

Soil heterotrophic respiration in *Casuarina equisetifolia* plantation at different stand ages

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Abstract: The soil respiration rates (R_h) in 6-year-old (young), 17-year-old (middle-age), 31-year-old (mature) *Casuarina equisetifolia* coastal plantations were measured using an LICOR-8100 automated soil CO₂ flux system from May 2006 to April 2007. Results show that R_h displayed an obvious seasonal pattern across the observed years. The maximum values of R_h occurred at June and July and the minimum at December and January. Soil temperature and soil moisture as well as their interaction had significant effects on the monthly dynamics of R_h . The analysis by one-way ANOVA showed that R_h had a significantly exponential relation ($p < 0.05$) to soil temperature at soil depth of 5 cm, and had a linear relation ($p < 0.05$) to soil water content of the upper 20 cm. The result estimated by the two-factor model shows that soil temperature at soil depth of 5 cm and soil moisture at soil depth of 20 cm could explain 68.9%–91.9% of seasonal variations in R_h . The order of R_h rates between different stand ages was middle-age plantation > mature plantation > young-age plantation. With the increase of growth age of plantation, the Q_{10} of R_h increased. The contribution of R_h to total soil surface CO₂ flux was 71.89%, 71.02% and 73.53% for the young, middle-age and mature plantation, respectively. It was estimated that the annual CO₂ fluxes from R_h were 29.07, 38.964 and 30.530 t·ha⁻¹·a⁻¹ for the young, middle-age and mature plantation, respectively.

Keywords: soil heterotrophic respiration; coastal plantation; stand age; *Casuarina equisetifolia*

Introduction

Soil respiration rates (R_s), or surface soil CO₂ efflux, is the primary pathway of soil carbon release to the atmosphere (Raich & Tufekcioglu 2000). In recent decades, R_s has received more attention because it plays a critical role in global carbon cycle and in global warming feedback (Fang & Wang 2007). The R_s is generally defined as two components, autotrophic respiration (R_a)

and heterotrophic respiration (R_h) (Kuzyakov & Larionova 2005). In carbon cycle of terrestrial ecosystem, R_h is the main process of soil carbon loss to the atmosphere. Thus, quantifying contribution of R_h to total soil CO₂ effluxes is of great importance and prerequisite for estimating the carbon balance in terrestrial ecosystems (Hanson et al. 2000; Bond-Lamberty et al. 2004).

Most of the previous studies about R_h focus on the inland forest ecosystems (Hanson et al. 2000; Ngao et al. 2007; Saurette et al. 2008; Gaumont-Guay et al. 2008), and very few concerns coastal forest (Tyree et al. 2006). As so far, no available information has been found on R_h in coastal forest ecosystem in China. In coastal regions, forest plantation plays an important role in wind combating, sand fixation and local environment improvement. As the main plantation species, *Casuarinaceae* is widely distributed in southern China, especially in Zhejiang, Fujian, and Guangdong provinces. It was reported that the area of *Casuarina equisetifolia* plantations has been up to 3×10^5 ha in China (Zhong & Zhang 2003).

In this study, the R_h of *C. equisetifolia* plantation at 3 different growth stages was separated and measured in 2006 and 2007. The main points of the present study were addressed as: (1) seasonal dynamics of R_h ; (2) effects of temperature, soil moisture and forest age on R_h ; (3) contributions of R_h to total soil surface CO₂ flux; and (4) annual fluxes of R_h in the plantations at three different growth stages.

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Materials and method

Site description

The experiment site was situated in Chihu State-owned shelter forest of Huian county in Fujian Province, China (latitude 24°30′–24°55′N and longitude 118°40′E–118°55′E). The study site is at elevation of 5–400 m, with an average slope of <10°. The climate is maritime monsoon climate, with mean annual air temperature of 19.8 °C. The frost-free period is 320 d. Annual

mean precipitation is 1000–1500 mm, with annual mean evaporation of 2000 mm. The typical soil is homogeneity aeolian sand soil with low nutrients, and the layer thickness is 80–100 cm. The forest communities are dominated by *C. equisetifolia* pure plantations with few shrubs and grasses.

Three plantations in different growth stages (6-year-old, P6), (17-year-old, P17), and (31-year-old, P31) were chosen for the study. In each plantation, a plot of 20 m×20 m in size was set up. Before measurement of soil respiration, the characteristics of forest community and soil physiochemical properties were investigated (Table 1).

Table 1. The characteristics of forest community and soil physiochemical properties in *C. equisetifolia* plantations at three different stand ages

Stand ages	Community			Soil physiochemical properties (0–20 cm)				
	Density (trees·ha ⁻¹)	Average DBH (cm)	Average tree height (m)	pH	Soil bulk density (g·cm ⁻³)	SOC (g·kg ⁻¹)	TN (mg·kg ⁻¹)	C/N
P6(6-year-old)	2500	5.5	9.5	5.47±0.08	1.22±0.05	7.14±1.14	72.72±3.81	98.21±5.62
P17(17-year-old)	2202	12.2	10.8	5.24±0.05	1.24±0.05	12.00±1.65	76.90±4.12	156.05±7.73
P31(31-year-old)	1962	15.1	13.7	5.10±0.05	1.37±0.09	12.43±1.69	81.15±4.26	153.27±7.32

Measurement methods

In each plantation, the roots were excluded in three subplots (1 m×1 m) by using trenching method. Moreover, the total surface soil CO₂ effluxes were measured in one subplot (50 cm×50 cm).

R_s and R_h were measured at a certain day in the middle of every month from May 2006 to April 2007. Measurements were at hourly intervals throughout a daytime from 08:00 a.m. to 18:00 p.m. Every measurement lasted for 2 min with 3 replicates. In addition, four observations on daily variations (beginning at 08:00 in the morning and ending at 08:00 of next morning) were carried out for each site on July 2006, October 2006, January 2007 and April 2007, respectively.

R_s and R_h were measured using an LI-8100 Automated Soil CO₂ Flux System (LiCor Inc., Lincoln, NE, USA) equipped with a short-term chamber. Integrated pump circulates the headspace air from the chamber to the non-dispersive infrared (NDIR) gas analyzer during the closed state of the chamber and the CO₂ concentration data as well as the calculated flux rates are recorded in the system. A PVC soil collar (inner diameter 20.4 cm×7.5 cm tall) was inserted into soil at 5-cm depth on May 2006 in every R_s and R_h point, and left in the same locations throughout the study period. Meanwhile, all living vegetation aboveground in the plots was removed (don't disturb the litter layer) and the plots were kept free of living plants during the study period.

Simultaneously, air temperature (T_a), surface soil temperature (T_0), soil temperature at a depth of 5 cm (T_5) and soil water content at a depth of 20 cm were measured by a DHM2 ventilated thermometer, convoluted tubule geo-temperature table, the temperature probe belong to the LI-8100 and drying method, respectively.

Data analysis

Daily average flux of R_h according to the daily variations in 24 h

was observed in the four days on July 2006, October 2006, January 2007 and April 2007, respectively. It was found that the effluxes measured during 10:00–12:00 in the morning can basically represent for the daily average flux of R_h . Kessavalou et al. (1998), Du et al. (2006) and Jin et al. (2007) also used similar treatments in studying the seasonal dynamics of soil greenhouse gases. Therefore, in this study, we used the R_h rates during 10:00–12:00 in the morning to represent for the R_h daily average flux in the other eight months.

The relationship between R_h and temperature (T) was simulated with an exponential equation and Q_{10} value (defined as the increase in respiration rate per 10°C increase in temperature).

$$R_h = ae^{bT}, Q_{10} = e^{10b} \quad (1)$$

where, a is the basic respiratory (the respiration rate when temperature at 0 °C), b is the temperature sensitive coefficient. The relationship between R_h and soil surface water content (W , volume to volume, 0–20 cm) is simulated with a linear equation.

$$R_h = cW + d \quad (2)$$

where, c and d are the regression coefficients.

Correlation analysis showed that R_h had a better correlation with T_5 than that of T_a and T_0 . So the dependence of T_5 and soil water content (W) on R_h was analyzed by a nonlinear regression model as follows.

$$R_h = a \times e^{bT} \times W^c \quad (3)$$

where, T is the soil temperature at 5-cm soil depth, W is the soil water content, and b and c are regression coefficients.

$$H_c = R_h / R_s \quad (4)$$

where, H_C is the contribution of R_h to total soil surface CO_2 flux.

Calculating of annual CO_2 flux of R_h : firstly, the average daily values of T_a (from an automatic weather station of the forest center) were transformed to T_5 according to their significant correlations and the average monthly values of T_5 were calculated; secondly, combining with the average monthly value W (3 days per month for each forest type), T_5 was inputted into the nonlinear model (Equation (3)) to get the CO_2 emission of R_h in every month during the study period, and finally, the fluxes in the 12 months were cumulated to get the annual flux of R_h .

We used SPSS 13.0 (SPSS Inc. 2001) to do the statistical analyses. Relationships between R_h and T_a , T_0 , T_5 and W , and the differences among 3 plots were analyzed using the One-way ANOVA. Two-factor regression analyses were made under the nonlinear module. Graphs were prepared by using Excel (Microsoft Corp. 2003).

Results

Seasonal dynamics of R_h

The R_h rate in *C. equisetifolia* plantation had an obvious seasonal pattern, with a similar single peak curve for each of three growth stages (Fig. 1). In March and April, soil CO_2 efflux stayed at a low level, and increased slowly with the increase of microbial activity resulted by the risen temperature and the improvement of soil water content. It reached the maximum at June (for P17) or July (for P6 and P31) and decreased gradually with the minimum at December (for P6 and P17) or January (for P31). The coefficients of seasonal variation were 23.68%, 36.96% and 30.46% for P6, P17 and P31, respectively.

The mean R_h rate ($\mu mol CO_2 m^{-2} s^{-1}$, mean \pm SE) in the three

plots was ranked as: P17 (2.76 ± 0.085) > P31 (2.17 ± 0.055) > P6 (2.00 ± 0.039). Mean multiple comparisons in SPSS, at the 0.05 significant level, showed that there was a significant difference of R_h rates between P6 and P17 ($p=0.018$), but no significant difference was found between P17 and P31 ($p=0.060$), and between P6 and P31 ($p=0.593$).

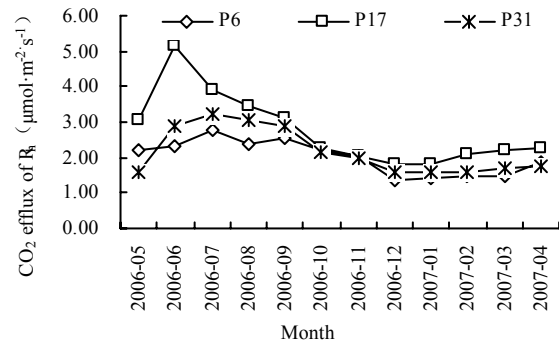


Fig. 1 Seasonal variations of R_h rates for the *C. equisetifolia* plantations at three different stand ages

Models of R_h to temperature and soil surface water

In this paper, T_a , T_0 , T_5 and W were measured and their effects on R_h were analyzed (Table 2). The influences of temperature on R_h were different among T_a , T_0 and T_5 . The R_h rates had no significant correlations with T_a and T_0 for the three plots except for the relationship between T_a and R_h in the young-age plantation, but had significant relationships with T_5 in all the three growth stages ($P<0.05$). T_5 could account for 51.73%– 62.87% of R_h variations.

Table 2. Models of R_h to temperature and surface soil water content for the *C. equisetifolia* plantations at three different stand ages

Stand age	$R=a e^{bT}$			$R=aW+b$	$R=a \times e^{bT} \times W^c$			
	T_a	T_0	T_5		a	b	c	R^2
	R^2	R^2	R^2					
P6(6-year-old)	0.5309 ($p=0.082$)	0.5908 ($p=0.116$)	0.6287 ($p=0.000$)	0.5926 ($p=0.000$)	1.973	0.014	0.114	0.689 ($p=0.000$)
P17(17-year-old)	0.3239 ($p=0.002$)	0.4510 ($p=0.067$)	0.5173 ($p=0.000$)	0.7865 ($p=0.003$)	3.731	0.020	0.422	0.919 ($p=0.000$)
P31(31-year-old)	0.4750 ($p=0.157$)	0.5806 ($p=0.719$)	0.6230 ($p=0.000$)	0.8207 ($p=0.001$)	2.262	0.038	0.552	0.901 ($p=0.000$)

Notes: all models were tested at the significance level of 0.05 (two-tailed). T_a , T_0 , T_5 indicate air temperature, surface soil temperature, soil temperature at 5-cm depth, respectively. W means soil water content at 20-cm depth. The temperature indicator in the two-factor model was based on T_5 .

In addition, the Q_{10} values of R_h based T_5 were 1.39, 1.44, and 1.95 for P6, P17 and P31, respectively, which showed that the temperature sensitivity of microbial respiration improved with the increase of stand ages. Kirschbaum (2006) and Fan et al. (2008) summarized the influencing factors of R_h , and pointed out that temperature, moisture and the soil substrates were the main factors. In our study, no obvious differences between temperature and soil moisture conditions were found among *C. equisetifolia* plantation at three stand ages; furthermore, soil organic carbon

contents were similar. Soil heterotrophic respiration is the decomposition process of soil organic matter and litter dominated by soil microbial. Hence we hypothesized that the differences of Q_{10} values in our study were mostly likely caused by different temperature sensitivity of litter decomposition in the plantations at different stand ages. Of course, the hypothesis must be tested and confirmed in subsequent study.

The linear models showed that the R_h rate had significant correlation with surface soil water content. Soil water content ac-

counted for 59.26%–82.07% of R_h variations, which stated clearly that the effect extent of soil moisture on microbial respiration was larger than that of temperature because of the sandy soil of coastal protective forest.

The biological processes in R_s can be considerably influenced by soil temperature and moisture, and their interactions correlated to R_s must be considered, so do R_h (Zhang et al. 2005; Wang & Yang 2007), which also was confirmed in our study. Soil temperature at soil depth of 5 cm and moisture at soil depth of 20 cm had better prediction on R_h than using respective independent variable, and R^2 values were 0.689, 0.919 and 0.901 for P6, P17 and P31, respectively.

Contribution of R_h to total surface soil CO_2 efflux and its seasonal pattern

In P6, the H_C varied from 65.50% to 78.56%, with an average of 71.89%; while in P31, the H_C varied from 66.96% to 81.26%, with the average of 73.53%. In P17, the H_C varied from 64.82% to 77.74%, with an average of 71.02% (Fig. 2). However P31 had higher H_C than P6 and P17 in most months, and there was no significant difference of H_C among the three plantations at different growth ages at 0.05 significant level ($P=0.326$).

The seasonality of H_C showed a similar pattern for the 3 plots (Fig. 2). From May 2006 to August 2006, the H_C maintained at a low level with an average of 67.51%. H_C increased and peaked on February–March 2007, which showed a reverse seasonality pattern comparing to the air temperature.

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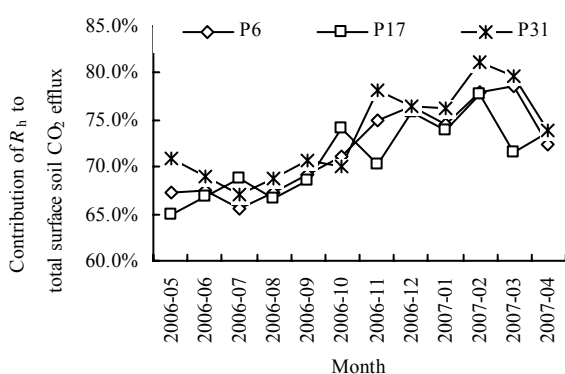


Fig. 2 Seasonal dynamics of the contribution of R_h to total surface soil CO_2 efflux in the *C. equisetifolia* plantations at different growth ages (P6 means 6-year-old; P17, 17-year-old; P31, 31-year-old)

The seasonality of H_C showed a similar pattern for the plantations at three different stand ages (Fig. 2). From May 2006 to August 2006, the H_C maintained at a low level with an average of 67.51%. H_C increased and peaked on February–March 2007,

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Annual soil CO_2 flux of R_h

The results showed that T_a was significantly correlated to T_s , and the average of R^2 values in the plantations at three different growth ages reached 0.92 ($p<0.05$), so the air temperature could be converted to the soil temperature at a soil depth of 5 cm. According to the method mentioned above, we estimated the annual soil CO_2 effluxes of R_h and the values were 29.072, 38.964 and 30.530 $\text{t}\cdot\text{ha}^{-1}\cdot\text{a}^{-1}$ for the young, middle age and mature plantations, respectively (Table 3).

Table 3. Linear relationships between T_s and T_a and annual CO_2 flux of R_h of the *C. equisetifolia* plantations at three different stand ages ($\text{t}\cdot\text{ha}^{-1}\cdot\text{a}^{-1}$)

Stand age	Linear relationship between T_s and T_a	Annual CO_2 flux of R_h
P6(6-year-old)	$y = 0.927x + 1.724$ ($R^2=0.9333$, $P<0.05$, $n=36$)	29.072
P17(17-year-old)	$y = 0.927x + 1.724$ ($R^2=0.9333$, $P<0.05$, $n=36$)	38.964
P31(31-year-old)	$y = 0.895x + 2.851$ ($R^2=0.8850$, $P<0.05$, $n=36$)	30.530

Discussion

R_h and the environmental variables

Our study shows that the mean R_h rate in *C. equisetifolia* plantations was $2.31\mu\text{mol CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, which was higher than that of native *Castanopsis kawakamii* forest ($1.88\mu\text{mol CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$), monoculture *Castanopsis kawakamii* plantations ($0.82\mu\text{mol CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) and Chinese fir plantations ($0.73\mu\text{mol CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) in Fujian Province (Yang et al. 2006). Zhou et al. (2000) reported that the mean R_h rate was $0.94\mu\text{mol CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in deciduous broad-leaved forest and $1.30\mu\text{mol CO}_2\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in evergreen broad-leaved forest.

Many reports about soil respiration pointed out that temperature and moisture were the main corresponding environmental factors (Fang & Moncrieff 2001; Rodeghiero & Cescatti 2005). R_h is also influenced by both climatic factors and human activity (Jia et al. 2005). Temperature is the main controlling factor of R_h in terrestrial ecosystems and would be regulated by soil moisture (Davidson et al. 2000). When soil moisture was the limiting factor, R_h was not controlled by temperature, and instead, soil moisture became the major driver of R_h . This phenomenon was confirmed by Kucera & Kirkham (1971), Zhou et al. (2004) and Raich & Mora (2005). In this study, air and soil temperatures were maintained at a relatively high level and had little variation, but soil moisture in the coastal sandy soils was very low during the study period (averaged 4.33%). Therefore, surface soil moisture was regarded as the limiting factor of R_h and presented a high relationship with soil heterotrophic respiration. The maximums of R_h appeared on June or July when there were better soil

moisture conditions, compared with soil moisture on August with the highest air temperature.

Contribution of R_h to total soil CO_2 flux

The contributions of R_h to total surface soil CO_2 flux varied greatly with different plantations and sampling time. Raich & Tufekcioglu (2000) revealed that H_C had a low value in cold biomes, ranging from 7% to 50% in arctic tundra and 11% to 38% in boreal forests. In temperate regions, the H_C ranged from 50% to 67% in broad-leaved forests from 38% to 65% in pine forests. Therefore, the trend of increase of H_C with temperature increasing in macro-scale is quite clearly. The annual mean air temperature is 19.8°C at this field, so the average H_C at 72.15% in the three *C. equisetifolia* plantations was reasonable.

However trenching method was a mature approach to separate R_a and R_h , which also could affect R_h and its contribution to total surface soil CO_2 flux; for example, Uchida et al. (1998) had pointed that, in the trenched plots, the decomposition of the root residuals could largely enhance the soil heterotrophic respiration and elevate the contribution of soil microbial respiration to total soil CO_2 effluxes. Moreover, temperature and soil moisture would change in the trenched plots, which could also alter the soil microbial activity (Wang & Yang 2007). So far, however, the estimated range of trench effect and how to reduce or even eliminate the trench effect have been not well studied.

In the *C. equisetifolia* plantations, the H_C varied at different growth seasons. In winter and spring, H_C presented a high value; while in summer and autumn, H_C was decreased. The results were agreed with the findings by Zhang XQ et al. (2005) and Chang et al. (2007). Li et al. (2000) reported that soil respiration was mainly contributed by the decomposition of soil organic matter when air temperature was below 15°C; but when air temperature was above 15°C, the contribution of plant root respiration was larger than microbial respiration. In our study, the H_C decreased when air temperature and soil temperature increased. This phenomenon may be attributed to the booming of root metabolic activity and the higher Q_{10} of the root respiration than that of microbial respiration (Boone et al. 1998; Epron et al. 2001). In addition, the higher H_C in winter and spring was partly due to the frequent windy weathers which increased the litter of branches.

Effects of stands at different ages on R_h

Previous studies about the influence factors of R_h in forest communities focused upon soil temperature and moisture (Wang & Yang, 2007; Bauer et al. 2008; Gaumont-Guay et al. 2008; Graf et al. 2008), the extent and mechanism of the effect of stand age on R_s and R_h were not well known. Saurette et al. (2008) reported that there was no significant difference of R_h between plantations at different growth ages. However, the significant difference of R_h was found in this study ($p < 0.05$), the R_h was highest in the middle-age plantation and then followed by the mature plantation. The lowest R_h rate occurred at the young-age plantation. Different growth ages of plantation can influence R_s and R_h by the type

and number of microbial communities (Gao et al. 2007), above-ground biomass (Chapin & Ruess 2001), Quantity and quality of litters (Liu & Fang 1997) and soil quality (Shao et al. 2005). The differences of R_h rates among different forest communities that had similar climatic conditions and soil environment were mainly driven by different quantities and quality of soil respiratory substrates (Ma et al. 2003; Yang et al. 2005). In our study, the *C. equisetifolia* plantations at three stand ages have the same meteorological conditions and soil physiochemical properties; especially there were no obvious differences of soil organic carbon contents which had low background levels and low contribution of soil mineral respiration to soil heterotrophic respiration. Moreover, the litter quality among the three plantations was similar. Thus, we hypothesize that the differences of R_h rates between the plantations at three growth ages may be attributed to the different litterfall quantity (especially different quantities of fresh litter). This hypothesis need proof in the future study.

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References

- Bauer J, Herbst M, Huisman JA, Weiermueller L, Vereecken H. 2008. Sensitivity of simulated soil heterotrophic respiration to temperature and moisture reduction functions. *Geoderma*, **145**(1-2): 17–27.
- Bond-Lamberty B, Wang CK, Gower ST. 2004. A global relationship between the heterotrophic and autotrophic components of soil respiration? *Global Change Biology*, **10**(10): 1756–1766.
- Boone RD, Nadelhoffer KJ, Canary JD, Kaye JP. 1998. Roots exert a strong influence on the temperature sensitivity of soil respiration. *Nature*, **396**(6711): 570–572
- Chang Jianguo, Liu Shirong, Shi Zuomin, Chen Baoyu, Zhu Xueling. 2007. Soil respiration and its components partitioning in the typical forest ecosystems at the transitional area from the northern subtropics to warm temperate, China. *Acta Ecologica Sinica*, **27**(5): 1791–1802. (In Chinese with English Abstract)
- Chapin FS, Ruess RW. 2001. Carbon cycle - The roots of the matter. *Nature*, **411**(6839): 749–752.
- Davidson EA, Verchot LV, Cattanio JH, Ackerman IL, Carvalho JEM. 2000. Effects of soil water content on soil respiration in forests and cattle pastures of eastern Amazonia. *Biogeochemistry*, **48**(1): 53–69.
- Du R, Lu DR, Wang GC. 2006. Diurnal, seasonal, and inter-annual variations of N_2O fluxes from native semi-arid grassland soils of Inner Mongolia. *Soil Biology & Biochemistry*, **38**(12): 3474–3482.
- Epron D, Le Dantec V, Dufrene E, Granier A. 2001. Seasonal dynamics of soil carbon dioxide efflux and simulated rhizosphere respiration in a beech forest. *Tree Physiology*, **21**(2-3): 145–152.
- Fan Zhiping, Wang Hong, Deng Dongzhou, Sun Xuekai, Gao Jungang, Zeng Dehui. 2008. Measurement methods of soil heterotrophic respiration and key factors affecting the temperature sensitivity of the soil heterotrophic respiration. *Chinese Journal of Ecology*, **27**(7): 1221–1226. (In Chinese)

- Fang C, Moncrieff JB. 2001. The dependence of soil CO₂ efflux on temperature. *Soil Biology & Biochemistry*, **33**(2): 155–165.
- Fang Jingyun, Wang Wei. 2007. Soil respiration as a key belowground process: issues and perspectives. *Acta Phytoecologica Sinica*, **31**(3): 345–347. (In Chinese)
- Gao Yongjian, Yuan Yuxin, Liu Siwei, Wang Ying, Liu Linlin. 2007. The effects of different stand ages on microbe populations and enzyme activities. *Chinese Agricultural Sciences Bulletin*, **23**(7): 185–189. (In Chinese)
- Gaumont-Guay D, Black TA, Barr AG, Jassal RS, Nesic Z. 2008. Biophysical controls on rhizospheric and heterotrophic components of soil respiration in a boreal black spruce stand. *Tree Physiology*, **28**(2): 161–171.
- Graf A, Weihermuller L, Huisman JA, Herbst M, Bauer J, Vereecken H. 2008. Measurement depth effects on the apparent temperature sensitivity of soil respiration in field studies. *Biogeosciences*, **5**(4): 1175–1188.
- Hanson PJ, Edwards NT, Garten CT, Andrews JA. 2000. Separating root and soil microbial contributions to soil respiration: A review of methods and observations. *Biogeochemistry*, **48**(1): 115–146.
- Jia Bingrui, Zhou Guangsheng, Wang Yufeng, Wang Yuhui. 2005. Affecting factors of soil microorganism and root respiration. *Chinese Journal of Applied Ecology*, **16**(8): 1547–1552. (In Chinese)
- Jin Zhao, Qi Yuchun, Dong Yunshe. 2007. Diurnal and seasonal dynamics of soil respiration in desert shrubland of *Artemisia Ordosica* on plateau of Inner Mongolia, China. *Journal of Forestry Research*, **18**(3): 231–235.
- Kessavalou A, Doran JW, Mosier AR, Drijber RA. 1998. Greenhouse gas fluxes following tillage and wetting in a wheat fallow cropping system. *Journal of Environmental Quality*, **27**(5): 1105–1116.
- Kirschbaum MUF. 2006. The temperature dependence of organic-matter decomposition - still a topic of debate. *Soil Biology & Biochemistry*, **38**: 2510–2518.
- Kucera CL, Kirkham DR. 1971. Soil respiration studies in tallgrass prairie in Missouri. *Ecology*, **52**(5): 912–915.
- Kuzyakov Y, Larionova AA. 2005. Root and rhizomicrobial respiration: A review of approaches to estimate respiration by autotrophic and heterotrophic organisms in soil. *Journal of Plant Nutrition and Soil Science*, **168**(4): 503–520.
- Li Linghao, Wang Qibing, Bai Yongfei. 2000. Soil respiration of a *leymus chinensis* grassland stand in the Xilin River Basin as affected by over-grazing and climate. *Acta Phytoecologica Sinica*, **24**(6): 680–686. (In Chinese)
- Liu Shaohui, Fnag Jingyun. 1997. Effect factors of soil respiration and the temperature effect on soil respiration in the global scale. *Chinese Journal of Ecology*, **17**(5): 469–476. (In Chinese)
- Ma Hongliang, Zhu Jianguo, Xie Zubin. 2003. Effect of elevated atmospheric CO₂ on below-ground. *Soils*, **35**(6): 465–472. (In Chinese)
- Ngao J, Longdoz B, Granier A, Epron D. 2007. Estimation of autotrophic and heterotrophic components of soil respiration by trenching is sensitive to corrections for root decomposition and changes in soil water content. *Plant and soil*, **301**(1–2): 99–110.
- Raich JW, Mora G. 2005. Estimating root plus rhizosphere contributions to soil respiration on annual croplands. *Soil Science Society of American Journal*, **69**(3): 634–639.
- Raich JW, Tufekcioglu A. 2000. Vegetation and soil respiration: Correlations and controls. *Biogeochemistry*, **48**(1): 71–90.
- Saurette DD, Chang SX, Thomas BR. 2008. Autotrophic and heterotrophic respiration rates across a chronosequence of hybrid poplar plantations on northern Alberta. *Canadian Journal of Soil Science*, **88**(3): 261–272.
- Shao Yuehong, Pan Jianjun, Xun Bo. 2005. Study on Characteristics of Soil Organic Carbon Decompositions and Carbon Pool under Different Vegetations. *Journal of Soil Water Conservation*, **19**(3): 24–28. (In Chinese)
- Tyree MC, Seiler JR, Aust WM, Sampson DA, Fox TR. 2006. Long-term effects of site preparation and fertilization on total soil CO₂ efflux and heterotrophic respiration in a 33-year-old *Pinus taeda* L. plantation on the wet flats of the Virginia Lower Coastal Plain. *Forest Ecology and Management*, **234**(1–3): 363–369.
- Uchida M, Nakatsubo T, Horikoshi T, Nakane K. 1998. Contribution of micro-organisms to the carbon dynamics in black spruce (*Picea mariana*) forest soil in Canada. *Ecological Research*, **13**(1): 17–26.
- Wang CK, Yang JY. 2007. Rhizospheric and heterotrophic components of soil respiration in six Chinese temperate forests. *Global Change Biology*, **13**(1): 123–131.
- Yang Yusheng, Chen Guangshui, Wang Xiaoguo, Xie Jinsheng, Dong Bin, Li Zhen, Gao Ren. 2005. Effect of clear-cutting on soil respiration of Chinese fir plantation. *Acta Pedologica Sinica*, **42**(4): 584–590. (In Chinese)
- Yang Yusheng, Chen Guangshui, Xie Jinsheng, Wang Xiaoguo, Niu Zhipeng, Han Yonggang, Zhang Youli. 2006. Soil heterotrophic respiration in native *castanopsis kawakamii* forest and monoculture *castanopsis kawakamii* plantations in subtropical China. *Acta Pedologica Sinica*, **43**(1): 53–61. (In Chinese)
- Zhang Dongqiu, Shi Peili, Zhang Xianzhou. 2005. Some advance in the main factors controlling soil respiration. *Advance in Earth Sciences*, **20**(7): 778–785. (In Chinese)
- Zhang Xianquan, Wang Wenjie, Zu Yuangang, Zhang Wanli. 2005. The difference between different components of soil respiration in several types of forests in northeastern China. *Journal of Northeast Forestry University*, **33**(2): 46–47, 73. (In Chinese)
- Zhong Chonglu, Zhang Yong. 2003. Introduction, cultivation and operation of *Casuarina* plants in China. *China Forestry Science and Technology*, **17**(2): 3–5. (In Chinese)
- Zhou Cunyu, Zhou Guoyi, Zhang Deqiang. 2004. Soil surface CO₂ efflux and its influencing factors in Dinghushan forest ecosystem. *Science in China (Series D, Earth Sciences)*, **34**(S2): 175–182. (In Chinese)
- Zhou Yurong, Yu Zhenliang, Zhao Shidong. 2000. Carbon storage and budget of major Chinese forest types. *Acta Phytoecologica Sinica*, **24**(5): 518–522. (In Chinese)

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09-04-01

墨西哥哈里斯科 (Jalisco) 林分结构小面积估测=Small-area estimation of forest stand structure in Jalisco, Mexico[刊, 英]/Robin M. Reich¹; Celedonio Aguirre-Bravo² (1. Department of Forest, Rangeland and Watershed Stewardship, Colorado State University, Fort Collins, Colorado 80523 USA; 2. Rocky Mountain Research Station, USDA Forest Service, Fort Collins, Colorado 80523 USA) //Journal of Forestry Research. -2009(4). -285~292

对小的生态经济区开展自然资源统计是很困难的, 政府决策人员只能依靠州水平的数据库来评价一定区域或局部的自然资源(森林、牧场、草地、农田等)状况。小面积评估技术可以用于评定这些资源。然而, 哪一种小面积估测法可以给出最可靠、最准确的结果还不得而知。本研究检测了小面积评估分析常用的两种方法(即综合估计法和回归估计法)的可靠性、准确性。运用这两种方法分析墨西哥哈里斯科(Jalisco)州全州的自然资源数据, 从而检测每种方法对所选择的森林林分结构特征预测结果的好坏。研究表明, 回归方法在多个地理尺度上, 对森林林分结构特征预测的可靠性和准确性均最好。因此, 推荐州或地方资源管理者, 在没有其他适当的辅助信息资料的情况下, 可运用回归分析法来评估小区域内自然资源状况。图4表5参14。

关键词: 森林结构; 回归估计法; 综合估计法; 空间格局; 分层随机抽样; 卫星图; 清单和监测

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09-04-02

墨西哥瓦州州东北部橡树商品材材积核算系统= A system for calculating the merchantable volume of oak trees in the northwest of the state of Chihuahua (Mexico)[刊, 英]/Marín POMPA-GARCÍA¹, José Javier CORRAL-RIVAS^{1*}, José Ciro HERNÁNDEZ-DÍAZ², Juan Gabriel ALVAREZ-GONZÁLEZ³ (1. Universidad Juárez del Estado de Durango, Río Papaloapan y Blvd. Durango S/N, Col. Valle del Sur, 34120 Durango, México.; 2. Instituto de Silvicultura e Industria de la Madera. Universidad Juárez del Estado de Durango. Km 5.5, Carretera a Mazatlán, CP 34120, Durango, Dgo., México; 3. Departamento de Ingeniería Agroforestal, Universidad de Santiago de Compostela. Escuela Politécnica Superior, Campus universitario, 27002 Lugo, Spain) //Journal of Forestry Research. -2009(4). -293~300

比较了Kozak (1988), Bi (2000) 和 Fang 等 (2000)的削度模型, 形成了橡树商品材材积核算系统。在所研究的不同树龄、不同密度的松树-橡树混交林内, 采集了298株橡树削度数据。结果表明, Fang 等 (2000)建立的可兼容分段模型可以很好地描述实验数据, 可用于估测树木在指定树高的直径、树高-指定直径比、商品材材积量和总体积。本研究建立的公

式可作为该研究区内森林调查的基本工具, 简便、易于操作。统计分析结果表明, 尽管本研究所建立的公式具有局地适应性, 但仍可应用于其他区域。图4表4参53。

关键词: 商品材材积; 栎属; 削度模型; 变化拐点

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不同林龄的木麻黄林下土壤异养呼吸特征= Soil heterotrophic respiration in *Casuarina equisetifolia* plantation at different stand ages[刊, 英]/肖胜生^{1,2}, 叶功富^{3*}, 张立华⁴, 金钊⁵, 刘丽香^{1,2} (1. 中国科学院地理科学与资源研究所, 北京 100101; 2. 中国科学院研究生院, 北京 100039; 3. 福建省林业科学研究院, 福州 350012; 4. 厦门大学生命科学学院, 厦门 361005; 5. 中国科学院地球环境研究所, 西安 710075) //Journal of Forestry Research. -2009(4). -301~306

以中国亚热带木麻黄沿海防护林为研究对象, 在2006年5月-2007年4月, 利用LI-8100土壤呼吸自动观测系统对不同林龄(幼林、中林、成林)木麻黄人工林生态系统的土壤异养呼吸特征进行了监测。结果表明, 不同林龄木麻黄林地土壤异养呼吸季节动态均呈单峰曲线, 最大值出现在6~7月份, 最小值则出现在12~1月份。土壤温度和水分对土壤异养呼吸的季节变化存在显著影响, 并有明显的交互作用, 进行单因素方差分析发现, 土壤异养呼吸季节变化与5cm深的土温存在着较好的指数相关关系 ($p < 0.05$), 与土壤表层含水量存在较好的线性相关关系 ($p < 0.05$)。双因素模型模拟结果显示, 5cm土温和土壤表层含水量能够共同解释土壤异养呼吸变异的68.9%~91.9% ($p < 0.05$)。不同林龄下土壤异养呼吸速率差异显著 ($p < 0.05$), 其平均土壤异养呼吸速率呈现为成林>幼林。随着林龄的增大, 土壤异养呼吸对土壤温度的敏感性增强。土壤异养呼吸对土壤总呼吸的贡献在幼林、中林、成林中分别达到71.89%、71.02%和73.53%。幼林、中林、成林样地土壤异养呼吸CO₂年释放量分别为29.072、38.964和30.530 t·hm⁻²·a⁻¹。图2表3参39。

关键词: 土壤异养呼吸; 沿海防护林; 林龄; 木麻黄

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09-04-04

不同土壤水分胁迫格局对印度沙漠中的黄檀种苗营养吸收和生物量的影响= Effect of varying soil water stress regimes on nutrient uptake and biomass production in *Dalbergia sissoo* seedlings in Indian desert [刊, 英]/G. Singh, Bilas Singh (Division of Forest Ecology, Arid Forest Research Institute, New Pali Road, Jodhpur- 342005, India) // Journal of Forestry Research. -2009(4). -307~313