

2-2021

Impact of Increased Nutrients and Lowered pH on Photosynthesis and Growth of Three Marine Phytoplankton Communities From the Coastal South West Atlantic (Patagonia, Argentina)

Takako Masuda

Ondřej Prášil

Virginia E. Villafañe

Macarena S. Valiñas

Keisuke Inomura

See next page for additional authors

Authors

Takako Masuda, Ondřej Prášil, Virginia E. Villafañe, Macarena S. Valiñas, Keisuke Inomura, and E. Walter Helbling



Impact of Increased Nutrients and Lowered pH on Photosynthesis and Growth of Three Marine Phytoplankton Communities From the Coastal South West Atlantic (Patagonia, Argentina)

OPEN ACCESS

Edited by:

Eva Ortega-Retuerta,
UMR 7621 Laboratoire
d'Océanographie Microbienne
(LOMIC), France

Reviewed by:

Haimanti Biswas,
National Institute of Oceanography
(CSIR), India
Antonietta Quigg,
Texas A&M University at Galveston,
United States

***Correspondence:**

Takako Masuda
takako@alga.cz

Specialty section:

This article was submitted to
Aquatic Microbiology,
a section of the journal
Frontiers in Marine Science

Received: 24 September 2020

Accepted: 07 January 2021

Published: 12 February 2021

Citation:

Masuda T, Prášil O, Villafañe VE,
Valiñas MS, Inomura K and
Helbling EW (2021) Impact
of Increased Nutrients and Lowered
pH on Photosynthesis and Growth
of Three Marine Phytoplankton
Communities From the Coastal South
West Atlantic (Patagonia, Argentina).
Front. Mar. Sci. 8:609962.
doi: 10.3389/fmars.2021.609962

Takako Masuda^{1*}, **Ondřej Prášil**¹, **Virginia E. Villafañe**^{2,3}, **Macarena S. Valiñas**^{2,3},
Keisuke Inomura⁴ and **E. Walter Helbling**^{2,3}

¹ Institute of Microbiology, The Czech Academy of Sciences, Třeboň, Czechia, ² Estación de Fotobiología Playa Unión, Chubut, Argentina, ³ Consejo Nacional de Investigaciones Científicas y Técnicas (CONICET), Chubut, Argentina, ⁴ Graduate School of Oceanography, University of Rhode Island, Narragansett, Rhode Island, RI, United States

Effect of global change variables on the structure and photosynthesis of phytoplankton communities was evaluated in three different sites of the Patagonian coast of Argentina: enclosed bay (Puerto Madryn, PM), estuarine (Playa Unión, PU), and open waters (Isla Escondida, IE). We exposed samples to two contrasting scenarios: Present (nutrients at *in situ* levels) vs. Future (with lowered pH and higher nutrients inputs), and determined growth and photosynthetic responses after 2 days of acclimation. Under the Future condition phytoplankton growth was higher in the estuarine site compared to those in PM and IE. This effect was the most pronounced on large diatoms. While the increase of photosynthetic activity was not always observed in the Future scenario, the lower photosynthetic electron requirement for carbon fixation ($\Phi_{e,C} = \text{ETR}/P_m^B$) in this scenario compared to the Present, suggests a more effective energy utilization. Long-term experiments were also conducted to assess the responses along a 4 days acclimation period in PU. Diatoms benefited from the Future conditions and had significantly higher growth rates than in the Present. In addition, $\Phi_{e,C}$ was lower after the acclimation period in the Future scenario, compared to the Present. Our results suggest that the availability, frequency and amount of nutrients play a key role when evaluating the effects of global change on natural phytoplankton communities. The observed changes in diatom growth under the Future scenario in PU and IE and photosynthesis may have implications in the local trophodynamics by bottom up control.

Keywords: global change, plankton community structure, carbon incorporation, oxygen production, PSII photochemistry

INTRODUCTION

Global change is a phenomenon commonly associated with an increase in the average temperature that brings about significant changes in the ecosystems, affecting the marine life and its environment in various ways (IPCC, 2019). Other processes, directly or indirectly related to global change, will alter coastal environments, either through significantly decreased pH (associated with higher CO₂ levels released into the atmosphere) and/or due to higher nutrient inputs resulting from increasing agricultural and industry activities and urban development (IPCC, 2019). The effects of such alterations as well as the capacity to adapt to such changes are species-specific among phytoplankton communities (Gao et al., 2012). Higher nutrient inputs usually favor phytoplankton growth and photosynthesis (Falkowski et al., 1998; Behrenfeld et al., 2006), especially under nutrient-limiting conditions (Howarth, 1988; Cloern, 1999). Many organisms on the other hand, benefit from increasing CO₂ levels by investing less energy into carbon concentrating mechanisms (Levitán et al., 2007; Rost et al., 2008; Hopkinson et al., 2011). In nature, different environmental variables interact usually in a synergistic or antagonistic manner, enhancing or reducing their individual effects (Crain et al., 2008). For example, Domingues et al. (2016) found that nutrient enrichment suppressed the negative effects of high CO₂ and UVR on cyanobacteria and diatoms in an estuarine community. Hama et al. (2016), on the other hand, found significant effects of nutrient addition on phytoplankton growth only during the early stages of a 1 month mesocosm experiment, while the negative effects of higher acidification were evident only during the post-bloom period.

Intensive efforts have been put in evaluating and potentially predicting the effects of global change on the growth and photosynthetic performance of phytoplanktonic organisms, due to their key role as primary producers in the aquatic food webs. Such efforts are evident from large available datasets (Dutkiewicz et al., 2015; Bach and Taucher, 2019) that also highlight the vulnerability of organisms and ecosystems to global change. Some geographical regions, however, have been poorly explored and represented, and thus there is an urgent need to fill the knowledge gaps in how global change may impact primary producers in these areas. One clear example of these gaps are the South West Atlantic Ocean coastal waters, in which relatively little is known about the effects of global change variables on these organisms, despite the fact that 6–11% of global primary productivity occurs in this area (Uitz et al., 2010; Rousseaux and Gregg, 2014). Patagonian waters constitute one of the most important fishery areas of the Atlantic Ocean Basin (Carli et al., 2012; Góngora et al., 2012). In this region, some studies have determined significant effects on different global change variables on the size structure, taxonomic composition and phytoplankton responses of natural communities [Villafañe et al. (2021) and references therein]. Only few studies (Villafañe et al., 2015, 2018; Bermejo et al., 2020) assessed the effects of both increased nutrients and lowered pH, conditions that are already present in Patagonian coastal waters (Bermejo et al., 2018; Orsell et al., 2018), on phytoplankton communities, however, since these studies were

conducted in an estuarine site with unique characteristics it made the extrapolation of these results to other ecosystems unfeasible. To get a broader range of responses, we selected three geographical sites with similar temperature and radiation conditions but different geomorphological characteristics, and thus with different water exchange frequencies, on the Patagonian coast. The aim of our study is to assess the response and efficiency of photosynthesis in these three phytoplankton communities exposed to simulated conditions of global change (lowered pH and increased nutrients), using an experimental approach (Morris, 1991; Quinn and Keough, 2002; Boyd et al., 2010). We believe that these evaluations may have important implications on our understanding of current and potentially future local and regional ecosystem services.

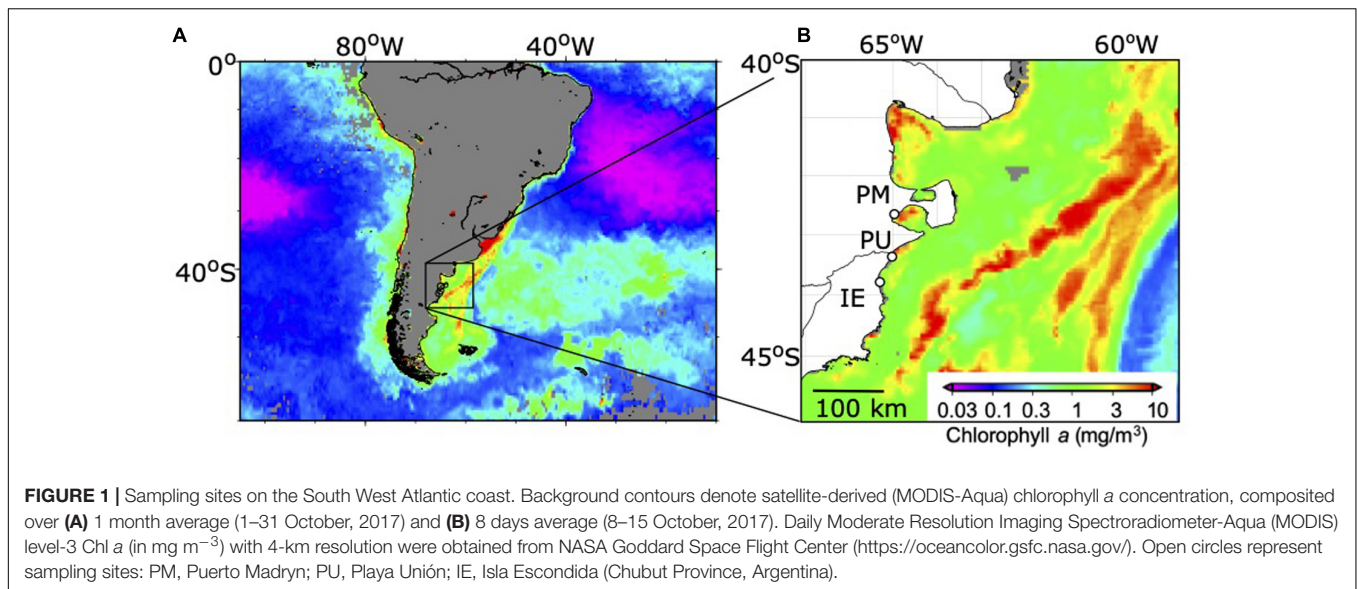
MATERIALS AND METHODS

Study Area

This study was carried out with seawater collected at three sites in the Patagonian coast, in the Chubut Province, Argentina: (a) Puerto Madryn (42° 39' S; 64° 58' 40" W, hereafter PM), located in the Golfo Nuevo; (b) Playa Unión, at the mouth of the Chubut River estuary (43° 20.5' S, 65° 0.20' W, hereafter PU), and (c) Isla Escondida (43° 47' S; 65° 18' 17" W, hereafter IE) (**Figures 1A,B**). These sites were selected because of their different physical, biological and chemical characteristics. The PM site is located within an enclosed gulf, with little exchange with open waters (Rivas and Beier, 1990) which results in high Chl *a* concentration at surface waters throughout the year (D'Agostino et al., 2018). Only a few studies have focused on the effects of global change variables in this site; namely, one investigated the impact of solar ultraviolet radiation (UVR) on photosynthesis, and another the interactive effects of UVR and nutrients on phytoplankton growth and species composition (Villafañe et al., 2004; Marcoval et al., 2008). The PU site is under the influence of riverine waters that carry a heavy load of anthropogenic nutrients (Helbling et al., 2010; Bermejo et al., 2018), which also results in high Chl *a* concentration at the plume of the river (**Figure 1B**). Several global change studies have been done on phytoplankton communities from this site, with variables such as lowered pH, nutrients, temperature, input of terrigenous materials and mixing (Helbling et al., 2015b; Bermejo et al., 2018; Villafañe et al., 2018; Bermejo et al., 2020). Contrary to the PU site, IE is located in the open waters, and is a highly productive area, being the main nursery location for the hake (*Merluccius hubbsii*) (Álvarez-Colombo et al., 2011). To the best of our knowledge, in this site no research on the effects of global change on phytoplankton communities has yet been conducted.

Experimental Set up

Surface seawater (<0.5 m) samples (ca. 20 and 100 L, for the short- and long-term experiments, respectively) from the sites were collected with an acid-cleaned (1N HCl) bucket the day before the start of the experiments, and immediately taken to the laboratory at the Estación de Fotobiología Playa Unión (EFPU, 1 h, 10 and 40 min away from PM, PU, and IE,



respectively). To remove large zooplankton, samples were pre-screened (200 μm mesh) and stored in the dark overnight in an incubator set at *in situ* temperature (14°C; **Table 1A**) until the start of experiments in the following morning. Two types of experiment were done: (a) Short-term experiments (PM, PU, and IE), where samples were allowed to acclimate for 2 days, and then measurements and analyses were performed during the day (see below) and (b) Long-term experiment at PU, where the analyses and measurements were performed along different acclimation periods (i.e., no acclimation, 2- and 4 days acclimation). A schematic representation of the experiments indicating the sampling and measurements is presented in **Figure 2**. The PU site for the long-term experiment was the most dynamic from the three selected sites, with large gradients in its biotic and abiotic variables due to the riverine input. As a consequence, the organisms here had different acclimation capacities compared to those from PM to IE, which came from more stable conditions. For this long-term experiments, we selected 2 and 4 days of acclimation that were defined mid- and long-term acclimation, respectively, as these periods were enough to acclimate to the different scenarios as seen in previous

studies carried out at PU (Villafañe et al., 2015). Prior to these experiments, we collected samples for Chl *a*, nutrients, pH, and taxonomic composition to assess the main characteristics of the study sites.

The short-term experiments consisted of two phases: Phase I (acclimation), in which the cells were placed in microcosms (500 mL of UV-transparent Teflon bottles) and were acclimated to solar radiation for 2 days in an outdoor water-bath. The temperature of the water-bath (14°C) was controlled manually by running water and the addition of ice when necessary. During the acclimation period, the Teflon bottles were mixed manually every 2–3 h to prevent cells from sinking. This process also ensured a homogeneous irradiance to the cells. The short-term experiments and acclimation periods started on 15, 16, and 17th of October 2017, for PM, IE, and PU, respectively (**Figure 2**). Two scenarios, Present and Future, were implemented in the microcosms (triplicate for each condition): (i) Present scenario: Samples with *in situ* nutrients concentration and pH; (ii) Future scenario: Samples with increased nutrients concentration and lower pH values. Here, pH was reduced by 0.4 units to reach the levels expected by the end of the century (IPCC, 2019). Increased nutrient concentration was achieved by adding macronutrients as in the *f/20* medium (Guillard and Ryther, 1962). Nutrient enrichment mimicked a higher input (10 times compared with the Present scenario) due to human activity (Depetris et al., 2005) and increase in both atmospheric dust deposition (Wagener et al., 2008; Cloern et al., 2016), and river runoff (Rabalais et al., 2009; Bermejo et al., 2018). To lower the pH, CO_3^{2-} (as Na_2CO_3), HCO_3^- (as NaHCO_3), and HCl (0.01N) were added to the seawater samples to increase pCO_2 and the dissolved inorganic carbon (DIC) content, as recommended by Gattuso et al. (2010). To determine the variations that may have occurred during the day, pH measurements in the microcosms were performed early in the morning, on a daily basis, using a pH meter (Hanna, model HI-2211, United States). Adjustments were done as required to maintain a difference of 0.4 units between the Present and the

TABLE 1A | Chemical and biological characteristics of seawater collected for each experiment.

Experiment	Short-term			Long-term
	PM	IE	PU	PU
Nitrate + nitrite (μM)	0.18	0.24	0.55	7.89
Phosphate (μM)	1.29	0.76	1.33	1.68
Silicate (μM)	2.69	2.51	26.6	22.2
Temperature (°C)	14	14	13.8	13.5
pH	8.102*	8.34	8.07	8.28
Chl <i>a</i> ($\mu\text{g L}^{-1}$)	0.92	0.39	1.17	1.15

*Data from Kahl (2018).

TABLE 1B | Incident solar radiation (PAR, irradiances, $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$) during the short- and long-term experiments.

Experiment		Short-term			Long-term	
		PM	IE	PU	No acc	Long acc
Sites/Acclimation						
Date of measurement in October 2017		17th	18th	19th	16th	20th
9–10 h	PAR ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$)	457	617	777	759	786
	PAR for FRRf ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$)	508	419	553	765	638
12–13 h	PAR ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$)	1,421	1,480	1,120	1,458	1,504
	PAR for FRRf ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$)	1,630	1,131	929	1,476	1,169
15–16 h	PAR ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$)	969	1,133	731	1,074	443
	PAR for FRRf ($\mu\text{mol photons m}^{-2} \text{ s}^{-1}$)	1,207	929	464	1,207	1,131

No acc, Long acc denote samples with no acclimation or long acclimation of 4 days, respectively.

Future scenarios as described before. Phase II measurements started after 2 days of acclimation and consisted of oxygen evolution, carbon incorporation and fluorescence measurements (see below), three times of the day: morning (9–10 am), noon (12–1 pm), and afternoon (3–4 pm).

For the long-term (5 days) phytoplankton experiment from PU, the acclimation period started on the 16th of October (Figure 2). The collected sample (c.a. 100 L) was distributed into six microcosms—three for each scenario, as described above—in 10 L (LPDE Cubitainers, Nalgene), and put inside the outdoor water-bath for temperature control. Oxygen evolution, carbon incorporation and fluorescence parameter measurements were measured three times a day on October 16 and 20th, as described above for the short-term experiments (Figure 2). For the long-term PU experiment, Chl *a* was sampled on a daily basis, while samples for nutrient and taxonomic composition were collected on the 16, 18, and 20th of October.

Analyses and Measurements

Solar Radiation Measurements

Solar radiation was continuously monitored using a European Light Dosimeter Network broadband filter radiometer (ELDONET, Real Time Computers, Germany), which measures UV-B (280–315 nm), UV-A (315–400 nm), and PAR (400–700 nm) every second, averages the data over a 1 min interval, and stores them in a computer. This radiometer was permanently installed on the roof-top of EFPU, and routinely calibrated once a year using a solar calibration procedure. For this calibration, the irradiance data during a clear sky condition was compared with the output of radiation transfer models such as STAR (Ruggaber et al., 1994) and Daylight (Björn and Murphy, 1985).

Nutrients

Samples for the determination of macronutrient concentrations ($\text{NO}_3^- + \text{NO}_2^-$, PO_4^{3-} , SiO_3^{2-} ; in triplicates for each scenario during the long-term at PU) were collected in 100 mL High Density Polyethylene (HDPE) bottles and kept frozen (-20°C) until the analyses. Nutrient concentrations were measured by colorimetric techniques (Strickland and Parsons, 1972) using a spectrophotometer (Hewlett Packard, model 8453E, United States). Specifically, $\text{NO}_3^- + \text{NO}_2^-$ concentrations were measured by the Cd-Cu column method, while PO_4^{3-} and

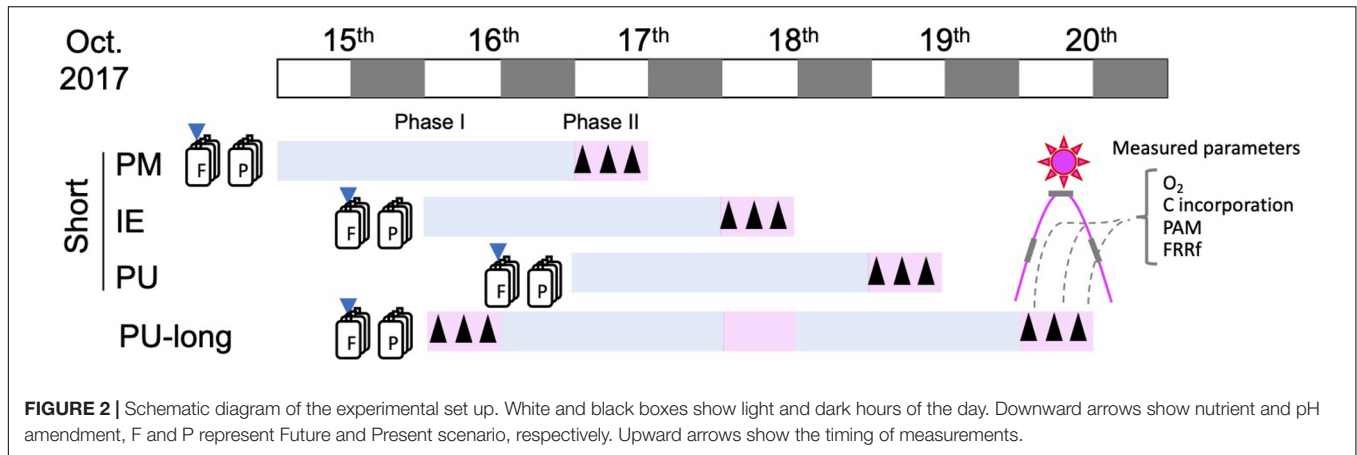
SiO_3^{2-} concentrations were measured by the molybdenum blue method. Detection limit ($3 \times$ standard deviation of the blank) for each nutrient was 0.09, 0.18, and 0.11 μM for $\text{NO}_3^- + \text{NO}_2^-$, PO_4^{3-} , SiO_3^{2-} , respectively.

Counting and Identification of Phytoplankton Cells

Samples for the identification and counting of phytoplankton cells ($>2 \mu\text{m}$, i.e., nano- and microplankton) were placed in 60 mL brown glass bottles and fixed with buffered formalin (final concentration of 0.4% of formaldehyde in the sample). Sub-samples of 10 or 25 mL were allowed to settle for 24 h in a sedimentation chamber (HydroBios GmbH, Germany). A drop of Rose Bengal was added to the chamber to better distinguish small cells from detritus and sediment. Species were identified and enumerated using an inverted microscope (Leica, model DM IL, Germany) following the technique described by Villafañe and Reid (1995). At least 200 cells were counted and identified at the species or genus level, although in some cases only the level group (e.g., flagellates) could be determined. To estimate cell size- and biovolume-based carbon biomass, unidentified species were sorted by size (e.g., flagellates in the 2–5 μm , 5–10 μm categories; diatoms in 10–20 μm ; 20–30 μm etc.). Diatoms were sorted in centric and pennates. The biovolumes of the phytoplankton cells were calculated according to Hillebrand et al. (1999) and converted into carbon content (i.e., biomass) using the equations of Strathmann (1967), considering the abundance of cells in the samples.

Chlorophyll *a* (Chl *a*) Concentration

For the extraction of photosynthetic pigments, cells from 30 to 100 mL microcosm aliquots were collected on Munktell MG-F glass fiber filters (25 mm), and put in 15 mL centrifuge tubes with 5 mL of absolute methanol (Holm-Hansen and Riemann, 1978). The tubes containing the filters were then placed in a sonicator for 20 min at 20°C , and in dark for 40 min. The extracted samples were centrifuged for 15 min at 1,500 rpm (20°C), and the supernatants were collected. Samples were scanned at 250 and 750 nm using a spectrophotometer (Hewlett Packard, model HP 8453E, United States) and Chl *a* concentration was calculated using the equations of Porra (2002). Chl *a* concentrations of the same samples were also determined fluorometrically (Holm-Hansen et al., 1965) using



a fluorometer (Turner Designs, Model Trilogy, United States) routinely calibrated against spectrophotometric measurements. There were no differences between the Chl *a* concentration calculated from both instruments, so the fluorometer data was used for all calculations.

Carbon Incorporation

Sub-samples for the determination of carbon incorporation were collected from the microcosms (in Phase II of the short-term, and on days 1 and 5 of the long-term experiment), and put into 30 mL Teflon bottles, inoculated with 0.185 MBq of radiocarbon (Holm-Hansen and Helbling, 1995) and placed inside a water-bath for temperature control. Samples were exposed to solar radiation (1 h) in the morning (9–10 am), noon (12–1 pm), and afternoon (3–4 pm). The inoculated samples were then immediately filtered through 0.7 μm Munktell GF/F filters (25 mm in diameter) under low pressure (<100 mm Hg). Filters were then put into 20 mL scintillation vials, and the excess of inorganic carbon was removed by exposing the filters to HCl fumes for 24 h. After acidification and following the addition of a scintillation cocktail (Optiphase Hisafe 3), samples were counted using a liquid scintillation counter. The rate of carbon incorporation was normalized by the Chl *a* concentration to calculate the assimilation number (P_m^B , in $\mu\text{mol C } \mu\text{g Chl } a^{-1} \text{ h}^{-1}$).

Variable Chlorophyll Fluorescence Parameters

Sub-samples from Phase II of the short-term, and from days 1 and 5 of the long-term experiment (Figure 2) were collected from the microcosms (triplicate for each scenario), put into 30 mL Teflon bottles in the outdoor water-bath for temperature control and exposed to solar radiation for 1 h in the morning (9–10 am), noon (12–1 pm), and afternoon (3–4 pm). Aliquots (2 mL) were taken from the Teflon bottles at the beginning, middle and end of the 1 h incubation period, and transferred to the measuring FastAct head of the benchtop FastOcean FRR (Chelsea Technologies Group, United Kingdom) for the photo-physiological measurements via fast repetition rate fluorometry (FRRf). Samples in the measuring chamber were kept at the outdoor temperature (14°C) using a laboratory thermostat. First, dark measurements were performed to derive the intrinsic and maximum Chl fluorescence F_o and F_m and the effective

absorption cross-section of Photosystem II (σ). In order to analyze the *in situ* photo-physiology of the phytoplankton assemblage, samples were then exposed to actinic light intensity provided by the white LEDs of the FastAct illumination system attached to FastOcean that was set to the *in situ* light intensity, as measured by the local outside radiometer. The measurements in the presence of actinic light provided parameters F' and F_m' . All measurements were conducted using a FastPro8 software. The single turnover FRR protocol consisted of 90 flashlets of 1 μs duration using the blue (450 nm) LED illumination; 70 sequences were averaged and then fitted using the equations of Kolber et al. (1998). A blank was measured every day by filtering the sample through a 0.2 μm filter. The F_o , F_m , F' , F_m' values, measured for the blank, were then subtracted from the respective parameters measured in the samples. The maximum efficiency of absorbed light used for photochemistry (F_v/F_m) was estimated by the FRRf ST protocol during the dark-regulated state as $((F_m - F_o)/F_m)$. The absolute electron transport rate (ETR) normalized to Chl *a* [$\mu\text{mol electrons } (\mu\text{g Chl } a)^{-1} \text{ h}^{-1}$] was calculated using the “absorption” method of Oxborough et al. (2012) as:

$$ETR_{\text{Chl}} = \frac{F_m \cdot F_o}{(F_m - F_o)} \cdot \frac{(1 - F')}{F_m} \cdot \frac{E \cdot K_A}{[\text{Chl } a]} \cdot 1/3600 \quad (1)$$

where F_m , F_o , F' , F_m' are the blank corrected Chl variable fluorescence parameters defined above, E is the intensity of the set irradiance (in $\mu\text{mol photons } \text{m}^{-2} \text{ s}^{-1}$), K_A is the instrument-specific calibration factor ($11,800 \text{ m}^{-1}$), 3,600 is factor to convert seconds to hours and $[\text{Chl } a]$ is the Chl *a* concentration (in $\text{mg } \text{m}^{-3}$).

To estimate the overall responses to actinic light, photosynthesis-irradiance (P-E) curves were measured during the morning (9–10 am), noon (12–1 pm), and afternoon (3–4 pm). ETR vs. E curves were obtained by exposing the cells to increasing irradiances from 0 to 2,278 $\mu\text{mol photons } \text{m}^{-2} \text{ s}^{-1}$ using 11 light steps at 20 s interval between each steps. Each measurement was conducted in triplicates. Light use efficiency (α), light saturation index (E_k), and the maximum electron transport rate (ETR_{max}) were then estimated by fitting the data to the model using equations of (Eilers and Peeters, 1988).

Net Community Productivity

Sub-samples taken from the microcosms were placed in 30 mL Teflon bottles (Nalgene) and exposed to solar radiation for 1 h inside the outdoor water-bath for temperature control. Oxygen concentration was measured using an optode-probe system (Minioxy-10 Presens GmbH, Germany) equipped with fiber optics and sensor-spots (SP-PSt3-NAU-D5-YOP), together with the Oxyview 6.02 software, to register the data. The system was calibrated before the measurements using a two-point calibration for 100 and 0% oxygen saturation, at the desired temperature and at atmospheric pressure. Oxygen concentration measurements were done three times in each period (i.e., at 9, 9:30 and at 10 a.m.; at 12, 12:30 a.m. and 1 p.m.; and at 3, 3:30 and 4 p.m.), with each measurement lasting 1 min. Hence, ca. 20 data were obtained during each measured period. Net community productivity (NCP, in $\mu\text{mol O}_2 \mu\text{g Chl } a^{-1} \text{ h}^{-1}$) was calculated by normalizing the change of oxygen concentration during the 1 h of exposure by Chl *a* concentration.

Computation of C Chemistry

We have applied the following equations for the computation of C chemistry (Emerson and Hedges, 2008):

$$[\text{CO}_2] = \frac{[\text{DIC}]}{1 + \frac{K_1}{[\text{H}^+]} + \frac{K_1 K_2}{[\text{H}^+]^2}} \quad (2)$$

$$[\text{HCO}_3^-] + \frac{[\text{DIC}]}{\frac{[\text{H}^+]}{K_1} + 1 + \frac{K_2}{[\text{H}^+]}} \quad (3)$$

$$[\text{CO}_3^{2-}] = \frac{[\text{DIC}]}{1 + \frac{[\text{H}^+]^2}{K_1 K_2} + \frac{[\text{H}^+]}{K_2}} \quad (4)$$

where [] represents concentrations. We have computed $[\text{H}^+]$ from pH and used 2191 (μM) for $[\text{DIC}]$, which is the expected value based on the method we used (Gattuso et al., 2010). For the pH values in PM, we used a value (8.102) from an earlier study (Kahl, 2018) close to PM; K_1 and K_2 were obtained by the following equations based on (Lueker et al., 2000; Emerson and Hedges, 2008):

$$K_1 = 10^{-pK_1} \quad (5)$$

$$K_2 = 10^{-pK_2} \quad (6)$$

where

$$pK_1 = \frac{3633.86}{T} - 61.2172 + 9.6777 \ln(T) - 0.011555S + 0.0001152S^2 \quad (7)$$

$$pK_2 = \frac{471.78}{T} + 25.9290 - 3.16967 \ln(T) - 0.01781S + 0.0001122S^2 \quad (8)$$

where T is temperature (in K) and S is salinity.

$p\text{CO}_2$ is approximated by the following equation (Weiss, 1974; DOE, 1994; Emerson and Hedges, 2008):

$$p\text{CO}_2 \sim \frac{[\text{CO}_2]}{K_{H, \text{CO}_2}} \quad (9)$$

where

$$K_{H, \text{CO}_2} = \exp \left\{ \frac{9345.17}{T} - 60.2409 + 23.3585 \ln \left(\frac{T}{100} \right) + S \left[0.023517 - 0.00023656T + 0.0047036 \left(\frac{T}{100} \right)^2 \right] \right\} \quad (10)$$

Data Treatment and Statistics

Specific growth rates (μ) in each scenario were calculated using the total cell abundances as:

$$\mu = \frac{\ln \left(\frac{N_x}{N_0} \right)}{(t_x - t_0)} \quad (11)$$

where N_0 and N_x represent the initial and final cell concentration at the initial time (t_0) and at the end (t_x) of each sampling period.

For the short-term experiments, differences in carbon incorporation, NCP, ETR, P_m^B and $\Phi_{e,C}$, as the ratio of ETR and P_m^B (Lawrenz et al., 2013), were evaluated using two-way repeated measures ANOVAs with sites (PM, PU, and IE) and scenarios (Present and Future) as factors, and period of the day (i.e., morning, noon, and afternoon) as the repeated measures. When significant differences were obtained, *a posteriori* Tukey test was performed. In these experiments, limitation of sample volume precluded statistical analysis among samples for Chl *a* concentration, nutrients and phytoplankton abundance.

For the long-term experiment, differences in biomass and abundance between micro and nanoplankton cells, or among the most representative phytoplankton groups (i.e., centric diatoms, pennate diatoms, and flagellates) were evaluated using three-way ANOVAs, with size (micro vs. nanoplankton), groups (centric diatoms, pennate diatoms, or flagellates), scenarios (Present and Future), and acclimation (no, mid, or long-acclimation), as factors. The same type of analysis was done for the growth rates, but to compare μ for days 1–3 and 3–5, to assess the differences in mid- and long-acclimation. For carbon incorporation, NCP, ETR, P_m^B , and $\text{ETR} (P_m^B)^{-1}$, two-way repeated measures ANOVAs were performed to evaluate differences in the measured parameters during the period of the day (i.e., morning, noon, and afternoon), with the scenario (i.e., Present and Future) and the acclimation (no-, mid-, and long-acclimation) as factors. When significant differences were determined, *a posteriori* Tukey test was performed. For comparisons of Chl *a* among acclimation time and between scenarios, we performed a repeated measures ANOVA.

In all cases, the data sets were tested for normality, homoscedasticity and sphericity using Kolmogorov-Smirnov, Cochran and Mauchly's tests, respectively. If the data did not meet sphericity, we adjusted them using Greenhouse-Geisser correction. Although in some cases the data did not meet the assumptions, we decided to continue with the ANOVA's tests, as

for a large and balanced design, ANOVA are highly robust to such deviations (Underwood, 1997).

RESULTS

In situ Conditions at the Sampling Sites and Radiation Levels During the Experiments

The initial conditions at sampling time, as well as the irradiance during the acclimation and measurement periods for each site are shown in **Table 1**. In the short-term experiments, the temperature was rather similar among the study sites (13.5–14°C), while pH was lower at PU (8.07) compared to IE (8.34). Nutrient concentrations presented some differences among the sites, with PM having the lowest nitrate + nitrite values (0.18 μM), whereas PU the highest (0.55 μM), during the short-term experiment; phosphate concentrations, on the other hand, were similar in PM and PU (1.29 and 1.33 μM , respectively), and almost two times higher than that in IE (0.76 μM). Silicate concentrations displayed the largest differences among sites, with PU having ca. 10 times more than the other (26.6 μM in PU, 2.69 and 2.51 μM in PM, and IE, respectively, **Table 1A**). Initial Chl *a* concentrations were low in the three sampling sites, with IE being the lowest (0.39 $\mu\text{g L}^{-1}$) and PU the highest (1.17 $\mu\text{g L}^{-1}$). For the long-term experiment, the initial conditions were similar to that of short-term experiment, except for the nitrate + nitrite concentrations, which were much higher in the former case (7.89 μM ; **Table 1A**), suggesting an extra input via riverine waters. The influence of clouds on solar PAR was generally insignificant during the experiments, so the samples were exposed to rather similar irradiances at the same time of the day for the different experiments. An exception to this was the afternoon of October 20th, when the long-term experiment had much lower PAR values (**Table 1B**).

Carbon Chemistry

The estimated CO_2 concentrations at the sampling time varied between 7.6 and 15.2 μM , during both short- and long-term experiments. pH reduction (by 0.4 units) increased these estimated CO_2 concentrations from 20.6 to 39.6 μM (**Table 2**).

TABLE 2 | Estimated concentrations of CO_2 , HCO_3^- , CO_3^{2-} before and after the pH-reducing treatment.

Molecules/Site		PM	IE	PU (Short-term)	PU (Long-term)
CO_2 ($\mu\text{mol kg}^{-1}$)	Before	13.8	7.6	15.2	9.1
	After	36.2	20.6	39.6	24.3
HCO_3^- ($\mu\text{mol kg}^{-1}$)	Before	1996.3	1887.6	2011.8	1929.8
	After	2079.8	2043.0	2083.8	2059.5
CO_3^{2-} ($\mu\text{mol kg}^{-1}$)	Before	180.9	295.8	164.0	252.2
	After	75	127.5	67.7	107.1
pCO_2 (μatm)	Before	356.3	194.8	387.5	228.7
	After	932.4	529.5	1008.1	613.0

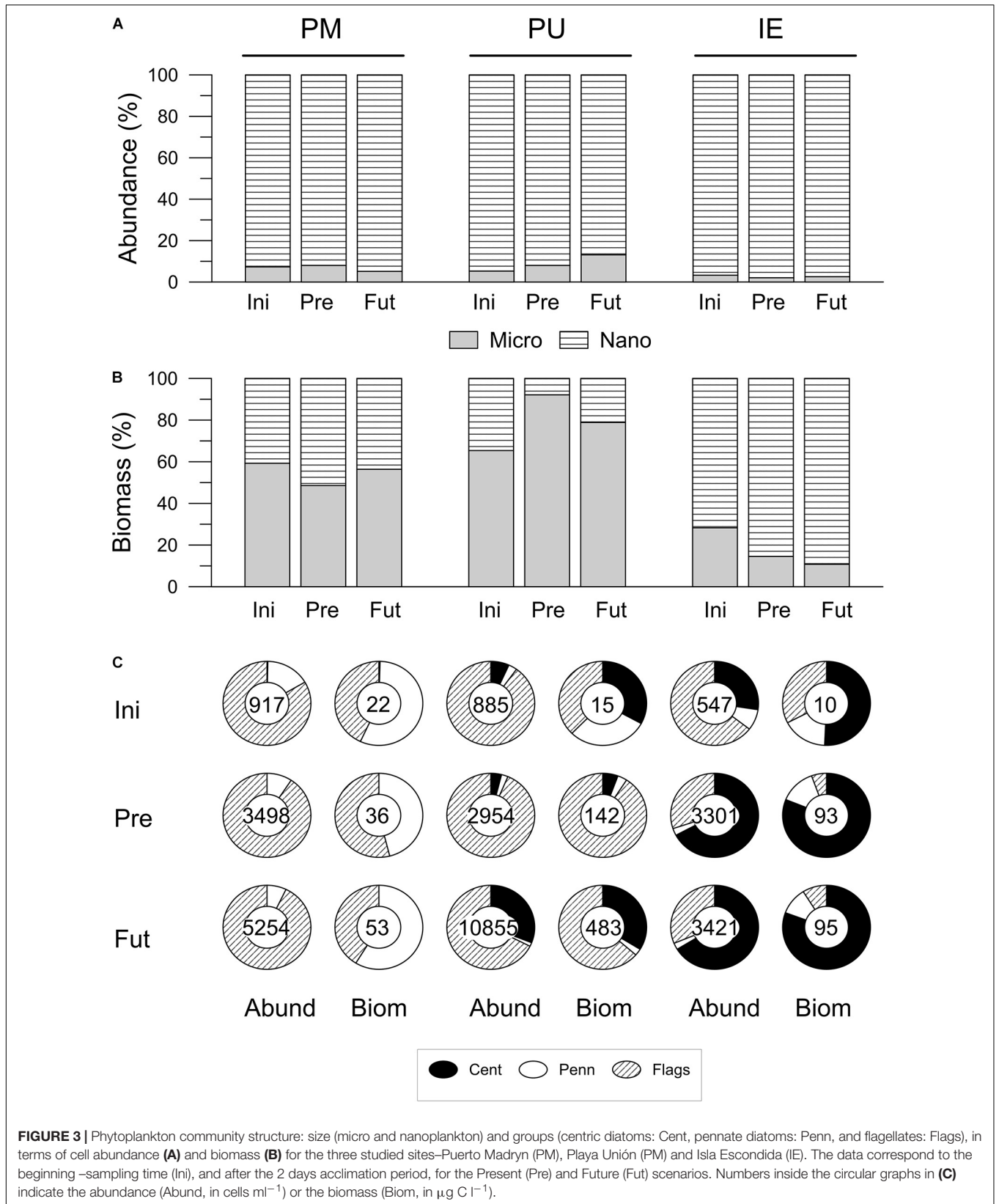
The biggest change (24.4 μM) was observed at PU during short-term experiment. pCO_2 increased from 387.5 to 1,008 μatm as expected (**Table 2**). The pCO_2 at all sites and experiments increased by a comparable amount as the ratio of increased pCO_2 (i.e., values after divided by values before the acid addition) varied between 2.60 and 2.71.

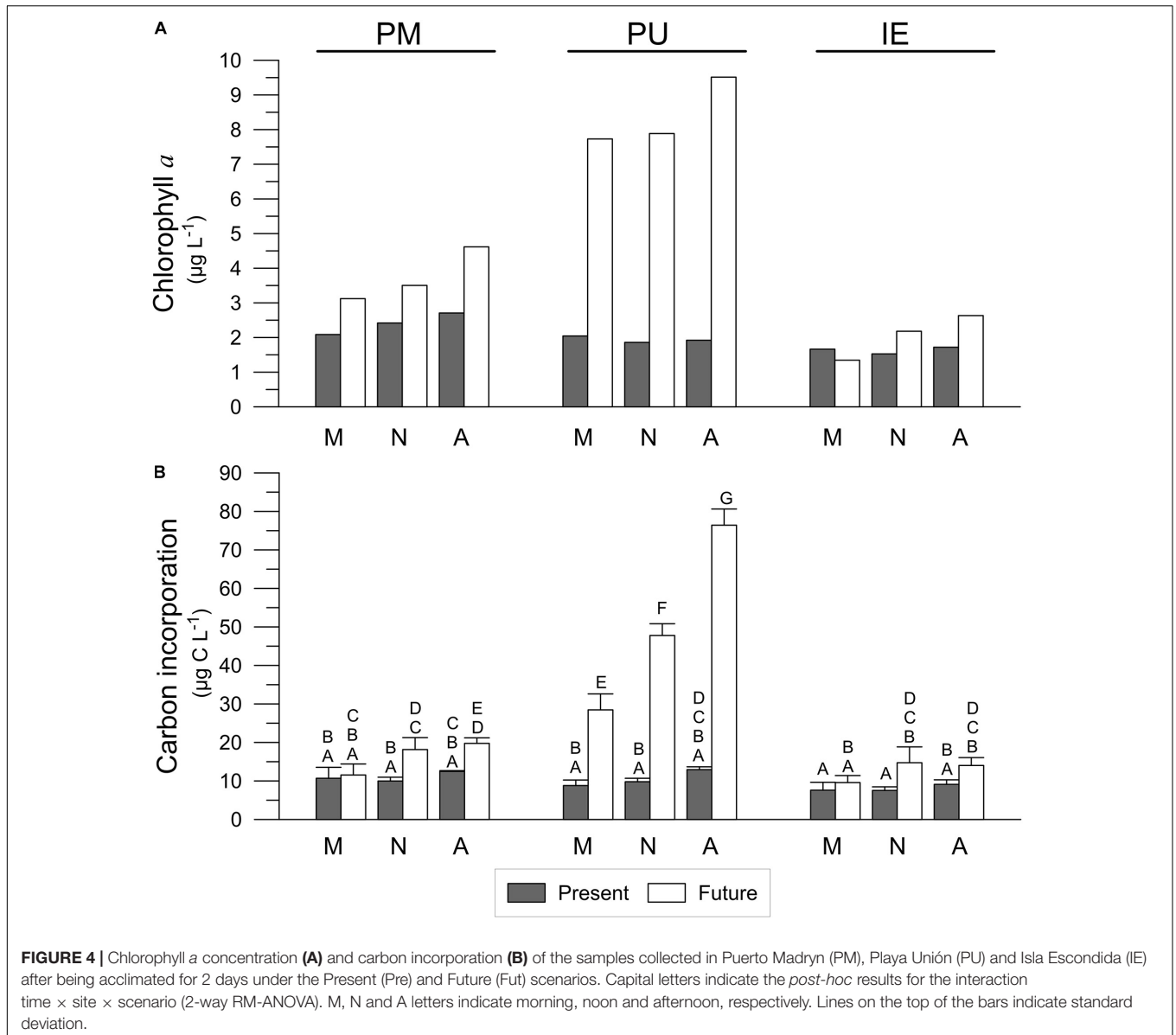
Short-Term Experiments (PM, PU, IE) Relative Abundance, Growth, and Changes in the Phytoplankton Community Structure

In terms of abundance, nanoplanktonic cells (<20 μm) dominated (92%) in all the three sites (**Figure 3A**); however, in terms of carbon biomass, microplanktonic cells (>20 μm) had the largest share (>50%) in PM and PU, whereas nanoplankton biomass was higher in IE (**Figure 3B**). The three sites were initially characterized by different proportions of phytoplankton groups (**Figure 3C**), but in terms of abundance, unidentified flagellates dominated in all sites (65–90%). The dominance of unidentified flagellates in terms of abundance was more evident at PU, compared to PM and IE. The rest of the community in PM was composed of pennate diatoms only (e.g., *Grammatophora marina*, *Nitzschia longissima*, and *Navicula* spp.), while this group were present together with centric diatoms (e.g., *Thalassiosira* spp.) in PU and IE (**Figure 3C**). In PM there were small changes in the size composition, both in terms of abundance (**Figure 3A**) or biomass (**Figure 3B**) in the two scenarios, at the end of the 2 days acclimation period. Nevertheless, the total abundance and biomass increased in both experimental scenarios (numbers inside the circular graphs; **Figure 3C**). A similar situation was observed in PU, but the Future scenario stimulated the growth of centric diatoms (e.g., *Thalassiosira* spp. and *Chaetoceros* spp.) compared to the Present (e.g., mainly *Thalassiosira* spp.), reaching the highest abundances and carbon biomass of the three sites (**Figure 3C**). Similarly, in IE, centric diatoms (*Thalassiosira* spp.) increased in abundance and biomass from the initial value, reaching similar values in the Present and Future scenarios (**Figure 3C**). The dominance of centric diatoms in abundance and biomass was the highest in IE, and the lowest in PM (**Figure 3C**). Overall, small cells dominated at the end of the acclimation period in PM, with the carbon concentration per cell decreasing from 24 ng C cell⁻¹ in the initial sample to ~10 ng C cell⁻¹ in the Present and Future scenarios. On the other hand, large cells dominated in PU, with the carbon concentration per cell increasing from 17 ng C cell⁻¹ in the initial sample to 48 and 44.5 ng C cell⁻¹, for the Present and Future scenarios, respectively, having the largest change throughout the experimental period as compared to the other sites. Although cellular carbon content also increased in IE after the acclimation period, this change was smaller compared to PU, as in the initial sample the cellular carbon content was 18 ng C cell⁻¹ and at the end these values reached 28.2 and 27.8 ng C cell⁻¹, for both the Present and Future scenarios, respectively.

Photosynthetic Rates

Photosynthetic rates were evaluated using three different and independent measurements: (a) net community production NCP [in $\mu\text{mol O}_2$ ($\mu\text{g Chl } a$)⁻¹ h⁻¹], the net budget of photosynthetic





oxygen production and active respiratory processes by all microorganisms present, (b) electron transport rates, ETR (in $\mu\text{mol e}^- (\mu\text{g Chl } a)^{-1} \text{ h}^{-1}$), the electrons passed through photosystem II (PSII), and (c) assimilation numbers, as P_m^B [in $\mu\text{mol C} (\mu\text{g Chl } a)^{-1} \text{ h}^{-1}$], the inorganic carbon incorporated in the photosynthetic Calvin–Benson–Bassham cycle. All photosynthetic parameters were normalized to the Chl *a* concentrations measured at the same sampling periods. Chl *a* was in general higher in the Future scenario compared to the Present (Figure 4A) in all sites, being this increase especially evident in PU, and less in IE. In addition, Chl *a* concentrations increased along the day in the Future scenario in all sites, suggesting an active growth during the light period (Figure 4A). A similar trend was observed for bulk carbon incorporation (Figure 4B and Table 3), with values increasing during the day (especially in samples under the Future scenario) and these being

significantly higher in the Future scenario (28.5 ± 3.4 – $76.4 \pm 3.4 \mu\text{g C L}^{-1}$) than in the Present (8.8 ± 1.2 – $12.9 \pm 0.6 \mu\text{g C L}^{-1}$) for PU during all the days. For PM the significant differences between Present and Future scenarios were observed at noon and afternoon, while for IE only at noon. Compared to PM and IE, PU had the highest values of carbon incorporation in the Future scenario ($76.4 \pm 3.4 \mu\text{g C L}^{-1}$; Figure 4B and Table 3).

For NCP, there was a significant interaction between the period of the day, site and scenario (Figure 5A and Table 3), with higher values in the morning, which decreased during the day. In PM and IE, negative values were determined, indicating that community respiration surpassed oxygen production by phototrophs (Figure 5A). Significant differences between Future and Present scenarios were observed only for the PU samples, with NCP values being higher in the Present [$3.21 \pm 0.12 \mu\text{mol O}_2 (\mu\text{g Chl } a)^{-1} \text{ h}^{-1}$] than in the Future [1.04 ± 0.05

$\mu\text{mol O}_2$ ($\mu\text{g Chl } a$) $^{-1} \text{ h}^{-1}$] during the morning measurements. A significant interaction among the period of the day, site and scenario was also determined for ETR (Figure 5B and Table 3), but it increased during the day in both scenarios. Differences among the period of the day were only significant in PU, with values during the noon and the afternoon being higher than in the morning. ETR values were similar between the Present and Future scenarios in all sites with the exception of samples measured during the morning in IE, where higher values were registered in the Future (Figure 5B and Table 3). Finally, P_m^B values were similar along the day for each site; however, there was a significant interaction between the period of the day and site, with values in PU increasing from morning to afternoon, especially when compared to PM and IE (Figure 5C and Table 3). Moreover, a significant interaction between sites and scenario was determined in the Present scenario, with higher P_m^B values in PU than in PM and IE (Figure 5C and Table 3). There was also a significant interaction among the period of the day, site and scenario for the electron requirement for carbon fixation $\Phi_{e,C}$ (calculated as ETR/P_m^B at *in situ* irradiance) (Figure 5D and Table 3) with values varying between 5.8 and 11.8, and being generally (but not significantly) higher in the Present than in the Future scenario.

Long-Term Experiment (PU) Abundance of Cells, Growth and Changes in the Phytoplankton Community Structure

Throughout the long-term experiment at PU, nanoplanktonic cells were dominant in terms of abundance in the both scenarios, accounting for > 86% of the phytoplankton community (Figure 6A). In terms of biomass, microplanktonic cells were the most represented group, except for the samples at the end of the acclimation period (day 5), where nanoplankton biomass (~53%) surpassed it slightly (Figure 6C). Unidentified flagellates and centric diatoms (e.g., *Thalassiosira* spp.) were the most representative groups in terms of both, abundance (Figure 6B) and biomass (Figure 6D); pennate diatoms (mainly *Asterionellopsis glacialis*) only contributed in a small proportion

(<19%, Figures 6B,D). Cell abundances, either based on size or groups, increased throughout the experiment, reaching their highest values in the Future scenario at the end of the acclimation period (Figures 6A,B). In day 5 significant differences between scenarios were detected, with higher nano- and microplankton abundances under the Future scenario (Figure 6A and Table 4), and with significant differences only for centric diatoms (Figure 6B). Both cell size- and group-based biomass (Figures 6C,D) increased throughout the acclimation period, having significant differences between the Present and the Future scenarios after 5 days of acclimation. Growth rates based either on size (Figure 6E) or groups (Figure 6F) decreased toward the end of the acclimation period (long acclimation). During the first half of the experiment (mid-acclimation) no differences were detected between Present and Future scenarios; however, at the end of the acclimation period, growth rates were significantly higher under the Future (Figures 6E,F) compared to the Present. There were significant interactions between scenarios and groups (Figure 6F and Table 4); for example, higher growth rates were recorded for centric diatoms in the Future than in the Present scenario (Figure 6F).

Photosynthetic Rates

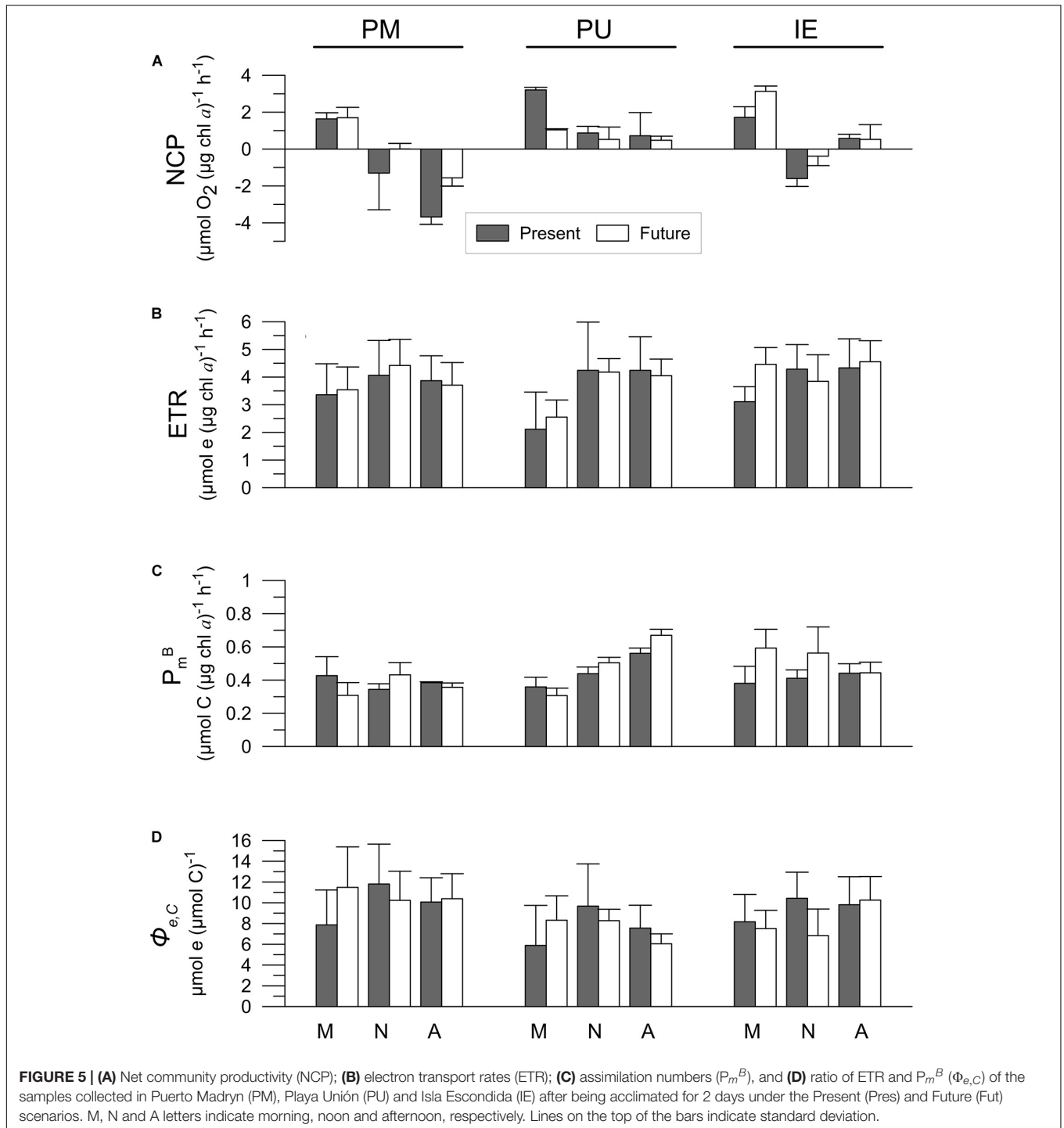
Chl *a* concentration increased during the long-term experiment at PU until the third day of experiment from an initial value of 2.79 and 2.57 $\mu\text{g L}^{-1}$ in the Present and Future scenario, respectively, but decreased afterward in the Present, while it kept increasing in the Future, reaching a concentration of ca. 60 $\mu\text{g Chl } a \text{ L}^{-1}$ (Figure 7A). Carbon incorporation (Figure 7B) was similar on day 1 (no acclimation) in both scenarios, but was significantly higher in the Future scenario as compared to the Present at the end of the acclimation period (long acclimation). Moreover, carbon incorporation also significantly increased on this day from morning to afternoon (Figure 7B and Table 5).

We found significant differences in the NCP for no-acclimated cells, with decreasing values toward noon and increasing again during the afternoon (Figure 8A and Table 5). In the no acclimation condition, respiration surpassed oxygen production only at noon (i.e., negative values) and also NCP was significantly lower in the Future than in the Present scenario (Figure 8A). Under long acclimation, NCP rates were much lower than at the beginning of the experiment, with no significant differences between scenarios, except for the Present in the afternoon, where NCP was significantly lower than in the Future (Figure 8A and Table 5). ETR values were similar in both scenarios in the no acclimation condition, but they were significantly higher in the Present than in the Future scenario at the end of the experiment (Figure 8B and Table 5). The highest ETR values were registered at the end of the experiment, during the afternoon measurements (Figure 8B). P_m^B showed a clear diurnal pattern of increasing values from morning to afternoon both at the beginning (no acclimation) and end (long acclimation) of the experiment in both scenarios (Figure 8C and Table 5). P_m^B values were higher in the no acclimated compared to the long acclimated cells, but only for the Present scenario (Figure 8C and Table 5). Finally, the electron requirement for carbon fixation $\Phi_{e,C}$ ratio was similar in both scenarios when cells were not acclimated (Figure 8D), but cells in the Present showed higher values compared to those

TABLE 3 | p -values of the two-way RM-ANOVAs for carbon incorporation (Figure 4B), net community production (NCP) (Figure 5A), electron transport rates (ETR) (Figure 5B), assimilation number (P_m^B) (Figure 5C), and ratio ETR to P_m^B ($\Phi_{e,C}$; Figure 5D).

	Carbon incorporation	NCP	ETR	P_m^B	$\Phi_{e,C}$
Sites	<0.001	<0.001	<0.05	<0.001	0.272
Scenario	<0.001	0.102	0.150	0.759	0.871
Sites × Scenario	<0.001	<0.05	0.345	<0.05	0.811
Period	<0.001	<0.001	<0.001	0.055	<0.001
Period × Sites	<0.001	<0.001	<0.001	<0.05	<0.001
Period × Scenario	<0.001	0.080	<0.001	0.154	<0.001
Period × Sites × Scenario	<0.001	<0.05	<0.05	0.194	<0.05

Sites represent PM, PU, and IE; Scenarios represent Present and Future, and Period represents those parts of the day when measurements were conducted (i.e., morning, noon, afternoon).

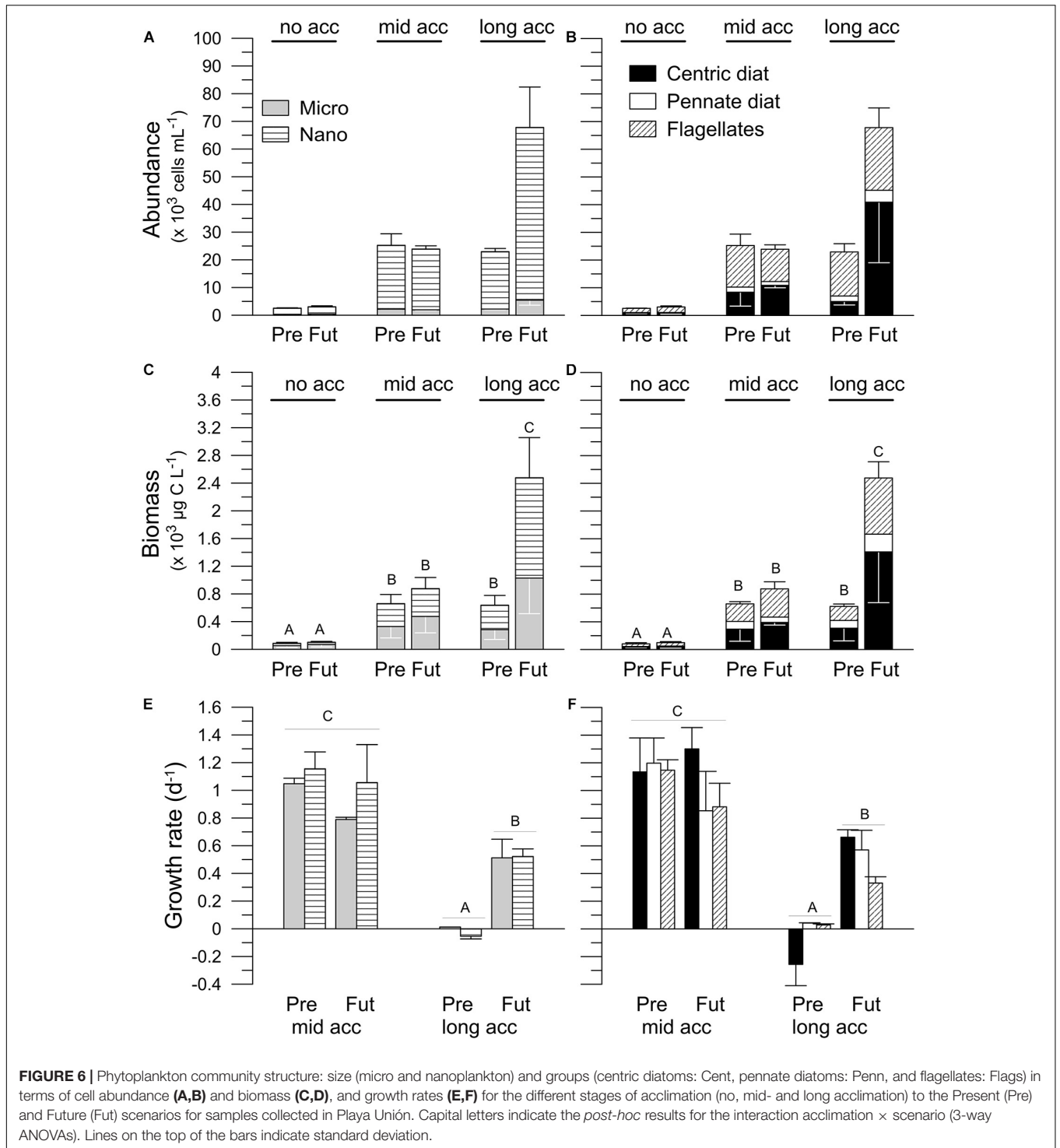


in the Future scenario at the end of the experiment (Figure 8D and Table 5).

DISCUSSION

In this study, we aimed to enhance our knowledge about the effects of global change variables on phytoplankton growth

and photosynthesis in three different sites of the Patagonian coast, by applying an experimental cluster (Future scenario). We especially focused on key components of the photosynthetic processes; the water splitting complex (by oxygen evolution measurements), PSII photochemistry (by PAM and FRRf measurements), and Calvin-Benson-Bassham cycle (by carbon incorporation measurement), in an attempt to determine the extents of their vulnerability to increased nutrient and lowered



pH conditions. This is a novel approach, as most field studies have generally considered the effects of global change variables on single photosynthesis targets. Our previous studies have determined that Future scenarios had differential impacts on oxygen dynamics and PSII photochemistry (Villafañe et al., 2018; Vizzo et al., 2021), and that acclimation over time changed the rETR vs. NCP relation (Villafañe et al., 2018).

We used the same photosynthetic parameters as in these previous studies to get a more realistic response to global change. We hope that our results, together with those obtained for the phytoplankton community structure (e.g., taxonomic changes, effects on different size fractions), will provide a better assessment of the secondary production in the Argentinean Sea. Consequently, we discuss our findings together with the potential

mechanisms that could explain our observations, and consider some of the implications for this area, which is believed to sustain one of the highest productivities of the World's oceans (Longhurst et al., 1995; Gregg and Rousseaux, 2019).

Future Conditions Stimulate Growth and Photosynthetic Activity, Especially in Estuarine Waters

The experimental sites in the South West Atlantic Ocean (SWAO) are normally exposed to intense winds (Gaiero et al., 2003; Gassó and Stein, 2007), and the water column over the continental shelf is generally well mixed (Neale et al., 2003). There is, however, a marked seasonal change, with intense winds in spring-summer and a relatively calm period during winter (Villafañe, 2004; Helbling et al., 2005; Bermejo et al., 2018). Previous reports for the area (Villafañe, 2004) indicate that phytoplankton blooms occur during winter (June–August), when wind speeds are low (Helbling et al., 2005; Fitch and Moore, 2007; Bermejo et al., 2018). Thus, our sampling period during October corresponded to post bloom (Villafañe, 2004). At PU, nutrient concentrations are relatively high during the whole year due to riverine input (i.e., from the Chubut river), which not only carries nutrients, but also particles and dissolved matter (Bermejo et al., 2018; Vizzo et al., 2021). It has been determined that the nutrient inputs due to riverine runoff are increasing over time (Bermejo et al., 2018) because of anthropogenic activity (i.e., farming, uses of land, fisheries by-products, etc.) and extreme rain events that are more frequent (Vizzo et al., 2021). On the other hand, PM and IE do not receive a constant input of nutrients, and their dynamics are more subjected to biological activity rather than anthropogenic input via a riverine runoff. Our results also highlight the different strategies of phytoplankton used for growth in the three sites for the effects of nutrient inputs.

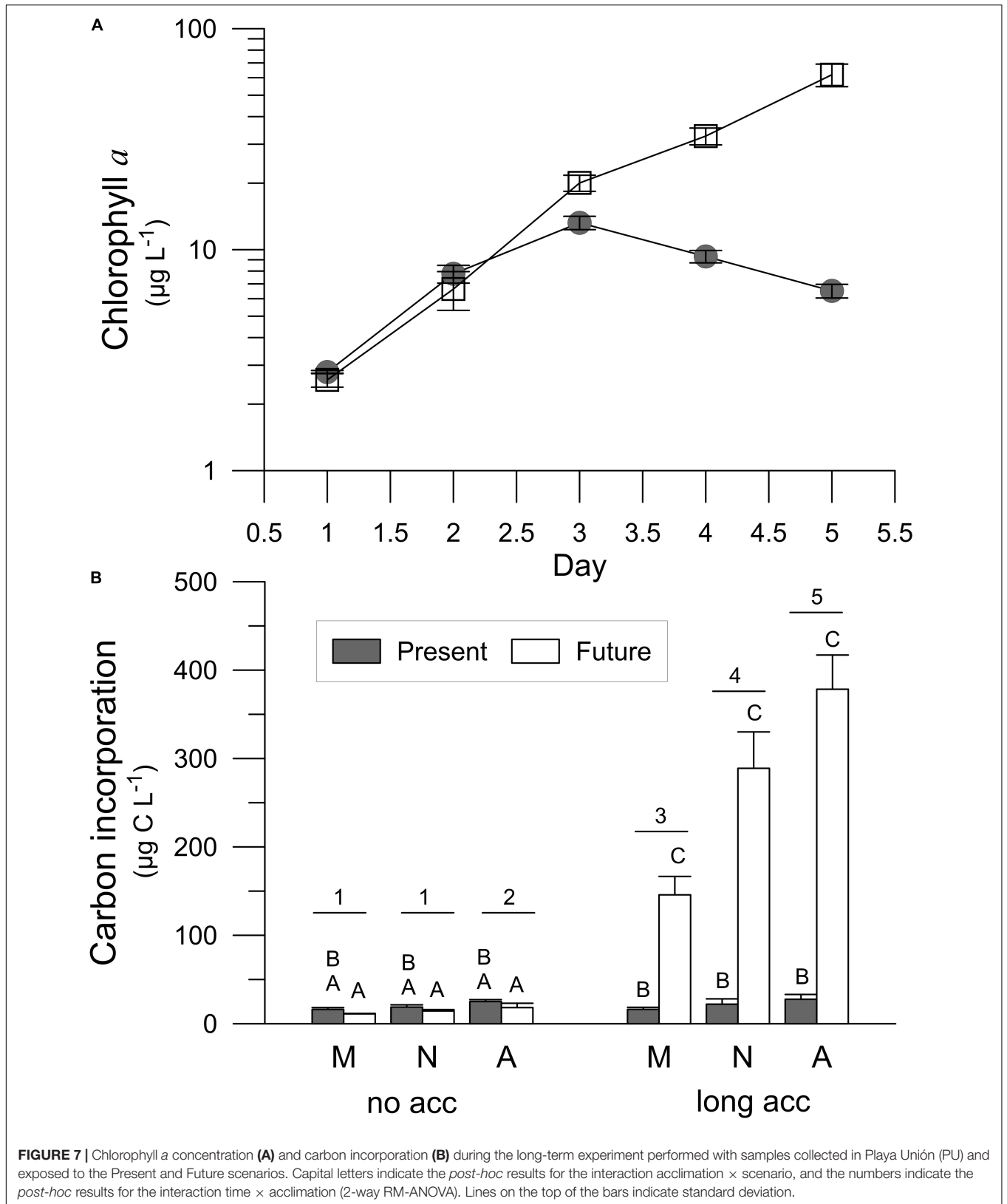
Fresh nutrient inputs seldom occur in PM and IE, while nutrients from riverine origin are received twice a day in PU (Helbling et al., 1992; Bermejo et al., 2018) due to the semi-diurnal tide regime (Piccolo and Perillo, 1999). It is known that, diatoms and larger cells are *r*-strategists, characterized by rapid growth under a nutrient-replete environment (Margalef, 1978; Thomsen et al., 1994; Litchman, 2007). On the other hand, dinoflagellates and most flagellates are expected to be *K*-strategists (Smayda and Reynolds, 2001; Litchman, 2007; Glibert, 2016), with slow growth and with higher nutrient affinity. It is recognized that the *r*-strategists allocate a large portion of resources to cell division, which is advantageous in a variable environment such as PU, whereas *K*-strategists would be favored in less variable environments such as PM and IE.

Cabrerizo et al. (2018) studied the impact of different nutrient sources (aeoline vs. riverine) on community structure, and determined that the combination of high solar UVR and nutrient addition decreased CO₂ sink capacity of coastal ecosystems in the SWAO by 27%, by changing the phytoplankton community toward heterotrophic nanoflagellates. Our Future generally stimulated the increase of biomass and carbon incorporation in the three sites, in spite of the dominance of different phytoplankton groups (Figures 3, 4). Nevertheless, some differences were observed among the sites; for example in PM, where new nutrient inputs are limited due to little water exchange (Rivas and Beier, 1990), we observed an increase in the abundance of small flagellates (Figure 3A) and pennate diatom biomass (Figure 3C), with only small changes in the photosynthetic rates (Figure 5). A similar behavior was observed at IE, with only small changes between the Present and Future scenarios (Figures 3, 5). In PU, however, growth—especially that of large centric diatoms such as *Thalassiosira* spp. (Figure 3C), and carbon incorporation (Figure 4) were significantly stimulated by the Future scenario. The differential growth of the large taxonomic groups in our study agrees with those previous findings where different communities were exposed to lowered pH conditions. For example, a study that compiled the results of 69 experiments on ocean acidification (Bach and Taucher, 2019) determined that diatoms responded to high CO₂ in ca. 60% of the experiments, and from this total, 56% indicated a positive response while 32% responded in a negative way. A recent study reported no major changes in the species composition of a dinoflagellate-dominated community under lowered pH in the Derwent River estuary, Australia (Nielsen et al., 2012). A similar study in the Godavari River Estuary (India) on the other hand reported that not only lowered pH but also nutrient supply play key role in shaping the community structure during microcosms experiments (Biswas et al., 2011). The growth rates at PU determined during the mid-acclimation (Figure 6E) were similar to those reported by Durán-Romero et al. (2017), while the growth rates for the long acclimation (Figure 6E) agrees with those reported by Villafañe et al. (2018); in this latter study the Future condition had higher rates than the Present, as observed in our study. The differential responses observed among these studies are most likely related to the acclimation period that not only change community structure but also condition the physiology of cells.

TABLE 4 | *p*-values of the three-way ANOVAs for phytoplankton abundance (Figures 6A,B), biomass (Figures 6C,D), and growth rates (Figures 6E,F) based on cell sizes (i.e., nanoplankton, microplankton) or groups (i.e., centric diatoms, pennate diatoms or flagellates).

	Abundance	Biomass	Growth rate
Cell size	<0.001	0.161	0.282
Acclimation	<0.001	<0.001	<0.001
Scenario	<0.05	<0.001	0.068
Cell size × Acclimation	<0.001	<0.05	0.280
Cell size × Scenario	<0.05	0.643	0.466
Acclimation × Scenario	<0.05	<0.001	<0.05
Cell size × Acclimation × Scenario	<0.05	0.569	0.691
Group	<0.001	<0.001	0.511
Acclimation	<0.001	<0.001	<0.001
Scenario	<0.001	<0.05	<0.05
Group × Acclimation	0.511	0.199	0.469
Group × Scenario	<0.05	0.443	<0.05
Acclimation × Scenario	<0.001	<0.05	<0.001
Group × Acclimation × Scenario	<0.05	0.695	0.743

Scenarios are Present and Future. Acclimation refers to the duration (no, mid- and long-acclimation).



Villafañe et al. (2018) addressed the dual role of DOM carried by the Chubut river, and determined that the photosynthesis of cells exposed to solar radiation with an input of DOM inside the microcosms (i.e., physico-chemical role, attenuation of solar radiation and chemical exposure) were less efficient compared to cells that received an extra input of DOM outside the microcosms (i.e., physical role, attenuation of solar radiation). In our study, using the same global change variables, NCP values at PU were positive throughout the day, with significant diurnal increases of ETR from morning to noon and continuous increases in P_m^B throughout a day, unlike in the other sites (Figure 5). The estuarine area around PU had high primary productivity as it was highlighted both in our previous (Villafañe, 2004) and current study (Figure 4), and in a recent review about primary production in the SWAO (Lutz et al., 2018). Better fitness of phytoplankton observed in PU could also be due to the fact that here, compared to the other sites, phytoplankton were less stressed by the high-light conditions, also evidenced by the positive NCP values and the generally lower electron requirement for carbon fixation ($\Phi_{e,C}$) during the day (Tables 1, 3 and Figures 5A–C). On the other hand, no significant effect on the biomass was observed, when pCO_2 was manipulated, but competition among species and its growth phase lead succession in phytoplankton species composition and abundance (Osma et al., 2020). Better diatom fitness in growth under lowered pH at the PU site may also be due to their adaptation to a larger carbonate chemistry variability related to carbon concentration mechanisms (CCMs) (Burkhardt et al., 2001; Reinfelder, 2011), which often occurs in estuarine environments (Duarte et al., 2013). Although increasing temperatures stimulate gene expression and RuBisCO activity in diatoms (Helbling et al., 2011), free ocean CO_2 is a limiting factor for RuBisCO, the primary enzyme of carbon fixation (Beardall and Raven, 2004). To overcome this constraint, many phytoplankton species have developed an intracellular mechanisms that converts bicarbonate ions (HCO_3^-) to CO_2 , supporting higher carbon fixation rates than what would be

possible if photosynthesis only relied on diffusive CO_2 uptake (Giordano et al., 2005). Since the efficiency and regulation of CCMs differ among the species (Burkhardt et al., 2001; Trimborn et al., 2008), the proportion of direct CO_2 uptake and HCO_3^- uptake also varies among the phytoplankton species (Matsuda and Kroth, 2014). Increased CO_2 may benefit those phytoplankton species that rely on CO_2 uptake, as well as CCM efficient species that can downregulate CCM (Rost et al., 2008; Eggers et al., 2014; Taucher et al., 2015). The downregulation of CCM under elevated CO_2 has been observed for two diatom species (*Thalassiosira weissflogii* and *Dactyliosolen fragilissimus*) under nutrient limited condition (Taucher et al., 2015).

Influence of Acclimation Period on the Synergistic Interaction of CO_2 and Nutrient Enrichment

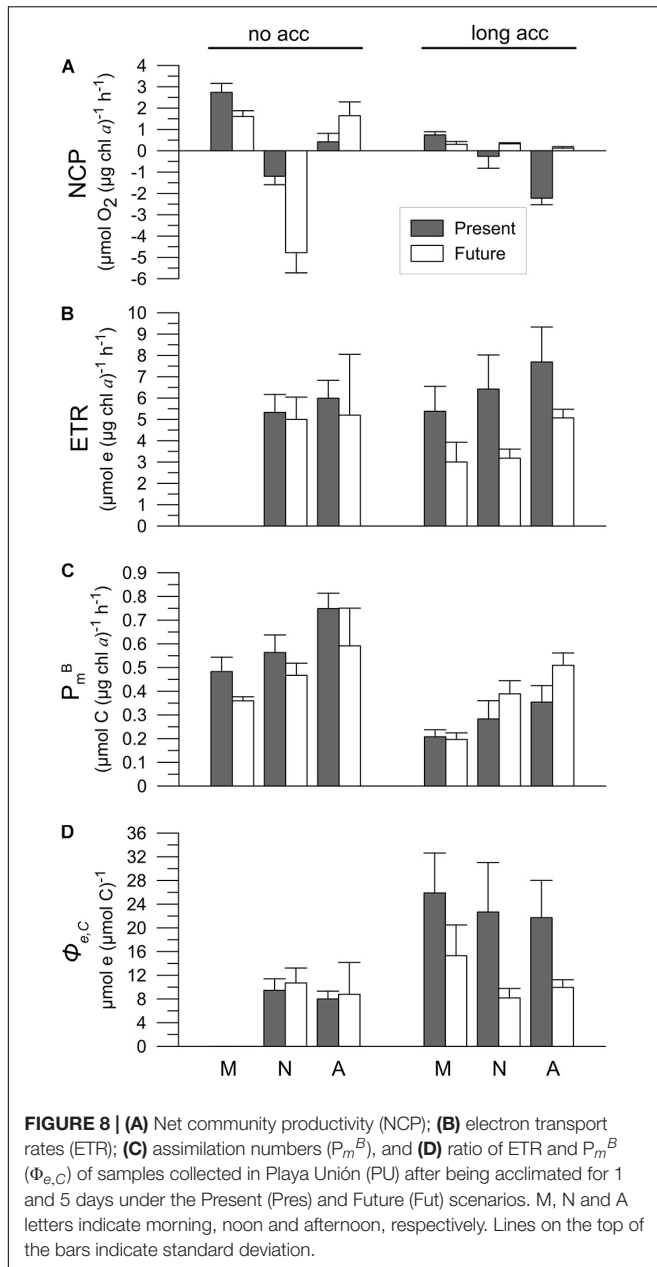
Earlier studies suggested a controversial synergistic interaction among increased CO_2 and nutrient concentrations. For example, beneficial effects of high CO_2 on phytoplankton (Domingues et al., 2016) including cyanobacteria (Hutchins et al., 2007; Fu et al., 2008), diatoms (Sun et al., 2011; Tatters et al., 2012) and dinoflagellates (Fu et al., 2010) were reported under nutrient-replete conditions, whereas Hama et al. (2016) reported significant effect of acidification during post-blooming period, when ambient nutrient concentrations are negligible. In our study, the positive interaction of lowered pH and increased nutrients was observed after 2 days of acclimation during the short-term experiment (Figures 3, 5). The positive synergistic effect of the Future treatment on abundance, biomass, Chl concentrations and carbon incorporation was observed also after 4 days of acclimation during the long-term experiment (Figures 6, 7). On the other hand, the lower growth rates after 3 days in both treatments (Figure 6E) suggest that the cells were stressed due to a gradual exhaustion of nutrients and/or inorganic carbon at the end of the long-term experiment. Stress was more pronounced in the Present treatment of this long-term experiment, and was also manifested as Chl *a* decline and a stagnant C incorporation between the days 1 and 5 (Figure 7). Stress in the later phase of the long-term experiment caused the loss of NCP diurnal dynamics (Figure 8A), the uncoupling between the photosynthetic membrane electron transport (ETR, Figure 8B), and a declining C-fixation (P_m^B , Figure 8C). This uncoupling was more pronounced under the Present conditions, resulting in the increase of $\Phi_{e,C}$ values (Figure 8D). The most probable reason behind this decline in the Present scenario was nutrient limitation, as after only 2 days the amount of nitrate decreased from its initial value of $7.89 \mu M$ to $0.05 \mu M$ (Table 1).

We noted differences in the dominance of the phytoplankton groups, but also in the growth and other photosynthetic parameters measured between the short- and long-term experiments (cf. Figures 3–5 with Figures 6–8). These differences between the experiments carried out in the same site (PU) highlight the role of nutrients and/or light intensity in the observed responses, as seen in the higher nitrate availability at the beginning of the long-term and compared with the

TABLE 5 | p -values of the two-way RM-ANOVAs for carbon incorporation (Figure 7B), net community production (NCP) (Figure 8A), electron transport rates (ETR) (Figure 8B), assimilation number (P_m^B) (Figure 8C), and ratio of ETR and P_m^B ($\Phi_{e,C}$; Figure 8D).

	Carbon incorporation	NCP	ETR	P_m^B	$\Phi_{e,C}$
Acclimation	<0.001	0.222	0.357	<0.001	<0.05
Scenario	<0.001	0.380	<0.001	0.490	<0.05
Acclimation × Scenario	<0.001	<0.001	<0.001	<0.05	<0.05
Period	<0.001	<0.001	<0.001	<0.001	0.253
Period × Acclimation	<0.001	<0.001	<0.05	0.486	0.077
Period × Scenario	0.053	<0.001	0.845	0.291	0.304
Period × Acclimation × Scenario	0.170	<0.001	0.138	0.155	0.162

Scenarios are Present and Future; Period represents those parts of the day when measurements were performed (i.e., morning, noon, afternoon). Acclimation refers to the duration (no, mid- and long-acclimation).



short-term experiment (Table 1). These results suggest that the effect of acidification might be much more pronounced when ambient nutrient concentrations are negligible, as it was observed by Hama et al. (2016), or could be because of downregulation of CCMs.

The high initial nitrate concentrations in PU are due to an active exchange between the river and the sea, which are the characteristics of estuarine waters (Helbling et al., 1992; Bermejo et al., 2018), which include anthropogenic nutrient loading (Kennish, 1992; Paerl et al., 2006). As reported by Bermejo et al. (2018), the main nutrient source in PU is the riverine water, which is influenced by rainfall (Vizzo et al., 2021) and possibly also from the fish processing factories

(Arvanitoyannis and Kassaveti, 2008). Since anthropogenic nutrient loading impacts marine ecosystems and decreases biodiversity, it eventually decreases also its ecosystem services (Halpern et al., 2008; Vörösmarty et al., 2010; Barbier et al., 2011). Periodical evaluation, especially of nutrient concentrations as well as plankton community composition, is vital to maintain biodiversity and ecosystem services of the studied area.

The Energy Transfer Efficiency Under Lowered pH and Nutrient Enrichment

The electron requirement for carbon fixation, $\Phi_{e,C}$, calculated as $\text{ETR} (P_m^B)^{-1}$ was in general higher (i.e., less efficient) in the Present as compared to the Future scenario, especially at noon, suggesting a better performance in the later scenario, with more carbon fixed per mole of electron transferred via PSII. This was more relevant at noon maybe because of the high light stress at this period of the day, widely reported along a long latitudinal range e.g., from tropical to polar environments (Rivkin and Putt, 1988; Levy et al., 2004; Aikawa et al., 2009). As a result of high light stress, the diurnal hysteresis—asymmetric response of photosynthetic entities to the similar irradiance levels in the morning vs. afternoon—was observed as many earlier studies (Villafañe et al., 2013; Helbling et al., 2015a). The $\Phi_{e,C}$ values were in all cases higher (ca. 6–12 $\text{mol e}^- (\text{mol C})^{-1}$) after 2 days acclimation (Figure 5D) than the theoretical values of 4–6 $\text{mol e}^- (\text{mol C})^{-1}$ (Genty et al., 1989; Lawrenz et al., 2013). This slight inefficiency suggests an extra use of electrons by other processes such as photorespiration (Badger et al., 2000), or alternative electron transport (Wagner et al., 2006; Suggett et al., 2009; Waring et al., 2010; Thamatrakoln et al., 2013). Energy transfer efficiency in our phytoplankton communities (Figures 5D, 8D) were within the range of earlier field and laboratory studies, namely from 1.2 to 54.3 $\text{mol e}^- (\text{mol C})^{-1}$, with a mean value of ca. 10 $\text{mol e}^- (\text{mol C})^{-1}$ (Suggett et al., 2009; Lawrenz et al., 2013). While the detailed mechanisms driving the variations of $\Phi_{e,C}$ are not well understood, we conclude that the less energy required by carbon concentrating mechanisms operating under higher nutrient availability coupled with acidification conditions, will result in a more efficient reductant energy flow from the primary photosynthetic processes to the biochemical carbon fixation. Since the variability of $\Phi_{e,C}$ occurs also on taxonomic level (Hughes et al., 2018), the combination of microscopic identification and molecular techniques, such as next generation sequencing, together with physiological investigation using freshly isolated monoclonal cultures from the ecosystems will help to understand the role of key players in energy transfer in various ecosystems.

Implications for Primary Production in the Future Patagonia

Our study clearly showed the different growth strategies of phytoplankton communities under increased nutrient availability and lowered pH in three different sampling sites of the Patagonian coastal waters (Argentina). The largest growth stimulation in the estuarine site indicate that future global change conditions will significantly influence environments

with high fluctuations and gradients in their abiotic variables (Cloern, 2001). Our results also demonstrate the importance of carbon-nutrient history of the phytoplankton community, namely that planktons in PU under frequent nutrient availability have advantage using enriched nutrient for cell growth. Finally, since we observed an important growth of the large diatoms that are in general tightly connected to larger zooplankton (Landry, 1977), we believe that future global conditions may also change the trophodynamics of certain areas by favoring large-sized species.

DATA AVAILABILITY STATEMENT

The original contributions presented in the study are included in the article/supplementary material, further inquiries can be directed to the corresponding author/s. The code used for the computation of C chemistry is written in Python 3 and freely available in Zenodo at <https://zenodo.org/record/4304878>.

AUTHOR CONTRIBUTIONS

TM, OP, VV, MV, and EWH designed the experiments, and performed measurements and analysis. TM, VV, EWH, and OP interpreted data. EWH prepared the figures. MV analyzed statistical data. KI computed carbon chemistry. TM and VV

prepared the original draft, which was revised by all the co-authors.

FUNDING

This work was supported by the Consejo Nacional de Investigaciones Científicas y Técnicas (CONICET)-The Czech Academy of Sciences (CAS) and Bilateral Mobility project #16-10 awarded to OP and EWH. Agencia Nacional de Promoción Científica y Tecnológica-ANPCyT (PICT 2015-0462) and Fundación Playa Unión.

ACKNOWLEDGMENTS

We thank the comments and suggestions of two reviewers that helped us to greatly improve the manuscript. KI was supported by the Simons Foundation (Simons Postdoctoral Fellowship in Marine Microbial Ecology, Award 544338). We thank P. Bermejo and R. Blum for nutrient analyses, and Dr. T. Shiozaki for satellite image analysis. The Cooperativa Eléctrica y de Servicios de Rawson provided the building's infrastructure to carry out this work. This is Contribution No. 188 of Estación de Fotobiología Playa Unión.

REFERENCES

- Aikawa, S., Hattori, H., Gomi, Y., Watanabe, K., Kudoh, S., Kashino, Y., et al. (2009). Diel tuning of photosynthetic systems in ice algae at Saroma-ko Lagoon, Hokkaido, Japan. *Polar Sci.* 3, 57–72. doi: 10.1016/j.polar.2008.12.001
- Álvarez-Colombo, G., Dato, C., Macchi, G., Palma, E., Machinandiarena, L., Christiansen, H., et al. (2011). Distribution and behavior of Argentine hake larvae: evidence of a biophysical mechanism for self-recruitment in northern Patagonian shelf waters. *Cienc. Mar.* 37, 633–657. doi: 10.7773/cm.v37i4b.1777
- Arvanitoyannis, I. S., and Kassaveti, A. (2008). Fish industry waste: treatments, environmental impacts, current and potential uses. *Int. J. Food Sci.* 43, 726–745. doi: 10.1111/j.1365-2621.2006.01513.x
- Bach, L. T., and Taucher, J. (2019). CO2 effects on diatoms: a synthesis of more than a decade of ocean acidification experiments with natural communities. *Ocean Sci. J.* 15, 1159–1175. doi: 10.5194/os-15-1159-2019
- Badger, M. R., Von Caemmerer, S., Ruuska, S., and Nakano, H. (2000). Electron flow to oxygen in higher plants and algae: rates and control of direct photoreduction (Mehler reaction) and rubisco oxygenase. *Philos. Trans. R. Soc. Lond. B. Biol. Sci.* 355, 1433–1446. doi: 10.1098/rstb.2000.0704
- Barbier, E. B., Hacker, S. D., Kennedy, C., Koch, E. W., Stier, A. C., and Silliman, B. R. (2011). The value of estuarine and coastal ecosystem services. *Ecol. Monogr.* 81, 169–193.
- Beardall, J., and Raven, J. A. (2004). The potential effects of global climate change on microalgal photosynthesis, growth and ecology. *Phycologia* 43, 26–40. doi: 10.2216/i0031-8884-43-1-26.1
- Behrenfeld, M. J., O'malley, R. T., Siegel, D. A., McClain, C. R., Sarmiento, J. L., Feldman, G. C., et al. (2006). Climate-driven trends in contemporary ocean productivity. *Nature* 444, 752–755. doi: 10.1038/nature05317
- Bermejo, P., Durán-Romero, C., Villafañe, V. E., and Helbling, E. W. (2020). Influence of fluctuating irradiance on photosynthesis, growth and community structure of estuarine phytoplankton under increased nutrients and acidification. *J. Exp. Mar. Biol. Ecol.* 526, 151348. doi: 10.1016/j.jembe.2020.151348
- Bermejo, P., Helbling, E. W., Durán-Romero, C., Cabrerizo, M. J., and Villafañe, V. E. (2018). Abiotic control of phytoplankton blooms in temperate coastal marine ecosystems: a case study in the South Atlantic Ocean. *Sci. Total Environ.* 612, 894–902. doi: 10.1016/j.scitotenv.2017.08.176
- Biswas, H., Cros, A., Yadav, K., Ramana, V. V., Prasad, V. R., Acharyya, T., et al. (2011). The response of a natural phytoplankton community from the Godavari River Estuary to increasing CO2 concentration during the pre-monsoon period. *J. Exp. Mar. Biol. Ecol.* 407, 284–293. doi: 10.1016/j.jembe.2011.06.027
- Björn, L. O., and Murphy, T. M. (1985). Computer calculation of solar ultraviolet radiation at ground level. *Physiol. Veg.* 23, 555–561.
- Boyd, P. W., Strzepak, R., Fu, F., and Hutchins, D. A. (2010). Environmental control of open-ocean phytoplankton groups: now and in the future. *Limnol. Oceanogr.* 55, 1353–1376. doi: 10.4319/lo.2010.55.3.1353
- Burkhardt, S., Amoroso, G., Riebesell, U., and Sültemeyer, D. (2001). CO2 and HCO3⁻ uptake in marine diatoms acclimated to different CO2 concentrations. *Limnol. Oceanogr.* 46, 1378–1391. doi: 10.4319/lo.2001.46.6.1378
- Cabrerizo, M. J., Carrillo, P., Villafañe, V. E., Medina-Sánchez, J. M., and Helbling, E. W. (2018). Increased nutrients from aeolian-dust and riverine origin decrease the CO2-sink capacity of coastal South Atlantic waters under UVR exposure. *Limnol. Oceanogr.* 63, 1191–1203. doi: 10.1002/lno.10764
- Carli, P. D., Braccalenti, J. C., García-De-León, F. J., and Gómez, E. P. A. (2012). La pesquería del langostino argentino *Pleoticus muelleri* en Patagonia Un único stock? *Anales. Inst. Patagonia.* 40, 103–112. doi: 10.4067/s0718-686x2012000200009
- Cloern, J. E. (1999). The relative importance of light and nutrient limitation of phytoplankton growth: a simple index of coastal ecosystem sensitivity to nutrient enrichment. *Aquatic. Ecology.* 33, 3–16.
- Cloern, J. E. (2001). Our evolving conceptual model of the coastal eutrophication problem. *Mar. Ecol. Prog. Ser.* 210, 223–253. doi: 10.3354/meps210223
- Cloern, J. E., Abreu, P. C., Carstensen, J., Chauvaud, L., Elmgren, R., Grall, J., et al. (2016). Human activities and climate variability drive fast-paced change across the world's estuarine-coastal ecosystems. *Glob. Chang. Biol.* 22, 513–529. doi: 10.1111/gcb.13059

- Crain, C. M., Kroeker, K., and Halpern, B. S. (2008). Interactive and cumulative effects of multiple human stressors in marine systems. *Ecol. Lett.* 11, 1304–1315. doi: 10.1111/j.1461-0248.2008.01253.x
- D'Agostino, V. C., Degradi, M., Santinelli, N., Sastre, V., Dans, S. L., and Hoffmeyer, M. S. (2018). The seasonal dynamics of plankton communities relative to the foraging of the southern right whale (*Eubalaena australis*) in northern Patagonian gulfs, Península Valdés, Argentina. *Continental Shelf Res.* 164, 45–57. doi: 10.1016/j.csr.2018.06.003
- Depetris, P. J., Gaiero, D. M., Probst, J. L., Hartmann, J., and Kempe, S. (2005). Biogeochemical output and typology of rivers draining Patagonia's Atlantic seaboard. *J. Coast. Res.* 214, 835–844. doi: 10.2112/015-nis.1
- DOE (1994). *Handbook of Methods for the Analysis of the Various Parameters of the Carbon Dioxide System in Sea Water, Version 2*, eds A. G. Dickson and C. Govet (Washington, DC: DOE).
- Domingues, R. B., Guerra, C. C., Galvão, H. M., Brotas, V., and Barbosa, A. B. (2016). Short-term interactive effects of ultraviolet radiation, carbon dioxide and nutrient enrichment on phytoplankton in a shallow coastal lagoon. *Aquat. Ecol.* 51, 91–105. doi: 10.1007/s10452-016-9601-4
- Duarte, C. M., Hendriks, I. E., Moore, T. S., Olsen, Y. S., Steckbauer, A., Ramajo, L., et al. (2013). Is ocean acidification an open-ocean syndrome? Understanding anthropogenic impacts on seawater pH. *Estuaries Coasts* 36, 221–236. doi: 10.1007/s12237-013-9594-3
- Durán-Romero, C., Villafañe, V. E., Valiñas, M. S., Gonçalves, R. J., Helbling, E. W., and Sathyendranath, S. (2017). Solar UVR sensitivity of phyto- and bacterioplankton communities from Patagonian coastal waters under increased nutrients and acidification. *ICES J. Mar. Sci.* 74, 1062–1073. doi: 10.1093/icesjms/fsw248
- Dutkiewicz, S., Morris, J. J., Follows, M. J., Scott, J., Levitan, O., Dyhrman, S. T., et al. (2015). Impact of ocean acidification on the structure of future phytoplankton communities. *Nat. Clim. Change* 5, 1002–1006. doi: 10.1038/nclimate2722
- Eggers, S. L., Lewandowska, A. M., Barcelos, E. R. J., Blanco-Ameijeiras, S., Gallo, F., and Matthiessen, B. (2014). Community composition has greater impact on the functioning of marine phytoplankton communities than ocean acidification. *Glob. Chang. Biol.* 20, 713–723. doi: 10.1111/gcb.12421
- Eilers, P. H. C., and Peeters, J. C. H. (1988). A model for the relationship between light-intensity and the rate of photosynthesis in phytoplankton. *Ecol. Model.* 42, 199–215. doi: 10.1016/0304-3800(88)90057-9
- Emerson, S. R., and Hedges, J. I. (2008). *Chemical Oceanography and the Marine Carbon Cycle*, 1 Edn. Cambridge: Cambridge University Press.
- Falkowski, P. G., Barber, R. T., and Smetacek, V. (1998). Biogeochemical controls and feedbacks on ocean primary production. *Science* 281, 200–206. doi: 10.1126/science.281.5374.200
- Fitch, D. T., and Moore, J. K. (2007). Wind speed influence on phytoplankton bloom dynamics in the Southern Ocean Marginal Ice Zone. *J. Geophys. Res.* 112:C08006. doi: 10.1029/2006jc004061
- Fu, F. X., Mullholand, M. R., Garcia, N. S., Beck, A., Bernhardt, P. W., Warner, M. E., et al. (2008). Interactions between changing pCO₂, N₂ fixation, and Fe limitation in the marine unicellular cyanobacterium *Crocosphaera*. *Limnol. Oceanogr.* 53, 2472–2484. doi: 10.4319/lo.2008.53.6.2472
- Fu, F. X., Place, A. R., Garcia, N. S., and Hutchins, D. A. (2010). CO₂ and phosphate availability control the toxicity of the harmful bloom dinoflagellate *Karlodinium veneficum*. *Aquat. Microb. Ecol.* 59, 55–65. doi: 10.3354/ame01396
- Gaiero, D. M., Probst, J. L., Depetris, P. J., Bidart, S. M., and Leleyter, L. (2003). Iron and other transition metals in Patagonian riverborne and windborne materials: geochemical control and transport to the southern South Atlantic Ocean. *Geochim. Cosmochim. Acta* 67, 3603–3623. doi: 10.1016/s0016-7037(03)00211-4
- Gao, K., Helbling, E. W., Häder, D. P., and Hutchins, D. A. (2012). Responses of marine primary producers to interactions between ocean acidification, solar radiation, and warming. *Mar. Ecol. Prog. Ser.* 470, 167–189. doi: 10.3354/meps10043
- Gassó, S., and Stein, A. F. (2007). Does dust from Patagonia reach the sub-Antarctic Atlantic Ocean? *Geophys. Res. Lett.* 34:L01801.
- Gattuso, J.-P., Gao, K., Lee, K., Rost, B., and Schulz, K. G. (2010). "Approaches and tools to manipulate the carbonate chemistry," in *Guide to Best Practices for Ocean Acidification Research and Data Reporting*, eds U. Riebesell, V. J. Fabry, L. Hansson, and J. P. Gattuso (Brussels: European Union), 41–52.
- Genty, B., Briantais, J.-M., and Baker, N. (1989). The relationship between the quantum yield of photosynthetic electron transport and quenching of chlorophyll fluorescence. *BBA Bioener.* 990, 87–92. doi: 10.1016/s0304-4165(89)80016-9
- Giordano, M., Beardall, J., and Raven, J. A. (2005). CO₂ concentrating mechanisms in algae: mechanisms, environmental modulation, and evolution. *Annu. Rev. Plant Biol.* 56, 99–131. doi: 10.1146/annurev.arplant.56.032604.144052
- Glibert, P. M. (2016). Margalef revisited: a new phytoplankton mandala incorporating twelve dimensions, including nutritional physiology. *Harmful Algae* 55, 25–30. doi: 10.1016/j.hal.2016.01.008
- Góngora, M. E., González Zevallos, D., Pettovello, A., and Mendía, L. (2012). Caracterización de las principales pesquerías del golfo San Jorge Patagonia. *Argentina. Lat. Am. J. Aquat. Res.* 40, 1–11. doi: 10.3856/vol40-issue1-fulltext-1
- Gregg, W. W., and Rousseaux, C. S. (2019). Global ocean primary production trends in the modern ocean color satellite record (1998–2015). *Environm. Res. Lett.* 14:124011. doi: 10.1088/1748-9326/ab4667
- Guillard, R. R. L., and Ryther, J. H. (1962). Studies of marine planktonic diatoms. I. *Cyclotella nana* Hustedt and *Detonula confervacea* Cleve. *Can. J. Microbiol.* 8, 229–239. doi: 10.1139/m62-029
- Halpern, B. S., Walbridge, S., Selkoe, K. A., Kappel, C. V., Micheli, F., D'agrosa, C., et al. (2008). A global map of human impact on marine ecosystems. *Science* 319, 948–952.
- Hama, T., Inoue, T., Suzuki, R., Kashiwazaki, H., Wada, S., Sasano, D., et al. (2016). Response of a phytoplankton community to nutrient addition under different CO₂ and pH conditions. *J. Oceanogr.* 72, 207–223. doi: 10.1007/s10872-015-0322-4
- Helbling, E. W., Banaszak, A. T., and Villafañe, V. E. (2015a). Differential responses of two phytoplankton communities from the chubut river estuary (Patagonia, Argentina) to the combination of UVR and elevated temperature. *Estuaries Coasts* 38, 1134–1146. doi: 10.1007/s12237-013-9752-7
- Helbling, E. W., Banaszak, A. T., and Villafañe, V. E. (2015b). Global change feed-back inhibits cyanobacterial photosynthesis. *Sci. Rep.* 5:14514.
- Helbling, E. W., Barbierilg, E. S., Gonc, M. A. M., Alves, R. J., and Villafañe, V. E. (2005). Impact of solar ultraviolet radiation on marine phytoplankton of Patagonia, Argentina. *Photochem. Photobiol.* 81, 807–818. doi: 10.1562/2005-03-02-ra-452r.1
- Helbling, E. W., Buma, A. G. J., Boelen, P., Van Der Strate, H. J., Giordano, M. V. F., and Villafañe, V. E. (2011). Increase in Rubisco activity and gene expression due to elevated temperature partially counteracts ultraviolet radiation-induced photoinhibition in the marine diatom *Thalassiosira weissflogii*. *Limnol. Oceanogr.* 56, 1330–1342. doi: 10.4319/lo.2011.56.4.1330
- Helbling, E. W., Pérez, D. E., Medina, C. D., Lagunas, M. G., and Villafañe, V. E. (2010). Phytoplankton distribution and photosynthesis dynamics in the Chubut River estuary (Patagonia, Argentina) throughout tidal cycles. *Limnol. Oceanogr.* 55, 55–65. doi: 10.4319/lo.2010.55.1.0055
- Helbling, E. W., Santamarina, J. M., and Villafañe, V. E. (1992). Chubut river estuary (Argentina): estuarine variability under different conditions of river discharge. *Rev. Biol. Mar. (Valparaíso)* 27, 73–90.
- Hillebrand, H., Dürselen, C.-D., Kirschtel, D., Pollinger, U., and Zohary, T. (1999). Biovolume calculation for pelagic and benthic microalgae. *J. Phycol.* 35, 403–424. doi: 10.1046/j.1529-8817.1999.3520403.x
- Holm-Hansen, O., and Helbling, E. (1995). "Técnicas para la medición de la productividad primaria en el fitoplancton," in *Manual de Métodos Ficológicos*, eds K. Alveal, M. E. Ferrario, E. C. Oliveira, and E. Sar (Concepción: Universidad de Concepción), 329–350.
- Holm-Hansen, O., Lorenzen, C. J., Holmes, R. W., and Strickland, J. D. H. (1965). Fluorometric determination of chlorophyll. *J. Cons. Perm. Int. Explor. Mar.* 30, 3–15. doi: 10.1093/icesjms/30.1.3
- Holm-Hansen, O., and Riemann, B. (1978). Chlorophyll *a* determination: improvements in methodology. *Oikos* 30, 438–447. doi: 10.2307/3543338
- Hopkinson, B. M., Dupont, C. L., Allen, A. E., and Morel, F. M. (2011). Efficiency of the CO₂-concentrating mechanism of diatoms. *Proc. Natl. Acad. Sci. U.S.A.* 108, 3830–3837. doi: 10.1073/pnas.1018062108
- Howarth, R. W. (1988). Nutrient limitation of net primary production in marine ecosystems. *Ann. Rev. Ecol. Evol.* 19, 89–110. doi: 10.1146/annurev.es.19.110188.000513

- Hughes, D. J., Campbell, D. A., Doblin, M. A., Kromkamp, J. C., Lawrenz, E., Moore, C. M., et al. (2018). Roadmaps and detours: Active chlorophyll-*a* assessments of primary productivity across marine and freshwater systems. *Environ. Sci. Technol.* 52, 12039–12054. doi: 10.1021/acs.est.8b03488
- Hutchins, D. A., Fu, F.-X., Zhang, Y., Warner, M. E., Feng, Y., Portune, K., et al. (2007). CO₂ control of *Trichodesmium* N₂ fixation, photosynthesis, growth rates, and elemental ratios: Implications for past, present, and future ocean biogeochemistry. *Limnol. Oceanogr.* 52, 1293–1304. doi: 10.4319/lo.2007.52.4.1293
- IPCC (2019). *IPCC Special Report on the Ocean and Cryosphere in a Changing Climate*, eds H.-O. Pörtner, D. C. Roberts, V. Masson-Delmotte, P. Zhai, M. Tignor, E. Poloczanska, et al. (Geneva: IPCC).
- Kahl, L. (2018). *Dinámica del CO₂ en el Océano Atlántico Sudoccidental*. Ph.D. thesis, Universidad de Buenos Aires, Buenos Aires.
- Kennish, M. J. (1992). *Ecology of Estuaries: Anthropogenic Effects*. Boca Raton, FL: CRC Press.
- Kolber, Z. S., Prášil, O., and Falkowski, P. G. (1998). Measurements of variable chlorophyll fluorescence using fast repetition rate techniques: defining methodology and experimental protocols. *Biochim. Biophys. Acta-Bioenerg.* 1367, 88–106. doi: 10.1016/s0005-2728(98)00135-2
- Landry, M. R. (1977). A review of important concept in the trophic organization of pelagic ecosystems. *Helgol. Meeres.* 30, 8–17. doi: 10.1007/bf02207821
- Lawrenz, E., Silsbe, G., Capuzzo, E., Ylostalo, P., Forster, R. M., Simis, S. G., et al. (2013). Predicting the electron requirement for carbon fixation in seas and oceans. *PLoS One* 8:e58137. doi: 10.1371/journal.pone.0058137
- Levitán, O., Rosenberg, G., Setlik, I., Setlikova, E., Grigel, J., Klepetar, J., et al. (2007). Elevated CO₂ enhances nitrogen fixation and growth in the marine cyanobacterium *Trichodesmium*. *Glob. Chang. Biol.* 13, 531–538. doi: 10.1111/j.1365-2486.2006.01314.x
- Levy, O., Dubinsky, Z., Schneider, K., Achituv, Y., Zakai, D., and Gorbunov, M. Y. (2004). Diurnal hysteresis in coral photosynthesis. *Mar. Ecol. Prog. Ser.* 268, 105–117. doi: 10.3354/meps268105
- Litchman, E. (2007). “Resource competition and the ecological success of phytoplankton,” in *Evolution of Primary Producers in the Sea*, eds P. G. Falkowski and A. H. Knoll (Cambridge, MA: Academic Press), 351–375. doi: 10.1016/b978-012370518-1/50017-5
- Longhurst, A., Sathyendranath, S., Platt, T., and Caverhill, C. (1995). An estimate of global primary production in the ocean from satellite radiometer data. *J. Plankton Res.* 17, 1245–1271. doi: 10.1093/plankt/17.6.1245
- Lueker, T. J., Dickson, A. G., and Keeling, C. D. (2000). Ocean pCO₂ calculated from dissolved inorganic carbon, alkalinity, and equations for K₁ and K₂: validation based on laboratory measurements of CO₂ in gas and seawater at equilibrium. *Mar. Chem.* 70, 105–119. doi: 10.1016/S0304-4203(00)00022-0
- Lutz, V., Segura, V., Dogliotti, A., Tavano, V., Brandini, F. P., Calliari, D. L., et al. (2018). “Overview of primary production in the Southwestern Atlantic,” in *Plankton ecology of the Southwestern Atlantic – From the Subtropical to the Subantarctic Realm*, eds M. S. Hoffmeyer, M. E. Sabatini, F. P. Brandini, D. L. Calliari, and N. H. Santinelli (Heidelberg: Springer), 101–126. doi: 10.1007/978-3-319-77869-3_6
- Marcova, M. A., Villafane, V. E., and Helbling, E. W. (2008). Combined effects of solar ultraviolet radiation and nutrients addition on growth, biomass and taxonomic composition of coastal marine phytoplankton communities of Patagonia. *J. Photochem. Photobiol. B* 91, 157–166. doi: 10.1016/j.jphotobiol.2008.03.002
- Margalef, R. (1978). Life-forms of phytoplankton as survival alternatives in an unstable environment. *Oceanol. Acta* 1, 493–509.
- Matsuda, Y., and Kroth, P. G. (2014). “Carbon fixation in diatoms,” in *The Structural Basis of Biological Energy Generation*, ed. M. F. Hohmann-Marriott (Heidelberg: Springer), 335–362. doi: 10.1007/978-94-017-8742-0_18
- Morris, M. D. (1991). Factorial sampling plans for preliminary computational experiments. *Technometrics* 33, 161–174. doi: 10.1080/00401706.1991.10484804
- Neale, P. J., Helbling, E. W., and Zagarese, H. E. (2003). “Modulation of UVR exposure and effects by vertical mixing and advection,” in *UV Effects in Aquatic Organisms and Ecosystems*, eds E. W. Helbling and H. E. Zagarese (Cambridge: Royal Society of Chemistry), 108–134.
- Nielsen, L. T., Hallegraeff, G. M., Wright, S. W., and Hansen, P. J. (2012). Effects of experimental seawater acidification on an estuarine plankton community. *Aquat. Microb. Ecol.* 65, 271–286. doi: 10.3354/ame01554
- Orselli, I. B. M., Kerr, R., Ito, R. G., Tavano, V. M., Mendes, C. R. B., and Garcia, C. A. E. (2018). How fast is the Patagonian shelf-break acidifying? *J. Marine Syst.* 178, 1–14. doi: 10.1016/j.jmarsys.2017.10.007
- Osma, N., Latorre-Melín, L., Jacob, B., Contreras, P. Y., Von Dassow, P., and Vargas, C. A. (2020). Response of phytoplankton assemblages from naturally acidic coastal ecosystems to elevated pCO₂. *Front. Mar. Sci.* 7:323. doi: 10.3389/fmars.2020.00323
- Oxborough, K., Moore, C. M., Suggett, D. J., Lawson, T., Chan, H. G., and Geider, R. J. (2012). Direct estimation of functional PSII reaction center concentration and PSII electron flux on a volume basis: a new approach to the analysis of Fast Repetition Rate Fluorometry (FRRF) data. *Limnol. Oceanogr.-Meth.* 10, 142–154. doi: 10.4319/lom.2012.10.142
- Paerl, H. W., Valdes, L. M., Peierls, B. L., Adolf, J. E., Lawrence, W., and Harding, J. (2006). Anthropogenic and climatic influences on the eutrophication of large estuarine ecosystems. *Limnol. Oceanogr.* 51, 448–462. doi: 10.4319/lo.2006.51.1_part_2.0448
- Piccolo, M. C., and Perillo, G. M. E. (1999). “Estuaries of Argentina: a review,” in *Estuaries of South America: Their Geomorphology and Dynamics*, eds G. M. E. Perillo, M. C. Piccolo, and M. P. Quivira (Berlin: Springer-Verlag), 101–132. doi: 10.1007/978-3-642-60131-6_6
- Porra, R. J. (2002). The chequered history of the development and use of simultaneous equations for the accurate determination of chlorophylls *a* and *b*. *Photosynth. Res.* 73, 149–156.
- Quinn, G. P., and Keough, M. J. (2002). *Experimental Design and Data Analysis for Biologists*. Cambridge: Cambridge University Press.
- Rabalais, N. N., Turner, R. E., Diaz, R. J., and Justić, D. (2009). Global change and eutrophication of coastal waters. *ICES J. Mar. Sci.* 66, 1528–1537. doi: 10.1093/icesjms/fsp047
- Reinfelder, J. R. (2011). Carbon concentrating mechanisms in eukaryotic marine phytoplankton. *Ann. Rev. Mar. Sci.* 3, 291–315. doi: 10.1146/annurev-marine-120709-142720
- Rivas, A. L., and Beier, E. J. (1990). Temperature and salinity fields in the Northpatagonian Gulfs. *Oceanol. Acta* 13, 15–20.
- Rivkin, R. B., and Putt, M. (1988). Seasonal pattern of diel periodicity in photosynthesis by polar phytoplankton: species-specific responses. *J. Phycol.* 24, 369–376. doi: 10.1111/j.1529-8817.1988.tb00186.x
- Rost, B., Zondervan, I., and Wolf-Gladrow, D. (2008). Sensitivity of phytoplankton to future changes in ocean carbonate chemistry: current knowledge, contradictions and research directions. *Mar. Ecol. Prog. Ser.* 373, 227–237. doi: 10.3354/meps07776
- Rousseaux, C., and Gregg, W. (2014). Interannual variation in phytoplankton primary production at a global scale. *Remote Sens.* 6, 1–19. doi: 10.3390/rs6010001
- Ruggaber, A., Dlugi, R., and Nakajima, T. (1994). Modelling of radiation quantities and photolysis frequencies in the troposphere. *J. Atmos. Chem.* 18, 171–210. doi: 10.1007/bf00696813
- Smayda, T. J., and Reynolds, C. S. (2001). Community assembly in marine phytoplankton: application of recent models to harmful dinoflagellate blooms. *J. Plankton Res.* 23, 447–461. doi: 10.1093/plankt/23.5.447
- Strathmann, R. R. (1967). Estimating the organic carbon content of phytoplankton from cell volume or plasma volume. *Limnol. Oceanogr.* 12, 411–418. doi: 10.4319/lo.1967.12.3.0411
- Strickland, J. D. H., and Parsons, T. R. (1972). A practical handbook of seawater analysis. *J. Fish Res. Board Can.* 167, 1–310.
- Suggett, D. J., Moore, C. M., Hickman, A. E., and Geider, R. J. (2009). Interpretation of fast repetition rate (FRR) fluorescence: signatures of phytoplankton community structure versus physiological state. *Mar. Ecol. Prog. Ser.* 376, 1–19. doi: 10.3354/meps07830
- Sun, J., Hutchins, D. A., Feng, Y., Seubert, E. L., Caron, D. A., and Fu, F.-X. (2011). Effects of changing CO₂ and phosphate availability on domoic acid production and physiology of the marine harmful bloom diatom *Pseudo-nitzschia multiseries*. *Limnol. Oceanogr.* 56, 829–840. doi: 10.4319/lo.2011.56.3.0829

- Tatters, A. O., Fu, F. X., and Hutchins, D. A. (2012). High CO₂ and silicate limitation synergistically increase the toxicity of *Pseudo-nitzschia fraudulenta*. *PLoS One* 7:e32116. doi: 10.1371/journal.pone.0032116
- Taucher, J., Jones, J., James, A., Brzezinski, M. A., Carlson, C. A., Riebesell, U., et al. (2015). Combined effects of CO₂ and temperature on carbon uptake and partitioning by the marine diatoms *Thalassiosira weissflogii* and *Dactyliosolen fragilissimus*. *Limnol. Oceanogr.* 60, 901–919. doi: 10.1002/lno.10063
- Thamatrakoln, K., Bailleul, B., Brown, C. M., Gorbunov, M. Y., Kustka, A. B., Frada, M., et al. (2013). Death-specific protein in a marine diatom regulates photosynthetic responses to iron and light availability. *Proc. Natl. Acad. Sci. U.S.A.* 110, 20123–20128. doi: 10.1073/pnas.1304727110
- Thomsen, H. A., Buck, K. R., and Chavez, F. P. (1994). “Haptophytes as components of marine phytoplankton,” in *Haptophyte Algae*, eds J. C. Green and B. S. C. Leadbeater (Oxford: Clarendon Press), 187–208.
- Trimborn, S., Lundholm, N., Thoms, S., Richter, K. U., Krock, B., Hansen, P. J., et al. (2008). Inorganic carbon acquisition in potentially toxic and non-toxic diatoms: the effect of pH-induced changes in seawater carbonate chemistry. *Physiol. Plant.* 133, 92–105. doi: 10.1111/j.1399-3054.2007.01038.x
- Uitz, J., Claustre, H., Gentili, B., and Stramski, D. (2010). Phytoplankton class-specific primary production in the world's oceans: Seasonal and interannual variability from satellite observations. *Global Biogeochem. Cy* 24:GB3016.
- Underwood, A. J. (1997). *Experiments in Ecology: Their Logical Design and Interpretation Using Analysis of Variance*. Cambridge: Cambridge University Press.
- Villafañe, V. E. (2004). Annual patterns of ultraviolet radiation effects on temperate marine phytoplankton off Patagonia. *Argentina. J. Plankton Res.* 26, 167–174. doi: 10.1093/plankt/fbh011
- Villafañe, V. E., Banaszak, A. T., Guendulain-García, S. D., Strauch, S. M., Halac, S. R., and Helbling, E. W. (2013). Influence of seasonal variables associated with climate change on photochemical diurnal cycles of marine phytoplankton from Patagonia (Argentina). *Limnol. Oceanogr.* 58, 203–214. doi: 10.4319/lno.2013.58.1.0203
- Villafañe, V. E., Cabrerizo, M. J., Carrillo, P., Hernando, M. P., Medina-Sánchez, J. M., Narvarte, M. A. (2021). “Global change effects on plankton from Atlantic Patagonian coastal waters: role of interacting drivers,” in *Global Change in Atlantic Coastal Patagonian Ecosystems: A Journey Through Time*, eds E. W. Helbling, M. A. Narvarte, R. A. González, and V. E. Villafañe (Heidelberg: Springer).
- Villafañe, V. E., Marcoval, M. A., and Helbling, E. W. (2004). Photosynthesis versus irradiance characteristics in phytoplankton assemblages off Patagonia (Argentina): temporal variability and solar UVR effects. *Mar. Ecol. Prog. Ser.* 284, 23–34.
- Villafañe, V. E., Paczkowska, J., Andersson, A., Duran Romero, C., Valinas, M. S., and Helbling, E. W. (2018). Dual role of DOM in a scenario of global change on photosynthesis and structure of coastal phytoplankton from the South Atlantic Ocean. *Sci. Total Environ.* 634, 1352–1361.
- Villafañe, V. E., and Reid, F. M. H. (1995). “Métodos de microscopía para la cuantificación del fitoplancton,” in *Manual de Métodos Ficológicos*, eds K. Alveal, M. E. Ferrario, E. C. Oliveira, and E. Sar (Concepcion: Universidad de Concepcion), 169–185.
- Villafañe, V. E., Valiñas, M. S., Cabrerizo, M. J., and Helbling, E. W. (2015). Physiological responses of Patagonian coastal marine phytoplankton in a scenario of global change: role of acidification, nutrients and solar UVR. *Mar. Chem.* 177, 411–420.
- Vizzo, J. I., Cabrerizo, M. J., Helbling, E. W., and Villafañe, V. E. (2021). Extreme and gradual rainfall effects on winter and summer estuarine phytoplankton communities from Patagonia (Argentina). *Mar. Environ. Res.* 163:105235.
- Vörösmarty, C. J., McIntyre, P. B., Gessner, M. O., Dudgeon, D., Prusevich, A., Green, P., et al. (2010). Global threats to human water security and river biodiversity. *Nature* 467, 555–561.
- Wagner, T., Guieu, C., Losno, R., Bonnet, S., and Mahowald, N. (2008). Revisiting atmospheric dust export to the Southern Hemisphere ocean: biogeochemical implications. *Global Biogeochem. Cy* 22, GB2006.
- Wagner, H., Jakob, T., and Wilhelm, C. (2006). Balancing the energy flow from captured light to biomass under fluctuating light conditions. *New Phytol.* 169, 95–108.
- Waring, J., Klenell, M., Bechtold, U., Underwood, G. J. C., and Baker, N. R. (2010). Light-induced responses of oxygen photoreduction, reactive oxygen species production and scavenging in two diatom species. *J. Phycol.* 46, 1206–1217.
- Weiss, R. F. (1974). Carbon dioxide in water and seawater: the solubility of a non-ideal gas. *Mar. Chem.* 4, 203–215.

Conflict of Interest: The authors declare that the research was conducted in the absence of any commercial or financial relationships that could be construed as a potential conflict of interest.

Copyright © 2021 Masuda, Prášil, Villafañe, Valiñas, Inomura and Helbling. This is an open-access article distributed under the terms of the Creative Commons Attribution License (CC BY). The use, distribution or reproduction in other forums is permitted, provided the original author(s) and the copyright owner(s) are credited and that the original publication in this journal is cited, in accordance with accepted academic practice. No use, distribution or reproduction is permitted which does not comply with these terms.