of plant roots (3, 14). Thus, in parallel  
313 with the availability of oxygen, methanotroph activity was highest at the surface and  
314 limited to the acrotelm and mesotelm. Although few studies have examined *pmoA*  
315 expression in wetlands, Freitag *et al.* (41) observed that transcript-to-gene ratios reflected  
316 methane dynamics in a UK peatland. Transcript-to-gene ratios of this study were in  
317 agreement with those determined by Freitag *et al.* (41). It should be noted that this study  
318 does not address microbial groups that mediate anaerobic methane oxidation (AOM),

319 during which methanotrophy is coupled to utilization of alternate electron acceptors such  
320 as sulfate, nitrate, or nitrite (42, 43, 44). Given the scarcity of these alternate electron  
321 acceptors in the S1 bog (14), anaerobic methane oxidation would not be favored.  
322 However, further studies of AOM are warranted in this ecosystem.

323 Independent lines of evidence revealed the identity of active methanotrophs in the  
324 surface (0 to 10 cm depth), where the highest methane oxidation activity was detected.  
325 Results from metagenomes and next generation sequencing of *pmoA* cDNA amplicons  
326 from field samples indicated a predominance of the Type II methanotroph, *Methylocystis*,  
327 at the surface. Since the sequencing of cDNA amplicons was conducted on the same  
328 samples as those used for qPCR, the community composition should be directly  
329 comparable to gene expression determinations. *Methylocystis* comprised over 75 % of the  
330 metabolically active methanotrophs, with the Type I methanotroph *Methylomonas*  
331 making up the remainder of the active community. Thus, these groups appear to be the  
332 most abundant and the most active in mediating aerobic methanotrophy in the S1 bog.  
333 Previous metagenomic analysis of field samples from the S1 peat bog suggested the  
334 potential involvement of *Methylocystis* and, to a lesser extent, *Methylomonas* in methane  
335 oxidation processes (14, 15). The data presented in this study suggest not only active  
336 involvement of both of these genera in methane oxidation but also the involvement of a  
337 second Gammaproteobacterial methanotroph, *Methylovulum*.

338 Field results were confirmed in the laboratory using a combination of stable  
339 isotope probing and next generation sequencing of 16S rRNA genes in a series of  
340 microcosm incubations. The active methanotrophic community was composed of a  
341 combination of *Methylocystis*, *Methylomonas*, and *Methylovulum*, which were observed

342 to shift with time (Figure 4). As indicated previously, the presence of *Methylocystis* was  
343 not surprising given the well-documented presence, activity, and cultivated isolates from  
344 this methanotroph in acidic peatland ecosystems (6, 17, 21, 24). The presence and  
345 abundance of *Methylomonas* and *Methylovulum* were more surprising. While  
346 *Methylomonas* has been detected in amplicon sequences and cultured from peatlands, this  
347 genus has not been definitively linked to active methane oxidation in acidic boreal  
348 peatlands (23, 26, 45, 46). Through the use of SIP, *Methylomonas* has been shown to be  
349 active in methane oxidation in other environments such as a cave system, a soda lake, and  
350 landfill cover soil (47, 48, 49) that are more neutral to alkaline in pH. Studies on  
351 methanotrophs in peatlands have utilized a variety of methods including both cultivation-  
352 dependent and cultivation-independent such as diagnostic microarrays, PLFA-SIP, clone  
353 libraries, and DNA-SIP; however, this study utilized DNA-SIP experiments where the  
354  $^{13}\text{C}$ -enriched DNA obtained was directly sequenced in combination with metagenomic  
355 and cDNA analysis of field samples. Within the top 10 cm of the S1 bog, potential rates  
356 of methanogenesis only reach  $0.025 \mu\text{mol CH}_4 \text{ g dwt}^{-1} \text{ d}^{-1}$  (27). If these potential rates are  
357 representative of *in situ* rates of methanogenesis, the methane concentrations in the  
358 headspace of SIP incubations after 14 days were more representative of the natural  
359 environment, lending greater significance to the observed shifts in populations of  
360 methanotrophs present (Figure 3, 4). The combination of amplicon sequencing of SIP  
361 enrichment samples and metagenomic sequence analysis of field samples, coupled with  
362 analysis of multiple time points enabled analysis of the Gammaproteobacterial (Type I)  
363 methanotroph community which can now be considered to be key active methane  
364 oxidizers in an acidic peatland ecosystem.



365 Perhaps most remarkable is the presence of *Methylovulum* in the active  
366 methanotrophic community. The first isolate of this genus, *Methylovulum miyakonense*,  
367 was obtained in 2011 and to date no new species within this genus have been  
368 characterized (50). Although originally isolated from forest soil, *M. miyakonense* was  
369 also isolated from peatland soil (51), suggesting *Methylovulum* is present in other  
370 peatlands. The cultivation of a *Methylovulum*-like methanotroph from another acidic peat  
371 bog by Kip *et al.* (45) further supports this possibility. However, the strains did not  
372 appear to grow under acidic conditions, begging the question of the extent of the role  
373 *Methylovulum* might be playing in acidic peatland soil (50, 51). To our knowledge, these  
374 are the first data directly linking *Methylovulum* to active methane cycling in peatlands.  
375 Although the relative abundance of *Methylovulum* was low in the SIP incubations, there  
376 was a distinct enrichment in the <sup>13</sup>C-enriched samples compared to the <sup>12</sup>C-enriched  
377 samples suggesting active methane consumption (Supplementary Figure 4). While  
378 microcosm experiments may induce enrichment of organisms normally low in abundance  
379 *in situ*, other methanotrophic organisms detected at low abundance in the metagenomes,  
380 such as *Methylosinus*, were not enriched with <sup>13</sup>C over the course of the incubation,  
381 suggesting *Methylovulum* actively participates in methane consumption, if at low  
382 abundance.

383 Several possibilities may explain a seemingly neutrophilic methanotroph participating  
384 actively in methane oxidation in an acidic soil environment. One previous suggestion is  
385 the existence of neutral microenvironments, such as the plant endosphere, within the bog  
386 system, providing a specific ecological niche for *Methylovulum* (51). The *Methylovulum*  
387 16S rRNA gene sequences detected in our experiments were not closely related to *M.*  
388 *miyakonense*, suggesting the existence of as yet uncultivated members of this genus that  
389 may be acido-tolerant or acidophilic. *Methylovulum* from other environments has also  
390 been identified as potentially psychrotolerant and capable of oxidizing methane at low  
391 concentrations, suggesting adaptability of this organism to the changing environmental

392 conditions at the surface of boreal peatlands (52). This would not be the first example of  
393 an organism seemingly suited to one particular environment playing a role in a wholly  
394 different environmental system. Rahman *et al.* (53) showed in 2011 that *Methylocella*, a  
395 facultative methanotroph isolated from acidic soil, resides in many diverse environments  
396 encompassing a pH range of 4.3 to 10.0. Based on this example, it is not necessarily  
397 surprising to find a methanotroph commonly found in more neutral environments actively  
398 participating in methane oxidation in the acidic peat soil. Rather this would encourage  
399 further probing of the active microbial community, potentially with a transcriptomic  
400 approach, to more fully assess which microorganisms are present and active in each  
401 environmental system.

402 An important step in analyzing the potential impacts of changing climate on the  
403 methane cycle in peatlands is to first identify the microorganisms actively involved in  
404 methane cycling. These data take a step toward that goal by identifying the active  
405 methane oxidizing bacteria at the S1 bog in the Marcell Experimental Forest. Active  
406 methane oxidizers include representatives from both Alphaproteobacteria and  
407 Gammaproteobacteria, and we show for the first time that *Methylovulum* and  
408 *Methylomonas* are directly involved in methane oxidation at the surface of the peat bog.  
409 Using these data, the key bacteria involved in methane oxidation can be targeted for  
410 cultivation for future studies to examine the physiology of these organisms and  
411 subsequently the potential effects of climate change on this methane oxidizing  
412 community in boreal peat bogs.

413

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598

599 **Figure Legends**

600 **Figure 1:** *pmoA* gene and transcript abundance in copies\*g dwt<sup>-1</sup> at depth from the S1

601 peatland. Samples tested were from duplicate soil cores collected in July 2013. Error bars

602 represent standard deviation.

603 **Figure 2:** Methanotroph community composition at depth in the Spruce S1 peat bog  
604 based on metagenomic and cDNA analysis of *pmoA*. Methanotrophs detected included  
605 *Methylocystis* (dark grey), *Methylomonas* (light grey), and *Methylosinus* (white). Samples  
606 for metagenomic analysis were collected at Spruce in July 2012 (SP0712) and samples  
607 for cDNA analysis were collected at Spruce in July 2013 (SP0713).

608 **Figure 3:** The consumption of methane with time in the stable isotope probing  
609 incubations. Circles represent  $^{12}\text{C}$ - $\text{CH}_4$  treatments whereas triangles represent  $^{13}\text{C}$ - $\text{CH}_4$   
610 amended treatments. The observed methane consumption rates ranged between 13.85 and  
611  $17.26 \mu\text{mol CH}_4 \text{ g dwt}^{-1} \text{ d}^{-1}$  (calculated based on three-point linear region within each  
612 sample distribution). Peat utilized was from the 0 – 10 cm depth interval in hollows from  
613 the S1 bog, collected in July 2012.

614 **Figure 4:** The relative abundance of Alphaproteobacterial (Type II) and  
615 Gammaproteobacterial (Type I) methanotrophs based on 16S rRNA genes in  $^{13}\text{C}$ -  
616 enriched fractions (H) compared to light fractions (L) after 8 days (T1) and 14 days (T2)  
617 of incubation. The difference between methanotrophic communities in heavy and light  
618 fractions was significant based on ANOVA analysis (F-value = 7.144, df=3, p-value =  
619 0.0439).

620 **Figure 5:** Phylogeny of methanotrophs within SIP fractions from 8 and 14 day  
621 incubations (diamonds) showing organisms within the Alphaproteobacteria,  
622 *Methylocystis sp.*, and the Gammaproteobacteria, *Methylomonas* and *Methylovulum sp.*  
623 based on 16S rRNA gene analysis. This phylogenetic tree was prepared with the  
624 maximum-likelihood method with bootstrap analysis of nucleic acid sequences (1000

625 replications).

626











