ANALYSIS OF UNRESOLVED COMPLEX MIXTURES OF HYDROCARBONS EXTRACTED FROM LATE ARCHEAN SEDIMENTS BY COMPREHENSIVE TWODIMENSIONAL GAS CHROMATOGRAPHY (GC×GC)

Gregory T. Ventura, Fabien Kenig, Christopher M. Reddy, Glenn S. Frysinger, Robert K. Nelson, Ben Van Mooy, and Richard B.Gaines

¹ University of Illinois at Chicago, Department of Earth and Environmental Sciences, M/C 186, 845 West Taylor Street, Chicago, IL 60607-7059, USA

² Department of Marine Chemistry and Geochemistry, MS#4, Woods Hole Oceanographic Institution, Woods Hole, MA 02543-1543, USA

³ Department of Science, U.S. Coast Guard Academy, New London, CT 06320-8101, USA

Abstract

Hydrocarbon mixtures too complex to resolve by traditional capillary gas chromatrography display gas chromatograms with dramatically rising baselines or "humps" of coeluting compounds that are termed unresolved complex mixtures (UCMs). Because the constituents of UCMs are not ordinarily identified, a large amount of geochemical information is never explored. Gas chromatograms of saturated/unsaturated hydrocarbons extracted from Late Archean argillites and greywackes of the southern Abitibi Province of Ontario, Canada contain UCMs with different appearances or "topologies" relating to the intensity and retention time of the compounds comprising the UCMs. These topologies appear to have some level of stratigraphic organization, such that samples collected at any stratigraphic formation collectively are dominated by UCMs that either elute early- (within a window of C₁₅-C₂₀ of *n*-alkanes), earlyto mid- $(C_{15}$ - C_{30} of *n*-alkanes), or have a broad UCM that extends through the entire retention time of the sample (from C_{15} - C_{42} of n-alkanes). Comprehensive two-dimensional gas chromatography time-of-flight mass spectrometry (GC×GC-MS) was used to resolve the constituents forming these various UCMs. Early- to mid- eluting UCMs are dominated by configurational isomers of alkyl-substituted and non substituted polycyclic compounds that contain up to six rings. Late eluting UCMs are composed of C₃₆-C₄₀ mono-, bi-, and tricyclic archaeal isoprenoid diastereomers. Broad UCMs spanning the retention time of compound elution contain nearly the same compounds observed in the early-, mid-, and late retention time UCMs. Although the origin of the polycyclic compounds is unclear, the variations in the UCM topology appear to depend on the concentration of initial compound classes that have the

potential to become isomerized. Isomerization of these constituents may have resulted from hydrothermal alteration of organic matter.

1. Introduction

The identification of individual molecular constituents within complex organic mixtures of petroleum and sediment solvent extracts is typically achieved by a series of chemical and chromatographic separations. The most common technique employs a capillary gas chromatography (GC), often coupled to a mass spectrometer, in order to facilitate the separation, quantitation, and identification of individual molecular components in complex mixtures. Unfortunately, the chromatographic resolution afforded by capillary GC is insufficient to resolve some complex mixtures, which appear as a pronounced rising baseline or a series of rising baselines in a gas chromatogram (e.g. Gough and Rowland, 1990). In oils, the number of unidentified compounds comprising such unresolved complex mixtures (UCMs) may amount to 250,000 compounds (Sutton *et al.*, 2005) indicating that an enormous amount of inaccessible, geochemical information is unexploited.

UCMs are frequently observed in petroleums and hydrocarbon (HC) extracts that have undergone biodegradation (Rowland and Maxwell, 1984; Killops and Al-Juboori, 1990; Gough and Rowland, 1990; Frysinger *et al.*, 2003) or hydrothermal alteration (Rushdi and Simoneit 2002; Simoneit *et al.*, 2004; Zárate-del Valle and Simoneit, 2005). UCMs of biodegraded oils typically contain a low abundance or complete absence, of aliphatic compounds (e.g. Connan, 1984; Rowland *et al.*, 1986; Palmer, 1993; Swannell *et al.*, 1995; Bost *et al.*, 2001; Reddy *et al.*, 2002). The removal of these constituents, which averages 32% of the total hydrocarbon

composition of most crude oils (Killops and Killops, 2005) concentrates a preexisting complex mixture to yield the amplified chromatographic baseline of a UCM (Killops and Al-Juboori, 1990).

Petroleum formed by hydrothermal alteration of sedimentary organic matter undergo accelerated diagenesis and catagenesis (Simoneit and Lonsdale, 1982) from exposure to high temperature fluids that reach up to 350 °C (Simoneit, 1984, 1994; Simoneit et al., 1992; Kvenvolden and Simoneit, 1990). The UCMs of these hydrothermal petroleums contain abundant low and high molecular weight n-alkanes (e.g. Ogihara and Ishiwatari, 1998; Yamanaka et al., 2000; Rushdi and Simoneit, 2002a and b; Simoneit et al., 2004). These petroleums occasionally have homologous or pseudohomologous series of compounds with a carbon number predominance, such as n-cycloalkanes and branched alkanes with quaternary carbon atoms (BACQs; Ogihara and Ishiwatari, 1998; Simoneit, 1994; Rushdi and Simoneit, 2002a and b; Simoneit et al., 2004). The UCMs of hydrothermal petroleums are thought to be composed of branched and cyclic compounds (e.g. Rushdi and Simoneit, 2002; Simoneit et al., 2004). Aside from this, their composition and formation has yet to receive the same level of attention as the UCMs produced by biodegradation. Because the compositional evolution of petroleum is strongly dependent on numerous variables such as temperature, time, source composition, water washing, fractionation during migration, and microbial degradation, the large inventory of organic compounds present in UCMs may yield critical information that can be used to understand factors regulating the occurrence of petroleum in subsurface environments.

Attempts to identify the constituents in UCMs have met with varying degrees of success and involved chemical oxidation (Gough and Rowland, 1990; Killops and Al-Juboori, 1990; Revill *et al.*, 1992; Warton, 1999; Warton *et al.*, 2000), thin layer chromatography (Liu *et al.*,

2005), chemical oxidation followed by treatment with molecular sieves (e.g. Armanios et al., 1994; Ellis et al., 1994; Fazeelat et al., 1994), statistical deconvolution methods (Pool et al., 1997; Dagan 2000; Demir et al., 2000) and field ionization mass spectrometry (Payzant et al., 1979). These attempts have led to the initial belief that UCMs were composed of hydrocarbons with similar chemical properties that include large numbers of branched and cyclic aliphatic and aromatic isomers (Eglinton et al., 1975; Payzant et al., 1979; Alexander et al., 1982; Sanders and Tibbetts, 1987; Killops and Al-Juboori, 1990). The use of oxidative degradation followed by GC enables the release of additional compounds, which led to the notion that UCMs might be mixtures of fairly simple compounds comprising unsubstituted alkyl chains such as isometric monoalkyl substituted "T"-branched alkanes (Gough et al., 1992; Warton et al., 2000). "T"branched alkanes were shown to be resistant to biodegradation (Gough et al., 1992). Warton et al. (1997) demonstrated that 3% of the alkanes from a biodegraded crude oil were "T" branched alkanes. Payzant et al. (1979) showed that UCM of the urea non adduct fractions of heavy oils was dominated by mono- and bicyclic compounds, with decreasing contribution of tri- to hexacyclic compounds. However, the structure of most of these compounds was not determined.

Even with excellent capillary GC columns and the extensive use of chemical degradation methods, some complex mixtures of compounds cannot be resolved and the nature of the compounds that form these UCMs remains unclear (Gough *et al.*, 1992; Gough & Rowland, 1990; Warton, 1999). This problem, which was latent in the study of biodegraded crude oil and crude oil contamination (i.e. Frysinger and Gaines, 2001; Reddy *et al.*, 2002; Frysinger *et al.*, 2003), is overcome with comprehensive two-dimensional gas chromatography (GC×GC; Liu & Phillips, 1991). GC×GC links two capillary columns, with different stationary phases, via a modulator that creates packets of analytes by temporarily focusing the effluent leaving the first

column. The entrance of these packets into the secondary column produces a chromatogram with a high signal-to-noise ratio. Furthermore, the separation power of the primary column is conserved into the secondary column, such that compounds not resolved by the first column may be resolved by the second column. GC×GC was successfully used to separate and identify biomarkers (molecular fossils) in crude oils (Frysinger and Gaines, 2001), modern and Holocene marine sediment extracts (Johnson *et al.*, 2003), and UCMs of crude oil contaminated sediments (Frysinger *et al.*, 2003; Reddy *et al.*, 2002). More recently, GC×GC has been used to distinguish drilling mud contaminants from naturally occurring compounds in oils (Reddy *et al.*, 2007).

Sediments from the 2.710-2.704 Ga Tisdale and 2.685 to 2.676 Ga Porcupine Assemblage (Ayer *et al.*, 2002) of the Abitibi greenstone belt of Ontario, Canada were analyzed to assess the composition, origin, and preservation potential of organic matter from Late Archaean hydrothermal depositional environments (Ventura, 2006; Ventura *et al.*, 2007). Bitumens were extracted from greywackes and argillites of the Tisdale Assemblage were collected from interflow sedimentary rocks deposited between tholeitic basalts and tholeitic dacites (Brisban, 1997). Bitumens were also extracted from distal margin turbidites of the overlying Porcupine Assemblage (Fig. 1; Rice *et al.*, 1992). These sediments were subjected to post burial hydrothermal alteration (Brisban, 1997; Kerrich and Ludden, 2000) and lower greenschist metamorphism (Dimroth *et al.*, 1983; Thompson, 2002).

The saturated and unsaturated hydrocarbon (s/u HC) fractions of these bitumens were observed to contain pronounced UCMs (Fig. 1; Ventura, 2006; Ventura *et al.*, 2007). These UCMs have different appearances or "topologies" relating to the intensity and retention time of the compounds comprising the UCMs. These topologies also appear to have some level of stratigraphic organization suggesting that the type of UCM topology likely dependents upon the

specific nature of the initial source of organic matter as well as diagenetic, or catagenetic processes (Ventura *et al.*, 2007). In this study GC×GC-MS is used to identify the hydrocarbons that represent the dominant compounds forming the different UCM topologies. These results are then used to explore potential mechanisms of UCM formation.

2. Samples and methods

2.1 Samples

Samples were collected near Timmins, Ontario in an area named the Porcupine Gold Camp (PGC). Samples were selected to represent the pronounced UCM topology of each stratigraphic formation. The s/u HC fractions of solvent extracts from seven samples included two greywackes from the Vipond Formation, one argillite from the Krist Formation, and four argillites from the Hoyle Formation. A complete description of the geologic setting, bulk organic carbon characterization, and biomarker composition is provided in Ventura (2006) and Ventura *et al.* (2007).

2.2 Methods

2.2.1. Bitumen extraction

The surfaces of all samples were either scrubbed first with a wire brush and rinsed with double distilled water before being washed with dichloromethane (DCM)/methanol (MeOH) in a 7.5:1, v/v ratio or ground off using a steel wire brush attached to a Dremmel. The samples were then ground to a powder using a shatter box. Between 100 to 180 g of powdered sediment was Soxhlet extracted with a mixture of dichloromethane: methanol (DCM:MeOH; 7.5:1 v/v) for 72 hr. The resulting bitumen was then separated into polar and apolar fractions by column chromatography on alumina oxide following the method described in Simons and Kenig (2001). The apolar fraction was collected using hexane/DCM (9:1 v/v) as an eluent. Polar fractions were recovered using DCM/MeOH (1:1 v/v) as an eluent. Elemental sulfur was removed from the apolar fractions by passing the apolar fraction through a Pasteur pipette filled with activated Cu (HCl 6N) using hexane as the eluent. Apolar fractions were then separated into s/u HC and aromatic fractions by column chromatography on silica gel. S/u HC fractions were recovered with hexane. Aromatic fractions were recovered using DCM/MeOH (7.5:1 v/v) as an eluent. Sub-fractions were rotary evaporated to ~1mL volume, transferred to tared vial, dried under a continuous low flow of N₂, and again weighed. A procedural blank was added to each batch of samples.

2.2.2. Gas Chromatography –Mass Spectroscopy (GC-MS)

The s/u HC and aromatic fractions were analyzed in full scan and selected ion monitoring (SIM) modes using a Hewlett-Packard 6890 gas chromatograph (GC) coupled to a Hewlett-Packard 5973 mass selective detector (MSD). The GC was fitted with a HP-5MS capillary column (30-m; 0.25-mm ID; 0.25-µm film thickness). The MSD was operated in electron ionization mode at 70 eV, scanning a mass range of *m/z* 40-650 at 2.44 scans/sec. Injection was

done in pulsed-splitless mode and helium was the carrier gas. Samples were injected at 60 °C (held for 1.5 min.). The oven temperature was programmed to 130 °C at 20 °C min⁻¹, and then at 4 °C min.⁻¹ to 315 °C, and held at 315 °C for 90 min.

2.2.3. Metastable reaction monitoring gas chromatography-mass spectrometry (MRM-GCMS)

Metastable reaction monitoring (MRM) was performed at the Department of Earth, Atmospheric and Planetary Sciences, Massachusetts Institute of Technology (Cambridge, MA, USA) following the method described in Brocks *et al.* (2003). Saturated and unsaturated fractions of 23 of the 26 samples were analyzed by GC-MS MRM using a VG Autospec Ultima-Q coupled to a CarloErba GC (8000 series) with a tandem high resolution, double focusing magnetic sector-quadroupole mass analyzer. Internal standard D⁴ (d₄ –C₂₉-ααα-ethylcholestane; Chiron Laboratories AS) was added to all samples prior to injection.

2.2.4. Comprehensive two-dimensional gas chromatography –mass spectroscopy (GC×GC-MS)

S/u HC fractions were analyzed by GC×GC-MS using a Leco Pegasus 4D system consisting of a Hewlett-Packard 6890 GC configured with a split injector, two chromatography columns, a liquid nitrogen-cooled pulsed jet modulator, and a time-of-flight (TOF) mass spectrometer. Samples were dissolved in 25 μL cyclohexane and 2.0 μL was injected into a 300 °C splitless injector (2 min. purge time). The first-dimension separation was performed on a nonpolar 5% phenyl polydimethylsiloxane phase (Agilent, DB-5, 30.0-m; 0.25-mm I.D.; 0.25-μm film thickness) and temperature programmed from 120 (1 min.) to 320°C at 2°C min.⁻¹. The

modulation column was deactivated column (1.0-m x 0.10-mm I.D.) and temperature programmed from 200 (1 min.) to 400 °C at 2 °C min.⁻¹. The second-dimension separation was performed on a polar 50% phenyl equivalent polysiloxane phase (BPX-50, SGE, 1.5-m; 0.10-mm I.D.; 0.1-μm film thickness) and temperature programmed from 120 (1 min.) to 320 °C at 2 °C min.⁻¹. A 1.0 m TOF detector transfer line of 0.10-mm I.D. deactivated column was used. Hydrogen was used as the carrier gas in constant flow mode (1.2 mL/min). The GC×GC modulator period was 8 s. The GC×GC was coupled with a Leco time of flight mass spectrometer that collected spectra from 45 to 600 u at 100Hz. A detector voltage of 1800 V was used with a solvent delay of 11 min. The similar column array used for GC×GC-FID provided equivalent compound separation.

Comparisons of retention times and mass spectral data collected by GC-mass spectrometry (GCMS), GC-MRM-MS, and GC×GC-MS biomarker compound classes were used to identify the presence of common biomarker compound classes such as *n*-alkanes, branched alkanes, acyclic isoprenoids, and polycyclic terpanes, such as hopanes and steranes (Ventura, 2007). C₄₀ acyclic irregular isoprenoids were identified by mass spectral analysis using GC×GC-MS and coinjection of an archaeal lipid standard containing acyclic, mono-, bi-, and tricyclic biphytane. Cycloalkanes were tentatively identified by mass spectral analysis. All reported mass spectra were obtained using the "Peak True" deconvolution algorithm of ChromaToF, the data processing tool of Leco's software package. The "Peak True" algorithm is based on an approach by Biller and Biemann (1974). "Peak True" mass spectra have greater fragmentation resolution of closely eluting peaks and represents a more consistent and tractable method than conventional caliper selected approaches.

2.2.5. Preparation of the archeael biphytane standard.

A standard mixture of archaeal biphytanes was prepared from membrane phospholipids of the archaeon *Thermoplasma acidophilum* (Matreya L.L.C) using protocols adapted from Pease *et al.* (1992) and DeLong *et al.* (1998). To remove the glyco- and phospho-lipid headgroups, the membrane lipid was hydrolyzed in 1.25 M HCl in MeOH at 100 °C for two hours. The reaction was neutralized with 3M NaOH and extracted three times with DCM. After drying over NaSO₄, an aliquot of the resultant glycerol dialkyl glycerol tetraethers (GDGTs) was examined for purity using high pressure liquid chromatography-mass spectrometry (HPLC/MS) as described by Hopmans *et al.* (2000). To cleave the ether bonds of the of GDGTs, the extract was refluxed in concentrated HI for 4 hours at 100 °C, extracted three times with hexane and dried via rotary evaporation. To reduce the resultant alkyl iodides to hydrocarbons, the extract was dissolved in 2M LiAlH₄ in THF and refluxed under N₂ for 2 hr. at 70 ° C. Ethyl acetate was added dropwise to quench excess LiAlH₄, and then the reaction was extracted with three additional aliquots of ethyl acetate, which were then dried over NaSO₄. The presence of biphytanes was confirmed by GC-MS and comparison with spectra of DeLong *et al.* (1998).

3. Results

All samples collected at the PGC have s/u HC fractions with UCMs that can be grouped in three categories based on chromatographic baseline topology (Fig. 1). Type I UCMs are raised baselines from early- to mid- retention times corresponding to the elution of C_{15} - C_{25} n-alkanes. Type II UCMs are dramatically rising baselines at late retention times that span the range of C_{30}

to C_{42} *n*-alkanes UCMs with a type III topology have a single massive UCM of early- to lateretention time (n- C_{15} to n- C_{42}).

UCM topologies vary stratigraphically and geographically (Fig. 1). The s/u HC fractions of the four samples collected from the Vipond Formation have a type III UCM topology. The two samples collected from the Gold Center Formation samples have a type II UCM topology, and the Krist Formation sample has a type I UCM topology. All ten Hoyle Formation samples collected away from areas of gold mineralization have s/u HC fractions with a type I UCM topology and all nine Hoyle Formation samples collected in areas of gold mineralization have both type I and II topologies.

GC-MS was unable to resolve the constituents forming the three UCM topologies. Averaged mass spectra across these UCMs display abundant mono- and bi-unsaturated fragment ions, such as m/z 97 and m/z 95 (Fig. 2). Due to the abundance of aliphatic compounds within the Type I topology the most intense fragment ion is m/z 57 (Fig. 2A). Mass chromatography of the fragment ions m/z 97, m/z 95 and m/z 93 illustrates the close elution of compounds producing mono-, bi-, and tri-unsaturated fragment ions upon electron ionization in all UCMs (Fig. 2A-D). Hydrogenation of these samples does not affect the UCM topology, suggesting the UCMs are dominated by cyclic compounds.

3.1 Composition of Type I UCMs

Type I UCMs are operationally defined as having a raised "baseline" in the n-C₁₅ - n-C₂₅ range (Fig. 1). The GC×GC-MS total ion current (TIC) chromatogram (Fig. 3A) displays some well resolved compounds (e.g. n-alkanes). However, the mass of cyclic compounds appears unresolved. Because compounds forming the type I UCM have similar fragmentation patterns,

monitoring of typical mono- and bicyclic fragment ions (e.g. m/z 69, m/z 83, m/z 97, and m/z 123) does little to resolve diversity of these compounds. Visualization of individual peaks within the UCM is greatly enhanced by sequentially monitoring mass chromatograms of molecular ions for compounds with different number of rings and different numbers of carbon atoms over the entire elution range of the UCM (Fig 3 B and C). For example, for the C_{16} and C_{17} compounds, mass chromatography of m/z 224 and m/z 238 (monocyclic), m/z 222 and m/z 236 (bicyclic), m/z 220 and m/z 234 (tricyclic), and m/z 218 and m/z 232 (tetracyclic) shows four sub-parallel lineament of peaks, and one peak for m/z 218 (tetracyclic; Fig. 3B and C). When this approach is applied to molecular ions of compounds from C_{15} to C_{29} (Fig. 3D) the bulk of the UCM is observed to consist of near uniformly distributed peaks of overlapping groups of mono- $(C_{13}-C_{29})$, bi- $(C_{15}-C_{25})$, tri- $(C_{16}-C_{25})$, tetra- $(C_{16}-C_{23})$, penta- $(C_{19}-C_{21})$, and hexacyclic $(C_{22}-C_{23})$ unsusbstituted and alkylsubstituted compounds (Fig. 3D). The tentative identification of some of these compounds is provided below.

3.1.1 Monocyclic alkanes

The overall distribution of monocyclic alkanes is monitored by the summed mass chromatography of fragment ions m/z 96, and m/z 110 (Fig. 4A). However, to resolve the distribution of monocyclic alkanes at each carbon number, the mass chromatogram of the molecular ion was added to the summed mass chromatogram. For example, monitoring of C_{16} alkylcycloalkanes was done on the summed chromatogram m/z 224 + m/z 96 + m/z 110 (Fig. 4B).

n-Alkylcyclohexane and *n*-alkylcyclopentane were identified on the basis of mass spectral fragmentation and by comparison with published spectra (e.g. Rubinstein and Strausz,

1979; Hoffmann *et al.* 1987; Simoneit *et al.* 2004). *n*-Alkylcyclopentane and *n*-alkylcyclohexane homologues having 16 or less carbon atoms nearly coelute (Fig. 4B). Homologues with greater than 16 carbon atoms become increasingly separated in the first dimension (Fig. 4A). Tentatively identified *n*-alkylcycloheptane (Fig. 5A) elutes at slightly earlier first dimension and later second dimension retention times relative to *n*-alkylcylohexane and *n*-alkylcyclopentane (Fig. 4A and B). The *n*-alkylcyclohexanes, *n*-alkylcyclopentanes, and *n*-alkylcycloheptanes of the Vipond and Krist Formation sediment extracts have extremely pronounced carbon number predominance. Such carbon number preference is inconsistent with the age and maturity of these sediments and is likely a product of more recent contamination (Grosjean and Logan, 2007). A similar carbon number predominance is not observed in the substituted cycloalkanes contributing to the formation of type I UCM.

A group of 7 peaks, tentatively identified as isomers of nonylmethylcyclohexane (Fig. 4B), were identified by mass spectral comparison and by the similarity of the first dimension isomers elution pattern to published data (Fowler *et al.*, 1986; Hoffmann *et al.*, 1987). The mass spectra of *trans* 1-methyl-2-nonylcyclohexane is shown in Figure 5B. The mass spectra of all observed alkylmethylcyclohexane stereoisomers were nearly identical and dominated by the fragment ion m/z 97. Alkylmethylcyclohexanes were observed in the C_{16} - C_{28} range, which extends past the zone of maximum intensity of the UCM.

Three groups of monocyclic compounds eluting at slightly earlier first and second dimension times to the cluster of alkylmethylcyclohexanes (Fig. 4A and B) are also present. The peaks in each group contain identical fragmentation patterns suggesting the members of each group have stereochemical differences similar to the alkylmethylcyclohexanes. However, slight difference in the fragmentation patterns between each group indicates a variation in the number

of alkyl-substituents bonded to the cyclohexane or cyclopentane ring. The first group containing three peaks that elute before the alkylmethylcyclohexanes is tentatively identified as isomers of alkylmethylcyclopentanes (Fig. 5). The mass spectra of these peaks display a pattern of skeletal ring decomposition which results in the formation of 1-alkene ions (McLafferty and Tureček, 1993) that is common to cycloalkanes with less than six-membered rings. The dominant m/z 83 fragment ion, characteristic of a 6 carbon ring, is likely due to the presence of a methyl substituent on the pentacyclic ring. The two other groups noted Unknown X_1 and X_2 in Figure 4B could not be tentatively identified on the sole basis of their mass spectral fragmentation (Fig. 5E and F), but are likely compounds with two or more methyl substituents on a cyclohexane or cyclopentane ring.

3.1.2. Bicyclic alkanes

Bicyclic alkanes above C_{18} occupy part of the same GC×GC chromatographic space as monocyclic alkanes that have one less carbon atom (Fig. 3D). Monitoring of bicyclic alkane distributions is thus facilitated by mass chromatography of molecular ions. In the mass chromatogram of the molecular ion of C_{16} bicyclic alkanes (m/z 222; Fig. 6A), three groups of peaks with very similar mass spectra are identified.

The first group (B in Fig. 6A) includes at least four peaks. The mass spectra of these compounds (Fig. 6B) include the fragment ion m/z 137, which corresponds to a cleaved C_{10} bicyclic ion (decalin) from the C_6 alkyl substituent of the molecular ion (m/z 220). The presence of a smaller fragment ion m/z 136 indicates a hydrogen rearrangement, formed when an alkyl substituent is lost via the cleavage of the C-C bond α to the ring system. The base peak m/z 81 is formed by the cleavage of the cyclohexane rings of the decalin ion α to the two tertiary carbon

atoms of the bicyclic system (e.g. Golovkina *et al.*, 1984). Thus, it can be hypothesized that the compounds forming the first group are isomers of hexyldecalins. However, the position of the alkyl substituent (1 or 2) cannot be determined and it is possible that both isomers are present.

The second group (C in Fig. 6A), includes four peaks with nearly identical mass spectra that are tentatively identified as alkyl-methyldecalins (Fig. 6C). Cleavage of a C_5 alkyl substituent from the molecular ion results in the formation of a C_{11} bicyclic ion represented by the fragment ion m/z 151. The base peak m/z 95 indicates fragmentation α to the tertiary carbon atoms of the methyl substituted cyclohexyl-ring. The position of both methyl and pentyl substituents cannot be determined on the sole basis of the mass spectral data.

The third group (D) includes three peaks with nearly identical mass spectra (Fig. 6 A). The mass spectra of these compounds have an m/z 81 base peak likely formed by the cleavage of a cyclohexane ring from a bicyclic system. However, the mass spectral data is unspecific and no structure can be proposed at this point.

3.1.3. Tricyclic compounds

Tricyclic compounds volumetrically contribute less to the type I UCMs than monocyclic and bicyclic compounds. Five groups of compounds are identified on the mass chromatogram of the molecular ion (m/z 220) of C₁₆ tricyclic alkanes (Fig. 7A). All of the peaks within each group have nearly identical mass spectra. The mass spectra of these compounds (Fig. 7) do not match those of tricyclic terpanes (e.g. they do not have a fragment ion m/z 123). Interpretation of the mass spectrometric data was facilitated by Kiselev's *et al.* (1984) study of perhydrophenanthrene and perhydroanthracene fragmentation. The mass spectrum of peak E (Fig. 7E) has a fragment ion M^+ -29 (m/z 191) of intensity equal to that of the molecular ion,

suggesting the loss of an ethyl substituent from one of the terminal rings of the perhydroanthracene or perhydrophenantrene. Otherwise, the mass spectrum of peak E is similar to those published by Kiselev *et al.* (1984). Detail of the fragmentation pattern is shown in figure 7E. The fragment ion m/z 150 corresponds to a perhydro-cyclopropa[a]naphthalene. Thus, peak E is tentatively identified as an ethyl-perhydroanthracene or ethyl-perhydrophenanthrene, though the location of the ethyl substituent remains unclear.

The compounds of the first group (B) of Fig. 7A are tentatively identified as isomers of dimethyl-substituted perhydroanthracene or perhydrophenanthrene on the basis of mass spectral data (Fig. 7B). The fragment ions M⁺-15 and M⁺-29 correspond to the loss of one and two methyl substituents, respectively. The fragment ions m/z 149 and m/z 164 correspond to an octahydro-methylnaphathalene and perhydro-cyclopropa[a]naphthalene, respectively. The position of the methyl substituents, located on terminal rings, could not be determined. The group of peak C and peak D (Fig. 7A) are tentatively interpreted to be mono- or bi-substituted dodecahydrofluorene (Fig. 7C, and D) on the basis of mass spectral fragmentation. No tentative structure for the group of peak F is proposed.

3.1.4. Tetracyclic alkanes

Tetracyclic alkanes are also part of the type I UCM (Fig. 8). The mass chromatogram of the molecular ion of C_{16} tetracyclic compounds (m/z 218) displays one peak, tentatively identified as hexadecahydropyrene by comparison of its mass spectra with that of published reports (McLafferty *et al.*, 1989). The mass chromatogram for the m/z 232 molecular ion of C_{17} tetracyclic compounds displays 5 major peaks (Fig. 8A). Peak B, C, and D are tentatively identified as methyl-hexadecahydro-pyrenes on the basis of mass spectral fragmentation patterns

(Fig. 8B, C and D). These peaks all have M^+ -15 fragment ions (m/z 217) suggesting the presence of a methyl substituent. For peak B and C, methyl substitution at C-1 or C-2 favors cleavage of the C₄H₈ (isobutylene) or C₄H₁₀ (isobutyl) from the molecular ion, resulting in the fragment ions m/z 175 and m/z 177, respectively (Fig. 8B and C). The fragment ions m/z 189 and m/z 191 are formed by cleavage of the C₃H₈ or C₃H₆ (propylene) from the ring opposed to that carrying the methyl substituent. Thus, peaks B and C are tentatively interpreted to be 1- or 2-methyl-hexadecahydro-pyrenes. For peak D, methyl substitution at C-4 explains the absence of the fragment ions m/z 175 and m/z 177 (no possible cleavage of C₄ units) and the presences of the fragment ions m/z 189 and m/z 191, via cleavage of a C₃H₈ or C₃H₆ unit from the molecular ion (Fig. 8D). Thus peak D is tentatively identified as a 4-methyl-hexadecahydro-pyrene. The mass spectral fragmentation for peak E (Fig. 8E) cannot be rationalized and no tentative structure is proposed.

There is no fragment ion M^+ -15 in the mass spectrum of peak F (Fig. 8F), suggesting that this compound is not methyl substituted. Thus, considering the abundance of the fragment ions m/z 175 and m/z 191, we tentatively identified peak F as hexadecahydro-benzo[de]anthracene, the only structure possible.

3.1.5. Pentacyclic alkanes

Pentacyclic alkanes were tentatively identified in the type I UCM where they are in low abundance compared to mono-. bi-, tri-, and tetracyclic alkanes. The distribution of C_{19} and C_{20} pentacyclic compounds is monitored on the summed mass chromatogram m/z 258 + m/z 272 (Fig. 9A). Peaks B, C, and D are C_{19} compounds (M^+ 258). Peaks E, F, G, and H are C_{20} compounds (M^+ 272).

Octadecahydro-benzo[cd]pyrene is the only possible structural isomer for an hexacyclic compound comprised solely of 6-membered rings having a mass of 258 amu (Fig. 9B). The mass spectrum of peak B (Fig. 9B) contains the fragment ions m/z 217 (M⁺-41) and m/z 215 (M⁺-43). These fragment ions correspond to tetradecahydropyrene and dodecahydropyrene ions formed by cleavage of propylene and propane, respectively, from the fragment ion (M⁺-258). Fragment ions m/z 173 and m/z 175 correspond to octahydrophenalene and decahydrophenalene ions formed by cleavage of saturated and unsaturated C_3 units from dodecahydropyrene ions. Thus, on the basis of mass spectral data, peak B is tentatively interpreted as octadecahydrobenzo[cd]pyrene.

Peak C and D must be pentacyclic compounds with at least one cyclopentane ring. The mass spectrum of peak C is weak and difficult to interpret and most probably includes fragment ions derived from more than one compound. However, the very high intensity of fragment ion m/z 215 (M⁺-43; 100%) suggests presence of a methyl substituent on the pentacyclic ring and cleavage of the bond between the two adjacent tertiary carbon atoms (Fig. 9C). Thus, very tentatively, peak C may be octadecahydro-methylcyclopenta[cd]pyrene. The mass spectra of peak D does not display an M⁺-15 fragment ion (Fig. 9D). The abundance of the fragment ion m/z 173, corresponding to an octahydrophenalene ion, suggests the cleavage of two rings and loss of two C₃ units. Thus, an octadedahydro-cyclopenta[e]pyrene structure is proposed for peak D.

Peaks E and F, both C_{20} pentacyclic hydrocarbons, have very similar mass spectra (Fig. 9E and F). The fragment ion m/z 257 (M⁺-15) suggests the presence of a methyl substituent on the pentacyclic ring system. The tetradecahydro-pyrene ion (m/z 215) indicate the cleavage of a C_4 unit, suggesting that the methyl substituent is located in position 3, 4 or 5 of an

octadecahydro-benzo[cd]pyrene. Thus, peak E and F are tentatively identified as octadecahydro-methylbenzo[cd]pyrene isomers.

Peak G and H are both C_{20} pentacyclic hydrocarbon. The mass spectra of both these peaks do not display an M+-15 fragment ion. The mass spectral fragmentation patterns for these compounds are typical of a condensed ring system but are not informative enough to hypothesize a structure.

3.1.6. Hexacyclic alkanes

Hexacyclic alkanes were also detected in type I UCM. Mass chromatography of the molecular ion for C_{22} hexacyclic compounds (m/z 298; Fig. 10A) has three peaks with mass spectra including m/z 298 as a molecular ion. All these mass spectra include the fragment ion m/z 257 (M^+ -41) corresponding to the typical loss of a propylene fragment from a condensed ring system, as observed for tetracyclic and pentacyclic alkanes. None of the spectra display fragment ions with m/z M^+ -15 nor M^+ -29 (Fig. 10B) suggesting that these compounds are not methyl substituted and, thus, do not include five membered rings.

Only three structural isomers are possible for hexacyclic C_{22} compounds made of cyclohexane rings having a mass 298 amu: Icosahydro-dibenzo[cd,mn]pyrene, icosahydro-dibenzo[def,mno]chrysene, and docosahydro-benzo[ghi]perylene (Fig. 10C). However, it is not possible to determine which of these three structures corresponds to each peak in the mass chromatogram m/z 298, the mass spectral data for peaks C and D (not shown) being too weak.

3.2. Composition of Type II UCMs

GC×GC-MS enables identification of compounds forming the bulk of the type II UCM. These UCMs are composed of acyclic and cyclic terpenoids with 1 to 5 rings, which form bands separated in the second dimension of the GC×GC chromatogram (Fig. 11). Each band is composed of one or more series of compounds. For example, the tetracyclic band include both steroids and secohopanes, which elute in the second dimension after a pseudohomologous series of tricyclic terpanes (cheilathanes) and before a later eluting band pentacyclic triterpanes (hopanoids; Fig. 11).

The dominant compounds forming the type II UCMs are C₃₆-C₄₁ acyclic, mono-, bi- and tricyclic irregular (head-to-head) isoprenoids. Acyclic 3,7,11,15,18,22,26,30octamethyldotriacontane (C₄₀H₈₂; biphytane) and C₄₀ mono-, bi- and tricyclic cyclic archaeal lipids were identified on the basis of mass spectral fragmentation and coelution with an archaeal lipid standard (Fig. 12). Lower carbon number derivatives of the acyclic and cyclic biphytanes were identified on the basis of mass spectral data, following the rational of the fragmentation of acylic and cyclic biphytane (Fig. 12 and 13). GC×GC-MS analysis of sample OC-114m enabled the resolution of additional acyclic irregular isoprenoid with more than 40 carbon atoms (Fig. 11 and 12). No cyclic isoprenoids with more than 40 carbons were detected and neither tetracyclic biphytane nor derivatives of tetracyclic biphytanes were observed. The C₄₀ mono, bi-, and tricyclic biphytane derivatives and their lower C₃₆-C₃₉ homologs are each, respectively, represented by doublets, triplets, and at least quadruplets of peaks with exactly the same mass spectra (Fig. 10). Within each group, the peaks are clearly separated in the 1st dimension, but have nearly the same 2nd dimension retention time, suggesting differences in volatility but similar polarity.

3.3. Composition of type III UCMs

The s/u HC fraction of all samples collected from the Vipond Formation have a type III UCM topology. GC×GC-MS showed that type III UCMs are made of nearly the same compounds as those observed for type I and II UCMs. GC×GC chromatograms are dominated by overlapping bands of alkyl-substituted mono-, bi-, and tri-cyclic compounds at early to midretention times. Although tetracyclic alkanes were not detected, minor contributions of penta- and hexacyclic compounds were also present. The compounds forming the mass of the late eluting part of the UCM are dominated by hopanoids, steroids, tricycic terpanes, secohopanes, and acyclic, mono-, bi- and tricyclic irregular isoprenoids (archaeal lipids) with carbon number ranges similar to those observed for compounds forming the type II UCM. Thus, the massive type III UCM appears to be a hybrid of type I and II UCMs.

4. Discussion

4.1. Origin of PGC UCMs

The UCMs observed in the PGC samples are unlikely to have arisen by selective preservation of bioresistant compounds during biodegradation as ascribed to UCMs of biodegraded oils (e.g. Killops and Al-Juboori, 1990). Easily biodegradable, low molecular weight (LMW) *n*-alkanes and LMW acyclic isoprenoids, such as 2,6,10-trimethyltridecane or 2,6,10-trimethylpentadecane (Peters and Moldowan, 1993), are abundant in all samples. This is concordant with the observed low concentrations of biodegraded C-25 norhopanes (Ventura, 2006). Furthermore, biodegradation can explain neither the presence of highly isomerized

archaeal lipids observed in the type II UCMs, nor the abundance of mono-, bi-, tri-, tetra-, pentaand hexacyclic alkanes in type I UCMs.

4.1.1. Origin of type I UCMs

The type I UCM forming compounds are not likely contributions from laboratory or field related contamination. Most of these compounds do not have a carbon number preference. Exceptions are *n*-alkylcyclopentanes, *n*-alkylcyclohexanes, *n*-alkylcycloheptanes of samples from the Vipond and Krist Formation, which have concentrations that are positively correlated with BAQCs. These compounds likely reflect contamination by products of commercial polyethylene products (Grosjean and Logan, 2007; Brocks *et al.*, 2008). Monocyclic alkanes are devoid of a carbon number preference in samples that do not contain BAQCs (Ventura *et al.*, 2007). The type I UCM forming compounds are also present in samples that had the surface of the core ground away prior to extraction (Ventura *et al.*, 2007) and thus are unlikely to derive from surface contamination of the sample. It is unclear, however, whether these compounds share a common origin with one another or with those forming type II UCMs.

Very little information is currently available about the occurrence or formation of the tentatively identified cyclic compounds. Payzant *et al.* (1979, 1980) showed that the UCMs of extractable hydrocarbon of the Mannville oils, Lloydminster heavy oil, Cold Lake Bitumen, Athabasca bitumen, and Athabasca asphaltene pyrolysate are dominated by mono- to hexacyclic hydrocarbons and identified the dominant compounds as substituted and unsubstituted monocyclic hydrocarbon and decalin, as is observed in the type I UCMs. Payzant *et al.* (1980) proposed that the tricyclic component of the UCM of Cold Lake bitumen is composed of alkyl

perhydrophenanthrenes. No structure was proposed for the tetracyclic, pentacyclic and hexacyclic component of the UCMs of Manville oils (Payzant *et al.*, 1980).

If the compounds forming type I UCM share a common origin, several hypothetical scenarios may explain their formation. These compounds appear unrelated to cyclic terpenoids and are not known biosynthetic products. Their presence may thus represent hydrogenation of polycyclic aromatic hydrocarbons (PAHs). Although many reductive techniques are known, they typically involve support from noble metal catalysis under high hydrogen pressure and/or elevated temperatures, either in the gas phase or in the presence of organic solvents (Nelkenbaum *et al.*, 2007). Saturation typically leads to the formation of additional isomers as observed by Yuan and Marshall (2005) who used a mixture of supercritical CO₂ with molecular hydrogen to form six tetradecahydroantracene isomers from the reduction of anthracene.

Alternatively, anaerobic degradation of PAHs has been demonstrated to yield hydrogenated derivatives of PAHs (Annweiler *et al.*, 2002; Mechenstock *et al.*, 2004).

Denitrification and sulfate reducing conditions can lead to the transformation of naphthalene and phenanthrene to naphthoic and phenanthroic acids (Mechenstock *et al.*, 2004). Aitken *et al.* (2004) proposed a reductive pathway involving carboxylation of naphthalene to 2-naphthoic acid, followed by several hydrogenation steps, which may account for the presence of decahydro-2-nahthtoic acid in many biodegraded oils. Additional steps to the reductive 2-naphthoic acid pathway include cleavage of a C2-fragment via beta-oxidation with another metabolite to yield carboxycyclohexylacetic acid (Mechenstock *et al.*, 2004). Alkylation following decarboxylation of the decahydro-2-nahthtoic and carboxycyclohexylacetic acid could potentially yield alkyldecalins and alkylcyclohexanes, respectively. However, little is known

about the biodegradability of higher PAH ring systems and as stated above, the PGC extracts display little evidence of biodegradation.

Lastly, Rubinstein and Strausz (1979) have proposed that alylcyclohexane and alkylmethylcyclohexane series of South Western Alberta oils were formed by the cyclization of fatty acids during diagenesis. This hypothesis was endorsed by Payzant *et al.* (1980) to explain the cyclohexane, decalin, and the perhydrophenanthrene series tentatively identified in Cold Lake bitumen. If such a process is viable for monocyclic and bicyclic alkanes, it appears increasingly unlikely to account for the formation of compounds with more than three rings. However, all but one of the tentative structures proposed here for the tetracyclic, pentacyclic and hexacyclic compounds of type I UCM can be formed from cyclization of a single aliphatic branched compound.

4.1.2. Origin of type II UCMs

Type II UCMs are dominated by acyclic and cyclic irregular head-to-head isoprenoids considered biomarkers of archaea (e.g. Chappe *et al.*, 1979 a and b, 1982; Brassell *et al.*, 1981; Petrov *et al.*, 1990; Vink *et al.*, 1998; Stefanova, 2000). Biphytane and cyclized derivatives of biphytane are the hydrocarbon skeletons formed by cleavage of GDGTs derived from archaea (e.g. Chappe *et al.*, 1982; Volkman and Maxell, 1986; Kates, 1997). Archaeal lipid contributions are highly variable in samples from the Hoyle Formation of the Porcupine Assemblage.

Archaeal lipids were not observed in the Krist Formation of the Porcupine Group. Hoyle Formation core samples OC-114m and HP12378GZ collected from the Owl Creek and Hoyle Pond Mines, respectively, display dramatically higher relative abundance of acyclic and cyclic biphytane, as well as biphytane derivatives, than the samples collected away from gold

producing centers and likely represent Late Archaean, secondary additions of organic matter from subsurface hydrothermal communities (Ventura *et al.*, 2007).

Thermophilic and methanogenic Euryarchaeota synthesize GDGTs with acyclic biphytane (e.g. caldarchaeol) as well as biphytane with up to 4 pentacyclic rings (De Rosa *et al.* 1988; Sprott *et al.*, 1997; Schouten *et al.*, 1998; Schouten *et al.*, 2000). The relative abundance of acyclic and cyclic archaeal lipids was calculated by integrating peak areas of the GC×GC TIC chromatogram. The percentage of acyclic archaeal lipids in sample OC-114m is 30%, the percentage of mono-, bi-, and tricyclic archaeal lipids was 27, 23, and 20%, respectively. The high relative abundance of cyclic compounds (70 %) relative to the acyclic compounds (30%) is consistent with hyperthermophillic archaea (Ernst *et al.*, 1998; Uda *et al.*, 2001; De Rosa *et al.*, 1980). The presence of > C₄₀ acyclic isoprenoids may indicate derivatives of H-type Caldarchaeols having C₈₀-isoprenoid carbon chains that derive from the hyperthermophilic methogen *Methanothermus fervidus* (Morri *et al.*, 1998). However, the existence of these molecules has been recently refuted (Gattinger *et al.*, 2002).

4.1.2.1. Thermal stress of type II UCM constituents

The archaeal lipids contributing to the type II and III UCM topologies have undergone substantial thermal stress resulting in the production of archaeal lipid derivatives. The distribution of archaeal lipid derivatives is consistent with a non-random pattern of cracking. The mass chromatograms of the fragment ions m/z 194 and m/z 165, which correspond to the fragmentation of cyclic archaeal lipids, do not display compounds with less than 30 carbon atoms (Fig. 12 and 13). The absence of archaeal lipid derivatives formed by cleavage of a carbon-carbon bond at the mid-section of the isoprenoid skeleton strongly suggests that

hydrocarbon cracking preferentially affects the terminal ends of the isoprenoids. Subsequently, the lower carbon number biphytane derivatives (<C₃₉) likely formed by several episodes of successive cracking of the terminal carbon atoms of biphytane and biphytane derivatives. This cracking pattern may be due, in part, to the higher thermal resistance of carbon-carbon bonds with one tertiary substituted atoms relative to that of a carbon-carbon bonds (e.g. Fierro *et al.*, 2002) as well as lower C-C bond strengths for carbon atoms located at the terminal ends of *n*-alkanes (e.g. Morrison and Boyd, 1973).

4.1.2.1. Isomerization of type II UCM constituents

The C₄₀ mono, bi-, and tricyclic biphytane derivatives and their lower C₃₆-C₃₉ homologs are each, respectively, represented by doublets, triplets, and at least quadruplets of peaks with exactly the same mass spectra (Fig. 10). The elution pattern and number of peaks suggests that they correspond to diastereomers formed by isomerization, most probably, at the alkyl-substituted chiral carbon of pentacyclic rings. Stereoisomers formed by changes in the stereochemistry of methyl substituents are not separated by the non-chiral chromatographic column used here. Biogenic cyclic biphytanes have a specific stereochemical configuration, which is *trans* for the 1,3 substitution pattern of cyclopentane rings (Sinninghe Damsté *et al.*, 2002; Montenegro *et al.*, 2003). The PGC sample monocyclic biphytanes may have both a *cis* and *trans* ring conformation, resulting in two distinct peaks. *Trans-trans*, *cis-trans*, and *cis-cis* are the three possible conformations the rings in bicyclic biphytane and may explain the triplets of peaks with identical mass spectra observed. Up to 8 different stereochemical ring conformations are possible for tricyclic biphytanes. It is can be hypothesized that some of these tricyclic diastereomers coelute, resulting in the quadruplets of peaks observed.

Nothing is known about the conditions required for the formation of diastereomers of cyclic archaeal isoprenoids. Diasteromers are commonly observed in geologic samples and are the products from either acid-clay catalyzed reactions and/or thermal stress (Rubinstien *et al.*, 1975). However, acid-clay catalyzed reactions are unlikely to have resulted in the formation the PGC archaeal lipid diastereomers. The Hoyle and Owl Creek mines were hydrothermally active (Phillips, 1986; Roberts, 1987; Kerrich *et al.*, 1987, Veizer, 1989, Brisban, 1997; Kerrich and Ludden, 2000). Illite, which later altered to mica and albite during greenschist metamorphism (Davies and Whitehead, 1993) is a weak Bronsted and Lewis acid (Tannenbaum *et al.*, 1986) and its catalyzing activity is suppressed in the presence of water (Goldstein, 1983; Tannenbaum and Kaplan, 1985; Seewald, 2003). Catalytic activity of clay involves carbonium-ion intermediaries that generate a predominance of branched products (Seewald, 2003), which is inconsistent with dominance of cyclic moieties in the PGC sample solvent-extracts.

Isomerization of a chiral ring center is less common due to the steric hindrance imposed by the ring. High temperatures associated with catagenesis causes the covalent bond between hydrogen and a tertiary substituted chiral carbon to undergo bond elongation, which if broken produces a hydrogen radical (H¹) or hydride ion (H-). Under such conditions, the broken bond may reform with equal probability that the same steric configuration is preserved or changed. Such transformations are favored at high temperatures and potentially require additions of metal catalysts (Akhmedov *et al.*, 1999), or the presence of high hydrogen pressure to increase the likelihood of isomerization without excessive hydrocarbon cracking (Seewald, 1994). These conditions are thought to be common in hydrothermal environments in which H₂O disproportionates forming H⁺ and OH buffered solutions (Rullkötter *et al.*, 1984; Seewald, 1994, 2003; Lewan, 1997; Price and DeWitt, 2001). It is noteworthy, that within the PGC, exogenous

hydrogen may have formed by serpentinization of mafic and ultramafic rocks located lower in the stratigraphic profile (Fyon *et al.*, 1983).

5. Conclusions

The questions driving Precambrian organic geochemistry not only includes what constitutes syngeneity and to what end can syngeneity be established, but also to what range of depositional environments and physical and chemical burial conditions is sedimentary organic matter preserved. The accelerated crustal development of Earth's early history likely resulted in hydrothermal processes being common features of Archean depositional environments. Understanding how the composition and preservation of organic matter is affected by these processes can help determine the diversity of geological localities that can yield extractable hydrocarbons.

This study analyzed bitumens extracted from hydrothermally altered late Achaean sediments. The s/u HC fractions of these bitumens contain three different UCM topologies, which were irresolvable by traditional GC-MS. The increased separation power of GC×GC-MS facilitated the tentative identification of multiple compound classes that were embedded within these UCMs. Type I UCMs are composed of mono- to hexacycloalkanes of unknown origin. Type II UCMs are mostly composed of C₃₅-C₄₀ archaeal lipids. Type III UCMs represent a mix of Type I and II UCM forming compounds.

This investigation demonstrates that the analysis of UCMs can provide important information about the biological source and diagenetic history of the sedimentary organic matter.

The II UCMs appear to result from the diagenetic alteration of biogenic material via a coupled process of isomerization and hydrocarbon cracking. The isomerization of these compounds was likely caused by thermal stress in the presence of high hydrogen pressure. Such conditions are common within hydrothermal environments.

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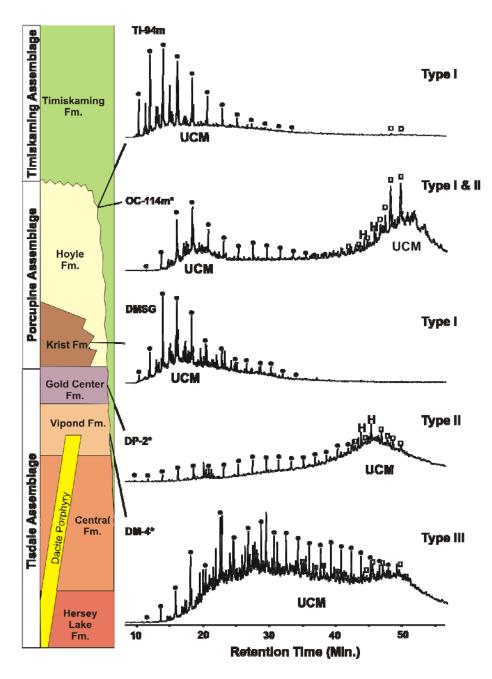


Figure 1. Stratigraphic profile of the PGC (left). Total ion current (TIC) chromatograms of the extractable hydrocarbon fraction of representative samples Tisdale and Porcupine Group (right). Unresolved complex mixtures (UCMs) are observed in samples from all the formations and are labeled "Type 'I-III" in reference to their topology (see text). Filled circles indicate *n*-alkanes. Number above peak indicates numbers of carbon atoms in the compound. H indicates hopanes. Open squares indicate acyclic irregular isoprenoids, biphytane and derivatives.

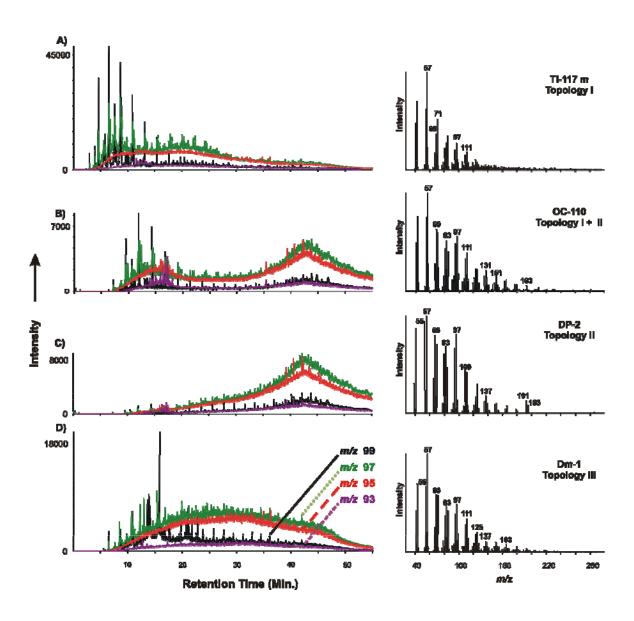


Figure 2. Mass chromatograms of fragment ions m/z 99, m/z 97, m/z 95 and m/z 93 (left) and mass spectra (right) of samples (A) TI-117 and (B) OC-110m of the Hoyle Formation; (C) DP-2 of the Gold Center Formation, and (D) DM-1 of the Vipond Formation.

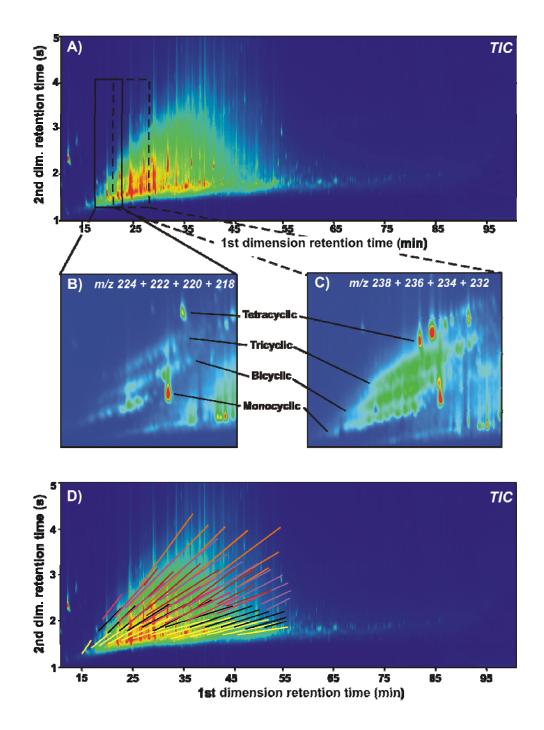


Figure 3. A) GC×GC-MS TIC chromatogram from the s/u HC fraction of the Krist Formation sample DMSD (Type I UCM). B) Partial, summed mass chromatogram of the fragment ions m/z 224+222+220+198 showing C₁₆ mono-, bi-, tri-, and tetracyclic compounds. C) Partial, summed mass chromatogram of the fragment ions m/z 238+236+234+232 showing C₁₇ mono-, bi-, tri-, and tetracyclic compounds. D) GC×GC-MS TIC chromatogram from the s/u HC fraction of the Krist Formation sample DMSD (Type I UCM) showing the lineaments of mono-, di-, tri-, tetra-, penta- and hexacyclic naphtenes.

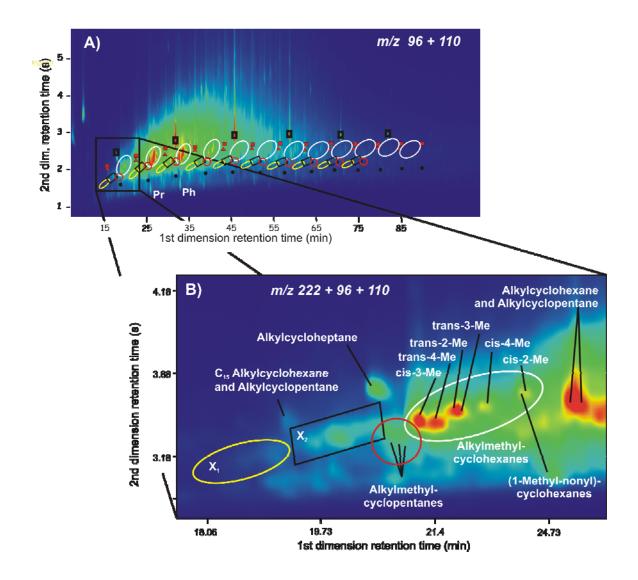


Figure 4. A) and B) GC×GC-MS mass chromatogram of the fragment ion m/z 224+96+110 of the s/u HC fraction of the Krist Formation sample DMSD. n-alkanes are labeled with black dots, alkylcyclopentanes and alkylcyclohexanes are marked by red triangles and squares, respectively. Tentatively identified homologous series of alkylethylcyclohexanes, alkylmethylcyclopentanes, alkylmethylcyclohexanes, and n-alkycycloheptanes are marked by yellow ellipses, black rectangles, red circles, larger white ellipses, small black boxes, respectively.

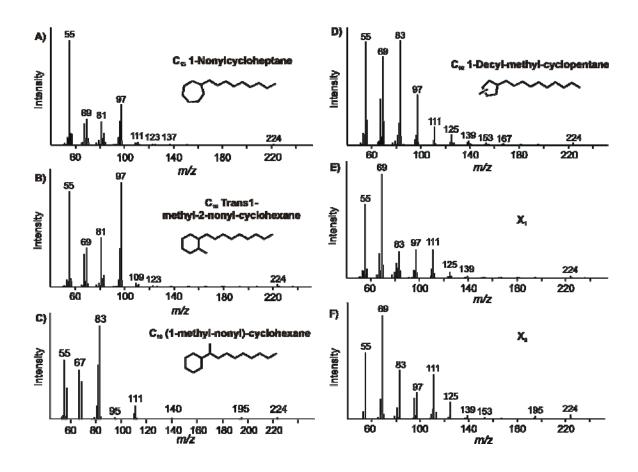


Figure 5. Mass spectra of tentatively identified C_{16} A) alkylcycloheptane, B) 1-methylalkylcyclohexane, C) the trans-2-alkylmethylcyclohexane D) alkyldimethylcyclopentane, and E and F) unidentified monocyclic labeled X_1 and X_2 , respectively.

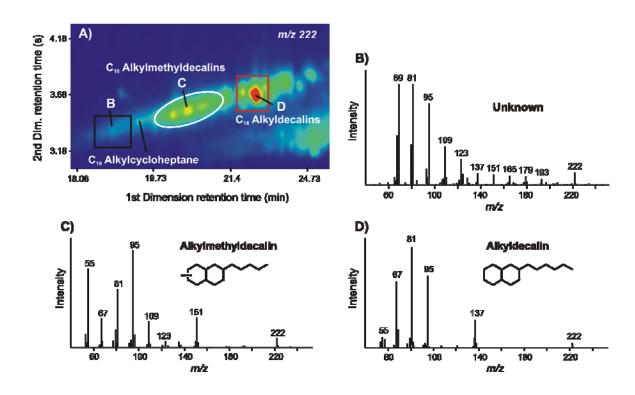


Figure 6. A) GC×GC-MS partial mass chromatogram of the fragment ion m/z 222 of the s/u HC fraction of the Krist Formation sample DMSD, B), C), and D) are mass spectra of peaks B, C, and D, respectively, in the mass chromatogram. C) and D) are representative mass spectra of peaks tentatively identified as alkylmethyldecalin and alkyldecalins.

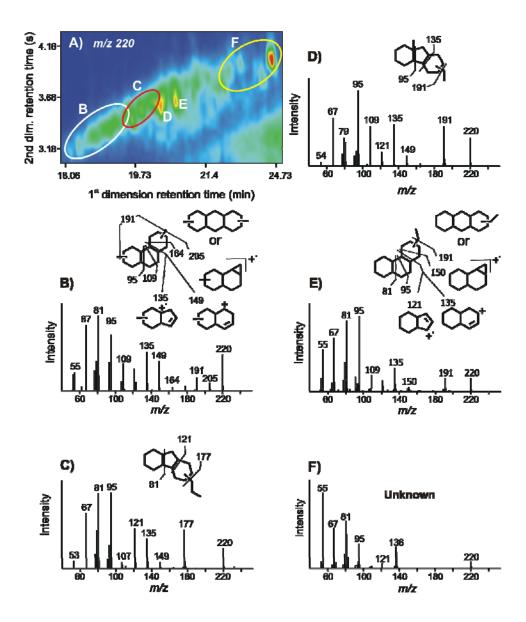


Figure 7. A) GC×GC-MS partial mass chromatogram of the fragment ion m/z 220 of the s/u HC fraction of the Krist Formation sample DMSD showing the distribution of C₁₄ tricyclic compounds. Letters B), C), D), and F) are mass spectra for peaks inscribed by white (B), red (C) and yellow (F) ellipses and peaks D and E, respectively of the m/z 220 mass chromatogram. Possible fragmentation patterns are provided for B) and E) mass spectra.

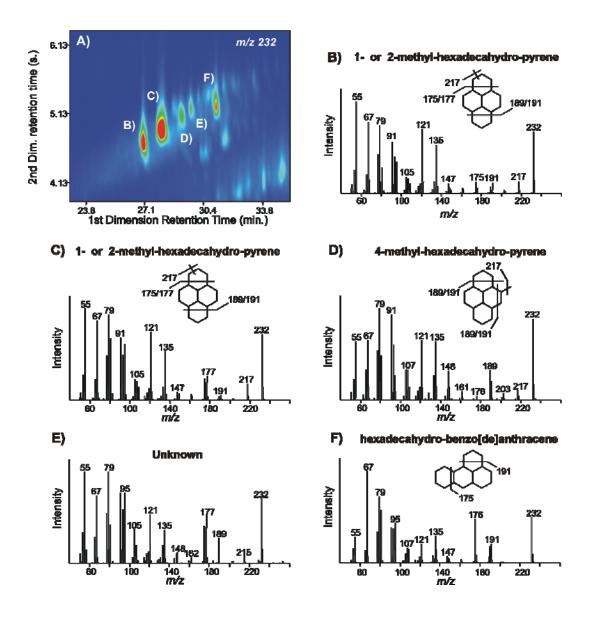


Figure 8. GC×GC-MS partial mass chromatogram of the fragment ion m/z 232 of the s/u HC fraction of the Krist Formation sample DMSD showing the distribution of C_{17} tetracyclic compounds. Letters marking the order of mass spectra correspond to the peaks labels of the mass chromatogram.

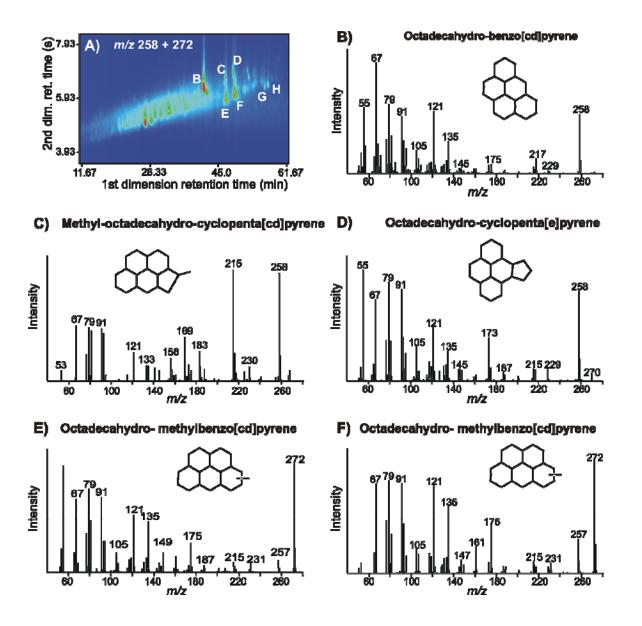


Figure 9. GC×GC-MS partial, summed mass chromatogram of the fragment ion m/z 258 + 272 of the s/u HC fraction of the Krist Formation sample DMSD showing the distribution of C_{19} and C_{20} pentacyclic compounds. Letters marking the order of mass spectra correspond to the peaks labels of the mass chromatogram.

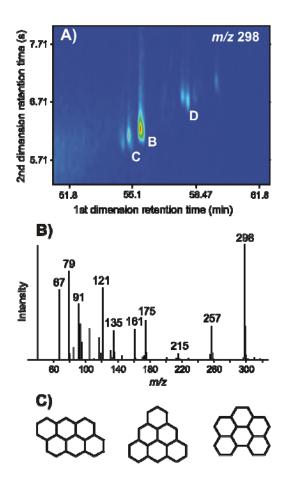


Figure 10. GC×GC-MS partial mass chromatogram of the fragment ion m/z 298 of the s/u HC fraction of the Krist Formation sample DMSD showing the distribution of C₂₂ hexacyclic compounds. B) Mass spectra of peak B. C) Tentative structures of hexacyclic compound corresponding to peak C, B and D.

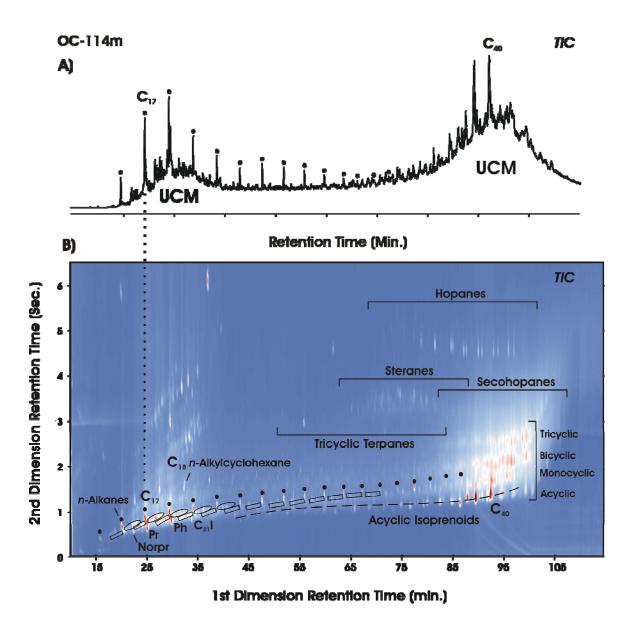


Figure 11. A) GC-MS TIC chromatogram of the s/u HC fraction of the bitumen from sample OC-114m (of the Hoyle Formation) containing a type I and II UCM. B) GC×GC-MS TIC chromatogram of the same sample with the labeled elution dimension of *n*-alkanes (black circles), mono-, bi-, tri-, tetra- (steranes), and pentacyclic (hopanes). Dotted line indicates irregular isoprenoids. Boxes encapsulate monomethylalkanes. Open ovals enclose monoethylalkanes.

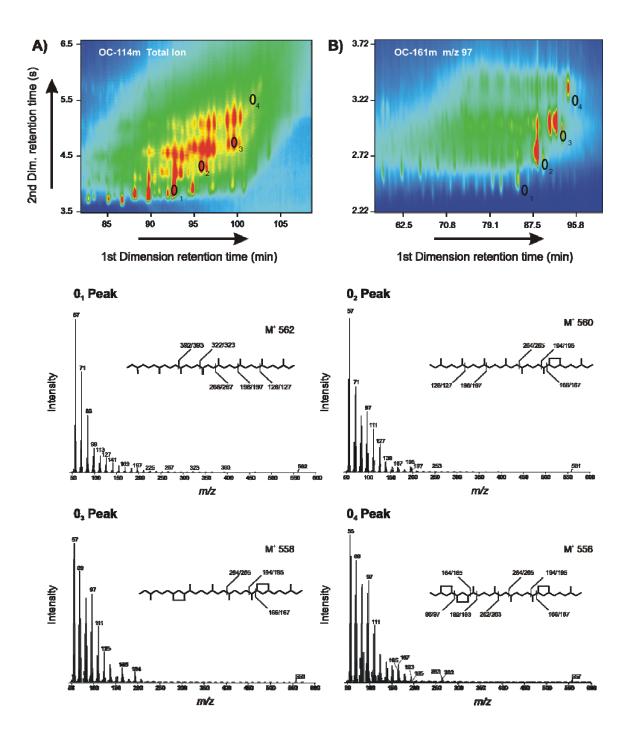


Figure 12. A) GC×GC-MS TIC chromatogram of sample OC-114m. C_{40} acyclic, mono-, bi-, and tricyclic biphytanes are labeled with black ovals. B) The m/z 97 GC×GC mass chromatogram of sample OC-161m coinjected with an archaeal lipid standard containing C_{40} acyclic, mono-, bi-, and tricyclic biphytanes. Below are mass spectra of peaks 0_1 - 0_4 corresponding to the C_{40} mono-, bi-, and tricyclic biphytane in the OC-144m sample s/u HC fraction.

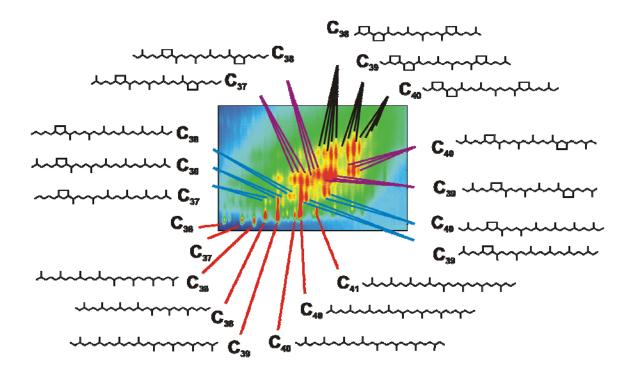


Figure 13. GC×GC-MS TIC chromatogram of the s/u HC fraction of Hoyle Formation sample OC-114m. Biphytane and acyclic derivatives of biphytane are labeled with red lines. The structures of mono-, bi-, and tricyclic biphytanes and their derivatives are marked by blue, purple, and black lines, respectively. Branched lines leading to the sample structures identify peaks in the chromatogram having the same mass spectra.

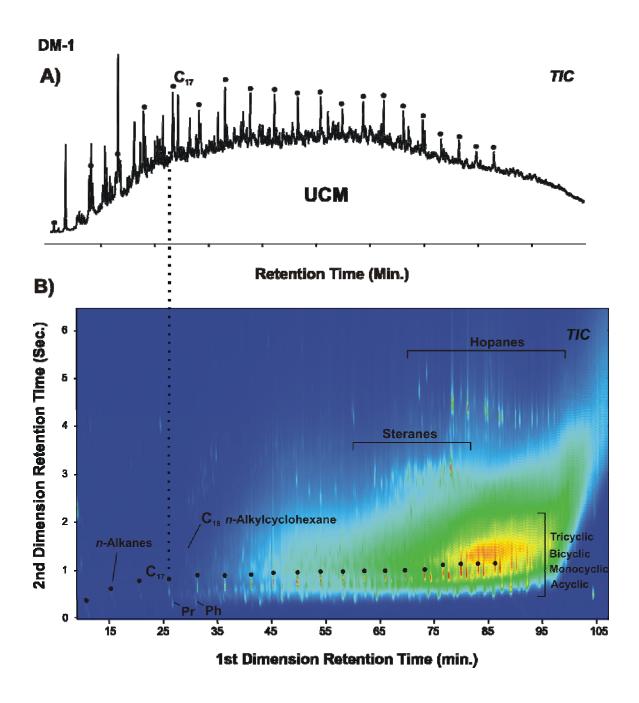


Figure 14. A) TIC chromatogram of the s/u HC fraction of the bitumen from sample DM-1 (of the Viapond Formation) containing a type III UCM. B) $GC \times GC$ -MS TIC chromatogram of the same sample with the labeled elution dimension of n-alkanes (black circles), mono-, bi-, tri-, tetra- (steranes), and pentacyclic (hopanes).