



Title	Estimation of paleo-firing temperatures using luminescence signals for the volcanic lava baked layer in Datong, China
Author(s)	Liu, Z; Zhao, H; Wang, CM; Li, SH
Citation	Quaternary Geochronology, 2015, v. 30 n. pt. B, p. 363-368
Issued Date	2015
URL	http://hdl.handle.net/10722/216830
Rights	Copyright © 2015 Elsevier B.V.; This work is licensed under a Creative Commons Attribution-NonCommercial-NoDerivatives 4.0 International License.

Estimation of Paleo-firing Temperatures Using Luminescence Signals for the Volcanic Lava Baked Layer in Datong, China

Zhe Liu^a, Hua Zhao^{a*}, Cheng-Min Wang^a, Sheng-Hua Li^{a, b}

^a Institute of Hydrogeology and Environmental Geology, CAGS, Shijiazhuang 050061, China

^b Department of Earth Sciences, The University of Hong Kong, Pokfulam Road, Hong Kong, China

*Corresponding author. Tel: +86 31167598578. E-mail address: zhaohua65@163.com (Hua Zhao)

Abstract: Eight paleo-fired samples from the baked layer in different depths under the lava and one unfired sample were collected from Datong, China. Fine quartz grains (4-11 μm) from samples were used for probing into relationship between luminescence signals and paleo-firing temperatures. Findings from the re-firing experiments indicated that using thermoluminescence (TL) and optically stimulated luminescence (OSL) sensitivity changes could estimate the paleo-firing temperatures of samples: (1) 110 °C TL sensitivity change rate against the re-firing temperature can tell whether the sample has been fired to temperatures above 500 °C or not; (2) 150 °C TL sensitivity against the re-firing temperature can indicate whether the sample has been fired to temperatures above 900 °C or not; (3) the more specific paleo-firing temperatures can be estimated by comparing the ratio of OSL and 150 °C TL sensitivities against re-firing temperatures. Results showed that the paleo-firing temperatures of the eight lava-baked samples decreased exponentially with the distance from the lava. Based on the estimated temperature profile, the temperature of the lava was estimated to be about 1100 °C.

Key words: Datong volcano; baked layer; temperature; luminescence; sensitivity; fine grain quartz

1. Introduction

Deducing the temperatures of the lava flows is of great significance for analyzing the flowing process of lava (Griffiths, 2000; Pinkerton et al., 2002), which helps to define the extents and degree affected by volcanoes. The paleo-firing temperature estimation of the baked layer beneath the lava can provide a possibility for assessing the temperature of the lava above. The estimation of paleo-firing temperatures also provides the information of the

22 degree of zeroing for heated samples, which is essential for luminescence dating.

23 Several studies showed that luminescence sensitivity changes of quartz were related to the thermal history
24 (Bøtter-Jensen et al., 1995; Wintle and Murray, 1999; Han et al., 2000; Poolton et al., 2000; Schilles et al., 2001; Li,
25 2002; Polymeris et al., 2006; Oniya et al, 2012). Therefore, many attempts using TL and OSL sensitizations of
26 quartz were made to acquire paleo-firing temperatures. Sunta and David (1982) measured the sensitivity of the 110
27 °C TL peak before and after the application of the pre-dose at different re-firing temperatures. They observed that
28 for pre-fired samples the ratio of the latter to the former remained constant until the firing temperature was attained.
29 The ratio began to increase significantly when the heating temperature was higher. Göksu et al. (1989) applied this
30 method to determine the ancient heat treatment of flint. However, Watson and Aitken (1985) observed that the
31 procedure developed by Sunta and David (1982) was not generally applicable. Their studies showed that the
32 sensitivity of the 110 °C and 230°C TL peak, respectively, may be associated with the maximum temperature the
33 quartz had experienced. Therefore, to some extent, the quartz can preserve a memory of its thermal history.
34 Polymeris et al. (2007) observed a relationship between TL and OSL sensitivities, and concluded that they could be
35 used to assess paleo-firing temperatures. It was reported that samples that have been heated to different
36 temperatures will display different in the plots of the sensitivity versus the re-firing temperature. However, most of
37 the studies were carried out on archaeological samples with small palaeo-doses. The potential for assessing the
38 paleo-firing temperatures based on luminescence sensitivity still needs further studies, particularly for geological
39 samples.

40 In this paper, we implemented experiments to acquire the relationship between the luminescence sensitivity
41 and firing temperatures. Methods determining the paleo-firing temperatures were established and applied to the
42 samples from a lava-baked layer in Datong, China.

43 2. Samples and equipment

44 Datong volcanic group is located in the Datong Basin in the northern part of China. More than 30 volcanoes
45 erupted in the area of about 900 km². The volcanic eruptions in this region lie between Quaternary loess layers or
46 lacustrine sediments. Samples used in this study were collected from Yujiashai profile. The lava in this area has an
47 age of about 300 ka (Li and Sun, 1984). The lacustrine sediments that have been baked by the lava flow had a
48 distinct red color compared with the unfired layer. Eight paleo-fired samples from the lava-baked layer and one
49 unfired sample beneath the baked layer (No: YJZC-7) were collected (Fig. 1). The unfired sample was used to heat
50 to different known temperatures for the investigation of the link between the luminescence sensitization and firing
51 temperatures in the laboratory. A portion of the unfired sample was used as a reference for comparisons.

52 Sample preparation was carried out under red light in the laboratory. All the samples were treated with 30%
53 H₂O₂ and 30% HCL to remove the organic materials and the carbonates. They were treated with 30% fluorosilicic
54 acid for 5 days for the removal of feldspar contamination. After that, fine grains in the range 4-11 μm were selected
55 according to Stoke's Law using acetone and deposited on stainless steel discs. The purity of the quartz was checked
56 with infrared stimulation (Li et al., 2002).

57 Luminescence measurements were performed using a Daybreak 2200 TL/OSL automatic system. A ⁹⁰Sr/⁹⁰Y
58 beta source, delivering ~0.05 Gy s⁻¹, to quartz in stainless steel discs, was used for irradiation. The luminescence
59 signals were detected through two 3 mm U-340 filters in all observations. A SX2-10-13 muffle oven with the
60 maximum temperature of 1350 °C was used for heating the samples to different temperatures in the laboratory.

61 Throughout the paper, laboratory heated samples are referred to as fired samples, the temperatures as the firing
62 temperatures. The samples from the lava-baked layer are called paleo-fired samples, heated to paleo-firing
63 temperatures. Heat treatment to the fired or paleo-fired samples in the laboratory is referred to as re-firing.

64 3. Experimental procedures

65 A portion of the unfired sample was heated to different temperatures. The unfired fine-grain quartz sample was
66 divided into 12 subsamples. Eleven were heated to various temperatures ranging from 100 °C to 1100 °C in steps of
67 100 °C for 15 minutes in the muffle oven, and then cooled naturally. The sample without any thermal treatment was
68 used as a reference for comparison.

69 The 12 fired samples and the 8 paleo-fired samples were used for re-firing experiments. Each sample was
70 divided into 8 subsamples for re-firing at temperatures ranging from 300 °C to 1000 °C in steps of 100 °C for 15
71 minutes in the muffle oven. Then, thermoluminescence (TL) and optically stimulated luminescence (OSL) signals
72 induced by a dose of 10 Gy were measured. A heating rate of 5 °C/s was used for the TL measurements to 260 °C.
73 The OSL signals were measured at 125 °C for 100 seconds. The sensitivity of 110 °C TL, 150 °C TL, and OSL
74 signals were obtained.

75 4. Results

76 4.1 110 °C TL sensitivity against re-firing temperature

77 The 110 °C TL sensitivity against re-firing temperature of fired samples are illustrated in Fig. 2a. For samples
78 that have been heated to 500 °C and lower, the 110 °C TL sensitivity increases sharply when the re-firing
79 temperature was around 500 °C. It decreases slightly for the temperature above 500 °C. For fired samples that have
80 been heated to more than 500 °C, it remains unchanged. Therefore, if there is an obvious peak around 500 °C in the
81 110 °C TL sensitivity curve, it indicates that the sample has not been heated to temperatures higher than 500 °C. On
82 contrary, if there is no obvious peak at about 500 °C in the 110 °C TL sensitivity curve, the sample might have been
83 heated to temperatures higher than 500 °C.

84

85 We applied the observations above to the paleo-fired samples. The 110 °C TL peak sensitivity against re-firing
86 temperature of the paleo-fired samples are shown in Fig. 2b. All the curves had a peak at about 500 °C. However,
87 from field **observations** and studies, some of the samples were certainly heated to temperatures above 500 °C.
88 Studies have shown that the TL sensitivity of quartz increases proportionally to the dose received before the
89 thermal activation around 500 °C, which is called pre-dose effect (Aitken, 1985). **After** being paleo-fired, the
90 **natural samples** were naturally given large pre-doses during the process of sedimentation, before the re-firing in the
91 laboratory. **This pre-dose effect causes the TL sensitivity to increase significantly at the re-firing temperature of 500**
92 **°C for the paleo-fired samples, even for the samples have been paleo-fired at temperatures higher than 500 °C.** The
93 pre-dose effect affects the judgment on whether the **paleo-firing** temperature **was** above 500 °C or not. Therefore,
94 we **conclude** that using the 110 °C TL sensitivity alone is not **sufficient** for indicating the **paleo-firing** temperatures,
95 because it is **not only affected** by the thermal history, but also the pre-dose effect (Li, 2002).

96 4.2 110 °C TL sensitivity change rate against re-firing temperature

97 Several studies have **investigated** the pre-dose effect **in** quartz (Zimmerman, 1971; McKeever, et al., 1985;
98 Yang and McKeever, 1990; Rendell, et al., 1994; Li, 2002). The pre-dose effect is particularly important for the
99 paleo-fired samples because they are about 300 ka of age, of equivalent to a dose of 1300 Gy. The pre-dose effect
100 gives rise to the different responses **for the fired and paleo-fired samples** as shown in results in section 4.1. **We**
101 **define** the sensitivity change rate as the ratio of sensitivity at a temperature and the sensitivity of the temperature
102 100 °C lower of the same sample. For example, the sensitivity change rate at 600 °C is the ratio of the 110 °C TL
103 sensitivity at 600 °C against the corresponding value at 500 °C.

104 The 110 °C TL sensitivity change rates against re-firing temperature for the fired samples are shown in Fig. 3a.
105 For samples that **have not** been heated to more than 500 °C, an obvious peak at about 500 °C **appears** in the curve of

106 110 °C TL sensitivity change rate versus re-firing temperature, but does not appear for samples that have been
107 heated to more than 500 °C.

108 The relationship between 110 °C TL sensitivity change rate and re-firing temperature for the paleo-fired
109 samples are shown in Fig. 3b. The curves have a peak at about 500 °C for the samples with a distance of more than
110 about 100 cm from the lava, but the peak does not appear for the samples of less than about 100 cm from the lava,
111 except for a sample of 10 cm away from the lava. We deduce that the samples of less than about 100 cm from the
112 lava had been heated to more than 500 °C, while the samples with a distance of more than about 100 cm from the
113 lava had not. The exception of the sample 10 cm from the lava will be discussed later.

114 4.3 150 °C TL peak sensitivity against re-firing temperature

115 The 150 °C TL peak sensitivity changes against re-firing temperature of fired samples are shown in Fig. 4a. A
116 large difference between the samples of the firing temperatures lower than 900 °C or above was observed. It has
117 indicated that curves have a distinct peak at about 900 °C for samples that have been heated to temperatures lower
118 than 900 °C. No peak was observed for samples that were heated to 900 °C or above. This difference offered a
119 means for distinguishing samples of the firing temperature below and above 900 °C.

120 The 150 °C TL peak sensitivity changes against re-firing temperature of the paleo-fired samples are plotted in
121 Fig. 4b. The curves of the three samples near the lava, with the distance of 10, 40, 55 cm from the lava, do not have
122 a peak at about 900 °C, which suggests that the samples have been heated to 900 °C or above. However, other
123 samples (except one 128 cm away from the lava) have a peak at about 900 °C, which implies that the samples have
124 not been heated to 900 °C.

125 4.4 OSL/150 °C TL sensitivity against re-firing temperature

126 The ratio of OSL sensitivity and 150 °C TL sensitivity (OSL/150 °C TL sensitivity) against re-firing

127 temperature of fired samples are displayed in Fig. 5. It shows that each curve of fired samples is separated from the
128 curve of unfired sample (YJZC-7) before the re-firing temperature reaches the temperature the sample has been
129 heated to. The two curves are overlapping with each other for the re-firing temperature above. The method is best
130 suitable for samples fired 500 °C and above.

131 OSL/150 °C TL sensitivity changes against re-firing temperature of the lava-baked samples are illustrated in
132 Fig. 6. Each curve of the paleo-fired samples is compared with the curve of the unfired sample. Some of them do
133 not have overlaps, however, there is an intersection, after which, the two lines have the similar trend of changing
134 with temperature. The temperature of the intersection is regarded as the paleo-firing temperature. Using this method,
135 the temperatures of the samples were acquired as shown in Table 1. We consider that there are differences between
136 natural and laboratory simulation. Complicated influencing factors were involved for natural samples compared to
137 the fired samples.

138 From the data illustrated in Table 1, the paleo-firing temperatures of the eight lava-baked samples decrease
139 exponentially with the distance from the lava (Fig. 7). The results show that the sample of 10 centimeters away
140 from the lava might have been fired to more than 1000 °C. The sample of 158 centimeters beneath the lava might
141 have been heated to about 300 °C. Based on the estimated temperature profile, the temperature of the lava above
142 was about 1100 °C (Fig. 7). We would like to point out that all the re-firing treatments were carried out in the
143 muffle oven for 15min in the laboratory, which was shorter than the baked time by lava in nature. This may lead to
144 overestimation of the firing temperature (Han et al., 2000).

145 5. Discussions

146 Our observations showed that the luminescence sensitization had a close relationship with the firing
147 temperature. The findings from the re-firing experiments indicated that TL and OSL sensitivity changes against the

148 re-firing temperature can be used to estimate the paleo-firing temperatures of samples. The results obtained by
149 using the three methods, **which utilizing** 110 °C TL sensitivity change rate, 150 °C TL sensitivity and the ratio of
150 OSL/150 °C TL sensitivity, **can be well compared** with each other. The 110 °C TL sensitivity is able to distinguish
151 whether the sample has been fired to temperatures above 500 °C or not **for the fired samples. However, the 110 °C**
152 **TL signal is pre-dose dependent (Li, 2002), therefore, the sensitivity change of each re-firing temperature may give**
153 **a spurious result in the temperature estimation for the paleo-fired samples.** The pre-dose effects can be minimized
154 **by** using the 110 °C TL sensitivity change rate, because the same pre-dose was applied for all re-firing temperatures
155 of a sample. The 150 °C TL sensitivity can be used to **identify** samples that have been heated to ≥ 900 °C.

156 **By** using the ratio of OSL sensitivity to 150 °C TL sensitivity, we can estimate the paleo-firing temperature **to**
157 which the sample has been heated. It can be obtained by comparing the OSL/150 °C TL sensitivity versus re-firing
158 temperature curve of the paleo-fired samples with that of the unfired sample. Both OSL and 150 °C TL sensitivity
159 have also been affected by pre-dose effect (Li and Chen, 2001). Both signals have different responses to thermal
160 treatment. **Ratio** matching provides a way of demonstrating the thermal effect. We interpret that the OSL/150 °C
161 sensitivity ratio has combined both effects of OSL sensitivity and 150 °C TL sensitivity to **re-firing** temperatures.
162 The combination can give a better resolution to the paleo-temperature of sample being heated. We **deduce** that the
163 ratio of OSL/150 °C TL sensitivity is dominated by the **maximum** temperature the sample had experienced. When
164 the re-firing temperature is lower than the fired temperature, the **maximum** temperatures of the unfired and fired
165 sample are different, so that the two lines are separating. When the re-firing temperature is higher than the fired
166 temperature, the max temperatures of the unfired and fired sample are the same, so that the two lines are
167 overlapping. Hence, the temperature can be constrained by comparing the curve of **fired/paleo-fired samples and**
168 **the unfired one.**

169 The peaks at curve of 110 °C TL and 150 °C TL sensitivity against re-firing temperature indicate that the phase
170 changes shift from α -quartz to β -quartz at 573 °C, and β -quartz to β -tridymite at 870 °C **occurred**. The 110 °C TL
171 sensitivity rate increases significantly to a peak around 500 °C for samples that have **heated at less than 500 °C**, i.e.
172 α -quartz. For samples heated to 500-900 °C, i.e. β -quartz, the 110 °C TL sensitivity change rates do not have a
173 peak at about 500 °C. For the sample that has been heated to above 900 °C, i.e. β -tridymite, the 110 °C TL
174 sensitivity change rates have a peak around 500 °C. This **may** explain the exception of the sample 10 cm away from
175 the lava, which has similar behavior as samples heated below 500 °C in the 110 °C TL sensitivity change rate
176 curves (Fig. 3b). The sample was heated to the highest temperature among the lava-baked samples, because it is the
177 closest to the lava. The result from the OSL/150 °C TL sensitivity against re-firing temperatures **suggests** that this
178 sample was heated to temperature above 1000 °C. Similarly, the phase changes of quartz have been demonstrated in
179 the 150 °C TL sensitivity. It is noted that quartz turns back to α -quartz when it cooled to temperature of 573 °C or
180 lower. However, the sensitivity changes of quartz luminescence signals are irreversible. It was explained as a result
181 of transferring holes from non-luminescence traps to luminescence traps (Zimmerman, 1971).

182 It has been demonstrated that the **paleo-firing temperature estimation using the 110 °C TL signal is precluded**
183 **by the pre-dose effect**. However, the pre-dose effect can be minimized when using the rate of the 110 °C TL
184 sensitivity change, because the same pre-dose is applied to the sample. Only the thermal effect will dominate the
185 110 °C TL sensitivity change rate. Similarly, the ratio of OSL sensitivity and 150 °C TL would have small impacts
186 of pre-dose, because the pre-dose effects affect both of the signals (Chen et al., 2000). Another advantage of using
187 the ratio is that no normalization is required for the aliquots.

188 6. Conclusions

189 Luminescence signals of 110 °C TL, 150 °C TL and OSL of quartz have close relationship with the

190 **paleo-firing** temperatures. The sensitivity of the signals can be affected by their thermal and ionizing radiation
191 histories. The pre-dose effects on the signals can be minimized when rate or ratios are used for thermal history
192 study. Significant changes in luminescence sensitivity happened at temperatures that are coincident with phase
193 change temperatures of the quartz.

194 Acknowledgements

195 We thank Prof. Yan-Chou Lu, Xu-Long Wang and Jin-Feng Liu for very helpful comments. We thank the
196 anonymous referee for the constructive comments. This work was supported by NSFC grant (40972208) and China
197 Geological Survey grant (1212011120147). This study was financially supported by the grants to Sheng-Hua Li
198 from the Research Grant Council of the Hong Kong Special Administrative Region, China (Project no. 7028/08P,
199 7033/12P and 17303014).

200 References

- 201 Aitken, M. J., 1985. Thermoluminescence Dating. Academic Press, London.
- 202 Bøtter-Jensen, L., Agersnap Larsen, N., Mejdahl, V., Poolton, N. R. J., Morris, M. F., McKeever, S. W. S., 1995. Luminescence
203 sensitivity changes in quartz as a result of annealing. *Radiation Measurements*, 24(4), 535-541.
- 204 Chen, G., Li, S-H., Murray A.S., 2000. Study of 110°C TL peak sensitivity in optical dating of quartz. *Radiation Measurements*,
205 32(5-6), 641-645.
- 206 Göksu, H., Weiser, A., Regulla, D., 1989. 110 °C TL peak records the ancient heat treatment of flint. *Ancient TL*, 7(1), 15-17.
- 207 Griffiths, R. W., 2000. The dynamics of lava flows. *Annual Review of Fluid Mechanics*, 32(1), 477-518.
- 208 Han, Z. Y., Li, S-H., Tso, M. Y. W., 2000. Effects of annealing on TL sensitivity of granitic quartz. *Radiation measurements*, 32(3),
209 227-231.
- 210 Li, H. H., Sun, J.Z., 1984. Research for Datong volcanic activity using TL age. *Science in China*, (7), 637-644 (in Chinese).

211 Li, S-H., 2002. Luminescence sensitivity changes of quartz by bleaching, annealing and UV exposure. *Radiation effects and defects in*
212 *solids*, 157(3), 357-364.

213 Li, S-H., Chen, G., 2001. Studies of thermal stability of trapped charges associated with OSL from quartz. *Journal of Physics D:*
214 *Applied Physics*, 34(4), 493-498.

215 Li, S-H., Sun, J. M., Zhao, H., 2002. Optical dating of dune sands in the northeastern deserts of China. *Palaeogeography,*
216 *Palaeoclimatology, Palaeoecology*, 181(4), 419-429.

217 McKeever, S. W., Chen, C. Y., Halliburton, L. E., 1985. Point defects and the pre-dose effect in natural quartz. *Nuclear Tracks* 10(4),
218 489-495.

219 Oniya, E. O., Polymeris, G. S., Tsirliganis, N. C., Kitis, G., 2012. Behavior of various Nigerian quartz samples to repeated irradiation
220 and heating. *Geochronometria*, 39(3), 212-220.

221 Pinkerton, H., James, M., Jones, A., 2002. Surface temperature measurements of active lava flows on Kilauea volcano, Hawai' i.
222 *Journal of volcanology and geothermal research*, 113(1), 159-176.

223 Polymeris, G., Kitis, G., Pagonis, V., 2006. The effects of annealing and irradiation on the sensitivity and superlinearity properties of
224 the thermoluminescence peak of quartz. *Radiation Measurements*, 41(5), 554-564.

225 Polymeris, G., Sakalis, A., Papadopoulou, D., Dallas, G., Kitis, G., Tsirliganis, N., 2007. Firing temperature of pottery using TL and
226 OSL techniques. *Nuclear Instruments and Methods in Physics Research Section A: Accelerators, Spectrometers, Detectors and*
227 *Associated Equipment* 580 (1), 747-750.

228 Poolton, N. R. J., Smith, G. M., Riedi, P. C., Bulur, E., Bøtter-Jensen, L., Murray, A. S., Adrian, M., 2000. Luminescence sensitivity
229 changes in natural quartz induced by high temperature annealing: a high frequency EPR and OSL study. *Journal of Physics D:*
230 *Applied Physics*, 33(8), 1007.

231 Rendell, H. M., Townsend, P. D., Wood, R. A., Luff, B. J., 1994. Thermal treatments and emission spectra of TL from quartz.

232 Radiation Measurements,23(2), 441-449.

233 Schilles, T., Poolton, N. R. J., Bulur, E., Bøtter-Jensen, L., Murray, A. S., Smith, G., Riedi, P. C., Wagner, G. A., 2001. A
234 multi-spectroscopic study of luminescence sensitivity changes in natural quartz induced by high-temperature annealing. Journal
235 of Physics D: Applied Physics, 34(5), 722.

236 Sunta, C., David, M., 1982. Firing temperature of pottery from pre-dose sensitization of TL, PACT 6, 460-467.

237 Watson, I., Aitken, M., 1985. Firing temperature analysis using the 110 °C TL peak of quartz. Nuclear Tracks and Radiation
238 Measurements 10 (4-6), 517-520.

239 Wintle, A. G, Murray, A. S., 1999. Luminescence sensitivity changes in quartz. Radiation Measurements, 30(1), 107-118.

240 Yang, X. H., McKeever, S. W. S., 1990. The pre-dose effect in crystalline quartz. Journal of Physics D: Applied Physics, 23(2), 237.

241 Zimmerman, J., 1971. The radiation-induced increase of the 100 °C thermoluminescence sensitivity of fired quartz. Journal of Physics
242 C: Solid State Physics, 4(18), 3265-3276.

243

244

245

246

247

248

249

250

251

252 Table 1

253 Paleo-firing temperatures of the paleo-fired samples from the baked layer at Datong

Distance from the lava(cm)	Results from 110 °C TL peak	Results from 150 °C TL peak	Results from OSL/150 °C TL peak
10	<500 °C	≥900 °C	>1000 °C
40	≥500 °C	≥900 °C	800 °C
55	≥500 °C	≥900 °C	700 °C
81	≥500 °C	<900 °C	550 °C
97	≥500 °C	<900 °C	500 °C
107	<500 °C	<900 °C	400 °C
128	<500 °C	>900 °C	400 °C
158	<500 °C	<900 °C	300 °C

254

255

256

257

258

259

260

261

262

263

264

265

266 **Figure captions:**

267 Fig. 1. Location of samples in the profile and distances of samples from the lava above.

268 Fig. 2. 110 °C TL sensitivity changes with re-firing temperatures for (a) the fired samples with different temperatures and (b) the
269 paleo-fired samples from baked layer.

270 Fig. 3. 110 °C TL sensitivity change rate with re-firing temperatures for (a) the fired samples with different temperatures and (b) the
271 paleo-fired samples from the baked layer.

272 Fig. 4. 150 °C TL sensitivity changes with re-firing temperatures for (a) the fired samples with different temperatures and (b) the
273 paleo-fired samples from the baked layer.

274 Fig. 5. OSL /150 °C TL sensitivity changes with re-firing temperatures of the fired samples with different temperatures.

275 Fig. 6. OSL/150 °C TL sensitivity changes with re-firing temperatures of the paleo-fired samples from the baked layer.

276 Fig. 7. Relation between paleo-firing temperatures and distance from lava.

277

278

279

280

281

282

283

284

285

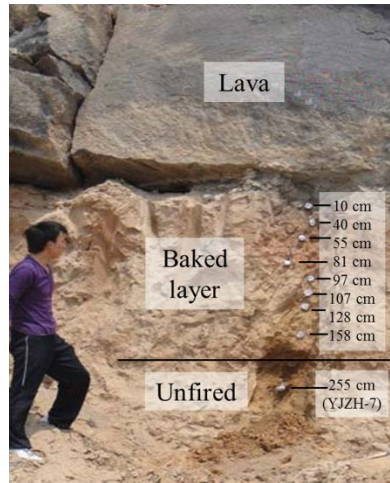
286

287

288

289

290



291

292

293 Fig. 1.

294

295

296

297

298

299

300

301

302

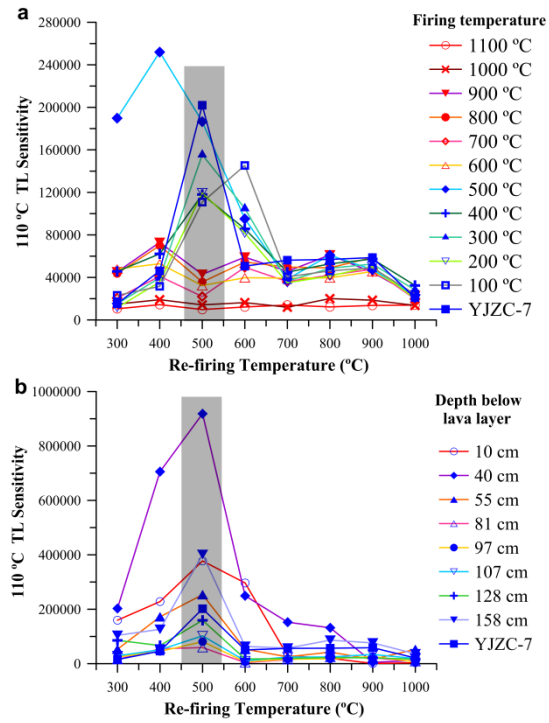
303

304

305

306

307



308

309

310 Fig. 2.

311

312

313

314

315

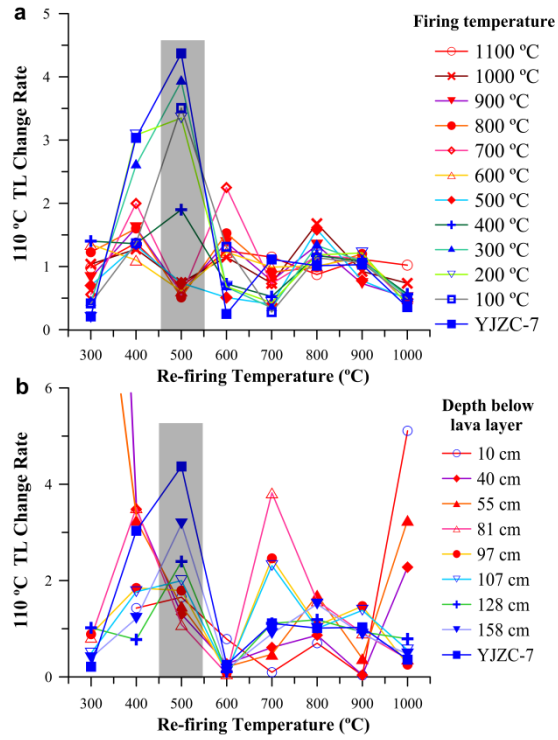
316

317

318

319

320



321

322

323 Fig. 3.

324

325

326

327

328

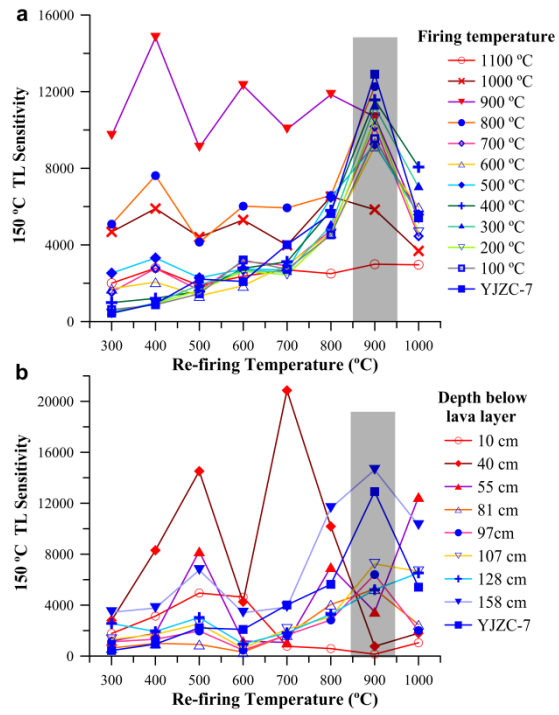
329

330

331

332

333



334

335

336 Fig. 4.

337

338

339

340

341

342

343

344

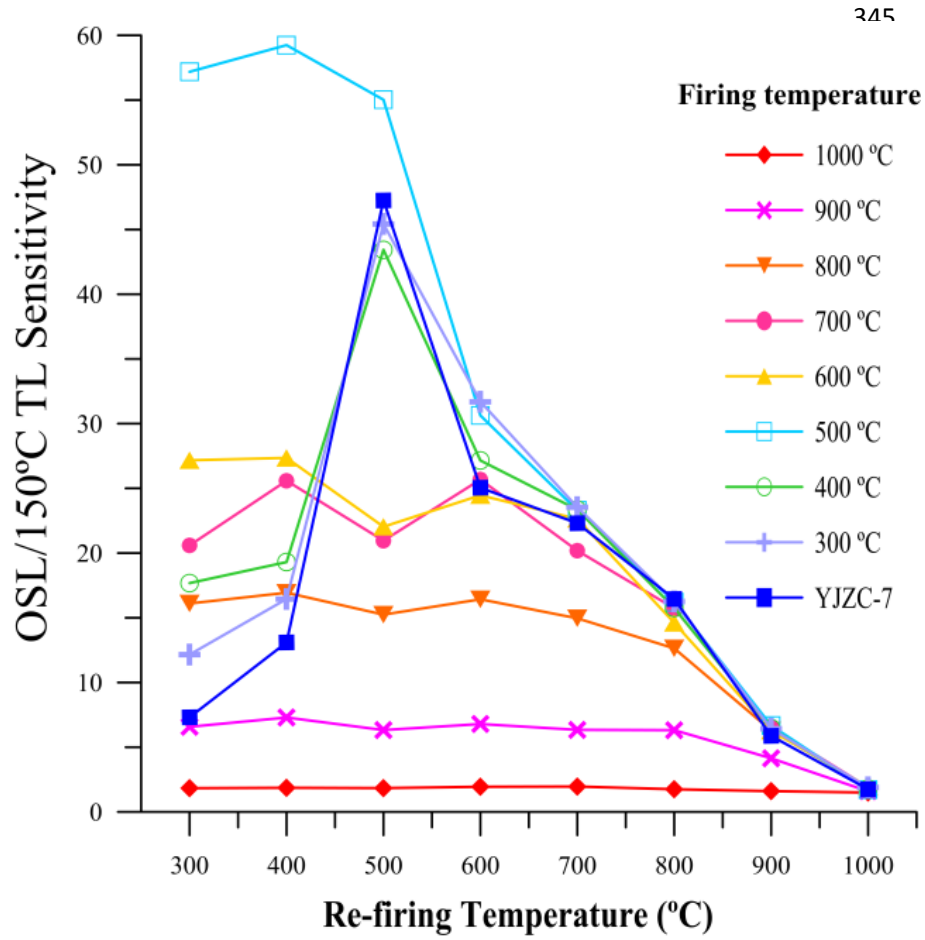


Fig. 5.

357

358

359

360

361

362

363

364

365

366

367

368

369

370

371

372

373

374

375

376

377

378

379

380

381

382

383

384

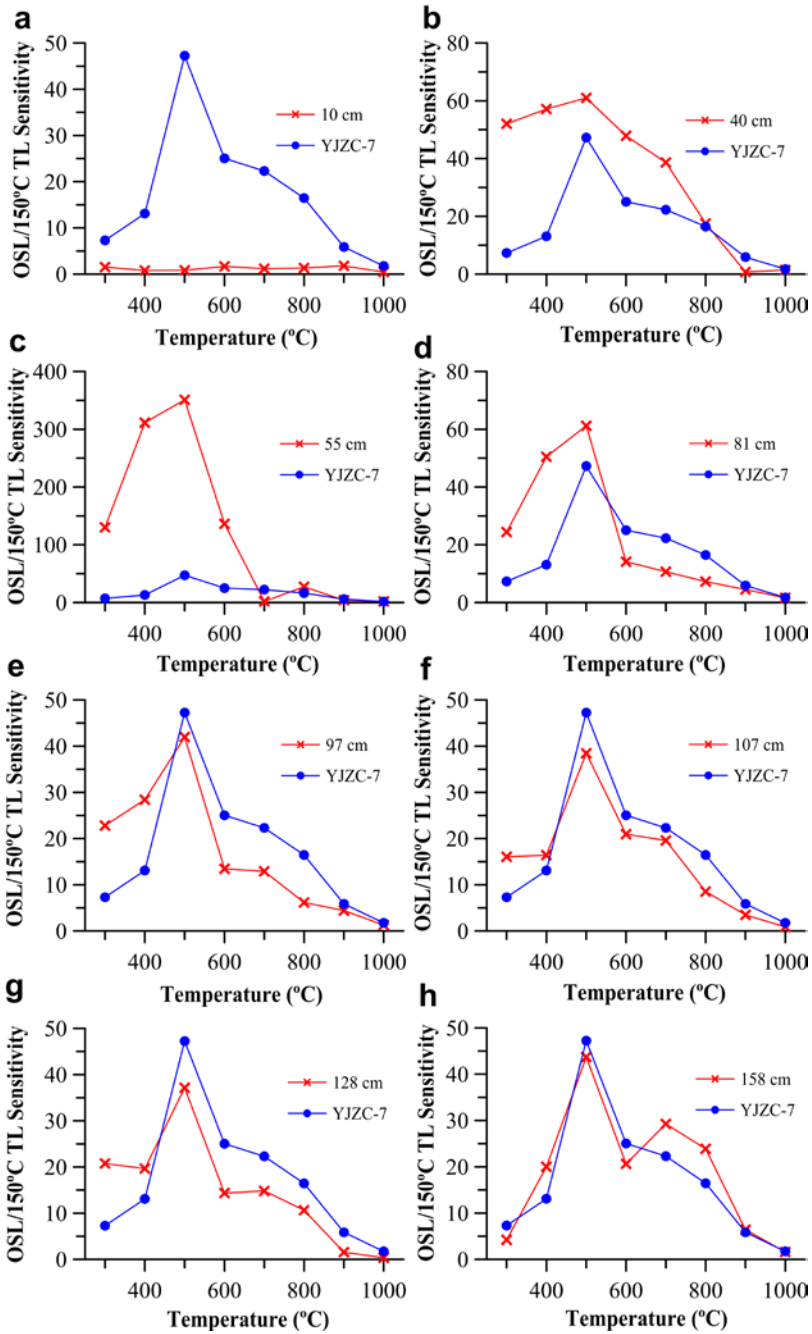
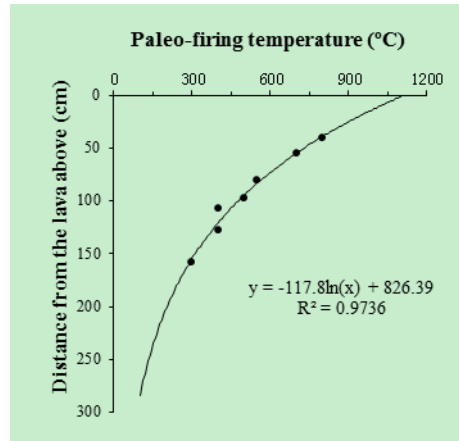


Fig. 6.

385



386

387

388 Fig. 7.

389