brought to you by CORE

1	
2	Running Title: Effect of CO ₂ sequestration on foraminifera
3	
4	
5	Impact of intentionally injected carbon dioxide hydrate on
6	deep-sea benthic foraminiferal survival
7	
8	
9	
10	Joan M. Bernhard ¹ , James P. Barry ² , Kurt R. Buck ² , Victoria R. Starczak ³
11	
12	¹ Geology and Geophysics Department, MS #52, Woods Hole Oceanographic Institution, Woods
13	Hole, MA 20543 USA
14	² Monterey Bay Aquarium Research Institute, 7700 Sandholdt Road, Moss Landing, CA 95039
15	USA
16	³ Biology Department, MS #50, Woods Hole Oceanographic Institution, Woods Hole, MA
17	20543 USA
18	
19	Corresponding Author: Joan M. Bernhard; Tel: 1 508 289 3480; Fax: 1 508 457 2076;
20	jbernhard@whoi.edu
21	
22	Keywords: carbon dioxide sequestration, CO ₂ injection, climate change, foraminifera,
23	experiment, hypercapnia, meiofauna, Monterey Bay, ocean acidification, protist

24 Abstract

25 Sequestration of carbon dioxide (CO_2) in the ocean is being considered as a feasible 26 mechanism to mitigate the alarming rate in its atmospheric rise. Little is known, however, about 27 how the resulting hypercapnia and ocean acidification may affect marine fauna. In an effort to 28 understand better the protistan reaction to such an environmental perturbation, the survivorship 29 of benthic foraminifera, which is a prevalent group of protists, was studied in response to deep-30 sea CO₂ release. The survival response of calcareous, agglutinated, and thecate foraminifera was determined in two experiments at ~3.1 and 3.3 km water depth in Monterey Bay (California, 31 32 USA). Approximately five weeks after initial seafloor CO₂ release, in situ incubations of the live-dead indicator CellTracker Green were executed within seafloor-emplaced pushcores. 33 Experimental treatments included direct exposure to CO₂ hydrate, two levels of lesser exposure 34 35 adjacent to CO₂ hydrate, and controls, which were far removed from the CO₂ hydrate release. 36 Results indicate that survivorship rates of agglutinated and thecate foraminifera were not 37 significantly impacted by direct exposure but the survivorship of calcareous foraminifera was 38 significantly lower in direct exposure treatments compared to controls. Observations suggest 39 that, if large scale CO₂ sequestration is enacted on the deep-sea floor, survival of two major 40 groups of this prevalent protistan taxon will likely not be severely impacted, while calcareous 41 foraminifera will face considerable challenges to maintain their benthic populations in areas 42 directly exposed to CO₂ hydrate.

Introduction

44 In both the scientific and public arenas, much attention has recently focused on the issues 45 of global warming and climate change. Although debate over the most appropriate mitigation 46 pathway continues, it is certain that the atmospheric concentration of carbon dioxide (CO_2) has 47 increased significantly in the recent past (e.g., Keeling et al., 1995; Keeling 1998). One option 48 to curtail the rapidly rising atmospheric CO_2 levels under consideration is to sequester waste CO_2 49 in the deep ocean (e.g., Caldeira, Akai et al., 2005). A number of scenarios have been proposed 50 for such ocean carbon storage, including fertilization of the sea surface with iron to promote 51 phytoplankton growth and accelerate the biological pump, thereby increasing dissolved inorganic 52 carbon (DIC) export to the deep sea (e.g., Buesseler et al., 2004), CO₂ injection at mid-ocean 53 water depths (e.g., Ozaki, 1997), and CO₂ injection onto the deep-sea floor (e.g., Brewer et al., 54 2000). Ongoing research is attempting to elucidate the benefits and drawbacks of each approach 55 if implemented on the large scale. In particular, impacts on both ocean chemistry and inhabitants 56 must be ascertained, especially given that CO₂ dissolution causes a concomitant decrease in pH 57 (Brewer et al., 2000). Physiological responses to elevated CO₂, or hypercapnia, and ocean 58 acidification are challenging to organisms in general (reviewed in, e.g., Siebel & Walsh, 2003; 59 Pörtner *et al.*, 2004); it is unclear which taxa, if any, will be unaffected by these environmental 60 pressures.

A series of in situ experiments have been conducted to assess the effects of direct
injection of CO₂ on the seafloor, concentrating on the lower bathyal zone since it theoretically
provides a longer period of sequestration compared to shallow water (Brewer *et al.*, 1999, 2000;
Barry *et al.*, 2005). Although studies have begun to describe the effects of such CO₂ disposal on
deep-sea fauna including fish (Tamburri *et al.*, 2000), crustaceans (Barry *et al.*, *submitted*),

66 echinoderms (Barry et al., submitted), and metazoan meiofauna (e.g., Carman et al., 2004; 67 Watanabe et al., 2006; Fleeger et al., 2006; Thistle et al., 2007), little is known about the effect 68 of such activities on protista. Because protists comprise a substantial portion of the deep-sea 69 benthos (e.g., Alongi & Pinchon, 1988; Coull et al., 1977; Snider et al., 1984; Gooday et al., 70 2000; Smith *et al.*, 2002), it is important to establish the effects of bathyal CO₂ release on these 71 single-celled eukaryotes. One study that surveyed the effects of CO_2 release on deep-sea 72 meiobenthos observed that 4.5 weeks after seafloor CO₂ injection many meiofaunal groups (i.e., 73 nematode metazoans and flagellate and amoebae protists) experienced elevated mortality 74 compared to sites removed from CO₂ manipulation (Barry *et al.*, 2004). Robust conclusions 75 about the response of some meiofaunal protists (i.e., allogromiid [thecate] foraminifera and 76 ciliates) to that increased pCO₂ exposure could not be drawn, however, due to low population 77 sizes (Barry et al., 2004) and because an accurate means to determine survival was not 78 implemented for those taxa.

79 In a continuing effort to better ascertain the response of meiofaunal protists to elevated 80 CO₂ exposure, a study was executed to determine in situ survival response of deep-sea benthic 81 foraminifera to CO₂ release simulating seafloor disposal of this greenhouse gas. For the study, 82 two experiments were conducted, each with a duration of 4.5 weeks in order to evaluate faunal 83 responses to quasi-chronic changes in ocean chemistry, rather than ephemeral changes over short 84 periods (i.e., days). Foraminifera were selected as the study taxon for two main reasons. First, 85 they are a critical link in marine food webs (e.g., Legendre & Le Févre, 1995; van Oevelen et al., 86 2006; Rowe *et al.*, *in press*). Second, while a large proportion of foraminiferal species secrete 87 calcium carbonate shells called tests, the majority lack inorganic tests (so-called allogromiid or 88 thecate forms, e.g., Gooday, 2002) or use detrital particles to construct agglutinated tests (Sen

Gupta, 1999), allowing comparative experimentation between calcifying and non-calcifying species of the same taxonomic group of Rhizarian protists (Adl *et al.*, 2005). This diversity in foraminiferal test composition is particularly useful when considering the biologic effects of ocean acidification. It may be hypothesized that the mortality of calcareous foraminiferal species will be higher in response to deep-sea CO₂ sequestration than the mortality of species of the two non-carbonate foraminiferan groups (i.e., thecate and agglutinated taxa).

95 Materials and Methods

96 Two replicate experiments were conducted on the seafloor of Monterey Bay, off the 97 California coast (USA). Using the surface support vessel RV Western Flyer and the ROV 98 Tiburon for each experiment, a set of 40.6-cm diameter PVC cylinders were placed on the 99 seafloor so that ~15 cm extended above the sediment-water interface; 15 cylinders (Experiment 100 1) and 7 cylinders (Experiment 2) were configured in a circle with a diameter of ~ 20 m. Then, 101 over two to three days, liquid CO₂ was injected into each cylinder using methods described in, 102 e.g., Barry et al. (2004), Carman et al. (2004), and Fleeger et al. (2006). CO₂ hydrate formed 103 immediately as a 'skin' on liquid CO₂ pools. The targeted pH decline within the \sim 20-m diameter 104 circle compared to in situ pH (~7.8; Thistle et al., 2007) was 0.2 pH units because data suggests 105 that pH declines of that magnitude can be an important physiological threshold (e.g., Seibel & 106 Walsh, 2003).

107 For both experiments, Conductivity-Temperature-Depth instruments (SeaBird Model 19+ 108 CTDs) were deployed on the seabed at locations near the margin, ~3-5 m from the margin, and 109 in the center of the circle of CO₂ cylinders (Fig. 1). Each CTD was equipped with up to 4 110 Seabird Model SBE 18 pH sensors, positioned from 3 to 50 cm above the seabed. These 111 instruments collected data at ~ 2 minute intervals throughout each experiment. Perturbations in 112 the pH of surficial sediments were also measured during Experiment 2 within CO₂ cylinders and 113 from 0.1 - 8 m from CO₂ pools, as reported in Barry *et al.* (*submitted*). 114 During the commencement of Experiment 1, a suite of control samples was obtained, as

explained below. The first experiment was initiated from 13-17 December 2004 at a water depth
of 3088 m (36.6985°N, 123.0020°W) and terminated from 18-21 January 2005. Experiment 2

117 was initiated from 12-16 December 2005 at a water depth of 3266 m (35°48.6105'N,

119 region of the experimental sites is typical of the deep Eastern Pacific (WOCE, 120 www.nodc.noaa.gov/WOCE; Sector p17N, St. 10, 38.23733N, 124.93833W, 3087 m depth; 121 DIC=2352.7; TA=2442.3; pH_{tot}=7.7834; Omega_{CA}=0.93; Omega_{AR}=0.6). 122 As part of each experiment's termination, suites of samples were obtained from a site 123 with no increase in CO_2 (i.e., ~100 m from CO_2 release), from inside the circle of PVC cylinders, 124 and from inside individual PVC cylinders. These three treatments are hereby referred to as 125 Control, Elevated, and Direct Contact, respectively. Elevated treatment samples were further 126 distinguished depending on their distance from the PVC cylinders due to experimental logistics 127 (see below): Elevated Center samples were collected from the approximate middle of the PVC-128 cylinder circle while Elevated Edge samples were obtained close to PVC cylinders (i.e., within 129 \sim 1-3 m; Fig. 1). The configuration of core locations was mandated by the fact that this 130 experiment included sampling for multiple additional purposes and by the logistical 131 considerations of working in the deep sea. For example, the ROV manipulator reach and 132 seafloor disturbance in the form of sediment resuspension were critical factors for core 133 placement. More details on the experimental design and concurrent science objectives are 134 presented in Ricketts et al. (2005; submitted) and Thistle et al. (2007). 135 The fluorogenic probe CellTracker™ Green CMFDA (5-chloromethylfluorescein 136 diacetate; Invitrogen, hereafter referred to as CellTracker Green) was used to distinguish living 137 from dead foraminifera (Bernhard et al., 2006) and, thus, establish survivorship. A fluorogenic 138 probe is a non-fluorescent compound that yields a fluorescent product after structural 139 modification; CellTracker Green accumulates intracellularly in live cells after enzymatic

122°34.0952' W) and terminated from 18-21 January 2006. The carbonate chemistry in the

118

hydrolysis cleaves the molecule, thereby producing fluorescence (reviewed in Bernhard *et al.*,141 1995).

142 Quantitative sediment samples were obtained using pushcores outfitted with a device to 143 allow injection of concentrated CellTracker Green after pushcore placement into the seafloor. 144 This approach was adopted to allow in situ seafloor incubations with CellTracker Green because 145 temperature and pressure changes during ascent could cause mortality due to sample recovery 146 rather than due to in situ experimental conditions. CellTracker Green is soluble in dimethyl 147 sulfoxide (DMSO); this stock solution was diluted with an equal volume of 0.2 µm-filtered 148 seawater (FSW) to prevent freezing in the injector capillary tubing prior to dispensation. 149 Initially, the pushcores (7-cm inner diameter) were emplaced into sediments (Fig. 2a) so 150 that the internal header space volume was within a targeted range. To do this, the pushcorer was positioned so that it penetrated the seafloor to a depth placing the sediment-water interface 151 152 between two core-tube markings. The ROV manipulator arm was then used to carefully squeeze 153 a large diameter (2-cm outer diameter) flexible tube filled with FSW, which flowed through a 154 check valve to displace the concentrated CellTracker Green from the capillary tubing producing 155 a final concentration of $\sim 1 \,\mu$ M CellTracker Green within each core's header space. These cores 156 incubated in situ until the following day, when they were recovered by the ROV. A negative 157 control core, which was injected with DMSO and FSW but not CellTracker Green, was similarly 158 incubated and collected from an area with no increase in CO₂, during Experiment 2. To maintain 159 the integrity of the sediment-water interface, the pushcores were not equipped with a stirring 160 device.

161 After ROV recovery by the support ship, the incubated pushcores were taken within ~ 10 162 minutes into an environmental room approximating in situ temperature (~ 5 °C). Within an hour,

the overlying water containing CellTracker Green was removed via siphon; the surface 1-cm of each core was sectioned from the underlying sediments, placed in a high density polyethylene (HDPE) container, and fixed in ~3.8% formalin buffered with sodium borate. All Direct Contact treatment cores became disturbed upon ascent due to degassing. In these instances, fine particles were suspended and the sediment-water interface was typically disturbed. Thus, for these cores, the overlying waters were retained in addition to the surface sediment interval or allowed to settle for approximately an hour prior to sectioning.

170 In the shore-based laboratory, samples were sieved over a 63-µm screen with tap water, 171 and the coarser fraction examined with a Leica MZ FLIII stereo dissecting microscope equipped 172 with appropriate epifluorescence optics (excitation 480 ± 20 nm, emission ≥ 510 nm). Fluorescent 173 foraminifera were isolated, segregated into brightly and dimly fluorescent groups, sorted by 174 species within those groups, and enumerated. Thecate and "soft-shelled" agglutinated taxa, 175 which were not identified to species, were archived in buffered formalin while calcareous and 176 robust agglutinated taxa were air dried and archived on micropaleontology slides. Because a 177 single individual of observed unilocular cylindrical agglutinated foraminiferal taxa (i.e., 178 Bathysiphon, Rhabdammina, Rhizammina) can easily break, those fragments were minimally 179 enumerated to prevent over estimates of population density. Within a sample, it is typically clear 180 which fragments likely arose from the same specimen due to test grain size, diameter, 181 composition, texture, and color. During the picking process, the treatment of each sample (i.e., 182 Control, Elevated, Direct Contact, negative control) was withheld from the microscopist. 183 Species counts were normalized to in situ sample volume to provide abundance estimates. 184 Abundance data from each experiment were statistically analyzed separately for each 185 taxonomic group. For thecate and agglutinated groups in Experiment 1, a one-way ANOVA was

used to test for differences in density among the Control, Elevated Center, Elevated Edge, and
Direct Contact treatments. For Experiment 2, a one-way ANOVA was conducted to test whether
mean density of either group differed between the Control, Elevated Edge, and Direct Contact
treatments. To homogenize variances, analysis was done on logged transformed thecate data and
on square root agglutinated data.

191 To determine the effects of CO₂ injection on calcareous foraminiferal abundance, the 192 mean abundance in the Control treatment was compared to the mean abundance in the Direct 193 Contract treatment. For this comparison, the mean of the Control treatment was compared to a 194 value of 0 under the null hypothesis with a one sample t-test because calcareous for a minifera 195 were absent in the Direct contact treatment cores. Additionally, a one-way ANOVA was used to 196 test whether mean densities of calcareous foraminifera differed between Control, Elevated 197 Center and Elevated Edge treatments. For Experiment 2, mean calcareous foraminiferal 198 densities in the Elevated Edge and Controls were compared with a two sample t-test and the 199 mean density of the Control treatment was tested in a single sample t-test with a hypothesized 200 mean of zero.

201 Results

202 Only cores that were visibly injected (Fig. 2a) were processed and analyzed. In one case, 203 the capillary tubing parted from a coupling, thus CellTracker Green was not injected into that 204 core so it was excluded from study. Replicate cores from Control and Elevated treatments of 205 each experiment were typically taken within 1 m of each other (Fig. 2b); Direct Contact cores 206 were obtained from different PVC cylinders. Careful manipulation by the ROV pilots resulted in 207 absence of visible sediment-water interface disturbance within the corers, except for those 208 collected from within the PVC cylinders (i.e., Direct Contact cores), which degassed during 209 ascent as noted above.

The negative control core (i.e., injected solely with DMSO, without CellTracker Green) lacked brightly fluorescent foraminifera, but had dimly fluorescent foraminifera. Thus, only brightly fluorescent foraminifera were considered to be living at the time of seafloor incubation with CellTracker Green. In CellTracker Green-incubated cores, the CellTracker Green diffused to at least 1-cm depth because some organisms in the 1-2 cm interval fluoresced brightly.

215 *pH changes*

216 The perturbations of pH measured in bottom waters during both experiments were 217 variable but not large. In an area equivalent to the Elevated Edge treatment, the maximum pH 218 reductions compared to ambient [pH reductions were calculated as perturbations from the 219 background pH measured at the site] values were 0.15 to 0.25 pH units, with average changes 220 less than 0.05 units. In the center of the circle (i.e., Elevated Center treatment), pH reductions 221 sometimes approached 0.1 - 0.2 pH units, but typical pH perturbations were less than 0.05 units. 222 In areas corresponding to the Elevated Edge treatment, pore-waters of surface sediments 223 at the beginning of Experiment 2 (i.e., 1-3 d after hydrate placement) were ~0.2 units lower than

at control locations (~50 m away) but the pH of Elevated Edge pore-waters at experimental
termination did not differ from the pore-water pH of control area sediments (Barry *et al.*, *submitted*). Changes in the pH of surficial sediment pore waters (i.e., upper ~10 cm) were large
within Direct Contact cylinders (~ -2.0 pH units compared to pore-water of control sites; Barry *et al.*, *submitted*).

229 Species composition of living foraminiferal assemblages

230 The living calcareous foraminiferal assemblage, as determined by bright CellTracker 231 Green labeling, in the Control and Elevated samples was dominated by rotaliid forms including 232 Uvigerina canariensis, Chilostomella oolina, and Globobulimina pacifica; the most common 233 miliolid form was *Quinqueloculina venusta* (Table 1; Fig. 3). The living non-calcareous 234 foraminiferal assemblage in Control and Elevated samples was dominated by thecate and "soft-235 shelled" agglutinated forms such as saccamminids, as well as Reophax dentaliniformis, R. 236 spiculifer, Hormosinella guttifera, and Paratrochammina challengeri (Fig. 4). Agglutinated 237 species such as *R. dentaliniformis* and *R. spiculifer* and thecate species dominated Direct Contact 238 samples.

239 Abundances of live foraminifera

The foraminiferal abundance between cores within a treatment was highly variable for some treatments but not for others (Table 1). Abundances of all live foraminifera ranged from 4.4 to 5.7 specimens·10cm³ in Initial Control cores of Experiment 1, and from 6.2 to 11.2 specimens·10cm³ in Control samples collected at the end of either experiment (Supplemental Data; Fig. 5). Thus, the average abundance of foraminifera in Control samples was 7.7 specimens·10cm³ (n = 8, SD = 3.0). Foraminiferal abundances within the Elevated samples spanned a wider range (2.6 to 26.0 specimens·10cm³; $\bar{x} = 12.1$, SD = 8.6) than those within

Direct Contact samples (1.3 to 11.7 specimens $\cdot 10$ cm³; $\bar{x} = 5.4$, SD = 4.0). The average of total abundance of Elevated Edge samples ($\bar{x} = 14.9$, SD = 10.4) was higher than all other treatments (Fig. 5a,b).

Abundances of the cate for a veraged 2.3 specimens 10 cm^3 (SD = 2.3) in all 250 Control cores and 3.3 specimens 10 cm³ both in Elevated samples and Direct Contact samples 251 252 (SD = 3.8, 2.5, respectively). Abundances of agglutinated foraminifera averaged 2.1-2.2 specimens 10 cm³ in both the Control and Direct Contact samples, but were nearly twice as 253 254 abundant in samples collected from within the circle (i.e., Elevated treatments combined; 4.1 specimens 10 cm^3 , SD = 3.8). The mean abundances of the cate and agglutinated groups were 255 256 generally higher in Elevated Edge samples than in other treatments (Fig. 5c-f), the exception 257 being thecate abundances for Experiment 2 (Fig. 5d). The mean density of thecate or 258 agglutinated foraminifera did not differ significantly between the Control, Elevated Center, Elevated Edge or Direct Contact treatments in Experiment 1 (One-way ANOVA, the cate, $F_{3,7}$ = 259 260 0.39, p = 0.762; agglutinates $F_{3,7} = 0.49$ p= 0.701) or in Experiment 2 between Control, Elevated Edge, and Direct Contact (One-way ANOVA, the cate, $F_{2,3} = 0.09$, p = 0.913; agglutinates, $F_{2,3} = 0.09$, F_{2 261 262 0.40, p = 0.699).

Calcareous foraminiferal abundances averaged 3.2 specimens·10cm³ and 4.7 specimens·10cm³ in Control and Elevated cores, respectively, but no calcareous foraminifera were living in the Direct Contact cores (Table 1; Fig. 5g,h). For Experiment 1, the mean density of calcareous foraminifera did not differ significantly between the Control, Elevated Center or Elevated Edge treatments (One-way ANOVA, $F_{2,5} = 0.48$, p=0.645) and, for Experiment 2, their mean density did not differ significantly between the Control and Elevated Edge treatments (Two sample t-test, -1/x transformed data, df = 2, t = 2.88, p = 0.102). In both experiments, the

- 270 mean density of calcareous foraminifera in the Control samples was significantly different from
- a mean of zero (One sample t-test, Experiment 1: t = 6.26, df = 2, p = 0.025, Experiment 2:
- t=25.96, df =1 p=0.025). Thus, there were significantly more calcareous foraminifera in the
- 273 Control treatments than in the Direct Contract treatment, which lacked calcareous foraminifera in
- every core of both experiments.

275 Discussion

The hypothesis that higher mortality would occur in calcareous foraminifera compared to thecate and agglutinated foraminifera in response to deep-sea CO_2 sequestration is supported by our data. Survivorship of the thecate and agglutinated foraminiferal populations at our Monterey Bay sites were not significantly affected by direct exposure to CO_2 hydrate on the experimental time scale.

281 Although the water depth at our experimental sites likely exceeded the regional calcite 282 saturation depth and the carbonate compensation depth (CCD), calcareous (calcitic) foraminifera 283 were living in the area as evidenced by Control treatment results. The existence of calcitic 284 foraminifera is not unexpected since most abyssal regions of today's oceans have live calcareous 285 foraminifera (e.g., Bernhard, 1992; Linke & Lutze, 1993). In general, the observed abundances 286 of foraminifera in control samples were comparable to abundances from similar water depths at 287 sites ~150 km to the south (Bernhard, 1992). Using a different viability method (i.e., Adenosine 288 Triphosphate (ATP) assay), abundances of live foraminifera integrated over the surface 1 cm ranged from 0 to ~12 specimens $\cdot 10$ cm³ in water depths from 3319-3728 m ($\overline{x} \sim 6$ 289 specimens $\cdot 10$ cm³; n = 4; Bernhard, 1992). Deep-sea benthic foraminiferal distributions are 290 291 known to be patchy on the scale of km to cm (e.g., Bernstein *et al.*, 1978; Bernstein & Meador, 292 1979), so it is not unusual to document considerable variations in foraminiferal abundance over 293 short distances. Indeed, Bernhard (1992) also noted large variations in foraminiferal abundances 294 between sites, although those cores were separated by kilometers, not meters, as in this study. 295 The species compositions of agglutinated and calcareous assemblages encountered in our 296 samples resemble those previously reported from the region at comparable water depths (e.g., 297 Bernhard, 1992).

298 Even though some bathyal and abyssal foraminiferal species' abundances exhibit 299 subsurface maxima (e.g., Corliss, 1985; Bernhard, 1992), a down-core analysis of our samples 300 was not feasible. Because, in general, the majority of bathyal foraminifera live in the top cm 301 (e.g., ~52-71%, Gooday, 1986; ~80%, Szarek et al., 2007) and because we expect the maximum 302 physical and chemical changes due to CO_2 hydrate release within the surface sediments, we feel 303 our data reflects the typical survival response of bathyal benthic foraminifera to such 304 environmental perturbations. It is possible that foraminiferans migrated in response to the 305 changing milieu, as observed in other experimental studies (e.g., Alve & Bernhard, 1995; 306 Moodley et al., 1998). In particular, if infaunal specimens migrated upward into the surface cm, 307 the average abundances in treatments affected by CO₂ would have been higher than those of 308 Controls. Indeed, the abundance data for agglutinated taxa were consistently higher, although 309 not significantly different, in Elevated Edge samples compared to Controls (Fig. 5e,f). This 310 trend was not evident for other groups (Fig. 5c,d,g,h). Why foraminifera may have migrated 311 toward potentially stressful concentrations of CO₂ is unclear, except the option of migrating 312 deeper into sediments would have been more stressful for aerobic migrants compared to 313 facultative anaerobes, which are known for foraminifera (e.g., Bernhard & Alve, 1996; Moodley 314 et al., 1998), or if the migrants were incapable of performing complete denitrification (Risgaard-315 Petersen et al., 2006; Høgslund et al., 2008). Recent findings suggest elevated pCO₂ exposure is 316 stressful to at least some meiofauna because higher numbers of harpacticoid copepods emerged 317 from sediments in response to elevated pCO₂ compared to copepod emergence rates at control 318 sites (Thistle et al., 2007).

Degassing of cores collected from PVC cylinders during ascent may have minimized the
 observed foraminiferal abundances in the Direct Contact cores. We discount this possibility

because the average density of thecate and agglutinated morphotypes in Direct Contact samples was near or equal to their density in Control cores. If significant abundance dilution occurred due to degassing disturbance, we would expect all three groups of foraminifera to be similarly affected, which was not the case.

325 Observed patterns and magnitudes of pH change resemble those reported for similar 326 experiments (Barry et al., 2005). Importantly, the observed bottom-water pH decreases in areas 327 corresponding to Elevated Center treatments (~ 0.05 units for Experiment 1) were small 328 compared to those declines corresponding to Elevated Edge treatments (initially 0.2 pH units, but 329 ~ 0.1 to 0.15 pH units at experiment end). Thus, it is not surprising that all three groups of 330 benthic foraminifera (i.e., calcareous, agglutinated, thecate) tolerated the Elevated Center 331 treatments because those sediments apparently did not experience extended large decreases in pH 332 and, by inference, extended large increases in pCO_2 . It is also noteworthy, however, that thecate 333 and agglutinated foraminiferal survival was not negatively impacted by the substantial pH 334 decrease (> 1 pH unit) in the Direct Contact treatments.

335 Economically viable protocols of deep-sea CO₂ sequestration include formation of deep-336 sea CO₂ lakes and near-bottom injection of CO₂ (reviewed in Herzog *et al.*, 2000; Caldeira & 337 Akai *et al.*, 2005). In all cases, a plume of CO_2 -rich seawater will lead from liquid CO_2 or CO_2 338 hydrate drifting with currents and diffuse, eventually being diluted (Caldeira and Akai et al., 339 2005). Thus, CO₂ and pH gradients will range from substantial (e.g., the lowest pH measured is 340 4.5, a value measured within cm of CO₂ hydrate; Brewer *et al.*, 2000) to mild (i.e., <-0.05 pH 341 units) at some distance, depending on specific mixing conditions. Our Direct Contact treatment 342 recreates the CO₂ lake approach, although on a much smaller scale. Benthos beneath any CO₂-343 hydrate pool will be directly exposed for long periods of time unless they can migrate from the

hydrate pool (Tamburri *et al.*, 2000). Deep-sea foraminifera move very slowly (<25 μm·min⁻¹;
Gross, 2000), so active avoidance at these spatial scales is an implausible escape mechanism for
them. The pH changes in our experiments are similar to those under any carbon sequestration
option: conditions of our Elevated treatments are within the magnitude of pH change expected
along the dilution gradient resultant from presently proposed mitigation scenarios.

349 The in situ response of foraminifera (and any other group of deep-sea organisms) to such 350 CO₂-hydrate exposure is unknown over time periods longer than our experiments. Our results 351 indicate that although the calcareous for a population tolerates short exposures to a 0.2 352 pH unit decrease, their survival response to direct CO₂-hydrate exposure results in a significant 353 increase in mortality. If, however, plumes of dissolved CO₂ hydrate are diffused from the 354 seafloor over a shorter time scale than our experiments, it is possible that calcareous 355 foraminiferal survival will not be significantly impacted by large-scale CO₂ sequestration. 356 Although the effects of localized deep-sea CO₂ sequestration appear to negatively impact 357 calcareous foraminifera (Ricketts et al., 2005; submitted; this study), two other protistan groups 358 (i.e., flagellates and amoebae; Barry et al., 2004), and a number of metazoan meiofauna (Barry et 359 al., 2004; Carman et al., 2004; Watanabe et al., 2006; Fleeger et al., 2006; Thistle et al., 2007),

360 our findings suggest that survival of at least some protistan meiofauna (i.e., thecate and

361 agglutinated foraminifera) are not similarly influenced by direct exposure to CO₂ hydrate.

362 Because the abundance and diversity of thecate and agglutinated foraminifers are considerable in

bathyal and abyssal sediments (e.g., Gooday et al., 1998; 2000; Smith et al., 2002) and their

abundance can exceed that of other meiofauna (e.g., Gooday et al 2000), at least one major group

365 of deep-sea meiofauna will likely not collapse if large-scale sequestration of CO₂ is implemented

366 on the deep-ocean floor. Furthermore, our data showing that thecate and agglutinated

- 367 for aminiferal abundances do not significantly decline in response to direct exposure to CO₂
- 368 hydrate at these spatial and temporal scales substantiate recent laboratory findings documenting a
- 369 shallow-water thecate for a species survives ~2-week exposure to extremely high pCO₂
- 370 (200,000 ppm) where some specimens even reproduced (Bernhard et al., *in press*).

371 Acknowledgements

Appreciation is extended to Eric Pane for synopsis of pore-water pH data; Craig Okuda for designing the push-core injectors; Patrick Whaling, Chris Lovera, Linda Kuhnz, Kevin Carman, John Fleeger, David Thistle, and Erin Ricketts for sampling assistance; the Captain and crew of the RV *Western Flyer* and the ROV *Tiburon*; Marti Jeglinski for laboratory assistance; David Smith, Subject Editor and the two anonymous reviewers for their comments on a previous manuscript version. This work was funded by the Monterey Bay Aquarium Research Institute (project 200002; to JPB), US Department of Energy grant # DE-FG02-03ER63696 (to J. P.

379 Kennett and J.M.B.), and NSF OCE-0725966 (to J.M.B.).

380 **References**

- 381 Adl SM, Simpson AGB, Farmer MA, Andersen RA, Anderson OR, Barta JR, Bowser SS,
- 382 Brugerolle G, Fensome RA, Fredericq S, James TY, Karpov S, Kugrens P, Krug J, Lane
- 383 CE, Lewis LA, Lodge J, Lynn DH, Mann DG, McCourt RM, Mendoza L, Moestrup O,
- 384 Mozley-Standridge SE, Nerad, TA, Shearer CA, Smirnov AV, Spiegel FW, Taylor MFJR
- 385 (2005) The new higher level classification of eukaryotes with emphasis on the taxonomy
 386 of protists. *Journal of Eukaryotic Microbiology*, **52**, 399-459.
- 387 Alongi DM, Pinchon M (1988) Bathyal meiobenthos of the western Coral Sea: distribution and
- abundance in relation to microbial standing stocks and environmental factors. *Deep-Sea Research*, **35**, 491-503.
- Alve E, Bernhard JM (1995) Vertical migratory response of benthic foraminifera to controlled
 decreasing oxygen concentrations in an experimental mesocosm. *Marine Ecology Progress Series*, **116**, 137-151.
- 393 Barry JP, Buck KR, Lovera CF, Kuhnz L, Whaling PJ, Peltzer ET, Walz P, Brewer PG (2004)
- 394 Effects of direct ocean CO₂ injection on deep-sea meiofauna. *Journal of Oceanography*,
 395 **60**, 759-766.
- Barry JP, Buck KR, Lovera C, Kuhnz L, Whaling PJ (2005) Utility of deep sea CO₂ release
- experiments in understanding the biology of a high-CO₂ ocean: Effects of hypercapnia on
 deep sea meiofauna. *Journal of Geophysical Research—Oceans*, **110**, Issue: C9 Article
 Number: C09S12.
- 400 Barry JP, Buck KR, Lovera C, Brewer PG, Seibel BA, Drazen JC, Tamburri MN, Whaling PJ,
- 401 Kuhnz L, Pane E (*submitted*) Sensitivity of deep-sea animals to a high-CO₂ ocean.
- 402 *Marine Ecology Progress Series.*

403	Bernhard JM (1992) Benthic foraminiferal distribution and biomass related to pore-water
404	oxygen content: Central California Continental Slope and Rise. Deep-Sea Research, 39:
405	585-605.
406	Bernhard JM, Alve E (1996) Survival, ATP pool, and ultrastructural characterization of benthic
407	foraminifera from Drammensfjord (Norway): response to anoxia. Marine
408	Micropaleontology, 28, 5-17.
409	Bernhard JM, Newkirk SG, Bowser SS (1995) Towards a non-terminal viability assay for
410	foraminiferan protists. Journal of Eukaryotic Microbiology, 42, 357-367.
411	Bernhard JM, Ostermann DR, Williams DS, Blanks JK (2006) Comparison of two methods to
412	identify live benthic foraminifera: a test between Rose Bengal and CellTracker Green
413	with implications for stable isotope paleoreconstructions. <i>Paleoceanography</i> , 21 ,
414	PA4210, doi:10.1029/2006PA001290.
415	Bernhard JM, Mollo-Christensen E, Eisenkolb N, Starczak VR (in press) Tolerance of
416	allogromiid foraminifera to severely elevated carbon dioxide concentrations: Implications
417	to future ecosystem functioning and paleoceanographic interpretations. Global and
418	Planetary Change.
419	Bernstein BB, Meador JP (1979) Temporal persistence of biological patch structure in an
420	abyssal benthic community. Marine Biology, 51, 179-183.
421	Bernstein BB, Hessler RR, Smith CR, Jumars PA (1978) Spatial distribution of benthic
422	foraminifera in the abyssal central North Pacific. Limnology and Oceanography, 23,
423	401-416.
424	Brewer PG, Friederich G, Peltzer ET, Orr FM (1999) Direct experiments on the ocean disposal
425	of fossil fuel CO ₂ . Science, 284 , 943-945.

426	Brewer PG, Peltzer ET, Friederich G, Aya I, Yamane K (2000) Experiments on the ocean
427	sequestration of fossil fuel CO2: pH measurements and hydrate formation. Marine
428	<i>Chemistry</i> , 72 , 83-93.
429	Buesseler KO, Andrews JE, Pike SM, Charette MA (2004) The effects of iron fertilization on
430	carbon sequestration in the Southern Ocean. Science, 304, 414-417.
431	Caldeira, K, Akai M, et al. (2005) Ocean Storage. In: IPCC Special Report on Carbon Dioxide
432	Capture and Storage, B. Metz, O. Davidson, H. de Coninck, M. Loos, L. Meyer (Eds),
433	Cambridge U. Press, Cambridge.
434	Carman KR, Thistle D, Fleeger JW, Barry JP (2004) Influence of introduced CO ₂ on deep-sea
435	metazoan meiofauna. Journal of Oceanography, 60, 767-772.
436	Corliss BH (1985) Microhabitats of benthic foraminifera within deep-sea sediments. Nature, 314,
437	435-438.
438	Coull BC, Ellison RL, Fleeger JW, Higgins RP, Hope WD, Hummon WD, Rieger RM, Sterrer
439	WE, Thiel H, Tietjen JH (1977) Quantitative estimates of the meiofauna from the deep
440	sea off North Carolina, USA. Marine Biology, 39, 233-240.
441	Fleeger JW, Carman KR, Welsenhorn PB, Sofranko H, Marshall T, Thistle D, Barry JP (2006)
442	Simulated sequestration of anthropogenic carbon dioxide at a deep-sea site: Effects on
443	nematode abundance and biovolume. Deep-Sea Research I, 53, 1135-1147.
444	Gooday AJ (1986) Meiofaunal foraminiferans from the bathyal Porcupine Seabight (northeast
445	Atlantic): size structure, standing stock, taxonomic composition, species diversity and
446	vertical distribution in the sediment. Deep-Sea Research, 33, 1345-1373.
447	Gooday AJ (2002) Organic-walled allogromiids: aspects of their occurrence, diversity and
448	ecology in marine habitats. Journal of Foraminiferal Research, 32, 384-399.

449	Gooday AJ, Bett BJ, Shires R, Lambshead PJD (1998) Deep-sea benthic foraminiferal species
450	diversity in the NE Atlantic and NW Arabian sea: a synthesis. Deep-Sea Research II, 45,
451	165-201.
452	Gooday AJ, Bernhard JM, Levin LA, Suhr S (2000) Foraminifera in the Arabian Sea oxygen
453	minimum zone and other oxygen-deficient settings: taxonomic composition, diversity,
454	and relation to metazoan faunas. Deep-Sea Research II, 47, 25-54.
455	Gross O (2000) Influence of temperature, oxygen and food availability on the migrational
456	activity of bathyal benthic foraminifera: evidence by microcosm experiments.
457	Hydrobiologia, 426 , 123-137.
458	Herzog H, Eliasson B, Kaarstad O (2000) Capturing greenhouse gases. Scientific American, 282,
459	72-79.
460	Høgslund S, Revsbech NP, Cedhagen T, Nielsen LP, Gallardo VA (2008) Denitrification, nitrate
461	turnover, and aerobic respiration by benthic foraminiferans in the oxygen minimum zone
462	off Chile. Journal of Experimental Marine Biology and Ecology, 359, 85-91.
463	Keeling CD (1998) Rewards and penalties of monitoring the earth. Annual Review of Energy
464	and the Environment, 23, 25-82.
465	Keeling CD, Whorf TP, Wahlen M, Van der Plicht J (1995) Interannual extremes in the rate of
466	rise of atmospheric carbon dioxide since 1980. Nature, 375, 666-670.
467	Legendre L, Le Févre J (1995) Microbial food webs and the export of biogenic carbon in oceans.
468	Aquatic Microbial Ecology, 9, 69-77.
469	Linke P, Lutze GF (1993) Microhabitat preferences of benthic foraminifera—a static concept or

470 a dynamic adaptation to optimize food acquisition? *Marine Micropaleontology*, 20, 215471 234.

472	Moodley L, van der Zwaan GJ, Rutten GMW, Boom RCE, Kempers AJ (1998) Subsurface
473	activity of benthic foraminifera in relation to porewater oxygen content: laboratory
474	experiments. Marine Micropaleontology, 34, 91-106.
475	Ozaki M (1997) CO ₂ injection and dispersion in mid-ocean depth by moving ship. Waste
476	Management, 17, 369-373.
477	Pörtner HO, Langenbuch M, Reipschlager A (2004) Biological impact of elevated ocean CO ₂
478	concentrations: Lessons from animal physiology and earth history. Journal of
479	<i>Oceanography</i> , 60 , 705-718.
480	Ricketts ER, Kennett JP, Hill TM, Barry JP (2005) Effects of CO ₂ hydrate on deep-sea
481	foraminiferal assemblages. Proceedings of the Fifth International Conference on Gas
482	Hydrates, Trondheim, Norway, 3, (3020): 839-847.
483	Ricketts ER, Kennett JP, Hill TM, Barry JP (submitted) Effects of CO ₂ hydrate emplacement on
484	deep-sea foraminiferal assemblages: 3600 m on the California Margin. Marine
485	Micropaleontology.
486	Risgaard-Petersen N, Langezaal AM, Ingvardsen S, et al. (2006) Evidence for complete
487	denitrification in a benthic foraminifer. Nature, 443, 93-96.
488	Rowe GT, Wei C, Nunnally C, et al. (in press) Comparative Structure and Dynamics of Food
489	Webs in the Deep Gulf of Mexico. Deep-Sea Research II.
490	Sen Gupta BK (1999) Modern Foraminifera. Kluwer Academic Publishers, Dordrecht, The
491	Netherlands.
492	Seibel BA, Walsh PJ (2003) Biological impacts of deep-sea carbon dioxide injection inferred
493	from indices of physiological performance. Journal of Experimental Biology, 206, 641-
494	650.

495	Smith KL Jr, Baldwin RJ, Karl DM, Boetius A (2002) Benthic community responses to pulses in
496	pelagic food supply: North Pacific Subtropical Gyre. Deep-Sea Research I, 49, 971-990.
497	Snider LJ, Burnett BR, Hessler RR (1984) The composition and distribution of meiofauna and
498	nanobiota in a central North Pacific deep-sea area. Deep-Sea Research, 31, 1225-1249.
499	Szarek R, Nomaki H, Kitazato H (2007) Living deep-sea benthic foraminifera from the warm
500	and oxygen-depleted environment of the Sulu Sea. Deep-Sea Research II, 54, 145-176.
501	Tamburri MN, Peltzer ET , Friederich GE, Aya I, Yamane K, Brewer PG (2000) A field study of
502	the effects of CO ₂ ocean disposal on mobile deep-sea animals. <i>Marine Chemistry</i> , 72 ,
503	95-101.
504	Thistle D, Sedlacek L, Carman KR, Fleeger JW, Brewer PG, Barry JP (2007) Exposure to carbon
505	dioxide-rich seawater is stressful for some deep-sea species: an in situ, behavioral study.
506	Marine Ecology Progress Series, 340 , 9-16.
507	van Oevelen D, Soetaert K, Middelburg JJ, Herman PJM, Moodley L, Hamels I, Moens T, Heip
508	CHR (2006) Carbon flows through a benthic food web: Integrating biomass, isotope and
509	tracer data. Journal of Marine Research 64, 453-482.
510	Watanabe Y, Yamaguchi A, Ishidai H, et al. (2006) Lethality of increasing CO2 levels on deep-
511	sea copepods in the western North Pacific. Journal of Oceanography, 62, 185-196.

- 512 Table 1. Mean abundances of live foraminifera [(#·10cm³); listed for all foraminifera and by foraminiferal group] for each treatment,
- 513 presented by experiment. The number of replicates (# cores) is also listed. SD = Standard deviation.
- 514

515	Treatment	Experiment	# cores	Mean (SD)			
516				Total	Thecate	Agglutinated	Calcareous
517	Initial Control	1	3	4.8 (0.8)	0.9 (0.5)	1.6 (0.8)	2.3 (0)
518	Control	1	3	9.5 (2.9)	3.3 (2.8)	3.0 (0.8)	3.2 (1.8)
519		2	2	9.4 (1.8)	2.9 (3.3)	2.0 (0.9)	4.6 (0.5)
520	Elevated Center	1	3	8.3 (4.5)	2.4 (2.2)	3.0 (2.6)	2.9 (1.6)
521	Elevated Edge	1	2	14.3 (16.6)	6.5 (7.0)	6.1 (7.5)	1.7 (2.0)
522		2	2	15.6 (6.9)	1.4 (0.6)	3.6 (1.5)	10.5 (5.0)
523	Direct Contact	1	3	4.7 (1.7)	3.0 (0.9)	1.7 (1.4)	0
524		2	2	6.5 (7.4)	4.0 (4.6)	2.5 (2.8)	0

525 Figure Legends.

526 **Figure 1.** Schematic representation of experimental configurations (a. Experiment 1; b.

527	Experiment 2). The CO ₂ cylinders are shown as gray circles, and CTDs equipped with pH
528	sensors are shown as squares. Pushcores analyzed for this study are shown as smaller circles;
529	pushcore treatment categorizations are: open = Control; / = Initial Control, black = Direct
530	Contact, $+ =$ Elevated Center, $* =$ Elevated Edge; $\sim =$ DMSO.
531	

Figure 2. Photographs of injector pushcores placed in the seafloor. a. Representative pushcore from Elevated Center treatment immediately after the ROV manipulator released its grip on the thick injector tubing (upper left). Inset: Note the stream of whitish fluid (i.e., CellTracker Green and DMSO) entering the header space from the capillary tubing, which extends through the corer top, approximately to the arrow. b. Suite of three Control pushcores incubating in situ during Experiment 1.

538

Figure 3. Reflected light (a, c, e, g) and corresponding epifluorescence (b, d, f, h) micrographs
of calcareous foraminifera collected at experimental termination. Those labeled with * were
from a Direct Contact treatment core (935-JB2); those without * were from a Control treatment
core (937-JB7). Note that none of the Direct Contact specimens fluoresce brightly. a, b. Top
row: *Quinqueloculina venusta*, bottom row: *Pyrgo murrhina*; c, d. *Cassidulina* cf. *delicata*; e, f
(left to right): *Hoeglundina elegans, Cibicidoides* sp.; g, h. *Uvigerina canariensis*. Scale bars:
100 μm.

547	Figure 4. Reflected light (a, c, e, g) and corresponding epifluorescence (b, d, f, h) micrographs
548	of agglutinated and thecate foraminifera collected at experimental termination. Those labeled
549	with * were from Direct Contact treatment cores (935-JB2 or 935-JB3); those without * were
550	from Control treatment cores (937-JB6 or 937-JB7). Note that the Direct Contact specimen in
551	c,d fluoresces brightly; other Direct Contract specimens were deemed dead. a, b (left to right).
552	Unidentified allogromiid, saccamminid; c, d (left to right). Pelosina sp., unidentified
553	allogromiid; e, f. Hormosinella guttifera; g, h. Veleroninoides wiesneri. Scale bars: 100 µm.
554	
555	Figure 5. Histograms of CellTracker Green labeled benthic foraminifera, presented as total

density (specimens·10cm³) and by group and treatment for each Experiment. Error bars reflect
standard error.



Figure 1.





Figure 2.





Figure 3.





Figure 4.



Figure 5.