

PREFACE

Long-term studies of the marine ecosystem along the west Antarctic Peninsula

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Articles in this volume focus on longer-term studies of the marine ecosystem of the continental shelf west of the Antarctic Peninsula, principally by the Palmer, Antarctica Long-Term Ecological Research project (Ross *et al.*, 1996; Ducklow *et al.*, 2007). There is a rich history of oceanographic and ecological research in the Bellingshausen Sea region and on the continental shelf dating back to the 19th and early 20th centuries (El-Sayed, 1996). The modern era of scientific research started with the British *Discovery* Investigations of 1925-37 (Hardy, 1967), and included classic studies of phytoplankton (Hart, 1934) and krill (Marr, 1962). Hart's report presciently suggested primary producers could be limited by iron availability. El-Sayed (1996) dissects the subsequent history of oceanographic research up to the advent of the Southern Ocean GLOBEC (Hofmann *et al.*, 2001; Hofmann *et al.*, 2004) and JGOFS (Anderson and Smith Jr., 2001) programs. The period from the 1970's to the mid-90's was dominated by expeditionary and process-level studies of particular regions and processes extending over a few seasons to a few years at most. The Research on Antarctic Coastal Ecosystem Rates (RACER) Program (Huntley *et al.*, 1991; Karl, 1991) is the outstanding example of this mode of research, having focused on determination of key rate processes as a new approach to understanding ecosystem dynamics (Karl *et al.*, 1991a; Karl *et al.*, 1991b). RACER was a direct predecessor and major influence on Palmer LTER, GLOBEC and JGOFS. What was lacking in Antarctic waters, as in most other regions and ocean provinces were sustained, long-term observations of a variety of ocean properties and rates, conducted in the context of hypothesis-driven, experimental science (Ducklow *et al.*, 2008a). The creation of the US LTER Network in 1980 (Magnuson, 1990) made this possible.

Palmer LTER was started in 1990, based on the hypothesis that biological processes in the Antarctic marine environment are strongly affected by physical factors, particularly the

annual advance and retreat of sea ice (Smith *et al.*, 1995). A key early development was creation of a fixed sampling grid of ~ 50 hydrographic stations oriented along the axis of the Peninsula (**Figure 1**) and extending from the nearshore region of persistent pack ice across the continental shelf into the open sea influenced by the Antarctic Circumpolar Current (Waters and Smith, 1992). To date, Palmer LTER has conducted sixteen annual oceanographic cruises in the Austral summer (January-February; **Table 1**). Most sampling within the grid is limited by shiptime constraints to the 200 to 600 lines shown in Figure 1. The principal objective of each midsummer cruise is conducting a comprehensive sequence of core measurements at each station. The current listing of core measurements may be accessed at <http://pal.lternet.edu/publications/documents/protocols/>. The sampling grid addresses multiple spatial scales and permits repeated sampling over seasonal to annual time scales, covering short and long-term ecological phenomena. The grid unifies the measurement components and facilitates data integration. This volume reports on the first decade of cruises and includes the first systematic analyses at the grid scale. The cruise-derived data comprise annual snapshots of midsummer conditions. Satellite observations of ocean color and sea ice concentration provide seasonal and year-round coverage, respectively, in which the summer observations are embedded.

Palmer LTER investigators adopted multivariate statistical approaches to analyze the three-dimensional array of measurements (X,Y,T) obtained from satellite imagery and on the midsummer cruises. The paper by Martinson *et al.* (2008) outlines the Empirical Orthogonal Function (EOF) approach used herein by Smith *et al.* (2008), Martinson *et al.* (2008), Stammerjohn *et al.* (2008), Vernet *et al.* (2008) and Ross *et al.* (2008) to analyze the time-varying spatial distributions of properties and processes within the Palmer LTER grid. Smith *et al.* (2008) report on distributions and interannual variability of phytoplankton pigments derived

from SeaWiFS for 1997-2004. Their analyses extend 400 km seaward beyond the nominal sampling grid to encompass the Permanently Open Ocean Zone beyond the ACC, and also cover the entire Peninsula, thus providing a large-scale context for other observations (Figure 1 in Smith et al. 2008). Smith et al. demonstrate considerable interannual and regional variability in the size of the annual phytoplankton bloom. The bloom is generally initiated offshore in association with the Antarctic Circumpolar Current Front and not necessarily with the retreating ice edge. Montes-Hugo et al. (2008) also used SeaWiFS data (1997-2006) in conjunction with *in situ* data to generate a time series of phytoplankton size structure. They found that transitions in the size composition of the phytoplankton community were associated with changes in ENSO state, mirroring variability in other properties (see below).

The paper by Stammerjohn et al. (2008) utilizes satellite microwave data to derive measurements of sea ice concentration over a 1000 x 400 km region along the Peninsula (within the larger study region of Smith et al. 2008; see Figure 1 of Stammerjohn et al. 2008). Stammerjohn et al. (2008) define the covariability of sea ice, climate and ecosystem properties and identify different ice-ocean marine habitats based on the advance, retreat, duration and persistence of sea ice in different areas along the Peninsula. Interannual variations in the seasonality of sea ice-defined habitats are driven by ENSO and Southern Annular Mode (SAM)-associated variations in local winds. Longer-term trends in declining sea ice duration are associated with enhanced warming in the late autumn and early winter seasons of sea ice formation and advance. Sea ice is an important, though still poorly-sampled and not well-understood habitat for primary producers and consumers (the sea ice microbial community), as well as a key physical factor influencing biological processes in open water. In their paper Fritsen et al. (2008) explore 10-fold differences in biomass (chlorophyll) in sea ice in late winter

2001 and 2002. The timing of sea ice formation (June 2001 and April 2002) is proposed as a key factor regulating differences in accumulated biomass. The atmosphere-ocean-ice-biomass coupling is a mechanism possibly generating longer-period cycles in upper trophic levels characteristic of the Antarctic marine ecosystem.

Martinson et al. (2008) provide analysis of physical oceanographic properties (heat, salt) and define the three distinct, bathymetrically-defined regions shown in Figure 1. They also quantify a heat flux of $\sim 28 \text{ W m}^{-2} \text{ a}^{-1}$ associated with flooding onto the shelf of warm, Upper Circumpolar Deep Water. Variations in UCDW upwelling are driven by the state of ENSO and strength of the SAM, with La Niña and +SAM enhancing the heat flux. This heat flux is the proximate cause of sea ice and glacier melting in the region.

ENSO- and SAM-forced variability that is seen in the LTER study region also impacts on the physical and ecological characteristics of the broader Southern Ocean. Meredith et al. (2008) examined the roles of ENSO and the SAM in determining interannual changes in sea surface temperature around South Georgia, and showed that both fast atmospheric teleconnections and slower oceanic (advective) teleconnections are important. The phasing between these is important for sustaining the anomalies in surface temperature as they propagate across the South Pacific and into the Atlantic. As part of this, rapid atmospheric teleconnections imprint ENSO and SAM signals on the ocean surface immediately west of the Peninsula, offering an additional (more direct) mechanism for coupled climate modes to impact the regional ecosystem, and the ocean downstream. Accordingly, Meredith et al. hypothesize that variability in krill stocks around South Georgia is derived from ENSO- and SAM-forced variations in the advection of krill populations from the Peninsula downstream in the ACC.

Vernet et al. (2008) and Ross et al. (2008) each analyze twelve year time series of in situ biological measurements made within the central LTER grid (200-600 lines of Figure 1). Vernet et al. define spatial patterns and interannual variations in rates of primary production derived from 24-hour ^{14}C incubations. Primary production averaged $\sim 60 \text{ Mol C m}^{-2} \text{ d}^{-1}$ in the grid during January-February 1995-2006. There is a 5-6 year cycle in primary production rates, associated with ENSO forcing of sea ice seasonality. Late sea ice retreat enhanced summer primary production. The numerically dominant macro- and mesozooplanktonic species $> 2 \text{ mm}$ captured in net tows over 1993-2004 included euphausiids (*Euphausia superba*, *Thysanoëssa macrura*; and *Euphausia crystallorophias*), the shelled pteropod *Limacina helicina*, and the salp, *Salpa thompsoni* (Ross et al. 2008). There were both common (e.g., north-south abundance gradients) and taxon-specific features (e.g., inshore to offshore gradients) in their distributions. Like sea ice, plant pigments and primary production, interannual anomalies in abundance and distributions responded to anomalies in the ENSO cycle, suggesting causal connections among these variables. Time-series sediment traps offer a year-round window into water column processes not directly observable from satellites. Ducklow et al. (2008b) report on the 14-year series of sedimentation fluxes captured in a 170-meter deep trap in the Palmer region (Figure 1). Sedimentation approximating export from the surface layer is strongly periodic with $\sim 75\%$ of the annual flux exported during an early-summer pulse. Neither the timing nor magnitude of the pulses are directly related to timing of sea ice or phytoplankton blooms, an indication of the complexity of processes responsible for exporting production into the deep sea. Ducklow et al. show that the annual flux event has advanced by about 40 days later since 1993.

The southern area of the LTER sampling grid is in Marguerite Bay, the focus of the Rothera Oceanographic and Biological Time-Series (RaTS) station of the British Antarctic

Survey at Rothera Station on Adelaide Island (Figure 1). Clarke et al. (2008) reports on eight years of biological and physical observations that extend year-round with weekly resolution. Like other properties already mentioned, the RaTS observations exhibit interannual variability in the same spectral range as ENSO, enforcing the importance of this climate driver in the region. Clarke et al. (2008) demonstrate that UCDW is present even at this inshore location, pointing up the magnitude of irrigation of the shelf ecosystem with offshore waters. Wallace et al. (2008) investigated in detail the shorter-term processes of coastal upwelling and downwelling driven by local winds, and internal tides, that may contribute to the injection of UCDW at the RaTS site. Diurnal internal tides were detected despite the location being poleward of the critical latitude for their generation. The near-coastal location, rough bottom topography and local wind forcing are probably causes of these important physical processes influencing biological productivity in the region.

A key conclusion emerging from analysis of these decade-long observations is the strong teleconnection in the Peninsula region with ENSO variability and its interaction with SAM. It has been hypothesized that anthropogenic warming may generate a more persistent El Niño state (Meehl *et al.*, 2000; Wang and Schimel, 2003) and that the Southern Hemisphere ozone hole intensifies the positive SAM state (Thompson and Solomon, 2002). Better understanding of regional climate, physical oceanographic and ecosystem responses to anomalies in these climate modes will help us to predict responses of the Antarctic marine ecosystem to anthropogenic climate change. Such understanding requires, as one building block, continued time series observations and sustained experimental studies at the interannual and regional scale. Such studies can only be analyzed efficiently if the resulting data and attendant metadata are easily available, accessible and well-cared for. In the concluding paper in this volume, Baker and

Chandler (2008) discuss evolution in the oceanographic community from simple data storage and provision to interdisciplinary information management, emphasizing the demands of increasing volumes of data, more multidisciplinary and interdisciplinary data types and increased collaboration.

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Table 1. Palmer LTER Cruise Summary Information 1990-2014.

Cruise	Vessel*	Begin	End	# days	Year	Lines	Chief Sci
	PD	7-Nov	21-Nov	15	1991	600-700	Quetin
PAL-1	PD	5-Jan	8-Feb	35	1993	200-600	Quetin
	NBP	25-Mar	15-May	52	1993	200-600	Ross
ICE-1	PD	29-Aug	25-Sep	28	1993	200-600	Quetin
	PD	8-Oct	5-Nov	29	1993	200-600	Prezelin
PAL-2	PD	11-Jan	7-Feb	28	1994	300-600	Ross
PAL-3	PD	7-Jan	8-Feb	33	1995	200-600	Smith
PAL-4	PD	8-Jan	10-Feb	34	1996	200-600	Vernet
PAL-5	PD	11-Jan	13-Feb	34	1997	200-600	Ross
PAL-6a	AJ	13-Jan	24-Jan	11	1998	600	Quetin
PAL-6b	LMG	28-Jan	13-Feb	16	1998	200-600	Karl
PAL-7	LMG	8-Jan	12-Feb	36	1999	200-600	Ross
ICE-2	NBP	15-Jun	11-Jul	27	1999	200&600	Smith
PAL-8	LMG	8-Jan	1-Feb	25	2000	200-600	Vernet
PAL-9	LMG	5-Jan	31-Jan	27	2001	200-600	Ross
ICE-3	NBP	11-Sep	22-Oct	42	2001	***.100	Smith
PAL-10	LMG	6-Jan	28-Jan	23	2002	200-600	Ross
PAL-11	LMG	5-Jan	2-Feb	28	2003	200-600	Ross
PAL-12	LMG	7-Jan	2-Feb	26	2004	200-600	Ducklow
PAL-13	LMG	4-Jan	2-Feb	29	2005	200-600	Ducklow
PAL-14	LMG	7-Jan	5-Feb	29	2006	200-600	Quetin
PAL-15	LMG	7-Jan	7-Feb	31	2007	100-600	Ducklow
PAL-16	LMG	5-Jan	4-Feb	28	2008	100-600	Quetin
PAL-17	LMG	TBA	TBA	28	2009	-100-600	Martinson
PAL-18	LMG	TBA	TBA	28	2010	-100-600	TBA
PAL-19	LMG	TBA	TBA	28	2011	-100-600	TBA
PAL-20	LMG	TBA	TBA	28	2012	-100-600	TBA
PAL-21	LMG	TBA	TBA	28	2013	-100-600	TBA
PAL-22	LMG	TBA	TBA	28	2014	-100-600	TBA

* Vessel: PD, Polar Duke; AJ, Abel J; LMG, Laurence M Gould; NBP, Nathaniel B Palmer.

Figure 1. Palmer LTER study area along the west Antarctic Peninsula. The black dots are hydrographic lines 20 km apart in the cross-shelf direction and 100 km apart alongshore. The core study region encompassing the 200-600 grid lines has been occupied each January 1993-2008. The solid black and heavy grey lines identify the three bathymetrically-defined watermass regions (left to right, Slope-Shelf-Coastal). P: Palmer Station on Anvers Island; R: Rothera Station on Adelaide Island; MB, Marguerite Bay; ST, sediment trap mooring. The shaded bathymetry (white >750 m, light-grey, 450-750 m, dark-grey <450 m. Contours at the shelf break are greater than or equal to 1500 m at 750 m intervals.

