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# Assessment of canopy photosynthetic capacity and estimation of GPP by using spectral vegetation indices and the light–response function in a larch forest

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## ABSTRACT

Integration of CO<sub>2</sub> flux observations with remote sensing technique and ecosystem modeling is expected to be useful for estimation of gross primary production (GPP). We focused on the changes in the two main parameters for the canopy-scale light–response curve— $P_{max}$  (maximum GPP at light saturation) and  $\phi$  (initial slope)—as indicators to represent canopy photosynthetic capacity. We hypothesized that  $P_{max}$  and  $\phi$  could be evaluated by using spectral reflectance related to the changes in the levels of canopy nitrogen and chlorophyll. We analyzed the relationships between  $P_{max}$  and  $\phi$ , derived from tower-based CO<sub>2</sub> flux observations, and ground-based spectral vegetation indices (VIs) in a temperate deciduous coniferous forest.

The canopy-scale  $P_{max}$  and  $\phi$  showed clear seasonal changes accompanying phenological stages. Both the variations in  $P_{max}$  and  $\phi$  were strongly correlated with VIs, especially with the ratio vegetation index (RVI) and enhanced vegetation index (EVI), independent of the growth stages. Moreover, day-to-day short-term variations of  $P_{max}$  and  $\phi$  were affected by meteorological conditions such as vapor pressure

deficit (VPD) and relative solar radiation which was calculated as the ratio of monitored radiation per theoretical maximum radiation.

Thus, seasonal changes of  $P_{max}$  and  $\phi$  were effectively assessed by RVI or EVI, and their short-term variations were evaluated by the empirical relationships with VPD and relative solar radiation. We propose a new simple method for estimating GPP with good precision; by fitting the light–response function with the evaluated parameters, the estimated GPP reflects 3 types of temporal variation: diurnal, day-to-day, and seasonal.

**Keywords:** GPP, light–response curve, maximum photosynthesis ( $P_{max}$ ), initial slope ( $\phi$ ), spectral vegetation index, vapor pressure deficit (VPD), relative solar radiation, eddy covariance

## 1. Introduction

Estimating the uptake of CO<sub>2</sub> by terrestrial ecosystems is an important consideration in the forecasting of global warming. The eddy covariance method is used worldwide to measure CO<sub>2</sub> flux as a precise means for continuous measurement of CO<sub>2</sub> exchange. The data have accumulated in recent years, and regional comparison studies are being performed (e.g., Falge et al., 2002; Saigusa et al., 2008). To estimate CO<sub>2</sub> exchange over wide areas, the CO<sub>2</sub> flux measurements from distributed points must be scaled up to spatially continuous estimates. Remote sensing technique and terrestrial ecosystem models should be useful for this purpose. Because satellite remote sensing allows simultaneous repetitive observations over wide areas, information can be obtained with extensive temporal and spatial coverage. The spectral data from optical remote sensing can be used to evaluate parameters in ecological models based on photosynthetic mechanisms. The spectral data should reflect physiological characteristics independent of vegetation types or growth stages.

One of the methods that utilize remote sensing data to estimate GPP is the light use efficiency (LUE) model (Monteith, 1972, 1977; Prince and Goward, 1995). The LUE model expresses GPP as the product of absorbed photosynthetically active radiation (APAR) and the LUE. APAR is calculated by multiplication of observed incident PAR by the fraction of absorbed PAR (FAPAR), which is estimated from vegetation indices (VIs) such as the normalized difference vegetation index (NDVI) or the enhanced vegetation index (EVI), obtained by remote sensing (e.g., Huete et al., 2002). However, the relationship between FAPAR and NDVI shows seasonal hysteresis, being affected by phenological stages (Jenkins et al., 2007). On the other hand, although the LUE, the ratio of GPP to APAR, is often estimated by using VIs such as photochemical reflectance index (PRI) (Nichol et al., 2006; Nakaji et al., 2008), there are some problems on the estimation of LUE from VIs in the whole time scales. The LUE is affected by light environments such as variations in radiation intensity and sky condition (Sims et al. 2005; Nakaji et al. 2007).

The phenological development of vegetation strongly affects GPP of deciduous forests (Saigusa et al., 2008). Temporal variations of photosynthetic capacity and efficiency at canopy scale have not been adequately assessed in many models, even though the evident seasonal changes of those at leaf scale are shown (Muraoka and Koizumi, 2005). Therefore, our objective in this study was to assess the variations of canopy photosynthetic capacity and efficiency using spectral information and a plant physiological approach, to allow a scale-up of tower-based CO<sub>2</sub> flux observation data to a CO<sub>2</sub> balance over wide areas. For this purpose, we focused on the light–response function at canopy scale and variations of its two main parameters, the maximum GPP under light saturation ( $P_{max}$ ) and the initial slope ( $\phi$ ) of light–response curve, instead of on LUE and FAPAR.

A number of physiological studies at foliage scale have pointed out that the maximum photosynthetic capacity ( $A_{max}$ ) is correlated with leaf nitrogen (N) level (Reich et al., 1995, 1999). Furthermore, the correlations between mass-based  $A_{max}$  and mass-based N can be described with one general equation independent of plant functional type or biome (Wright et al., 2004). Also, the initial slope of the light–response curve is related to chlorophyll concentration, which is associated with the electron transport rate. Remote sensing studies have proven that foliage N and chlorophyll content are evident in the spectral reflectance (e.g., Yoder and Pettigrew-Crosby, 1995). Therefore, we hypothesized that  $P_{max}$  and  $\phi$  could be evaluated by using spectral reflectance as related to changes in canopy N and chlorophyll levels.

However, due to the limit of temporal and spatial resolution of remote sensing, it is difficult to detect short-term responses of plants or small physiological changes, such as stomatal conductance. Then we analyzed the relationship between day-to-day short-term variations of  $P_{max}$  and  $\phi$  with daily meteorological conditions. We combined our results for both seasonal and short-term variations to estimate the canopy-scale  $P_{max}$  and  $\phi$ , and thereby demonstrate and validate a new method for estimating GPP by using spectral vegetation indices and the light–response function.

## 2. Background and hypothesis

### 2.1. Light–response curve, $P_{max}$ and $\phi$

$P_{max}$  and  $\phi$  at canopy scale, as at leaf scale, are important parameters of photosynthesis, and they can be calculated from CO<sub>2</sub> flux measurements (Ruimy et al., 1995). GPP is obtained as the sum of net ecosystem production (NEP) and ecosystem respiration (RE):

$$\text{GPP} = \text{NEP} + \text{RE} \quad (1)$$

Here, NEP is assumed equal to the negative of net ecosystem CO<sub>2</sub> exchange (NEE). GPP has been approximated as a function of PAR by the light–response curve. Curve-fitting of the light–response curve by using rectangular hyperbolic functions is generally used as an effective method to gap-fill for missing values in NEE (Falge et al., 2001). Also the non-rectangular hyperbolic equation (Thornley, 1976) is widely applied for gap-filling of NEE (Kosugi et al., 2005; Saigusa et al., 2008), and for predicting canopy photosynthesis (e.g., Saito et al., 2009). We compared these equations and selected the following non-rectangular hyperbolic equation to estimate GPP because of the best fit to GPP among them at this site (Appendix A):

$$\text{GPP} = \frac{\phi \cdot \text{PAR} + P_{max} - \sqrt{(\phi \cdot \text{PAR} + P_{max})^2 - 4\phi \cdot \text{PAR} \cdot \theta \cdot P_{max}}}{2\theta} + Rd \quad (2)$$

Here,  $P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) is the maximum GPP at light saturation,  $\phi$  ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ ) is the initial slope of the curve, and  $\theta$  represents the convexity of the curve.  $Rd$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) is daytime respiration.

### 2.2 Vegetation indices

We analyzed the relationships between photosynthetic parameters and several indices (Table 1). NDVI, EVI, and the ratio vegetation index (RVI) are indices for green leaf quantity based on

Table 1
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differences in reflectance between red and near-infrared bands. Chlorophyll has high absorbance in the visible red wavelength region (around 630–690 nm), whereas there is strong reflectance from cell walls and intracellular water of plants in the near-infrared region (around 740–900 nm). Thus, an increase in the quantity of healthy green leaves is accompanied by a decrease in red reflectance and a rise in near-infrared reflectance. RVI is a simple ratio of the reflectance of red and near-infrared bands.

NDVI is widely used for assessing parameters such as the vegetation cover ratio, leaf area index (LAI), and FAPAR (e.g., Sellers et al., 1992; Potter et al., 1993; Myneni et al., 1997); however, because it is a normalized index, it has the disadvantage of saturating at high biomass (Goward and Huemmrich, 1992). EVI corrects this disadvantage and can indicate the status of vegetation with mitigating effects of aerosols and soil by incorporating the reflectance for blue wavelengths (Huete et al., 2002). We anticipated that EVI and RVI could detect the canopy photosynthetic capacity as indices for N content.

The photochemical reflectance index (PRI) has been proposed as an optical indicator for detecting epoxidation and de-epoxidation changes of xanthophyll related to heat dissipation (Gamon et al., 1997). This index is calculated from the reflectance on both sides of the green absorbance band (531 and 570 nm). It also corresponds to the carotenoid:chlorophyll pigment ratio (Sims and Gamon, 2002). In this study, we expected that PRI might detect the “daytime depression” phenomenon during foliage period as well as autumn color changes (Nakaji et al., 2006).

The canopy chlorophyll index (CCI) has been proposed as an index for reflecting changes in chlorophyll content through the red edge shift (Sims et al., 2006). CCI is calculated by the ratio of the first derivatives of reflectance at wavelengths around the red edge (700 and 720 nm).

### **2.3 Summary of GPP estimation methods**

Various environmental (stress) factors affect the photosynthetic rate of larch canopy. Diurnal and daily changes in GPP are mainly controlled by irradiance, as expressed by the light–response curve. The photosynthetic capacity  $P_{max}$  and  $\phi$ , which regulate the light–response curve, have seasonal trends in relation to the seasonality in temperature and LAI. Spectral vegetation indices are expected to

reflect these seasonal trends in relation to the changes in canopy N level. Moreover, photosynthetic capacity would be affected by daily environmental conditions such as temperature, VPD, and soil water.

Therefore, our goal was to provide a precise estimation of GPP with a combination of 3 types of temporal resolution: (1) seasonal variation, by using VIs; (2) short-term variation, by using meteorological factors; and (3) diurnal variation, by using the light–response function (Fig. 1).

Fig. 1

First, we assessed the seasonality of the parameters using VIs in Process 1. From the relationships with the VIs ( $VI_1$  and  $VI_2$ , showing the best correlations with  $P_{max}$  and  $\phi$ , respectively), seasonal variations for daily-scale  $P_{max}$  ( $P_{max\_season}$ ) and  $\phi$  ( $\phi_{season}$ ) could be empirically described as:

$$P_{max\_season} = f_1(VI_1), \quad \phi_{season} = f_2(VI_2) \quad (3)$$

In addition to the seasonality of the parameters, we defined  $\Delta P_{max}$  and  $\Delta\phi$  as the day-to-day short-term variations in daily  $P_{max}$  and  $\phi$ . These were expressed as differences between the approximated original parameters ( $P_{max}$  and  $\phi$  in Eq. 2) and the estimated seasonal parameters ( $P_{max\_season}$  and  $\phi_{season}$  in Eq. 3) as follows:

$$\begin{aligned} \Delta P_{max} &= P_{max} - P_{max\_season}, \\ \Delta\phi &= \phi - \phi_{season} \end{aligned} \quad (4)$$

$\Delta P_{max}$  and  $\Delta\phi$  were expressed as empirical functions of meteorological factors in Process 2. Then, the estimated daily  $P_{max}$  and  $\phi$  ( $est\_P_{max}$ ,  $est\_phi$ ) were determined as the sums of seasonal and short-term variations of parameters.

Finally in Process 3, GPP was estimated by fitting the daily fluctuating parameters  $est\_P_{max}$  and  $est\_phi$ ,  $R_d$  and PAR to the light–response function (Eq. 2). When half-hourly PAR data are used, half-hourly GPP estimates are obtained.



### 3. Methods

#### 3.1 Study site

The study site was a temperate deciduous coniferous forest at Tomakomai Flux Research Site (lat. 42°44'N, long. 141°31'E, 125 m above sea level), in the northern part of Japan. About 81% of the approximately 100-ha site was dominated by a stand of Japanese larch (*Larix kaempferi*). The trees were about 45 years old, 15 m tall, with mean diameter at breast height (DBH) of 19 cm. Other trees on the site included deciduous trees such as birch (*Betula ermanii*, *B. platyphylla*) and Japanese elm (*Ulmus japonica*) mixed with evergreen conifer species such as spruce (*Picea jezoensis*). The forest floor was thickly covered with Japanese spurge (*Pachysandra terminalis*), ferns (*Dryopteris crassirhizoma*), and other species.

The maximum LAI for the canopy was  $5.6 \text{ m}^2 \text{ m}^{-2}$  and for the lower-layer vegetation,  $3.6 \text{ m}^2 \text{ m}^{-2}$ , for a total maximum of  $9.2 \text{ m}^2 \text{ m}^{-2}$ , as determined by a leaf collection method. The seasonal change of LAI was estimated from Beer's law using an extinction coefficient of 0.58 and a plant area index of  $1.4 \text{ m}^2 \text{ m}^{-2}$  (Hirata et al., 2007). The soil was a highly permeable oligotrophic Regosol of volcanic origin (Fujinuma et al., 2001). The mean air temperature was 6.2 °C and mean annual precipitation was 1043 mm during 2001–2003.

The flux tower and forest at the site were completely destroyed by a typhoon on 8 September 2004.

#### 3.2 Measurement of CO<sub>2</sub> flux and micrometeorological parameters

CO<sub>2</sub> flux ( $F_c$ ) was measured using the eddy covariance technique with a closed-path infrared gas analyzer (Li6262; Li-Cor, Lincoln, Nebraska, USA); wind velocity and air temperature were measured simultaneously using a three-dimensional sonic anemometer-thermometer (DA-600-3TV; Kaijo, Tokyo, Japan) at 27 m on the tower. The sampled air was drawn from an inlet at 27-m height and pumped into the gas analyzer. Raw data were recorded at 10 Hz using a data logger (DR-M3; TEAC, Tokyo, Japan). CO<sub>2</sub> measurements were calibrated once a day by flowing two CO<sub>2</sub> standard

gases of 320 and 420 ppmv (Hirano et al., 2003). After removal of noise spikes, corrections were applied for three-dimensional coordinate rotation, for time lag, and for air density fluctuations, and then half-hourly mean values for  $F_c$  were calculated. In addition, the rate of change in CO<sub>2</sub> storage ( $F_s$ ) below the  $F_c$  measurement height was obtained from the vertical profile of CO<sub>2</sub> concentrations at eight levels.

Net ecosystem exchange (NEE) was calculated as the sum of  $F_c$  and  $F_s$ :

$$NEE = F_c + F_s \quad (5)$$

Meteorological factors such as global solar radiation, PAR, air temperature, precipitation on the flux tower, soil temperature, and soil water content were measured. As we did not monitor diffuse solar radiation, for this study we substituted the relative solar radiation (rSR). We defined rSR as the ratio of observed irradiance to theoretical irradiance of full sunlight, which was estimated from the solar zenith angle at the time, atmospheric turbidity, albedo, and precipitable water vapor (Kondo and Miura, 1983). We distinguished clear sky and cloudy conditions in this study by values of  $rSR \geq 75\%$  or  $rSR < 75\%$ , respectively.

### 3.3 Calculation of photosynthetic parameters

The data from January 2003 to August 2004 were available for analysis. When the friction velocity ( $u^*$ ) was less than  $0.3 \text{ m s}^{-1}$ , nighttime  $F_c$  tended to be underestimated at this site (Hirata et al., 2007), so under these conditions the values for NEE were excluded. NEE during nighttime and winter snowfall periods was assumed to be equal to RE. The observed RE was regressed using the equation of Lloyd and Taylor (1994):

$$RE = RE_{ref} \exp \left\{ \frac{E_0}{R} \left( \frac{1}{T_k + T_{ref} - T_0} - \frac{1}{T_k + T_s - T_0} \right) \right\} \quad (6)$$

Here,  $RE_{ref}$  is the ecosystem respiration rate ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) at the reference temperature (10 °C),  $E_0$  is activation energy ( $\text{J mol}^{-1}$ ),  $R$  is the gas constant ( $R = 8.314 \text{ J mol}^{-1} \text{ K}^{-1}$ ),  $T_s$  is the soil temperature (°C, 0.05 m depth), and  $T_0$  and  $T_k$  are constants with  $T_0 = 273.15 \text{ K}$  and  $T_k = 227.13 \text{ K}$ . The two regression parameters,  $RE_{ref}$  and  $E_0$ , were determined each day with a 91-day moving window, considered to reflect the seasonality of RE. Gaps in nighttime RE were filled and daytime RE was estimated from the soil temperature by fitting Eq. 6. GPP was then calculated by using Eq. 1.

GPP was fitted to the light–response curve (Eq. 2) with a 3-day (2 previous days included) moving window throughout a year, and the daily parameters  $P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ),  $\phi$  ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ ), and  $Rd$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) were determined. Here, the convexity ( $\theta$ ) of the light–response curve was set at 0.9, following previous studies (Kosugi et al., 2005; Saigusa et al., 2008; Saito et al., 2009). The Gauss-Newton least-squares method was used in the non-linear regression. Although a regression period of approximately 10 days is recommended for use in flux gap-filling (Falge et al., 2001; Baldocchi and Wilson, 2001), we used a moving window of 3 days to investigate the short-term effects of weather conditions. Because weather conditions showed a 3- or 4-day cycle variation especially in spring in Japan due to migratory High (e.g., Nemoto et al., 1982). The window size of 3 days was sufficiently short to remove seasonal effects and was long enough to provide sufficient data points for stable regression analysis.

Parameter values were excluded from analysis when the number of valid data points were too few for regression analysis (<20 for a 3-d period), when convergence was impossible (the coefficient of determination [ $r^2$ ] was less than 0.1), when the parameters were out of range ( $P_{max} \geq 50.0$  or  $\leq 0.1 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ;  $\phi \geq 1.0$  or  $\leq 0.001 \text{ mol CO}_2 [\text{mol photon}]^{-1}$ ), or when PAR was too low to reach light saturation.

### 3.4 Spectral reflectance and vegetation indices

Upward and downward spectral radiation was observed with two grating spectroradiometers (MS-700; EKO Instruments, Tokyo, Japan) mounted on the tower at 40-m height. Spectral radiation for 256 wavelength bands ranging from 305 to 1150 nm at about 3.3-nm intervals was measured every

minute, and the spectral reflectance of larch canopy was derived from the upward radiation divided by the downward radiation. Both spectroradiometers were simultaneously calibrated once a year for the radiation level of each band.

Each VI was calculated from the averaged spectral reflectance at three bands around each wavelength in Table 1. We used the data for half-hourly mean values around noon as the representative values for each day. If there was rain within the preceding 12 hours, the data were excluded from the analysis. Assuming the proposed application of this method for satellite remote sensing, we only used the data from under clear sky conditions ( $rSR > 75\%$ , see Section 3.2).

### **3.5 Data set**

In this study, we used the data of CO<sub>2</sub> flux, meteorological parameters and spectral reflectance which were collected from January 2003 to August 2004, just before the destruction of the tower. In order to analyze the seasonal variation of the photosynthetic parameters and VIs, the data from the autumn of 2003 were added as a continuation after August 2004. We selected 50% of the days at random for "training data" to construct the model, and used the rest of the data set as "test data" to examine the model validity.

### **3.6 Monitoring of phenology**

To compare the seasonal changes in  $P_{max}$  and  $\phi$  with the actual phenology of the vegetation, we visually assessed the phenology by using photographic images from digital cameras mounted on the flux tower.

The Japanese larch has two types of shoots: short-shoots and long-shoots. First, buds of the short-shoots break and grow, developing their leaves on the old branches of the previous year, which here we call leaf "flush". Subsequently, the long-shoots grow from the tips of the previous year's branches.

## 4. Results and discussion

### 4.1 Seasonal variations in $P_{max}$ and $\phi$

Clear seasonal variations of canopy  $P_{max}$  and  $\phi$  were observed related to changes in the vegetation phenology (Fig. 2). The flush of leaves in the larch forest started at the end of April, about 2 weeks after the disappearance of snow, and short-shoot leaves had developed for about 3 weeks. Both  $P_{max}$  and  $\phi$  increased along with the flush of leaves.

Fig. 2

From mid-June to the beginning of August, long-shoots elongated.  $P_{max}$  and  $\phi$  continued to increase and reached their maximum values at the end of June, during the elongation period.  $P_{max}$  and  $\phi$  gradually decreased from August with large daily fluctuations, while temperature and LAI remained at their maximum levels.

The leaf color change started in mid-October and reached peak coloration at the end of October. The leaves had fallen within the first 10 days of November.  $P_{max}$  and  $\phi$  continued to decline, reaching minimum values during the period of leaf color change.

### 4.2 Seasonal variations of vegetation indices (VIs)

Each VI had a different seasonal pattern, with different times of increase, maximum, and decrease (Fig. 3). At the time of snow disappearance, NDVI showed a jump up from its minimum level, and EVI and RVI rose a little, on the contrary PRI dropped down. NDVI sensitively reflected the condition of the ground surface. NDVI, EVI, and RVI increased linearly over 1 month just after the start of flush, whereas PRI and CCI increased more gradually over 2 months. Although NDVI maintained maximum values over 2 months during summer, EVI and RVI had shorter periods of maximum values, from mid-June to July. In early July of 2004, after a few days of heavy rain and low temperatures, LAI decreased even though it was still the leaf extension period (Fig. 2). VIs except for NDVI detected this

Fig. 3

change and decreased at the same time (Fig. 3). NDVI became saturated during summer and did not follow the decrease of LAI.

PRI and CCI peaked in late summer in 2003. Whereas EVI and RVI declined gradually from late summer, PRI declined later, in early autumn. PRI would be affected by the decrease of chlorophyll and the increase of carotenoid pigments (Sims and Gamon, 2002) and is therefore more sensitive to senescence than other VIs. PRI had greater daily fluctuation than the other VIs. In early June,  $\phi$  decreased during a few sunny days, and the drop was followed only by PRI. The strong solar radiation probably caused de-epoxidation of xanthophyll pigments for photoprotection, which was possibly detected by PRI (Gamon et al., 1992; Nakaji et al., 2006). PRI showed the unique pattern of declining in autumn far below the spring baseline and rising again during winter.

### 4.3 Relationships of $P_{max}$ and $\phi$ to VIs

$P_{max}$  and  $\phi$  were significantly correlated with VIs, LAI and temperature, and the correlations with VIs were higher than with LAI or temperature (Table 2).  $P_{max}$  was strongly correlated with EVI and RVI ( $r = 0.92$  for both). NDVI and  $P_{max}$  had an exponential relationship, because NDVI saturated during the period of maximum  $P_{max}$ . PRI showed a lower correlation with  $P_{max}$  because of its different pattern of decline in autumn.

Table 2
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$\phi$  had highest correlation with RVI ( $r = 0.81$ ) and next highest with EVI ( $r = 0.79$ ), and the correlation between  $P_{max}$  and  $\phi$  was relatively high ( $r = 0.74$ , data not shown). NDVI, PRI, and CCI had lower correlations with  $\phi$ , owing to different peak periods and patterns of decline. Variations of both  $\phi$  and PRI coincided well only during the period from spring to summer.

Seasonal variations of  $P_{max}$  and  $\phi$  at canopy scale have also been readily observed in other temperate mixed forest sites, whereas they are not clearly observed in subtropical evergreen coniferous forests or evergreen broadleaf forests (Hollinger et al., 1999; Zhang et al., 2006). These studies suggested that the changes in temperature or LAI mainly controlled the seasonal changes of canopy  $P_{max}$  and  $\phi$ . However, at our site we found that the correlations between LAI or temperature

and  $P_{max}$  and  $\phi$  were not higher than those between VIs and  $P_{max}$  or  $\phi$ . The ratio of LAI to  $P_{max}$  was different depending on the growth stage of the vegetation. It was higher in autumn than in spring, because leaves that had changed color, with reduced photosynthetic capacity, were still present on the plants in autumn.

Smith et al. (2002) observed a linear relationship between forest productivity and N concentration of canopy leaves independent of vegetation type. Furthermore, canopy N concentration correlates with the absorption rate for chlorophyll in the red region (around 680 nm) and blue region (around 490 nm), and the reflectance rate in the near-infrared region. This indicates the utility of VIs based on these bands, such as EVI, for estimating canopy N content and forest productivity.

Inoue et al. (2008) investigated the relationship between seasonal changes in  $P_{max}$ ,  $\phi$  and various spectral indices at an irrigated rice field. They report that  $P_{max}$  is correlated with RVI and other indices related to the wavelengths in chlorophyll a and b absorption region, and for  $\phi$ , related to the wavelengths in blue region (450 nm) and infrared region (1330 nm).

Therefore, our results, together with those of others, identified RVI and/or EVI as available indices to assess photosynthetic parameters more directly, significantly, and effectively than LAI or air temperature, because they reflect the changes in canopy N or chlorophyll content. Either RVI or EVI can be used for estimating the seasonal variation of  $P_{max}$ , and here we selected RVI. The same analysis was conducted for EVI (Appendix B). We obtained the following empirical equations for the relationships between  $P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) and RVI, and  $\phi$  ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ ) and RVI, from the training data (Fig. 4):

$$P_{max} = 3.65 \times \text{RVI} - 3.35 \quad (7)$$

$$\phi = 0.005 \times \text{RVI} - 0.001 \quad (8)$$

Fig. 4

#### 4.4 Short-term variations of $P_{max}$ and $\phi$ , and the effects of meteorological factors

$P_{max}$  and  $\phi$  had large short-term variations ( $\Delta P_{max}$  and  $\Delta\phi$ ) during photosynthetically active periods (herein,  $P_{max} > 20 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ , and  $\phi > 0.03 \text{ mol CO}_2 [\text{mol photon}]^{-1}$ ). We analyzed the effects of four meteorological parameters on  $\Delta P_{max}$  and  $\Delta\phi$  for 3-day periods: daily temperature anomaly ( $\Delta T$ , °C), relative solar radiation (rSR, %), maximum vapor pressure deficit (VPD, kPa), and soil water content (SWC, %). We used  $\Delta T$  instead of daily mean temperature to remove the seasonal trend. Here, we defined  $\Delta T$  as the difference between daily mean temperature and its 14-day running mean. We substituted rSR for the ratio of direct to diffuse radiation (Table 3).

Table 3

We found a significant negative correlation between  $\Delta P_{max}$  and rSR ( $r = -0.49$ ,  $P < 0.0001$ ) and with VPD ( $r = -0.77$ ,  $P < 0.0001$ ; Fig. 5a), and no significant relationship with either  $\Delta T$  or SWC. When VPD was higher than about 1 kPa in summer, many examples of ‘daytime depression’ were observed at this site (Wang et al., 2004), because the higher VPD caused lower stomatal conductance and lower  $P_{max}$ . When both VPD and rSR were high on sunny days, photosynthesis became saturated in the upper layer of the canopy. Meanwhile, when VPD was low and rSR was low,  $\Delta P_{max}$  was high. This suggested that on cloudy days diffuse radiation reached more deeply into the canopy and the photosynthetic capacity of the entire canopy was enhanced more than on sunny days. Our results agree with the diffuse enhancement reported by Gu et al. (2002). Both VPD and rSR had strong synergetic effects on  $\Delta P_{max}$ . As VPD and rSR are correlated with each other,  $\Delta P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) can be described by the effects of only VPD (kPa) (Fig. 5a) as:

Fig. 5

$$\Delta P_{max} = -11.54 \times \text{VPD} + 18.69 \quad (9)$$

According to previous studies, for European beech forests,  $\text{CO}_2$  uptake is high at VPD of 0.5-1 kPa, and  $\text{CO}_2$  uptake is reduced when VPD becomes higher (Herbst et al., 2002). For Chinese evergreen coniferous forests,  $P_{max}$  reaches its maximum value when VPD is about 1.26 kPa and rapidly declines when VPD exceeds 2 kPa (Zhang et al., 2006). The optimum and threshold values of VPD probably differ for each vegetation type or canopy structure.



For  $\Delta\phi$ , significant correlation was found only with rSR ( $r = -0.42$ ,  $P < 0.0001$ ; Fig. 5b), and less significant correlations were shown with  $\Delta T$  and VPD (Table 3). After several consecutive sunny days,  $\Delta\phi$  typically dropped suddenly. Exposure under strong direct solar radiation for few days possibly inhibited the canopy photosynthesis. In contrast,  $\phi$  was enhanced when rSR was low, that is, when the proportion of diffuse radiation was high, similar to previous observations (Hollinger et al., 1994; Gu et al., 2002).

Thus, we verified that both direct irradiance and diffuse radiation affected  $\Delta\phi$ . We obtained the relationship between  $\Delta\phi$  ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ ) and rSR (%) (Fig. 5b) as:

$$\Delta\phi = -0.0002 \times \text{rSR} + 0.0138 \quad (10)$$

Although we expected that soil water content would have some effect on  $\Delta\phi$ , no significant correlation was observed. This may be because there was no severe drought stress during growing periods at this site during this study.

#### 4.5 Estimation of $P_{max}$ , $\phi$ and GPP

To determine  $P_{max}$  and  $\phi$  using our new approach, we first estimated the seasonal variations ( $P_{max\_season}$  and  $\phi_{season}$ ) from their respective relationships with RVI (Eqs. (7) and (8)) on sunny days, using linear interpolation for other days.

Next,  $\Delta P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) and  $\Delta\phi$  ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ ) were evaluated from VPD (kPa) and rSR (%) respectively (Eqs. (9) and (10)). Daily estimated  $P_{max}$  ( $\text{est\_}P_{max}$  [ $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ]) was then determined as the sum of  $P_{max\_season}$  and  $\Delta P_{max}$ . The same process was used to determine  $\text{est\_}\phi$ .

Here, we define  $P_{max}$  and  $\phi$  as previously approximated by fitting the light–response curve (Eq. (2)) as the "original"  $P_{max}$  and  $\phi$ .  $P_{max\_season}$  as determined from RVI was an underestimate compared to the original  $P_{max}$  (slope = 0.72 for  $P_{max\_season}$  vs. original  $P_{max}$ ). After applying the correction ( $\Delta P_{max}$ ),  $\text{est\_}P_{max}$  was closer to the original (slope = 0.95 for  $\text{est\_}P_{max}$  vs. original  $P_{max}$ ; Table 4). Similarly,

Table 4
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$\phi_{\text{season}}$  was underestimated by RVI compared with the original  $\phi$ , and  $\text{est\_}\phi$  was an improvement after applying the  $\Delta\phi$  correction (slope of comparison increased from 0.66 before the correction to 0.78 after; Table 4). Thus, the effectiveness of the meteorological correction for both  $\Delta P_{\text{max}}$  and  $\Delta\phi$  was confirmed.

Finally, GPP was estimated by fitting  $\text{est\_}P_{\text{max}}$  and  $\text{est\_}\phi$  to the light–response function at half-hour time steps. Diurnal changes of GPP were well expressed in relation to the light intensity (Fig. 6).

Fig. 6

GPP was precisely estimated from the parameters compared with observed GPP at half-hour time steps during the growing season in 2003-2004 (Table 5, Fig. 7). The slope of linear regression between observations and estimated GPP was 0.96 ( $R^2 = 0.83$ ), and the standard error (SE) for the estimation was  $4.21 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ . Also when the parameters were estimated using EVI, GPP was well estimated (slope = 0.95,  $R^2 = 0.84$ , SE =  $4.23 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ), almost at the same precision as that by using RVI.

Table 5

Fig. 7

The GPP approximated by using the original parameters had a slope of 0.94 ( $R^2 = 0.87$ ) when regressed against observed GPP; the SE was  $3.67 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ . This is the implied regression error from using the light–response function. By our method, GPP tended to be about 4% of underestimate, and 83% of the half-hourly variations in observed GPP were explained by our model.

## 5. Conclusions

We propose a method for estimating GPP based on the light–response function that links remote sensing techniques and  $\text{CO}_2$  flux measurement, from a plant physiological standpoint. With this method we parameterized  $P_{\text{max}}$  and  $\phi$  as photosynthetic indicators consisting of both seasonal and short-term fluctuating components. The seasonal variations of  $P_{\text{max}}$  and  $\phi$  had higher linear correlations with those of VIs such as RVI and EVI than with air temperature or LAI. Our results indicated that  $P_{\text{max}}$  and  $\phi$  could be directly and adequately determined using either RVI or EVI. The

seasonality of  $P_{max}$  and  $\phi$  seemed to correspond to the changing phenology of the vegetation, presumably reflecting canopy N or chlorophyll content.

In addition to seasonal patterns, we found that the short-term variations of  $P_{max}$  and  $\phi$  were affected by VPD and the relative solar radiation, and we suggest the necessity of incorporating both seasonal and short-term variations into the estimation of  $P_{max}$  and  $\phi$ . Assuming the use of satellite data, which are obtained under clear sky conditions,  $P_{max}$ ,  $\phi$ , and GPP would be underestimated without corrections for the effects of meteorological conditions.

Large numbers of parameters in models often cause uncertainty in estimated results, and ecosystem modeling has been in need of effective parameterization. We believe that the ability to generalize the relationships between  $P_{max}$ ,  $\phi$ , and RVI and/or EVI will contribute to ecosystem modeling. Both RVI and EVI showed strong correlations with  $P_{max}$ , and similarly with  $\phi$ , and the GPP estimates from the parameters assessed by RVI or EVI were not very different. Further investigations involving different species and vegetation types should help to determine which VI is better to use for parameterization.

In this study, by using parameters of the light–response function that fluctuated daily, we demonstrated an improvement in providing accurate GPP estimates with high temporal resolution in response to diurnal, short-term, and seasonal changes. To further improve this method, we will also need additional studies of meteorological effects on the short-term variations in  $P_{max}$  and  $\phi$  across various biomes and under severe conditions.

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This study referred to some results from the Ministry of Environment Comprehensive Promotion Program for Global Environmental Research S1: “Research to develop a comprehensive carbon balance for Asian terrestrial ecosystems in preparation for 21<sup>st</sup> Century carbon management” (Principal Investigator: Takehisa Oikawa).

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ZeGraph (<http://www.zegraph.com/>) was used for non-linear regression calculations, and we deeply appreciate the guidance and cooperation provided by Jiye Zeng.

## Appendix A

We compared the RMSE from following three regression equations with the observed GPP:

1. *Michaelis-Menten's* rectangular hyperbolic equation:

$$\text{GPP} = \frac{\phi \cdot \text{PAR} \cdot P_{max}}{(\phi \cdot \text{PAR} + P_{max})} + Rd \quad (\text{A.1})$$

2. *Misterlich's* rectangular hyperbola:

$$\text{GPP} = P_{max} \left( 1 - e^{-(\phi \cdot \text{PAR} / P_{max})} \right) + Rd \quad (\text{A.2})$$

Eq. A.1 and A.2 are used by Falge et al. (2001).

3. The non-rectangular hyperbola (Thornley, 1976; Eq. (2)) with  $\theta$  fixed at 0.9.

The data set used and the method of calculation were the same as those mentioned in Section 3.3.

The averaged RMSE for each equation was, respectively, 3.16, 2.87, and 2.84  $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ . Therefore, we chose the non-rectangular hyperbola, which had the smallest regression errors. With the *Michaelis-Menten* equation, values of regression coefficients often became unrealistic. The parameters of the *Misterlich* equation were highly correlated and comparable to those of the non-rectangular hyperbola. Regressions of parameters of the *Misterlich* equation and the non-rectangular hyperbola yielded slopes and intercepts of 1.06 and  $-0.36$  ( $R^2 = 0.99$ ), respectively, for  $P_{max}$  and 1.32 and 0.0 ( $R^2 = 0.95$ ), respectively, for  $\phi$ .

## Appendix B

We obtained the following empirical equations for the relationships between  $P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) and EVI, from the training data:

$$P_{max} = 70.50 \times \text{EVI} - 9.03 \quad (\text{B.1})$$

We found a significant negative correlation between  $\Delta P_{max}$  (estimated by using EVI) and rSR ( $r = -0.52$ ,  $P < 0.0001$ ) and with VPD ( $r = -0.73$ ,  $P < 0.0001$ ), and no significant relationship with either  $\Delta T$  ( $r = -0.18$ ) or SWC ( $r = -0.03$ ).  $\Delta P_{max}$  ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) can be described by the effects of only VPD (kPa) as:

$$\Delta P_{max} = -11.00 \times \text{VPD} + 19.28 \quad (\text{B.2})$$

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Table 1. Vegetation indices (VIs) used in this study.

$R_\lambda$  indicates spectral reflectance at wavelengths around  $\lambda$  nm.  $R_{\text{nir}} = R_{857}$ ,  $R_{\text{red}} = R_{647}$ ,  $R_{\text{blue}} = R_{464}$ .  $D_\lambda$  is the first derivative of reflectance at wavelengths around  $\lambda$  nm.

Table 2. Correlation coefficients ( $r$ ) between  $P_{\text{max}}$ ,  $\phi$  and VIs, LAI and temperature under clear sky conditions ( $n = 88$ ). \*\*\* $p < 0.0001$  (Pearson's correlation test).

Table 3. Relationships between  $\Delta P_{\text{max}}$ ,  $\Delta \phi$  and four meteorological factors:  $\Delta T$  (temperature anomaly, °C), rSR (relative solar radiation, %), VPD (max vapor pressure deficit for 3-day periods, kPa), and SWC (soil water content, %). \*\*\* $p < 0.0001$ , \*\* $p < 0.001$ , \* $p < 0.01$ ; ns: not significant (Pearson's correlation test).

Table 4. Relationship between estimated parameters and the original parameters (see text). Seasonal parameters ( $P_{\text{max\_season}}$ ,  $\phi_{\text{season}}$ ) were estimated by using RVI and corrected (est\_ $P_{\text{max}}$ , est\_ $\phi$ ) using meteorological factors. ( $n = 171$  for  $P_{\text{max}}$ , and  $n = 176$  for  $\phi$ ).

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Fig. 6. Diurnal changes of estimated GPP and observed GPP at half-hourly time steps (from day of year 229-234 in 2004).

Fig. 7. Relationship between observed GPP and GPP estimated using parameters derived from RVI at half-hourly time steps (for test data;  $n = 4025$ ). Solid line is linear regression line; dashed line indicates  $y = x$ .

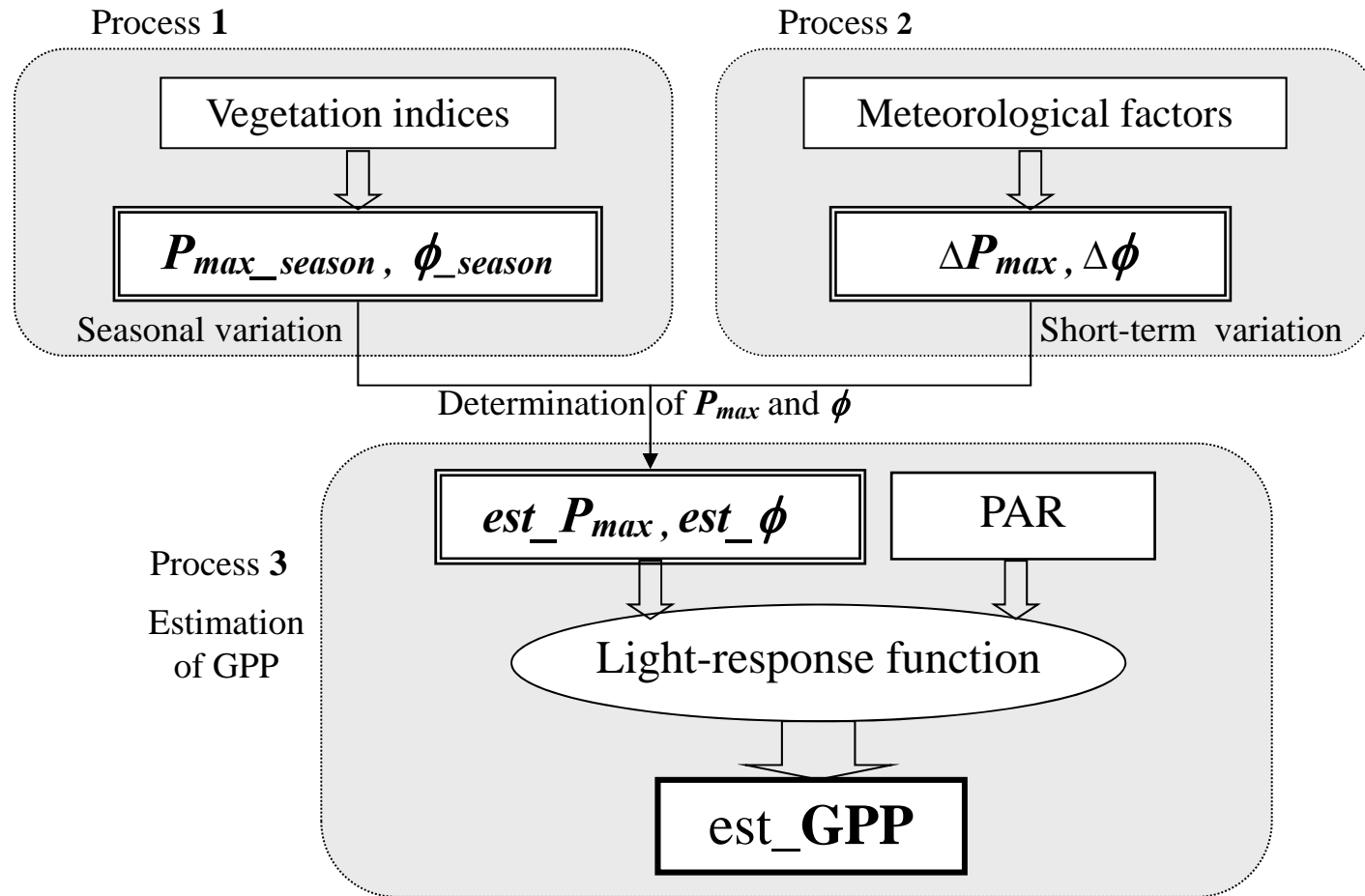


Fig. 1. Concept behind the determination of  $P_{max}$  and  $\phi$  and the estimation of GPP.  : input data,  : parameters.

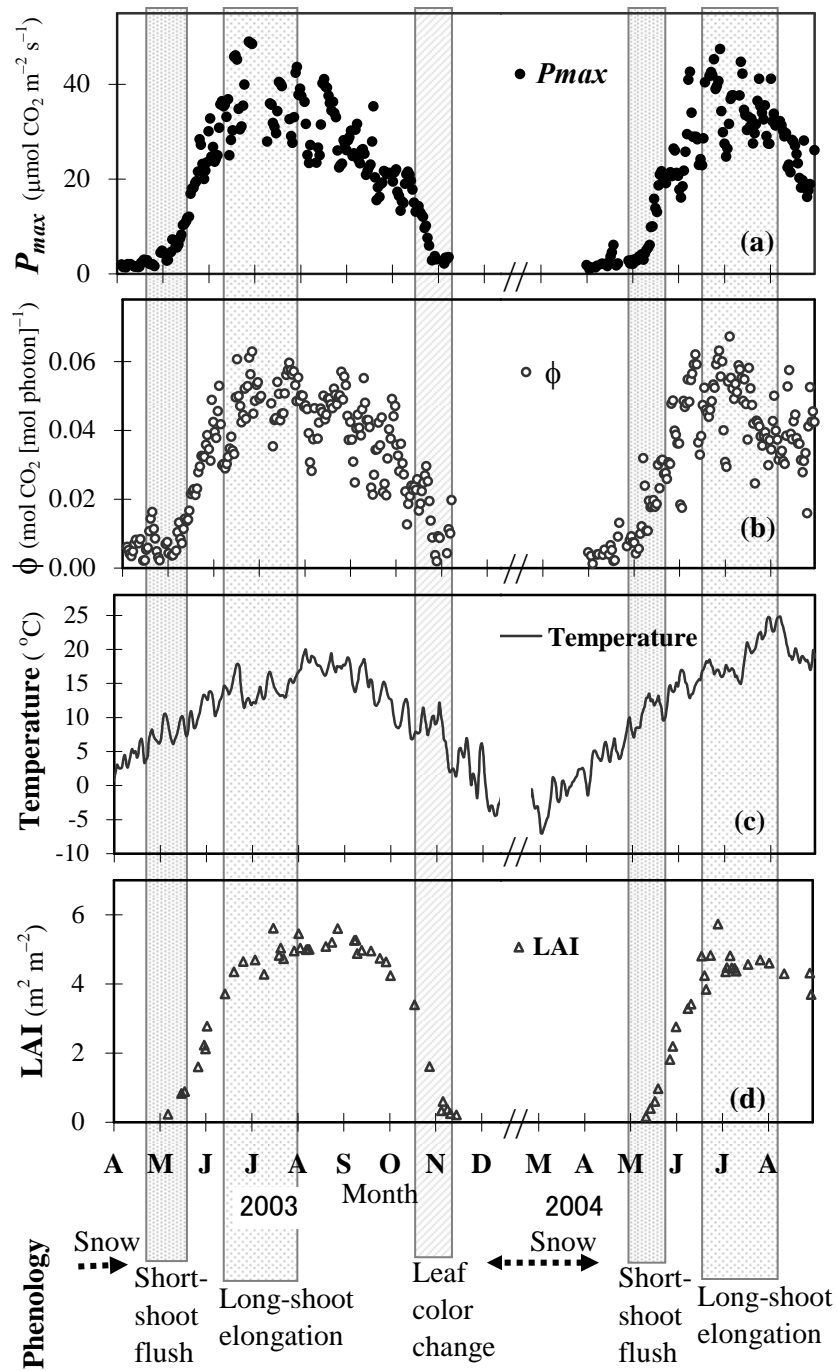


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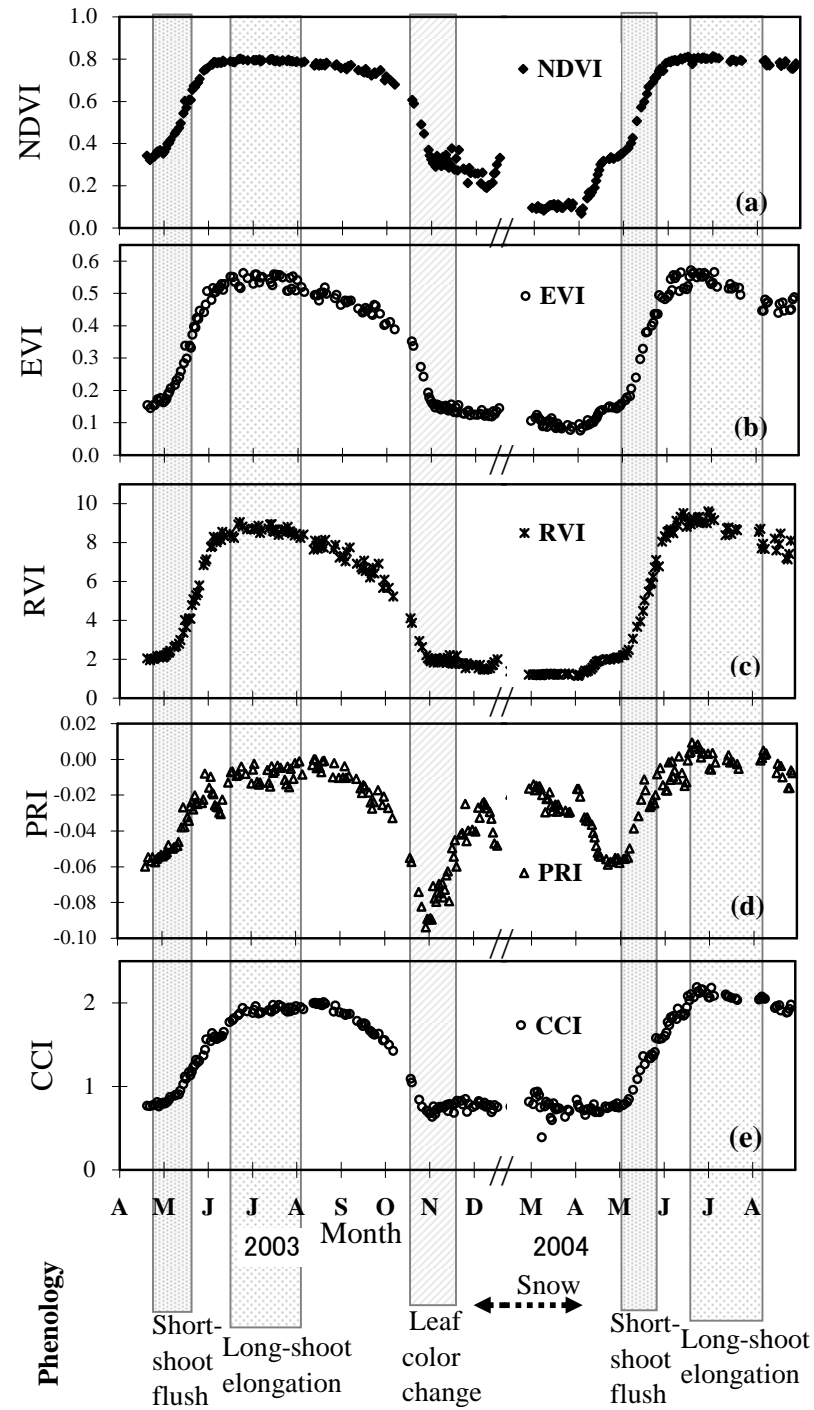


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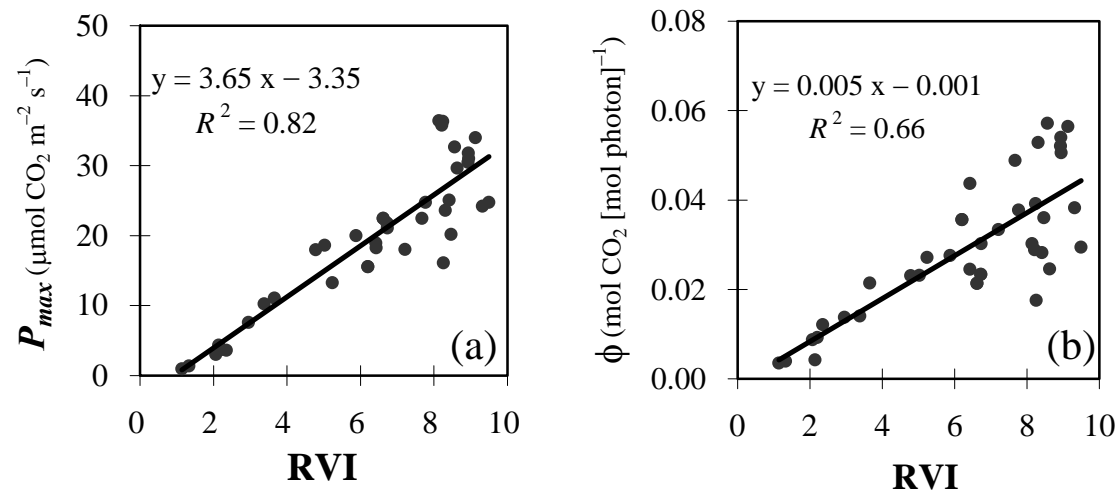


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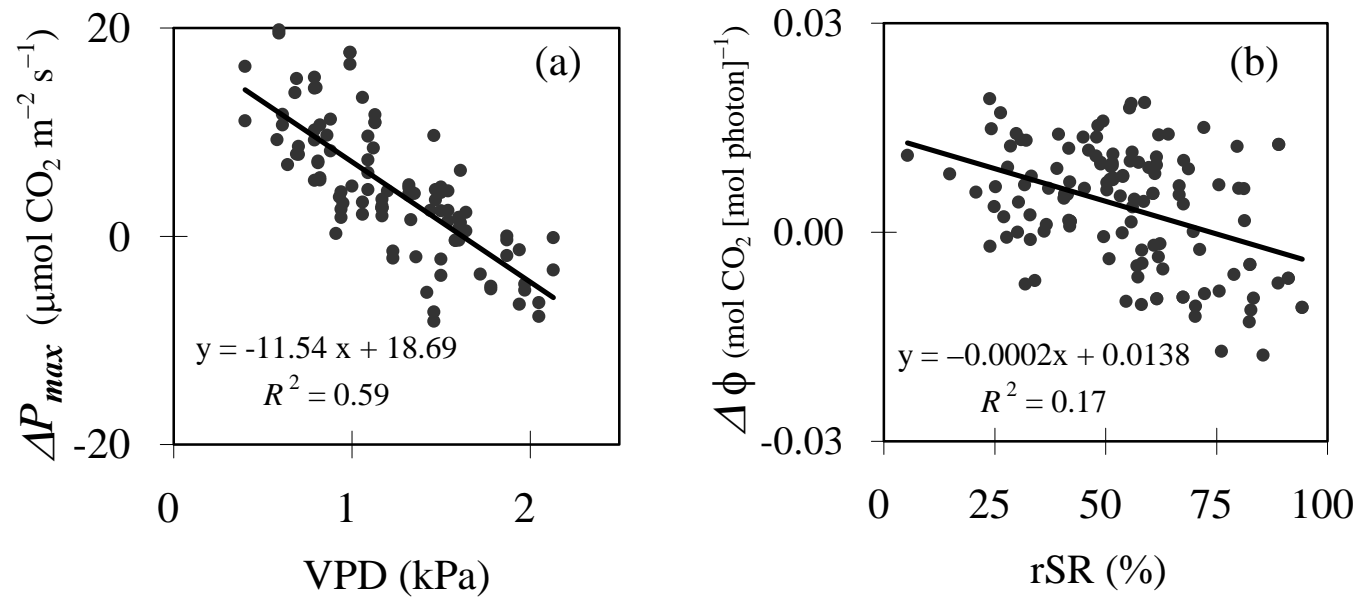


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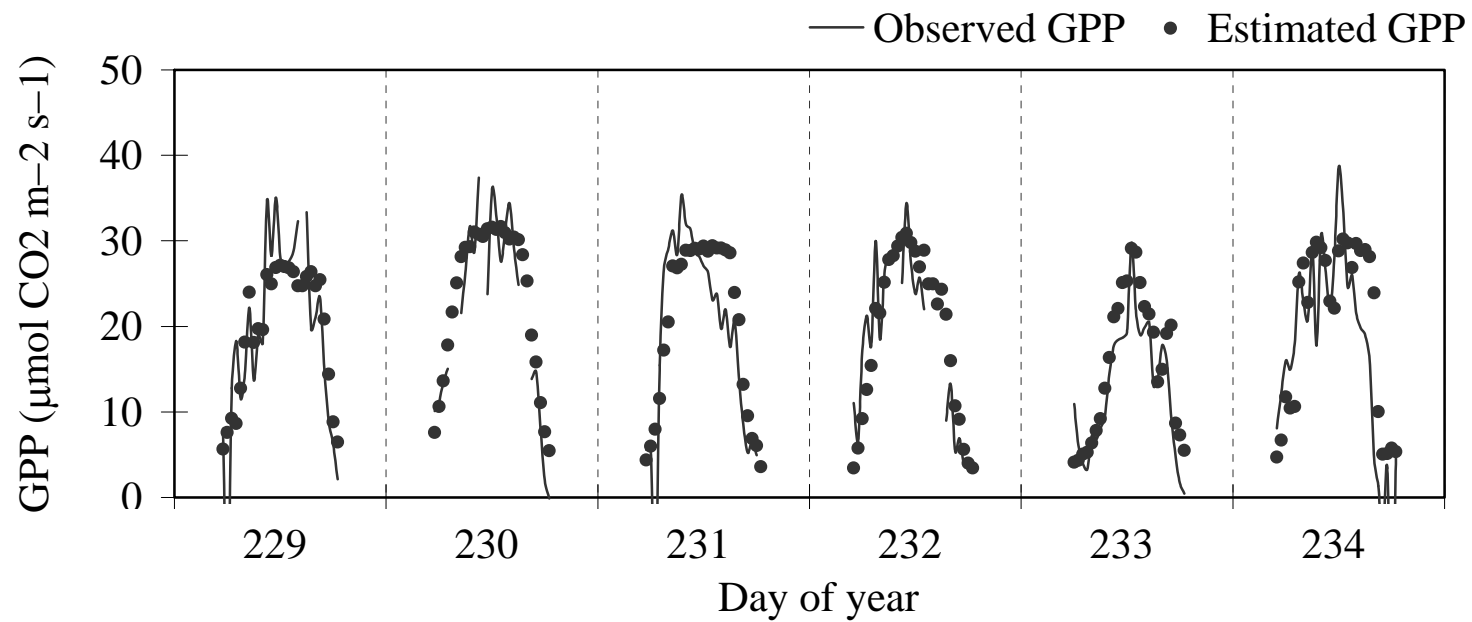


Fig. 6. Diurnal changes of estimated GPP and observed GPP at half-hourly time steps (from day of year 229-234 in 2004).

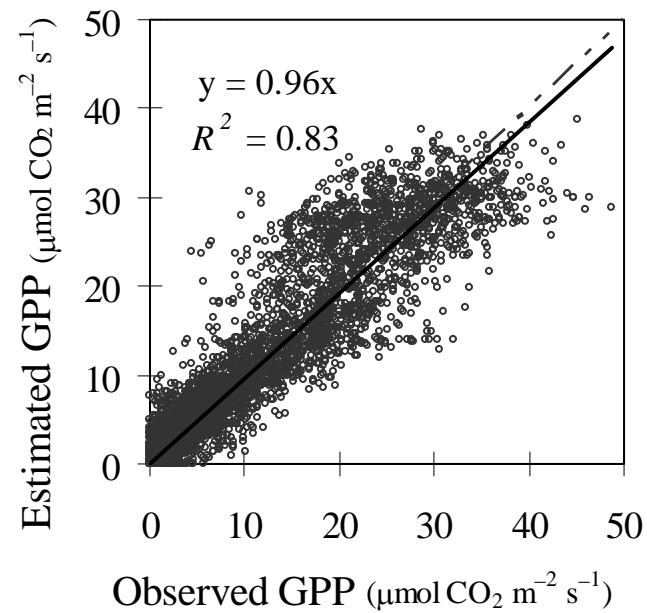


Fig. 7. Relationship between observed GPP and GPP estimated using parameters derived from RVI at half-hourly time steps (for test data;  $n = 4025$ ). Solid line is regression line, and dashed line indicates  $y = x$ .

Table 1. Vegetation indices (VIs) used in this study.

VI	Formulation	Reference
NDVI (normalized difference vegetation index)	$(R_{nir} - R_{red}) / (R_{nir} + R_{red})$	Tucker (1979)
EVI (enhanced vegetation index)	$2.5[(R_{nir} - R_{red}) / (R_{nir} + 6.1R_{red} - 7.5R_{blue} + 1.0)]$	Huete et al. (2002)
RVI (ratio vegetation index, simple ratio)	$R_{nir} / R_{red}$	Rouse et al. (1974)
PRI (photochemical reflectance index)	$(R_{531} - R_{570}) / (R_{531} + R_{570})$	Gamon et al. (1997)
CCI (canopy chlorophyll index)	$D_{720} / D_{700}$	Sims et al. (2006)

$R_{\lambda}$  indicates spectral reflectance at wavelength around  $\lambda$  nm.  $R_{nir} = R_{857}$ ,  $R_{red} = R_{647}$ ,  $R_{blue} = R_{464}$ .

$D_{\lambda}$  is the first derivative of reflectance at wavelengths around  $\lambda$  nm.

Table 2. Correlation coefficients ( $r$ ) between  $P_{max}$ ,  $\phi$  and VIs, LAI and temperature under clear sky conditions ( $n = 88$ ).

Parameter	$P_{max}$	$\phi$
VI		
NDVI	0.88 <sup>***</sup>	0.78 <sup>***</sup>
EVI	0.92 <sup>***</sup>	0.79 <sup>***</sup>
RVI	0.92 <sup>***</sup>	0.81 <sup>***</sup>
PRI	0.74 <sup>***</sup>	0.63 <sup>***</sup>
CCI	0.88 <sup>***</sup>	0.77 <sup>***</sup>
LAI	0.67 <sup>***</sup>	0.60 <sup>***</sup>
Temperature	0.75 <sup>***</sup>	0.59 <sup>***</sup>

\*\*\*  $p < 0.0001$  (Pearson's correlation test).

Table 3. Relationships between  $\Delta P_{max}$ ,  $\Delta\phi$  and four meteorological factors:  $\Delta T$  (temperature anomaly, °C), rSR (relative solar radiation, %), VPD (max vapor pressure deficit for 3-day periods, kPa), and SWC (soil water content, %).

Meteorological factor	$\Delta P_{max}$ ( $n = 105$ )	$\Delta\phi$ ( $n = 126$ )
$\Delta T$	-0.16 <i>ns</i>	-0.24 *
rSR	-0.49 ***	-0.42 ***
VPD	-0.77 ***	-0.25 *
SWC	0.10 <i>ns</i>	-0.05 <i>ns</i>

\*\*\*  $p < 0.0001$ , \*  $p < 0.01$ ; *ns*: not significant (Pearson's correlation test).

Table 4. Relationships between estimated parameters and the original parameters (see text).

Estimated parameters	Slope	Intercept	$R^2$
$P_{max\_season}$	0.72	3.56 ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ )	0.83
est_ $P_{max}$	0.95	1.54 ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ )	0.88
$\phi_{season}$	0.66	0.009 ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ )	0.78
est_ $\phi$	0.78	0.007 ( $\text{mol CO}_2 [\text{mol photon}]^{-1}$ )	0.82

Seasonal parameters ( $P_{max\_season}$ ,  $\phi_{season}$ ) were estimated by using RVI and corrected (est\_ $P_{max}$ , est\_ $\phi$ ) using meteorological factors ( $n = 171$  for  $P_{max}$ , and  $n = 176$  for  $\phi$ ).

Table 5. Regression parameters for comparison by linear regression of GPP estimated using three different parameter sets with observed GPP. GPP was estimated at half-hourly time steps ( $n = 4025$ ).

Parameters used	Slope	$R^2$	SE ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ )
Original parameters by light-response curve-fitting	0.94	0.87	3.67
Estimated $P_{max}$ by RVI and $\phi$ by RVI	0.96	0.83	4.21
Estimated $P_{max}$ by EVI and $\phi$ by RVI	0.95	0.84	4.23