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Leaching of nitrate, calcium and magnesium under maize cultivation on an oxisol in Togo

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Abstract

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In tropical regions, crop yields generally decrease with time, partly due to a decline in the levels of exchangeable bases linked to acid:fication of the upper layers of the soil. Nitrogen, calcium and magnesium balances were studied on an oxisol in southern Togo under continuous maize cropping with mineral fertilization and crop residue return, by measuring inputs and outputs.

The balance was near equilibrium for nitrogen (urea dressing, 120 kg N ha⁻¹ y⁻¹) and negative for calcium and magnesium. Leaching represented between 29% and 35% of the total output for nitrogen and accounted for practically all the calcium and magnesium outputs.

In the upper horizons, nitrate, calcium and magnesium contents in the soil solution were correlated. This was interpreted as an exchange between the calcium and magnesium ions adsorbed on the soil and the protons released by the transformation of ammonium into nitrate. The nutrient content of the soil solution increased in the case of no potassium fertilization, probably because of a lower nitrogen immobilization by microorganisms, the crop residue yield being reduced by the induced deficiency. An insufficient potassium fertilization mereases leaching losses of calcium, magnesium and nitrogen.

In the deeper horizons, the nitrate content of the soil solution was not correlated with the calcium and magnesium contents. Nitrogen was probably taken up by roots and some calcium and magnesium ions exchanged during their transport through the soil.

Résumé

En régions tropicales, les rendements déclinent généralement au cours du temps. en partie à cause d'une diminution des teneurs en bases échangeables du sol couplée à une acidification des horizons superficiels. Les bilans de l'azote, cu calcium et du magnésium ont été étudiés dans un oxisol du Togo méridional, pour une culture continue de mais avec fertilisation minérale et restitution des résidus de récolte, en mesurant les entrées et les sorties du système.

Le bilan est proche de l'équilibre pour l'azote (apport d'urée de 120 kg N ha⁻¹ an⁻¹). négatif pour le calcium et pour le magnésium. La lixiviation représente 29% à 85% des sorties du système pour l'azote, la quasi totalité pour le calcium et le magnésium.

Dans les horizons superficiels, les teneurs en nitrate, en calcium et en magnésium de la solution du sol sont liées. Cette liaison semble résulter de l'échange d'ions calcium et magnésium fixés sur le complexe d'échange contre des protons libéres lors de la transformation de l'azote de la forme ammoniacale à la forme nitrique. Les teneurs sont accrues par l'absence de fertilisation potassique,

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probablement en raison d'une moindre immobilisation de l'azote dans les micro-organismes, la masse de résidus de récolte produite étant plus faible pour une culture carencée. Une fertilisation potassique insuffisante augmente les entraînements de calcium, de magnésium et d'azote par lixiviation.

Dans les horizons profonds, les teneurs en nitrate ne sont pas liées à celles du calcium et du magnésium. L'azote est probablement prélevé par les racines, des ions calcium et magnésium échangés au cours du transfert à travers le sol.

Introduction

In West Africa, the challenge of a population growth rate in excess of 3% per annum with an increasing food deficit cannot be met without the efficient use of chemical fertilizers [16, 28, 29]. The quantities of nutrients to be applied are usually calculated from an agronomic and economic analysis of crop response to NPKS fertilizers [1, 5]. However, the yield of crops usually declines with time for a given level of fertilization. This evolution results partly from acidification of the upper layers of the soil associated with a decrease in exchangeable cations (mainly calcium and magnesium) and an increase in exchangeable aluminium for low pH soils [22, 16].

Lysimeter and ceramic cup studies have shown that in West African tropical environments leaching of nitrate. calcium and magnesium ions always represents a large part of agrosystem outputs, as soon as drainage occurs [11, 10, 14, 4, 27]. Although the dynamics of soil nutrients is only partially understood, some research on the relationships between ions in the soil solution, both in the field [15, 24] and in soil columns [17], have shown that nitrate ions move in the soil solution in association with other cations. Under temperate climates, losses of calcium and magnesium from the agrosystem alter the chemical properties of the soil only in the long term and they are compensated by liming. Under tropical climates, where cation exchange capacity (CEC) of the soil is generally low, these losses produce a rapid change in the level of exchangeable cations seldom corrected by the application of fertilizers or amendments.

In a previous paper [24] changes with time of the nutrient contents of the soil solution were studied in a continuous maize cropping system in Togo, and losses by leaching were estimated. This paper reports on the nutrient balances of nitrogen, calcium and magnesium and on the relationships between nutrients in the soil solution, in connection with K fertilisation.

Materials and methods

Experimental site

The experiments were carried out at the IRAT (Institut de Recherches Agronomiques Tropicales) experimental station of Davié (6° 22'N. 1° 13'E). 35 km north of Lomé. The annual weather pattern in this part of Togo (Altitude. 50 m) is tropical with two wet seasons (mean annual rainfall, 1040 mm; mean annual temperature, 27°C). The soil is a Ferralsol (Soil Taxonomy, Typic Eutrustox) representative of those of southern Togo locally called 'Terres de Barre'. The main chemical and physical properties of the soil are in Table 1.

Treatments

Two plots of a long-term NPK factorial experiment were used to estimate nutrient balances in 1985 and 1986. Both plots had been cropped twice a year with maize since 1976 (variety NH1. 110 days duration, 42,000 plants per hectare). One of the plot (K2), had received a NPK fertilization since 1976, the other (K0) had received N and P but no K. In 1985, the fertilization level was 20 kg P ha⁻¹ and -5 kg K ha^{-1} for the first cropping seasons, 13 kg P ha^{-1} and 50 kg K ha⁻¹ for the second, the nitrogen fertilization being 60 kg N ha⁻¹ in both seasons. In 1986, the same level (60 kg N ha⁻¹, 20 kg P ha^{-1} and 75 kg K ha^{-1}) was applied in each season. The fertilizers, urea (46% N), triple superphosphate (21% P, 14% Ca) and KC1 (52%. water soluble K. 47% chloride). were incorporated with a hoe to a depth of 5 cm of soil.

	oil layer at the Davié experimental	

	· · · ·	Fallow (after 13 years)	Treatment K ⁴ Treatment K2 (after 11 years of cultivation)
Clay (<2 μm) Silt (2–50 μm) Sand (50–2000 μm)		11 9 80	13 1^{-}
pH (water) Organic carbon (약)		5.4 ().99	5.2 0.50 0.2
C 'N Exchangeable cations (meq 1 Ca	100 g)	2,1	15 1.0 1.4
Mg K Na	· · · · · · · · · · · · · · · · · · ·	1.5 0.20 0.04	0.5 0.08 0.04 0.04
Al Cation exchange capacity		0.00 3.9	0.03 2.0 2.7
P Olsen-Dabin* (ppm)		40	46 25

*Extraction by NaHCO₃ 0.5 M and FNH, 0.5 M at pH 8.5

Phosphorus and potassium were applied at sowing, nitrogen was applied half at thinning and half at tasselling. Crop residues remained on the plots and were incorporated once a year before the first cropping season.

Experimental design and measurements

The mineral content of the rains was determined on 8 samples collected from 198- to 1986 (Table 2). The annual amount of nutrients brought by the rains was estimated by multiplying the average content by the year's total rainfall. Both plots $(8.0 \pm 4.5 \text{ m})$ were equipped with 4 neutron access tubes and 14 tensiometers to monitor soil water content and movement. Measurements were made every day during the period of cultivation and twice a week for the rest of the year. Drainage below the root zone (1.5 m) was estimated using Darcy's law and suction-conductivity relationships [19]. Statistical calculations were made to estimate the standard error on drainage [18]. 4.5 cm in diameter) was placed at 30 and 180 cm depths to collect soil solutions. Samples were taken twice a month through the year by applying a 600 hPa suction for 24 hours. Nitrate content was measured by the phenoldisulfonic method, calcium and magnesium by atomic absorption spectrophotometry and potassium by flame spectrophotometry. As the roots never reached below 150 cm and as the clay content of the soil did not change with depth at this level. the nutrient content of the soil solution was assumed to be the same at 180 cm and at 150 cm. Leaching of each element between two sampling dates was calculated by multiplying its content in the soil solution by the amount of drainage during the period.

Dry matter yields and crop removals were determined by collecting 10 plants from each of the plots for 3 years prior to the experiment [23]. The dry matter weight of each organ (root, stem, leaves, leaf sheaths and grain) was determined after drying at 80°C, and nutrient content measured on, 2 sub-samples. A relationship was established between the grain yield, and the

A battery of 4 ceramic cups (Soil Moisture

Table 2. Mean rainfall content and corresponding amounts supplied in precipitation (\$ samples collected from 1984 to 1986, average rainfall 1934-1986; 1043 mm)

	· .	· · · · · · ·			6 t	i		
Nutrient	NO.	Çu 👘	Mg		K	PO	Cl	,
Content (meg 1 ⁽¹⁾)	< 0.03	0.5	0.92	- 13 .	0.01	0.005	0 Q8	
Input kg ha y)	4.4 .	10	1.5	`//T_D^,	- 1	1.i '	29.5	1

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from the grain vield of the plot.

Results

Nutrient balances

The nutrient balance sheet of each element was calculated taking into account the amounts supplied in precipitation and fertilizer dressings as inputs. leaching below the root zone (150 cm) and crop removals as outputs. Losses by runoff were disregarded, as the site was flat (slope less than 1%). Ammonium content in the soil solution, between 0.06 and 0.11 meq 1^{-1} , was neglected. Losses of nitrogen in the gaseous phases were not measured. Ammonia volatilization was probably low, being negligible in tropical regions as soon as the fertilizer is incorporated [8]. Losses by denitrification were generally low, although a value of 30% for applied nitrogen has been reported [29].

removals for each plot and each cropping season

Rainfall in 1985 and 1986 (Table 3 and Fig. 1), was close to the annual average. The grain yields obtained were also near the average: 4.0 tha^{-1} y^{-1} in treatment K2 and 1.4 tha^{-1} y^{-1} in K0 between 1976 and 1989 [25]. The years 1985 and 1986 were therefore considered as representative of general conditions in the region.

Yield was lower in treatment K0 than in K2 because of the critical deficiency of potassium that occurs for all crops on the Terres de Barre after some years of cultivation without K fertilization [6]. This deficiency also produced stunted stems with short internodes, and many dead plants (41%). As a result, the mean weight of the crop residues returned to the soil was 2000 kg $ha^{-1} v^{-1}$ in treatment K0 and 4600 kg $ha^{-1} v^{-1}$ in K2. Despite this large difference in growth. annual evapotranspiration from both plots differed by less than $10^{\circ}c$, as evaporation from the soil remained high throughout the year (maximum LAI, 2.5). Consequently, the difference in drainage between both plots was less than 20%. This result is consistent with experiments carried out in the Niger [29] where NPK fertilization increased grain yields of pearl millet, but had little effect on water use. In 1986, drainage was higher than 1985 due to both the distribution and quantity of rainfall (Fig. 1). The precision of the drainage value was poor (CV 31%), as direct measurement was impossible in the field [18]. For nitrogen, the balance was near equilibrium in both treatment plots (Table 4), except in treatment K0 in 1986 where the confidence interval was high because of heavy drainage. If denitrification occurred, the balance was slightly negative. For calcium and magnesium, the balance was negative in all cases despite intensive NPK fertilization containing some calcium. Thisagrosystem is therefore not sustainable. If crop residues had been removed from the fields, as is usual in this region. losses would have been higher, as maize residues contained 54 kg N ha⁻¹. 23 kg Ca ha⁻¹ and 20 kg Mg ha⁻¹ in 1986 in treatment K2.

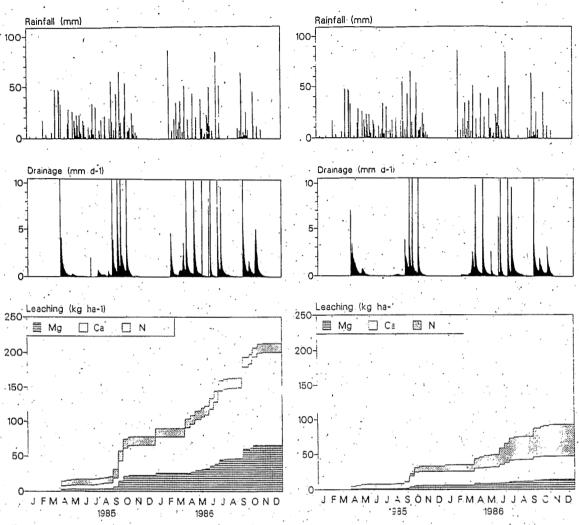
Nitrate leaching

In the agrosystem studied, nitrate in the leachates represented 29% (1985) and 38% (1986) of the nitrogen output in treatment K2, 73% (1985) and 85% (1986) in K0. Leaching was 30% (1985) and 42% (1986) of the applied amount of urea in

Year Rainfa Treatment	x	1985 1045 mm K0	. 1	K2	• `	1-36 1-51 mm K0		K2
Maize yield (t ha ⁻¹ y) Evapotranspiration (mm) Drainage (mm	• • •	2.0 781* 264 = 85**	· · · ·	4.5 829* 216 = 67**		1.5 c14 ± 161** -30 ± 1-1**		4.4 672 ± 145** 376 ± 115**
Water content variation in the 0-150 cm soil (2) er (mm)	а. Т				· . ·	-1=-		
"Estimated ""Standard errors	• -			· •.	· .	• • •	• • • • • • •	· •

Table 3. Maize yield and water balance of treatment K0 and K2 in 1985 and 1986 (two cross per year)

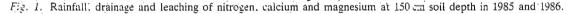
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K0 treatment

K2 treatment

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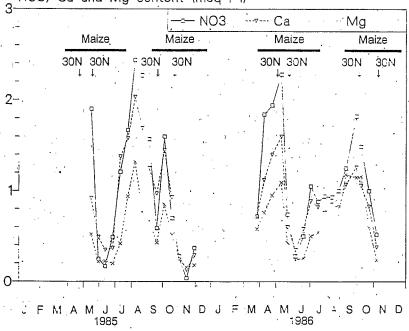
treatment K2, 76% (1985) and 128% (1986) in K0.

Nitrate leaching depends on the distribution of rainfall. In 1985, leaching in the K0 plot was $21 \text{ kg N} \text{ ha}^{-1}$ during the first growing seasons (April to July) with 377 mm of rainfall. It was 95 kg N ha⁻¹ in 1986 during the same period with only 308 mm (Fig. 1). There were many light rains in 1985, and few by heavier rains in 1986.

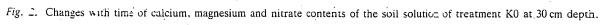
In both years, nitrate leaching was higher at the start of the rainy season than during the rest of the rainy period. This phenomenon was partly due to the decrease in drainage during the rainy seasons because of plant transpiration. However plant transpiration did not increase the evapotranspiration much during the cropping season. On the contrary, the nitrate content in the soil solution varied with time between 0.04 and $2.4 \text{ meq } 1^{-1}$ in treatment K0 at a depth of 30 cm (Fig. 2).

The highest nitrate contents in the soil solution was recorded at the very beginning of the rainy seasons, and this increase preceded the application of nitrogen fertilizers. The lowest nitrate content in the soil solution was recorded in June and November for both years. The quantities of nitrate leached were about three times higher in

Nitrogen balance (kg N hæ ⁻¹)	· · · ·	,	Treatment K0 1985	1986	Treatment K2 1985	1986
Inputs ·	Raintai		· 4 .	· + .	4.	4
	Fertilizer (urea)		120	120	120	120
Outputs	Maize grain	÷ .	33 .	28	90	· 82 ·
,	Leaching at 150 cm	•	91 ± 28	153 ± 48	36 = 12	50 = 15
Balance			0 ± 28	-57 ± 48	-2 = 12	-8 = 15
Calcium balance			Treatment K0		Treatment K2	· .
(kg Ca ha 1)		• •	1985	1986	1985	1986
Inputs	Rainfall	•	9	10	· 9 .	10
· · ·	Fertilizer (superphosphate)		<u>. רָכ</u>	26	22	26 .
Outputs	Maize grain		0.4 .	0.2	· 0.8	í 0.0
·	Leaching at 150 cm		_5 ± 22	148 = 47	47 = 14	91 = 25
Balance	•		-44 ± 22	-112 ± 47	-17 = 14	-56 = 28 .
Magnesium baland	°e · · •		. Treatment K0 .		Treatment K2	
(kg Mg ha ⁻¹)		•-•-	1985	1986	1985	1986
Inputs	Rainfall	• ,	2 **	.2	2	2
	Fertilize:		û ·	- 0	• 0 '	. 0
Outputs	Maize grain		1 <u>2</u>	. : 2	5	5
	Leaching at 150 cm		<u>29 ± 9</u>	60 = 19	. 13 = 4	28 = -
Balance	-		-29 ± 9	$^{1}-60 = 19$	-16 = -16	-31 = -1



NO3, Ca and Mg content (meq I-1)



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treatment K0 than in K2 during a given year (Table 4). By increasing dry matter production, potassium fertilization decreased drainage. But, as noted previously, the differences in drainage remained less than 20% between the two plots. The higher leaching in treatment K0 was therefore mainly due to a higher nitrate content of the soil solution (average nitrate content at 180 cm: $2.4 \text{ meq } l^{-1}$ for K0 treatment and $1.3 \text{ meq } l^{-1}$ for K2).

Calcium and magnesium leaching

After 11 years of cultivation, exchangeable calcium and magnesium in the upper layers of the soil had sharply decreased (Table 1) and almost all outputs of calcium and magnesium from the agrosystem came from leaching (Table 4). Calcium and magnesium concentrations in the soil solution were correlated:

Plot K0 (30 cm)

$$[Mg] = 0.58 (\pm 0.08) [Ca] - 0.08 (\pm 0.20)$$

$$(r^2 = 0.65, n = 3)$$

Plot K2 (30 cm)

$$[Mg] = 0.70 (\pm 0.04) [Ca] - 0.01 (\pm 0.08)$$

1)

$$n^2 = 0.90. n = 34)$$
 (2)

(Ca and Mg contents expressed in meq 1^{-1})

These relationships are close to those already observed in the same climatic area [15]. At 30 cm depth, the Mg Ca ratio in the soil solution of treatment K0 was about 0.59 and 0.71 in K2, whereas it was respectively 0.50 and 0.44 in exchangeable forms in the 0-20 cm soil layer (Table 1). The Mg/Ca ratio of the exchangeable ions of the soil increased with depth between 20 and about 70 cm (Figs 3 and 4). Exchangeable magnesium was therefore preferentially leached from the upper layers of the soil, compared with exchangeable calcium. It can be concluded that exchangeable magnesium was adsorbed by the

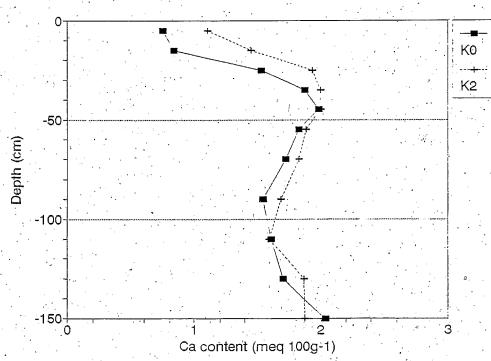


Fig. 3. Exchangeable calcium profile distribution for treatment K0 and K2 (mean values of 6 samples).

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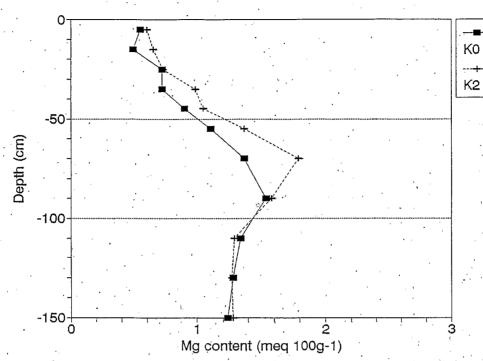


Fig. 4. Exchangeable magnesium profile distribution for treatment K0 and K2 (mean values of 6 samples).

solid phase of the soil less than exchangeable calcium. Changes with time of calcium and magnesium contents in the soil solution followed the same pattern as the nitrate content (Fig. 2). For the K0 treatment, calcium and magnesium contents were correlated with the nitrate content but there was no significant correlation with potassium (Table 5):

 $[Ca + Mg] = 1.03 (\pm 0.08) [NO_3] + 0.42 (\pm 0.30)$

 $(r^2 = 0.85, n = 31)$ (3)

(NO₃, Ca and Mg contents expressed in meq l^{-1})

For the K2 treatment, calcium and magnesium concentrations in the soil solution were correlated with both nitrate and potassium contents

Plot K0 (30 cm)

Table 5: Correlation matrix between the concentration of ions in the soil solution as affected by fertilizer treatment and depth of sampling

5 ampi						•						
	NÖ,	Ca	Mg	K	NO,	Ca	Mg	, K	NO ₃	Ca	Mg	K
NO Ca Mg K	1.00 0.92*** 0.82*** 0.33*	1.00 0.81*** 0.24*	1.00 0.27*	1.00	'1.00 0.81*** 0.82*** 0.76***	1.00 0.95** 0.67***	1.00 0.63***	1.00	1.00 -0.08* -0.03* 0.02*	1.00 0.70*** 0.40**	1.00 0.45**	1.00
Treati Treati ***P =	nent K0 at nent K2 at nent K2 at ≤ 0.01 $1 < P \leq 0.05$	30 cm (n = 180 cm (n =	3∸)	•					• •	•	· ' ·	۰.

*0.05 < P

(Table 5). The potassium content of the soil solution was always between 0.04 and 0.08 meq l^{-1} at 30 cm, except at the time of K fertilizer application when it reached up to 0.25 meq l^{-1} . At a 180 cm depth (below the root zone), there was no relationship between nitrate, calcium and magnesium contents (Table 5).

Discussion

As the sharp increase in the nitrate content of the soil solution at the beginning of each rainy season preceded the application of urea, the nitrate ions formed must stem from the mineralisation of the organic matter and the nitrification of the resulting ammonium:

organic N \rightarrow R-NH₂ + CO₂

(proteolysis) (4)

(5)

$$R-NH_2 + H_2O \rightarrow R-OH + NH_3$$
(ammonification)

$$\mathrm{NH}_3 + \mathrm{H}^- \longrightarrow \mathrm{NH}_4^- \tag{6}$$

 $NH_{1}^{-} + 3/2 O_{2} \rightarrow NO_{2}^{-} + H_{2}O + 2 H^{+}$ (nitrification by Nitrosomonas) (7)

$$NO_2^- + 1/2 O_2 \rightarrow NO_3^-$$

(nitrification by Nitrobacter) (8)

This widely described process which is common in tropical climates [3, 9, 8] results from intense microbiological activity following the first rains along with high soil temperature. In the soil, the applied urea is converted to ammonia by urease:

$$CO(NH_2)_2 + H_2O \rightarrow 2NH_3 + CO_2$$
 (9)

The ammonia is then transformed into nitrate ions according to the former relations. Both organic matter and urea therefore contribute to increasing the nitrate content of the soil solution.

The transformation of ammonia into ammonium requires one proton. The process of nitrification (ammonium to nitrate ions) produces two protons. For each nitrate ion formed from

either organic matter or urea, one proton is released in the soil solution. As the pH in the soil solution was between 7.2 and 5.9, and as there were few primary minerals to be attacked by the protons in this oxisol, almost all protons must have been exchanged with cations adsorbed by the solid phases of the soil. This interpretation is in agreement with the ratio between the nitrate content and the calcium plus magnesium content in the K0 soil solution that was almost equal to 1 (relation 3). It is consistent with other reports on the interrelations of ion content in the soil solution [12] although, depending on the acidity of the soil, calcium and magnesium in the soil solution can be partly replaced by aluminium [17]. It is supported by the fact that calcium content in the soil solution is lower when ammonia nitrification is blocked either by frost or by a nitrification inhibitor [7]. The transformation of organic matter nitrogen and applied urea into nitrate ions is therefore a direct cause of soil acidification and of the release of calcium and magnesium from the exchange complex of the soil.

The observed correlation between calcium and magnesium concentrations and potassium content in the K2 soil solution can be interpreted as competition between potassium, calcium and magnesium ions for exchange sites at the time of potassium fertilization. The potassium concentration in the soil solution decreased rapidly after addition of fertilizer and remained low, with most of the exchange of potassium ions from the soil solution to the solid phase of the soil occurring in less than one minute [20]. Consequently, potassium ions no longer interfered with the calcium and magnesium ions a few days after the dressing. Unlike proton exchanges, this process would not involve acidification of the soil, and it may be of short duration given that potassium ions are taken up by crops during the growing season.

The higher nitrate content of the soil solution at 30 cm in treatment K0 before any fertilization was probably due to the amount of crop residues returned to the soil. Mineralisation of crop residues needs nitrogen taken from the soil solution. More nitrogen was needed in K2 treatment than in K0, as the dry matter of crop residues is about 2600 kg ha⁻¹ y⁻¹ higher in K2

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than in K0. As a great part of the leaching occurred at the beginning of the cropping seasons, it seems that the decrease in nitrate leaching associated with potassium fertilization (Table 4) resulted mainly from the larger amounts of crop residues returned to the soil. Cropping decreases the nutrient content of the soil solution, and thus leaching losses, both by root uptake and by fixing nitrogen in microorganisms to mineralize the crop residues. Improving plant growth, especially by a balanced fertilization, would alleviate nitrate, calcium and magnesium leaching and acidification of the soil for a given amount of applied urea.

This interpretation agrees with the effect of fertilization on the decrease in soil acidification. The pH is 0.4 units higher after 11 years of cultivation in K2 than in K0 treatment (Table 1). With the same potassium dressings (85 kg K ha⁻¹ y⁻¹), the difference in pH between 0 kg P ha⁻¹ y⁻¹ and 23 kg P ha⁻¹ y⁻¹ treatment was 0.5 units for 0 kg N ha⁻¹ y⁻¹ (5.1 versus 5.6) and 0.9 units for 120 kg N ha⁻¹ y⁻¹ (4.3 versus 5.2).

The very low nitrate concentration levels observed in June and November cannot be a consequence of leaching, as drainage was only 31 mm between April and July in 1985. They must therefore result from the absorption of nitrate ions by maize roots. During the growth period, the crop absorbed large quantities of nitrogen (3.3 kg N ha⁻¹ day⁻¹ at tasselling in treatment K2). The root system in moderately fertilized agrosystems efficiently draws almost all the available nitrate ions from the soil solution in. the root zone, both in tropical [2] and temperate zones [21, 13].

Whenever drainage occurs, the soil solution formed in the upper layers of the soil passes through the deeper layers. Most of the nitrate ions are taken by roots and some calcium and magnesium ions are exchanged with other cations on the exchange complex. These processes can explain why the nitrate, calcium and magnesium contents are not correlated when the soil solution moves to the deeper layers (Table 5). The variations in the soil solution with depth depend on soil properties and on plant uptake. In the situation studied, calcium and magnesium leaching was higher and deeper in treatment K0 than in K2 (Fig. 3 and 4), as the nitrate content was higher in the soil solution and crop uptake was lower. Magnesium leaching went deeper than calcium leaching. This is in agreement with results obtained from magnesium-calcium exchange on kaolinitic soil clay [26].

Conclusion

As the nitrate content was almost the same as the calcium plus magnesium content in the soil solution of the upper horizons of the K0 treatment, the transformation of ammonia into nitrate seemed to be the main factor influencing the movement of calcium and magnesium ions from the solid phase of the soil to its solution. This release was increased by urea application or the addition of potassium ions: these apparently took some exchange sites previously occupied by calcium and magnesium ions. In tropical soils where CECs are low, the fertility of the agrosystems is therefore not sustainable in the long run without calcium and magnesium dressings.

The amounts of calcium and magnesium to be added must exceed losses by leaching. Evaluation of leaching rates below the root zone involves the measurements of both the nutrient content in the soil solution and the water balance in the deeper part of the profile. However the composition of the soil solution varies with depth from the surface, probably due to root extraction and ion exchange with the solid phase of the soil. The mere determination of the composition of the soil solution in the deeper layers is therefore insufficient to clarify the processes involved, and to allow evaluation of leaching in the long term. A better understanding of nitrate, calcium and magnesium leaching requires studying: (i) nutrient released by fertilizers, organic matter, crop residues, microorganisms and the soil matrix in the soil solution of the upper horizons; and (ii) change in composition of the soil solution as it passes through deeper layers, in relation to plant. uptake and ion exchange between the soil solution and the solid phase of the soil.

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