



1 **Production of HONO from heterogeneous uptake of NO₂ on**
2 **illuminated TiO₂ aerosols measured by Photo-**
3 **Fragmentation Laser Induced Fluorescence**

4 Joanna E. Dyson¹, Graham A. Boustead¹, Lauren T. Fleming¹, Mark Blitz^{1,2}, Daniel Stone¹,
5 Stephen R. Arnold³, Lisa K. Whalley^{1,2}, Dwayne E. Heard^{1*}

6 ¹ *School of Chemistry, University of Leeds, LS2 9JT, UK.*

7 ² *National Centre of Atmospheric Science, University of Leeds, LS2 9JT, UK.*

8 ³ *School of Earth and Environment, University of Leeds, LS2 9JT, UK.*

9 **Corresponding Author. Email: D.E.Heard@leeds.ac.uk*

10

11 **Abstract**

12 The rate of production of HONO from illuminated TiO₂ aerosols in the presence of NO₂ was
13 measured using an aerosol flow tube coupled to a photo-fragmentation laser induced
14 fluorescence detection apparatus. The reactive uptake coefficient of NO₂ to form HONO,
15 $\gamma_{NO_2 \rightarrow HONO}$, was determined for NO₂ mixing ratios in the range 34 – 400 ppb, with $\gamma_{NO_2 \rightarrow HONO}$
16 spanning the range $(9.97 \pm 3.52) \times 10^{-6}$ to $(1.26 \pm 0.17) \times 10^{-4}$ at a relative humidity of 15 ± 1
17 % and for a lamp photon flux of $(1.63 \pm 0.09) \times 10^{16}$ photons cm⁻² s⁻¹ (integrated between 290
18 and 400 nm), which is similar to values of ambient actinic flux at midday. $\gamma_{NO_2 \rightarrow HONO}$
19 increased as a function of NO₂ mixing ratio at low NO₂ before peaking at $(1.26 \pm 0.17) \times 10^{-4}$
20 at 51 ppb NO₂ and then sharply decreasing at higher NO₂ mixing ratios, rather than levelling
21 off which would be indicative of surface saturation. The dependence of HONO production on
22 relative humidity was also investigated, with a peak in production of HONO from TiO₂ aerosol
23 surfaces found at ~25 % RH. Possible mechanisms consistent with the observed trends in both
24 the HONO production and reactive uptake coefficient were investigated using a zero-
25 dimensional kinetic box model. The modelling studies supported a mechanism for HONO
26 production on the aerosol surface involving two molecules of NO₂, as well as a surface HONO
27 loss mechanism which is dependent upon NO₂. In a separate experiment, significant production
28 of HONO was observed from illumination of mixed nitrate/TiO₂ aerosols in the absence of
29 NO₂. However, no statistically significant production of HONO was seen from the illumination



30 of pure nitrate aerosols. The rate of production of HONO observed from mixed nitrate/TiO₂
31 aerosols was scaled to ambient conditions found at the Cape Verde Atmospheric Observatory
32 (CVAO) in the remote tropical marine boundary layer. The rate of HONO production from
33 aerosol particulate nitrate photolysis containing a photocatalyst was found to be similar to the
34 missing HONO production rate necessary to reproduce observed concentrations of HONO at
35 CVAO. These results provide evidence that particulate nitrate photolysis may have a significant
36 impact on the production of HONO and hence NO_x in the marine boundary layer where mixed
37 aerosols containing nitrate and a photocatalytic species such as TiO₂, as found in dust, are
38 present.

39

40 **1 Introduction**

41 A dominant source of OH radicals in polluted environments is the photolysis of nitrous acid
42 (HONO) (Platt et al., 1980; Winer and Biermann, 1994; Harrison et al., 1996; Alicke et al.,
43 2002; Whalley et al., 2018; Crilley et al., 2019; Lu et al., 2019; Slater et al., 2020; Whalley et al.,
44 2020). During a recent study in Winter in central Beijing, HONO photolysis accounted for over
45 90 % of the primary production of OH averaged over the day (Slater et al., 2020). Oxidation
46 by OH radicals is the dominant removal mechanism for many tropospheric trace gases, such as
47 tropospheric methane, as well as the formation of secondary species, including tropospheric
48 ozone (Levy, 1971), nitric and sulphuric acids which condense to form aerosols, and secondary
49 organic aerosols. Understanding the formation of HONO in highly polluted environments is
50 crucial to fully understand both the concentration and distribution of key atmospheric radical
51 species, as well as secondary products in the gas and aerosol phases associated with climate
52 change and poor air quality.

53 Atmospheric concentrations of HONO range from a few pptv in remote clean environments
54 (Reed et al., 2017) to more than 10 ppb in highly polluted areas such as Beijing (Crilley et al.,
55 2019). The main gas-phase source of HONO in the troposphere is the reaction of nitric oxide
56 (NO) with the OH radical. HONO has also been shown to be directly emitted from vehicles
57 (Kurtenbach et al., 2001; Li et al., 2008), for which the rate of emission is often estimated as a
58 fraction of known NO_x (NO₂+NO) emissions. Many heterogeneous HONO sources have also
59 been postulated including the conversion of nitric acid (HNO₃) on ground or canopy surfaces
60 (Zhou et al., 2003; George et al., 2005), bacterial production of nitrite on soil surfaces (Su et
61 al., 2011; Oswald et al., 2013) and, more recently, particulate nitrate photolysis, thought to be



62 an important source in marine environments (Ye et al., 2016; Reed et al., 2017; Ye et al.,
63 2017a; Ye et al., 2017b). Rapid cycling of gas-phase nitric acid to gas-phase nitrous acid via
64 particulate nitrate photolysis in the clean marine boundary layer has been observed during the
65 2013 NOMADSS aircraft measurements campaign over the North Atlantic Ocean (Ye et al.,
66 2016). Ground-based measurements of HONO made at Cape Verde in the tropical Atlantic
67 Ocean (Reed et al., 2017) provided evidence that a mechanism for renoxification in low NO_x
68 areas is required (Reed et al., 2017; Ye et al., 2017a).

69 Recent model calculations show a missing daytime source of HONO, which is not consistent
70 with known gas-phase production mechanisms, direct emissions or dark heterogeneous
71 formation (e.g. prevalent at night). It has been suggested that this source could be light driven
72 and dependent on NO₂ (Kleffmann, 2007; Michoud et al., 2014; Spataro and Ianniello, 2014; Lee
73 et al., 2016).

74 It is estimated that between 1604 and 1960 Tg yr⁻¹ of dust particles are emitted into the
75 atmosphere (Ginoux et al., 2001). Titanium dioxide (TiO₂) is a photocatalytic compound found
76 in dust particles at mass mixing ratios of between 0.1 and 10 % depending on the location the
77 particles were suspended (Hanisch and Crowley, 2003). When exposed to UV light ($\lambda < 390$
78 nm) TiO₂ promotes an electron (e_{CB}^-) from the conduction band to the valence band leaving
79 behind a positively charged hole (h_{VB}^+) in the valence band (Chen et al., 2012):



80 which can then lead to both reduction and oxidation reactions of any surface adsorbed gas-
81 phase species such as NO₂ leading to HONO.

82 In previous studies of the reaction of NO₂ on TiO₂ aerosol surfaces, HONO was observed as a
83 major gas-phase product (Gustafsson et al., 2006; Dupart et al., 2014). Gustafsson *et al.*, (2006)
84 observed a yield of gas-phase HONO of ~ 75 % (for each NO₂ removed), and showed the rate
85 of the photoreaction of NO₂ on pure TiO₂ aerosols depended on relative humidity, emphasising
86 the superhydrophilic nature of TiO₂ surfaces under UV irradiation. Dupart *et al.* (2014) also
87 reported a relative humidity dependence of the uptake of NO₂ onto Arizona Test Dust
88 containing TiO₂ with the main gas-phase products measured being NO and HONO, with a
89 HONO yield of 30 % in experiments with 110 ppb NO₂. Dupart et al. (2014) postulated the
90 following mechanism of HONO production, which is consistent with the formation of the NO₂⁻
91 anion seen in a previous study on TiO₂ surfaces (Nakamura et al., 2000):



92 In areas with high mineral dust loading, such as desert regions, far from anthropogenic sources,
93 NO_2 concentrations are typically low. However, when dust is transported to urban areas, this
94 source of HONO may become significant. One study reported that TiO_2 composed 0.75-1.58
95 $\mu g\ m^{-3}$ when aerosol loadings were 250-520 $\mu g\ m^{-3}$ over the same time period in southeast
96 Beijing, when air had been transported from the Gobi desert (Schleicher et al., 2010).

97 In this study, the production of HONO on the surface of TiO_2 particles in the presence of NO_2
98 is investigated as a function of NO_2 mixing ratio, aerosol surface area density and relative
99 humidity using an aerosol flow tube system coupled to a photo-fragmentation laser induced
100 fluorescence detector (Boustead, 2019). The uptake coefficient of NO_2 to generate HONO is
101 then determined, and a mechanistic interpretation of the experimental observations is
102 presented. The production of HONO directly in the absence of NO_2 from the illumination of a
103 mixed sample of nitrate and TiO_2 aerosol is also presented. Using a similar apparatus, previous
104 work had showed that TiO_2 particles produce OH and HO_2 radicals directly under UV
105 illumination (Moon et al., 2019). The atmospheric implications of these results and the role of
106 photo-catalysts for the formation of HONO are also discussed.

107 2 Method

108 2.1 Overview of the Experimental Setup

109 The production of HONO from illuminated aerosol surfaces is studied using an aerosol flow
110 tube system coupled to a photo-fragmentation laser induced fluorescence (PF-LIF) cell which
111 allows the highly sensitive detection of the OH radical formed through photo-fragmentation of
112 HONO into OH and NO followed by Laser-Induced Fluorescence (LIF) detection at low
113 pressure. The experimental setup used in this investigation is described in detail in (Boustead,
114 2019), therefore only a brief description of the setup is given here. A schematic of the
115 experimental setup is shown in Figure 1.

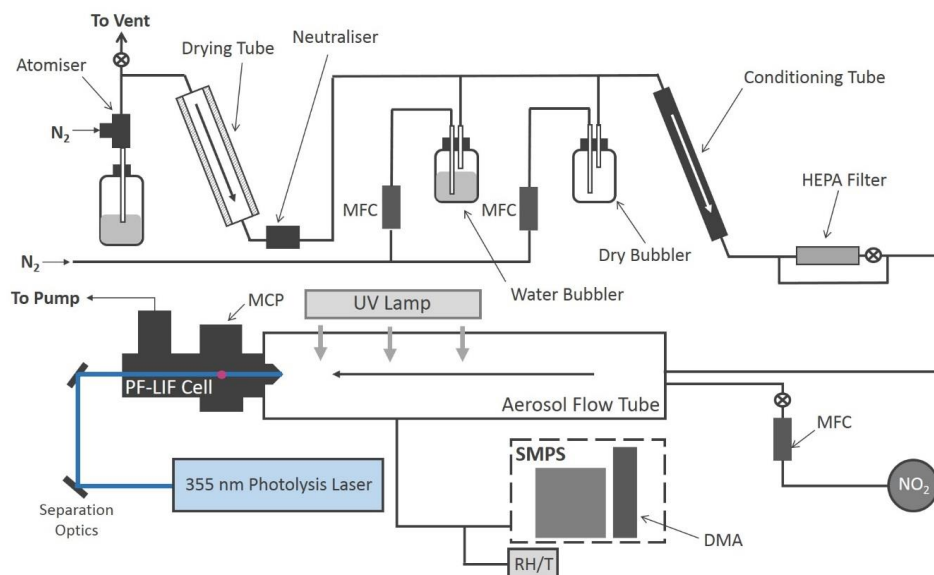


Figure 1. Schematic of the Leeds aerosol flow tube system coupled to a laser-fragmentation laser-induced fluorescence detector for HONO. The paths of the 355 nm (blue) and 308 nm (purple, depicted as travelling out of the page perpendicular to the 355 nm light) light are also shown. CPC: condensation particle counter; DMA: differential mobility analyser; HEPA: high efficiency particle air filter; FAGE: fluorescence assay by gas expansion; MCP: microchannel plate photomultiplier; MFC: mass flow controller; RH/T: relative humidity/ temperature probe; SMPS: scanning mobility particle sizer.

116

117 All experiments were conducted at room temperature (295 ± 3 K) using nitrogen (BOC, 99.998
118 %) or air (BOC, 21 ± 0.5 % O_2) as the carrier gas. A humidified flow of aerosols, ~ 6 lpm (total
119 residence time of 104 s in the flow tube), was introduced through an inlet at the rear of the
120 aerosol flow tube (Quartz, 100 cm long, 5.75 cm ID) which was covered by a black box to
121 eliminate the presence of room light during experiments. A 15 W UV lamp (XX-15LW Bench
122 Lamp, $\lambda_{\text{peak}}=365$ nm) was situated on the outside of the flow tube to illuminate aerosols and
123 promote the production of HONO (half the length of the flow tube was illuminated leading to
124 an illumination time of 52 s). The concentration of HONO is measured by PF-LIF with
125 sampling from the end of the flow tube via a protruding turret containing a 1 mm diameter
126 pinhole, through which the gas exiting the flow tube was drawn into the detection cell at 5 lpm.
127 The detection cell was kept at low pressure, ~ 1.5 Torr, using a rotary pump (Edwards, E1M80)
128 in combination with a roots blower (Edwards, EH1200). All gas flows in the experiment were
129 controlled using mass flow controllers (MKS and Brooks). The relative humidity (RH) and
130 temperature of the aerosol flow was measured using a probe (Rotronics HC2-S, accuracy ± 1



131 % RH) the former calibrated against the H₂O vapour concentration measured by a chilled
132 mirror hygrometer (General Eastern Optica), in the exhaust from the flow tube.

133 **2.2 Aerosol generation and detection**

134 Solutions for the generation of TiO₂ aerosol solutions were prepared by dissolving 5 g of
135 titanium dioxide (Aldrich Chemistry 718467, 99.5% Degussa, 80 % anatase: 20 % rutile) into
136 500 ml of milli-Q water. Polydisperse aerosols were then generated from this solution using an
137 atomiser (TSI model 3076) creating a 1 lpm flow of TiO₂ aerosol particles in nitrogen hereafter
138 referred to as the aerosol flow. This aerosol flow was then passed through a silica drying tube
139 (TSI 3062, capable of reducing 60 % RH incoming flow to 20 % RH) to remove water vapour,
140 then passed through a neutraliser to apply a known charge distribution and reduce loss of
141 aerosols to the walls. After the neutraliser the aerosol flow was mixed with both a dry and a
142 humidified N₂ flow (controlled by MFCs) to regulate the relative humidity of the system by
143 changing the ratio of dry to humid nitrogen flows. A conditioning tube was then used to allow
144 for equilibration of water vapour adsorption and re-evaporation to and from the aerosol surfaces
145 for the chosen RH, which was controlled within the range ~10-70 % RH. A portion of the
146 aerosol flow was then passed through a high efficiency particle filter (HEPA) fitted with a
147 bypass loop and bellows valve allowing control of the aerosol number concentration entering
148 the aerosol flow tube. Previous studies (George et al., 2013; Boustead, 2019) have shown the
149 loss of aerosol to the walls of the flow tube to be negligible. Aerosol size distributions were
150 measured for aerosols exiting the flow tube using a scanning mobility particle sizer (SMPS,
151 TSI 3081) and a condensation particle counter (CPC, TSI 3775) which was calibrated using
152 latex beads. Any aerosol surface area not counted due to the upper diameter range of the
153 combined SMPS/CPC (14.6 – 661.2 nm, sheath flow of 3 lpm, instrumental particle counting
154 error of 10-20 %) was corrected for during analysis by assuming a lognormal distribution,
155 which was verified for TiO₂ aerosols generated in this manner (Matthews et al., 2014).
156 However, the majority of aerosols, >90 %, had diameters in the range that could be directly
157 detected. Previously, we had determined by imaging the aerosols using a scanning electron
158 microscope (SEM) that the particles were spherical (Moon et al., 2018). In addition to the
159 experiments with single-component TiO₂, mixed ammonium nitrate/TiO₂ and single-
160 component ammonium nitrate aerosols were also generated using the atomiser for
161 investigations of HONO production from nitrate aerosols without NO₂ present. An example of
162 an aerosol size distribution from this work for single-component ammonium nitrate aerosols,
163 mixed ammonium nitrate/TiO₂ and single-component TiO₂ aerosols is shown in Figure 2.

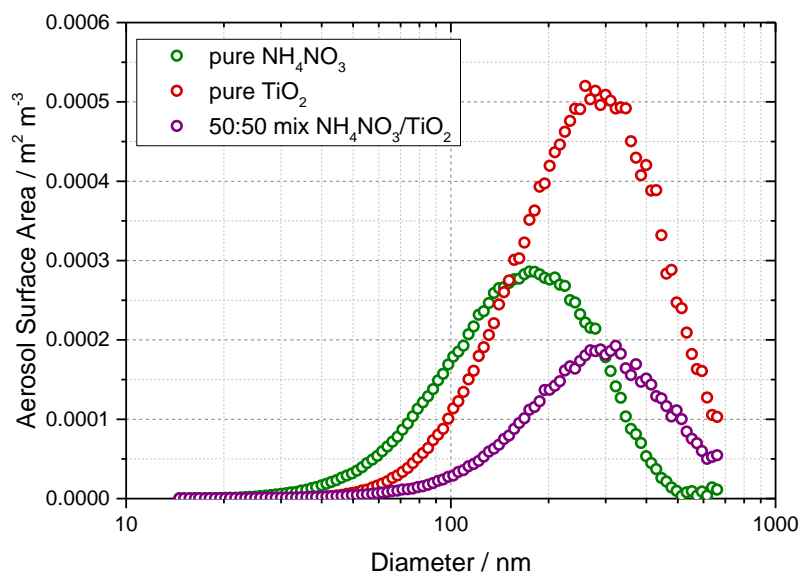


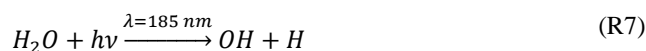
Figure 2 Typical aerosol surface area distribution for pure ammonium nitrate aerosols (green) and pure TiO₂ aerosols (red) and 50:50 mixed nitrate/TiO₂ aerosols (purple) measured after the flow tube.

164 2.3 Detection of HONO

165 As HONO is not directly detectable via LIF, it was necessary to fragment the HONO produced
166 into OH and NO (Liao et al., 2006), with detection of OH via LIF. A 355 nm photolysis laser
167 (Spectron Laser Systems, SL803) with a pulse repetition frequency (PRF) of 10 Hz and pulse
168 duration ~ 10 ns was used to fragment HONO into OH. This fragmentation wavelength was
169 chosen as HONO has a strong absorption peak at ~ 355 nm leading to the breakage of the HO-
170 NO bond to form NO and OH in their electronic ground states (Shan et al., 1989). A Nd:YAG
171 pumped dye probe laser (JDSU Q201-HD, Q-series, Sirah Cobra Stretch) with a PRF of 5000
172 Hz, was used for the detection of OH via the fluorescence assay by gas expansion (FAGE)
173 technique which employs the expansion of gas through a small pinhole into the detection cell.
174 The OH radical was measured using on-resonance detection by LIF via the excitation of the
175 $A^2\Sigma^+ (v' = 0) \leftarrow X^2\Pi_i (v'' = 0) Q_1(2)$ transition at 308 nm (Heard, 2006). A multi-channel plate
176 (MCP) photomultiplier (Photek, MCP 325) equipped with an interference filter at 308 nm (Barr
177 Associates, 308 nm. FWHM – 8 nm, ~50 % transmission) was used to measure the fluorescence
178 signal. A reference OH cell in which a large LIF signal could be generated was utilised to



179 ensure the wavelength of the probe laser remained tuned to the peak of the OH transition at
180 308 nm. OH measurements are taken both before and after each photolysis laser pulse allowing
181 measurement of any OH already present in the gas flow to be determined as a background
182 signal for subtraction. The OH generated from HONO photolysis was measured promptly
183 (~800 ns) after the 355 nm pulse to maximise sensitivity to OH before it was spatially diluted
184 away from the measurement region (Boustead, 2019). Offline measurements, with the probe
185 laser wavelength moved away from the OH transition (by 0.02 nm), were taken to allow the
186 signal generated from detector dark counts and scattered laser light to be measured and
187 subtracted from the online signal. To determine an absolute value of the HONO concentration,
188 [HONO], a calibration was performed, in order to convert from the HONO signal, S_{HONO} , using
189 $S_{\text{HONO}} = C_{\text{HONO}} [\text{HONO}]$, as described fully in (Boustead, 2019). A glass calibration wand was
190 used to produce OH and HO₂ in equal concentrations from the photolysis of water vapour at
191 185 nm:



192 An excess flow of NO was then added to generate HONO which was then detected as OH in
193 the cell. The excess flow of NO (BOC, 99.5 %) ensures rapid and complete conversion of OH
194 and HO₂ to HONO. The concentration of OH and HO₂ produced, and therefore the amount of
195 HONO produced in the wand, is calculated using:

$$[\text{OH}] = [\text{HO}_2] = [\text{H}_2\text{O}] \sigma_{\text{H}_2\text{O}} \phi_{\text{OH}} F t \quad (1)$$

196 where [H₂O] is the concentration of water vapour in the humidified gas flow, $\sigma_{\text{H}_2\text{O}}$ is the
197 absorption cross section of H₂O at 185 nm ($7.14 \times 10^{-20} \text{ cm}^2 \text{ molecule}^{-1}$ (Cantrell et al., 1997),
198 ϕ_{OH} is the quantum yield of OH for the photo-dissociation of H₂O at 185 nm (=1), F is the
199 lamp flux and t is the irradiation time (the product of which is determined using ozone
200 actinometry (Boustead, 2019).

201 A typical value of the calibration factor was $C_{\text{HONO}} = (3.63 \pm 0.51) \times 10^{-9} \text{ counts mW}^{-1}$ for N₂,
202 leading to a calculated limit of detection of 12 ppt for a 50 s averaging period and a signal-to-
203 noise ratio (SNR) of 1 (Boustead, 2019). The typical error in the HONO concentration was
204 15% at 1 σ , determined by the error in the calibration.



205 **2.4 Experimental procedure and data analysis**

206 The experiments were performed with a minimum flow of 6 lpm through the aerosol flow tube
207 giving a Reynolds number of ~ 150 which ensured a laminar flow regime. The HONO signal,
208 converted to an absolute concentration using a calibration factor, was measured over a range
209 of aerosol surface area densities, both in the presence and absence of illumination, and
210 background measurements without aerosols present, were also performed.

211 The HONO signal originates from several sources: the illuminated aerosol surface; the
212 illuminated quartz flow tube walls; dark reactions on aerosol surfaces; dark reactions on the
213 flow tube surface and finally from impurities in the NO₂ (Sigma Aldrich, >99.5 %, freeze pump
214 thawed to further remove any remaining NO or O₂) and N₂ flows (either HONO itself or a
215 species which photolyses at 355 nm to give OH). Of interest here is the HONO production
216 from both dark and illuminated aerosol surfaces which is atmospherically relevant. Following
217 transit through the flow tube, and in the presence of NO₂, the total concentration of HONO
218 measured by the PF-LIF detector is given by:

$$\begin{aligned} [\text{HONO}] = & [\text{HONO}]_{\text{illuminated aerosols}} + [\text{HONO}]_{\text{illuminated walls}} \\ & + [\text{HONO}]_{\text{dark aerosols}} + [\text{HONO}]_{\text{dark walls}} + [\text{HONO}]_{\text{impurities}} \end{aligned} \quad (2)$$

219 Any HONO seen without the presence of aerosol was therefore due to HONO impurities in the
220 N₂ or NO₂ gas, the dark production of HONO from the flow tube walls or from the production
221 of HONO from the illuminated reactor walls, which may include production from TiO₂
222 aerosols coating the flow tube in the presence of NO₂. This background HONO concentration
223 depended on the experimental conditions and on how recently the flow tube and PF-LIF cell
224 had been cleaned to remove any build-up of TiO₂ deposits. However, the build-up of TiO₂ on
225 the flow tube walls was relatively slow and back-to-back measurements were made in the
226 presence and absence of aerosols to obtain an accurate background. Even though the aerosol
227 surface area density (~0.02 m² m⁻³) was small compared to the surface area density of the
228 reactor walls (35 m² m⁻³), very little HONO signal was produced without the presence of
229 aerosols, and was always subtracted from the signal in the presence of aerosols. The HONO
230 signal was measured both with the lamp on and off for each aerosol surface area density to
231 investigate the production of HONO from illuminated aerosol surfaces. The HONO signal was
232 averaged over 50 s (average of 500 of the 355 nm photolysis laser pulses with a PRF of 10 Hz).
233 Once aerosols were introduced into the flow tube system a period of ~ 30 min was allowed for
234 equilibration and the measured aerosol surface area density to stabilise. In general, the relative



235 humidity of the system was kept constant at RH ~ 15 % for all experiments investigating
236 HONO production as a function of NO₂ mixing ratio over the range 34 - 400 ppb. In a number
237 of experiments, however, RH was varied in the range ~12-37 %.

238 The mixing ratio of NO₂ entering the flow tube was calculated using the concentration of the
239 NO₂ in the cylinder and the degree of dilution. The NO₂ mixing ratio within the cylinder was
240 determined using a commercial instrument based on UV-Vis absorption spectroscopy (Thermo
241 Fisher 42TL, limit of detection 50 pptv, precision 25 pptv) For each individual experiment, the
242 mixing ratio of NO₂ was kept constant (within the range 34 – 400 ppb) and the aerosol surface
243 area density was varied from zero up to a maximum of 0.04 m² m⁻³. In order to obtain the
244 HONO produced from illuminated aerosol surfaces in the flow tube for a given mixing ratio of
245 NO₂. As well as subtraction of any background HONO, a correction must be made for any loss
246 of HONO owing to its photolysis occurring within the flow tube.

247 In order to determine the rate of photolysis of HONO, the rate of photolysis of NO₂ was first
248 determined using chemical actinometry, and the known spectral output of the lamp and the
249 literature values of the absorption cross-sections and photo-dissociation quantum yields for
250 NO₂ and HONO were used to determine the rate of photolysis of HONO. When just flowing
251 NO₂ in the flow tube, the loss of NO₂ within the illuminated region is determined only by
252 photolysis and is given by:

$$-\frac{d[\text{NO}_2]}{dt} = j(\text{NO}_2)[\text{NO}_2] \quad (3)$$

253 where $j(\text{NO}_2)$ is the photolysis frequency of NO₂ for the lamp used in these experiments. From
254 the measured loss of NO₂ in the illuminated region, and with knowledge of the residence time,
255 the photolysis frequency, $j(\text{NO}_2)$, was determined to be $(6.43 \pm 0.30) \times 10^{-3} \text{ s}^{-1}$ for the set of
256 experiments using one lamp to illuminate the flow tube. $j(\text{NO}_2)$ is given by:

$$j(\text{NO}_2) = \int_{\lambda_1}^{\lambda_2} \sigma_\lambda \phi_\lambda F_\lambda d\lambda \quad (4)$$

257 where λ_1 and λ_2 represent the range of wavelengths over which the lamp emits, and σ_λ and ϕ_λ
258 are the wavelength-dependent absorption-cross section and photo-dissociation quantum yield
259 of NO₂, respectively, and F_λ is the flux of the lamp at a given wavelength. The flux of the lamp,
260 the spectral intensity of which was measured using a Spectral Radiometer (Ocean Optics QE-
261 Pro 500) as a function of wavelength, is shown in Figure 3.

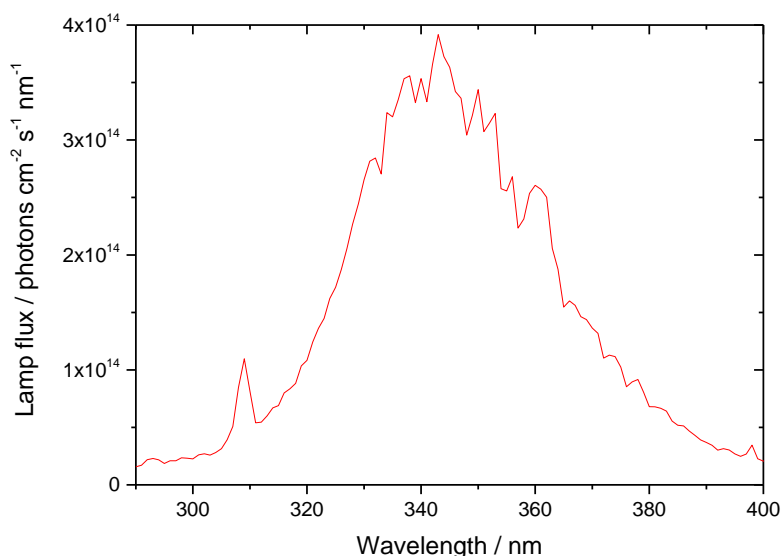


Figure 3. UVA emission spectrum for the 15 W bench lamp used in these experiments between 290-400 nm. The integrated photon flux over this wavelength range is $(1.63 \pm 0.09) \times 10^{16}$ photons $\text{cm}^{-2} \text{s}^{-1}$ determined from the measured $j(\text{NO}_2)$ of $(6.43 \pm 0.30) \times 10^{-3} \text{s}^{-1}$.

262 From the measured $j(\text{NO}_2)$, and with knowledge of σ_λ and ϕ_λ for NO_2 , the flux of the lamp was
263 determined to be $(1.63 \pm 0.09) \times 10^{16}$ photons $\text{cm}^{-2} \text{s}^{-1}$ integrated over the 290 – 400 nm
264 wavelength range of the lamp. Using this flux, and the known σ_λ and ϕ_λ for HONO over the
265 same wavelength range, $j(\text{HONO})$ was determined to be $(1.66 \pm 0.10) \times 10^{-3} \text{s}^{-1}$.

266 In the presence of aerosols under illuminated conditions, the rate of heterogeneous removal of
267 NO_2 at the aerosol surface to generate HONO is given by:

$$-\frac{d[\text{NO}_2]}{dt} = k[\text{NO}_2] \quad (5)$$

268 where k is the pseudo-first order rate coefficient for loss of NO_2 at the aerosol surface, and
269 which leads to the generation of HONO. The postulated mechanism for HONO production
270 from NO_2 is discussed in section 3.3.2 below, but for the definition of k it is assumed to be a
271 first order process for NO_2 . Integration of equation ((5) gives:

$$k = -\frac{\ln\left(\frac{[\text{NO}_2]_0 - [\text{HONO}]_t}{[\text{NO}_2]_0}\right)}{t} \quad (6)$$



272 where $[\text{NO}_2]_0 - [\text{HONO}]_t$ is the concentration of NO_2 at time t , assuming that each NO_2
273 molecule is quantitatively converted to a HONO molecule following surface uptake (see
274 section 3.3.2 for the proposed mechanism), and $[\text{NO}_2]_0$ is the initial concentration of NO_2 .
275 Hence k can be determined from equation ((6) using the measurement of the concentration of
276 HONO, $[\text{HONO}]$, that has been generated from TiO_2 aerosol surfaces for an illumination time
277 of t (and after subtraction of any background HONO produced from other sources and after
278 correction for loss via photolysis, see above), and with knowledge of $[\text{NO}_2]_0$.

279 The reactive uptake coefficient of NO_2 to generate HONO, $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$, defined as the
280 probability that upon collision of NO_2 with the TiO_2 aerosol surface a gas-phase HONO
281 molecule is generated, is given by:

$$\gamma_{\text{NO}_2 \rightarrow \text{HONO}} = \frac{4 \times k}{v \times SA} \quad (7)$$

282 where v is the mean thermal velocity of NO_2 , given by $v = \sqrt{(8RT/(\pi M))}$ with R , T and M as
283 the gas constant, the absolute temperature and the molar mass of NO_2 , respectively, SA is the
284 aerosol surface area density ($\text{m}^2 \text{m}^{-3}$) and k is defined as above. Rearrangement of equation ((7)
285 gives:

$$k = \frac{\gamma_{\text{NO}_2 \rightarrow \text{HONO}} \times SA \times v}{4} \quad (8)$$

286 Figure 4 shows the variation of k , determined from equation ((6) above with $t = 52$ s
287 (illumination time in the flow tube), against aerosol surface area density, SA , for $[\text{NO}_2]_0 = 200$
288 ppb and $\text{RH} = 15\%$, from which the gradient using equation (8) yields $\gamma_{\text{NO}_2 \rightarrow \text{HONO}} = (2.17 \pm$
289 $0.09) \times 10^{-5}$.

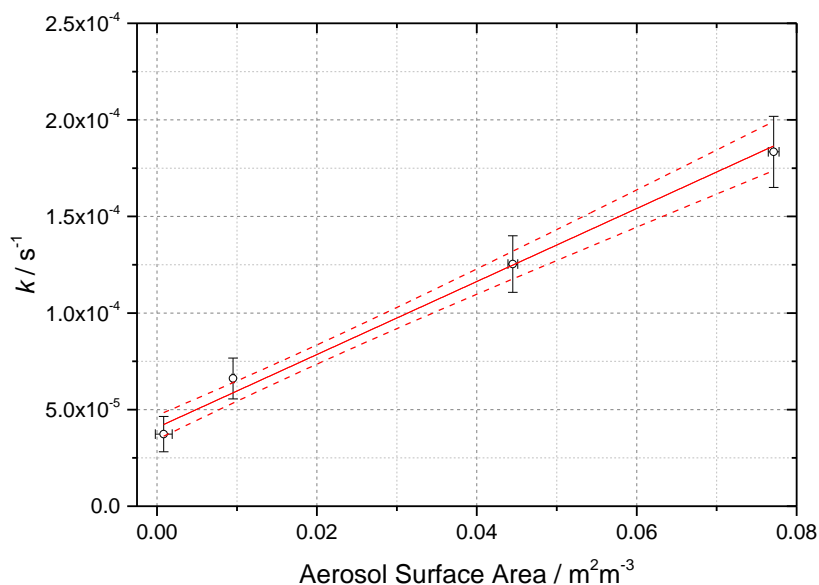


Figure 4. Pseudo-first-order rate coefficient for HONO production, k (open circles) as a function of aerosol surface area for $[\text{NO}_2]=200$ ppb and $\text{RH}=15 \pm 1$ %, $T = 293 \pm 3$ K and a photolysis time of 52 ± 2 seconds. The red line is a linear-least squared fit including 1σ confidence bands (dashed lines) weighted to both x and y errors (1σ), the gradient of which yields $\gamma_{\text{NO}_2 \rightarrow \text{HONO}} = (2.17 \pm 0.09) \times 10^{-5}$, with the uncertainty representing (1σ). The total photon flux of the lamp (see Figure 2 for its spectral output) = $(1.63 \pm 0.09) \times 10^{16}$ photons $\text{cm}^{-2} \text{s}^{-1}$.

290 The uncertainty in k (~ 20 %) shown in Figure 4 and determined by equation ((6) is mainly
291 controlled by the uncertainty in the HONO concentration (the HONO signal typically varies
292 between repeated runs for a given SA by ~ 10 % coupled with the 15 % error in calibration
293 factor), the initial NO_2 mixing ratio (10%), and the photolysis time, t (~ 3 %). The uncertainty
294 in SA is determined by the uncertainty in the SMPS (15%). The error in the value of
295 $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$ (typically 20%) is calculated from the 1σ statistical error of the weighted fit shown
296 in Figure 4. An experiment performed using air yielded an uptake coefficient value within 7 %
297 of the equivalent experiment done in N_2 , which is well within the experimental error.

298 2.5 Box model description

299 A kinetic scheme within the framework of a box model was used together with the differential
300 equation solver Facsimile 4.3.53 (MCPA software Ltd., 2020) to investigate the mechanism of
301 NO_2 adsorption on TiO_2 in the presence of light to produce HONO. The models were only
302 semi-explicit, focusing on determining the stoichiometric amounts of NO_2 needed to produce



303 a single HONO molecule in the gas-phase for comparison with the experimental dependence
304 of HONO production upon NO_2 mixing ratio, and to provide a predictive framework for
305 parameterising the HONO production rate with NO_2 mixing ratio in the atmosphere. Three
306 model scenarios were designed. The simplest model (Model 1) considered only the adsorption
307 of a single molecule of NO_2 to the TiO_2 surface, the surface conversion to HONO in the
308 presence of light and subsequent desorption of HONO, the latter assumed to occur rapidly. The
309 two further model scenarios investigated the effect of a 2:1 stoichiometric relationship between
310 the NO_2 adsorbed to the surface of TiO_2 and the HONO produced, via the formation of an NO_2
311 dimer. Model 2 incorporated an Eley-Rideal mechanism reliant on the adsorption of one NO_2
312 molecule to the surface followed by the subsequent adsorption of a second NO_2 molecule
313 directly onto the first (Figure 5). Model 3, however, features a Langmuir-Hinshelwood
314 mechanism of adsorption in which two NO_2 molecules adsorb to the surface, then diffuse to
315 one another before colliding on the surface and forming the *cis*-ONO- NO_2 dimer (Finlayson-
316 Pitts et al., 2003; de Jesus Madeiros and Pimentel, 2011; Liu and Goddard, 2012; Varner et al.,
317 2014). The formation of the asymmetric *cis*-ONO- NO_2 dimer followed by isomerisation to
318 form the asymmetric *trans*-ONO- NO_2 dimer has been suggested to have an enthalpic barrier
319 that is $\sim 170 \text{ kJ mol}^{-1}$ lower than for direct isomerisation to *trans*-ONO- NO_2 from the symmetric
320 N_2O_4 dimer (Liu and Goddard, 2012). The dimerisation of NO_2 and subsequent isomerisation
321 to form *trans*-ONO- NO_2 has been suggested under dark conditions to lead to the formation of
322 both HONO and HNO_3 in the presence of water vapour (Finlayson-Pitts et al., 2003; de Jesus
323 Madeiros and Pimentel, 2011; Liu and Goddard, 2012; Varner et al., 2014). Although the
324 interaction of light with TiO_2 with the concomitant production of electron-hole pairs (R1) is
325 central to HONO formation, we do not specify here the exact mechanism by which the electron-
326 hole pairs interact with surface-bound species to generate HONO. We propose that the
327 interaction with light speeds up the autoionisation of *trans*-ONO- NO_2 to form $(\text{NO}^+)(\text{NO}_3^-)$,
328 which is represented by reactions R13 and R15 in Models 2 and 3 respectively. $(\text{NO}^+)(\text{NO}_3^-)$
329 can then react rapidly with surface adsorbed water leading to HONO formation (Varner et al.,
330 2014).

331 A schematic of the proposed mechanism investigated with Models 2 and 3 is shown in Figure
332 5, and consists of (i) the adsorption of NO_2 onto a surface site, (ii) the conversion of NO_2 to
333 form HONO via the formation of an NO_2 dimer intermediate on the surface via either a Eley-
334 Rideal or Langmuir Hinshelwood- type mechanism, (iii) subsequent desorption of HONO from
335 the surface, and finally (iv) competitive removal processes for HONO both on the surface and



336 in the gas-phase that are either dependent or independent on the NO₂ mixing ratio. The model
 337 includes the gas-phase photolysis of NO₂ and HONO and the gas phase reactions of both
 338 HONO and NO₂ with OH and O(³P) atoms.

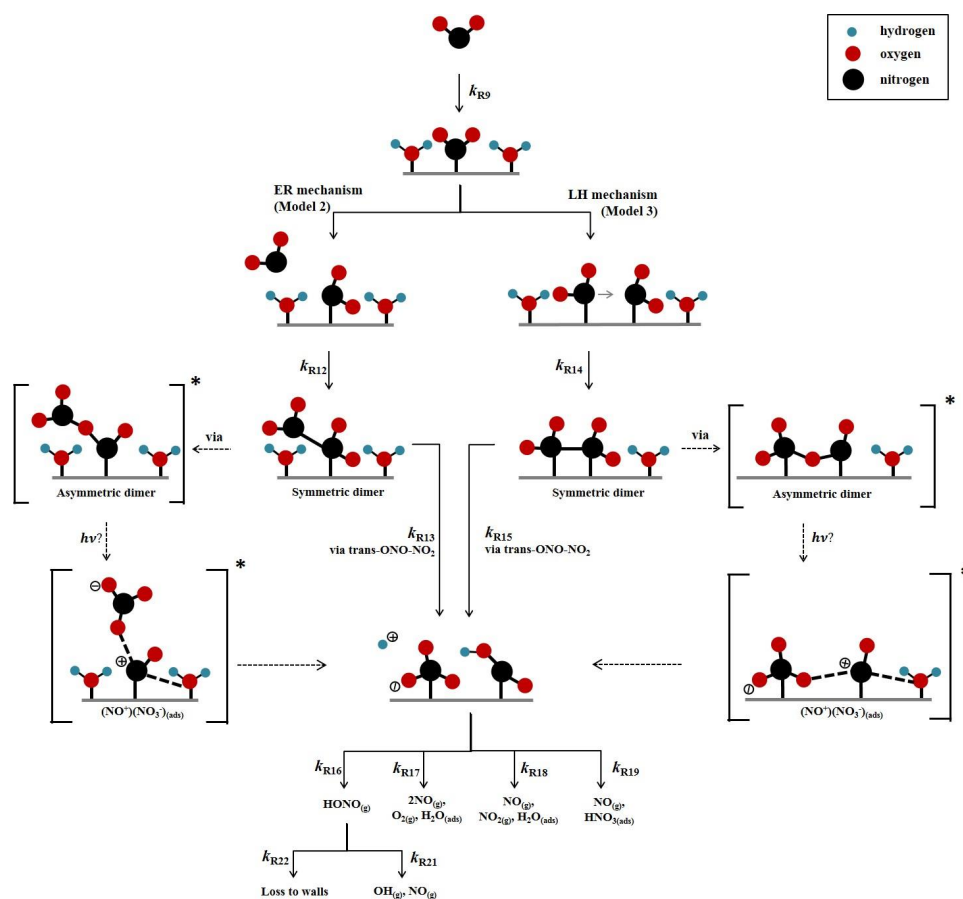


Figure 5. Schematic diagram of proposed mechanism of uptake of NO₂ on an aerosol surface in the presence of water to form HONO. Both Eley Rideal, Model 2, and Langmuir Hinshelwood, Model 3, mechanisms are shown with relevant estimated and calculated rate coefficients used in the models. NO₂ dependent and independent loss reactions of HONO are also depicted. Nitrogen shown in black, oxygen shown in red and hydrogen shown in blue. * denotes intermediate steps of the isomerisation of symmetric N₂O₄ to *trans*-ONO-NO₂ which is then predicted to form HONO.

339 To the best of our knowledge the enthalpy of adsorption of NO₂ onto a TiO₂ surface has not
 340 been determined, nor the bimolecular rate coefficients for the chemical steps on the surface
 341 shown in Figure 5. Hence, for each of the steps a rate coefficient (s⁻¹ or cm³ molecule⁻¹ s⁻¹) was
 342 assigned, as given in Table 1, and with the exception of the experimentally determined *j*(NO₂)



343 and the calculated $j(\text{HONO})$, and the gas-phase rate coefficients which are known, the rate
 344 coefficients were estimated, with the aim of reproducing the experimental NO_2 dependence of
 345 the HONO production and NO_2 reactive uptake coefficient; justification of chosen values is
 346 given below.

Reactions	Rate coefficient ^d
Model 1	
R9 $\text{NO}_{2(g)} + \text{surface} \rightarrow \text{NO}_{2(ads)}$	1×10^{-3}
R10 $\text{NO}_{2(ads)} \rightarrow \text{HONO}_{(ads)}$	1×10^{-3}
R11 $\text{HONO}_{(ads)} \rightarrow \text{HONO}_{(g)}$	1×10^{-2}
Model 2 and 3	
<i>Model 2 only – Eley-Rideal mechanism</i>	
R12 $\text{NO}_{2(g)} + \text{NO}_{2(ads)} \rightarrow \text{NO}_2 - \text{NO}_{2(ads)}$	1×10^{-2}
R13 $\text{NO}_2 - \text{NO}_{2(ads)} \xrightarrow{\text{via trans-ONO-NO}_2} \text{HONO}_{(ads)} + \text{HNO}_{3(ads)}$	5×10^{-3}
<i>Model 3 only – Langmuir-Hinshelwood mechanism</i>	
R14 $\text{NO}_{2(ads)} + \text{NO}_{2(ads)} \rightarrow \text{NO}_{2(ads)} - \text{NO}_{2(ads)}$	1×10^{-3}
R15 $\text{NO}_{2(ads)} - \text{NO}_{2(ads)} \xrightarrow{\text{via trans-ONO-NO}_2} \text{HONO}_{(ads)} + \text{HNO}_{3(ads)}$	5×10^{-3}
<i>Common to both Models 2 and 3</i>	
R9 $\text{NO}_{2(g)} + \text{surface} \rightarrow \text{NO}_{2(ads)}$	1×10^{-1}
R16 $\text{HONO}_{(ads)} \rightarrow \text{HONO}_{(g)}$	5×10^{-2}
R17 $\text{HNO}_{3(ads)} + \text{HONO}_{(ads)} \rightarrow 2\text{NO}_{(g)} + \text{O}_{2(g)} + \text{H}_2\text{O}_{(ads)}$	1×10^{-3}
R18 $\text{HONO}_{(ads)} + \text{HONO}_{(ads)} \rightarrow \text{NO}_{(g)} + \text{NO}_{2(g)} + \text{H}_2\text{O}_{(ads)}$	1×10^{-3}
R19 $\text{NO}_{2(g)} (\text{or species such as } \text{NO}_2^+) + \text{HONO}_{(ads)} \rightarrow \text{NO}_{(g)} + \text{HNO}_{3(ads)}$	5×10^{-3}
R20 $\text{NO}_{2(g)} + h\nu \rightarrow \text{NO}_{(g)} + \text{O}({}^3\text{P})_{(g)}$	6×10^{-3a}
R21 $\text{HONO}_{(g)} + h\nu \rightarrow \text{OH}_{(g)} + \text{NO}_{(g)}$	2×10^{-3b}
R22 $\text{HONO}_{(g)} \rightarrow \text{wall loss}$	1×10^{-4}
R23 $\text{HONO}_{(g)} + \text{OH}_{(g)} \rightarrow \text{NO}_{2(g)} + \text{H}_2\text{O}_{(g)}$	4.5×10^{-12c}
R24 $\text{NO}_{2(g)} + \text{OH}_{(g)} \xrightarrow{M} \text{HNO}_{3(g)}$	1×10^{-11c}
R25 $\text{O}({}^3\text{P})_{(g)} + \text{NO}_{2(g)} \rightarrow \text{O}_{2(g)} + \text{NO}_{(g)}$	1×10^{-11c}
R26 $\text{O}({}^3\text{P})_{(g)} + \text{O}_{2(g)} \xrightarrow{M} \text{O}_3$	1.5×10^{-14c}
R27 $\text{O}({}^3\text{P})_{(g)} + \text{NO}_{(g)} \xrightarrow{M} \text{NO}_{2(g)}$	1.7×10^{-12c}

347 **Table 1.** Reactions included in the chemical mechanism used to model NO_2 uptake onto TiO_2 aerosols. All rate
 348 coefficients are estimated, as described in Section 2.5, with the exception of the NO_2 and HONO photolysis rate
 349 coefficient and the gas phase rate coefficient which are known. ^aMeasured using chemical actinometry with the
 350 knowledge of the experimentally determined spectral output of the lamp and the cross-sections and quantum
 351 yields of NO_2 and HONO, see section 2.4 for more detail. ^bCalculated using a photon flux of $(1.63 \pm 0.09) \times$



352 10^{16} photons $\text{cm}^{-2} \text{s}^{-1}$.^c(Sander et al., 2003). ^dRate coefficients are in the units of s^{-1} for first-order processes or
353 $\text{cm}^3 \text{molecule}^{-1} \text{s}^{-1}$ for second-order processes. T for all k values is 298 K.

354 The modelled Gibbs free energy barrier for the isomerisation of N_2O_4 to form the asymmetric
355 ONO-NO_2 isomer (*cis* or *trans* conformation not specified) was estimated by Pimental et al.,
356 (2007) to be 87 kJ mol^{-1} with a rate coefficient as large as $2 \times 10^{-3} \text{ s}^{-1}$ in the aqueous phase at
357 298 K, stated in the study to confirm the Finlayson-Pitts model for the hydrolysis of NO_2 on
358 surfaces via the asymmetric *trans-ONO-NO}_2* dimer (Finlayson-Pitts et al., 2003). Using this
359 study as a guide, we estimated k_{R13} and k_{R15} as $5 \times 10^{-3} \text{ s}^{-1}$, slightly larger than that estimated
360 by Pimental et al., (2007) due to the presence of light. A study into the decomposition of HONO
361 on borosilicate glass surfaces suggested a rate coefficient for the loss HONO on the non-
362 conditioned chamber walls to be $(1.0 \pm 0.2) \times 10^{-4} \text{ s}^{-1}$ increasing to $(3.9 \pm 1.1) \times 10^{-4} \text{ s}^{-1}$ when
363 HNO_3 was present on the walls (Syomin and Finlayson-Pitts, 2003). From this we estimated a
364 light-accelerated loss rate coefficient of $1 \times 10^{-3} \text{ s}^{-1}$ for the loss of $\text{HONO}_{(\text{ads})}$ by reaction with
365 itself, k_{R18} , and through reaction with $\text{HNO}_{3(\text{ads})}$, k_{R17} . Both these reactions will occur on the
366 surface of the aerosol. We make the assumption that the rate of loss of HONO to the walls of
367 the chamber for this experiment is less than that of the heterogeneous loss reactions on the
368 photo-catalytic aerosol surface leading to a k_{R22} of $1 \times 10^{-4} \text{ s}^{-1}$ as reported by (Syomin and
369 Finlayson-Pitts, 2003). For $k_{\text{R12}}-k_{\text{R15}}$, initial values were adopted and were then adjusted to fit
370 the shape of the trend in experimental results of $[\text{HONO}]$ and $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$ versus $[\text{NO}_2]$,
371 discussed fully in Section 3.3.2. For completeness, gas-phase loss reactions of HONO and NO_2
372 with OH and the reactions of $\text{O}(^3\text{P})$ with NO, NO_2 and O_2 were also included in the model,
373 R23-R27, though their inclusion had no effect on the HONO concentration. The rates of R23-
374 R27 within the model are much smaller than HONO loss reactions on the surface (R17-R19)
375 and the photolysis reactions (R21). For both Models 2 and 3, the adsorption of an NO_2 molecule
376 to the surface, k_{R9} , was assumed to be rapid and not the rate determining step. Likewise, the
377 desorption of HONO was also assumed to be rapid, faster than the loss rates of adsorbed HONO
378 but slower than the adsorption of NO_2 ; this was necessary for the model to reproduce the trend
379 in the experimental results of $[\text{HONO}]$ versus $[\text{NO}_2]$, discussed fully in Section 3.3.2.

380 **3 Results and Discussion**

381 **3.1 HONO production from TiO_2 aerosol surfaces in the presence of NO_2**

382 The production of HONO on TiO_2 aerosol surfaces was measured as a function of the initial
383 NO_2 mixing ratio. Figure 6 shows the dependence of the HONO concentration, measured at



384 the end of the flow tube, on the initial NO₂ mixing ratio for an aerosol surface area of $(1.6 \pm$
385 $0.8) \times 10^{-2} \text{ m}^2 \text{ m}^{-3}$. A sharp increase in HONO production at a low mixing ratio of NO₂ was
386 seen followed by a more gradual reduction in HONO production after a peak production at \sim
387 $54 \pm 5 \text{ ppb NO}_2$.

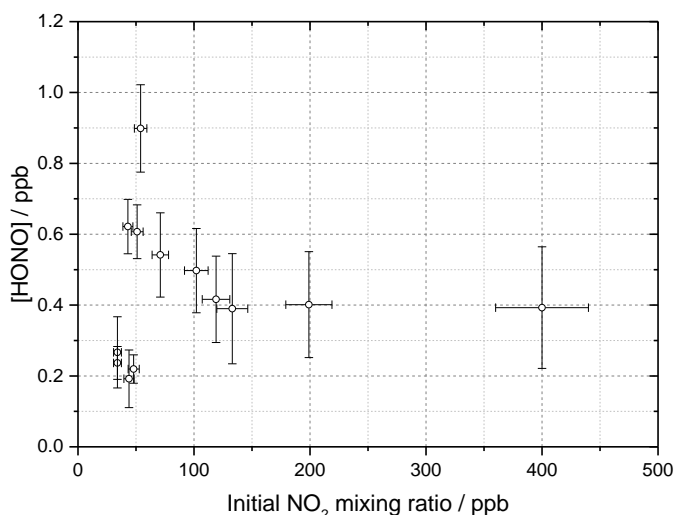


Figure 6. HONO concentration measured at the end of the flow tube as a function of the initial NO₂ mixing ratio, for the aerosol surface area density of $(1.6 \pm 0.8) \times 10^{-2} \text{ m}^2 \text{ m}^{-3}$, relative humidity $15 \pm 1 \%$, photon flux $(1.63 \pm 0.09) \times 10^{16} \text{ photons cm}^{-2} \text{ s}^{-1}$ (290–400 nm wavelength range), reaction time of 52 seconds and N₂ carrier gas. Each point is an average of up to 20 measurements at the same aerosol surface area and mixing ratio of NO₂. The highest concentration of HONO measured was $0.90 \pm 0.12 \text{ ppb}$ at $[\text{NO}_2] = 54 \pm 5 \text{ ppb}$. The *y* error bars represent 1σ while the *x* error bars represent the sum in quadrature of the errors in the N₂ and NO₂ gas flows and the NO₂ dilution. The SA varied over the experiments at different NO₂ mixing ratios leading to a larger error in the quoted SA.

388 Figure 7 shows the HONO concentration measured at the end of the flow tube over a range of
389 RH values for a fixed aerosol surface area density of $(1.59 \pm 0.16) \times 10^{-2} \text{ m}^2 \text{ m}^{-3}$ and at two
390 NO₂ mixing ratios, displaying a peak in HONO production between 25 – 30 % RH. Above \sim
391 37 % RH, for experiments including single-component TiO₂ aerosols, it was found that
392 significant aerosols were lost from the system before entering the flow tube, speculated to be
393 due to loss to the walls of the Teflon lines. As such the RH dependence was only studied up to



394 37 % RH, however a clear drop off in HONO production was seen for both NO₂ mixing ratios
 395 studied after ~ 30 % RH.

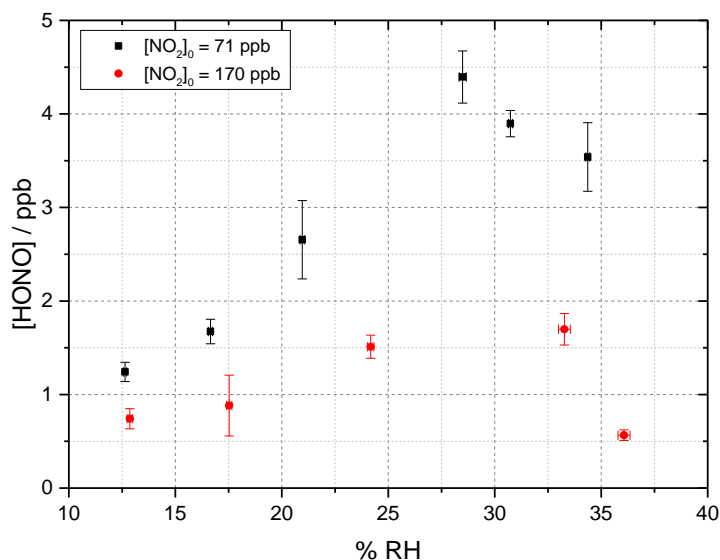


Figure 7. RH dependence of HONO production from illuminated TiO₂ aerosol surfaces at 295 K in N₂ at 71 (black) and 170 (red) ppb initial NO₂ mixing ratio. The aerosol surface area density was kept constant at $(1.59 \pm 0.16) \times 10^{-2} \text{ m}^2 \text{ m}^{-3}$ with a photon flux of $(1.63 \pm 0.09) \times 10^{16} \text{ photons cm}^{-2} \text{ s}^{-1}$ and an illumination time of 52 ± 2 seconds. The error bars represent 1σ .

396 A dependence of HONO production upon RH was expected due to the potential role of water
 397 as a proton donor in the production mechanism of HONO on TiO₂ surfaces (R2 and (R5, as
 398 shown in Figure 5 (Dupart et al., 2014). The fractional surface coverage of water on the TiO₂
 399 aerosol core, V/V_m , at 15 % RH and above was calculated using the parameterisation below,
 400 which was determined using transmission IR spectroscopy (Goodman et al., 2001):

$$\frac{V}{V_m} = \left[\frac{c \left(\frac{P}{P_0}\right)}{1 - \left(\frac{P}{P_0}\right)} \right] \left[\frac{1 - (n+1) \left(\frac{P}{P_0}\right)^n + n \left(\frac{P}{P_0}\right)^{n+1}}{1 + (c-1) \left(\frac{P}{P_0}\right) - c \left(\frac{P}{P_0}\right)^{n+1}} \right] \quad (9)$$

401 where V is the volume of water vapour adsorbed at equilibrium pressure P , V_m is the volume
 402 of gas necessary to cover the surface of TiO₂ particles with a complete monolayer, P_0 is the
 403 saturation vapour pressure, c is the temperature dependent constant related to the enthalpies of



404 adsorption of the first and higher layers (taken as 74.8 kJ mol^{-1} for TiO_2 (Goodman et al., 2001))
405 and n is the asymptotic limit of monolayers (8 for TiO_2 (Goodman et al., 2001)) at large values
406 of P/P_0 .

407 At 15 % RH, a fractional water coverage of 1.09 was calculated to be present on the surface,
408 increasing to 1.50 at 35 % RH. It has been shown in previous work that HONO can be displaced
409 from a surface by water, leading to an increase in gas-phase HONO with RH (Syomin and
410 Finlayson-Pitts, 2003). The increase in HONO with RH to ~25-30 % RH could therefore be
411 attributed to both an increase in the concentration of the water reactant leading to more HONO
412 formation and the increase in displacement of HONO from the surface due to preferential
413 adsorption of water. A decrease in HONO production seems to occur above ~ 30 % RH, which
414 could be due to the increased water adsorption inhibiting either NO_2 adsorption or the
415 electron/hole transfer process (Gustafsson et al., 2006). H_2O vapour adsorption is likely
416 enhanced by the superhydrophilic properties of TiO_2 surfaces under UV radiation meaning that
417 water monolayers form more quickly on the surface of TiO_2 owing to light-induced changes in
418 surface tension (Takeuchi et al., 2005; Gustafsson et al., 2006).

419 At the higher initial concentration of $\text{NO}_2 = 170 \text{ ppb}$, the RH dependence showed a similar
420 peak in HONO production between ~25 - 30 % RH but less HONO was produced overall, as
421 expected from Figure 6 given the higher NO_2 . Previous work on the production of HONO from
422 suspended TiO_2 aerosols reported a strong RH dependence of the uptake coefficient, γ , of NO_2
423 to form HONO with a peak at ~ 15 % RH and decreasing at larger RH (Gustafsson et al., 2006).
424 The same trend for the NO_2 uptake coefficient was observed by Dupart et al., 2014 on Arizona
425 test dust (ATD) aerosols with a peak in γ at ~ 25 % RH. This increase in the RH at which the
426 uptake coefficient for NO_2 in going from TiO_2 to ATD aerosols was ascribed to the lower
427 concentration of TiO_2 present in ATD aerosols as opposed to single-component TiO_2 aerosols
428 used by Gustafsson et al., 2006 as well as by differences in particle size distribution. Gustafsson
429 et al., 2006 reported a larger aerosol size distribution with a bimodal trend with mode diameters
430 of ~ 80 and ~ 350 nm for single-component TiO_2 aerosols whereas Dupart et al., 2014 reported
431 a smaller unimodal aerosol size distribution for ATD aerosols with a mode diameter of ~110
432 nm. In this work we also see a larger aerosol size distribution, with a lower mode diameter of
433 ~ 180 nm similar to Dupart et al., 2014 but for pure TiO_2 aerosols; aerosol size distribution
434 shown in Figure 2. Similar to the results of Dupart et al., 2014 we observe a trend inversion in
435 [HONO] vs RH at higher RH, between 25-30 %.



436 3.2 Dependence of reactive uptake coefficient on initial NO₂ mixing ratio

437 The reactive uptake coefficient, $\gamma_{NO_2 \rightarrow HONO}$ for NO₂→HONO on TiO₂ aerosol particles was
438 determined experimentally for 18 different initial NO₂ mixing ratios, and is shown in Figure 8.
439 For each initial NO₂ mixing ratio, the gradient of the first order rate coefficient for HONO
440 production, k , as a function of aerosol surface area density (e.g. Figure 4) and in conjunction
441 with equation ((8)), was used to obtain $\gamma_{NO_2 \rightarrow HONO}$. The uptake coefficient initially increases
442 with NO₂, reaching a peak at $\gamma_{NO_2 \rightarrow HONO} = (1.26 \pm 0.17) \times 10^{-4}$ for an initial NO₂ mixing ratio
443 of 51 ± 5 ppb, before sharply decreasing as the NO₂ mixing ratio continues to increase above
444 this value.

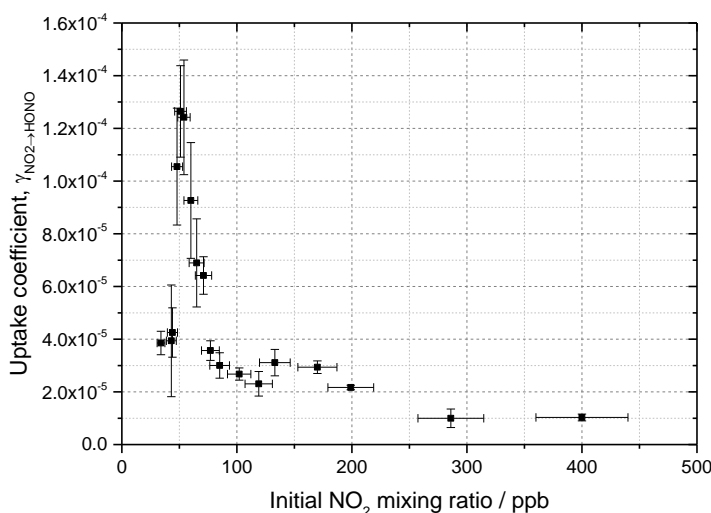


Figure 8 Experimental results showing the reactive uptake coefficients of NO₂ to form HONO, $\gamma_{HONO \rightarrow NO_2}$ onto TiO₂ aerosol surfaces as a function of the initial NO₂ mixing ratio. All experiments were conducted in N₂ at 295 K at 15 ± 1 % RH, a photon flux of $(1.63 \pm 0.09) \times 10^{16}$ photons cm⁻² s⁻¹ and an illumination time of 52 ± 2 seconds. $\gamma_{HONO \rightarrow NO_2}$ was determined for each NO₂ mixing ratio from the gradient of the pseudo-first-order rate coefficient for HONO production, k , versus aerosol surface area density varied from 0 - 0.04 m² m⁻³ (e.g. as shown in **Figure 4**) and Equation ((8)).

445 The increase in uptake coefficient with NO₂ at low NO₂ (< 51 ppb) has not been seen previously
446 in studies of HONO production from TiO₂ containing aerosols with similar [NO₂] ranges
447 (Gustafsson et al., 2006; Ndour et al., 2008; Dupart et al., 2014) or other aerosol surfaces
448 (Bröske et al., 2003; Stemmler et al., 2007). It is worth noting that several of these studies



449 reported the overall uptake of NO₂ onto aerosol surfaces and not specifically the uptake to form
450 HONO, although HONO was indirectly measured in all studies noted here (Gustafsson et al.,
451 2006;Ndour et al., 2008;Dupart et al., 2014). For single-component TiO₂ aerosols, Gustafsson
452 et al., (2006) reported a uptake coefficient, γ_{NO_2} , of 9.6×10^{-4} at 15 % RH and 100 ppb NO₂.
453 Taking into account the HONO yield of 0.75 given by (Gustafsson et al., 2006), an estimated
454 $\gamma_{NO_2 \rightarrow HONO} = 7.2 \times 10^{-4}$ is determined and can be compared to the value observed in this work
455 at 15 % RH and 100 ppb NO₂, ($\gamma_{NO_2 \rightarrow HONO} = (2.68 \pm 0.23) \times 10^{-5}$). The $\gamma_{NO_2 \rightarrow HONO}$ we
456 determine is 27 times smaller than reported by Gustafsson et al., (2006). This difference is
457 mostly due to the lower experimental photon flux in our setup, ~19 times less at $\lambda_{max} = 365$
458 nm owing to the use of one 15 W UV lamp to irradiate the flow tube (Boustead, 2019)
459 compared to Gustafsson et al., 2006 which utilised four 18 W UV lamps.

460 The origins of the increase in $\gamma_{NO_2 \rightarrow HONO}$, together with reaching a maximum and the
461 subsequent decrease at larger NO₂ mixing ratios was investigated using the kinetic box model
462 and postulated mechanism for HONO production described in Section 2.5. The aim was to
463 compare the observed production of HONO and $\gamma_{NO_2 \rightarrow HONO}$ with the modelled values, as a
464 function of NO₂ mixing ratio. The skill of the model to reproduce the observed behaviour
465 enables a validation of the postulated mechanism for HONO production, and variation of the
466 kinetic parameters enables the controlling influence of different steps in the mechanism on
467 HONO production to be evaluated.

468 **3.3 Modelling the HONO production mechanism on illuminated TiO₂** 469 **aerosol surfaces**

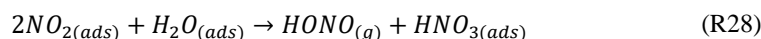
470 The HONO production on illuminated TiO₂ aerosol surfaces was investigated for each of the
471 mechanisms outlined in Table 1.

472 **3.3.1 Model 1**

473 Model 1 (see Table 1 and Figure 5), which contains the simplest mechanism, was designed to
474 reproduce the decreasing value of the NO₂ uptake coefficient to form HONO, $\gamma_{NO_2 \rightarrow HONO}$, with
475 increasing NO₂ and also the plateauing at higher NO₂ mixing ratios caused by NO₂ reaching a
476 maximum surface coverage, as seen by Stemmler et al., (2007). A decrease in the uptake
477 coefficient of NO₂, γ_{NO_2} , onto dust aerosol surfaces was also seen in studies where the
478 formation of HONO from NO₂ uptake was not directly studied (Ndour et al., 2008;Dupart et
479 al., 2014). The mechanism for Model 1 which is given in Table 1 describes the adsorption of



480 one NO₂ molecule to a surface site which then undergoes the reaction which forms HONO,
481 followed by desorption of HONO to the gas-phase, R9-R11. Any representation of the specific
482 chemical processes which convert NO₂ to HONO on the surface following the initial photo-
483 production of electron-hole pairs in the TiO₂ structure (R2) was not included here as the
484 primary focus was to produce the relationship between $\gamma_{NO_2 \rightarrow HONO}$ and the NO₂ mixing ratio.
485 Gustafsson et al., (2006) reported that the measured rate of photo-induced HONO production
486 is 75% that of the rate of NO₂ removal, whereas the dark disproportionation reaction (R28)
487 would predict a 50% yield, and hence that the HONO observed in their studies is not simply a
488 photo-enhancement of:



489 Gustafsson et al., (2006) suggests that an oxidant on the surface is produced following the
490 creation of the electron-hole pair (OH is generated in (R2)), and suggests H₂O₂ as a possibility,
491 which is consistent with the observation of OH and HO₂ radicals produced from the surface of
492 illuminated TiO₂ aerosols (Moon et al., 2019). For Model 1, outputs for the predicted
493 concentration of HONO and the reactive uptake coefficient, $\gamma_{NO_2 \rightarrow HONO}$, as a function of initial
494 NO₂ mixing ratio are shown in Figure 9.

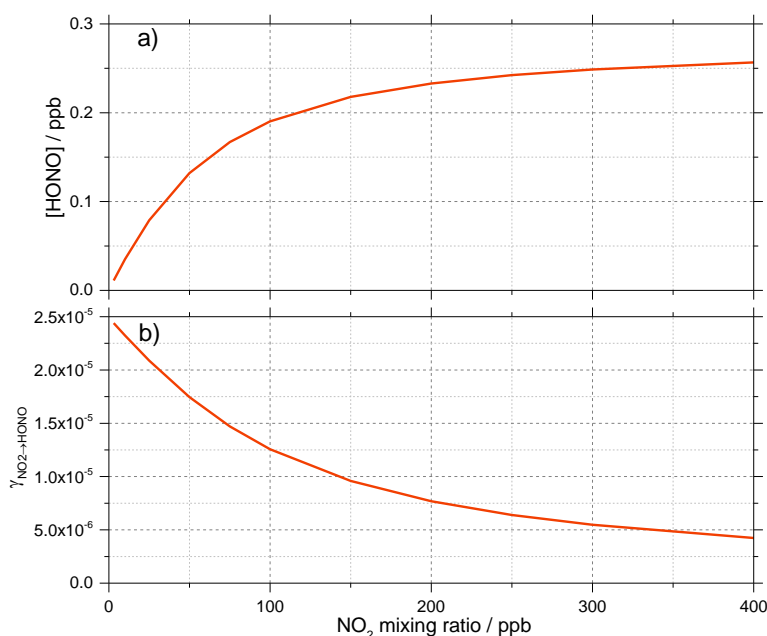




Figure 9 Model 1 calculations for (a) the concentration of HONO and (b) the reactive uptake coefficient to form HONO, $\gamma_{NO_2 \rightarrow HONO}$, as a function of NO_2 mixing ratio for a model run time of 52 s. The estimated rate coefficients used in this model are shown in Table 1.

495 For a run time of 52 s, equal to that of the experimental illumination time, Model 1 predicts an
496 increase in HONO production with increasing NO_2 mixing ratio until the HONO concentration
497 begins to plateau, reaching ~ 0.25 ppb at $[NO_2] = 400$ ppb, presumably owing to saturation on
498 active aerosol surface sites by NO_2 . This leads to the modelled reactive uptake coefficient,
499 $\gamma_{NO_2 \rightarrow HONO}$, monotonically decreasing with increasing NO_2 mixing ratio; a variation in NO_2
500 uptake coefficient similar to that seen in previous photo-enhanced NO_2 aerosol uptake studies
501 (Bröske et al., 2003; Stemmler et al., 2007; Ndour et al., 2008; Dupart et al., 2014). However,
502 the model predictions for Model 1 do not reproduce the experimental variations shown in
503 Figure 6 and Figure 8, in which there is an observed initial rise and then a fall in both the
504 HONO concentration and $\gamma_{NO_2 \rightarrow HONO}$ with increasing NO_2 mixing ratio. Hence, additional
505 processes were considered in the model in order to try to reproduce this behaviour.

506 **3.3.2 Models 2 and 3. Investigating the role of NO_2 dimerisation for the surface** 507 **formation of HONO, and including additional surface losses of HONO**

508 As the experimental $\gamma_{NO_2 \rightarrow HONO}$ increases with NO_2 at low NO_2 (Figure 8), we postulate in
509 Models 2 and 3 that the production of HONO under illuminated conditions is not fully first
510 order in NO_2 and requires more than one NO_2 molecule to form HONO, consistent with the
511 formation of the symmetric NO_2 dimer (N_2O_4) followed by isomerisation on the surface to
512 form the asymmetric *trans*-ONO- NO_2 dimer, which has been suggested to be more reactive
513 with water than the symmetric N_2O_4 dimer (Finlayson-Pitts et al., 2003; Ramazan et al.,
514 2004; Ramazan et al., 2006; Liu and Goddard, 2012) due to the autoionisation to form
515 $(NO^+)(NO_3^-)$ which we propose is accelerated by the presence of light; the full mechanism for
516 which is shown in Figure 5. A recent rotational spectroscopy study found that the *trans*-ONO-
517 NO_2 was better described as the ion pair $(NO^+)(NO_3^-)$ (Seifert et al., 2017). Reaction of the
518 $(NO^+)(NO_3^-)$ ion pair with surface adsorbed water can then lead to the formation of HONO and
519 HNO_3 , the feasibility of which is supported by molecular dynamics simulation studies (Varner
520 et al., 2014). While the symmetric N_2O_4 dimer is favoured as it is the most stable conformer,
521 the asymmetric forms have been experimentally observed in several studies (Fateley et al.,
522 1959; Givan and Loewenschuss, 1989b, a, 1991; Pinnick et al., 1992; Forney et al., 1993; Wang
523 and Koel, 1998, 1999; Beckers et al., 2010). A more recent *ab initio* study of NO_2 adsorption



524 at the air-water interface suggested an orientational preference of NO_2 on the surface, with both
525 oxygen atoms facing away from the interface which may imply that the asymmetric dimer
526 ONO-NO_2 can form directly, meaning the high barrier between the symmetric and asymmetric
527 forms does not need to be overcome (Murdachew et al., 2013).

528 The energy barrier to isomerisation of symmetric N_2O_4 in the gas-phase may be reduced due
529 to the interaction with water adsorbed on surfaces. We therefore rule out the dimer in the gas-
530 phase adsorbing onto the surface first, and then reacting to form HONO (Varner et al., 2014).
531 An interesting question is whether the first NO_2 molecule adsorbed to the surface dimerises via
532 the addition of a gaseous NO_2 via an Eley-Rideal (ER) type process, or whether a Langmuir-
533 Hinshelwood (LH) type mechanism is operating in which both NO_2 molecules are first
534 adsorbed and then diffuse together on the surface forming N_2O_4 . Both ER and LH mechanisms
535 to form the NO_2 dimer have been included in the model, denoted as Model 2 and Model 3,
536 respectively. The outputs for Models 2 and 3 (see Table 1 for details of the processes included)
537 for the HONO concentration and $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$ as a function of NO_2 are shown in Figure 10
538 together with the experimental data. The stoichiometric relationship of the requirement of two
539 NO_2 molecules forming HONO on the surface was key to reproducing the experimental trend
540 of first an increase and then a decrease in both the HONO concentration and the reactive uptake
541 coefficient with the initial NO_2 mixing ratio.

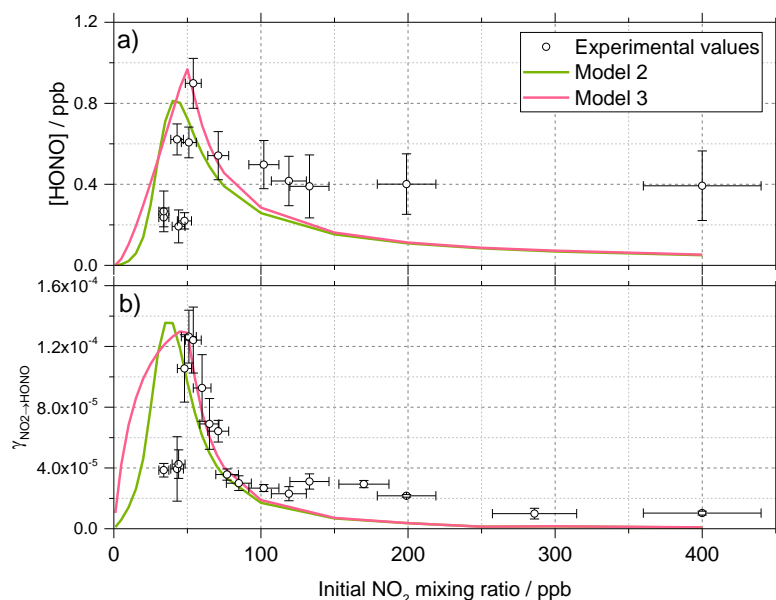
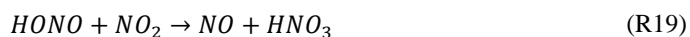
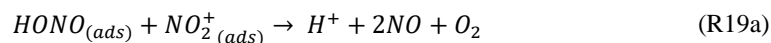


Figure 10. Experimental values (open circles with 1σ error bars), Model 2 (green line) and Model 3 (pink line) calculations for (a) HONO concentration after 52 s illumination and (b) NO_2 reactive uptake coefficient, $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$, as a function of the initial NO_2 mixing ratio. The mechanisms used for these model runs included a 2:1 stoichiometric relationship between the NO_2 adsorbed on the TiO_2 aerosol surface and the HONO produced, as well as additional HONO loss reactions which are dependent on NO_2 , see Table 1 for details. Models 2 and use an Eley-Rideal and Langmuir-Hinshelwood mechanisms, respectively, for the formation of the NO_2 dimer on the aerosol surface.

542 In previous work that investigated HONO production from humic acid aerosols, a saturation
 543 effect was seen with HONO production plateauing with increasing NO_2 mixing ratio (Stemmler
 544 et al., 2007), with the decreasing uptake coefficient, $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$, with increasing NO_2 being
 545 attributed to NO_2 fully saturating available surface sites. However, the observed decrease of
 546 $[\text{HONO}]$ at the high NO_2 mixing ratios shown in Figure 8 and Figure 10a suggests that
 547 additional reactions on the surface may remove HONO and result in the reduction of $[\text{HONO}]$
 548 that is measured. As $[\text{HONO}]$ decreases with the increase in the NO_2 mixing ratio, the removal
 549 process should either involve NO_2 directly:

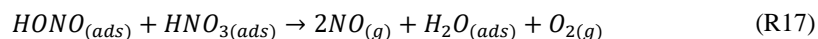


550 or involve species made rapidly from NO_2 on the surface, such as NO_2^+ :



551 which may be present at high enough concentrations of HNO₃ on the surface (Syomin and
552 Finlayson-Pitts, 2003) or following reaction with $h\nu_B^+$, or a product of the reaction of
553 O_2^- (or e_{CB}^-) with NO₂ (R4) i.e. NO₂⁻. Transition state theory (TST) studies of the gas-phase
554 reaction of HONO with NO₂ to form HNO₃ calculated a large activation energy which varied
555 depending on whether the reaction occurs via O abstraction by HONO (159 kJ mol⁻¹) or via
556 OH abstraction via NO₂ (~133-246 kJ mol⁻¹) (Lu et al., 2000). In the gas-phase these reactions
557 are too slow to be important but they could be enhanced on the surface, potentially more so on
558 a photoactive surface such as TiO₂. The NO₂ dependent loss reaction, k_{R19} in Table 1, was
559 necessary in the model to reproduce the sharp decrease in [HONO] versus NO₂ seen
560 experimentally after ~54 ppb NO₂. Without k_{R19} the modelled [HONO] continued to increase
561 to a plateau, as seen in Model 1 (see Figure 9). In order to observe the model output seen in
562 Figure 10 for model 2 and 3, k_{R19} also had to be slower than the desorption of HONO from the
563 surface, k_{R16} .

564 The addition of an NO₂ dependent loss reaction to both Model 2 and 3 had the most significant
565 effect on the trend in modelled HONO concentration. Though it is also possible that a
566 secondary product could remain adsorbed and therefore block active sites on the TiO₂ surface,
567 effectively poisoning the photo-catalyst, NO₂ independent loss reactions in the model, k_{R17} and
568 k_{R18} had little effect on the trend in [HONO] vs NO₂, only having an effect on the overall
569 [HONO]. HNO₃ has however been shown to remain adsorbed to surfaces once formed
570 (Sakamaki et al., 1983; Pitts et al., 1984; Finlayson-Pitts et al., 2003; Ramazan et al., 2004) and
571 may also react with adsorbed HONO, further reducing the product yield (Finlayson-Pitts et al.,
572 2003): these NO₂ independent loss reactions may therefore become more important at higher
573 NO₂ concentrations and hence surface concentrations of HONO and HNO₃:



574 The photolysis of particulate nitrate was not considered in Models 2 or 3, due to the lack of
575 particulate nitrate in the system at $t=0$. The gas-to-particle conversion of any HNO₃ formed
576 was not considered to be important due to the assumption that most HNO₃ formed would
577 remain adsorbed to the aerosol surface (Sakamaki et al., 1983; Pitts et al., 1984; Finlayson-Pitts
578 et al., 2003; Ramazan et al., 2004).



579 For Model 2, which includes the production of HONO via the Eley-Rideal mechanism, in order
580 to reproduce the experimentally observed sharp increase followed by a decrease in both
581 [HONO] and $\gamma_{NO_2 \rightarrow HONO}$ as a function of increasing NO_2 mixing ratio, the modelled rate
582 coefficient for the adsorption of a gas-phase NO_2 molecule to another the surface adsorbed
583 NO_2 to initially form the symmetric N_2O_4 dimer, k_{R12} , had to be larger than for the isomerisation
584 step to form HONO and HNO_3 via *trans*-ONO- NO_2 , k_{R13} . Interestingly, for HONO production
585 via the Langmuir-Hinshelwood mechanism, Model 3, the modelled rate coefficient for the
586 diffusion of one NO_2 molecule across the surface to form the dimer with another NO_2 molecule,
587 k_{R14} , had to be smaller than for the isomerisation step, k_{R15} , to more closely represent the
588 experimental results for the uptake coefficient. Additionally, in order to reproduce the
589 experimental trend in HONO formation as a function of NO_2 mixing ratio, the rate coefficient
590 for the NO_2 dependent loss reaction, k_{R19} , had to be larger than the NO_2 independent reactions,
591 k_{R17} and k_{R18} , leading to $k_{R19} = 5 \times 10^{-3} \text{ s}^{-1}$. The modelled HONO concentration also sensitive
592 to the active site surface concentration: Model 3 required an active site surface concentration
593 2.5 times that of Model 2 to reproduce the peak in [HONO] at ~ 51 ppb NO_2 observed in the
594 experimental results. The reason for this is due to the difference in active site occupation in the
595 2 models: one active site is being occupied by two NO_2 molecules per HONO formed in Model
596 2 as opposed to Model 3 where two active sites are occupied per HONO formed. Regardless
597 of the choice of an Eley Rideal or Langmuir Hinshelwood mechanism, both models reproduce
598 the general shape of [HONO] and $\gamma_{NO_2 \rightarrow HONO}$ with NO_2 , providing evidence that two NO_2
599 molecules are required to form HONO.

600 **3.4 HONO production from illumination of a mixed NH_4NO_3/TiO_2 aerosol** 601 **in the absence of NO_2**

602 The photolysis of particulate nitrate has been postulated as a source of HONO under ambient
603 sunlit conditions during several field campaigns, from both aircraft and ground based
604 measurements (Reed et al., 2017; Ye et al., 2017a; Ye et al., 2017b). Here, experiments were
605 carried out to investigate the formation of HONO from particulate nitrate photolysis, with and
606 without the addition of a photo-catalyst. This is of significant interest for marine environments
607 downwind of arid desert regions due to the availability of TiO_2 or other photocatalytic materials
608 within aerosols in dust plumes that are transported from these regions (Hanisch and Crowley,
609 2003).



610 Using the aerosol flow tube setup described in Sections 2.1-2.4, an aqueous solution of
611 ammonium nitrate (5 g NH_4NO_3 in 500 ml milli-Q water) was used to generate nitrate aerosols.
612 At the RH used in this experiment, $\sim 50\%$, the aerosols were still deliquesced. For these
613 experiments the residence time of the aerosols in the illuminated region of the flow tube was
614 30 seconds (flow rate ~ 6 lpm), with the production of HONO following illumination measured
615 as a function of aerosol surface area density. The number of lamps was increased from 1 to 4,
616 increasing the photon flux from $(1.63 \pm 0.09) \times 10^{16}$ to $(8.21 \pm 2.39) \times 10^{16}$ photons $\text{cm}^{-2} \text{s}^{-1}$
617 and $j(\text{NO}_2)$ from $(6.43 \pm 0.30) \times 10^{-3}$ to $(3.23 \pm 0.92) \times 10^{-2} \text{s}^{-1}$. The $j(\text{NO}_2)$, $j(\text{HONO})$ and flux
618 values for 4 lamps were more than 4 times that of 1 lamp only due to the lamp casings being
619 mirrored, and so with 4 lamps, with 2 lamps on either side of the flow tube, the casings reflected
620 the light back into the flow tube, increasing the effective light intensity. For these experiments,
621 no gaseous NO_2 was added to the gas entering the flow tube. As shown in Figure 11, for the
622 illumination of pure nitrate aerosols, although a small amount of HONO was observed at higher
623 aerosol loadings, no statistically significant production of HONO was seen.

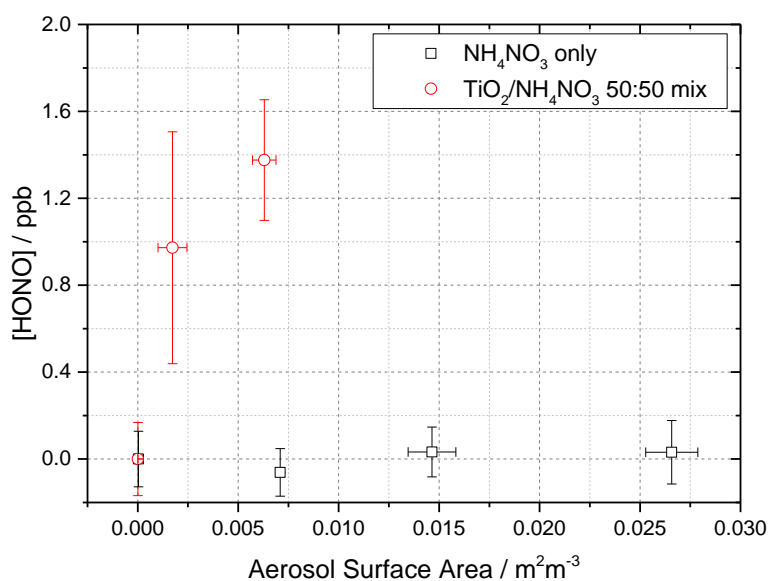


Figure 11. Dependence of the HONO concentration generated as a function of aerosol surface area density for pure NH_4NO_3 aerosol (black open squares, error bars represent 1σ) and 1:1 $\text{TiO}_2/\text{NH}_4\text{NO}_3$ mixed aerosol (red open circles, error bars represent 1σ). Both experiments were performed in N_2 at 295 K, an illuminated residence time of 30 s, and a lamp photon flux of $(8.29 \pm 2.39) \times 10^{16}$ photons $\text{cm}^{-2} \text{s}^{-1}$. The NH_4NO_3 only experiment was performed at $\sim 50 \pm 5\%$ RH while the $\text{TiO}_2/\text{NH}_4\text{NO}_3$ mix experiment was performed at 20 ± 2



% RH. For all points, the background HONO seen observed without illumination has been subtracted. At zero aerosol surface area density there is no HONO generated from the walls of the flow tube.

624 A second set of experiments were performed with an aqueous solution of titanium dioxide and
625 ammonium nitrate combined in a 1:1 mass ratio to give a $\text{TiO}_2/\text{NH}_4\text{NO}_3$ aerosol mixture (5 g
626 NH_4NO_3 and 5 g TiO_2 in 500 ml milli-Q water) to investigate if the photo-catalytic properties
627 of TiO_2 facilitate the production of HONO in the presence of nitrate. The RH was decreased to
628 ensure the maximum TiO_2 photocatalytic activity (Jeong et al., 2013). A recent study using
629 Raman micro spectroscopy to observe phase changes in salt particles reported an efflorescence
630 point of pure ammonium nitrate to be between 13.7-23.9 % RH (Wu et al., 2019). It is possible
631 therefore that at the RH used in this experiment, ~ 20 %, the aerosols were still deliquesced.
632 As shown in Figure 11, the presence of TiO_2 in the aerosol mixture showed a significant
633 production of HONO without the presence of NO_2 , a potentially significant result for the
634 production of HONO in low NO_x environments in the presence of mixed dust/nitrate aerosols,
635 for example in oceanic regions off the coast of West Africa, or in continental regions impacted
636 by outflow from the Gobi desert. Using the Aerosol Inorganic Model (AIM) (Clegg et al.,
637 1998;Wexler and Clegg, 2002), the nitrate content of the aerosol at 50 % RH was calculated.
638 From this and the aerosol volume distribution given by the SMPS, the $[\text{NO}_3^-]$ within the
639 aerosols could be calculated. The formation of HONO by photolysis of particulate nitrate is
640 given by:

$$\frac{d[\text{HONO}]}{dt} = j(\text{pNO}_3)[\text{NO}_3^-] \quad (10)$$

641 and hence:

$$[\text{HONO}] = j(\text{pNO}_3)[\text{NO}_3^-]t \quad (11)$$

642 where $j(\text{pNO}_3)$ is the photolysis frequency of nitrate for the lamps used in these experiments
643 and t is the illumination time of the experiment. With knowledge of $[\text{HONO}]$, $[\text{NO}_3^-]$ and $t =$
644 30 s, $j(\text{pNO}_3)$ can be calculated from a measurement of $[\text{HONO}]$ as a function of $[\text{NO}_3^-]$, as
645 shown in Figure 12, for the mixed nitrate/ TiO_2 experiment.

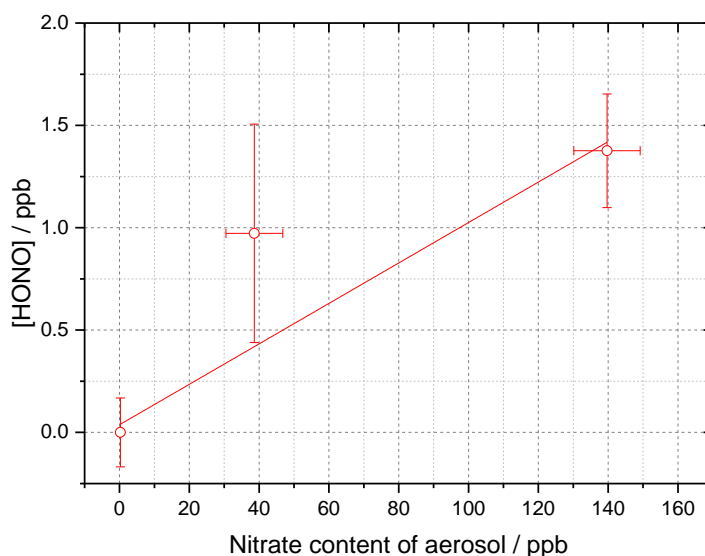


Figure 12. Dependence of [HONO] on the calculated nitrate concentration in the aerosol (using the AIM model) for the mixed TiO₂/ammonium nitrate aerosol experiment. Using Equation (10) and for $t = 30$ s, the gradient gives $j(\text{pNO}_3) = (3.29 \pm 0.89) \times 10^{-4} \text{ s}^{-1}$. Experiment performed at 15 ± 1 % RH, in N₂ at 295 K with a lamp photon flux of $(8.29 \pm 2.39) \times 10^{16} \text{ photons cm}^{-2} \text{ s}^{-1}$. For all points, the background HONO seen observed without illumination has been subtracted.

646 When using the 4 lamps together, the experimental particulate nitrate photolysis rate, $j(\text{pNO}_3)$,
647 was determined to be $(3.29 \pm 0.89) \times 10^{-4} \text{ s}^{-1}$ for the mixed nitrate/TiO₂ aerosol. From this, it
648 is possible to estimate $j(\text{pNO}_3)$ for ambient conditions typical of the tropical marine boundary
649 layer. Taking the ratio of the experimental $j(\text{HONO})$ for 4 lamps $((8.35 \pm 0.18) \times 10^{-3} \text{ s}^{-1})$ and
650 the measured $j(\text{HONO})$ from the RHaMBLe campaign held at the Cape Verde Atmospheric
651 Observatory [May-June, 2007] $(1.2 \times 10^{-3} \text{ s}^{-1})$ (Carpenter et al., 2010; Whalley et al., 2010; Reed
652 et al., 2017) and assuming that $j(\text{pNO}_3)$ and $j(\text{HONO})$ scale in the same way, ambient $j(\text{pNO}_3)$
653 can be determined from:

$$j(\text{pNO}_3)_N = j(\text{pNO}_3) \times \frac{1.2 \times 10^{-3}}{j(\text{HONO})} \quad (12)$$

654 where $j(\text{pNO}_3)_N$ is the photolysis rate coefficient of particulate nitrate at Cape Verde, $j(\text{pNO}_3)$
655 is the experimentally determined photolysis rate coefficient of particulate nitrate to form



656 HONO and $j(\text{HONO})$ is the HONO photolysis rate coefficient calculated from the
657 experimentally determined $j(\text{NO}_2)$.

658 Using $j(\text{pNO}_3) = (3.29 \pm 0.89) \times 10^{-4} \text{ s}^{-1}$, the rate of HONO production from nitrate photolysis
659 at Cape Verde was calculated to be $j(\text{pNO}_3)_{N=} = (4.73 \pm 1.01) \times 10^{-5} \text{ s}^{-1}$ from the mixed
660 nitrate/TiO₂ aerosol experiment. Although for pure nitrate aerosol in the absence of TiO₂ the
661 data were scattered and the HONO production small (Figure 11), an upper limit estimate of
662 $j(\text{pNO}_3)_{N=} = (1.06 \pm 1.15) \times 10^{-6} \text{ s}^{-1}$ under conditions at Cape Verde could be made using
663 equation (11), as done for rate of HONO production from mixed nitrate/TiO₂ aerosols. The
664 atmospheric implications of this will be considered below.

665 **4 Implications of HONO production from TiO₂ for tropospheric chemistry**

666 **4.1 Production of HONO from sunlight aerosols containing TiO₂ in the** 667 **presence of NO₂**

668 For the reactive uptake of NO₂ onto illuminated TiO₂ particles as a function of the initial NO₂
669 mixing ratio, as shown in Figure 8, a maximum value of $\gamma_{\text{NO}_2 \rightarrow \text{HONO}} = (1.26 \pm 0.17) \times 10^{-4}$
670 was determined at 51 ± 5 ppb NO₂ for a photon flux from the lamp of $(1.63 \pm 0.09) \times 10^{16}$
671 photons cm⁻² s⁻¹. These experiments were for single-component TiO₂ particles, and so for dust
672 aerosols a value of $\gamma_{\text{NO}_2 \rightarrow \text{HONO}} = (1.26 \pm 0.17) \times 10^{-5}$ is appropriate assuming a 10 % fraction
673 of TiO₂ and/or other photoactive materials (which behave similarly for HONO production) in
674 mineral dust (Hanisch and Crowley, 2003). Dust aerosols are transported from the Gobi desert
675 to urban areas of China where high NO_x and nitrate aerosol concentrations have been observed
676 and in these areas HONO production facilitated by photo-catalysts may be important (Saliba
677 et al., 2014).

678 Using an average daytime maximum for [NO₂], $j(\text{NO}_2)$ and aerosol surface area measurements
679 for a non-haze period in May-June in 2018 in Beijing, of 50 ppb, $1 \times 10^{-2} \text{ s}^{-1}$ and $2.5 \times 10^{-3} \text{ m}^2$
680 m⁻³ (of which a maximum of 0.3 % was assumed to be TiO₂, though this could be higher in
681 dust impacted events (Schleicher et al., 2010)) respectively, a production rate of HONO of 1.70
682 $\times 10^5$ molecules cm⁻³ s⁻¹ (~ 24.8 ppt h⁻¹) has been estimated using the maximum reactive uptake
683 coefficient measured in this work, $\gamma_{\text{NO}_2 \rightarrow \text{HONO}} = (1.26 \pm 0.17) \times 10^{-4}$. The lamp used to
684 illuminate the TiO₂ aerosols in these experiments gives rise to $j(\text{NO}_2) = (6.43 \pm 0.3) \times 10^{-3} \text{ s}^{-1}$,
685 and so $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$ has been scaled by a factor of 1.55 to match the noon $j(\text{NO}_2)$ measured in
686 May-June 2018 in Beijing (10^{-2} s^{-1}), to take into account the relatively small difference in



687 experimental and atmospheric photon flux for Beijing. The HONO production rate estimated
688 here for noontime summer [May-June 2018] in Beijing (~ 25 ppt hr^{-1}) is similar to the value for
689 the maximum production of HONO from urban humic acid aerosol surfaces in Europe, 17 ppt
690 h^{-1} at 20 ppb NO_2 reported by Stemmler et al., 2007. For comparison, the net gaseous
691 production rate of HONO at noon in May-June 2018 Beijing was determined from the
692 measured rate of gas-phase production and losses:

$$P_{\text{HONO}} = k_{\text{OH}+\text{NO}}[\text{OH}][\text{NO}] - (j(\text{HONO}) \times [\text{HONO}] + k_{\text{OH}+\text{HONO}}[\text{OH}][\text{HONO}]) \quad (13)$$

693 where $k_{\text{OH}+\text{NO}} = 3.3 \times 10^{-11} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Atkinson et al., 2004), $k_{\text{OH}+\text{HONO}} = 6 \times 10^{-12}$
694 $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Atkinson et al., 2004) and $j(\text{HONO}) = 1 \times 10^{-2} \text{ s}^{-1}$ for an average maximum
695 noontime OH concentration of $8 \times 10^6 \text{ molecules cm}^{-3}$ (Whalley et al., 2020), NO
696 concentration of 1.45 ppb (Whalley et al., 2020) and HONO concentration of 0.8 ppb (Whalley
697 et al., 2020).

698 The net gas-phase production of HONO from Equation (13) was calculated to be -3.8 ppt hr^{-1}
699 (a net loss) as expected due to HONO loss by photolysis peaking at solar noon, suggesting the
700 production of HONO heterogeneously from TiO_2 and NO_2 ($\sim 25 \text{ ppt hr}^{-1}$) would have little
701 effect on the overall HONO budget for Beijing summertime at noon.

702 4.2 Production of HONO from photolysis of mixed dust/nitrate aerosols

703 Oceanic environments, for example the Atlantic Ocean which is impacted by both dust aerosols
704 from the Sahara and high concentrations of mixed nitrate aerosols from sea spray, and despite
705 low NO_2 concentrations could be important for particulate nitrate photolysis as a source of
706 HONO (Hanisch and Crowley, 2003; Ye et al., 2017b). From the particulate nitrate photolysis
707 experiments in the absence of NO_2 conducted here, a $j(\text{pNO}_3)_N = (4.73 \pm 1.01) \times 10^{-5} \text{ s}^{-1}$ was
708 determined in the presence of the TiO_2 photo-catalysts (Section 3.4). Using the experimental
709 $j(\text{pNO}_3)$, scaled to typical ambient light levels, and a mean noon concentration of nitrate
710 aerosols of 400 ppt measured at Cape Verde (Reed et al., 2017), taken as an example marine
711 boundary layer environment with a high concentration of mineral dust aerosols, a rate of
712 HONO production from particulate nitrate at Cape Verde was calculated as 4.65×10^5
713 $\text{molecule cm}^{-3} \text{ s}^{-1}$ (68 ppt hr^{-1}). We note that this value would be ~ 50 times smaller for pure
714 nitrate aerosols. The missing rate of HONO production i.e. not taken into account by the gas
715 phase production and loss, P_{other} , from the Cape Verde RHaMBLe campaign, can be calculated
716 using the observed HONO concentration, $[\text{HONO}]$ and the known gas-phase routes for HONO
717 production and loss:



$$P_{other} = ([\text{HONO}](j(\text{HONO}) + k_{\text{OH}+\text{HONO}}[\text{OH}])) - (k_{\text{OH}+\text{NO}}[\text{OH}][\text{NO}]) \quad (14)$$

718 where $k_{\text{OH}+\text{NO}} = 3.3 \times 10^{-11} \text{ cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Atkinson et al., 2004), $k_{\text{OH}+\text{HONO}} = 6 \times 10^{-12}$
719 $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$ (Atkinson et al., 2004) and $j(\text{HONO}) = 2 \times 10^{-3} \text{ s}^{-1}$ for average maximum
720 measured concentrations of $1 \times 10^7 \text{ molecules cm}^{-3}$ for OH (Whalley et al., 2010), 5.41×10^7
721 molecule cm^{-3} for NO (Whalley et al., 2010) and $1.23 \times 10^8 \text{ molecule cm}^{-3}$ for HONO (Whalley
722 et al., 2010).

723 Using Equation (14) this missing HONO production rate for Cape Verde was 34.6 ppt hr^{-1} ,
724 which is within a factor of two of the rate of HONO production (68 ppt hr^{-1}) calculated from
725 nitrate photolysis using our experimental HONO production data for mixed nitrate/TiO₂
726 aerosols. These results provide further evidence that particulate nitrate photolysis in the
727 presence of photocatalytic compounds such as TiO₂ found in dust could be significant in
728 closing the HONO budget for this environment (Whalley et al., 2010; Reed et al., 2017; Ye et
729 al., 2017a).

730 5 Conclusions.

731 The experimental production of HONO from both illuminated TiO₂ aerosols in the presence of
732 NO₂ and from mixed nitrate/TiO₂ aerosols in the absence of NO₂ was observed, with the
733 HONO concentrations measured using photo-fragmentation laser-induced fluorescence
734 spectroscopy. Using experimental data, the reactive uptake of NO₂ onto the TiO₂ aerosol
735 surface to produce HONO, $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$, was determined for NO₂ mixing ratios ranging from
736 34 to 400 ppb, with a maximum $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$ value of $(1.26 \pm 0.17) \times 10^{-4}$ for single-component
737 TiO₂ aerosols observed at 51 ppb NO₂, and for a lamp photon flux of $(1.65 \pm 0.02) \times 10^{16}$
738 $\text{photons cm}^{-2} \text{ s}^{-1}$ (integrated between 290 and 400 nm). The measured reactive uptake
739 coefficient, $\gamma_{\text{NO}_2 \rightarrow \text{HONO}}$, showed an increase then subsequent decrease as a function of NO₂
740 mixing ratio, peaking at $51 \pm 5 \text{ ppb}$. Box modelling studies supported a mechanism involving
741 two NO₂ molecules on the aerosol surface per HONO molecule generated, providing evidence
742 for the formation of a surface-bound NO₂ dimer intermediate. The exact mechanism for HONO
743 formation, for examples the step(s) which are accelerated in the presence of light, remains
744 unclear, although previous studies would suggest the process occurs via the isomerisation of
745 the symmetric N₂O₄ dimer to give *trans*-ONO-NO₂, either via *cis*-ONO-NO₂ or directly,
746 suggested to be more reactive with water than the symmetric dimer (Finlayson-Pitts et al.,
747 2003; Ramazan et al., 2004; Ramazan et al., 2006; de Jesus Madeiros and Pimentel, 2011; Liu
748 and Goddard, 2012; Murdachaew et al., 2013; Varner et al., 2014). Investigations into the RH



749 dependence of the HONO production mechanism on TiO₂ aerosols showed a peak in
750 production between ~25-30 % RH, with lower HONO production at higher NO₂ mixing ratios
751 observed for all RHs tested. The increase in HONO production with increasing RH can be
752 attributed to a higher concentration of H₂O on the surface increasing its availability for the
753 hydrolysis reaction to give HONO, whereas a decrease in HONO production after RH ~ 30 %
754 could be due to the increased water surface concentration inhibiting the adsorption of NO₂.
755 Using the laboratory reactive uptake coefficient for HONO production, $\gamma_{NO_2 \rightarrow HONO}$, the rate of
756 production of HONO from illuminated aerosols in Beijing in summer for typical NO₂ mixing
757 ratios and aerosol surface areas was found to be similar to that estimated previously for the
758 production of HONO from urban humic acid aerosol surfaces in Europe.

759 In the absence of NO₂, significant HONO production from 50:50 mixed nitrate/TiO₂ aerosols
760 was measured. Using the experimental HONO concentrations observed, a rate of HONO
761 production from nitrate photolysis was calculated, which was then scaled to the ambient
762 conditions encountered at the Cape Verde Atmospheric Observatory in the tropical marine
763 boundary layer. A HONO production rate of 68 ppt hr⁻¹ for the mixed nitrate/TiO₂ aerosol was
764 found for CVAO conditions, similar in magnitude to the missing HONO production rate that
765 had been calculated previously in order to bring modelled HONO concentrations into line with
766 field-measured values at CVAO. These results provide further evidence that aerosol particulate
767 nitrate photolysis may be significant as a source of HONO, and hence NO_x, in the remote
768 marine boundary layer, where mixed aerosols containing nitrate and a photo-catalytic species
769 such as TiO₂, as found in dust, are present.

770 However, the production of HONO from pure, deliquesced ammonium nitrate aerosols alone
771 could not be definitively confirmed over the range of conditions used in our experiments,
772 suggesting that another component within the aerosol is necessary for HONO production.
773 Future work should be directed towards studying pure nitrate aerosols over a wider range of
774 conditions, for example varying the aerosol pH, and also adding other chemical species into
775 the aerosol which may promote HONO production.

776 *Data availability.* Data presented in this study can be obtained from authors upon request
777 (d.e.heard@leeds.ac.uk)

778 *Competing interests.* The authors declare that they have no conflict of interest.



779 *Acknowledgements.* We are grateful to the Natural Environmental Research Council for
780 funding a SPHERES PhD studentship (Joanna E. Dyson) and for funding the EXHALE project
781 (grant number NE/S006680/1).

782 **References**

- 783 Alicke, B., Platt, U., and Stutz, J.: Impact of nitrous acid photolysis on the total hydroxyl radical
784 budget during the Limitation of Oxidant Production/Pianura Padana Produzione di Ozono
785 study in Milan, *J. Geophys. Res. Atmos.*, 107, <https://doi.org/10.1029/2000JD000075>, 2002.
- 786 Atkinson, R., Baulch, D. L., Cox, R. A., Crowley, J. N., Hampson, R. F., Hynes, R. G., Jenkin,
787 M. E., Rossi, M. J., and Troe, J.: Evaluated kinetic and photochemical data for atmospheric
788 chemistry: Volume I - gas phase reactions of O_x, HO_x, NO_x and SO_x species, *Atmos. Chem.*
789 *Phys.*, 4, 1461-1738, <https://doi.org/10.5194/acp-4-1461-2004>, 2004.
- 790 Beckers, H., Zeng, X., and Willner, H.: Intermediates involved in the oxidation of nitrogen
791 monoxide: Photochemistry of the cis-N₂O₂-O₂ complex and of sym-N₂O₄ in Solid Ne Matrices,
792 *Chemistry—A European Journal*, 16, 1506-1520, <https://doi.org/10.1002/chem.200902406>,
793 2010.
- 794 Boustead, G. A.: Measurement of nitrous acid production from aerosol surfaces using Photo-
795 Fragmentation Laser-Induced Fluorescence, School of Chemistry, University of Leeds, 2019.
- 796 Bröske, R., Kleffmann, J., and Wiesen, P.: Heterogeneous conversion of NO₂ on secondary
797 organic aerosol surfaces: A possible source of nitrous acid (HONO) in the atmosphere?, *Atmos.*
798 *Chem. Phys.*, 3, 469-474, <https://doi.org/10.5194/acp-3-469-2003>, 2003.
- 799 Cantrell, C., Zimmer, A., and Tyndall, G. S.: Adsorption cross sections for water vapor from
800 183 to 193 nm, *Geophys. Res. Lett.*, 24, 2195-2198, <https://doi.org/10.1029/97GL02100>, 1997.
- 801 Carpenter, L. J., Fleming, Z. L., Read, K. A., Lee, J. D., Moller, S. J., Hopkins, J. R., Purvis,
802 R. M., Lewis, A. C., Müller, K., Heinold, B., Herrmann, H., Fomba, K. W., van Pinxteren, D.,
803 Müller, C., Tegen, I., Wiedensohler, A., Müller, T., Niedermeier, N., Achterberg, E. P., Patey,
804 M. D., Kozlova, E. A., Heimann, M., Heard, D. E., Plane, J. M. C., Mahajan, A., Oetjen, H.,
805 Ingham, T., Stone, D., Whalley, L. K., Evans, M. J., Pilling, M. J., Leigh, R. J., Monks, P. S.,
806 Karunaharan, A., Vaughan, S., Arnold, S. R., Tschritter, J., Pöhler, D., Frieß, U., Holla, R.,
807 Mendes, L. M., Lopez, H., Faria, B., Manning, A. J., and Wallace, D. W. R.: Seasonal
808 characteristics of tropical marine boundary layer air measured at the Cape Verde Atmospheric
809 Observatory, *Journal of Atmospheric Chemistry*, 67, 87-140, [https://doi.org/10.1007/s10874-](https://doi.org/10.1007/s10874-011-9206-1)
810 [011-9206-1](https://doi.org/10.1007/s10874-011-9206-1), 2010.
- 811 Chen, H., Nanayakkara, C. E., and Grassian, V. H.: Titanium dioxide photocatalysis in
812 atmospheric chemistry, *Chem. Rev.*, 112, 5919-5948, <https://doi.org/10.1021/cr3002092>,
813 2012.
- 814 Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic model of the system H⁺-
815 NH₄⁺- Na⁺- SO₄²⁻- NO₃⁻- Cl⁻- H₂O at 298.15 K, *J. Phys. Chem. A*, 102, 2155-2171,
816 <https://doi.org/10.1021/jp973043j>, 1998.
- 817 Crilley, L. R., Kramer, L. J., Ouyang, B., Duan, J., Zhang, W., Tong, S., Ge, M., Tang, K., Qin,
818 M., Xie, P., Shaw, M. D., Lewis, A. C., Mehra, A., Bannan, T. J., Worrall, S. D., Priestley, M.,
819 Bacak, A., Coe, H., Allan, J., Percival, C. J., Popoola, O. A. M., Jones, R. L., and Bloss, W. J.:
820 Intercomparison of nitrous acid (HONO) measurement techniques in a megacity (Beijing),
821 *Atmos. Meas. Tech.*, 12, 6449-6463, <https://doi.org/10.5194/amt-12-6449-2019>, 2019.



- 822 de Jesus Madeiros, D., and Pimentel, A. S.: New insights in the atmospheric HONO formation:
823 new pathways for N₂O₄ isomerisation and NO₂ dimerisation in the presence of water. , J. Phys.
824 Chem. A, 115, 6357-6365, <https://doi.org/10.1021/jp1123585>, 2011.
- 825 Dupart, Y., Fine, L., D'Anna, B., and George, C.: Heterogeneous uptake of NO₂ on Arizona
826 Test Dust under UV-A irradiation: an aerosol flow tube study, *Aeolian Res.*, 15, 45-51,
827 <https://doi.org/10.1016/j.aeolia.2013.10.001>, 2014.
- 828 Fateley, W. G., Bent, H. A., and Crawford Jr, B.: Infrared spectra of the frozen oxides of
829 nitrogen, *J. Chem. Phys.*, 31, 204-217, <https://doi.org/10.1063/1.1730296>, 1959.
- 830 Finlayson-Pitts, B. J., Wingen, L. M., Summer, A. L., Syomin, D., and Ramazan, K. A.: The
831 heterogeneous hydrolysis of NO₂ in laboratory systems in outdoor and indoor atmospheres: An
832 intergrated mechanism, *Phys.Chem.Phys.Chem*, 5, 223-242, <https://doi.org/10.1039/b208564j>,
833 2003.
- 834 Forney, D., Thompson, W. E., and Jacox, M. E.: The vibrational spectra of molecular ions
835 isolated in solid neon. XI. NO₂⁺, NO₂⁻, and NO₃⁻, *The Journal of Chemical Physics*, 99, 7393-
836 7403, <https://doi.org/10.1063/1.465720>, 1993.
- 837 George, C., Strekowski, R. S., Kleffmann, J., Stemmler, K., and Ammann, M.: Photoenhanced
838 uptake of gaseous NO₂ on solid organic compounds: a photochemical source of HONO?,
839 *Faraday Discuss.*, 130, <https://doi.org/10.1039/b417888m>, 2005.
- 840 George, I. J., Matthews, P. S. J., Whalley, L. K., Brooks, B., Goddard, A., Baeza-Romero, M.,
841 and Heard, D. E.: Measurements of uptake coefficients for heterogeneous loss of HO₂ onto
842 submicron inorganic salt aerosols., *Phys. Chem. Chem. Phys.*, 15, 12829-12845,
843 <https://doi.org/10.1039/c3cp51831k>, 2013.
- 844 Ginoux, P., Chin, M., Tegen, I., Prospero, J. M., Holben, B., Dubovik, O., and Lin, S. J.:
845 Sources and distributions of dust aerosols simulated with the GOCART model, *J. Geophys.*
846 *Res. Atmos.*, 106, 20255-20273, <https://doi.org/10.1029/2000JD000053>, 2001.
- 847 Givan, A., and Loewenschuss, A.: Fourier transform infrared and Raman studies on solid
848 nitrogen dioxide: Temperature cycling of ordered, disordered, and multicomponent layers, *The*
849 *Journal of Chemical Physics*, 90, 6135-6142, <https://doi.org/10.1063/1.456379>, 1989a.
- 850 Givan, A., and Loewenschuss, A.: On the intermolecularity or intramolecularity of nitrosonium
851 nitrate formation in thin films of nitrogen dioxide: A Fourier transform infrared study, *The*
852 *Journal of chemical physics*, 91, 5126-5127, <https://doi.org/10.1063/1.457609>, 1989b.
- 853 Givan, A., and Loewenschuss, A.: Fourier transform infrared study of amorphous N₂O₄ solid:
854 Destabilization with inert impurities, *The Journal of chemical physics*, 94, 7562-7563,
855 <https://doi.org/10.1063/1.460192>, 1991.
- 856 Goodman, A. L., Bernard, E. T., and Grassian, V. H.: Spectroscopic study of nitric acid and
857 water adsorption on oxide particles: enhanced nitric acid uptake kinetics in the presence of
858 adsorbed water, *J. Phys. Chem. A*, 105, 6443-6457, <https://doi.org/10.1021/jp0037221>, 2001.
- 859 Gustafsson, R. J., Orlov, A., Griffiths, P. T., Cox, R. A., and Lambert, R. M.: Reduction of
860 NO₂ to nitrous acid on illuminated titanium dioxide aerosol surfaces: implications for
861 photocatalysis and atmospheric chemistry, *Chem. Commun.*, 37, 3936-3938,
862 <https://doi.org/10.1039/b609005b>, 2006.
- 863 Hanisch, F., and Crowley, J. N.: Ozone decomposition on Saharan dust: an experimental
864 investigation, *Atmos. Chem. Phys.*, 3, 119-130, <https://doi.org/10.5194/acp-3-119-2003>, 2003.



- 865 Harrison, R. M., Peak, J. D., and Collins, G. M.: Tropospheric cycle of nitrous acid, *J. Geophys.*
866 *Res. Atmos.*, 101, 14429-14439, <https://doi.org/10.1029/96JD00341>, 1996.
- 867 Heard, D. E.: Atmospheric field measurements of the hydroxyl radical using Laser-Induced
868 Fluorescence spectroscopy, *Annu. Rev. Phys. Chem.*, 57, 191-216,
869 <https://doi.org/10.1146/annurev.physchem.57.032905.104516>, 2006.
- 870 Jeong, M.-G., Park, E. J., Seo, H. O., Kim, K.-D., Kim, Y. D., and Lim, D. C.: Humidity effect
871 on photocatalytic activity of TiO₂ and regeneration of deactivated photocatalysts, *Appl. Surf.*
872 *Sci.*, 271, 164-170, <https://doi.org/10.1016/j.apsusc.2013.01.155>, 2013.
- 873 Kleffmann, J.: Daytime sources of nitrous acid (HONO) in the atmospheric boundary layer,
874 *Chem. Phys. Chem.*, 8, 1137-1144, <https://doi.org/10.1002/cphc.200700016>, 2007.
- 875 Kurtenbach, R., Becker, K. H., Gomes, J. A. G., Kleffmann, J., Lörzer, J. C., Spittler, M.,
876 Wiesen, P., Ackermann, R., Geyer, A., and Platt, U.: Investigations of emissions and
877 heterogeneous formation of HONO in road traffic tunnel, *Atmos. Environ.*, 35, 3385-3394,
878 [https://doi.org/10.1016/S1352-2310\(0\)00138-8](https://doi.org/10.1016/S1352-2310(0)00138-8), 2001.
- 879 Lee, J. D., Whalley, L. K., Heard, D. E., Stone, D., Dunmore, R. E., Hamilton, J. F., Young,
880 D. E., Allan, J. D., Laufs, S., and Kleffmann, J.: Detailed budget analysis of HONO in central
881 London reveals a missing daytime source, *Atmos. Chem. Phys.*, 16, 2747-2764,
882 <https://doi.org/10.5194/acp-16-2747-2016>, 2016.
- 883 Levy, H.: Normal atmosphere: large radical and formaldehyde concentrations predicted,
884 *Science*, 173, 141-143, <https://doi.org/10.1126/science.173.3992.141>, 1971.
- 885 Li, S., Matthews, J., and Sinha, A.: Atmospheric hydroxyl radical production from
886 electronically excited NO₂ and H₂O, *Science*, 319, <https://doi.org/10.1126/science.1151443>,
887 2008.
- 888 Liao, W., Hecobian, A., Mastromarino, J., and Tan, D.: Development of a photo-
889 fragmentation/laser-induced fluorescence measurement of atmospheric nitrous acid,
890 *Atmospheric Environment*, 40, 17-26, <https://doi.org/10.1016/j.atmosenv.2005.07.001>, 2006.
- 891 Liu, W. G., and Goddard, W. A.: First-principle study of the role of interconversion between
892 NO₂, N₂O₄, *cis*-ONO-NO₂, and *trans*-ONO-NO₂ in chemical processes, *J. Am. Chem. Soc.*,
893 134, 12970-12978, <https://doi.org/10.1021/ja300545e>, 2012.
- 894 Lu, K., Fuchs, H., Hofzumahaus, A., Tan, Z., Wang, H., Zhang, L., Schmitt, S. H., Rohrer, F.,
895 Bohn, B., Broch, S., Dong, H., Gkatzelis, G. I., Hohaus, T., Holland, F., Li, X., Liu, Y., Liu,
896 Y., Ma, X., Novelli, A., Schlag, P., Shao, M., Wu, Y., Wu, Z., Zeng, L., Hu, M., Kiendler-
897 Scharr, A., Wahner, A., and Zhang, Y.: Fast Photochemistry in Wintertime Haze:
898 Consequences for Pollution Mitigation Strategies, *Environ. Sci. Technol.*, 53, 10676-10684,
899 <https://doi.org/10.1021/acs.est.9b02422>, 2019.
- 900 Lu, X., Park, J., and Lin, M. C.: Gas phase reactions of HONO with NO₂, O₃ and HCl: Ab initio
901 and TST study, *J. Phys. Chem. A*, 104, 8730-8738, <https://doi.org/10.1021/jp001610o>, 2000.
- 902 Matthews, P. S. J., Baeza-Romero, M. T., Whalley, L. K., and Heard, D. E.: Uptake of HO₂
903 radicals onto Arizona test dust particles using an aerosol flow tube, *Atmos. Chem. Phys.*, 14,
904 7397-7408, <https://doi.org/10.5194/acp-14-7397-2014>, 2014.
- 905 MCPA software Ltd.: Facsimile, 2020.
- 906 Michoud, V., Colomb, A., Borbon, A., Miet, K., Beekmann, M., Camredon, M., Aumont, B.,
907 Perrier, S., Zapf, P., Siour, G., Ait-Helal, W., Afif, C., Kukui, A., Furger, M., Dupont, J. C.,
908 Haefelin, M., and Doussin, J. F.: Study of the unknown HONO daytime source at a European



- 909 suburban site during the MEGAPOLI summer and winter field campaigns, *Atmos. Chem.*
910 *Phys.*, 14, 2805-2822, <https://doi.org/10.5194/acp-14-2805-2014>, 2014.
- 911 Moon, D. R., Taverna, G. S., Anduix-Canto, C., Ingham, T., Chipperfield, M. P., Seakins, P.
912 W., Baeza-Romero, M.-T., and Heard, D. E.: Heterogeneous reaction of HO₂ with airborne
913 TiO₂ particles and its implication for climate change mitigation strategies, *Atmos. Chem. Phys.*,
914 18, 327–338, <https://doi.org/10.5194/acp-18-327-2018>, 2018.
- 915 Moon, D. R., Ingham, T., Whalley, L. K., Seakins, P. W., Baeza-Romero, M. T., and Heard,
916 D. E.: Production of OH and HO₂ radicals from near-UV irradiated airborne TiO₂
917 nanoparticles, *Phys.Chem.Phys.Chem*, 21, 2325-2336, <https://doi.org/10.1039/C8CP06889E>,
918 2019.
- 919 Murdachaew, G., Varner, M. E., Philips, L. F., Finlayson-Pitts, B. J., and Gerber, R. B.:
920 Nitrogen dioxide at the air-water interface: trapping, adsorption, and solvation in the bulk and
921 at the surface, *Phys. Chem. Chem. Phys.*, 15, 204-212, <https://doi.org/10.1039/c2cp42810e>,
922 2013.
- 923 Nakamura, I., Sugihara, S., and Takeuchi, K.: Mechanism for NO photooxidation over the
924 oxygen-deficient TiO₂ powder under visible light irradiation, *Chem. Lett.*, 29, 1276-1277,
925 <https://doi.org/10.1246/cl.2000.1276>, 2000.
- 926 Ndour, M., D'Anna, B., George, C., Ka, O., Balkanski, Y., Kleffman, J., Stemmler, K., and
927 Ammann, M.: Photoenhanced uptake of NO₂ on mineral dust: Laboratory experiments and
928 model simulations, *Geophys. Res. Lett.*, 35, L05812, <https://doi.org/10.1029/2007GL032006>,
929 2008.
- 930 Oswald, R., Behrendt, T., Ermel, M., Wu, D., Su, H., Cheng, Y., Breuninger, C., Moravek, A.,
931 Mougín, E., Delon, C., Loubet, B., Pommerening-Röser, A., Sörgel, M., Pöschl, U., Hoffmann,
932 T., Andeae, M. O., Meixner, F. X., and Trebs, I.: HONO emissions from soil bacteria as a
933 major source of atmospheric reactive nitrogen, *Science*, 341, 1233-1235,
934 <https://doi.org/10.1126/science.1242266>, 2013.
- 935 Pinnick, D., Agnew, S., and Swanson, B.: Fluid dinitrogen tetroxide at very high pressure and
936 high temperature: observation of the nitrite isomer, *The Journal of Physical Chemistry*, 96,
937 7092-7096, <https://doi.org/10.1021/j100196a046>, 1992.
- 938 Pitts, J. N., Sanhueza, E., Atkinson, R., Carter, W. P. L., Winer, A. M., Harris, G. W., and
939 Plum, C. N.: An investigation of the dark formation of nitrous acid in environmental chambers,
940 *Int. J. Chem. Kinet.*, 16, 919-939, <https://doi.org/10.1002/kin.550160712>, 1984.
- 941 Platt, U., Perner, D., Harris, G. W., Winer, A. M., and Pitts, J. N.: Observations of nitrous acid
942 in an urban atmosphere by differential optical absorption, *Nature*, 285, 312-314,
943 <https://doi.org/10.1038/285312a0>, 1980.
- 944 Ramazan, K. A., Syomin, D., and Finlayson-Pitts, B. J.: The photochemical production of
945 HONO during the heterogeneous hydrolysis of NO₂, *Phys. Chem. Chem. Phys.*, 6, 3836-3843,
946 <https://doi.org/10.1039/B402195A>, 2004.
- 947 Ramazan, K. A., Wingen, L. M., Miller, Y., Chaban, G. M., Gerber, R. B., Xantheas, S. S., and
948 Finlayson-Pitts, B. J.: New experimental and theoretical approach to the heterogeneous
949 hydrolysis of NO₂: key role of molecular nitric acid and its complexes, *J. Phys. Chem. A*, 110,
950 6886-6897, <https://doi.org/10.1021/jp056426n>, 2006.
- 951 Reed, C., Evans, M. J., Crilley, L. R., Bloss, W. J., Sherwen, T., Read, K. A., Lee, J. D., and
952 Carpenter, L. J.: Evidence for renoxification in the tropical marine boundary layer, *Atmos.*
953 *Chem. Phys.*, 17, 4081-4092, <https://doi.org/10.5194/acp-17-4081-2017>, 2017.



- 954 Sakamaki, F., Hatakeyama, S., and Akimoto, H.: Formation of nitrous acid and nitric oxide in
955 the heterogeneous dark reaction of nitrogen dioxide and water vapor in a smog chamber, *Int.*
956 *J. Chem. Kinet.*, 15, 1013-1029, <https://doi.org/10.1002/kin.550151006>, 1983.
- 957 Saliba, N., Moussa, S., and Tayyar, G.: Contribution of airborne dust particles to HONO
958 sources, *Atmos. Chem. Phys. Discuss.*, 14, <https://doi.org/10.5194/acpd-14-4827-2014>, 2014.
- 959 Sander, S., Friedl, R., Barker, J., Golden, D., Kurylo, M., Wine, P., Abbatt, J., Burkholder, J.,
960 Kolb, C., and Moortgat, G.: Chemical kinetics and photochemical data for use in atmospheric
961 studies, evaluation number 14, *JPL Publ.*, 02, 25, 2003.
- 962 Schleicher, N., Norra, S., Chai, F., Chen, Y., Wang, S., and Stüben, D.: Anthropogenic versus
963 geogenic contribution to total suspended atmospheric particulate matter and its variations
964 during a two-year sampling period in Beijing, China, *J. Environ. Monit.*, 12, 434-441,
965 <https://doi.org/10.1039/B914739J>, 2010.
- 966 Seifert, N. A., Zaleski, D. P., Fehnel, R., Goswami, M., Pate, B. H., Lehmann, K. K., Leung,
967 H. O., Marshall, M. D., and Stanton, J. F.: The gas-phase structure of the asymmetric, trans-
968 dinitrogen tetroxide (N₂O₄), formed by dimerization of nitrogen dioxide (NO₂), from rotational
969 spectroscopy and ab initio quantum chemistry, *The Journal of Chemical Physics*, 146, 134305,
970 <https://doi.org/10.1063/1.4979182>, 2017.
- 971 Shan, J. H., Wategaonkar, S. J., and Vasudev, R.: Vibrational state dependence of the A state
972 lifetime of HONO, *Chem. Phys. Lett.*, 158, 317-320, [https://doi.org/10.1016/0009-2614\(89\)87343-9](https://doi.org/10.1016/0009-2614(89)87343-9), 1989.
- 974 Slater, E. J., Whalley, L. K., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F., Hopkins,
975 J. R., Dunmore, R. E., Shaw, M., Hamilton, J. F., Lewis, A. C., Crilley, L. R., Kramer, L.,
976 Bloss, W., Vu, T., Sun, Y., Xu, W., Yue, S., Ren, L., Acton, W. J. F., Hewitt, C. N., Wang, X.,
977 Fu, P., and Heard, D. E.: Elevated levels of OH observed in haze events during wintertime in
978 central Beijing, *Atmos. Chem. Phys. Discuss.*, 2020, 1-43, <https://doi.org/10.5194/acp-2020-979>
362, 2020.
- 980 Spataro, F., and Ianniello, A.: Sources of atmospheric nitrous acid: State of the science, current
981 research needs, and future prospects, *J. Air Waste Manage. Assoc.*, 64, 1232-1250,
982 <https://doi.org/10.1080/10962247.2014.952846>, 2014.
- 983 Stemmler, K., Ndour, M., Elshorbany, Y., Kleffmann, J., D'Anna, B., George, C., Bohn, B.,
984 and Ammann, M.: Light induced conversion of nitrogen dioxide into nitrous acid on submicron
985 humic acid aerosol, *Atmos. Chem. Phys.*, 7, 4237-4248, <https://doi.org/10.5194/acp-7-4237-986>
2007, 2007.
- 987 Su, H., Cheng, Y., Oswald, R., Behrendt, T., Trebs, I., Meixner, F. C., Andreae, M. O., Cheng,
988 P., Zhang, Y., and Pöschl, U.: Soil nitrate as a source of atmospheric HONO and OH radicals,
989 *Science*, 333, 1616-1618, <https://doi.org/10.1126/science.1207687>, 2011.
- 990 Syomin, D. A., and Finlayson-Pitts, B. J.: HONO decomposition on borosilicate glass surfaces:
991 implications for environmental chamber studies and field experiments, *Phys. Chem. Chem.*
992 *Phys.*, 5, 5236-5242, <https://doi.org/10.1039/b309851f>, 2003.
- 993 Takeuchi, M., Sakamoto, K., Martra, G., Coluccia, S., and Anpo, M.: Mechanism of
994 photoinduced superhydrophilicity on the TiO₂ photocatalyst surface, *J. Phys. Chem. B*, 109,
995 15422-15428, <https://doi.org/10.1021/jp058075i>, 2005.
- 996 Varner, M. E., Finlayson-Pitts, B. J., and Gerber, R. B.: Reaction of a charge-separated
997 ONONO₂ species with water in the formation of HONO: an MP2 molecular dynamics study,
998 *Phys. Chem. Chem. Phys.*, 16, 4483-4487, <https://doi.org/10.1039/c3cp55024a>, 2014.



- 999 Wang, J., and Koel, B. E.: IRAS studies of NO₂, N₂O₃, and N₂O₄ adsorbed on Au (111) surfaces
1000 and reactions with coadsorbed H₂O, *J. Phys. Chem. A*, 102, 8573-8579,
1001 <https://doi.org/10.1021/jp982061d>, 1998.
- 1002 Wang, J., and Koel, B. E.: Reactions of N₂O₄ with ice at low temperatures on the Au (111)
1003 surface, *Surf. Sci.*, 436, 15-28, [https://doi.org/10.1016/S0039-6028\(99\)00457-4](https://doi.org/10.1016/S0039-6028(99)00457-4), 1999.
- 1004 Wexler, A. S., and Clegg, S. L.: Atmospheric aerosol models for systems including the ions
1005 H⁺, NH₄⁺, Na⁺, SO₄²⁻, NO₃⁻, Cl⁻, Br⁻, and H₂O, *J. Geophys. Res. Atmos.*, 107, ACH 14-11-
1006 ACH 14-14, <https://doi.org/10.1029/2001JD000451>, 2002.
- 1007 Whalley, L. K., Furneaux, K. L., Goddard, A., Lee, J. D., Mahajan, A., Oetjen, H., Read, K.
1008 A., Kaaden, N., Carpenter, L. J., Lewis, A. C., Plane, J. M. C., Saltzman, E. S., Wiedensohler,
1009 A., and Heard, D. E.: The chemistry of OH and HO₂ radicals in the boundary layer over the
1010 tropical Atlantic Ocean, *Atmos. Chem. Phys.*, 10, 1555-1576, <https://doi.org/10.5194/acp-10-1555-2010>, 2010.
- 1012 Whalley, L. K., Stone, D., Dunmore, R., Hamilton, J., Hopkins, J. R., Lee, J. D., Lewis, A. C.,
1013 Williams, P., Kleffmann, J., Laufs, S., Woodward-Massey, R., and Heard, D. E.:
1014 Understanding in situ ozone production in the summertime through radical observations and
1015 modelling studies during the Clean air for London project (ClearLo), *Atmos. Chem. Phys.*, 18,
1016 2547-2571, <https://doi.org/10.5194/acp-18-2547-2018>, 2018.
- 1017 Whalley, L. K., Slater, E. J., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F., Hopkins,
1018 J. R., Dunmore, R. E., Shaw, M., Hamilton, J. F., Lewis, A. C., Mehra, A., Worrall, S. D.,
1019 Bacak, A., Bannan, T. J., Coe, H., Ouyang, B., Jones, R. L., Crilley, L. R., Kramer, L. J., Bloss,
1020 W. J., Vu, T., Kotthaus, S., Grimmond, S., Sun, Y., Xu, W., Yue, S., Ren, L., Acton, W. J. F.,
1021 Hewitt, C. N., Wang, X., Fu, P., and Heard, D. E.: Evaluating the sensitivity of radical
1022 chemistry and ozone formation to ambient VOCs and NO_x in Beijing, *Atmos. Chem. Phys.*
1023 *Discuss.*, 2020, 1-41, <https://doi.org/10.5194/acp-2020-785>, 2020.
- 1024 Winer, A. M., and Biermann, H. W.: Long pathlength differential optical absorption
1025 spectroscopy (DOAS) measurements of gaseous HONO, NO₂ and HCNO in the California
1026 South Coast Air Basin, *Res. Chem. Intermed.*, 20, 423-445,
1027 <https://doi.org/10.1163/156856794X00405>, 1994.
- 1028 Wu, L., Li, X., and Ro, C.: Hygroscopic behavior of ammonium sulfate, ammonium nitrate,
1029 and their mixture particles, *Asian J. Atmos. Environ.*, 13.3, 196-211,
1030 <https://doi.org/10.5572/ajae.2019.13.3.196>, 2019.
- 1031 Ye, C., Zhou, X., Pu, D., Stutz, J., Festa, J., Spolaor, M., Tsai, C., Cantrell, C., Mauldin, R. L.,
1032 Campos, T., Weinheimer, A., Hornbrook, R. S., Apel, E. C., Guenther, A., Kaser, L., Yuan,
1033 B., Karl, T., Haggerty, J., Hall, S., Ullmann, K., Smith, J. N., Ortega, J., and Knote, C.: Rapid
1034 cycling of reactive nitrogen in the marine boundary layer, *Nature*, 532, 489-491,
1035 <https://doi.org/10.1038/nature17195>, 2016.
- 1036 Ye, C., Heard, D. E., and Whalley, L. K.: Evaluation of novel routes for NO_x formation in
1037 remote regions, *Environmental Science Technology*, 51, 7442-7449,
1038 <https://doi.org/acs.est.6b06441>, 2017a.
- 1039 Ye, C., Zhang, N., Gao, H., and Zhou, X.: Photolysis of particulate nitrate as a source of HONO
1040 and NO_x, *Environ. Sci. Technol.*, 51, 6849-6856, <https://doi.org/10.1021/acs.est.7b00387>,
1041 2017b.



1042 Zhou, X., Gao, H., He, Y., Huang, G., Bertman, S. B., Civerolo, K., and Schwab, J.: Nitric acid
1043 photolysis on surfaces in low NO_x environments: Significant atmospheric implications,
1044 Geophys. Res. Lett., 30, 2217, <https://doi.org/10.1029/2003GL018620>, 2003.

1045

1046