



Arctic kelp forests

Diversity, resilience and future

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1	ARCTIC KELP	FORESTS:	DIVERSITY,	RESILIENCE	AND FUTURE.
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ABSTRACT. The Arctic is one of the most rapidly changing places on Earth and it is a sentinel 14 region for understanding the range and magnitude of planetary changes, and their impacts on 15 ecosystems. However, our understanding of arctic coastal ecosystems remains limited, and the 16 impacts of ongoing and future climate change on them are largely unexplored. Kelp forests are 17 the dominant habitat along many rocky Arctic coastlines, providing structure and food for 18 economically and ecologically important species. Here we synthesize existing information on 19 20 the distribution and diversity of arctic kelp forests and assess how ongoing changes in environmental conditions could impact the extent, productivity, and resilience of these 21 22 important ecosystems. We identify regions where the range and growth of arctic kelp are likely to undergo rapid short-term increase due to reduced sea ice cover, increased light, and warming. 23 However, we also describe areas where kelp could be negatively impacted by rising freshwater 24 25 input and coastal erosion due to receding sea ice and melting permafrost. In some regions, arctic kelp forests have undergone sudden regime shifts due to altered ecological interactions 26 or changing environmental conditions. Key knowledge gaps for arctic kelp forests include 27 28 measures of extent and diversity of kelp communities (especially northern Canada and northeastern Russia), the faunal communities supported by many of these habitats, and the role 29 of arctic kelp forests in structuring nearby pelagic and benthic food webs. Filling in these gaps 30 and strategically prioritizing research in areas of rapid environmental change will enable more 31 32 effective management of these important habitats, and better predictions of future changes in 33 the coastal ecosystems they support and the services that they provide.

35 Keywords (6): seaweed, climate change, polar, sea ice loss, borealization

36 1.1. INTRODUCTION

The effects of humans are pervasive and are transforming natural ecosystems and 37 biogeochemical cycles on global scales (Halpern et al. 2008; Waters et al. 2016). There is, 38 however, great regional variation in the nature, magnitude, and direction of these changes 39 (Burrows et al. 2011; Krumhansl et al. 2016), and it is only by understanding these 40 geographical intricacies that we can begin to grasp the full extent of our footprint on the planet. 41 Currently, the Arctic is warming 2 - 4 times faster than the global average and is now one of 42 the most rapidly changing regions in the world (IPCC 2014). Marine ecosystems along Arctic 43 coasts are experiencing increases in sea temperatures, dramatic declines in sea ice, and 44 increased input of freshwater (Wassmann and Reigstad 2011; Coupel et al. 2015; Acosta 45 Navarro et al. 2016; Ding et al. 2017). These changes are altering carbon cycling, affecting the 46 timing and magnitude of primary production, and driving shifts in the structure and function of 47 marine communities (Grebmeier et al. 2006; Nelson et al. 2014). As a result, the entire Arctic 48 region has been designated an ocean warming hotspot (Hobday and Pecl 2014). Impacts of 49 rapid environmental change on arctic ecosystems has broad significance due to both the global 50 uniqueness and large geographic extent of the region, and because it may act as a sentinel for 51 other ecosystems experiencing slower rates of change (Pecl et al. 2014; Hobday and Pecl 2014). 52 Despite this, most Arctic coasts remain relatively unexplored, and the extent and resilience of 53 coastal ecosystems are poorly understood, as are the ongoing and future impacts of climate 54 55 change on them. Understanding changes to arctic ecosystems is especially critical because borealization (i.e., the northward shift of temperate communities) could squeeze out high arctic 56 ecosystems altogether, resulting in the planetary loss of an entire climate zone (Fossheim et al. 57 2015; Kortsch et al. 2015). 58

Kelp are large brown seaweeds that occur on rocky coasts throughout the Arctic(Wernberg et al. 2018). Many (or most) kelps are important foundation species that create

habitat (forests) for numerous fish and invertebrates (Christie et al. 2009; Norderhaug and
Christie 2011; Teagle et al. 2017), provide food to marine communities through high
production and export of detritus and dissolved organic material (Krumhansl and Scheibling
2012; Renaud et al. 2015; Abdullah et al. 2017; Filbee-Dexter et al. 2018 in press), and store
and sequester carbon (Krause-Jensen and Duarte 2016). Currently, information on the
distribution, diversity, stability, and function of kelp forests is missing for large portions of the
Arctic (Wiencke and Clayton 2009; Krumhansl et al. 2016; Wilce 2016).

A recent global analysis of records of kelp abundance over the past 5 decades showed 68 69 that kelp forests are changing in many regions of the world (Krumhansl et al. 2016). At the warmest edges of their range, sudden shifts from kelp forests to reefs dominated by low-lying 70 turf-forming algae have been increasingly documented over the last decade (Filbee-Dexter and 71 Wernberg 2018). Along other temperate coasts, native kelps are being replaced by invasive 72 kelps or other seaweeds (Wernberg et al. 2018), or are being heavily overgrazed by sea urchins 73 (Filbee-Dexter and Scheibling 2014). In many of these regions, declines in kelp abundance are 74 partly explained by the direct and indirect effects of warming sea temperatures (Ling et al. 75 2009; Catton 2016; Filbee-Dexter et al. 2016; Wernberg et al. 2016). Considering the 76 widespread changes throughout the temperate and tropical range of kelp and the ongoing 77 environmental changes occurring in the Arctic, the fate of arctic kelps in this era of rapid 78 79 change is a critical gap in our knowledge of arctic marine ecosystems.

Here we synthesize existing information on the distribution, biomass, and dominant species of arctic kelp forests. We explore some of the services provided by arctic kelps and identify missing baseline measures of their extent. We analyze changes in the sea ice extent and temperature conditions for known locations of kelp, and explore how recent and future changes in these and other conditions could impact their growth, reproduction, and survival.

Finally, we highlight key gaps in our understanding of these ecosystems, and suggest strategiesfor future research.

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88 1.2. HIDDEN BLUE FORESTS OF THE ARCTIC

89 1.2.1. Bounds of arctic marine ecosystems

Arctic and temperate marine ecosystems are separated by a moving boundary, generally 90 91 defined by latitude, sea ice cover, light variability, and the locations of the polar front and other ocean currents (Piepenburg 2005). The locations of these boundaries can be seasonal, 92 93 unpredictable, and can shift with climate change. A precise and universally accepted geographical definition of 'Arctic marine ecosystems' therefore does not exist, and different 94 southern limits for arctic marine ecosystems are used in the literature (Zenkevitch 1963; 95 Piepenburg 2005; Gattuso et al. 2006; Wilce 2016). For example, so called 'Arctic conditions' 96 (ice scoured intertidal zones, ocean temperatures $< 0^{\circ}$ C, and months with little to no daylight) 97 extend below the Arctic circle along the coasts of Greenland and Eastern Canada, which are 98 influenced by the cold southward moving Labrador and Greenland currents, but are restricted 99 to above the Arctic circle along the coasts of northern Norway, Iceland and in the southern 100 Bering sea, which are influenced by the warmer northward moving Gulf Stream and North 101 Pacific currents, respectively (Wilce 2016). The convergence of cool waters from the Arctic 102 Ocean and warm waters from the Atlantic and Pacific Oceans occurs around 65°N on the east 103 104 coast of Greenland, 80°N west of Svalbard, 76°C in the Barents Sea, in the Bering Strait, 63°N in the eastern Canadian Arctic Archipelago, and then slightly north between Baffin Island and 105 the west coast of Greenland (AMAP 1998). However, other factors such as sea ice, light, and 106 glacial run-off also create Arctic conditions south of these limits (AMAP 1998). Here we define 107 'arctic kelps' as kelps occurring within the boundaries defined by the Arctic Monitoring and 108 Assessment Program (AMAP). AMAP originally defined Arctic boundaries in 1991 as regions 109

north of the 10°C July isotherm. These boundaries have since been expanded to include some 110 areas that correspond to political boundaries of member nations of the Arctic Council (e.g., 111 coastal shelf of Iceland, Norwegian northwest coast, Hudson Bay, and the Aleutian Islands) 112 (AMAP 2017). We used this definition because monitoring programs, assessments and 113 decision-making on pollution and climate change in Arctic regions often use AMAP 114 boundaries. However, despite our inclusive definition of the Arctic, much of this manuscript 115 116 focuses on kelp forests at higher latitudes within the AMAP region where kelps face the most extreme Arctic conditions and where globally unique species compositions are found. 117

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119 1.2.2. Distribution, growth forms and evolution of arctic kelps

Although kelps range along most Arctic coasts, sparse records of kelp in some parts of the 120 Arctic have been attributed to a lack of hard substrata (Kjellman 1883; Wilce 2016). Only about 121 35% of the Arctic basin is rocky substrate and shallow coastal areas and inner Arctic fjords are 122 often dominated by sediment due to glacial run off and river deposition (Leont'yev 2003; 123 Lantuit et al. 2012), which limits the presence of kelp. In areas with suitable substrate, dense 124 kelp forests can extend from the intertidal zone down to depths of 30-40 m depending on light 125 conditions, wave regime, and grazing intensity (Wernberg et al. 2018). The deepest recorded 126 kelp was observed at 60 m depth in Disko Bay, Greenland (Boertmann et al. 2013). In high 127 Arctic regions, available light and sea ice further restrict this depth range and the upper 128 129 sublittoral zone is a barren, low salinity environment that is constantly impacted by sea ice and meltwater (Wiencke and Clayton 2011). 130

The diversity of kelp in the high Arctic tends to be lower than in temperate kelp forests (Wiencke and Clayton 2011). Genetic evidence indicates that most kelps reinvaded the Arctic from the Atlantic Ocean ~8,000 years ago following the last ice age, which eliminated benthic flora from most current Arctic subtidal regions (Wulff et al. 2011). As a result, most arctic

kelps have optimal growth temperatures that exceed those experienced during the Arctic 135 summer and many of these species therefore also thrive along warmer, temperate coasts 136 (Wiencke and Amsler 2012). In the high Arctic especially, kelps tend to be morphologically 137 smaller compared to their southern range limits (e.g., Kuznetsov et al. 1994; Kuznetsov and 138 Shoshina 2003; but see Borum et al. 2002). However, kelps still form dense canopies in some 139 regions (e.g., western Alaska and northern Norway) and provide most of the algal biomass and 140 141 the largest three-dimensional biogenic structure on rocky coasts in Arctic regions (Wiencke and Amsler 2012). In fact, these lush underwater forests are particularly striking in the Arctic, 142 143 where terrestrial coasts are barren and ice scoured with little three-dimensional structure.



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Fig 1. Photographs of select kelps from high Arctic regions: a) *Laminaria solidungula*, b) *Alaria elliptica*, c) *Saccharina longicruris*, d) *Saccharina nigripes*, and e) *Saccorhiza dermatodea* (Guiry and Guiry 2017).

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The species pool is relatively young, with only one truly arctic endemic kelp, *Laminaria solidungula* (Kjellman 1883; Zenkevitch 1963; Wilce and Dunton 2014). All other kelp species
found in Arctic regions also extend into sub-arctic and northern temperate waters and include *Alaria esculenta, Agarum clathratum, Eualaria fistulosa, Laminaria digitata, Laminaria*

hyperborea, Nereocystis luetkeana, Saccharina latissima, Saccharina longicruris, Saccharina 153 nigripes, Saccorhiza dermatodea, Alaria elliptica, and Alaria oblonga (the latter 2 are only 154 found in Russia) (Fig 1, Table 1). There is currently taxonomic confusion regarding some arctic 155 species; S. nigripes, for example, has often been misidentified as L. digitata, and appears to be 156 restricted to Arctic or subarctic conditions, although more information on its distribution is 157 needed (McDevit and Saunders 2010). In 2006 a new species of kelp Aureophycus aleuticus 158 159 was collected from Kagamil Island, Aleutian Islands, but its classification within the order Laminariales is still unclear (Kawai et al. 2013). New DNA barcoding techniques show 160 161 promise for clearing up misidentifications caused by diverse growth morphologies of kelps in arctic conditions (McDevit and Saunders 2010; Bringloe et al. 2017). 162

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164 1.2.1. Adaptations to Arctic conditions

Kelps in arctic environments are challenged by extremely low water temperatures, periods of 165 low salinity, and extreme variability in light caused by large annual variations in day length, 166 light intensity, and sea ice cover. In their northernmost range, kelps live in temperatures at the 167 point of freezing sea water during polar nights (e.g., NE Greenland, Borum et al. 2002; Franz 168 Joseph Land, Shoshina et al. 2016). Day-length ranges from 24-hour sunlight in mid-summer 169 to several months of total darkness during winter (Hanelt 1998). The low angle of the sun and 170 periods of complete darkness mean that high Arctic areas only receive 30 - 40 % of the light 171 received in the tropics on an annual basis. The long period of darkness during winter is further 172 extended in areas with partial or complete sea ice cover, especially if the ice is thick or covered 173 by snow (Mundy et al. 2007). Subtidal habitats in the Arctic can therefore be without light for 174 much of the year. Studies from NE Greenland illustrates this; the annual surface irradiance 175 (PAR) in Young Sound (74° 18' N) amounts to ca. 6100 mol photons m⁻², but the ice-free 176

period is limited to August and September so that the amount of available light at 10 and 20 m depth is only 234 and 40 mol photons m^{-2} yr⁻¹, respectively (Borum et al. 2002).

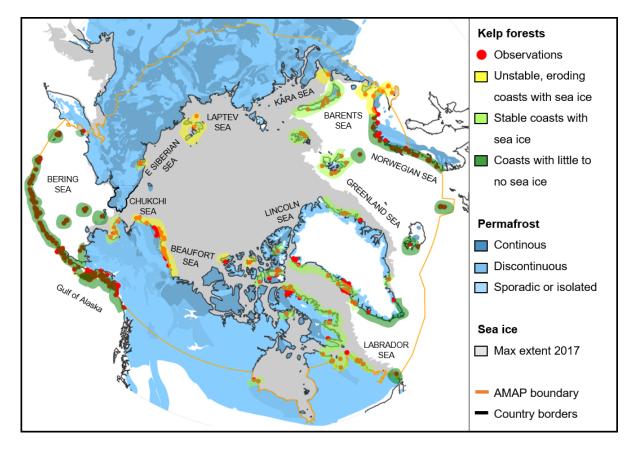
The marked seasonal variation in light availability in the Arctic concentrates primary 179 production into a short period and creates strong seasonality in the growth of kelp (Chapman 180 and Lindley 1980; Dunton and Jodwalis 1988; Borum et al. 2002; Makarov et al. 2008). Arctic 181 kelps are well adapted to these long periods of darkness or low light conditions. Studies on S. 182 183 latissima and L. solidungula show that these species store most of the carbon obtained during the short summer period and subsequently use these reserves to form new blades during the 184 185 succeeding period of almost darkness (Chapman and Lindley 1980; Dunton and Jodwalis 1988; Borum et al. 2002). Remarkably, the peak growth period for Alaskan L. solidungula was from 186 February to April under full ice cover (Dunton 1985), and the production of new lamina in S. 187 latissima from Young Sound (NE Greenland) occurred under ice cover and in complete 188 darkness, likely based on re-allocation of C from the old lamina or stipe (Borum et al. 2002). 189

Many kelp species can also cope with multi-year sea ice, which can cause severe 190 mechanical damage to benthic organisms in the intertidal and upper subtidal zone (Krause-191 Jensen et al. 2012; Dayton 2013; Shoshina et al. 2016). Most kelp forests recover from sea ice 192 damage through high reproduction and recolonization of the scoured substrate. Keats et al. 193 (1985) found, for example, that populations of A. esculenta recovered within a few years after 194 having been removed by ice-scour in the uppermost reaches of its range. However, Konar 195 196 (2013) found slow recolonization in clearing experiments on kelps in the Boulder Patch (< 10 % recolonization after 7 years), which is much slower than rates in many temperate kelp forests. 197 198

199 1.3. KNOWN LOCATIONS OF ARCTIC KELPS

Data on the current extent and distribution of kelps in the Arctic is not available. To overviewthe observational data record of kelp in subarctic and Arctic seas we compiled records of kelp

over the last 2 centuries, within the AMAP boundaries, from primary literature, museum 202 collections, dive logs, Arctic expeditions, coastal monitoring, and local ecological knowledge 203 from Inuit and northern communities (N = 1179 records, Fig 3). The spatial extent of these 204 ecosystems ranged from 100s of km² of kelp forests to small patches of kelp within inner fjords 205 and boulder patches along sedimentary coasts. The number of kelp records decreased with 206 latitude, with the northernmost observations of kelp forests > 80° N at Svalbard, Norway and 207 208 Franz Joseph Land, Russia (Shoshina et al. 1997; Bartsch et al. 2016). Most records were from northern Norway, western Greenland, eastern Canada, and northwestern USA. The earliest 209 210 records of arctic kelp were from the Canadian high Arctic during expeditions in search of the Northwest passage (Lee 1980). Other early records come from Kjellman (1883), who published 211 the first comprehensive review of polar benthic algae based on expeditions from Sweden via 212 Norway to Novaya Zemlya, and into the Siberian sea, Russia, and Rosenvinge (1893, 1899), 213 who described the algal flora in Greenland a decade later. Dive research on arctic kelp forests 214 was first conducted in Greenland, Canada and USA by Wilce (1963), Chapman and Lindley 215 (1980), and Dunton et al. (1982). It is worth noting that these historical records represent a 216 baseline and may not reflect current kelp distributions. 217

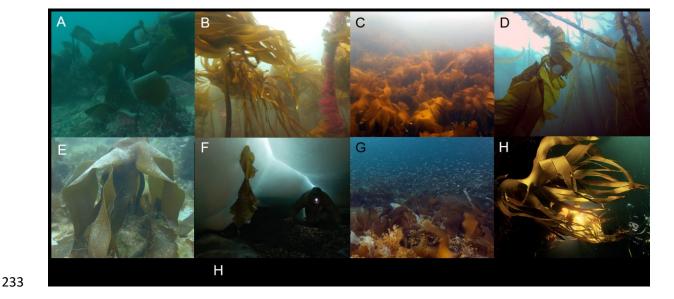


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Fig. 2. Kelp locations (red) within AMAP Arctic boundary line (orange). Gray shading shows 219 maximum sea ice extent, blue shading shows continuous permafrost (90 - 100 % cover), 220 discontinuous permafrost (50-90 %), and sporadic and isolated patches of permafrost (< 50 %) 221 222 (2016 National Snow and Ice Data Centre, https://nsidc.org/data/docs/fgdc/ggd318_map_circumarctic/). Eroding coasts (yellow) and 223 stable coasts (light green) in regions with sea ice were differentiated according to the Arctic 224 coastal classification scheme developed by Lantuit et al. (2012). 225

226

Extreme variation in environmental conditions occur within the AMAP arctic boundaries. Large regional differences in coastal conditions are strongly driven by the cover of sea ice and the presence of permafrost (frozen soil, rock, or sediment) (Lantuit et al. 2012). To capture this variability in our description of arctic kelps, we grouped information from our observational data into 3 general categories: (1) kelps on stable coasts with sea ice, (2) kelps



on unstable, eroding coasts with sea ice, and (3) kelps on coasts with little to no sea ice.

Fig 3. Photographs show examples of arctic kelp forests: (A) Laminaria solidungula in the 234 Beaufort Sea, Alaska, USA (Ken Dunton), (B and C) Laminaria hyperborea in Malangen fjord, 235 Norway (Thomas Wernberg, Karen Filbee-Dexter), (D) Eularia fistulosa Aleutian Islands, 236 Alaska (Pike Spector), (E) Saccharina latissima under sea ice in Kangiqsujuaq, Canada (PBS, 237 2017), (F) Laminaria digitata in Svalbard, Norway (Max Schwanitz), (G) Saccharina 238 latissima, S. longicruris, Alaria esculenta, Laminaria solidungula in northern Baffin Island, 239 Canada (Frithjof Küpper), and (H) Laminaria hyperborea along the Murmansk coast, Russia 240 (Dalnie Zelentsy). 241

- 242
- 243 1.3.1 Kelps on stable Arctic coasts with sea ice

Stable, rock bound coasts and fjord systems in Arctic areas with seasonal cover of sea ice can support luxurious kelp forests, although their vertical distribution is limited by ice scour (shallow) and light. These areas are expected to experience pronounced changes in environmental conditions when sea ice retreats. Although this should increase overall primary productivity along these coasts, the species composition of algae currently found in these Arctic regions may be lost permanently if more temperate-adapted algal communities push northwardand outcompete kelps that are adapted to seasonal sea ice (Krause-Jensen and Duarte 2014).

In the northern Barents Sea, kelp forests of mixed A. esculenta, L. digitata and S. 251 latissima occur within high latitude fjords off Svalbard, the western White Sea, and Franz 252 Joseph Land (Kuznetsov et al. 1994; Cooper et al. 1998; Bartsch et al. 2016; Fig 3fh). Luxuriant 253 stands of L. digitata, L. solidungula, S. dermatodea, and A. clathratum were observed within 254 255 fjords in western Novaya Zemlya (Shoshina and Anisimova 2013). In the northernmost regions around Svalbard and Novaya Zemlya, the arctic endemic kelp L. solidungula is found in inner 256 257 fjords and areas that receive cold polar currents (Svendsen 1959; Hop et al. 2012; Shoshina and Anisimova 2013). 258

The west coast of Greenland is largely rockbound and dominated by sub-littoral kelp 259 forests from Cape Farewell in the south (59° N) to Smiths Sound in the north (>80° N, 260 Rosenvinge 1893, 1899). The western Greenland kelp forests are dominated by S. longicruris 261 north of 62° N and by S. latissima south of this latitude, while other species such as L. 262 solidungula, A. esculenta, Agarum clathratum, S. nigripes and S. dermatodea are present, but 263 less conspicuous (Rosenvinge 1899; Krause-Jensen et al. 2012). The kelp forests in western 264 Greenland are narrow and shallow in the north, but become broader, more abundant, and extend 265 deeper in the south due to less ice cover (Krause-Jensen et al. 2012). In some parts of 266 Greenland, high densities of sea urchins or a lack of hard bottom restricts the extent of the kelp 267 forests (Krause-Jensen et al. 2012). The kelp populations in eastern Greenland tend to be 268 situated deeper, have less biomass per unit area and grow more slowly than those on the west 269 coast (Borum et al. 2002; Krause-Jensen et al. 2012), which may be due to lower water 270 temperatures, longer periods with ice-cover, and more heavy scour by pack ice. S. latissima 271 and A. esculenta appear to be the dominant species along most of the east coast (recorded as 272

high as Danmarks Havn (75° N)), while *L. solidungula, S. nigripes, S. longicruris and A. clathratum* are present, but less abundant (Rosenvinge 1899).

In Hudson Bay and Eastern Canada, sea ice extends below the Arctic circle due to the 275 influence of the cold Labrador current. S. latissima, A. clathratum, A. esculenta, and L. 276 solidungula have been documented between Ellesmere Island and Labrador, and along coasts 277 in Lancaster Sound, Ungava Bay, Hudson Bay, Baffin Bay, and Resolute Bay (Table 1). These 278 279 ecosystems can be highly productive in some areas, with luxuriant beds of 15-m long S. *latissima* observed in Frobisher Bay, and beds containing a biomass of 19 kg wet weight m⁻² 280 281 of A. esculenta measured in Ungava Bay (Sharp et al. 2008). Kelp forests have also been documented in eastern Chukchi Sea from Norton Sound to north of the Bering Strait along the 282 west coast of Alaska (70 and 71° N; Phillips and Reiss 1985). 283

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1.3.2. Kelps on eroding, permafrost bound Arctic coasts with sea ice

Scattered low relief, rocky coasts in the eastern Siberian, Laptev, Beaufort, and Chukchi 286 seas, and the Canadian high Arctic have temperatures and light conditions that should support 287 kelp (Krumhansl and Scheibling 2012), but observations are rare in these regions (Zenkevitch 288 1963; Lee 1973; Wilce and Dunton 2014; Wilce 2016). These coasts are more permanently 289 icebound compared to other Arctic regions- especially in the Beaufort, eastern Siberian, and 290 Laptev seas – and the seafloor is often covered in sediment due to intense glacial run off. Low 291 292 salinity, high levels of sedimentation, and sparse substrate make kelps and other macroalgae poorly developed (Taylor 1954; Leont'yev 2003; Dayton 2013). As a result, kelps along these 293 coasts face 'uniquely Arctic conditions' such as extensive sea ice scour, long periods of 294 darkness, variable salinity, turbidity, and low temperatures (Wilce 2016). The associated 295 macroalgal communities in these regions have distinct species compositions compared to other 296 regions of the Arctic, possibly because they are less connected to nearby temperate 297

communities due to outflow of polar currents from the north to south along their coasts (Wilce 298 and Dunton 2014). In the Alaskan Beaufort Sea, kelps are found in scattered rocky habitats in 299 shallow waters (5 - 10 m depth) along the mainly sedimentary coast. Research on kelps in this 300 area are from the 'Boulder Patch' (71° N), where L. solidungula forms beds intermixed with 301 A. esculenta and S. latissima on shallow cobbles and boulders (Wilce and Dunton 2014; Fig. 302 3a). These isolated kelp communities contain about half of the 140 macroalgal species found 303 304 in the Arctic. The Boulder Patch has been studied since 1978 and revisited in 14 separate years between 1978 - 2012, over which time the species composition has remained relatively static 305 306 (Wilce and Dunton 2014).

In the northwestern high Canadian Arctic, low availability of rocky substrate and a harsher climate support smaller, fragmented kelp forests (Lee 1980). This region of the Canadian Arctic commonly supports *L. solidungula*, which has been observed as high as 74.5° N.

In northeastern Russia, observations of kelp are limited to a handful of records along these sedimentary coasts, namely, *S. latissima* off Amderma, mainland Russia, Kotel Nyy Island (Cooper et al. 1998), and along the Russian coast of Chukchi Sea (Zenkevitch 1963); *L. solidungula* on islands in the Laptev Sea and within bays in the Siberian Sea (Cooper et al. 1998), and *S. latissima*, *L. solidungula*, *S. nigripes*, *A. elliptica* and *A. oblonga* in the Kara sea (Zenkevitch 1963; Guiry and Guiry 2017).

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318 1.3.3. Kelps in Arctic regions with little to no sea ice

Kelp forests in the Norwegian Sea, the Barents Sea, and the northern Pacific (Aleutian Islands and northern Gulf of Alaska) have high upper limits of biomass compared to other arctic kelp forests (Table 1; Fig 3bcd). These regions have little to no sea ice and ocean temperatures that are warmer than other Arctic regions due to the influence of the Gulf Stream

or the Pacific Current. Kelp forests in some of these regions (e.g., the Gulf of Alaska) are highly 323 influenced by environmental conditions on land, namely high freshwater inputs from melting 324 325 permafrost and melting glaciers that creates strong clines in salinity in coastal areas (Spurkland and Iken 2011; Lind and Konar 2017). Kelp in other regions with little to no sea ice appear to 326 be more influenced by biological factors than by environmental conditions. Many kelp forests 327 are strongly influenced by the density of herbivorous sea urchins, which increase with the loss 328 329 of higher level predators (e.g., crabs, cod, otters) (Doroff et al. 2003; Filbee-Dexter and Scheibling 2014). Importantly, kelps currently found in areas with little to no sea ice may 330 331 represent future scenarios for other Arctic regions.

Along the western and northern coast of Norway, and along low-lying, rock-bounded 332 coasts within the Murmansk region of Russia, Laminaria hyperborea dominates the exposed 333 coasts (Fig 3bc, Table 1) and kelp forests can obtain biomasses up to 21 kg fresh weight m⁻² 334 (Fig S1). In the mid-1970s, high densities of the green sea urchin Strongylocentrotus 335 droebachiensis destructively grazed kelp forests and created extensive urchin barrens, 336 restricting the distribution of kelp to exposed regions or shallow surf zones (Leinaas and 337 Christie 1996). Currently, regional recovery of kelp forests is occurring following decreases in 338 sea urchin populations due to reduced urchin recruitment in the south (Fagerli et al. 2013) and 339 increased crab predation in the north (Fagerli et al. 2015). 340

In the North Pacific Ocean, surface canopy forming kelps *Eualaria fistulosa* and *Nereocystis luetkeana* and subsurface kelps (*Agarum clathratum, Alaria esculenta, Costaria costada, Laminaria digitata,* and *Saccharina latissima*) form forests along the Aleutian Island chain, the northern Gulf of Alaska coast and the northeastern coast of Russia. *Eularia fistulosa* dominates surface canopies in the Aleutian Islands and *E. fistulosa* and *N. leutkeana* in southeast Alaska that can grow from > 30 m depth. Subsurface kelps tend to be competitively dominant in both regions (Duggins 1980, Dayton 1975). Kelp forests in the northern Gulf of

Alaska occur within the largest freshwater discharge system in North America, and experience 348 strong gradients of salinity due to substantial glacial inputs. The amount of glacial melt is 349 increasing with climate change, further lowering salinity and negatively effecting kelps in these 350 areas (Lind and Konar 2017). In contrast, kelp forests along the shores of the Aleutian Islands 351 are more influenced by biotic interactions. These coasts have alternated between kelp forests 352 and urchin barrens for over a century (Estes et al. 2004). Shifts between these two ecosystem 353 354 states are driven by changing abundances of sea otters, which are major predators of the sea urchin Strongylocentrotus polyacanthus (Estes and Duggins 1995). Evidence from the region 355 356 suggests that kelp forests established in 1911 after protection of sea otters enabled their populations to rebound (Estes et al. 1978). The recovered kelp forests (Eualaria fistulos and 357 Laminaria spp.) were maintained for decades, until otter populations declined again due to 358 predation by killer whales in the 1990s (Doroff et al. 2003; Estes et al. 2004), once again 359 limiting kelp forests to exposed areas and shallow depths, which serves as refuges from grazing 360 (Konar and Estes 2003). 361

362

363 1.4. ECOSYSTEM SERVICES PROVIDED BY ARCTIC KELP

Kelps can provide extensive substrate for colonizing organisms, and their canopies create 364 habitat for a number of marine plants, fish, and invertebrates (Teagle et al. 2017). The flora in 365 arctic kelp forests can be diverse and has been described in detail for some high Arctic regions 366 (e.g., Wilce and Dunton 2014; Küpper et al. 2016). Diverse fish, invertebrate and epiphytic 367 communities are found in kelp forests in Svalbard, Norway, the Aleutian Islands, the Gulf of 368 Alaska, and the Boulder Patch, USA (Hamilton and Brenda 2007; Włodarska-Kowalczuk et al. 369 2009; Wilce and Dunton 2014). Kelp canopies can create favourable conditions for some 370 understory species and were shown to provide predation refuge for juvenile cod in 371 Newfoundland, Canada (Gotceitas et al. 1995) and rockfish and ronquils in the Gulf of Alaska 372

(Dean et al. 2000b). Traditional knowledge from northern communities in Greenland reported
higher arctic cod catches in areas near kelp forests compared to other areas (Krause-Jensen and
Duarte 2014). Despite these reports, the smaller size and patchy nature of kelps in some Arctic
regions may reduce their importance as habitat forming species compared to temperate forests.
Kelp also has cultural value for northern peoples and features in their traditions and stories. It
is a traditional food for Inuit, who harvest it from under sea ice during low tide (Wein et al.
1996) and can be used by farmers as fertilizer or to cattle feed (Reedy and Katherine 2016).

Kelp-derived organic material constitutes a significant component of coastal primary 380 381 production, often forming the base of benthic food webs in nearby habitats (Dunton and Schell 1987; Fredriksen 2003; Krumhansl and Scheibling 2012). Direct consumption rates on most 382 high arctic kelps are unknown, but are likely lower than those along temperate and subarctic 383 coasts, as herbivores tend to be less abundant and the digestion of algae hypothesized to be less 384 energy efficient in colder ecosystems compared to warmer ecosystems (Floeter et al. 2005; 385 Konar 2013; Wilce 2016). Konar (2007) deployed grazer exclusion cages in experimental 386 clearings in kelp forests in the Beaufort Sea, Alaska, and found that the overall increase in algal 387 recruitment due to grazing was < 1% of the total area cleared. Similarly, the sea urchin 388 Strongylocentrotus droebachiensis, a key grazer of kelps along temperate coasts in the North 389 Atlantic (Filbee-Dexter and Scheibling 2014), is confined to shallow waters in the south 390 western Barents Sea (Murman coast), localized patches in Jan Mayen (Gulliksen et al. 1980), 391 392 Novaya Zemlya (Nordenskiøld 1880) and southern parts of Svalbard (Gulliksen and Sandnes 1980), and is rare or absent around Franz Josefs Land and the Laptev and Kara Sea (Levin et 393 al. 1998). Clear exceptions to this pattern of low grazing pressure at higher latitudes include 394 kelp forests in the Aleutian islands and northern Norway, where high consumption rates by sea 395 urchins have been recorded (Estes and Duggins 1995; Leinaas and Christie 1996). 396

Kelp carbon contributions to marine organisms in coastal environments can be 397 substantial. On average, around 80% of the kelp production globally (91% for the Boulder 398 399 Patch in the Beaufort Sea) enters coastal food webs as detritus, through detachment or exudation of dissolved organic carbon, which is exported to adjacent ecosystems on beaches 400 and deeper offshore areas (Krumhansl and Scheibling 2012). Macroalgal-derived carbon can 401 be used by benthic herbivores and predators, while upper trophic level fishes and marine 402 403 mammals generally use phytoplankton-derived carbon (McMeans et al. 2013). Stable isotope analyses show kelp carbon contributed 57% to nearshore fish populations in the Gulf of Alaska 404 405 (von Biela et al. 2016), 15 to 75% to rock greenling, predatory sea stars, and cormorants in the Aleutian Islands (Duggin et al. 1989), 0 to 42% for diverse marine predators in Baffin Island, 406 Canada (McMeans et al. 2013), and 50% to mysid crustaceans in the Beaufort Sea (Dunton and 407 Schell 1987). The latter predatory snails are a critical food source for higher trophic levels such 408 as fish, whales, and birds, indicating the high importance of kelp as a primary producer (Dunton 409 and Schell 1987). 410

A comprehensive understanding of the nature and extent of kelp subsidy to other arctic 411 benthic, pelagic, and terrestrial ecosystems is still lacking, and the magnitude and importance 412 of kelp exported from shallow coasts to deeper habitats is a debated topic of on-going research 413 (Renaud et al. 2015). In the subarctic and Arctic regions, most research has focused on the 414 vertical influx of phytoplankton- or zooplankton-derived organic matter as the main source of 415 416 carbon in benthic systems. In Greenland, Krause-Jensen et al. (2007) showed that primary production of kelps and other benthic algae can contribute to > 20% of the total primary 417 production in shallow coastal areas. However, at depths > 15 m this production was largely 418 insignificant compared to that of phytoplankton and benthic microalgae (Krause-Jensen et al. 419 2007). The magnitude of, and timing by which, kelp-derived carbon enters arctic ecosystems 420 is especially interesting because climate change is triggering earlier phytoplankton blooms in 421

the Arctic, creating temporal mismatch between pelagic primary production and some higher trophic level species that synchronize their life cycle or behaviour to this pulsed source of energy (van Leeuwe et al. 2018). In light of this mismatch, understanding other sources of arctic primary production during food-limited periods is becoming critical.

Knowing the residence time of kelp detritus in Arctic environments is important in light 426 of increased interest in blue carbon sequestration worldwide (Krause-Jensen and Duarte 2016). 427 428 In the Canadian High Arctic, large amounts of macroalgal detritus have been observed on the seafloor in sheltered fjords (Küpper et al. 2016). In northern Norway (70°N), pulses of whole 429 430 kelp blades rapidly reached deep-fjord communities (> 400 m depth) during the spring shedding of old L. hyperborea lamina (Filbee-Dexter et al. 2018). If kelp material degrades 431 slower and remains intact longer in colder arctic environments, it may be more likely to be 432 sequestered in ocean sediments than kelp carbon produced at lower latitudes. 433

434

435 1.5. KELPS IN A SENTINAL REGION OF CHANGE

Key changes that will influence kelps in the Arctic include elevated temperatures (Najafi et al. 436 2015; Wang et al. 2017), decreased cover and thickness of sea ice (Arctic Monitoring and 437 Assessment Programme. 2011; Parkinson and Comiso 2013; Ding et al. 2017), reduced 438 salinity, and increased turbidity (IPCC 2014; Günther et al. 2015). Other environmental 439 changes that could impact kelps are altered nutrients levels and increased UV radiation. 440 Reduced sea ice and warming could also bring in invasive species by increasing shipping traffic 441 or warm water species migration (Miller and Ruiz 2014), which could impact kelp 442 communities. The cumulative impact of these stressors will likely affect kelp growth rates and 443 periods severely, but ultimately depends on their nature and strength, the interactions between 444 them, and the ways in which different kelp species acclimate and/or adapt to new conditions 445 (Harley et al. 2012). 446

447

448 1.5.1. Temperature

Temperatures in the Arctic are projected to increase by $3 - 4^{\circ}$ C by the end of the 21^{st} Century 449 under realistic warming scenarios (IPCC 2014; Huang et al. 2017). Currently, kelps in Arctic 450 waters experience low temperatures with little seasonal variation. Water temperatures rarely 451 exceed 5°C in summer in the high Arctic, but may reach 10°C during summer in the southern-452 453 most parts of Arctic or where warm ocean currents affect local climate. Average temperatures may be below 0° C with a variation as small as $\pm 1^{\circ}$ C in high latitude places affected by cold 454 455 currents (e.g., Igloolik, Northwest Territories, Canada (Bolton and Lüning 1982); Young Sound, eastern Greenland (Borum et al. 2002); Franz Joseph Land, Russia (Shoshina et al. 456 2016)). 457

To explore prior and ongoing temperature changes in the vicinity of documented 458 locations of arctic kelp, we related these to maps of surface temperature for the region. We 459 calculated average temperature measures from 1986 and 2016 at each of our kelp locations 460 using historical IPCC temperature maps (IPCC 2014. accessed 461 through gisclimatechange.ucar.edu). Around each kelp location we averaged the mean summer (July to 462 September) temperature over this 20-year period within a buffer radius of 1° latitude, which 463 corresponded to the spatial error associated with locations of early records. We also calculated 464 the magnitude and rate of the predicted increase in mean summer temperature at each location 465 using climate model forecasts for 2016 to 2036 (IPCC 2014). We used the model based on the 466 conservative greenhouse gas emission scenario B1, which predicted a conservative increase of 467 1.1 to 2.9 °C by 2090-2099 relative to 1980-1999 (SRES 2000). 468

The mean summer temperature across all kelp locations has increased by 0.35° C (± 0.20) per decade over the period from 1986 to 2016 (Fig. 4a) and is predicted to increase by 1.09° C (± 0.59) per decade over the next century (Fig 4b). Predicted temperature increases are

least pronounced for kelps along the coasts of Greenland and eastern Siberia, and most
pronounced in the Barents Sea, Beaufort Sea, and Canadian High Arctic, suggesting that
changes to kelp forests due to warming will first occur in these regions.

Based on temperature tolerance and growth optima of most arctic kelp species, warmer 475 temperatures should increase growth rates (Müller et al. 2009; Shoshina et al. 2016). The 476 optimum growth temperature for most arctic and cold-temperate kelp species range from 10 to 477 478 15°C (Wiencke and Amsler 2012; Roleda 2016), and growth at 0 to 5°C is typically only 25 – 30% of growth at their optimum temperature (e.g., Bolton and Lüning 1982). Upper 479 480 temperature limits on growth of arctic kelps ranges from 16°C to 21°C (Assis et al. 2018), which are well above conditions found along Arctic coasts. This suggests warming could more 481 than double kelp production in some regions the next 2 - 3 decades. Warming may also 482 improve recruitment; for example, germination of spores, fertility (Golikov and Averintsev 483 1977), and survival of arctic kelp gametophytes are limited by temperatures below -1° C 484 (Sjøtun and Schoschina 2002; Müller et al. 2008; Assis et al. 2018) (Table 2). Such changes 485 will vary across kelp species and will likely alter their competitive interactions. In the northern 486 Gulf of Alaska, spore settlement and gametophyte growth of Eualaria fistulosa were more 487 negatively impacted by elevated temperatures and low salinity, than that of the more widely 488 distributed N. luetkeana and S. latissima (Lind and Konar 2017). A. esculenta is best adapted 489 to low temperatures and cannot survive in waters warmer than 16°C (Sundene 1962). 490 491 Likewise, recruitment of L. solidungula becomes limited when temperatures exceed 10° C. Other, more warm adapted temperate kelps such as L. hyperborea, L. digitata and Saccharina 492 polyschides may extend their range northward, following the trend of boreal species moving 493 into the Arctic (Fossheim et al. 2015; Hargrave et al. 2017; Stige and Kvile 2017). However, 494 kelps produce short-lived zoospores that disperse slowly (current patterns of kelp diversity and 495 structure can still be related to glacial cycles (Neiva et al. 2018), so any temperature-driven 496

497 northern expansion of temperate kelp species into polar regions is likely to be slow (Konar
498 2007; Wilce 2016).

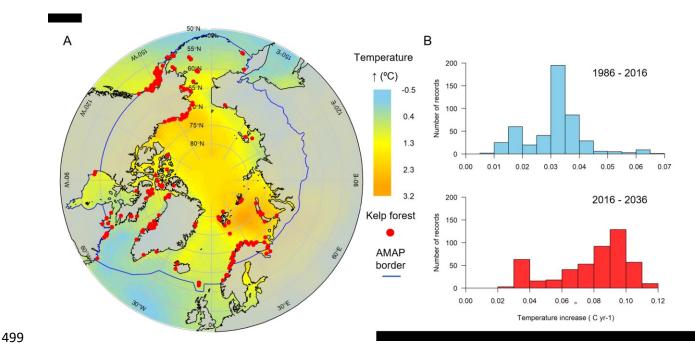


Fig. 4. a) Global trends in predicted increase in mean summer (July 21 to Sept 21) surface temperature from 2016 to 2036 according to IPCC models. Kelp locations are shown in red within AMAP Arctic boundary line (blue). b) Rate (y⁻¹) of historic and c) rate of projected warming of peak summer temperature (Aug to Sept) calculated on basis on linear trend analysis for all for all 1° latitude radius buffers around each kelp forest record.

505

506 1.5.2. Sea ice and light

The amount of light reaching the benthos is a defining factor for benthic primary production and depends largely on the extent of sea ice cover. Sea ice is rapidly retreating in the Arctic (areal loss of 3.5 - 4.5% per decade, Fig 5a). Average sea ice extent (± SD) declined by 3.7%between 2006 and 2016 (from 16.2 ± 104 to 15.6 ± 105 M km²), and by 23% in 2016 compared to average sea ice measures from 1981 to 1989 (21.4 ± 2.4 M km²).

512 To examine ongoing changes in sea ice extent at locations with records of kelp, we 513 obtained the position of the ice edge (defined by a threshold of >15% sea ice cover) from NASA satellite images taken weekly from 2006 to 2016 (http://nsidc.org/, NOAA, accessed 2017).
We constrained our measures to this period because years prior to 2006 had lower resolution
spatial measures for coastal regions. At each kelp location we calculated the nearest distance
(m) to the sea ice edge each week over the 10-year period. To compare these trends over this
last decade with broader patterns of sea ice loss we obtained daily measures of areal sea ice
extent from NASA satellite data from 1980 to 2016 (Fig 5).

520 Of the total 1179 records of kelp, 2.6% occurred in locations where the ice-free period was < 1 week in 2006 and 0.12% occurred where the ice-free period was < 1 week in 2016 521 522 (mean 0.55 ± 0.99 SD), supporting evidence of survival and growth under extremely low light conditions (Wilce 2016). On average, the annual mean and minimum distance (km) to sea ice 523 (mean \pm SD) were highly variable at kelp locations (mean 221 \pm 156 km and minimum 30 \pm 524 62 km in 2006, and mean 274 ± 341 km and minimum 49 ± 138 km in 2016; Fig. S2). For 525 records that were under sea ice for at least 1 week during this period, the mean distance to the 526 sea ice edge increased from 45 ± 24 km to 88 ± 72 km and the minimum distance to sea ice 527 edge increased from 0.53 ± 1.52 km to 0.59 ± 1.88 km from 2006 to 2016. Increases in distance 528 to sea ice were largest in the White Sea and Novaya Zemlya, Russia and southeastern 529 Greenland, and lowest in northern Canada and northeastern Russia (Fig. 5b). 530

Available evidence indicates that the loss of sea ice currently occurring in the Arctic 531 will lead to the northward expansion of kelps (Müller et al. 2009), and an increase in the depth 532 range and productivity of these habitats due to increased light and reduced scour in the surf 533 zone, which narrows the vertical distribution of kelp (Krause-Jensen et al. 2012; Krause-Jensen 534 and Duarte 2014). Kelps cannot exist in areas with permanent sea ice (Shoshina et al. 2016), 535 so ice loss may open new habitats in the high Arctic. The effect of sea ice loss on kelps may 536 even be stronger than anticipated because day length increases rapidly during the period of ice 537 break-up (Clark et al. 2013), implying a slight reduction in ice cover will result in a 538

disproportionately large increase in the amount of light reaching kelp. These expectations are supported by correlative studies from along the west coast of Greenland showing that the extent of sea ice cover explained 92% of the variation in maximum depth distribution and 80% of the variation in kelp growth (Krause-Jensen et al. 2012). Hop et al. (2012) monitored the biomass and depth range of kelps in Svalbard, Norway between 1996 and 2014 and found that kelp biomass (mainly *L. digitata*) recently increased 2 - 4 fold in the shallow zone (2.5 m depth). They ascribed these changes to reductions in sea ice cover (Bartsch et al. 2016).

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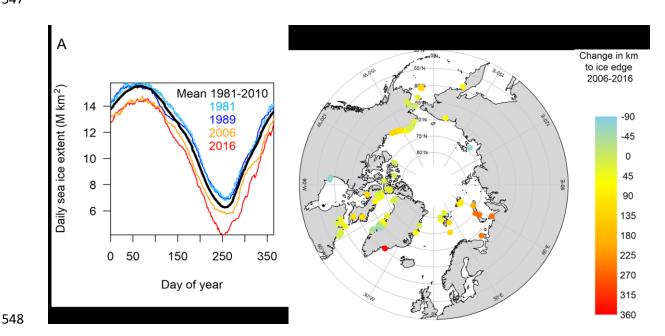


Fig 5. A) Daily sea ice extent in millions of km for entire Arctic region between 1981 and 2010.
B) Change in mean distance to sea ice edge (km) between 2006 and 2016, for locations of kelp
that occurred under ice for at least 1 week over this period.

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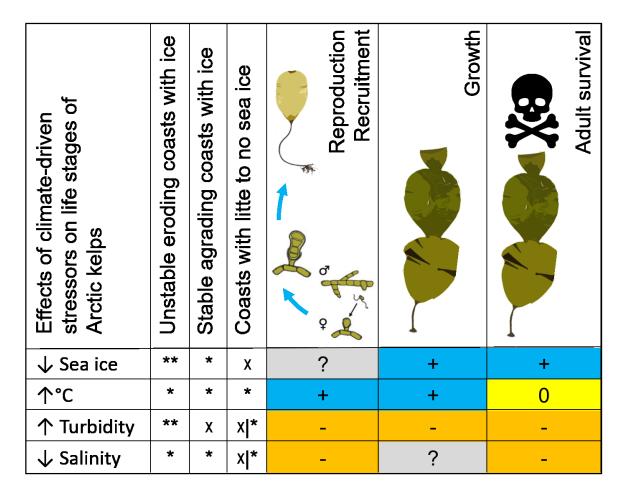
553 1.5.3. Salinity and turbidity

As a consequence of reduced sea ice and melting permafrost, many Arctic coastlines are breaking apart and eroding into the sea. These traditionally icebound coasts can be fragile because ice provides protection from storms and waves, and its loss can expose the ground to

the elements and make it unstable (Lantuit et al. 2012). Coastal environments near these 557 eroding regions are receiving higher amounts of sediment loading and freshwater inputs, 558 resulting in longer and more extreme periods of low salinity and intense turbidity and 559 sedimentation (Lantuit et al. 2012; McClelland et al. 2012; Fritz et al. 2017). Since 2000, 560 average erosion rate of permafrost-bound coasts was 0.5 m yr⁻¹, and reached 10 m per yr⁻¹ along 561 some segments. Inputs of sediment and particulate organic carbon (POC) from coastal erosion 562 are currently entering the Arctic ocean at rates \sim 430 Tg yr⁻¹ sediment and 4.9 – 14 Tg yr⁻¹ POC 563 (Fritz et al. 2017). Coastal erosion is most severe along the shallow coasts of the Laptev, East 564 565 Siberian and Beaufort Seas (Lantuit et al. 2012), but increased turbidity from melting ice can also be pronounced near the heads of Arctic fjords (Bartsch et al. 2016) and in areas receiving 566 glacial discharge (Traiger and Konar 2018). 567

Increased turbidity and reduced salinity is expected to reduce the performance and 568 lower depth limit of kelp by reducing light penetration and restricting photosynthesis (Aumack 569 et al. 2007; Fredersdorf et al. 2009; Spurkland and Iken 2011; Wiencke and Amsler 2012; 570 Traiger and Konar 2018) (Fig 6). Variable salinity reduced photosynthetic efficiency of L. 571 solidungula, S. dermatodea, L. digitata, A. esculenta and S. latissima (Karsten 2007). 572 Laboratory experiments on kelps collected from Svalbard, Norway found that sediment from 573 melting ice negatively impacted their recruitment (Zacher et al. 2016). Manipulative field 574 experiments on kelp forests in Alaska and found that glacier run-off reduced kelp settlement 575 and recruitment by increasing sedimentation in the coastal zone (Traiger and Konar 2018). 576 Research from Kola bay and anecdotal reports from areas along the Siberian shelf in Russia 577 describe declines in the lower depth limit of kelp forests due to low transparency of water (< 3 578 m visibility) caused by domestic pollution, sediment plumes and agricultural run-off 579 (Malavenda and Malavenda 2012). These negative impacts may offset the possible positive 580 effects of warming and increased light on kelp growth in some Arctic regions. This was evident 581

in the Beaufort Sea, where long-term records of annual growth of *L. solidungula* kelps showed no change in productivity since 1979, despite earlier sea ice break-up and a longer ice-free period in recent years (Bonsell and Dunton 2018). This pattern was explained by increasing resuspension of sediment and larger coastal erosion following sea ice break-up, which counter balanced the positive effect of longer ice-free periods.



587

Fig 6. Effects of environmental changes on arctic kelps from laboratory and field experiments.
+ is positive, - negative, 0 is no measurable effect, and ? is unknown. Relative importance of stressors for 3 different coastal regions (see Fig 2): ** = strong impact, * = moderate impact, and 'x' little to no impact. Note increased turbidity and decreased salinity can also occur along coasts with no sea ice that receive glacial melt or other freshwater inputs.

593

594 1.5.4. Nutrients

595 Nutrient concentrations are predicted to increase and change their seasonal timing along Arctic 596 coasts with increased (and earlier) spring melts, but the impacts of elevated nutrient richness 597 on arctic kelps are unclear. Nutrient availability is typically low in most Arctic waters, and 598 nutrient concentrations tend to increase during winter when primary production is low, but 599 decrease to extremely low levels during the short Arctic summer. Therefore, pelagic primary 600 production is therefore often limited by low nutrient availability in late summer.

601 This may not be the case for kelps. In a study of twenty-one different species of arctic macroalgae (including Laminaria spp.), none of them were significantly nitrogen-limited in 602 603 July (Gordillo et al. 2006). Kelps may be able to acquire and accumulate nutrients in winter when nutrient availability is relatively high. Nutrients can be translocated from the blade 604 towards the meristem (Davison and Stewart 1983) and nutrient reserves can subsequently be 605 606 used to support photosynthesis and, thus, prolong blade growth during summer when insolation is high and nutrient availability is low (Gagne et al. 1982; Henley and Dunton 1997; Pueschel 607 and Korb 2001). Most kelp species should therefore remain rather unaffected by increasing 608 nutrient availability, but studies have shown that the growth of at least some species, here L. 609 solidungula, decreases significantly in early spring as nutrient concentrations drop (Chapman 610 and Lindley 1980; Dunton et al. 1982). This suggests that some kelp species and/or kelps in 611 certain extremely nutrient poor areas can be limited by low nutrient availability, and therefore 612 would be stimulated by increased nutrient levels. 613

It is important to note that pelagic phytoplankton are more stimulated by increasing nutrient and light levels compared to benthic algae. Estimates predict thus that the pelagic production by phytoplankton in some Arctic waters will increase 3-fold within this century due to longer ice-free periods and increased run-off from land (e.g., Rysgaard and Glud 2007). This significant increase in phytoplankton biomass and productivity will decrease light penetration 619 in the water column, which will negatively affect kelp biomass and depth limit, possibly620 offsetting any benefits that higher nutrient levels could have on some kelp species.

621

622 1.5.5. UV radiation

Other changes in environmental conditions that could impact kelps include increased 623 UV radiation, which is especially pronounced at high latitudes (Garcia-Corral et al. 2014). 624 625 Increases in UV radiation negatively impacts photosynthesis of arctic kelps (Roleda et al. 2006; Müller et al. 2008; Roleda 2016) and reduces their performance (Heinrich et al. 2015). 626 However, research to date indicates that UV damage will have a minor impact on arctic kelps 627 compared other environmental changes, and will mainly affect early life stages (Roleda et al. 628 2006; Wiencke et al. 2006). In laboratory experiments on L. solidungula collected from 629 Svaldbard by Roleda (2016), high UV radiation disrupted the life cycle of meiospores and 630 gametophytes. UV exposure also caused significant declines in photosynthetic efficiency, and 631 increased transcription of DNA repair genes, but these effects were less pronounced in kelps 632 collected from the field compared to cultured plants (Heinrich et al. 2015). Fredersdorf et al. 633 (2009) examined combined effects of different temperatures, salinity, and UV radiation levels 634 on photosynthesis of A. esculenta collected from Svalbard. They found that A. esculenta 635 zoospores were sensitive to synergistic effects of temperature and salinity changes (Fredersdorf 636 et al. 2009), but that adults A. esculenta could tolerate a range of UV conditions. 637

638

639 1.6. PREDICTING CHANGES TO DISTRIBUTION OF ARCTIC KELPS

640 Predicting changes to arctic kelps under rapidly changing environmental conditions remains641 challenging. Assis et al. (2018) developed models that described the current distributions of

642 Alaria esculenta, L. solidungula, L. digitata, L. hyperborea, S. latissima, and S. dermatoada in

643 the northern Atlantic according to environmental parameters (mainly sea temperature, sea ice,

salinity, upwelling), and used these relationships to predict the impacts of climate change on 644 their future distribution. These models predicted large northward expansions of these species, 645 646 including the expansion of L. hyperborea to Svalbard, Norway, and further into the White Sea, the spread of S. dermatoada and L. digitata (or S. nigripes depending on source, S. Fredriksen 647 personal communication) along the northeastern coast of Greenland, and the expansion of A. 648 esculenta into the Canadian high Arctic. The models also predicted L. solidungula and S. 649 650 latissima would extend northward to cover the northernmost coasts of Greenland, Russia and Canada, suggesting that all Arctic coasts would have environmental conditions suitable for kelp 651 652 forests in the future. Similar range expansions have been predicted for L. solidungula and S. *latissima* with models by Müller et al (2009) and for a number of fucoid species by Jueterbock 653 et al. (2013, 2016). However, there is a discrepancy between these predictive models and long-654 term field observations of changes to arctic kelps. In Canada, Adey and Hayek (2011) were 655 unable to identify significant shifts in the distributions of subtidal algal species in the eastern 656 subarctic or boreal regions over the past 40 years. Likewise, Merzouk & Johnson (2011) 657 reviewed the distribution of kelp along the northwest Atlantic shores from records dating back 658 to the 1950s and were unable to document any significant change in dominant kelp species 659 composition or abundance over that period, despite increasing sea temperature, although, the 660 lack of sufficient spatially and temporally extensive datasets for this region prevented them 661 from concluding that no change had occurred. Northward range expansions of kelps may be 662 limited by extensive gaps between suitable substrate (e.g., from northern Norway to Svalbard) 663 and low dispersal potential of kelps (Wernberg et al. 2018). It is also possible that the spread 664 and performance of kelps may be more influenced by changes in turbidity, sea ice cover, and 665 light penetration compared to relatively small changes in sea temperatures. This suggests that 666 model predictions may overestimate northern range expansions of kelps, at least in the short-667 668 term.

669

670 1.7. CONCLUSIONS

The Arctic is at the epicenter of the global climate crisis, and emerging opportunities and 671 developments have increased international attention on changes to ecosystems in this area. 672 Long-term research from Greenland and Norway suggests a warmer Arctic with less sea ice 673 may support higher kelp productivity and biomass and expand the northern range and lower 674 675 depth limit of these species. However, the degree to which these changes will positively affect kelps will vary regionally and depend on the extent that melting sea ice and permafrost 676 677 increases turbidity in coastal areas, as well as the available substrate in the lower depth range (Bartsch et al. 2017; Bonsell and Dunton 2018). Predictive models and laboratory experiments 678 suggest the 'borealization' of arctic kelp forests will occur as temperatures warm, altering the 679 species composition of existing cold and ice-adapted kelp communities in high Arctic regions. 680 Although current predictions are highly uncertain, the possible expansion of kelp forests should 681 provide new habitats for fish and other marine organisms, and a suite of valuable ecosystem 682 services along Arctic coastlines. Interestingly, where data are available, kelp abundance 683 appears relatively stable, suggesting these changes are occurring slower than predicted or are 684 being buffered by other factors. Either way, anticipating these changes, and understanding 685 these new ecosystems will be a key priority for northern communities. 686

Our understanding of kelp forests is rapidly expanding in many regions of the Arctic. However, baseline measures of the extent of kelp communities are missing in northern and eastern Canadian Arctic, Siberia, the east Greenland Shelf, and Russia. This lack of data is not unique to kelp ecosystems. Despite the fact that over 28% of the world's coastlines are found in the Arctic (Lantuit et al. 2012), they remain largely unstudied, which jeopardizes current strategies to protect or conserve arctic environments and will have consequences for northern communities that rely on them. Lack of data has already greatly hindered our ability to detect

and understand the impacts of climate change on these and other ecosystems (e.g., Merzouk 694 and Johnson 2011). Exploring effects of ongoing and future climate changes will provide 695 important insight on the stability of these ecosystems. Maintaining and augmenting current 696 monitoring initiatives and time series data sets should be a priority. For kelp forests, 697 understanding how these ecosystems influence the structure and function of coastal arctic food 698 webs is an important focus for ongoing research. There is also a critical lack of knowledge on 699 700 the contribution of kelp forests to carbon cycling in the Arctic. Filling in these gaps and strategically prioritizing research in areas of rapid environmental variation will enable us to 701 702 more effectively understand and conserve these ecosystems.

Arctic coasts are in line to become one of the most impacted environments in the world 703 under changing climate. For this region to act as a sentinel for climate change it is critical to 704 705 monitor and understand the impacts of environmental stressors on arctic ecosystems. Kelp forests provide a key example of the regional diversity of responses to climate change, and 706 demonstrate the need for a mechanistic understanding of how multiple stressors and diverse 707 ecological processes influence ecosystem structure and function. Although it is tempting to 708 make generalized statements about broad-scale climate-driven impacts, the reality is much 709 more nuanced, regionally specific, and highly uncertain. What is clear is that extensive 710 ecological changes are likely to occur in these rapidly changing environments, with both 711 'positive' or 'negative' consequences for a range of species. 712

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1101

1102 Table 1/S1. Species composition, depth limit and biomass (wet weight per m²) of Arctic kelp forests. Bolded names indicate dominant species. (-

1103) is not reported.

Location	Site	Year	Depth	Species	Latitude, Long	Kelp WW	Reference
		Sam	limit			(g m ⁻²)	
		pled	(m)			Mean ± SE (n)	
Canada							
Hudson and Ungava	Kangirsuk			L. solidungula S. longicruris	60.0373, -70.1796	11.8 ± 1.3 (25)	(Sharp et al. 2008)
Bay							
Hudson and Ungava	Basking I		10	L. solidungula S. longicruris	59.9848, -69.9478	2.9 ± 0.2 (25)	(Sharp et al. 2008)
Bay							
Labrador sea	E. Port Markham	2003	30	A. clathratum A. esculenta	52.3667, -55.7333	801.8	(Adey and Hayek 2013)
Labrador sea	Tilcey I	2003	20	A. clathratum A. esculenta L. digitata S.	52.2167, -55.6333	1808.8	(Adey and Hayek 2013)
				dermatodea S. latissima			
Labrador sea	South Cove	2003	30	A. clathratum A. esculenta S.	53.2167, -55.6333	4109.8	(Adey and Hayek 2013)
				dermatodeaS. latissima S. longicruris			
Baffin Bay	Walls I, Cape St.	2003	12	A. clathratum A. esculenta L. digitata S.	52.2167, -55.6167	1903.4	(Adey and Hayek 2013)
	Charles			dermatodea S. latissima			
Hudson and Ungava	Tuvalik Pt.		12	A. clathratum A. esculenta L. solidungula	60.0568, -69.6745	8.4 ± 1.1 (25)	(Sharp et al. 2008)
Bay				S. groenlandica S. longicruris			

Hudson and Ungava	Pikyuluk I		12	A. esculenta L. digitata, L. solidungula, S.	59.9868, -69.9337	9.2 ± 2 (25)	(Sharp et al. 2008)
Bay				longicruris			
Greenland							
Baffin Bay	Qaanaaq	2009		A. clathratum S. latissima S. longicruris	77.4667, -69.2500	15.0 ± 2.6^{1}	(Krause-Jensen et al. 2012)
Baffin Bay	Dundas				77.5500, -68.8667	$14.9\pm0.8^{\rm 1}$	(Krause-Jensen et al. 2012)
Baffin Bay	Uummannaq	2009	33	A. clathratum S. latissima	70.6667, -51.6000	24.1 ± 4.0^1	(Krause-Jensen et al. 2012)
Labrador sea	Disko Bay				69.4833, -53.6333	$18.8\pm0.9^{\scriptscriptstyle 1}$	(Krause-Jensen et al. 2012)
Labrador sea	uuk	2008	30	A. clathratum A. esculenta S. longicruris	64.1333, -51.6167	18.0 ± 1.1^{1}	(Krause-Jensen et al. 2012)
Labrador sea	Eqip Sermia	2009	27	A. clathratum S. latissima	69.7500, -50.3500	12.6 ± 2.8^{1}	(Krause-Jensen et al. 2012)
Norway							
Norwegian Sea	Finnøy-Håvær V	2012	20	A. esculenta L. hyperborea S. latissima	62.8203, 6.5472	$1141.1 \pm 349,1$	(Christie et al. 2014 (NIVA
							report))
Norwegian Sea	Finnøy-Håvær N	2012		A. esculenta L. hyperborea S. latissima	62.8252, 6.5546	$1301.0\pm360{,}3$	(Christie et al. 2014)
Norwegian Sea	Vega-Ivarsbraken	2012	20	A. esculenta L. hyperborea S. latissima	65.6764, 11.5494	$1589.7 \pm 377,7$	(Christie et al. 2014)
Norwegian Sea	Vega-Bubraken	2012	20	A. esculenta L. hyperborea S. latissima	65.6802, 11.5984	$712.7 \pm 246,2$	(Christie et al. 2014)
Norwegian Sea	Vega-Igerøy	2012		A. esculenta L. hyperborea S. latissima	65.6901, 12.1310	788.3 ± 133,9	(Christie et al. 2014)
Norwegian Sea	Senja-Sjursvika	2012	20	A. esculenta L. hyperborea S. latissima	69.0956, 16.7792	$818.4 \pm 174,5$	(Christie et al. 2014)
Norwegian Sea	Senja-Stongeland	2012	20	A. esculenta L. hyperborea S. latissima	69.0427, 16.8795	$307.8\pm69{,}0$	(Christie et al. 2014)

Norwegian Sea	Senja-Halvardsøya	2012	20	A. esculenta L. hyperborea S. latissima	69.1599, 16.8958	$864.3 \pm 115,9$	(Christie et al. 2014)
Norwegian Sea	Senja-	2012	20	A. esculenta <i>L. hyperborea</i> S. latissima	69.3110, 16.8978	$741.8\pm135{,}9$	(Christie et al. 2014)
	Kjerringbergnes						
Norwegian Sea	Senja-Månesodden	2012	20	A. esculenta <i>L. hyperborea</i> S. latissima	69.3111, 16.8978	$1038.7\pm92{,}3$	(Christie et al. 2014)
Norwegian Sea	Senja-Lemmingsvær	2012	20	A. esculenta <i>L. hyperborea</i> S. latissima	69.0270, 16.9326	$561.2\pm125,\!3$	(Christie et al. 2014)
Norwegian Sea	Hekkingen I	2016	10	A. esculenta L. hyperborea S. latissima	69.6167, 17.8860	21976.0 ± 2967,0	(Filbee-Dexter et al. 2018)
Barents Sea	Kongsfjorden	2013	20	A. esculenta L. digitata L. solidungula S.	78.9833, 11.9632	4614.0	(Bartsch et al. 2016; Hop et
				dermatodea S. latissima			al. 2016)
Barents Sea	Finnmark-	2012	20	A. esculenta <i>L. hyperborea</i> S. latissima	70.6991, 29.4393	$691.7\pm110{,}7$	(Christie et al. 2014)
	Kongsfjord						
Barents Sea	Posangerfjord	-	-			4.1 ± 1.8	(Christie et al. 2014)
Barents Sea	Finnmark-Bøkefjord	2012	20	A. esculenta <i>L. hyperborea</i> S. latissima	69.8525, 30.1300	$703.5\pm163,\!9$	(Christie et al. 2014)
Russia							
Barents Sea	Cape Abram		15	S. latissima	69.0210, 33.0226	613.3	(Shoshina et al. 2016)
Barents Sea	Cape Mishukov		6	A. esculenta S. latissima	69.0595, 33.0429	183.3	(Malavenda and Malavenda
							2012)
Barents Sea	Belokamenka Bay		6	S. latissima	69.0777, 33.1807	836.7	(Malavenda and Malavenda

Barents Sea	Cape Retinskiy		6	A. clathratum L. digitata S. latissima	69.1122, 33.3793	1550.0	(Malavenda and Malavenda 2012)
							- /
White sea	Ostrov Asafiy	1973	9	S. latissima	66.4210, 33.6559	1922.0	(Myagkov 1975)
White sea	Nikolskaya Bay		8	L. digitata S. latissima	66.2167, 33.8333	$5232.0 \pm 1201,0$	(Plotkin et al. 2005)
USA							
Beaufort sea	Boulder patch	1980	7	A. esculenta L. solidungula S. latissima	70.3208, -147.5833	262.0	(Dunton and Schell 1986;
							Dunton et al. 1982)
Gulf of Alaska	Knight Island	1998		A. cribosum E. fistulosa L. spp. S.	60.3327, -147.7644	$900\pm200~SE$	(Dean et al. 2000a)
				latissima			
Aleutian Islands	Tanaga I, Adak I,	2016	-	A. clathratum E. fistulosa L. spp.	51.5521,-178.4067;	$1908\pm372~SE^2$	(Konar et al. 2017)
	Atka I, Chuginadak I			Ondonthalia setacea Ptilota serrata	51.6102,-177.0966;		
				Laminaria longipes	51.8619,-175.1848;		
					52.3509,-170.8579		
Aleutian Islands	Umnak I/Anangula I,	2016	-	A. clathratum E. fistulosa Laminaria spp.	52.7790,-169.3972;	$3523\pm674~SE^2$	(Konar et al. 2017)
	Unalaska I			Ondonthalia setacea Ptilota serrata	53.2908,-167.9203		
				Laminaria longipes			
Aleutian Islands	Adak I	1987	30	<i>E. fistulosa Laminaria</i> spp.	51.6102,-177.0966	2920 ± 1810	(Duggins et al. 1989)
	Amchitka I	1987	30	<i>E. fistulosa Laminaria</i> spp.	51.5043,178.4812	2628 ± 1912	(Duggins et al. 1989)
Aleutian Islands	Kiska I	_		A. cribosum <i>E. fistulosa Laminaria</i> spp.	51.5961, -178.6748	12645 ± 4999	(Wilmers et al. 2012)

	Aleutian Islands	Ogliuga I	- A. cribosur	m, E. fistulosa, Laminaria spp.	52.0563,177.4398	12645 ± 4999	(Wilmers et al. 2012)
1104	¹ Dry weight. ² SE of dor	minant species E. fistulos	<i>a</i>				
1105							
1106							
1107							