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Nitrogen deposition and climate effects on soil nitrogen availability: influences of habitat type and soil characteristics.

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17

18 **Abstract**

19

20 The amount of plant-available nitrogen (N) in soil is an important indicator of eutrophication of semi-
21 natural habitats, but previous studies have shown contrasting effects of N deposition on mineralisable
22 N in different habitats. The stock of readily mineralisable N (N_{rm}) was measured in 665 locations
23 across Britain from a range of intensively and extensively managed habitats, allowing N availability
24 to be studied in relation to soil and vegetation type, and also to variation in climate and in reactive N
25 deposition from the atmosphere. Mineralisable N contents were correlated with deposition in
26 extensively managed habitats but not in intensively managed habitats. The following statements apply
27 only to extensively managed habitats. All habitats showed a similar increase in N_{rm} with N deposition.
28 However, soil characteristics affected the relationship, and soil carbon content in particular was a
29 major control on mineralisation. The N_{rm} stock increased more with N deposition in organic than in
30 mineral soils. The nitrate proportion of N_{rm} also increased with N deposition but, conversely, this
31 increase was greater in mineral than in organic soils. The measurements could be used as indicators of
32 eutrophication, e.g. deposition rates of over $20 \text{ kg N ha}^{-1} \text{ y}^{-1}$ are associated with nitrate proportions of
33 $> 41\%$ in a mineral soil (2% carbon), and with N_{rm} stocks of over 4.8 kg N ha^{-1} in an organic soil (55
34 % carbon). Both N_{rm} and nitrate proportion increased with mean annual temperature of the sampling
35 location, despite consistent incubation temperature, suggesting that increasing temperatures are likely
36 to increase the eutrophying effects of N pollution on semi-natural ecosystems.

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38 Keywords: deposition; eutrophication; mineralization; nitrate; nutrient; pollution; production

39

40 **1. Introduction**

41

42 The progressive eutrophication of terrestrial ecosystems by reactive nitrogen (N) from fertilisers and
43 atmospheric pollution has been implicated in widespread changes in productivity (Hungate et al.,
44 2003), losses of biodiversity (Phoenix et al., 2006; Bobbink et al., 2010) and declines in water quality
45 (Magee, 1982). There is strong evidence that floristic change towards more eutrophic assemblages is
46 occurring (Braithwaite et al., 2006; Maskell et al., 2010), but these changes have not been easy to
47 ascribe to N pollution, in part due to lack of clear evidence that the availability of N in soil has
48 increased. Studies of effects of N deposition rate on soil mineralisable N have shown inconsistent
49 effects in similar habitats (e.g. Rao et al., 2009; Vourlitis et al., 2007). We used a simple measure of
50 soil mineralisable N to investigate patterns of N availability across Britain, in different soil and habitat
51 types, and related these patterns to rates of atmospheric N deposition.

52

53 Large amounts of available or readily-mineralisable N in soil reflect increased plant exposure to
54 mineral N, which is likely to increase productivity in many terrestrial habitats (LeBauer and Treseder,
55 2008). This increased productivity is beneficial for ecosystem services such as agricultural or forest
56 production, and is likely to increase net carbon (C) storage, at least in the short term (Wamelink et al.,
57 2009). However, increased productivity in semi-natural habitats is likely to decrease ground-level
58 light-availability and lead to the loss of low-growing plants and associated invertebrate species
59 (Bobbink et al., 2010; Hautier et al., 2009; Wallisdevries and Van Swaay, 2006), reducing
60 biodiversity value at a landscape scale. Such changes will be more pronounced where N is the main
61 limiting factor, for example in higher-precipitation regions (Lee et al., 2010). Large proportions of
62 nitrate in mineralisable N are also associated with floristic change (Diekmann and Falkengren-Grerup,
63 1998). Increased mineral N contents, particularly when combined with high rates of nitrification, are
64 also likely to increase N leaching and reduce downstream water quality (Gundersen et al., 2006).
65 Measurements of N availability are therefore useful for several areas of research and policy.

66 Predictions made by mechanistic models of C and N cycling used at ecosystem and global scales need
67 to be tested against measures of medium-term N fluxes (Finzi et al., 2011; Magid et al., 1997).
68 Agronomic researchers require measures of N availability to predict productivity (Ros et al., 2011). In
69 semi-natural systems, niche occupancy models that predict likely species occurrence in relation to
70 changing environmental factors (Latour and Reiling, 1993; Smart et al., 2010; Sverdrup et al., 2007)
71 sometimes require abiotic indicators of eutrophication; such models are used to inform pollution
72 abatement policy such as the Convention on Long-Range Transboundary Air Pollution (de Vries et
73 al., 2010).

74

75 Plant-available N is not straightforward to define or measure, and is thus a major source of uncertainty
76 in current ecosystem models (Wamelink et al., 2002). Soil total C / N ratio has been used as an
77 indicator of N availability, as it provides some indication of the degree to which the capacity of an
78 ecosystem to absorb excess N has been exhausted (Gundersen et al., 1998). However, variation in the
79 reactive proportion of soil N means that this ratio has limited capacity to predict the onset of N
80 leaching (Rowe et al., 2006). Direct measurements of soluble N in lysimeters or by KCl extraction
81 provide a snapshot measurement of plant-available N, but such measurements are inherently variable
82 due to short-term variations in the rate of efflux (leaching and uptake) from the soluble pool.
83 Measurements of potentially mineralisable N seem likely to be more robust indicators of N
84 availability (Ros et al., 2011), for reasons explained below.

85

86 Mineralisation is the conversion of organic residues into mineral forms, initially to ammonium, and
87 then when conditions allow nitrification, to nitrate. The rate of mineralisation of N and the total stock
88 of mineralisable N in soil are likely to be affected by factors such as soil pH, temperature, moisture,
89 stocks of C and N, and the recalcitrance of this organic matter. In a study of 31 widely differing soils
90 incubated under standard conditions, the stock of potentially mineralisable N was found to be highly
91 variable, although the proportion of this stock that was mineralised per week was similar (Stanford
92 and Smith, 1972). Studies on single soil types have shown that incubation temperature exerts greater
93 control over N mineralisation than water content, over ranges from 10 to 25 °C and from -30 to -1700

94 kPa (Sierra, 1997), and that increasing pH leads to decreased ammonification, but an increase in
95 nitrification (Pietri and Brookes, 2008). However, while incubation temperature effects on soil
96 organic matter mineralisation have been widely studied (von Lutzow and Kogel-Knabner, 2009),
97 effects of the typical temperature of the sampling location are less well understood.

98

99 Plants are able to take up N even from soils with no net mineralisation flux (Dyck et al., 1987), since
100 they can intercept available N before it can be re-immobilised (Schimel and Bennett, 2004). Net
101 mineralisation measurements therefore probably underestimate plant-available N, at least in low-N
102 systems. Gross mineralisation fluxes can be measured using isotopic dilution or by adsorption onto
103 strong ion-exchange resins, but these measurements probably overestimate plant-available N (Fierer
104 et al., 2001). Soluble organic forms of N may also be produced during the decomposition of organic
105 matter, and may themselves be significant sources of N for plant nutrition (Chapin et al., 1993; Hill et
106 al., 2011; Schimel and Chapin, 1996). Plant growth can also decrease or more commonly increase
107 mineralisation, with plant cultivation changing mineralisation rates to 70 – 500 % of rates in controls
108 without plants (Kuzyakov, 2002). This implies that there can be no definitive measure of plant-
109 available N, but net mineralisation measurements remain useful to distinguish soils across a range of
110 N availability (Schimel and Bennett, 2004). Net mineralisation flux has most commonly been
111 measured by comparing the amounts of extractable nitrate and ammonium before and after a period of
112 incubation, using paired soil samples (e.g. Keeney, 1980; Waring and Bremner, 1964). Disturbance
113 can change mineralisation and immobilisation rates, so *in situ* or intact core methods are preferred
114 (Raison et al., 1987). As well as introducing some error due to analysis of two spatially separated
115 cores, the paired core method for measuring net mineralisation may be unsuitable when cores cannot
116 be transferred rapidly to the laboratory under controlled conditions, since mineralisation in transit is
117 likely to lead to large variation in mineral N contents on arrival at the laboratory. In the current study
118 we therefore used a single-extraction method, in which soils were flushed through before incubation
119 with approximately four pore-volumes of an artificial rain solution, to remove any accumulation of
120 mineral N during transit. Mineralisable N measured using this method helped explain the occurrence
121 of plant species in associated plots, predicting a component of the variation in mean Ellenberg “N”

122 score (Ellenberg, 1974) that was largely orthogonal to soil properties such as pH, moisture content
123 and total N/C ratio (Rowe et al., 2011).

124

125 Different plant species may be adapted to use oxidised, reduced or dissolved organic N (Miller and
126 Bowman, 2003). Dissolved organic N uptake may become less prevalent and nitrate uptake more
127 prevalent in more productive systems (Nordin et al., 2001), perhaps because competitively superior
128 species are able to take up whichever is currently the most abundant form (Ashton et al., 2010). The
129 availability of nitrate may be more important than total available N for explaining species occurrence
130 (Andrianarisoa et al., 2009; Bengtson et al., 2006; Wilson et al., 2005). Soil nitrate concentrations
131 tend to increase with N enrichment (Corre et al., 2007), perhaps because nitrate immobilisation is
132 inhibited by greater ammonium concentrations (Bradley, 2001). High nitrate concentrations are also
133 associated with greater N leaching losses (MacDonald et al., 2002).

134

135 The effects of atmospheric N deposition on N mineralisation rates have been studied in selected
136 ecosystems, with varying results. In a study of sixteen sites in Californian deserts, N deposition was
137 found to be correlated with soil total C and total N, but mineralisation rates were unrelated to N
138 deposition (Rao et al., 2009). However, a similar study of semiarid Californian shrublands found that
139 N mineralisation increased linearly with N deposition rate (Vourlitis et al., 2007). Observed
140 relationships between mineralisable N and N deposition were positive in southern Swedish deciduous
141 forests (Falkengren-Grerup et al., 1998), Appalachian deciduous forests (Boggs et al., 2005) and pine
142 stands in Nanchang, China (Chen et al., 2010); negative in sugar maple stands in Ontario (Watmough,
143 2010); unrelated in pine stands in Alberta (Laxton et al., 2010); and unimodal for spruce stands across
144 Germany (Corre et al., 2007). A study of forest plots across the northeastern US showed a positive
145 relationship between mineralisable N and N deposition in maple stands, but no relationship in beech
146 stands (Lovett and Rueth, 1999). This variation suggests the need for more studies in which the same
147 survey and analytical techniques are used across different habitats, to clarify whether there are indeed
148 differences in responses to N deposition, and to explore potential reasons for these differences (Nave
149 et al., 2009).

150

151 The aim of the current study was to examine variation in soil net N mineralisation and net nitrification
152 across a range of British habitats in relation to soil properties, habitat type, temperature of the
153 sampling site, and the gradient of N deposition, to address the hypothesis that increased N deposition
154 leads to increases in available N and in the nitrate proportion of this available N.

155

156 **2. Methods**

157

158 Soil cores for analysis were taken in summer 2007 during the UK Countryside Survey, a large
159 stratified random survey of 1 km² squares across Britain, i.e. England, Wales and Scotland (Firbank et
160 al., 2003). The stratification is based on 32 land use classes, each sampled using eight squares, giving
161 a total of 256 squares. The survey has been repeated five times since 1978, and has expanded, but the
162 current study was restricted to the original set of squares for which there is a long history of repeat
163 measurements. Samples for mineralisable N were taken from three of the five randomly located main
164 plots in each square. Access to some sites was restricted, however, and of the planned 768 analyses
165 only 665 were carried out, from plots located within 237 of the squares. In the Countryside Survey,
166 the squares were mapped in terms of “Broad Habitat” on the basis of floristic and structural
167 characteristics (Maskell et al., 2008), meaning that each sample could be related to a specific Broad
168 Habitat (Table 1).

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Table 1. Number of mineralisable nitrogen analyses carried out per Broad Habitat. I = habitat assumed to be intensively managed; E = habitat assumed to be extensively managed.

Broad Habitat	N	Broad Habitat	N
Improved Grassland (I)	149	Fen, Marsh and Swamp (E)	12
Arable and Horticulture (I)	148	Bracken (E)	6
Bog (E)	78	Urban (E)	4
Neutral Grassland (I)	76	Littoral Sediment (E)	3
Acid Grassland (E)	56	Calcareous Grass (E)	2
Coniferous Woodland (E)	48	Supralittoral Rock (E)	2
Dwarf Shrub Heath (E)	42	Supralittoral Sediment (E)	2
Broadleaf, Mixed and Yew Woodland (E)	37		

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Coarse litter was removed before sampling, and soil cores were taken by pressing a 5 cm diameter by 15 cm long plastic pipe into the soil until the end was level with the soil surface. The plastic tube was carefully extracted, and cores were returned to the laboratory by normal post, taking 1-5 days. Cores were kept at 4 °C for a further 1-5 days until sufficient cores had been received for an analytical batch. Mineralisable N analyses were carried out after first flushing out soil solution by laying the core horizontally on a perforated rack and repeatedly spraying with a dilute salts solution, then incubating for 28 days at 10 °C, by extracting mineral N from the incubated core using 1M KCl. This procedure was designed to reduce variability in initial mineral N concentrations due to pre-sampling rain events and uncertain conditions during transfer to the lab, and was described in detail in Rowe et al. (2011), except that the flushing solution was based on concentrations of major ions except ammonium and nitrate in average UK rain in 2007 as estimated using the FRAME model (Rognvald Smith, CEH Edinburgh, pers com.). These concentrations were: 17.6 µeq Ca²⁺ L⁻¹; 30.1 µeq Mg²⁺ L⁻¹; 125 µeq Na⁺

188 L^{-1} ; $140 \mu eq Cl^{-1} L^{-1}$ and $57.2 \mu eq SO_4^{2-} L^{-1}$, resulting in a solution with a pH of approximately 4.6. The
189 total net mineral N production during the incubation (N_{mm}) was expressed as $kg N ha^{-1}$ in the top 15
190 cm of soil, using bulk density measurements made on soil cores taken from adjacent locations. This
191 unit was chosen for two reasons. Firstly, the rate of mineralisation of N in a given sample declines
192 with time (Stanford and Smith, 1972), so a single measurement cannot be used to calculate flux
193 during shorter or longer periods of time, but is better viewed as an indicator of the stock of readily
194 mineralisable N. Secondly, since soils vary widely in their organic C content, expressing
195 mineralisable N concentrations per g soil or per g organic matter gives the impression of high
196 availability on mineral or organic soils, respectively. The stock of available N in the top 15 cm of soil,
197 by contrast, is a measure of N availability within the plant rooting zone that is comparable across a
198 variety of habitats.

199

200 Nitrification was calculated as the net nitrate production during the incubation, and was expressed as a
201 proportion of N_{mm} rather than as a total amount, to separate this signal from that of the overall quantity
202 of mineralisable N. After incubation, a subsample was analysed for total C content by mass loss on
203 ignition ($375^{\circ}C$ for 16 hours) using a ratio of 0.55, which was the mean ratio of elementally analysed
204 C to loss-on-ignition in the main Countryside Survey dataset (Emmett et al., 2010). Soil pH was
205 measured in samples from adjacent soil cores, in a slurry of 10 g fresh soil with 25 ml de-ionised
206 water. Soil moisture content was measured gravimetrically in samples from adjacent soil cores and
207 expressed as % of fresh weight.

208

209 Estimates of atmospheric N deposition fluxes were obtained using the CBED model (Smith et al.,
210 2000), which predicts fluxes based on atmospheric concentrations, fertiliser application rates, and the
211 interception characteristics of vegetation. Deposition estimates for woodland were used for woodland
212 habitats, and deposition estimates for open moorland were used for all other habitats. Effects of N
213 deposition were not examined within habitats as defined the Countryside Survey (Maskell et al.,
214 2008) that were considered to be intensively managed (Improved grassland, Neutral grassland, and
215 Arable), but only within extensively managed habitats where little or no N fertiliser is likely to have

216 been applied and where more than 10 analyses were carried out, i.e., for samples from: Broadleaf,
217 mixed and yew woodland; Coniferous woodland; Acid grassland; Dwarf Shrub Heath, Fen, marsh and
218 swamp, and Bog. Mean annual temperature for each Countryside Survey square was estimated as the
219 average of monthly average air temperatures in the years preceding the survey, 2001-2006 (Met
220 Office, 2009).

221

222 Correlations between variables were analysed using Spearman's rank-correlation test. Linear mixed-
223 effects models were fitted to N_{rm} stock and nitrate proportion data by maximum likelihood (ML) using
224 the lme procedure of R (Pinheiro & Bates 2004; R Development Core Team, 2007). The Countryside
225 Survey square was included as a random effect. Effects of Broad Habitat and N deposition rate on N_{rm}
226 stock and nitrate proportion were examined by fitting these two explanatory variables and the
227 interaction between them as fixed effects. Effects of continuous variables (N deposition, annual mean
228 temperature, soil C content and soil pH) on N_{rm} stock and nitrate proportion were examined by fitting
229 these variables and interactions among them as fixed effects. In both cases, a maximal model
230 including all interactions was fitted, and terms were then removed in ascending order of influence on
231 model likelihood, until further simplification caused an increase in Akaike's information criterion
232 (AIC). To reduce heteroscedasticity, stock data were log transformed before analysis, first adding half
233 the detection limit to zero values, and nitrate proportion was logit transformed, first adding half the
234 detection limit to zero values and subtracting half the detection limit from values of one. Nitrate
235 proportions could not be calculated for samples with no detectable N_{rm} . Back-transformed means and
236 standard errors are presented.

237

238 **3. Results**

239

240 **3.1 Mineralisable N stock and nitrate proportion**

241

242 The log-average N_{rm} stock measured across all British soils was 8.8 kg ha^{-1} in 0-15 cm depth soil. The
243 distribution of N_{rm} by Broad Habitat is shown in Figure 1a. The measurement clearly distinguished
244 habitats that are considered fertile and productive from those considered unproductive, although
245 variability was greater for the 'Broadleaf, mixed and yew woodland' and 'Fen, marsh and swamp'
246 habitats, both of which can occur on a wide range of soil types in Britain. The intensively managed
247 habitats 'Arable and Horticulture' and 'Improved Grassland' had consistently large N_{rm} stocks, and
248 Bog and 'Dwarf shrub heath' had consistently small stocks.

249

250 The mean proportion of nitrate in N_{rm} across all British soils was $0.52 \text{ g NO}_3\text{-N g}^{-1}$ total mineralisable
251 N, and there was considerable variation in nitrate proportion among Broad Habitats (Figure 1b). The
252 greatest proportion of nitrate was in the Arable and Horticulture habitat, and there were small nitrate
253 proportions in less fertile habitats such as Bog, Acid Grassland and Dwarf Shrub Heath.

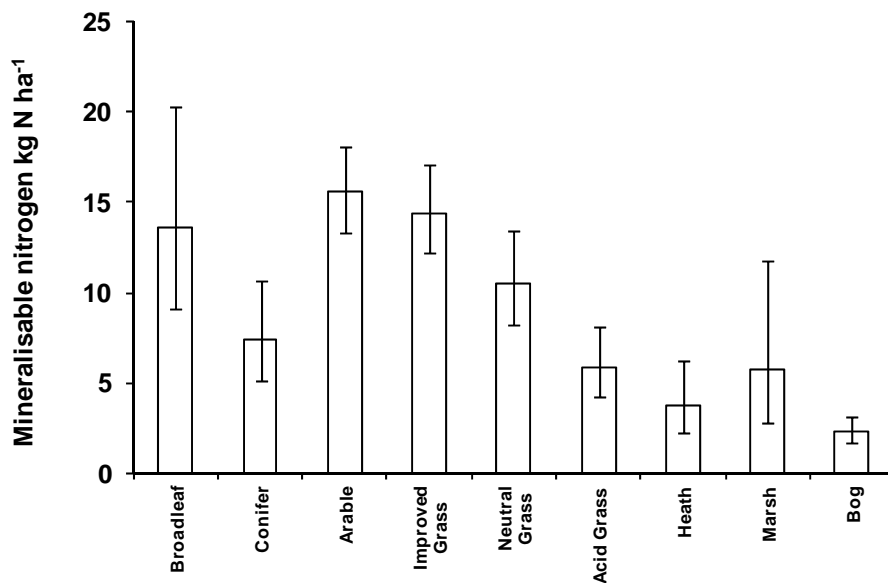
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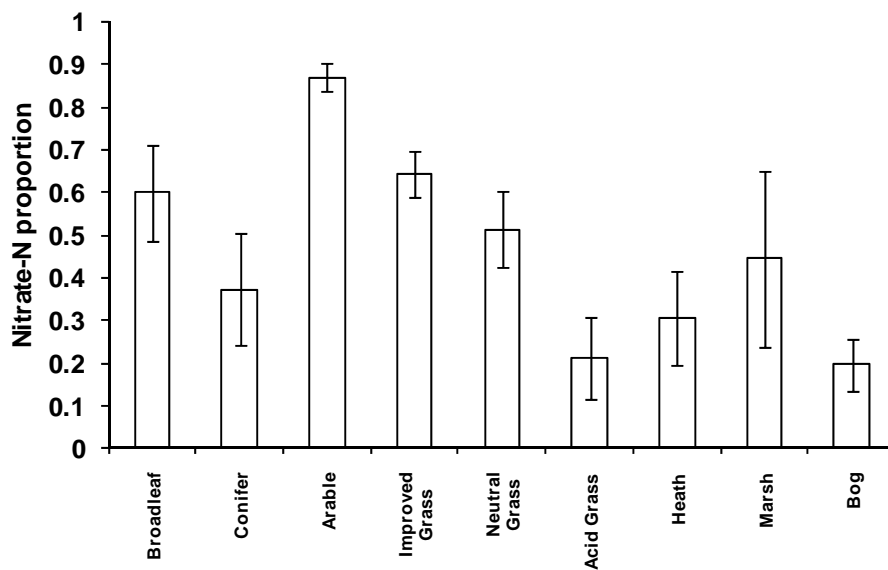
257 Figure 1: Mean (\pm one standard error) values for: a) stock of total readily-mineralisable nitrogen (kg
258 N ha⁻¹); and b) nitrate proportion of total readily-mineralisable nitrogen, in the top 15 cm of soil in
259 different Broad Habitats across Britain: Broadleaf = Broadleaf, mixed and yew woodland; Conifer =
260 Coniferous woodland; Arable = Arable and horticulture; Improved grassland; Neutral grassland; Acid
261 grassland; Heath = Dwarf shrub heath; Marsh = Fen, marsh and swamp; Bog.

262 a)



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264 b)



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266 **3.2 Factors affecting mineralisable N**

267

268 Correlation analysis in extensively managed habitats showed a close association between N_{rm} and N
269 deposition (Table 2a). The stock of N_{rm} was also strongly correlated with soil C content, moisture
270 content at sampling, and mean annual temperature. Significant correlations also illustrated spatial
271 associations, for example between higher temperatures towards the south of Britain and greater N
272 deposition rates and lower soil C contents. The proportion of nitrate in mineralisable N was positively
273 correlated with N deposition rate, mean annual temperature and soil pH, and negatively correlated
274 with soil moisture and C contents. Within intensively managed habitats, N_{rm} was not correlated with
275 N deposition rate (Table 2b). Neither N_{rm} nor nitrate proportion were correlated with mean annual
276 temperature in intensively managed habitats. The N_{rm} stock in intensively managed habitats was also
277 not correlated with intrinsic soil properties, but nitrate proportion still tended to be greater with
278 greater N_{rm} . Nitrate proportion in intensively managed habitats also increased with N deposition rate
279 and soil pH, and decreased with greater soil moisture and C contents. Since soil C content was very
280 strongly associated with soil moisture in both extensively managed and intensively managed habitats
281 (Spearman's $\rho = 0.881$ and 0.811 , respectively), and soil C content was expected to have a more
282 direct effect on N_{rm} , soil moisture was left out of subsequent regression analyses.

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285

286 Table 2. Spearman's rank correlation coefficients among readily-mineralisable N (N_{rm}), proportion
 287 nitrate in mineralisable N (pNO_3), N deposition (N_{dep}), soil total carbon (C_{tot}), soil moisture, soil pH
 288 and mean annual temperature (Temp), in: a) extensively managed habitats ($N = 290$); and b)
 289 intensively managed habitats ($N = 375$). *** = $P < 0.001$; ** = $P < 0.01$; * = $P < 0.05$; ^{ns} = $P > 0.05$.

	N_{rm}	pNO_3	N_{dep}	C_{tot}	Moisture	pH
a) extensively managed habitats						
pNO_3	0.296***					
N_{dep}	0.604***	0.280***				
C_{tot}	-0.502***	-0.402***	-0.489***			
Moisture	-0.613***	-0.390***	-0.577***	0.881***		
pH	0.110 ^{ns}	0.214***	-0.087 ^{ns}	-0.426***	-0.273***	
Temp	0.477***	0.327***	0.741***	-0.475***	-0.482***	-0.024 ^{ns}
b) intensively managed habitats						
pNO_3	0.259***					
N_{dep}	0.001 ^{ns}	-0.148**				
C_{tot}	0.016 ^{ns}	-0.298***	0.117*			
Moisture	-0.093 ^{ns}	-0.380***	0.118*	0.811***		
pH	-0.054 ^{ns}	0.274***	0.071 ^{ns}	-0.395***	-0.405***	
Temp	0.000 ^{ns}	0.045 ^{ns}	0.246***	-0.041 ^{ns}	-0.053 ^{ns}	0.240***

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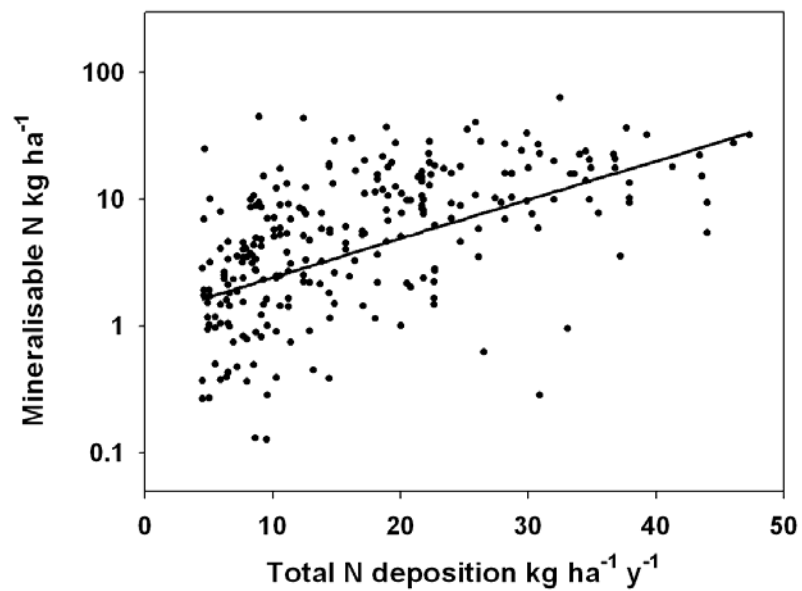
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293 Within extensively managed habitats, there was an increase in the stock of N_{rm} with more N
 294 deposition ($P < 0.001$; Figure 2). Neither the intercept nor the slope of the fitted relationship differed
 295 among habitats ($P > 0.05$). The nitrate proportion in N_{rm} also increased with total N deposition ($P <$
 296 0.001), and there were significant differences among habitats in the intercept ($P < 0.05$), but not the
 297 slope ($P > 0.05$) of this relationship (Figure 3).

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300 Figure 2. Response of total readily-mineralisable N stock to total N deposition.



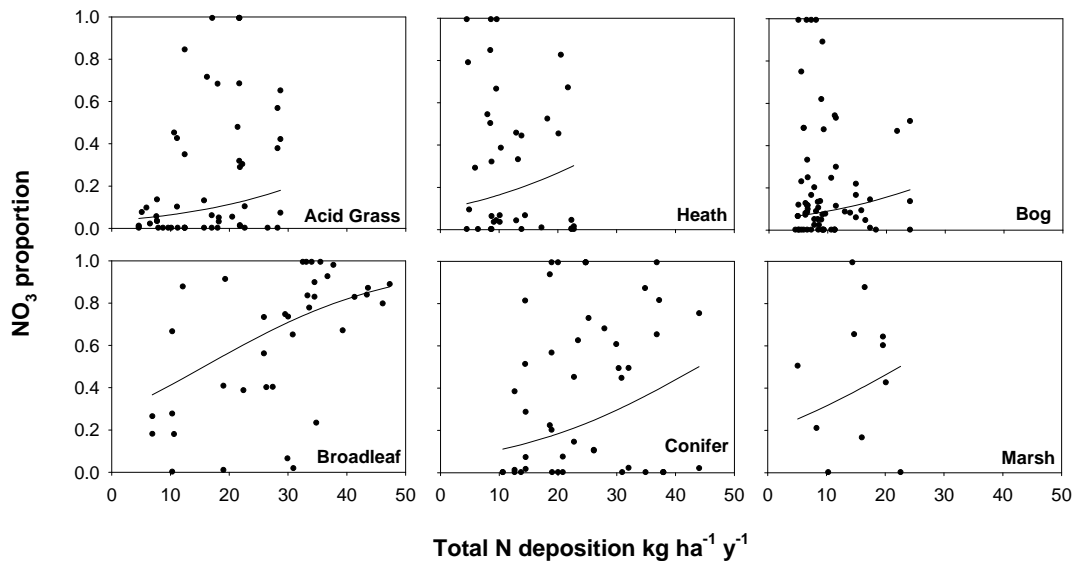
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305 Figure 3. Responses of the proportion of nitrate in total readily-mineralisable N stock to total N
306 deposition in selected extensively managed Broad Habitats: Broadleaf = Broadleaf, mixed and yew
307 woodland; Conifer = Coniferous Woodland; Acid Grass = Acid Grassland; Heath = Dwarf Shrub
308 Heath; Marsh = Fen, Marsh and Swamp; Bog = Bog. Lines are from a linear mixed model fit to logit-
309 transformed data, with different intercepts for different Broad Habitats.



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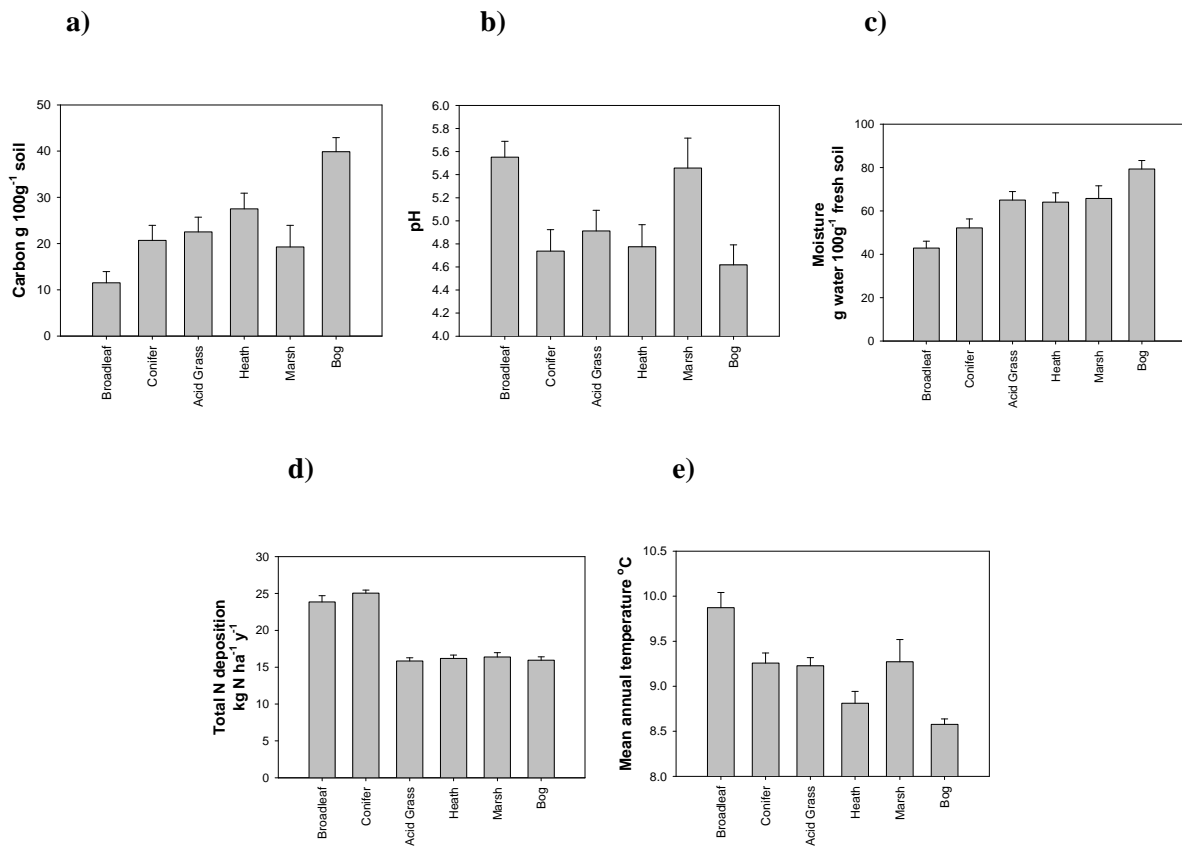
312 Potential explanatory variables for variation in N_{rm} were analysed for the subset of plots from
313 extensively managed habitats that were included in the current study. The extensive Broad Habitats
314 differed in their mean soil C content ($P < 0.001$; Figure 4a), soil pH ($P < 0.001$; Figure 4b), soil
315 moisture content at sampling ($P < 0.001$; Figure 4c), N deposition rate ($P < 0.001$; Figure 4d) and
316 annual mean temperature ($P < 0.001$; Figure 4e). The best model for N_{rm} based on continuous
317 measurements (rather than habitat category) is given in Table 3, and illustrated in Figure 5. The main
318 explanatory factors for N_{rm} were soil C ($P < 0.001$), mean annual temperature ($P < 0.001$), and N
319 deposition ($P < 0.001$). Interactions between soil C and N deposition ($P = 0.062$; Figure 6a) and
320 between soil C and soil pH ($P = 0.252$; Figure 6b) were retained in the model, since removal of these
321 terms increased AIC. The nitrate proportion of N_{rm} was best predicted (Table 4) by soil C ($P < 0.001$),
322 mean annual temperature ($P < 0.001$), and interactions between soil C and total N deposition ($P <$
323 0.05 ; Figure 7a) and between soil C and mean annual temperature ($P < 0.05$; Figure 7b).

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325

326 Figure 4. Soil properties, N deposition and temperature for plots from extensively-managed broad
327 habitats included in the current study: a) soil total carbon content; b) soil pH; c) soil moisture content;
328 d) total N deposition; and e) mean annual air temperature. Broadleaf = Broadleaf, mixed and yew
329 woodland; Conifer = Coniferous Woodland; Acid Grass = Acid Grassland; Heath = Dwarf Shrub
330 Heath; Marsh = Fen, Marsh and Swamp; Bog = Bog.

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339 Table 3. ANOVA table for fixed effects in a linear mixed-effects model predicting \log_{10} (readily-
340 mineralisable N, kg ha^{-1} in 0-15 cm soil) in extensively managed habitats, from soil total carbon
341 content (C_{tot} , $\text{g C } 100 \text{ g}^{-1}$ dry soil), soil pH, nitrogen deposition rate (N_{dep} , $\text{kg ha}^{-1} \text{ y}^{-1}$) and mean annual
342 temperature (Temperature, $^{\circ}\text{C}$). F- and p- values computed for Type I (sequential) sums-of-squares;
343 numDF = numerator degrees of freedom, denDF = denominator degrees of freedom.

	Value	numDF	denDF	F-value	p-value
Intercept	-0.135	1	141	244.3	<0.001
C_{tot}	-0.0329	1	141	45.9	<0.001
Soil pH	0.0214	1	141	2.7	0.100
Temperature	0.0574	1	123	18.0	<0.001
N_{dep}	0.0155	1	141	16.0	<0.001
Soil pH : C_{tot}	0.00477	1	141	1.3	0.252
N_{dep} : C_{tot}	0.000396	1	141	3.5	0.062

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349 Table 4. ANOVA table for fixed effects in a linear mixed-effects model predicting logit(proportion
350 nitrate in mineralisable N) in extensively managed habitats, from soil total carbon content (C_{tot} , g C
351 100 g^{-1} dry soil), nitrogen deposition rate (N_{dep}) and mean annual temperature (Temperature, °C). F-
352 and p- values computed for Type I (sequential) sums-of-squares; numDF = numerator degrees of
353 freedom, denDF = denominator degrees of freedom.

	Value	numDF	denDF	F-value	p-value
Intercept	-9.24	1	130	50.8	<0.001
C_{tot}	0.0546	1	130	29.5	<0.001
Temperature	0.747	1	121	12.7	<0.001
N_{dep}	0.101	1	130	0.6	0.452
$C_{tot} : \text{Temperature}$	-0.00376	1	130	5.8	0.018
$C_{tot} : N_{dep}$	-0.00361	1	130	6.4	0.013

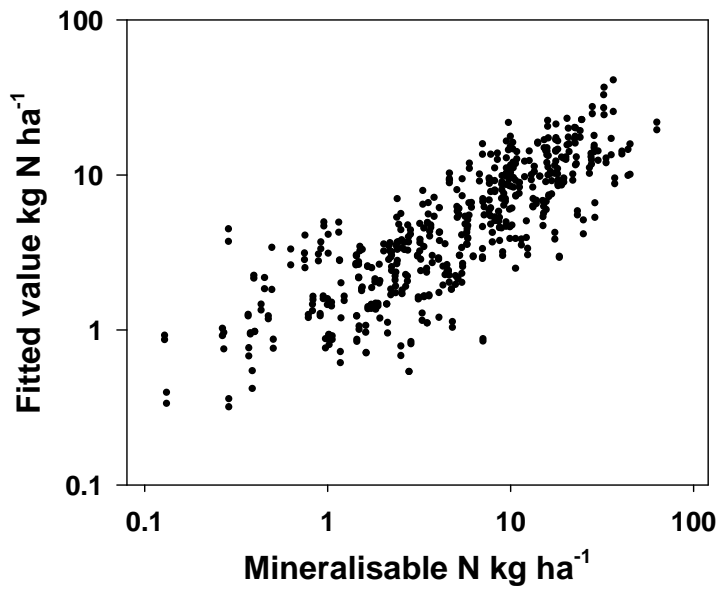
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358 Figure 5. Comparison of observed total readily-mineralisable N stock with within-group fitted values
359 from a mixed-effects model with fixed effects: $\log_{10}(\text{total readily-mineralisable N stock} + 0.07, \text{ kg N}$
360 $\text{ha}^{-1} \text{ y}^{-1}) = -0.135 - 0.0329 \times \text{soil C (\%)} + 0.0213 \times \text{soil pH} + 0.0574 \times \text{mean annual temperature (}^\circ\text{C)}$
361 $+ 0.0155 \times \text{total N deposition (kg N ha}^{-1} \text{ y}^{-1}) + 0.00477 \times \text{soil C} \times \text{soil pH} + 0.000396 \times \text{total N}$
362 $\text{deposition} \times \text{soil C}$, and Countryside Survey 1 km² square as a random effect.



363

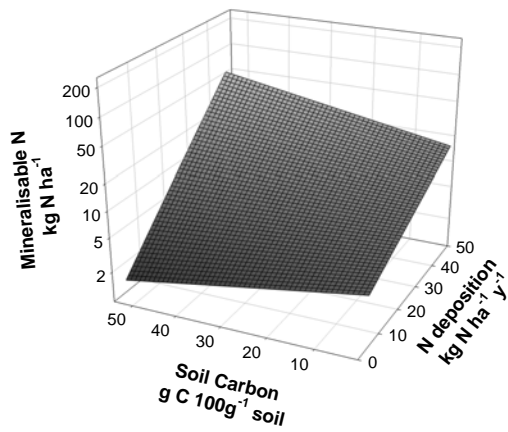
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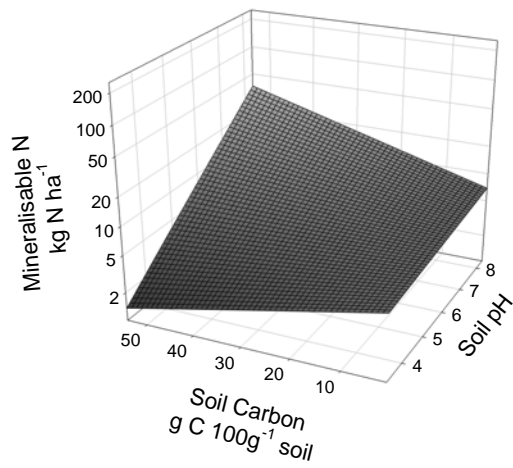
366

367 Figure 6. Fitted models for readily-mineralisable N stock in extensively managed habitats, in relation
368 to: a) soil total carbon, and total nitrogen deposition, at the mean values for pH (4.82) and annual
369 mean temperature (9.1 °C) within the dataset; and b) soil total carbon and soil pH at the mean values
370 for total nitrogen deposition (16.9 kg N ha⁻¹ y⁻¹) and mean annual temperature within the dataset.

371 **a)**



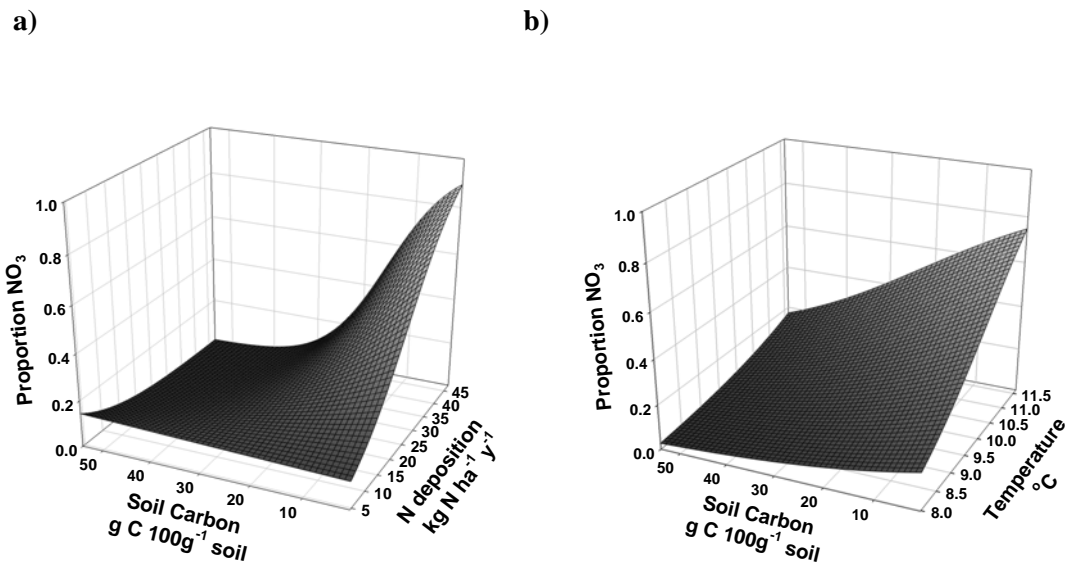
b)



374

375 Figure 7. Fitted models for nitrate proportion of mineralisable N in extensively managed habitats, in
376 relation to: a) soil total carbon and total nitrogen deposition, at the mean value for annual mean
377 temperature (9.1 °C) within the dataset; and b) soil total carbon and temperature, at the mean value for
378 total nitrogen deposition (16.9 kg N ha⁻¹ y⁻¹) within the dataset .

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385 4. Discussion

386

387 The variation among habitats in mineralisable N that was revealed in the current study is consistent
388 with the picture of greater N availability and a greater proportion of nitrate in more intensively
389 managed agricultural habitats, due to inherent soil properties, climatic differences and/or direct effects
390 of more intensive management. However, N_{rm} stocks in extensively managed habitats were of
391 comparable magnitude, particularly in the woodland and ‘Fen, Marsh and Swamp’ habitats. There
392 was considerable variability in both N_{rm} stock and nitrate proportion within individual habitats. While
393 extensively managed habitats differed significantly in their relationships between N deposition and

394 N_{rm} (Figure 2) and between N deposition and nitrate proportion (Figure 3), much unexplained
395 variance remained in these relationships.

396

397 The continuous environmental variables examined in the study were considerably more useful than
398 categorical differences among habitats for explaining variation in N_{rm} . In extensively managed
399 habitats in the current study, variation in N_{rm} was clearly related to soil characteristics, but was
400 strongly affected by N deposition rate. The N_{rm} stock increased with total N deposition, and there was
401 an interaction with soil C content. Increasing N deposition also increased N_{rm} in more mineral soils,
402 but in completely organic soils was associated with a greater increase in N_{rm} across the observed range
403 of N deposition. Larger values of N_{rm} were also associated with greater sampling location
404 temperature, implying that any increase in mean annual temperature is likely to increase N
405 availability, whether directly or by increasing the proportion of plant species with rapid growth rates
406 and more decomposable litter. Nitrate proportions were also greater in soils from warmer locations. A
407 significant negative interaction with soil C content suggests that temperature effects on nitrate
408 proportion will be more pronounced in more mineral soils.

409

410 In a meta-analysis of experimental N addition studies in north temperate forest, Nave et al. (2009)
411 found no differences between mineral and organic horizons in the response of mineralisable N to N
412 deposition, but did find differences in this response between different biogeographical regions, and
413 highlighted the importance of the proportions of recalcitrant and labile pools in soil organic matter. In
414 contrast to the current study, Booth et al. (2005) found in a meta-analysis covering a wide range of
415 ecosystems that mineralisable N was correlated with substrate concentrations of organic matter. The
416 negative correlation of N_{rm} with soil total C found in the current study may differ because many of the
417 soils had large organic matter contents (mean C content for the extensively managed habitats included
418 was 27%). The greater effect of N deposition flux on N availability in organic soils than in more
419 mineral soils shown in the current study may be because a larger proportion of the organic matter is
420 recalcitrant in the very organic soils that were included. In soils from a temperature gradient in the
421 Great Plains region, Barrett and Burke (2000) found that while C mineralisation increased with soil

422 organic matter content, gross N immobilisation also increased; a similar result to that found in the
423 current study. While the overall effect of increasing soil C content was a decrease in mineralisable N
424 stock in our study, an interaction between C content and N deposition rate suggests that this N
425 immobilisation flux may become saturated under chronically elevated N deposition. However, Hartley
426 and Mitchell (2005) found that experimental N additions increased mineralisable N more in a more
427 mineral soil (20% organic matter) than in a more organic soil (70% organic matter). This suggests that
428 there may be differences between effects observed after short-term additions and after chronic high N
429 deposition rates.

430

431 Several explanations are possible for the greater increase in N_{rm} with N deposition rate in more
432 organic soils. Proposed effects of increased N deposition include productivity stimulation (LeBauer
433 and Treseder, 2008) and inhibition of litter decomposition, at least on sites that are not greatly N-
434 limited (Craine et al., 2007; Knorr et al., 2005), either of which might increase the stock of readily-
435 mineralisable organic matter. Productivity stimulation by N may have been greater in more organic
436 soils that are generally less water-limited than mineral soils.

437

438 The proportion of nitrate in N_{rm} was strongly affected by soil C content and N deposition rate, and
439 was only large in soils with low C content and a large rate of N deposition. Nitrification is affected by
440 aeration (Sahrawat, 2008), and the texture of the soil on fine scales (e.g. clay, silt and sand fractions,
441 or the degree of humification of organic matter) and medium scales (e.g. porosity and aggregation)
442 undoubtedly affected the diffusion of air into the soil core during the incubation. However, both
443 organic and mineral soils can vary considerably in aggregation development and porosity, and hence
444 the increase in nitrate proportion with decreasing organic matter content (where there is a large rate of
445 N deposition) may not be related to effects of soil structure. The large-scale spatial pattern of nitrate
446 proportion suggests little influence of soil texture, which varies at a smaller scale. Nitrification has
447 been found in previous studies to be correlated with total N mineralised (Booth et al., 2005) and with
448 soil pH (Andrianarisoa et al., 2009; Sahrawat, 2008; Ste-Marie and Pare, 1999). We also found
449 evidence of correlation between nitrate proportion and both total N_{rm} and soil pH ($P < 0.001$ for both

450 correlations, in extensively and intensively managed habitats; Table 2). Nitrification rates may also
451 indicate the size of the nitrifying bacteria population, and hence greater nitrate proportions may be
452 related to a history of elevated N inputs. The strong increase in nitrate proportion with N deposition in
453 more mineral soils suggests that N deposition has increased nitrifier activity in these soils, whereas
454 factors such as limited aeration may have prevented an increase in nitrifiers in more organic soils.

455

456 The N_{m} measurement reflects an amount of N that was insoluble at the start of the study but was
457 readily mineralised during the incubation. The net N mineralisation during an equivalent period under
458 field conditions would likely have been different, due to differences in disturbance, temperature,
459 aeration, interactions with plant roots, and other factors. The measurement nevertheless provides some
460 indication of the rate of N release from soil organic matter into the soil solution, whence it may be
461 available for plant uptake, or may be leached. Chen et al. (2006) found that gross N mineralisation
462 remained elevated 14 years after cessation of N additions, despite recovery of mineral N
463 concentrations and leaching rates. Although large amounts of N in readily-mineralisable organic
464 matter are not as immediate a cause for concern (in semi-natural systems susceptible to
465 eutrophication) as are large mineral N concentrations in soil solution, they reflect a pool of N that is
466 likely to lead to long-term increases in plant production and/or increased leaching of mineral N.

467

468 **5. Conclusions**

469

470 In extensively managed habitats, mineralisable N stock and nitrate proportion of mineral N were both
471 strongly influenced by N deposition rate, and by interactions with soil C content. Habitats varied in
472 mean mineralisable N stock, but did not show evidence of differential effects of N deposition, perhaps
473 due to variation in soil type within each habitat. The effect of N deposition on mineralisable N stock
474 was more apparent in more organic soils, whereas the effect on nitrate proportion was more apparent
475 in more mineral soils. With the proviso that responses also depend on soil C content and site
476 temperature, the study supports the use of both mineralisable N and nitrate proportion as indicators of

477 ecosystem eutrophication due to N pollution. The increase in mineralisable N stock with temperature
478 implies that climate change and N deposition are likely to have synergistic effects, accelerating the
479 change of semi-natural habitats to a more eutrophic state.

480

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482

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487

488 **7. References**

- 489 Andrianarisoa KS, Zeller B, Dupouey JL, Dambrine E. Comparing indicators of N status of 50 beech
490 stands (*Fagus sylvatica* L.) in northeastern France. *Forest Ecology and Management* 2009;
491 257: 2241-2253.
- 492 Ashton IW, Miller AE, Bowman WD, Suding KN. Niche complementarity due to plasticity in
493 resource use: plant partitioning of chemical N forms. *Ecology* 2010; 91: 3252-3260.
- 494 Barrett JE, Burke IC. Potential nitrogen immobilization in grassland soils across a soil organic matter
495 gradient. *Soil Biology & Biochemistry* 2000; 32: 1707-1716.
- 496 Bengtson P, Falkengren-Grerup U, Bengtsson G. Spatial distributions of plants and gross N
497 transformation rates in a forest soil. *Journal of Ecology* 2006; 94: 754-764.
- 498 Bobbink R, Hicks K, Galloway J, Spranger T, Alkemade R, Ashmore M, Bustamante M, Cinderby S,
499 Davidson E, Dentener F, Emmett B, Erisman JW, Fenn M, Gilliam F, Nordin A, Pardo L, De
500 Vries W. Global assessment of nitrogen deposition effects on terrestrial plant diversity: a
501 synthesis. *Ecological Applications* 2010; 20: 30-59.
- 502 Boggs JL, McNulty SG, Gavazzi MJ, Myers JM. Tree growth, foliar chemistry, and nitrogen cycling
503 across a nitrogen deposition gradient in southern Appalachian deciduous forests. *Canadian*
504 *Journal of Forest Research-Revue Canadienne De Recherche Forestiere* 2005; 35: 1901-1913.
- 505 Booth MS Stark JM, Rastetter E. Controls on nitrogen cycling in terrestrial ecosystems: A synthetic
506 analysis of literature data. *Ecological Monographs* 2005; 75: 139-157.
- 507 Bradley RL. An alternative explanation for the post-disturbance NO_3^- flush in some forest ecosystems.
508 *Ecology Letters* 2001; 4: 412-416.
- 509 Braithwaite ME, Ellis RW, Preston CD. *Change in the British Flora 1987-2004*: Botanical Society of
510 the British Isles., 2006.
- 511 Chapin FS, Moilanen L, Kielland K. Preferential use of organic nitrogen for growth by a
512 nonmycorrhizal arctic sedge. *Nature* 1993; 361: 150-153.

513 Chen FS, Fahey TJ, Yu MY, Gan L. Key nitrogen cycling processes in pine plantations along a short
514 urban-rural gradient in Nanchang, China. *Forest Ecology and Management* 2010; 259: 477-
515 486.

516 Chen Y, Hogberg P. Gross nitrogen mineralization rates still high 14 years after suspension of N input
517 to a N-saturated forest. *Soil Biology & Biochemistry* 2006; 38: 2001-2003.

518 Corre MD, Brumme R, Veldkamp E, Beese FO. Changes in nitrogen cycling and retention processes
519 in soils under spruce forests along a nitrogen enrichment gradient in Germany. *Global Change*
520 *Biology* 2007; 13: 1509-1527.

521 Craine JM, Morrow C, Fierer N. Microbial nitrogen limitation increases decomposition. *Ecology*
522 2007; 88: 2105-2113.

523 de Vries W, Wamelink W, van Dobben H, Kros H, Reinds G-J, Mol-Dijkstra J, Smart S, Evans C,
524 Rowe E, Belyazid S, Sverdrup H, van Hinsberg A, Posch M, Hettelingh J-P, Spranger T,
525 Bobbink R. Use of dynamic soil-vegetation models to assess impacts of nitrogen deposition
526 on plant species composition and to estimate critical loads: an overview. *Ecological*
527 *Applications* 2010; 20: 60-79.

528 Diekmann M, Falkengren-Grerup U. A new species index for forest vascular plants: development of
529 functional indices based on mineralization rates of various forms of soil nitrogen. *Journal of*
530 *Ecology* 1998; 86: 269-283.

531 Dyck W, Mees C, Hodgkiss P. Nitrogen availability and comparison to uptake in two New Zealand
532 *Pinus radiata* forests. . *New Zealand Journal of Forest Science* 1987; 17: 338–352.

533 Ellenberg H. Zeigerwerte der gefasspflanzen mitteleuropas. *Scripta Geobotanica* 1974; 9: 1-97.

534 Emmett BA, Reynolds B, Chamberlain PM, Rowe E, Spurgeon D, Brittain SA, Frogbrook Z, Hughes
535 S, Lawlor AJ, Poskitt J, Potter E, Robinson DA, Scott A, Wood C, Woods C. Countryside
536 Survey: Soils report from 2007. NERC/Centre for Ecology & Hydrology, 2010, pp. 192.

537 Falkengren-Grerup U, Brunet J, Diekmann M. Nitrogen mineralisation in deciduous forest soils in
538 south Sweden in gradients of soil acidity and deposition. *Environmental Pollution* 1998; 102:
539 415-420.

540 Fierer N, Schimel J, Cates R, Zou J. The influence of balsam poplar tannin fractions on carbon and
541 nitrogen dynamics in Alaskan taiga floodplain soils. *Soil Biology and Biochemistry* 2001; 33:
542 1827–1839.

543 Finzi AC, Austin AT, Cleland EE, Frey SD, Houlton BZ, Wallenstein MD. Responses and feedbacks
544 of coupled biogeochemical cycles to climate change: examples from terrestrial ecosystems.
545 *Frontiers in Ecology and the Environment* 2011; 9: 61-67.

546 Firbank LG, Barr CJ, Bunce RGH, Furse MT, Haines-Young R, Hornung M, Howard DC, Sheail J,
547 Sier A, Smart SM. Assessing stock and change in land cover and biodiversity in GB: an
548 introduction to Countryside Survey 2000. *Journal of Environmental Management* 2003; 67:
549 207-218.

550 Gundersen P, Callesen I, de Vries W. Nitrate leaching in forest ecosystems is related to forest floor
551 C/N ratios. *Environmental Pollution* 1998; 102: 403-407.

552 Gundersen P, Schmidt IK, Raulund-Rasmussen K. Leaching of nitrate from temperate forests - effects
553 of air pollution and forest management. *Environmental Reviews* 2006; 14: 1-57.

554 Hartley SE, Mitchell RJ. Manipulation of nutrients and grazing levels on heather moorland: changes
555 in *Calluna* dominance and consequences for community composition. *Journal of Ecology*
556 2005; 93: 990-1004.

557 Hautier Y, Niklaus PA, Hector A. Competition for light causes plant biodiversity loss after
558 eutrophication. *Science* 2009; 324: 636-638.

559 Hill PW, Quilliam RS, DeLuca TH, Farrar J, Farrell M, Roberts P, Newsham KK, Hopkins DW,
560 Bardgett RD, Jones DL. Acquisition and assimilation of nitrogen as peptide-bound and D-
561 enantiomers of amino acids by wheat. *PLoS ONE* 2011; 6: e19220.

562 Hungate BA, Dukes JS, Shaw MR, Luo YQ, Field CB. Nitrogen and climate change. *Science* 2003;
563 302: 1512-1513.

564 Keeney DR. Prediction of soil nitrogen availability in forest ecosystems: a review. *Forest Science*
565 1980; 26: 159-171.

566 Knorr M, Frey SD, Curtis PS. Nitrogen additions and litter decomposition: A meta-analysis. *Ecology*
567 2005; 86: 3252-3257.

568 Kuzyakov Y. Review: Factors affecting rhizosphere priming effects. *Journal of Plant Nutrition and*
569 *Soil Science-Zeitschrift Fur Pflanzenernahrung Und Bodenkunde* 2002; 165: 382-396.

570 Latour JB, Reiling R. MOVE: a multiple-stress model for vegetation. *The Science of the Total*
571 *Environment Supplement* 1993: 1513-1526.

572 Laxton DL, Watmough SA, Aherne J, Straker J. An assessment of nitrogen saturation in *Pinus*
573 *banksiana* plots in the Athabasca Oil Sands Region, Alberta. *Journal of Limnology* 2010; 69:
574 171-180.

575 LeBauer DS, Treseder KK. Nitrogen limitation of net primary productivity in terrestrial ecosystems is
576 globally distributed. *Ecology* 2008; 89: 371-379.

577 Lee M, Manning P, Rist J, Power SA, Marsh C A global comparison of grassland biomass responses
578 to CO₂ and nitrogen enrichment. *Philosophical Transactions of the Royal Society B-*
579 *Biological Sciences* 2010; 365: 2047-2056.

580 Lovett GM, Rueth H. Soil nitrogen transformations in beech and maple stands along a nitrogen
581 deposition gradient. *Ecological Applications* 1999; 9: 1330-1344.

582 MacDonald JA, Dise NB, Matzner E, Armbruster M, Gundersen P, Forsius M. Nitrogen input
583 together with ecosystem nitrogen enrichment predict nitrate leaching from European forests.
584 *Global Change Biology* 2002; 8: 1028-1033.

585 Magee PN. NITROGEN AS A POTENTIAL HEALTH-HAZARD. *Philosophical Transactions of the*
586 *Royal Society of London Series B-Biological Sciences* 1982; 296: 543-550.

587 Magid J, Mueller T, Jensen LS, Nielsen NE. Modelling the measurable: Interpretation of field-scale
588 CO₂ and N-mineralization, soil microbial biomass and light fractions as indicators of oilseed
589 rape, maize and barley straw decomposition. In: Cadish G, Giller KE, editors. *Driven by*
590 *nature: Plant litter quality and decomposition*. CAB International, Wallingford, 1997, pp. 349-
591 362.

592 Maskell LC, Norton LR, Smart SM, Carey PD, Murphy J, Chamberlain PM, Wood CM, Bunce RGH,
593 Barr CJ. *Field Mapping Handbook*. CS Technical Report No.1/07. Countryside Survey, 2008,
594 pp. 143.

595 Maskell LC, Smart SM, Bullock JM, Thompson K, Stevens CJ. Nitrogen deposition causes
596 widespread loss of species richness in British habitats. *Global Change Biology* 2010; 16: 671-
597 679.

598 Met Office, 2009. UKCP09: Gridded data sets of monthly values. [http://www.metoffice.gov.uk/
599 climatechange/science/monitoring/ukcp09/download/access_gd/index.html#lta](http://www.metoffice.gov.uk/climatechange/science/monitoring/ukcp09/download/access_gd/index.html#lta).

600 Miller AE, Bowman WD. Alpine plants show species-level differences in the uptake of organic and
601 inorganic nitrogen. *Plant and Soil* 2003; 250: 283-292.

602 Nave LE, Vance ED, Swanston CW, Curtis PS. Impacts of elevated N inputs on north temperate
603 forest soil C storage, C/N, and net N-mineralization. *Geoderma* 2009; 153: 231-240.

604 Nordin A, Hogberg P, Nasholm T. Soil nitrogen form and plant nitrogen uptake along a boreal forest
605 productivity gradient. *Oecologia* 2001; 129: 125-132.

606 Phoenix GK, Hicks WK, Cinderby S, Kuylenstierna JCI, Stock WD, Dentener FJ, Giller KE, Austin
607 AT, Lefroy RDB, Gimeno BS, Ashmore MR, Ineson P. Atmospheric nitrogen deposition in
608 world biodiversity hotspots: the need for a greater global perspective in assessing N
609 deposition impacts. *Global Change Biology* 2006; 12: 470-476.

610 Pietri JCA, Brookes PC. Nitrogen mineralisation along a pH gradient of a silty loam UK soil. *Soil
611 Biology & Biochemistry* 2008; 40: 797-802.

612 Pinheiro JC, Bates DM 2004. *Mixed-effects models in S and S-PLUS*. Springer. 582p.

613 R Development Core Team. *R: A language and environment for statistical computing*. R Foundation
614 for Statistical Computing, Vienna, Austria, 2007.

615 Raison RJ, Connell MJ, Khanna PK. Methodology for studying fluxes of soil mineral-N *in situ*. *Soil
616 Biology & Biochemistry* 1987; 19: 521-530.

617 Rao LE, Parker DR, Bytnerowicz A, Allen EB. Nitrogen mineralization across an atmospheric
618 nitrogen deposition gradient in Southern California deserts. *Journal of Arid Environments*
619 2009; 73: 920-930.

620 Ros GH, Temminghoff EJM, Hoffland E. Nitrogen mineralization: a review and meta-analysis of the
621 predictive value of soil tests. *European Journal of Soil Science* 2011; 62: 162-173.

622 Rowe EC, Emmett BA, Smart SM, Frogbrook ZL. A new net mineralizable nitrogen assay improves
623 predictions of floristic composition. *Journal of Vegetation Science* 2011; 22: 251-261.

624 Rowe EC, Evans CD, Emmett BA, Reynolds B, Helliwell RC, Coull MC, Curtis CJ. Vegetation type
625 affects the relationship between soil carbon to nitrogen ratio and nitrogen leaching. *Water Air
626 and Soil Pollution* 2006; 177: 335-347.

627 Sahrawat KL. Factors affecting nitrification in soils. *Communications in Soil Science and Plant
628 Analysis* 2008; 39: 1436-1446.

629 Schimel JP, Bennett J. Nitrogen mineralization: Challenges of a changing paradigm. *Ecology* 2004;
630 85: 591-602.

631 Schimel JP, Chapin FS. Tundra plant uptake of amino acid and NH₄⁺ nitrogen in situ: Plants compete
632 well for amino acid N. *Ecology* 1996; 77: 2142-2147.

633 Sierra J. Temperature and soil moisture dependence of N mineralization in intact soil cores. *Soil
634 Biology & Biochemistry* 1997; 29: 1557-1563.

635 Smart SM, Scott WA, Whitaker J, Hill MO, Roy DB, Critchley CN, Marini L, Evans C, Emmett BA,
636 Rowe EC, Crowe A, Le Duc M, Marrs RH. Empirical realised niche models for British higher
637 and lower plants - development and preliminary testing. *Journal of Vegetation Science* 2010;
638 21: 643-656.

639 Smith RI, Fowler D, Sutton MA, Flechard C, Coyle M. Regional estimation of pollutant gas
640 deposition in the UK: model description, sensitivity analyses and outputs. *Atmospheric
641 Environment* 2000; 34: 3757-3777.

642 Stanford G, Smith SJ. Nitrogen mineralization potentials of soils. *Soil Science Society of America
643 Proceedings* 1972; 36: 465-472.

644 Ste-Marie C, Pare D. Soil, pH and N availability effects on net nitrification in the forest floors of a
645 range of boreal forest stands. *Soil Biology & Biochemistry* 1999; 31: 1579-1589.

646 Sverdrup H, Belyazid S, Nihlgård B, Ericson L. Modelling change in ground vegetation response to
647 acid and nitrogen pollution, climate change and forest management in Sweden 1500-2100
648 A.D. *Water, Air, and Soil Pollution: Focus* 2007; 7: 163-179.

649 von Lutzow M, Kogel-Knabner I. Temperature sensitivity of soil organic matter decomposition-what
650 do we know? *Biology and Fertility of Soils* 2009; 46: 1-15.

651 Vourlitis GL, Zorba G, Pasquini SC, Mustard R. Chronic nitrogen deposition enhances nitrogen
652 mineralization potential of semiarid shrubland soils. *Soil Science Society of America journal*
653 2007; 71: 836-842.

654 Wallisdevries MF, Van Swaay CAM. Global warming and excess nitrogen may induce butterfly
655 decline by microclimatic cooling. *Global Change Biology* 2006; 12: 1620-1626.

656 Wamelink GWW, Joosten V, van Dobben HF, Berendse F. Validity of Ellenberg indicator values
657 judged from physico-chemical field measurements. *Journal of Vegetation Science* 2002; 13:
658 269-278.

659 Wamelink GWW, van Dobben HF, Mol-Dijkstra JP, Schouwenberg E, Kros J, de Vries W, Berendse
660 F. Effect of nitrogen deposition reduction on biodiversity and carbon sequestration. *Forest*
661 *Ecology and Management* 2009; 258: 1774-1779.

662 Waring SA, Bremner JM. Ammonium production in soil under waterlogged conditions as an index of
663 nitrogen availability. *Nature* 1964; 104: 951-952.

664 Watmough SA. An assessment of the relationship between potential chemical indices of nitrogen
665 saturation and nitrogen deposition in hardwood forests in southern Ontario. *Environmental*
666 *Monitoring and Assessment* 2010; 164: 9-20.

667 Wilson SM, Pyatt DG, Ray D, Malcolm DC, Connolly T. Indices of soil nitrogen availability for an
668 ecological site classification of British forests. *Forest Ecology and Management* 2005; 220:
669 51-65.

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671