

HOKKAIDO UNIVERSITY

Title	A Computer Program for the Calculation of Piezomagnetic Field due to a Spherical Pressure Source (Mogi Model) in the Inhomogeneously Magnetized Crust
Author(s)	UTSUGI, Mitsuru
Citation	Journal of the Faculty of Science, Hokkaido University. Series 7, Geophysics, 11(4), 739-751
Issue Date	2000-03-24
Doc URL	http://hdl.handle.net/2115/8859
Туре	bulletin (article)
File Information	11(4)_p739-751.pdf



Jour. Fac. Sci., Hokkaido Univ., Ser. VII (Geophysics), Vol. 11, No. 4, 739-751, 2000.

A Computer Program for the Calculation of Piezomagnetic Field due to a Spherical Pressure Source (Mogi Model) in the Inhomogeneously Magnetized Crust

Mitsuru Utsugi

Division of Earth and Planetary Sciences, Graduate School of Science, Hokkaido University, Sapporo 060-0810, Japan

(Received November 30, 1999)

Abstract

We developed an analytical method for evaluating the piezomagnetic field considering the inhomogeneously magnetized crust. In this paper, we present a FOR-TRAN program to calculate the piezomagnetic effect due to a spherical pressure source (the Mogi model) in the inhomogeneously magnetized crust.

1. Introduction

In the previous studies (Sasai, 1991; Utsugi et al., 1999), the analytical solution of piezomagnetic effect has been obtained in the case of uniformly magnetized crustal model. However, it is self-evident that the earth's crust is inhomogeneously magnetized. Then, to represent the inhomogeneity of the crustal magnetization, we divide the crust into a number of compartments. Each compartment is assumed to have its own uniform magnetic properties such as the magnetization and the stress sensitivity. The geomagnetic field change at a certain point on the earth's surface may be approximated by the sum of the piezomagnetic field derived from each compartment. The piezomagnetic field is generally expressed by the surface integral of displacement and its derivation over the boundary surface of magnetized region. In the case of the compartment model, the surface integral becomes finite one. This integral cannot be solved analytically. However the line integral with respect to either coordinate can be represented analytically using elliptic integrals. There are several simple algorithms to make rapid and exact evaluation of the elliptic integrals. Using these algorithms, the piezomagnetic field can be expressed by line integral. Through numerical evaluations of this line integral by using the double exponential method (e.g. Takahashi and Mori, 1974), we can obtain the piezomagnetic field considering the inhomogeneity of the crustal magnetization.

In this paper, we present a FORTRAN program which calculates the piezomagnetic effect due to a spherical pressure source (the Mogi model : Mogi, 1958) in the inhomogeneously magnetized crust.

2. Geomagnetic field changes due to Mogi model

We consider the coordinate system as shown in Fig. 1. A semi-infinite elastic medium occupies z>0. We assume that the region V_1 is uniformly magnetized cube and outside V_2 is demagnetized. We also assume that the elastic properties such as Lame constants λ , μ are common in the regions V_1 and V_2 . The analytical solution of the displacement **u** due to the Mogi model is obtained by Mindlin and Cheng (1950) and Yamakawa (1955) as follows:

$$u_x = \frac{C}{2\mu} \left[\frac{x}{R_1^3} + 2\frac{x}{R_2^3} - \frac{6xz(z+D)}{R_2^5} \right]$$
(1)

$$u_{y} = \frac{C}{2\mu} \left[\frac{y}{R_{1}^{3}} + 2\frac{y}{R_{2}^{3}} - \frac{6yz(z+D)}{R_{2}^{5}} \right]$$
(2)

$$u_{z} = \frac{C}{2\mu} \left[\frac{z - D}{R_{1}^{3}} - 2\frac{D}{R_{2}^{3}} - \frac{6z(z + D)^{2}}{R_{2}^{5}} \right]$$
(3)

where $R_1 = \sqrt{x^2 + y^2 + (z - D)^2}$, $R^2 = \sqrt{x^2 + y^2 + (z + D)^2}$ and D is the source depth. The moment C is given by



Fig. 1. Coordinate system, source and crustal model are shown. A semi-infinite elastic medium occupies z > 0. The cubic region $V_1(L \times W \times H \text{ km}^3)$ is uniformly magnetized and outside V_2 is demagnetized. In this medium, a spherical pressure source (the Mogi model) is assumed.

$$C = -\frac{1}{2}a^3 \varDelta P,$$

where *a* is the radius of the sphere and ΔP is the hydrostatic pressure acting on the surface of the sphere.

According to the representation theorem for the piezomagnetic field (Sasai, 1991), the geomagnetic change $\Delta \mathbf{M}$ is expressed by the integral over the boundary surface S_1 (Fig. 1) of the cubic region:

where δ_{ij} is the Kronecker's delta, J_k is the *k*-th component of initial magnetization within V_1 and **u** is the displacement vector. \mathbf{r}_0 and **r** indicate the observation point and the arbitrary point within the medium, respectively.

Substituting eqs. (1) to (3) into eq. (4), the geomagnetic field changes are written in the form of surface integral of function $f(1/R, 1/\rho)$. In the present case, we have to make the finite surface integral because S_1 is finite. As mentioned in the previous chapter, we cannot solve this finite surface integral analytically. However, using elliptic integrals, the integral with respect to either coordinate can be represented by the analytical form. Then we can transform the surface integral of eq. (4) to the line integral as shown in the following chapter.

3. Line integrals of $f(1/R, 1/\rho)$

In eq. (4), we see the following integrals of function $f(1/R, 1/\rho)$:

$$\Phi_{ij}(y, z, x_0, y_0, z_0; D) = \int \frac{1}{R^i \rho^j} dx, \qquad (5)$$

$$\Psi_{ij}(y, z, x_0, y_0, z_0; D) = \int \frac{x}{R^i \rho^j} dx, \qquad (6)$$

These integrals can be solved analytically using the following elliptic integrals :

$$F(\varphi, k) = \int_0^{\varphi} \frac{1}{\sqrt{1 - k^2 \sin^2 \varphi'}} d\varphi' \quad \text{(First Kind)}, \tag{7}$$

M. Utsugi

$$E(\varphi, k) = \int_0^{\varphi} \sqrt{1 - k^2 \sin^2 \varphi'} d\varphi' \quad \text{(Second Kind)}. \tag{8}$$

For example, we consider the following integral:

For the convenience, we denote R and ρ as follows:

$$\begin{cases} R^2 = x^2 + \zeta^2, \\ \rho^2 = x^2 - 2x_0 x + c_0^2, \\ \zeta^2 = y^2 + (z - D)^2, \\ c_0^2 = x_0^2 + c^2, \\ c^2 = (y_0 - y)^2 + (z_0 - z)^2, \end{cases}$$

where (x, y, z) is the arbitrary point within the magnetized region of the medium, and (x_0, y_0, z_0) is the observation point.

To express eq. (9) using the elliptic integrals, we introduce the following variable transform :

$$x \to \frac{\alpha t + \beta}{t + 1},\tag{10}$$

where

$$\begin{cases} \alpha = \frac{c_0^2 - \zeta^2 - D}{2x_0}, \\ \beta = \frac{c_0^2 - \zeta^2 + D}{2x_0} \quad (\alpha \le \beta), \\ D^2 = (c_0^2 - \zeta)^2 + 4x_0^2 \zeta^2. \end{cases}$$

Using eq. (10), R and ρ become more simply:

$$R^{2} = \frac{1}{(t+1)^{2}} \{ (\alpha^{2} + \zeta^{2})t^{2} + (\beta^{2} + \zeta^{2}) \},$$

$$\rho^{2} = \frac{1}{(t+1)^{2}} \{ (\alpha^{2} - 2x_{0}\alpha + c_{0}^{2})t^{2} + (\beta^{2} - 2x_{0}\beta + c_{0}^{2}) \}$$

Therefore the terms 1/R and $1/\rho$ are written as follows:

$$\frac{1}{R} \rightarrow \frac{1}{\sqrt{p(\alpha)}} \frac{|t+1|}{\sqrt{t^2 + \xi^2}}, \quad \frac{1}{\rho} \rightarrow \frac{1}{\sqrt{q(\alpha)}} \frac{|t+1|}{\sqrt{t^2 + \eta^2}},$$

where

$$\begin{cases} p(s) = s^{2} + \zeta^{2}, \\ q(s) = (x_{0} - s)^{2} + c^{2}, \\ \xi^{2} = p(\beta)/p(\alpha), \\ \eta^{2} = q(\beta)/q(\alpha) \quad (\xi^{2} > \eta^{2}), \end{cases}$$

and eq. (9) is rewritten as follows:

$$\Phi_{11} = \frac{\alpha - \beta}{\sqrt{p(\alpha)q(\alpha)}} \int \frac{1}{\sqrt{t^2 + \xi^2} \sqrt{t^2 + \eta^2}} dt.$$
(11)

743

Transforming $t \to \varphi = \tan^{-1}\left(\frac{t}{\eta}\right)$, we obtain the following result :

$$\Phi_{11} = \frac{\alpha - \beta}{\sqrt{p(\alpha)q(\alpha)}} \frac{1}{\xi} F(\varphi, k),$$

where $k^2 = (\xi^2 - \eta^2)/\xi^2$. To evaluate the elliptic integrals, there are some algorithms (Cayley, 1961; Byrd and Friedman, 1954) without solving eqs. (7) and (8) directly. Using these algorithms, we can evaluate φ_{11} easily. With the same manner, φ_{ij} and Ψ_{ij} are solved analytically. The exact forms of φ_{ij} and Ψ_{ij} are given by Utsugi (1999).

The geomagnetic change is written by the following line integral of **g** which is an arbitrary function of ϕ_{ij} and ψ_{ij} :

$$\Delta \mathbf{M}^{k}(\mathbf{r}_{0}) = C_{k} \int \mathbf{g}(\boldsymbol{\varphi}_{ij}(\mathbf{r}_{0},\mathbf{r}), \boldsymbol{\Psi}_{ij}(\mathbf{r}_{0},\mathbf{r})) dl_{\mathbf{r}}.$$
(12)

From numerical calculation of this integral, we can evaluate the piezomagnetic change considering the inhomogeneously magnetized crust. To evaluate eq. (12), we use the double exponential integral method (DEM).

4. Programs

The source list of the programs for calculating the piezomagnetic field is given in Appendix. The subroutine 'MGINHOMO' calculates the geomagnetic

	Table 1. In	nput paramete	ers.
X0, Y0, Z0	x0, y0, Z0	(km)	Observation point
CX0, CY0, CZ0		(km)	Location of the center of the cube
CL, CW, CH	L, W, H	(km)	Length, width and height of the cube
C0	$a^{3} \Delta P/2$	(km³•bar)	Moment of Mogi model
D0	D	(km)	Source depth
AMU	μ	(cgs)	Rigidity
POI	v		Poisson ratio
CMZX, CMZY, CMZZ	$\mathbf{J}=(J_x,J_y,J_x)$	(A/m)	Magnetization vector within the region V ₁
BETA	β	(bar ⁻¹)	Stress sensitivity

change due to a magnetized cubic block. This subroutine requires the parameters as shown in Table 1 and returns eastward (DMX), northward (DMY), downward (DMZ) and total force (DMF) components of geomagnetic change.

The double precision functions 'MGXY X_i ', 'MGXZ X_i ' and 'MGYZ X_i ' ($X_i = X, Y \text{ or } Z$) calculate the contributions from the x - y, x - z and y - z plane of V_i , respectively. The subroutine 'PHSIIJ' calculates Φ_{ij} and Ψ_{ij} . The elliptic integrals which appear in Φ_{ij} and Ψ_{ij} are calculated by the subroutine 'ELLIP-FE'. The subroutine 'DEMINT' calculates the numerical line integral of the functions of Φ_{ij} and Ψ_{ij} numerically using DEM.

5. Numerical example

In Fig. 2, we show a numerical example. This figure shows the profiles of the total force of geomagnetic change along the *y* axis $(x_0=0, z_0=1)(m)$. The case A in Fig. 2 is based on the uniformly magnetized crust : a layer $0 < z < H_c$ =15 km (H_c indicates the Curie depth) is uniformly magnetized by 1 A/m. The case B is based on the inhomogeneously magnetized crust as shown in Fig. 1. The intensity of magnetization within the cube is assumed as 1 A/m. In both



Fig. 2. Profiles of the total force of geomagnetic change along the y axis ($x_0=0$, $z_0=1$ m). Cases A and B indicate the geomagnetic changes based on an uniformly magnetized crustal model and a cubic model, respectively.

1 4010 51 4	suble model parameters.	
(X0, Z0)	(0, 1)	(m)
(CX0, CY0, CZ0)	(10, 0, 7.5)	(km)
(CL, CW, CH)	(10, 10, 15)	(km)
C0	10^{3}	(km³·bar)
D0	5	(km)
AMU	3.5×10^{11}	(cgs)
POI	0.25	
(CMZX, CMZY, CMZZ)	$(0, 1/\sqrt{2}, 1/\sqrt{2})$	(A/m)
BETA	10-4	(bar ⁻¹)

Table 2. Cubic model parameters.

cases, the magnetic inclination and declination are assumed as 45° and 0° , respectively. The model parameters are given in Table 2. Through this calculations, it becomes clear that the piezomagnetic effect is enhanced around the edges of the cubic block (at $y_0=5$ and $y_0=15$ km). This is caused by the fact that, unlike a uniform medium, the magnetic fields arising from stress-induced magnetic dipoles do not cancel with one another around the edges. The mechanism of this enhancement is well discussed in Oshiman (1990) and Utsugi (1999).

References

- Byrd, P.F. and M.D. Friedman, 1954. Handbook of elliptic integrals for engineers and physicists, Springer Verlag, pp. 335.
- Cayley, A., 1961. An elementary treatise on elliptic functions, Dover Publications, Inc, pp. 386.
- Mindlin, R.D. and D.H. Cheng, 1950. Nuclei of strain in the semi-infinite solid. J. Applied. Phys, **21**, 926-930.
- Mogi, K., 1958. Relations between the eruptions of various volcanoes and the deformations of the ground surfaces around them. Bull. Earthq. Res. Inst., Univ. Tokyo, 36, 99-134.
- Oshiman, N., 1990. Enhancement of tectonomagnetic change due to non-uniform magnetization in the Earth's crust -two dimensional case studies. J. Geomag. Geoelectr., 42, 607-619.
- Sasai, Y., 1991. Tectonomagnetic modeling on the basis of linear piezomagnetic effect. Bull. Earthq. Res. Inst., Univ. Tokyo, **66**, 585-722.
- Takahashi, R. and M. Mori, 1974. Double exponential formulas for numerical integration, Publ. R.I.M.S., Kyoto Univ., 9, 721-741.
- Utsugi, M., 1999. A theoretical study on seismomagnetic effect considering the inhomogeneously magnetized Earth's crust. Ph.D. Thesis, Hokkaido Univ., pp. 124.
- Utsugi, M., Y. Nishida, and Y. Sasai, 1999. Piezomagnetic potentials due to an inclined rectangular fault in a semi-infinite medium. Geophys. J. Int., **133** (in press).
- Yamakawa, N., 1955. On the strain produced in a semi-infinite elastic solid by an interior source of stress. J. Seism. Soc. Japan, (II), 8, 84-98 (in Japanese with English abstract).

1	SUBROUTINE WGINHOMO(X0, Y0, Z0, CX0, CY0, CZ0,
Z	& CL, CW, CH, CO, DO, AMU, POI, CMZX, CMZY, CMZZ,
4 (***	& UMA, UME, UME, UME)
śč	
6 C	GEOMAGNETIC FIELD CHANGE AT OUTSIDE THE MEDIUM
80	RASED ON THE CUBIC BLOCK MODEL
9 č	CODED BY M. UTSUGIDEC 1999
10 C	
11 (***	TNDIT DARAMETERS
iàc	X0, Y0, Z0 : LOCATION OF THE OBSERVATION POINT
14 C	CX0, CY0, CZ0 : LOCATION OF CENTER OF CUBIC BLOCK
15 C	CL, CW, CH : LENGTH, WIDTH AND HEIGHT UP CUBE
17 (DØ : SOURCE DEPTH
18 Č	AMU : RIGIDITY
19 C	POI : POISSON RATIO
20 0	SURFACE OF CUBE
22 C	BETA : STRESS SENSITIVITY
23 (
25 0	DNX.DNY.DMZ.DMF ; EAST, NORTH, DOWNWARD COMPONENTS
26 C	AND TOTAL FORCE OF GEOMAGNETIC
27 C	FIELD CHANGE (NT)
28 (THPLICTT REAL*R (A-H. O-Z)
30	INTEGER NEND, NPOW
31	COMMON /COM1/ NEND, NPOW, AO(2),
33	CONMON /COM2/ (00, C0X, C0Y, C0Z
34 C	
35	EXTERNAL NGXYX, NGXYY, NGXYZ
30	EXTERNAL MGXZX, MGXZY, MGXZZ EXTERNAL MGYZX, MGYZY, MGYZZ
38 C	
39	EPS = 1.0-15
41	(NZ0 - SORT(CMZX**2+CMZY**2+(MZZ**2)
42	SMM-CMZ0*BETA*AMU*(1.D0+POI)
43	C00-0.5D0*SMM*1.D-7
45	C0Y=-CHZY*C00
46	COZ-CMZZ*COO
47 (CALL ETAR (NEND NPOW, AO, AM, AP, BO, BB)
49 C	
50 C***	DMXY : CONTRIBUTIONS FROM X-Y PLANE ***
52 0	DMXYX : EASTWARD COMPONENTS
53 C	
54	CALL DEMINT (MGXYX, X0, Y0, Z0, D0, CZ-0.5D0°CH,
35 56	& CY-0,5D0*CH, CY+0.5D0*CH,
57	& EPS, L, DMXYX0)
58	CALL DEMINT (NGXYX, XØ, YØ, ZØ, DØ, CZ+0.5D0*CH,
59	& CX-0.500°CL, CX+0.500°CL, & CY-0.500°CW,
61	& EPS, L, DMXYXH)
62 C	
63 64 (UNATA=UNATA0-UNATAH
65 č	DMXYY: NORTHWARD COMPONENTS
66 C	
67	CALL DEMINE (MGXYY, X0, Y0, Z0, D0, CZ-0.5D0°CH, $z = CY_0 5D0°CE (Y_0, SD0°CE)$
69	& CY-0.5D0*CW, CY+0.5D0*CW,

70		
70	°	EF3, C, UMATTU)
<u>/1</u>	_CALL	UPRIAL (MONTH, AD, 10, 20, 00, C270.300 CH, $(x, y) = (x, y)$
12	<u><u></u></u>	
13	•	Enc i Nuvvu
1.	•	EFS, L, UMATTRY
75 0		
10	DMXY	F=DHATTO-DHATTN
77 0		DOMINING CONDONICHIEC
78 C	DMXY.	Z: DOWNWARD COMPONENTS
79 C		
80	CALL	DEMINI (MGXT2, X0, T0, 20, D0, C2-0.500°CH,
81	8	CX-0.500*CL, CX+0.500*CL,
82	8	CY-0.5D8*CW, CY+0.5D8*CW,
83	8	EPS, L, DMXYY0)
84	CALL	DEMINT (MGXYZ, X0, Y0, Z0, D0, C2+0.500°CH,
85	Ł	CX-0,5D0*CL, CX+0.5D0*CL,
86	8	CY-0.5D0°CW, CY+0.5D0°CW,
87	8	EPS, L, DWXYYH)
88 C		
89	DMXY	Y-DHXYY0-DHXYYH
90 C		
91 (***	DMXZ	: CONTRIBUTIONS FROM X-Z PLANE ***
97 C		
93 č	DMX7	X : FASTWARD COMPONENTS
94 C		
95	CALL	DEMINT (MGXZX, X0, Y0, Z0, D0, CY-0,500°CW,
96	2	CX-0.5D0*CL. CX+0.5D0*CL.
07		C7-0 500°CH C7-0 500°CH
	a .	
30	°	EF3, L, UMALAU, DA VA 70 DA (V.A SDA*CH
379		CY = COACT (CALA, AS, 10, 20, 00, CITOLODO CH, CALASSIC (CH, CALASSIC) (CT, ASSACL
100	•	
101	ě.	
102	5	EPS, L, DMAZAH)
103 C		
104	DMXZ:	X=DMXZXØ-DMXZXH
105 C		
106 C	DMXZ	Y: NORTHWARD COMPONENTS
107 C		
108	CALL	DEMINT (NGXZY, X0, Y0, 20, D0, CY-0.5D0°CW,
109	8	CX-0.5D0*CL, CX+0.5D0*CL,
110	8	CZ-0.5D0°CH, CZ+0.5D0°CH,
111	8	EPS, L, DMXZY0)
112	CALL	DEMINT (MGXZY, X0, Y0, Z0, D0, CY+0.500°CW,
113	8	CX-0.5D0*CL, CX+0.5D0*CL,
114	8	CZ-0, SD0*CH, CZ+0, SD0*CH,
115	8	EPS, L, DMXZYH)
116 C		
117	DMXZ	Y=DMXZY@-DMXZYH
118 C		
119 č	DMX7	7: DOWNWARD COMPONENTS
170 0	Distance of	
121	CALL	DEMINT (NGX77 X0, Y0, 70, D0, CY-0,500°CW,
177		(Y-0 SD0+(1 (Y+0 SD0+(1
122	a	
123	2	
124	°	EF3, L, UMALLUJ DEULINT (UCV77 VA VA 74 DA CV.A 5000()
125		C_{Y} a Spart (C_{Y} a Spart)
120	•	CATE CRAFTLE CATEGORIE
127	e .	
128	ð.	EPS, L, UMALLE)
129 C	-	NIX776 DIX770
130	DMXZ	<u></u>
131 C		
132 (***	DMYZ	: CONTRIBUTIONS FROM Y-Z PLANE ***
133 C		
134 C	DMYZ	K : EASTWARD COMPONENTS
135 C		
136	CALL	DEMINT (MGYZX, X0, Y0, Z0, D0, CX-0.5D0*CL,
137	•	CY-0.500*CW. CY+0.500*CW.
101		

Appendix

139	EPS, L, DWYZX0)
140	CALL DEMINT (MGYZX, X0, Y0, Z0, D0, CX+0.5D0*CL,
142	
143	2 EPS DWY2YH)
144 C	a cro, c, owiekily
145	DMYZX-DMYZX0-DMYZXH
146 C	
147 C	DWYZY: NORTHWARD COMPONENTS
148 C	
149	CALL DEMINT (MGYZY, X0, Y0, Z0, D0, CX-0.5D0*CL,
150	& CY-0.5D0*CW, CY+0.5D0*CW,
151	& CZ-0.5D0*CH, CZ+0.5D0*CH,
152	& EPS, L, DMYZYØ)
153	CALL DEMINI (MGYZY, X0, Y0, Z0, D0, CX+0.5D0*CL,
155	
156	
157 C	
158	DMYZY-DMYZYØ-DMYZYH
159 C	
160 C	DMYZZ: DOWNWARD COMPONENTS
161 C	
162	CALL DEMINT (MGYZZ, X0, Y0, Z0, D0, CX-0.5D0°CL,
163	& CY-0.5D0*CW, CY+0.5D0*CW,
164	& CZ-0.5D0*CH, CY+0.5D0*CW,
105	& EPS, C, UNITED)
167	CALL DEMINI (MGTZZ, X0, T0, Z0, D0, CX+0.500*(L,
168	2 (7-0 500°CH, C1+0.500°CH,
169	& EPS. L. DWYZZH)
170 C	
171	DWYZZ=DWYZZ0-DWYZZH
172 C	
173	DMX=DMXYX+DMXZX+DWYZX
174	DWY-DMXYY+DMXZY+DWYZY
175	DMZ=DMXYZ+DMXZZ+DMYZZ
176 C	
179 C	UMF=(UMA*(MZX+UMY*(MZY+UMZ*(MZZ)/(MZØ
179	RETURN
180	END
181 C	
182 C	
183 C	
184	SUBROUTINE DEMINT (FUNC, X0, Y0, Z0, D0,
185	& Z, X1, XZ, A, B, EPS, L, V)
186 (
188	IMPLICIE REALTO (A-H, U-Z) INTECER NEWD NDOW
189	CONNON /CON1/ NEND NPOW AD(2)
190	& AP(608, 2), AM(608,2) AO BR(608)
191 C	
192	DATA HALF, EPSO / 0.5D0, 1.0D-32 /
193	DATA EPSM, EPSP / 0.D0, 0.D0 /
194 C	
195	FAC = (B - A) * HALF
196	IF (L .EQ. 0) THEN
109/	SHEM = (B + A) * HALF
199	
200	SHEW = 0.00
201	SHEP = 0.00
202	ENDIF
203 C	
204	L1 = L + 1
205 C	
200	IF (ABS(EPS) .GE. EPSO) THEN
201	EMOA = WRP(Eb2)

208 EL SE 209 EPSV - EPSO 210 ENDIF 211 C 212 EPSQ = 0.2D0 * SQRT(EPSV) 213 C 214 215 C H - HALF 216 217 218 C IS = 2**NPOW IM - IS 219 220 221 222 KM - 0 KP = 0 NM - 0 223 C 224 225 VNEW = FUNC(X0, Y0, Z0, D0, & AO(L1)*FAC+SHFP, Z, X1, X2) * BO 226 • 227 * ---- INITIAL STEP -----228 • INTEGRATE WITH MESH SIZE . 0.5 229 • 230 • 231 AND CHECK DECAY OF INTEGRAND DO 10 I - IS, NEND, IM 232 C 233 234 IF (KM .LE. 1) THEN WM = FUNC(X0, Y0, Z0, D0, MA(I, L1)*FA(+ SMFM, Z, X1, X2) * BB(I) VNEW = VNEW + MM IF (ABS/KM) .LE. EPSV) THEN KM = KM + 1 FIST (CM .CE. 2) NM = I - IM FIST (CM .CE. 2) NM = I - IM 234 235 236 237 238 239 8 240 241 ELSE KM - 0 242 243 244 (245 246 247 ENDIF ENDIF IF (KP .LE. 1) THEN MP - FUNC(X0, Y0, Z0, D0, AP(I, L1)*FAC + SHFP, Z, X1, X2) * BB(I) VNEW - VNEW + WP 8 248 249 250 IF (ABS(WP) .LE. EPSV) THEN KP = KP + 1 251 252 253 254 255 256 C IF (KP .GE. 2) NP = I - IM ELSÊ KP - 0 ENDIF ENDIF 257 258 C IF (KM .EQ. 2 .AND. KP .EQ. 2) GOTO 11 259 10 CONTINUE 260 C 261 11 CONTINUE 262 C 263 IF (NM - N 264 NM - N 266 EPSM -266 ENDIF 266 ENDIF 267 C 268 TE (ND E IF (NM .EQ. 0) THEN NM - NEND EPSM - SQRT (ABS(WM)) 268 269 270 271 272 C IF (NP .EQ. 0) THEN NP = NEND EPSP = SQRT (ABS(WP)) ENDIF 273 EPSQ - MAX (EPSQ, EPSM, EPSP) 274 • 275 • ----- GENERAL STEP -----276 •

```
277 VOLD - H * FAC * VNEW

278 C

279 D 20 WSTEP = 1, NPOW

281 C

282 VNEW - 0.0

283 I H = IS

284 IS = IS / 2

285 C D0 540 I = IS, NN, IH

287 VNEW - VNEW

288 & + FUNC(X0, Y0, Z0, D0,

289 S40 CONTINUE

298 C ONTINUE

298 C ONTINUE

294 & + FUNC(X0, Y0, Z0, D0,

295 S40 CONTINUE

295 C CONTINUE

296 S50 C ONTINUE

296 S50 CONTINUE

297 C CONVERCED AND RETURN -----

308 V = VNEW

306 E HOLF

307 C H = H * HALF

308 VOLD = VNEW

308 VOLD = VNEW

310 C

200 CONTINUE

312 C VNEW - VNEW

314 RETURN

316 C END

317 C SUBROUTINE FIAB (NEND, NFOW, AO, AM, AP, BO, BE)

318 C

319 C

320 SUBROUTINE FIAB (NEND, NFOW, AO, AM, AP, BO, BE)

321 C VNEW = VNEW

332 PARAMETER (ALF = ONE / 2)

323 PARAMETER (MALF = ONE / 2)

324 PARAMETER (MALF = ONE / 2)

325 PARAMETER (MALF = ONE / 2)

326 A 9 = 0.9999 9999 9998 B0

331 H = 2 * ATAN (ONE)

332 O IND = NG

333 A O(1) = 0.D0

334 C H = H E P(H)

335 C MOM = NG

336 A 0(2) = 1.D0

337 NEND = 608

339 A O(1) = 0.D0

331 B C

339 A O(1) = 0.D0

341 B O = PH

342 C H = EXP (H)

344 E H = 1.D0

345 C
                                                                                  D0 540 I = IS, NM, IH

VNEW - VNEW

+ FUNC(X0, Y0, Z0, D0,

AN(I, L1)*FAC + SHFM, Z, X1, X2) * BB(I)

CONTINUE
                                                                               VNEW - VNEW
+ FUNC(XØ, YØ, ZØ, DØ,
AP(I, L1)*FAC + SHFP, Z, X1, X2) * B8(I)
CONTINUE
                                                                      SUBROUTINE FIAB (NEND, NPOW, AO, AM, AP, BO, BB)
```

346	DO 10 I = 1. NEND
347	EN - EH + EN
348	ENI - 1.D0 / EN
349	SH = (EN - ENI) • HALF
350	CH = (EN + ENI) + HALF
351	EXS = EXP (PH + SH)
352	EXSI = 1.00 / EXS
353	CHSI = 2.00 / (EXS + EXSI)
354	AP(I,1) = ((EXS - EXSI) + HALF) + CHSI
355	IF $(AP(1,1), GE, A9) AP(1,1) = A9$
350	AP(1,2) = EXST + CHST
357	AM(1,1) = -AP(1,1) AW(1,2) AD(1,2)
359	AR(1, c) = - Ar(1, c) $BR(1) = DW + (U + c) (U + c)$
360 10	
361 (CONTINUE
362	AP(508.2) = AP(507.2)
363	AN(608,2) - AN(607,2)
364 C	
365	RETURN
366	END
367 C	
368 C	
369 C	
370	DOUBLE PRECISION FUNCTION MGXYX(X0, Y0, Z0, D0, Y, Z, X1, X2)
3/1	IMPLICIT REAL'S (A~H, U-Z)
372	DIRENSION DI (10,10), PRI (10,10), PSI (10,10)
374 (
375	712-1447-0014(7-00)
376	$(2^{-}(Y_{0},Y_{1})^{+}(Y_{0},Y_{1})^{+}(7_{0},7)^{+}(7_{0},7)$
377	2T-SORT(2T2)
378	(C-SORT(CC2)
379 C	
380	CALL PHSIIJ(X0, ZT, CC, X1, X2, PHI, PSI)
381 C	
382	MGXYX=C0X*(3.D0*(Z0-Z)*(X0*PHI(3,5)-PSI(3,5)))
383	& +C0Y*(3.D0*(Z0-Z)*Y*PHI(3,5))
384	& +C0Z*(3.D0*(Z0-Z)*Z*PHI(3,5))
385	KE LUKN
397 C	END
388 0	
389 C	
390	DOUBLE PRECISION FUNCTION MEXYY(X0, Y0, 20, 00, Y, 2, X1, X2)
391	IMPLICIT REAL*8 (A-H. 0-Z)
392	DIMENSION DI(10,10), PHI(10,10), PSI(10,10)
393	COMMON /COM2/ CÓO, COX, COY, CÓZ
394 C	
395	ZT2=Y*Y+(Z-D0)*(Z-D0)
396	<u>CC2=(Y0-Y)*(Y0-Y)+(Z0-Z)*(Z0-Z)</u>
397	ZT=SQRT(ZTZ)
398 200 c	((=SQK)(((2)
399 (
401 0	CALL PHOILD(NO, 2), (C, XI, X2, PHI, POI)
402	MCXYY_CAX*(3 DA*(70_7)*(YA_Y)*(PHT(3 5)_PST(3 5)))
403	\mathbf{A} \mathbf{A} \mathbf{A} \mathbf{A}
404	A +(07*(3,D0*(20-7)*7*(Y0-Y)*PHI(3,S))
405	RETURN
406	END
407 C	
408 C	
409 C	
410	DOUBLE PRECISION FUNCTION MGXYZ(X0, Y0, Z0, D0, Y, Z, X1, X2)
411	IMPLICII REAL®8 (A-H, O-Z)
412	UIMENSION UI(10,10), PHI(10,10), PSI(10,10)
413	CUMMCUN / CUM2/ CUM3, CUX, CUY, CU2
414 (

415	772=¥*¥+(7-00)*(7-00)
416	
417	
417	21-50KT(212)
418	CC=SQRT(CC2)
419 C	
470	CALL PHISTING TT CC Y1 Y2 PHIL DCL)
421 6	CALC FISTIS(NO, 21, CC, X1, X2, FRI, FSI)
421 0	
422	MGXYZ=C0X*(3.D0*(Z0-Z)*(Y0-Y)*(PHI(3.3)-PSI(3.5)))
423	\$ +(0Y*(3, D0*(70-7)*Y*(Y0-Y)*PHI(3,5))
474	
435	a +(02 (3.00 (PHI(1,1)-(20-2) (20-2) PHI(3,5)))
425	RETURN
426	END
427 C	
478 6	
400 0	
429 L	
430	DOUBLE PRECISION FUNCTION MGXZX(X0, Y0, Z0, D0, Z, Y, X1, X2)
431	THPLTCTT REAL *8 (A-H Q-7)
427	DTHENET ON DI (10 10) DUT (10 10) DET (10 10)
136	DIMENSION DI (10,10), PHI (10,10), PSI (10,10)
433	COMMON /COM2/ C00, C0X, C0Y, C0Z
434 C	
435	ZT2=Y*Y+(7-D0)*(7-D0)
436	
430	_=\j=-j-(10-1)+(20-2)*(20-2)
43/	21=5QKT(212)
438	CC=SQRT(CC2)
439 C	
440	
	CHLL FIDILIJ(NU, 21, CC, A1, A2, PH1, PS1)
441 C	
44Z	MGXZX+C0X*(3,D0*(Y0-Y)*(X0*PHI(3,3)-PST(3,5)))
443	\$ +(0Y*(3 D0*(Y0-Y)*Y*PHT(3 5))
444	
	a +(02*(3.D0*(10-1)*2*PHI(3,5))
445	RETURN
446	END
447 C	
A49 C	
440 0	
449 C	
450	DOUBLE PRECISION FUNCTION MGX7Y(XA, YA 7A DA 7 Y X1 X2)
451	INDITITI DEALESS AL (-7)
465	DIMENSION DI (10 (0) DIT(10 (0) DET(10 (0)
434	DIMENSION DI(10,10), PHI(10,10), PSI(10,10)
453	COMMON /COM2/ COO, COX, COY, COZ
454 C	
455	712-1414 (7-00)+(7-00)
456	
450	$(22=(10-1)^{-}(10-1)+(20-2)^{-}(20-2)$
457	ZT-SQRT(ZT2)
458	CC-SORT(CC2)
459 C	
460	
464 6	CHLL PHILID(AW, ZF, CC, AI, AZ, PHI, PSI)
401 C	
462	MGXZY=C0X*(3,D0*(Y0-Y)*(Y0-Y)*(PHI(3,3)-PSI(3,5)))
463	4 +C0Y*(3, D0*(Y0-Y)*Y*(Y0-Y)*PHT(3,5))
464	
466	
403	REIUKN
466	ENU
467 C	
468 C	
460 0	
+09 (
470	DOUBLE PRECISION FUNCTION MGXZZ(X0, Y0, Z0, D0, Z, Y, X1, X2)
471	IMPLICIT REAL®8 (A-H. O-Z)
472	DIMENSION DICIA 10) PHICIA 10) PSICIA 10)
472	(0,0,0,0,0,0,0,0,0,0,0,0,0,0,0,0,0,0,0,
113	CUMHEUN / CUM2/ COO, COX, COT, CO2
474 C	
475	ZT2=Y*Y+(Z-D0)*(Z-D0)
476	(C2-(Y0-Y)+(Y0-Y)+(70-7)+(70-7)
477	
470	41=30K1(412)
410	CC=SQKT(CC2)
479 C	
480	CALL PHSTTICX0 7T CC X1 X2 PHT DST
481 C	
487	
402	MUALE COA (3.00°(Y0-Y)*(Y0-Y)*(PHI(3,3)-PSI(3,5)))
483	4 +LUT*(3.D0*Y*(Y0-Y)*(Y0-Y)*PHI(3,5))

484	\$ +(07*(3 Da*(Ya_Y)*PUT/3 3))
485	RETURN
486	FND
487 C	
488 C	
489 6	
490	DOUBLE PRECISION FUNCTION NEXTYCER VA 20 DA Y 7 V1 Y2)
401	THOLE IT DEALES (A U O T)
402	$\mathbf{D}_{\mathbf{M}} = \mathbf{D}_{\mathbf{M}} = $
402	COMPANY (CON12 (
493	Cummon / Cum2/ Cwa, Cwx, Cwr, Cw2
494 0	771 242 (7 00)4(7 00)
495	212=***+(2-00)*(2-00)
490	$(2 - (x - x)^{*}(x - x) + (2 - x)^{*}(2 - x))$
497	Z1-SQR1(212)
498	
499 C	
500	CALL PHSIIJ(YØ, ZT, CC, Y1, Y2, PHI, PSI)
501 C	
502	MGYZX=C0X*(3.D0*(X0-X)*X*PHI(3,5))
503	& +C0Y*(3.D0*(X0-X)*(X0*PHI(3,3)-PSI(3,5)))
504	& +C0Z*(3.D0*(X0-X)*Z*PHI(3.5))
505	RETURN
506	END
567 C	
508 C	
509 C	
510	DOUBLE PRECISION EUNCTION MCYZYCXA YA ZA DA Z Y X1 X2)
511	TWO IT (TT REAL * 4 4 7.)
512	$\frac{1}{10000000000000000000000000000000000$
512	COMPARISON (X, 10), (1
514 C	CUMMUM / CUM2/ CO0, C07, C07, C02
214 (
212	212-A*A+(2-00)*(2-00)
516	((2=(X0-X)*(X0-X)+(Z0-Z)*(Z0-Z)
517	ZT-SQRT(ZT2)
518	CC-SQRT(CC2)
51 9 C	1
520	CALL PHSIIJ(YØ, ZT, CC, Y1, Y2, PHI, PSI)
521 C	
522	MGYZY-C0X*(3.D0*(X0-X)*(X0-X)*PHI(3,5))
523	& +C0Y*(3.D0*(X0-X)*(Y0*PHI(3.3)-PSI(3.5)))
524	& +C0Z*(3,D0*(X0-X)*Z*(X0-X)*PHI(3,5))
525	RETURN
526	END
527 C	
528 C	
529 č	4
530	DOUBLE PRECISION FUNCTION MOV77CY0 VO 70 DO 7 V X1 X2)
531	TWOILT TT DEALTRE (A.W. 0.7)
537	DIMENSION DIVIA IA) DIVIA IA) DIVIA IA)
532	COMPON = C
534 6	CONHUNE / COM2 / COM, COX, COY, CO2
535	772 Y4Y (7 DA)8(7 DA)
533	
530	L(2=(A0-A)-(A0-A)+(20-2)*(20-2)
33/	
338	(C=SQRT(UC2)
539 C	
540	CALL PHSIIJ(YØ, ZT, CC, Y1, Y2, PHI, PSI)
541 C	
542	MGYZZ=C0X*(3.D0*X*(X0-X)*(X0-X)*PHI(3.5))
543	& +C0Y*(3.D0*(X0-X)*(Y0*PHI(3.3)-PSI(3.5)))
544	& +C0Z*(3,00*Ž*(Z0-Z)*(X0-X)*PHI(3,5))
545	RETURN
546	FND
547 C	Silv
SAR C	
540 C	
249 (
220	SUBRUDITNE PHISILJ(XØ, ZI, CC, X1, X2, PHI, PSI)
351	IMPLICI KEALTS (A-H,O-Z)
552	DIMENSION DI(10,10), PHI(10,10), PSI(10,10)

553 C		
554	712-21*21	
\$55	CC2=CC*CC	
556 C		
557	CALL PQAB(X0,ZT2,CC2,X1,X2,ALP,BET,PALP,PBET,	
558	& QALP, QBET, DK, XI, ET, T1, T2, PH1, PH2)	
559	CALL ELLIPFE(PH1,DK,ELPF1,ELPE1)	
560	CALL ELLIPPE(PHZ, BK, ELPFZ, ELPEZ)	
561	ELPF=ELPF2-ELPF1	
562	ELPE-ELPEZ-ELPEI	
503 (V12_V14V1	
565		
566	TX1-T1+T1+X12	
567	TX2-T2*T2+XI2	
568	TE1-T1*T1+ET2	
569	TE2=T2+ET2	
570 C		
571 C	111	
572	DI(1,1)=ELPF/XI	
573 C	113	
5/4	DI(1,3)=(X12*ELPE-E12*ELPE)/(X1*E12*(X12-E12))	
575 0	131 NT/2 11./ELDE ELDE1//VT#/VT2 ET211	
577	$U_1(3,1)=(CLFT*CLFC)/(A1*(A12*C12))$ $L_1(3)(Y12*C0PT/TY2*TC2)) = 1/(Y12*C0PT/TY1*TC1))$	
578 C		
579	DI(3,3)+((XI2+ET2)*ELPE-2,D0*ET2*ELPE)/(XI*ET2*	
580	6 (X12-FT2)*(X12-FT2))	
581	4 +(T2/(XI2*SORT(TX2*TE2))	
582	& -T1/(X12*SQRT(TX1*TE1)))/(X12*(X12-ET2))	
583 C	151	
584	DI(5,1)=((3.00*XI2-ET2)*ELPF	
585	4 -2.D0*(2.D0*XI2-ET2)*ELPE)/(3.D0*XI2*XI*(XI2-ET2)*(XI2-ET2))	
586	+(3.00*X12-2.00*ET2)*(12/(X12*SQRT(1X2*1E2))	
587	6 -11/(X12*SQR1(1X1*1E1)))/(3.D0*X12*X12*(X12-E12))	
200	$\mathbf{x} + (12)(X12^{-5}UR^{+}(1X2^{-1}X2^{-1}E2))$	
500 (a -11/(X12-50K+(1X1-1X1-1X1-1E1)))/(3.00-X12)	
591	DI(3 5)-(DI(3 3)-DI(5 1))/(XI2-FI2)	
592 C		
593	AB-ALP-BET	
594 C		
595 C	PHI11, PSI11	
596	PHI(1,1)=AB*DI(1,1)/SQRT(PALP*QALP)	
597	PSI(1,1)-AB*(BET*DI(1,1)	
598	40.5D0*ALP*(LOG(X12+E12+2.D0*F2*T2+2.D0*SQRT(1E2*T22))	
599	& -LOG(X12+E12+2.00*11*11+2.00*SQKI(1E1*(X1))))/SQKI(PALP*QALP)	
600	PH133 PH1/32 22_40#(01/1 1)_(V12,ET2 & DA\#01/1 2)	
692	$f(x_1, y_2) = f(x_1, y_2) - (x_1, y_2) - ($	
603		
604	<pre>4 +(XI2-1,D0)/((ET2-XI2)*(ET2-XI2)*TX2)</pre>	
605	<pre>k +(ET2-1,D0)/((ET2-XI2)*(ET2-XI2)*TE2))</pre>	
606	& -2.D0*(SQRT(TX1*TE1)	
607	& +(XI2-1.D0)/((ET2-XI2)*(ET2-XI2)*TX1)	
608	4 +(ET2-1.DØ)/((ET2-XI2)*(ET2-XI2)*TE1)))/(SQRT(PALP*QALP)**3)	
609 C	PHI35 PSI35	
610	PH1(3,5)-AB*(D1(1,1)-(X12+2.D0*12-15.D0)*D1(3,1)	
612	6 +3.00°(t)2°t)2-10.00°t)2(2+3.00)°U(3,3)	
613	<pre>a -(ti2::3-13.U0*ti2*ti2+13.U0*ti2+1.U0)*U1(3,3)) DST/3 5_AR*(ALD*D1/1 1)_(ALD*(YT2+7 NA*ET2))</pre>	
614	\$5 D0*(2 D0*AI PARET))*DT(3 1)	
615	A +(3, D0*ALP*ET2*ET2-10, D0*ET2*(2, D0*ALP*RET)	
616	& +5.D0*(ALP+2.D0*BET))*DI(3.3)	
617	6 -(ALP*ET2**3-5.D0*ET2*ET2*(2.D0*ALP+BET)	
618	& +5.D0*ET2*(ALP+2.D0*BET)+BET)*DI(3,5))	
619	RETURN	
670		
020	END	

622 C	
623 C	
624	SUBROUTINE POAB(X0,ZT2,CC2,X1,X2,ALP,BET,PALP,PBET,
626	INPLICIT REAL*8 (A-H.O-Z)
627	DATA F0,F1/0.D0,1.D0/
628 C	
629	X02=X0*X0 TC(YA NE CA) TUEN
631	DFT2=(((7+x02-7T2))*(((2+x02-7T2))+4, D0*x02*7T2
632	DET=SQRT(DET2)
633	ALP = -2.D0 * X0 * ZT2 / (CC2+X02-ZT2+DET)
634	BET=(CC2+X02-ZT2+DET)/(2.D0*X0)
636	RET2=RET*RET
637 C	
638	PALP=ALP2+ZT2
639	PBET-BETZ+ZTZ
641	QALF=(X0-ALF)*(X0-ALF)+(L2 QAET_(X0-ALF)*(X0-ALF)+(L2
642 C	
643	T1=(X1-BET)/(ALP-X1)
644	T2-(X2-BET)/(ALP-X2)
040 646	
647	XI=SORT(XI2)
648	ET=SQRT(ET2)
649	ELSE
650	†1≖λ1 T2_¥2
652	IF(ZTZ.GT.CC2) THEN
653	XI2=ZT2
654	ET2=CC2
656	
657	ET2=ZT2
658	ENDIF
659	XI-SORT(XI2)
661	EI=SUKI(EIZ) AID_F1
662	BET=F0
663	PALP=F1
664	QALP=F1
666	ENUIF DK2_(XT2_ET2)/XT2
667	DK=SQRT(DKZ)
668	PH1=DATAN(T1/ET)
669	PH2=DATAN(T2/ET)
671 C	RETIION
672	END
673 C	
674 C	
676	SUBROUTINE ELLIPEECONO CO ELPE ELPEN
677	IMPLICIT REAL*B (A-H,O-Z)
678	DATA F0,F1,PI,EPS,INF/0.D0,1.D0,3.1415926535897932D0,
679	£ 2.22044604925031D-16,2.D16/
681	G=ABS(G0)
682	G2=G*G
683	PH-PH0
084 (695	TE(C E0 E1) THEN
686	ELPE = DSIN(PH)
687	IF(ABS(PH0).LT.PI) ELPF-DLOG((F1+DSIN(PH))/DCOS(PH))
688	IF(ABS(PH0).GE.PI) ELPF=INF
689	
030	

M. Utsugi

691 C		
692	49-F1	
602		
055	B0=D5QR1(+1-62)	
694	P=F1	
695	ARGEFO	
696	CE1 E1	
0.00		
697	CE2=FØ	
698 1	$100 A = 0.500^{\circ}(A0+B0)$	
600	B DECRT(AA*PA)	
0000	B = DSQRT(A0-B0)	
100	C = 0.500*(A0-B0)	
701	$(E1 = (E1 - (2, D0^{+0}P)) (A^{+}C)$	
792	TECHA NE DO DOL THEM	
702	17(PN0.NE. 90.D0) [HEN	
703	ARG = REAL(NINT(PH/PI))	
704	TECHLOT FO AND NOD CHIPTLA SDA FILL FO FOLLARC ADC.	
705		
705	FIT + FITHWATAN(DO-DTAN(PH)/AU)+PT*AKG	
706	CE2 = CE2+C*DSIN(PH)	
707	ENDIF	
70.9 C		
700 (
/09	IF(ABS(A-A0).LE.EPS.AND.ABS(B-B0).LE.EPS.	
710	& AND, ABS(A-B), LF, EPS) THEN	
711	(0T0 200	
111	0010 200	
/12	ELSE	
713	AQ - A	
714	PA - P	
745		
/15	P = P+FI	
716	GOTO 100	
717	ENDIE	
710 -	L1021	
718 C		
719 2	200 IF(PH0,NE.PI) AGMPH = PH/(2.D0**P)	
770	TECONA EO DEL ACADA - DE 12 DA	
7774	11(FIN:EQ:FI) AUMENT = FI72.00	
(21	ELPE = AGMPH/A	
722	ELPE = CE1*AGMPH/A+CE2	
773 3	300 PETIEN	
724	END	

•