



**This electronic thesis or dissertation has been  
downloaded from Explore Bristol Research,  
<http://research-information.bristol.ac.uk>**

*Author:*

**Nash, Maisie V**

*Title:*

**Metagenomic insights into microbial communities in proglacial landscapes**

**General rights**

Access to the thesis is subject to the Creative Commons Attribution - NonCommercial-No Derivatives 4.0 International Public License. A copy of this may be found at <https://creativecommons.org/licenses/by-nc-nd/4.0/legalcode>. This license sets out your rights and the restrictions that apply to your access to the thesis so it is important you read this before proceeding.

**Take down policy**

Some pages of this thesis may have been removed for copyright restrictions prior to having it been deposited in Explore Bristol Research. However, if you have discovered material within the thesis that you consider to be unlawful e.g. breaches of copyright (either yours or that of a third party) or any other law, including but not limited to those relating to patent, trademark, confidentiality, data protection, obscenity, defamation, libel, then please contact [collections-metadata@bristol.ac.uk](mailto:collections-metadata@bristol.ac.uk) and include the following information in your message:

- Your contact details
- Bibliographic details for the item, including a URL
- An outline nature of the complaint

Your claim will be investigated and, where appropriate, the item in question will be removed from public view as soon as possible.

---

# Metagenomic insights into microbial communities in proglacial landscapes

---

Maisie Victoria Nash

A thesis submitted to the University of Bristol in accordance with the requirements for award of the degree of Doctor of Philosophy in the Faculty of Science

School of Geographical Sciences

April 2019



Word count: 46,231

---

# Abstract

Environmental DNA analysis using metagenomics can provide an insight into the taxonomy and functional potential of microbial communities *ex situ*, without the need for culturing or DNA amplification. However, metagenomics has had limited application to environmental microbial ecology, in particular, to microbial communities in proglacial regions. This thesis aims to contribute to the body of literature on environmental metagenomics through evaluating assemblers for soil microbial ecologists, and subsequently applying metagenomics to investigate microbial communities in proglacial environments.

Assembly of metagenome sequencing reads can improve sequence alignment to taxonomic and functional databases, thereby improving ecological conclusions. However, limited guidance is available for assembler choice by microbial ecologists. The first study in this thesis compares assemblers for soil metagenome data, demonstrating the importance of assembler evaluation and parameterization. The guidance produced was applied to investigate microbial communities in proglacial regions, including fjords and forefields. Proglacial forefields present a unique opportunity to understand microbial colonization in land exposed by glacier retreat. Here, metagenomics was used to investigate microbial diversity and functional potential during forefield succession, alongside comparing the diversity of nitrogen-fixing bacteria between Arctic forefields. This work contributes to our understanding of Arctic microbial ecology, which has significance given the continued exposure of forefield soils during global warming. In addition, metagenomics was used to investigate microbial communities in oligotrophic, dark, saline fjord waters, fed by glacial meltwater. This work highlights the potential of metagenomics to understand uncultured microbial samples and demonstrate areas for further analysis, such as targeting novel genomes.

This thesis has contributed to the literature on metagenomics by providing methodological guidance for microbial ecologists, alongside enhancing understanding of microbial diversity in proglacial regions. It is hoped that this work will inspire others to use metagenomics to explore uncultured microbial samples and to target further analysis or exploration for unique genomes.

## Acknowledgements

I would like to thank the following people who have provided guidance and support over the course of my PhD. Firstly, I would like to thank my advisors, Patricia Sánchez-Baracaldo who has been a fantastic collaborator and a huge source of support, Alex Anesio for providing me with the opportunity to enter the world of microbiology, and Gary Barker for pushing me to become a better bioinformatician. Thank you to NERC for funding my project through a GW4<sup>+</sup> DTP scholarship and the PISCES project for supporting my fieldwork. Many thanks to my cohort of PhD students and the Browns community, in particular Nathan Christmas for being my technical help desk, and Steve Chuter, Ale Urrea and Claire Donnelly for supporting me through the ups and downs. I would like to thank my fieldwork team, including Jon Hawkings, Rebecca Huggett, Rory Burford, Alex Beaton, Helena Pryer, Hong Chin and all those on the PISCES project. Thank you to my team and coaches at CrossFit 605 who have kept my competitive spirit alive and motivated me to improve in all aspects of my life. I would like to thank my best friend, Alice Haworth and my sisters, Amy and Libby, for supporting me through my crazy pursuit in science. Finally, I would like to thank my parents, Trevor and Julie, without their support I would never have made it this far, and their work ethic and positive attitude inspires me every day.

## Authors declaration

I declare that the work was carried out in accordance with the requirements of the University's regulations and Code of Practice for Research Degree programmes and that it has not been submitted for any other academic award. Except where indicated by specific reference in the text, the work is the candidate's own work. Work done in collaboration with, or the assistance of others, is indicated as such. Any views expressed in the dissertation are those of the author.

Signed: ..... Date:.....

# Table of Contents

Abstract.....	2
Acknowledgements.....	3
Authors declaration.....	4
List of Figures.....	9
List of Tables.....	10
Key terms.....	11
Abbreviations.....	12
<b>Chapter 1: Literature Review.....</b>	<b>13</b>
1.1 Introduction to glacial systems.....	13
1.2 Microbial communities in glacial systems.....	14
1.3 Proglacial forefields and microbial communities.....	14
1.3.1 Using forefields to investigate microbial succession.....	15
1.3.2 Plant succession and soil development.....	16
1.3.3 Microbial succession in glacial forefields.....	17
1.3.4 Microbial nitrogen cycling in glacial forefields.....	19
1.3.5 Carbon, phosphorous and sulfur in forefields.....	20
1.4 Proglacial fjords and microbial communities.....	22
1.4.1 Fjord systems.....	22
1.4.2 The influence of glacial meltwater.....	22
1.4.3 Microbial communities and fjord systems.....	23
1.5 Proglacial environments and climate change.....	24
1.6 Methodological considerations for investigating microbial communities...25	
1.6.1 Methods available.....	25
1.6.2 Culture based studies.....	25
1.6.3 First generation DNA sequencing.....	26
1.6.4 Taxonomic marker genes.....	27
1.6.5 Next generation DNA sequencing.....	27
1.6.6 Metagenomics.....	28
Metagenome sequence assembly and annotation.....	29
Sequence assembly.....	29
Sequence annotation.....	31

The choice of assembler.....	32
Selecting an appropriate assembler.....	33
1.7 Summary and research gaps.....	34
1.8 Aims and objectives.....	35
<b>Chapter 2: Comparison of publicly available metagenome assemblers for soil bacterial communities.....</b>	<b>38</b>
2.1 Introduction.....	38
2.2 Methodology.....	40
2.3 Results.....	46
2.3.1 Assembly contiguity and completeness.....	46
2.3.2 Modifying community complexity.....	51
2.3.3 Assembly parameterization.....	52
2.3.4 Assembly chimeras.....	56
2.3.5 Taxonomic distribution of assemblies.....	57
2.4 Discussion.....	59
2.4.1 The premise of the study.....	59
2.4.2 Assembly evaluation metrics.....	59
2.4.3 Simulated bacterial communities.....	60
2.4.4 The choice of assembler.....	61
2.4.5 The importance of parameterisation.....	62
2.4.6 Best performing assemblers.....	62
2.5 Conclusion.....	63
2.6 Observations.....	64
2.7 Recommendations for best practice.....	65
2.8 Limitations and future work.....	66
<b>Chapter 3: Metagenomic insights into diazotrophic communities across Arctic glacier forefields.....</b>	<b>68</b>
3.1 Introduction.....	69
3.2 Materials and methods.....	71
3.2.1 Field sampling.....	71
3.2.2 Soil organic carbon and total nitrogen content.....	72
3.2.3 DNA extraction, library preparation and sequencing.....	72
3.2.4 Metagenome assembly and annotation.....	73
3.2.5 Nif taxonomy.....	74

3.2.6 Gene phylogeny.....	74
3.3 Results and Discussion.....	76
3.3.1 Soil carbon and nitrogen.....	76
3.3.2 Rarefaction analysis.....	78
3.3.3 Nif genes recovered .....	79
3.3.4 Nitrogenase clusters .....	81
3.3.5 Diazotroph community structure.....	84
3.4 Conclusions .....	89
3.5 Limitations and future work.....	90
<b>Chapter 4: Microbial community development over a gradient of soil succession along the Midtre Lovénbreen glacier forefield, Svalbard.....</b>	<b>92</b>
4.1 Introduction.....	92
4.2 Methodology.....	97
4.2.1 Field sampling.....	97
4.2.2 DNA extraction, library preparation and sequencing.....	98
4.2.3 Rarefaction curves.....	99
4.2.4 Metagenome assembly and mapping.....	100
4.2.5 Taxonomic annotation.....	100
4.2.6 Metagenome binning.....	100
4.2.7 Soil metadata.....	101
4.3 Results .....	102
4.3.1 Rarefaction curves.....	102
4.3.2 Soil metadata.....	102
4.3.3 Microbial community composition.....	106
4.3.4 Genome bins.....	109
4.4 Discussion.....	114
4.4.1 Total nitrogen and total organic carbon.....	114
4.4.2 Microbial community composition.....	115
4.4.3 Microbial metabolisms – Carbon.....	117
4.4.4 Microbial metabolisms – Nitrogen.....	118
4.4.5 Microbial metabolisms – Sulfur.....	120
4.4.6 Unique genomes.....	121
4.5 Conclusions.....	121
4.6 Limitations and future work.....	124



<b>Chapter 5: Sequencing the deep – microbial communities recovered from benthic fjord sediments, Chilean Patagonia.....</b>	<b>125</b>
5.1 Introduction.....	125
5.2 Methods.....	128
5.2.1 Field sampling.....	128
5.2.2 Metagenome DNA extraction and sequencing.....	129
5.2.3 Metagenome read annotation.....	130
5.2.4 Metagenome assembly, genome binning and annotation.....	130
5.2.5 16s rRNA and dsrAB phylogenies.....	130
5.2.6 Sediment organic carbon and total nitrogen content.....	131
5.3 Results and discussion.....	132
5.3.1 Fjord metadata and influence of glacial meltwater.....	132
5.3.2 Sediment TN and TOC.....	135
5.3.3 Metagenome assembly and annotation.....	136
5.3.4 Genome binning.....	142
5.3.5 Phylogenetic analysis – 16s rRNA.....	148
5.3.6 Phylogenetic analysis – dissimilatory sulfite reductase.....	150
5.3.7 Implications for biogeochemical cycling.....	152
5.4 Conclusions.....	153
5.5 Limitations and future work.....	154
<b>Chapter 6: Concluding discussion.....</b>	<b>156</b>
6.1 Summary.....	156
6.2 Limitations and future opportunities.....	160
6.3 Implications.....	161
References.....	163
Appendix 1.....	188
Appendix 2 .....	194
Appendix 3.....	204
Appendix 4.....	220

## List of Figures

1.1	Schematic of forefield environments.....	15
2.1	Methodology schematic.....	46
2.2	Maximum contig length and percentage coverage.....	48
2.3	Number of contigs and percentage coverage for assemblies.....	49
2.4	Percentage coverage for all assemblies.....	54
2.5	Maximum contig length for all assemblies.....	55
2.6	Chimeric contigs.....	56
2.7	Assembly taxonomic heat maps.....	58
3.1	Normalized nif gene abundance, soil TN and TOC.....	80
3.2	nifH maximum likelihood phylogeny.....	82
3.3	Taxonomic distribution of forefield nif genes.....	85
4.1	Soil nitrogen cycle.....	95
4.2	Rarefaction curves for Midtre Lovénbreen metagenomes.....	102
4.3	Soil TN and TOC content.....	105
4.4	Read based microbial community composition for soil samples.....	108
4.5	SEED function profile for nitrogen metabolism in genome bins.....	113
4.6	Forefield succession schematic.....	122
5.1	Sampling sites in the Steffen and Baker fjords.....	129
5.2	CTD depth profiles for sediment sampling sites.....	134
5.3	Underway sampling data (salinity, temperature, pH and turbidity).....	135
5.4	TN and TOC content of Steffen fjord sediments.....	136
5.5	Kaiju taxonomic classification of reads, at the class level.....	139
5.6	16s rRNA phylogeny of <i>Nitrosopumilus</i> sample sequences.....	149
5.7	Phylogeny for dissimilatory sulfite reductase (dsrAB).....	151

## List of Tables

2.1	Metagenome assemblers tested.....	41
2.2	Genomes used in artificial metagenomes.....	42
2.3	Artificial metagenome description.....	44
2.4	Metagenome assembly summary.....	47
2.5	Average metagenome assembly quality.....	50
3.1	Accession numbers of reference nifH sequences.....	75
3.2	TN and TOC content in forefield soils.....	76
3.3	TN and TOC ANOVA between forefields.....	77
3.4	Post Hoc Tukey between forefields.....	77
3.5	Forefield soil TN and TOC content.....	78
3.6	NCBI BLASTn matches for sample nif sequences.....	84
4.1	Sampling site locations and soil age.....	98
4.2	Summary statistics for metagenomes.....	99
4.3	TOC, TN, Temperature and pH for soil samples.....	104
4.4	One-way ANOVA comparing TN and TOC between all soils.....	105
4.5	Read-based microbial community composition.....	107
4.6	Percentage classified and unclassified reads at the genus level.....	107
4.7	Genome binning summary for assembled metagenomes.....	110
5.1	Benthic sediment sampling site locations.....	129
5.2	Metadata for sediment sampling sites.....	132
5.3	Metagenome assembly statistics.....	137
5.4	Percentage unassigned sequences.....	137
5.5	Read based taxonomic composition.....	141
5.6	Sample 1 assembled bins and BLAST matches.....	144
5.7	Sample 2 assembled bins and BLAST matches.....	145
5.8.	Sample 4 assembled bins and BLAST matches.....	146
5.9	Sample 5 assembled bins and BLAST matches.....	147
5.10	Sample 7 assembled bins and BLAST matches.....	148
5.11	BLAST matches for dsrAB gene sequences.....	152

# Key terms

<b>Allochthonous</b>	Originating outside the location it is identified in
<b>Anoxic</b>	Without oxygen
<b>Assembly bubbles</b>	Two de Bruijn assembly paths that begin and end at the point on the graph
<b>Assembly completeness</b>	How much of the input DNA (reads) are incorporated in the output assembly
<b>Assembly contiguity</b>	Low long DNA contigs are in an assembly
<b>Assembly coverage</b>	The percentage of organisms, genes or sequences captured by sequencing or assembly
<b>Autochthonous</b>	Formed in situ
<b>Autotroph</b>	An organism which can produce complex organic compounds from inorganic material
<b>Bioinformatics</b>	The computational analysis of complex genetic data e.g. from genomes or metagenomes
<b>Chemo-lithotroph</b>	An organism which can use inorganic substances for growth
<b>Chemo-organotroph</b>	An organism which obtains its energy from organic molecules
<b>Chronosequence</b>	A set of soil sites which share similar attributes but differ in soil development due to age
<b>Contig</b>	A consensus set of DNA reads, merged into a continuous sequence using overlaps
<b>Diazotroph</b>	A microorganism which can fix dinitrogen ( $N_2$ ) to ammonia ( $NH_3$ )
<b>Facultative anaerobes</b>	An organism which can respire anaerobically or aerobically
<b>Forefield</b>	Land immediately in front of a glacier terminus, exposed by ice retreat
<b>Heterotroph</b>	An organism which uses fixed organic substances for growth
<b>Homopolymer regions</b>	Repeat regions in genomic DNA
<b>Hypoxic</b>	Deprived of oxygen
<b>K-mer</b>	A short DNA sequence that reads are fragmented in to prior to assembly
<b>Metagenome</b>	Environmental genetic material (DNA) sourced directly from a sample, without amplification
<b>Methanogen</b>	Microbes which produce methane ( $CH_4$ ) during anaerobic respiration
<b>N50</b>	The length (N) over which 50% of contigs are above this length
<b>Oligotrophic</b>	Very low nutrient
<b>Phototroph</b>	An organism which utilizes energy from the sun to create organic molecules
<b>Phylogeny</b>	The evolutionary history of a group of organisms and their relationship to one another
<b>Primary succession</b>	Gradual development of land through microbial and plant colonization
<b>Proglacial</b>	The area in front of a glacier terminus (ocean, lake, soil)
<b>Proteome (proteomics)</b>	The entire set of proteins expressed by a genome or metagenome
<b>Read coverage</b>	The frequency a sequenced DNA read appears in a metagenome
<b>Scaffold</b>	A non-contiguous set of contigs and gaps, formed to create a draft genome
<b>Taxonomy</b>	The classification of organisms
<b>Transcriptome</b>	All the mRNA expressed by a genome or metagenome

# Abbreviations

<b>16s rRNA</b>	16s ribosomal ribonucleic acid
<b>ANOVA</b>	Analysis of variance
<b>AOA</b>	Ammonia oxidizing archaea
<b>AS</b>	Alignment score
<b>ATP</b>	Adenosine triphosphate
<b>BLAST (n&amp;p)</b>	Basic local alignment search tool (nucleotide & protein)
<b>BP</b>	Base pairs
<b>CTD sensor</b>	Conductivity, temperature and depth sensor
<b>ddNTP's</b>	Dideoxynucleotides triphosphates
<b>DNA</b>	Deoxyribonucleic acid
<b>dsrAB gene</b>	Dissimilatory sulfite reductase
<b>EPS</b>	Exopolymeric substances
<b>INDELS</b>	Insertions and deletions
<b>LCA</b>	Last common ancestor
<b>ML</b>	Midtre Lovénbreen glacier
<b>NCBI</b>	National Centre for Biotechnology Information
<b>NGS</b>	Next generation sequencing
<b>Nif genes</b>	Nitrogen fixation genes
<b>NPP</b>	Net primary productivity
<b>OLC</b>	Overlap layout consensus
<b>OTU</b>	Operational taxonomic unit
<b>PCR</b>	Polymerase chain reaction
<b>PP</b>	Primary productivity
<b>RB</b>	Rabots glacier
<b>RL</b>	Russell glacier
<b>SD</b>	Standard deviation
<b>ST</b>	Storglaciären glacier
<b>TN</b>	Total nitrogen
<b>TOC</b>	Total organic carbon
<b>UNSECO</b>	United Nations Educational, Scientific and Cultural Organization

# Chapter 1: Literature Review

## 1.1 Introduction to glacial systems

At present, it is estimated that ice sheets and glaciers account for ~10% of global land coverage (16 million km<sup>2</sup>) (Knight, 1999). These systems comprise of subglacial (beneath glacier), englacial (within glacier), supraglacial (glacier surface) and proglacial (in front of glacier) environments (Benn and Evans, 2014). Ice sheets and glaciers have been the focus of scientific research over recent years due to ongoing ice melt and glacier retreat with warming global temperatures (IPCC, 2013). The importance of these regions with regard to meltwater fluxes, sea level rise, sediment and nutrient transport and global surface albedo have been documented (Tranter *et al.*, 2002; Hood *et al.*, 2009; Wadham *et al.*, 2010; Wadham *et al.*, 2013; Dutton *et al.*, 2015). The importance of glaciers and ice sheets in global biogeochemical cycles has also been noted, through supplying limiting nutrients to downstream fjord and ocean ecosystems, such as iron, carbon and nitrogen, and thereby stimulating productivity in near shore regions (Tranter *et al.*, 1994; Stratham *et al.*, 2008; Lawson *et al.*, 2014; Hawkings *et al.*, 2015).

However, the study of microbial communities, in both glaciers and ice sheets has only come to the forefront of research in recent years (Hodson, 2006; Hodson *et al.*, 2008; Anesio *et al.*, 2009). The potential implications of these communities on global biogeochemical cycles, such as carbon, nitrogen and methane, are substantial (Anesio *et al.*, 2009; Boyd *et al.*, 2010; Stibal *et al.*, 2012; Telling *et al.*, 2012; Wadham *et al.*, 2012). Most research on microbial communities in glacial habitats has focused on supraglacial environments, such as cryoconite holes and subglacial sediments (Anesio *et al.*, 2009; Boyd *et al.*, 2010; Telling *et al.*, 2010). However, limited research has investigated microbial community diversity and function in proglacial regions, such as forefields and fjords (Duc *et al.*, 2009; Zumsteg *et al.*, 2013). Thus far, these regions have been shown to harbour diverse, active communities, and may therefore have vital roles in local biogeochemical cycles (Duc *et al.*, 2009). Additionally, these microbial communities may be influenced by warming temperatures and increased glacial melt, in line with climate change (Alison and Tresder, 2008; Rinnan *et al.*, 2009).

## 1.2 Microbial communities in glacial systems

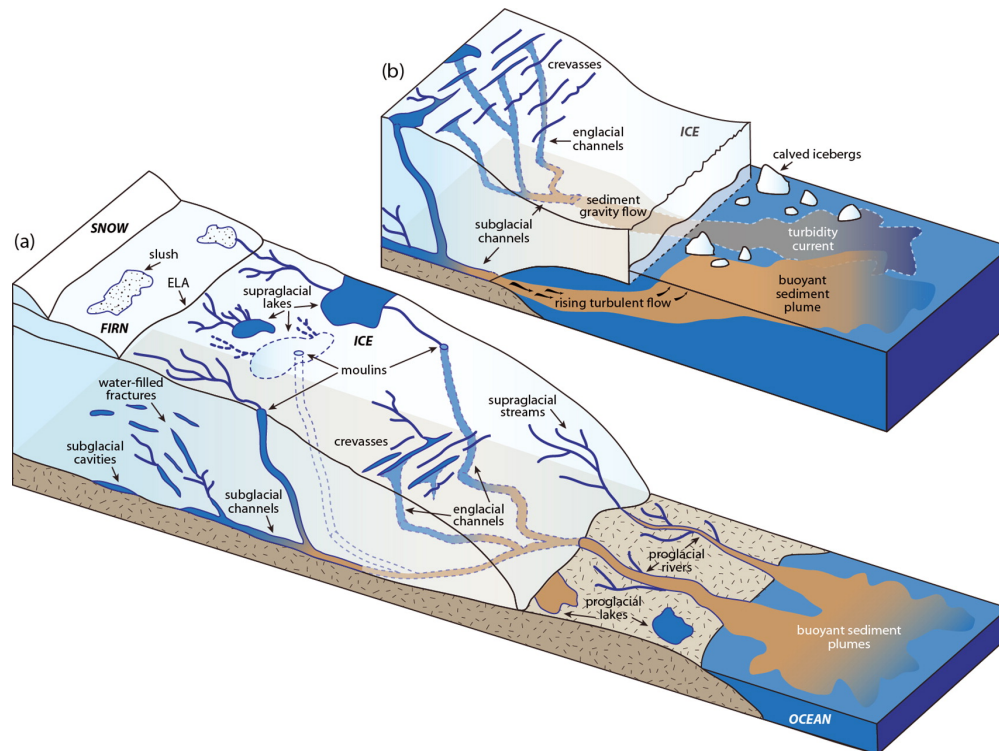
In early work, glaciers were considered purely abiotic systems, due to the low temperatures, high UV exposure and the oligotrophic nature of the glacial environments (Collins, 1979; Raiswell, 1984). However, recent developments have identified active and diverse microbial communities in supraglacial, subglacial and proglacial habitats (Skidmore *et al.*, 2000; Sigler and Zeyer, 2002; Bhatia *et al.*, 2006; Stibal *et al.*, 2006; Malard and Pearce, 2018). A wide range of organisms have been identified, such as diazotrophs, methanogens, heterotrophs and chemolithoautotrophs (Skidmore *et al.*, 2000; Christner *et al.*, 2008; Hodson *et al.*, 2008; Stibal *et al.*, 2012). The adaptation of these organisms to survive in such extreme conditions, and their importance in both local and global biogeochemical cycles has sparked a wave of research interest (Mindl *et al.*, 2007; Wadham *et al.*, 2013). These organisms can often be found living across environmental gradients, for example at stages of increasing soil development in glacier forefields (Bradley *et al.*, 2014) or along salinity gradients in fjords draining glacial meltwater (Gutiérrez *et al.*, 2015). Consequently, it is interesting to understand the diversity of these communities, how they contribute to biogeochemical cycles, and how this varies over environmental gradients. Investigating microbial communities in proglacial regions is particularly important, as these may be modified with climate change, for example, increased glacial meltwater fluxes may modify salinity balances in fjords (Davila *et al.*, 2002). Here, focus is placed on proglacial microbial communities in forefield soils and fjord sediments (Figure 1.1).

## 1.3 Proglacial forefields and microbial communities

The proglacial forefield of land terminating glaciers facilitates the growth of a diverse range of microorganisms, due to the presence of proglacial rivers, lakes, soils and vegetation, all of which harbour distinct, active communities (Liu *et al.*, 2006; Duc *et al.*, 2009; Reddy *et al.*, 2009; Zumsteg *et al.*, 2013; Figure 1.1). Proglacial rivers are important for the export of labile organic matter and nutrients to downstream ecosystems (Hood & Scott, 2008; Hood *et al.*, 2009; Hawkings *et al.*, 2014; Hopwood *et al.*, 2014).

Proglacial soils have been the focus of an abundance of plant-based research, in particular, looking at the succession of plant species on newly exposed soil, following glacier retreat (Frenot *et al.*, 1998; Hodkinson *et al.*, 2003; Breen and Levesque, 2006). However, more recently, these soils have been the focus of microbiologists, investigating how microbial communities are initially established, and their subsequent role in soil, nutrient and plant development (Schutte *et al.*, 2009; Göransson *et al.*, 2011; Wojcik *et al.*, 2019). These

microbial communities are present across an environmental gradient of soil succession, and therefore may be diverse in taxonomy and function across the proglacial region (Bradley *et al.*, 2014). The following section will discuss the research surrounding proglacial soil succession in depth, and in particular, its application within microbiology.



**Figure 1.1:** Schematic of glacier forefields (a) land terminating glacier with forefield soils and (b) marine terminating glacier with proglacial fjord. Source: Chu *et al.*, (2014).

### 1.3.1 Using forefields to investigate microbial succession

Terrestrial soils, exposed following glacial retreat, pose an interesting opportunity to study initial soil colonisation and succession (Edwards and Cook, 2015). As these soils have been beneath ice for thousands of years, soil development is limited, alongside the absence of established flora or microbial communities (Tscherko *et al.*, 2003). Furthermore, low temperatures and slow weathering rates mean that soil development occurs over longer timescales, and therefore the successional pathways are more identifiable (Bradley *et al.*, 2014). Consequently, it is possible to use these soils to investigate how microbial communities first colonise, and subsequently advance over time, whilst modifying soil development (Schulz *et al.*, 2013). This is typically carried out by utilizing a chronosequence approach, whereby a space-for-time substitution is used, moving away from the glacier terminus along a perpendicular transect (Tscherko *et al.*, 2003). The bulk of chronosequence-based studies have focused on the succession of soil structure and plant matter, whilst limited attention has



been paid to the advance of microbial communities (Edwards and Cook, 2015; Matthews and Vater, 2015).

Whilst chronosequence studies are fundamental to our current understanding of soil succession, it is important to acknowledge that chronosequences only show a record for a single location and time point (Fastie, 1995). Forefields are heterogeneous, with many micro-environments, soil structure disparities, differences in water and nutrient availability, and subsequently variations in floral and faunal diversity, all occurring within a single foreland (Duc *et al.*, 2009). Furthermore, these environments may change with seasonal environmental modifications or disturbances (Fastie, 1995; Bradley *et al.*, 2015). Consequently, considerations must be taken when extrapolating information obtained from a single chronosequence to other environments.

### 1.3.2 Plant succession and soil development

Investigating soil development is key to understanding how terrestrial environments change over time, the driving factors behind this, and the causes of heterogeneous succession both within and between environments (Hodkinson *et al.*, 2003). Soil development is mediated by key environmental factors, such as the local bedrock and topography, the climate, the time available, alongside both plant and microbial communities (Schulz *et al.*, 2013). Cold environments, which exhibit slow rock weathering rates, are preferentially used for the study of soil succession, due to the slow (and therefore traceable) developments over chronosequences (D'Amico *et al.*, 2014; Bradley *et al.*, 2015). Much of the research on succession has been focused on plant establishment and development (Frenot *et al.*, 1998; Hodkinson *et al.*, 2003).

The key trends identified with soil succession are an increase in nutrients, organic carbon and reduction in pH, in line with the development of vascular plants (Ohtonen *et al.*, 1999; Strauss *et al.*, 2009; Knelman *et al.*, 2012). Plants have been deemed as fundamental in enhancing soil stabilisation, modifying the biogeochemical properties of soils (i.e. building nutrient and organic matter pools through litter and root exudates) and subsequently facilitating the colonisation of higher plant communities (D'Amico *et al.*, 2014; Matthews and Vater, 2015).

Soil and vegetation development do not occur at the same rate, nor follow the same successional pattern between different environments (Hodkinson *et al.*, 2003). This relates to the initial soil conditions, local climatic factors and acting disturbances (Moreau *et al.*, 2008; Sattin *et al.*, 2009). Some plant communities may be rapidly established, whilst for others,

typically in cold, nutrient limited environments, plant free periods exceeding 50 years may be identified (Sattin *et al.*, 2009).

### 1.3.3 Microbial Succession in glacial forefields

Microbial succession in newly exposed soils is a more recent subject of interest, in comparison to the study of soil structure and plant colonisation (as reviewed in Bradley *et al.*, 2014). Microbial communities play an important role in developing soil structure and biogeochemical cycles, therefore influencing soil physicochemical characteristics and vegetation (Tscherko *et al.*, 2003; Hahn and Quideau, 2013). Microbial communities often act as the primary colonisers of forefield soils exposed by glacier retreat (Schmidt *et al.*, 2008). Examining this initial colonisation is key for our understanding of how life may first be established and may act as analogous to our understanding of extra-terrestrial life (Bradley *et al.*, 2015). Furthermore, initial colonisation, and subsequent succession is important for highlighting the pathways to fertile soil, the key factors involved in the heterogeneity between environments and if these may be modified with global climate change (Schulz *et al.*, 2013). Much of the research surrounding the importance of initial microbial colonisers has been based on glacial forefields, in locations such as the Damma Glacier, Switzerland and Midtre Lovénbreen glacier, Svalbard (Mindl *et al.*, 2007; Schmidt *et al.*, 2008; Schulz *et al.*, 2013).

The pioneer microbial colonisers may be significant in the development of soil nutrient stocks, regulating soil pH and promoting stability for plant colonisation through the release of EPS (Sattin *et al.*, 2009; Schulz *et al.*, 2013; Wojcik *et al.*, 2019). This may facilitate the colonisation of further microbial communities and higher plant life, which depend on the presence of labile nutrient pools and a degree of soil stability (Chapin *et al.*, 1994; Hahn and Quideau, 2013). The release of nutrients is particularly significant in extreme locations, such as glacial forefields, whereby initial soils are likely to be oligotrophic (Chapin *et al.*, 1994). However, the significance of the initial colonizers in developing labile nutrient pools has been debated (Nicol *et al.*, 2005). Jumpponen (2003) identified aeolian deposition as the most significant source of nutrients to newly exposed soils, alongside nutrients present in the initial soils themselves, whilst suggesting that microbial communities were dormant (Jumpponen, 2003). However, further research has identified the active contribution of initial microbial colonisers to nutrient stocks during the first 20 years of colonisation (Schmidt *et al.*, 2008). The action of photoautotrophic cyanobacteria and free living diazotrophs can build carbon and nitrogen pools in the soils, however aeolian deposition may also be a significant contribution to nutrient stocks (Nicol *et al.*, 2005; Schmidt *et al.*, 2008).

Furthermore, the composition of the initial coloniser communities has been the subject of debate (Kastovska *et al.*, 2005). One research body identifies autotrophic organisms as the dominant entities following soil exposure, facilitating the build-up of nutrients and organic carbon and the subsequent colonisation by heterotrophs and higher plants (Hodkinson *et al.*, 2002). However, more recently, studies have indicated the presence of heterotrophic communities, prior to the establishment of autotrophs (Tscherko *et al.*, 2003; Bardgett *et al.*, 2007). These heterotrophs can deplete the ancient autochthonous organic matter and nutrients that are already present in the exposed soils or use allochthonous nutrients received by aeolian deposition (Kastovska *et al.*, 2005; Bardgett *et al.*, 2007). Following the depletion of these stocks, photoautotrophic and diazotrophic organisms gain a competitive advantage and therefore become prevalent (Chapin *et al.*, 1994; Sattin *et al.*, 2009).

It has been shown that microbial communities in forefields become more abundant, active and diverse with succession, due to the enhanced availability of organic matter, nutrients, and moisture (Kastovska *et al.*, 2005; Nicol *et al.*, 2005; Mindl *et al.*, 2007; Schütte *et al.*, 2010). However, some studies do not show this trend, which may be attributed to an increase in competition between microbes (Sigler and Zeyer, 2002; Zumsteg *et al.*, 2013). Distinct shifts in community composition and functional diversity have been identified, indicating that the environmental factors present at each stage of succession are effective in selecting the dominant microbial communities (Tscherko *et al.*, 2003; Nemergut *et al.*, 2007; Frey *et al.*, 2013). These environmental factors include nutrient and organic matter availability, disturbances, water flow pathways, soil pH and mineralogy (Sakata Bekku *et al.*, 2004; Frey *et al.*, 2013). Consequently, whilst these factors drive changes along the chronosequence within a forefield, they also stimulate differences between forefields (Sigler and Zeyer, 2002; Liu *et al.*, 2012). Research by Schütte *et al.*, (2010) identified significant differences between bacterial communities in two adjacent glaciers, which share bedrock and dominant climate. These differences were attributed to UV exposure and moisture availability, largely in relation to topography, alongside the presence of vegetation (Schütte *et al.*, 2010). This is supported by Meola *et al.*, (2014), who show that environmental conditions are key drivers between successional differences, however stresses the importance of soil mineralogical properties in mediating bacterial community composition (Meola *et al.*, 2014).

More recently, microbial succession in glacial forefields has been represented by biogeochemical models, with the aim of delineating the complex nature of microbial dynamics and their role in soil development (Bradley *et al.*, 2014). Bradley *et al.*, (2015) presents the mathematical model SHIMMER, with the aim of simulating the initial microbial colonisation, prior to plant establishment. Whilst there are key uncertainties, such as the degree of aeolian

deposition, organic matter bioavailability and microbial activity constants, models such as SHIMMER provide new insights, supporting field studies on microbial succession (Bradley *et al.*, 2015). Furthermore, incorporating biogeochemical models will become increasingly important for understanding modifications with global climate change (Bradley *et al.*, 2014).

The above discussion has highlighted the importance of environmental variables in controlling the rate, abundance and diversity of microbial succession (Göransson *et al.*, 2011). These factors include the soil pH, oxygen, salinity, temperature, soil moisture, organic matter and nutrients, disturbances and competition (Hodkinson *et al.*, 2003; Mindl *et al.*, 2007). In particular, studies have noted nitrogen limitation as prevalent in newly exposed glacial soils (Kastovska *et al.*, 2005; Mindl *et al.*, 2007). Whilst initial soils are likely to be oligotrophic, nutrient concentrations following microbial succession are likely to be a result of the in-situ biotic and abiotic nutrient cycling, alongside the prevalence of aeolian deposition (Kastovska *et al.*, 2005; Schmidt *et al.*, 2008). Consequently, the following section will discuss nitrogen dynamics and the role of soil microbial communities during primary succession.

#### 1.3.4 Microbial nitrogen cycling in glacial forefields

Nitrogen is a fundamental nutrient in forefield soils, required by plants and microbes for protein synthesis (Treseder, 2008). Bioavailable nitrogen, which can be readily assimilated, consists of organic nitrogen (ON), ammonia, nitrite ( $\text{NO}_2^-$ ) and nitrate ( $\text{NO}_3^-$ ) (Barber, 1995; Bremner, 1965). The main sources of bioavailable nitrogen to forefield soils are from bedrock weathering, nitrogen-fixing microorganisms (diazotrophs), allochthonous deposition, degradation of organic material and in washing of snowmelt (Bradley *et al.*, 2015). Typical nitrogen concentrations range between  $0.2 - 2 \text{ mg g}^{-1}$  in newly exposed soils, with studies observing a general increase with soil development (Strauss *et al.*, 2009; Bradley *et al.*, 2014).

Nitrogen-fixing microorganisms are often shown to be prevalent in the initial pioneer communities (Duc *et al.*, 2009; Strauss *et al.*, 2009). If nitrogen levels in newly exposed soils are limited, nitrogen fixers obtain a competitive advantage, and readily colonise (Ohtonen *et al.*, 1999; Strauss *et al.*, 2009). This nitrogen fixation builds the bioavailable soil nitrogen stocks and facilitates the succession of higher microorganisms and vascular plants (Bradley *et al.*, 2015). Nemergut *et al.*, (2007) exemplifies this, utilising a chronosequence approach in three un-vegetated Peruvian successional soils. Low initial nitrogen concentrations were developed by nitrogen-fixing microbes, which increase the habitability of the soils for plant colonisation (Nemergut *et al.*, 2007). Studies have identified a plenitude of nitrogen fixation, mineralisation and denitrification associated genes in glacial forefields (Deiglmayer *et al.*,

2006; Duc *et al.*, 2009). Duc *et al.*, (2009) built upon this, and identified active nitrogen fixation was possible in forefield soils, through using assays sourced from Damma glacier soils (Duc *et al.*, 2009). This therefore indicates the importance of free-living diazotrophs to nitrogen cycling, prior to the establishment of plants (Duc *et al.*, 2009).

However, additional sources, alongside nitrogen fixation have been identified as significant contributors to forefield nitrogen stocks (Bradley *et al.*, 2014). Brankatschk *et al.*, (2011) suggests that allochthonous deposition and remineralisation of overridden organic matter contributes significantly more than microbial communities to soil nitrogen stocks, supported by the limited number of *nifH* genes (encoding nitrogen fixation), found in Damma Glacier soils (Brankatschk *et al.*, 2011). A significant increase in soil nitrogen may also be observed following the colonisation of soils by vascular plants, attributed to nitrogen from symbiotic root associated nitrogen fixers, and degraded plant litter (Bradley *et al.*, 2014).

Alongside nitrogen fixation, other components of the microbial nitrogen cycle have been identified in forefield soils (Brankatschk *et al.*, 2011). Kandeler *et al.*, (2006) identified an increase in denitrification-associated gene copies with soil development, which may be related to anoxic conditions from high moisture, plant dominated soils (Schluz *et al.*, 2011). Furthermore, Brankatschk *et al.*, (2011) identified nitrogen mineralisation, from decomposing organic matter as the most significant component of the nitrogen cycle in forefield soils. The findings also indicated low nitrogen fixation, nitrification and denitrification rates initially, however increased with soil age, related to increases in bioavailable nitrogen stocks (Brankatschk *et al.*, 2011). Importantly, gene copy numbers were high throughout, and even during periods of low activity, highlighting the mismatch between gene abundance and activity (Brankatschk *et al.*, 2011).

Consequently, nitrogen dynamics within forefield soils are complex, and likely to vary both within and between locations, largely in relation to the prevalent environmental characteristics (Deiglamayer *et al.*, 2006). For example, in areas which nitrogen is restricted in the bedrock mineralogy, such as Damma glacier, Switzerland, the soils rely on nitrogen fixation and aeolian deposition as the primary sources (Schulz *et al.*, 2013). However, in areas whereby nitrogen is readily sourced in initial soils from bedrock weathering, may depend less on microbial and allochthonous sources (Bradley *et al.*, 2014).

### 1.3.5 Carbon, phosphorous and sulfur in glacial forefields

Nutrient and organic matter inputs to glacial forefields can come from allochthonous sources, such as supraglacial or subglacial runoff, bird and mammal droppings and aeolian deposition (Zumsteg *et al.*, 2012; Bradley *et al.*, 2014). Additionally, in situ microbial activity can provide autochthonous sources of carbon, nitrogen, phosphate and sulfur (Hahn and Quideau, 2013).

Carbon is fundamental to microbial and plant growth, providing the backbone for biological molecules. A range of carbon compounds can be utilized by microbes, such as CO<sub>2</sub> which is fixed by autotrophic bacteria and plants, organic carbon molecules which are bioavailable for heterotrophic organisms, wind-blown hydrocarbons and ancient carbon from overridden soils (Hodkinson *et al.*, 2002; Guelland *et al.*, 2013). The total organic carbon (TOC) content of forefield soils typically ranges between 0.1 – 40 mg g<sup>-1</sup> and increases with soil succession (Guelland *et al.*, 2013; Bradley *et al.*, 2014). Consequently, TOC content in forefields has been shown to increase in relation to soil age, attributed to the development of microbial communities and vegetation (Conen *et al.*, 2007). Autotrophic microbes have been proposed as key facilitators for the build-up of organic carbon in forefield soils through CO<sub>2</sub> fixation, particularly in newly exposed soils (Strauss *et al.*, 2012). However, TOC can be supplied by aeolian deposition of materials such as soot or may already be present in soils in the form of ancient organic carbon, exposed by glacier retreat (Stibal *et al.*, 2008; Guelland *et al.*, 2013).

Phosphorous is a key limiting nutrient for microbes and plants, used in the synthesis of nucleic acids and adenosine triphosphate (ATP). Phosphorous is commonly sourced from weathering of underlying bedrock in forefields and therefore soil mineralogy can be a key control on microbial growth potential in initial soils (Egli *et al.*, 2012; Bradley *et al.*, 2014). Similarly, to TOC, phosphorous content in forefield soils has been shown to increase with succession, ranging between 2 – 8 ug g<sup>-1</sup> (Bradley *et al.*, 2014). Furthermore, whilst not a required nutrient for all microbial growth, sulfur is cycled in soils through microbial redox reactions (Koltz *et al.*, 2011). Sulfur oxidation is used by microbes for energy and is coupled to the reduction of oxygen or nitrate, in aerobic or anaerobic conditions, respectively (Wainwright, 1978). The process oxidises elemental sulfur (S<sup>0</sup>) or reduced sulfur such as sulfide (H<sub>2</sub>S), to sulfate (SO<sub>4</sub><sup>2-</sup>) (Wainwright, 1978). Conversely, SO<sub>4</sub><sup>2-</sup> can be used as a terminal electron acceptor for microbial oxidation of organic matter, in the absence of oxygen, forming reduced H<sub>2</sub>S (Widdel and Hansen, 1992). Evidence for microbial sulfide oxidation has been found across glacier forefields (Szykiewicz *et al.*, 2013; Wolicka *et al.*, 2014). Bedrock mineralogy and subsequent weathering is a key source of sulfide to the soil microbial sulfur cycle (Synkiewicz *et al.*, 2013).

## 1.4 Proglacial fjords and microbial communities

### 1.4.1 Fjord systems

Proglacial fjords are at the interface between glacial meltwater and the open ocean. Glaciers supply freshwater, sediment and nutrients to saline nearshore waters (Davila *et al.*, 2002; Lawson *et al.*, 2014; Gutiérrez *et al.*, 2015; Hawkings *et al.*, 2015; Figure 1.1). This freshwater flux often stratifies the water column, creating a surface freshwater lens on top of dense saline waters (Prado-Fiedler, 2009). Fjord systems draining glacier ice fields, such as those in Southern Chile, are hotspots of primary productivity (Iriarte *et al.*, 2007). These regions have been shown to have high rates of primary productivity, diverse ecosystem structures and to harbour unique surface and benthic communities (Iriarte *et al.*, 2007). Fjords may be subject to gradients of meltwater influence, with reduced salinity, increased nutrients and suspended sediment close to the glacier terminus (Davila *et al.*, 2002; Cowton *et al.*, 2012; Lee *et al.*, 2013; Wadham *et al.*, 2013). The biological implications of these gradients have yet to be fully understood. Bacterial community composition and function are likely to modify along these gradients, in line with salinity, light and nutrient content (Gutiérrez *et al.*, 2015). Changes to bacterial communities may have impacts on local biogeochemical cycles, and the wider food chain (Meerhoff *et al.*, 2013).

Understanding the influence of glaciers on microbial communities is key in the context of global climate change. Glaciers, such as those in the Patagonian ice fields, are retreating and therefore proglacial fjords and in-situ microbial communities are likely to be influenced by short term increases in ice melt (Iriarte *et al.*, 2014). By investigating and understanding microbial communities in current fjord systems, the future impacts of rising freshwater fluxes can be better understood.

### 1.4.2 The influence of glacial meltwater

When glacial meltwater drains through fjord systems, both vertical and horizontal environmental gradients in nutrients, light, salinity and sediment can be created (Iriarte *et al.*, 2014). This is due to the mixing of saline ocean water, and fresh glacial meltwater, and subsequent stratification of the water column (Iriarte *et al.*, 2014). In Chilean fjords, high nutrient, saline, Sub-Antarctic water flows at depth, below a freshwater layer of glacial meltwater (Aracena *et al.*, 2011). Mixing between the layers is crucial to distribute nutrients from deep water masses (Aracena *et al.*, 2011). Environmental oxygen gradients may occur due to the stratified water column, and deep hypoxic zones may be created due to the lack of

replenishment of oxygen from surface waters (Silva and Vargas, 2014). Distinct gradients in nutrient content are probable due to the stratification, with surface waters likely to be low in nutrient content due to lack of mixing with deeper high nutrient bottom waters (González *et al.*, 2013). This has been exemplified in the saline Sub-Antarctic bottom waters in Chilean fjords, which are high in both nitrogen and phosphate compared to surface waters (González *et al.*, 2013). Consequently, glacial runoff may be a key source of nutrients, such as iron, silicon, phosphate and nitrogen to nutrient limited surface waters in Chilean fjords (Hood *et al.*, 2009; Wadham *et al.*, 2013). Glacial meltwaters have been shown to supply significant levels of silicon to surface waters, however may be low in nitrate (González *et al.*, 2013). High nutrient bottom waters may therefore also be crucial in stimulating primary productivity through the supply of nitrate (and other limiting nutrients) to surface waters (Iriarte *et al.*, 2014). This has been exemplified by spring blooms stimulated by vertical mixing of the water column (González *et al.*, 2013). Distinct gradients in light attenuation may be created vertically and horizontally in glacially-fed fjords, due to the outflow of sediment in glacial runoff (Keck *et al.*, 2000). Glacier runoff has been shown to carry high sediment loads from glacial weathering, which are exported to downstream ecosystems (Hawkings *et al.*, 2015). This sediment load may restrict light levels in surface waters close to the glacier outflow, however this effect is likely to reduce with distance from the runoff source (Keck *et al.*, 2000; Hawkings *et al.*, 2015).

#### 1.4.3 Microbial communities and fjord systems

The vertical and horizontal environmental gradients in nutrients, salinity, light and sediment created by glacial outflow may influence microbial community structure and biogeochemical function (González *et al.*, 2013; Gutiérrez *et al.*, 2015). A stratified water column may encourage distinct regions of microbial growth depending on salinity tolerance, with organisms associated with freshwater environments more likely to be isolated in surface waters (Gutiérrez *et al.*, 2015). Research by Gutiérrez *et al.*, (2015) using 16s rRNA found distinctive seasonal meltwater influences on fjord microbial community structure from the Jorge Montt glacier (SW Chile). In Autumn, the high availability of meltwater was shown to encourage the presence of freshwater-tolerant organisms, alongside a greater bacterial richness in surface waters (Gutiérrez *et al.*, 2015). Additionally, research by Dethier and Schoch (2005) found that changes in salinity to benthic populations reduced species richness. Nutrient and temperature gradients have also been shown to be key controls on the composition of fjord communities under the influence of glacial meltwater (Renner *et al.*, 2012). The outflow of particulate matter in glacial runoff may have a significant influence on microbial community composition, function and primary productivity (Iriarte *et al.*, 2014). Although particulate matter may supply additional nutrients to surface waters (Hawkings *et al.*, 2015), the sediment may reduce light for



photosynthetic organisms (González *et al.*, 2013). The reduction of light attenuation, alongside the increase in turbidity caused by glacial outflow, has been shown to limit primary productivity, reduce microbial abundance and modify community composition in glacially fed fjords (Keck *et al.*, 2000; Iriarte *et al.*, 2014). Research by Aracena *et al.*, (2011) on the primary productivity of Chilean fjords showed that glacially fed waters had the lowest levels due to increased turbidity and reduced light levels. Overall, glacier outflow may have a significant influence on microbial community structure and function, through the development of environmental gradients (Iriarte *et al.*, 2014). The impacts of this may be identified at the scale of the microbial community structure but also on the wider ecosystem function, such as the distribution and growth of fish (Landaeta *et al.*, 2012; Gutiérrez *et al.*, 2015).

The microbial community of fjord benthic sediments may also be influenced through horizontal gradients in salinity, nutrients and temperature, created by glacial outflow (Keck *et al.*, 2000). In particular, organic matter has been shown to increase from oceanic sediments to inner fjords, in relation to discharge from terrestrially fed rivers (Aracena *et al.*, 2011). However, glacially fed fjords are likely to have reduced organic matter contents, largely due to the high inorganic sediment fluxes in rivers created from glacial weathering (Silva *et al.*, 2008; Aracena *et al.*, 2011). The limited organic matter content of glacier fed fjord sediments may also relate to the reduced primary productivity in surface waters from turbidity and limited light levels (Aracena *et al.*, 2011). Consequently, glacier meltwaters may influence the structure, function and productivity of microbial communities in both the water column and the sediments of fjords.

### **1.5 Proglacial environments and climate change**

The structure and function of proglacial microbial communities may be substantially modified with global climate change (Yde *et al.*, 2011). Enhanced warming of polar regions will act to accelerate glacier melt rates and extent, promoting deglaciation (IPCC, 2013). Continued ice melt will expose undeveloped soils in proglacial forefields and drive increased meltwater fluxes into fjords (IPCC, 2013).

It is important to understand forefield microbial communities and their succession, alongside how this may change with global climate change (Schulz *et al.*, 2013; Bradley *et al.*, 2014). Climate warming may modify soil successional pathways, the dominant microbial and plant communities, the biogeochemical cycling in these environments and the composition and fertility of developed soils (Davidson and Janssens, 2006; Schulz *et al.*, 2013). Fjords draining glacial meltwater will experience enhanced freshwater fluxes with ice melt (Davila *et al.*, 2002;

IPCC 2013). Changes to the salinity balance may modify the dominant microbial community structure and function, enhance water column stratification and subsequently modify local biogeochemical cycles (Meerhoff *et al.*, 2013; Gutiérrez *et al.*, 2015). Additionally, enhanced meltwater fluxes may increase sediment inputs to fjords and modify nutrient content (Iriarte *et al.*, 2010; González *et al.*, 2013). Not only does this have implications on the in-situ microbial communities, but also on nutrient dynamics and wider ecosystem functioning (Gutiérrez *et al.*, 2015). This is particularly significant in regions such as South West Chile, due to the commercial importance of Salmon fisheries, which rely on the supporting ecosystem to maintain productivity (Iriarte *et al.*, 2010). By Investigating the current microbial community structure function and diversity, we can better understand how these may change in future years.

## **1.6 Methodological considerations for investigating microbial communities**

### 1.6.1 Methods available

A variety of methods are available to investigate microbial community composition and functional diversity. These techniques range from traditional microbial culture-based studies, to more recent molecular methods, such as metagenomics (Ward *et al.*, 1990; Janssen *et al.*, 2002; Riesenfeld, 2004; Daniel, 2005; Teeling and Glockner, 2012). Whilst single organisms can be investigated through culture-based methods, developments in DNA sequencing have enabled whole environmental samples to be evaluated at once, using 16s amplicon sequencing and more recently, metagenomics (Tringe *et al.*, 2005; Fierer *et al.*, 2012). These methods are particularly interesting as they enable researchers to isolate the taxonomic diversity (16s), or both the taxonomic and functional potential of microbial communities (metagenomics) (Handelsman, 2004; Daniel, 2005; Fierer *et al.*, 2012). This provides a more comprehensive insight into natural microbial communities, in comparison to growing a single organism independently (Torsvik and Øvreås, 2002). The available techniques to investigate microbial communities are discussed in this section, in the context of scientific developments and recent advances.

### 1.6.2 Culture-based studies

Microbial cultures have been extensively used in soil research to investigate single organisms, isolated from environmental samples (Teeling and Glockner, 2012). Cultures are used to identify new organisms, optimum growth conditions, the ability of organisms to grow under extreme stressors (such as darkness and low temperatures), metabolic functions and more

recently, used in DNA sequencing (Kirk *et al.*, 2004). For example, Frey *et al.*, (2013) carried out microbial cultures of bacteria isolated from the Damma glacier forefield to understand their role in granite weathering and therefore soil formation. However, less than 1% of organisms can be successfully cultured in laboratory conditions (Riesenfeld, 2004; Teeling and Glockner, 2012). This was first identified as the 'great plate count anomaly' whereby laboratory cultures are unable to grow population sizes equal to those observed in natural samples (Handelsman, 2004). Consequently, the majority of organisms and their metabolisms, remain unidentifiable through culture-based techniques (Reed *et al.*, 2014; Riesenfeld, 2004). This has led to the emergence of molecular techniques, to investigate microbial genetic and functional diversity, without the need to culture (Streit and Schmitz, 2004; Thomas *et al.*, 2012).

### 1.6.3 First generation DNA sequencing

The emergence and subsequent developments of DNA sequencing technologies have revolutionised our understanding of microbial communities (Rondon *et al.*, 2000; Torsvik and Øvreås, 2002). Understanding the DNA sequence of a single organism (genomics), or multiple organisms (metagenomics) enables microbial taxonomy, phylogenetics and functional diversity to be inferred (Riesenfeld *et al.*, 2004; Tringe *et al.*, 2005; Tringe and Rubin, 2005). This is particularly significant for uncultured organisms, as sequencing allows an examination of the genetic structure, metabolic pathways and community diversity, without the need for laboratory growth experiments (Handelsman, 2004; Daniel, 2005; Xu, 2006).

First generation DNA sequencing was based on the chain termination (or Sanger) method (Sanger *et al.*, 1977; Swerdlow *et al.*, 1990; Wooley *et al.*, 2010). The isolated DNA or genes of interest were first amplified prior to sequencing, either by cloning plasmid vectors, or through artificial replication in using polymerase chain reaction (PCR) (Erlich, 1989; Weisburg *et al.*, 1991; Newton *et al.*, 1997). In PCR, the DNA strands are denatured, a primer annealed and subsequently extended by DNA polymerase, to amplify the number of fragments (Shendure and Ji, 2008; Wooley *et al.*, 2010). Di-deoxynucleoside triphosphates (ddNTPs) are incorporated, which act to halt DNA chain extension (Fierer *et al.*, 2005). The DNA sequence is identified through electrophoresis, based on the fluorescent tag of the terminating ddNTPs (Shendure and Ji, 2008). Sequencing is repeated to ensure all areas of the gene (or genome) of interest are covered, however the read lengths returned by Sanger sequencing (up to 1,000 base pairs) may still leave unresolved long repeat sequences (homopolymer regions) (Wooley *et al.*, 2010).

#### 1.6.4 Taxonomic marker genes

PCR-based amplification is now fundamental in DNA sequencing and has given rise to improved taxonomic assignment of environmental samples (Riesenfeld *et al.*, 2004; Blazewicz *et al.*, 2013). Universal genes, such as 16s rRNA in bacteria, can be used as taxonomic marker genes, whereby gene specific primers are used in PCR, and the target regions are sequenced, rather than full genomes (Nübel *et al.*, 1997; Handelsman, 2004). Sequences are subsequently compared to databases, and operational taxonomic units (OTUs) defined based on rRNA sequence similarity (Streit and Schmitz, 2004). The genetic structure and function of organisms in environmental samples is therefore inferred from the nearest sequenced neighbour, identified on the database (Riesenfeld *et al.*, 2004). Additionally, non-universal, function specific, target genes can be used, such as *nifH* for nitrogen fixation, to understand the taxonomic diversity of a microbial population involved in a certain metabolic pathway (Gaby and Buckley, 2012). Molecular markers have become a popular mechanism for understanding taxonomic diversity, however do not isolate overall community microbial function or abundance, as the remaining (un-targeted) functional genes are not sequenced (Thomas *et al.*, 2012).

#### 1.6.5 Next generation DNA sequencing

Advances in sequencing technologies have led to the emergence of next generation (NGS) (or massively parallel) sequencing (Mardis, 2008). These technologies vastly improve on the speed, depth and cost of traditional Sanger sequencing methods (Metzker, 2010). This is due to the high throughput of NGS machines and the ability to sequence millions of DNA fragments at one time (Mardis, 2008; Shendure and Ji, 2008). The popular NGS platforms include Illumina, PacBio and Ion Torrent (Metzker, 2010; Liu *et al.*, 2012). The use of these platforms has risen with developments in bioinformatics, such as increased data storage capacities and techniques for large dataset analysis (Mardis, 2008; Horner *et al.*, 2009). Each platform provides a unique method for DNA sequencing, however they are all based on a 'cyclic array' approach (Shendure and Ji, 2008). This is where DNA is cycled through repeated steps of enzyme-initiated manipulation and DNA bases are read by imaging (Shendure and Ji, 2008). NGS techniques can be applied to both single genomes and metagenomes, to sequence specific genes, or the whole genomes/communities (Quaiser *et al.*, 2002; Tringe and Rubin, 2005; Delmont *et al.*, 2011). All NGS sequencing platforms can produce mate-pair or paired-end reads, whereby each DNA strand is sequenced in two directions, thereby providing the distance between each paired read (Shendure and Ji, 2008). This helps resolve structural rearrangements, such as INDELS (insertions or deletions), and repeat regions, especially

when a fully sequenced reference genome is not available (Hajirasouliha *et al.*, 2010; Metzker, 2010; Miller *et al.*, 2010). NGS platforms produce varying read lengths, however the longer the read length provided, the higher the sequencing error rate (Teeling and Glockner, 2012). However, longer reads are beneficial for resolving large repeat regions, enhancing taxonomic and functional annotations and revealing structural rearrangements (Wommack *et al.*, 2008). Illumina sequencing is currently widely used due to reasonable read lengths (150 – 300bp) and minor error rate (Illumina, 2018).

However, NGS sequencing technologies still retain several limitations (Mardis, 2008; Ansorge, 2009; Alkan *et al.*, 2011; Teeling and Glockner, 2012). The cost of purchasing the sequencing platforms may hinder their use in some labs, alongside accounting for the cost of each round of DNA sequencing (Ansorge, 2009; Grada and Weinbrecht, 2013). Furthermore, as with Sanger sequencing, homopolymer regions may not be accurately resolved, due to spanning longer lengths than the short reads returned through sequencing (Alkan *et al.*, 2011; Grada and Weinbrecht, 2013). Additionally, base call errors generally increase towards the 3' end of read fragments, and therefore the error rate is not consistent throughout the sequencing run (Miller *et al.*, 2010). Furthermore, the datasets returned by NGS place high computational demands on downstream analysis, often requiring complex bioinformatic pipelines to assemble, annotate, and investigate the resulting DNA sequences (Horner *et al.*, 2009; Grada and Weinbrecht, 2013).

Further technological advances from NGS platforms have enabled the development of third generation sequencing platforms (Schadt *et al.*, 2010). Platforms have been provided by Pacific Biosciences (SMRT), Oxford Nanopore, Life Technologies (FRET) and Ion Torrent (Branton *et al.*, 2008; Wash and Image, 2008; Oszolak, 2012; Quail *et al.*, 2012; Roberts *et al.*, 2013). Whilst these technologies are still being established, they aim to improve the time and cost efficiency of DNA sequencing (Wash and Image, 2008; Schadt *et al.*, 2010). SMRT (Single Molecule sequencing in Real Time), produced by PacBio, is based on sequencing individual fragments of DNA, and can provide read lengths of up to 15kb, however error rates are still high (Roberts *et al.*, 2013).

#### 1.6.6 Metagenomics

Metagenomics is a more recent approach to DNA analysis and involves sequencing the complete DNA of a microbial community without PCR amplification (Handelsman, 2004). As multiple genes are sequenced, this approach allows both taxonomic and functional annotation of the microbial community (Tringe *et al.*, 2005). Consequently, metagenomics provides an

alternative to 16s amplicon sequencing, as the presence of functional genes allows an insight into the community's metabolic potential, and thus, biogeochemical importance (Handelsman, 2004; Daniel, 2005). Additionally, because this approach does not require DNA amplification, it is not susceptible to the bias typically introduced through PCR (Risenfeld *et al.*, 2004). However, much greater quantities of DNA are required prior to sequencing, which restricts the analysis of low biomass samples (Yilmaz *et al.*, 2010). As metagenomics supplies substantially more data than amplicon sequencing and requires greater downstream processing, the post-sequencing analysis process is more demanding (Schmieder and Edwards, 2011). The relatively recent introduction of this technique means there is a lack of formalised guidance or software for the analysis of metagenomic data, hindering its uptake by the scientific community.

### Metagenome sequence assembly and annotation

NGS metagenome sequencing supplies short DNA fragments, which depending on the sequencing platform, can range from 100bp (SOLiD) to 1000bp (Sanger) in length (Miller *et al.*, 2010). These raw reads can be directly interpreted using analysis packages such as MG-RAST, for taxonomic and functional annotation (Glass *et al.*, 2010). As these annotations are often made by aligning the DNA fragments to genes in databases, such as NCBI GenBank, the short length of the fragments limits the quality and magnitude of the possible alignments (Howe *et al.*, 2014; Vázquez-Castellanos *et al.*, 2014). Consequently, to provide better sequence annotations, the read fragments can be assembled into longer contigs using a sequence assembler (Nagarajan and Pop, 2013). This is particularly important for accurately understanding the metabolic function or phylogenetic diversity of the sequenced organisms (Pignatelli and Moya, 2011). This is because annotating genes with functional and taxonomic identity involves alignment to databases, and sequences which are longer in length will provide more accurate alignments (Van der Walt *et al.*, 2017).

### Sequence assembly

Sequence assembly aims to reconstruct the gene, genome or metagenome that was sequenced (Miller *et al.*, 2010). This involves identifying overlaps in the raw DNA read fragments, using an assembly algorithm, generating longer fragments (contigs) (Narzisi and Mishra, 2011). This is particularly important when using NGS platforms which supply short read fragments that are difficult to directly interpret (Mende *et al.*, 2012). The length of overlaps between raw reads are defined by the user, based on a specified k-mer value (a sequence of k- base calls) (Narzisi and Misha, 2011). The contigs generated can be built into a scaffold,

which define the orientation and order that the contigs occurred in the original genome (Myers *et al.*, 2000; Peng *et al.*, 2012). Sequence assemblers assume that similar fragments occur at proximity within a genome, however this may be fundamentally undermined by the presence of repeats that occur throughout the genome, or between genomes (Nagarjan and Pop, 2013). These repeats are less problematic when they fall within the read length, however become harder to resolve when they span longer than the sequencing read length (Narzisi and Misha, 2011). Repeat sequences, short reads and sequencing errors are the key issues that face assembly algorithms (Narzisi and Misha, 2011).

There are two dominant approaches for NGS sequence assembly; reference guided and de novo (Miller *et al.*, 2010; Zhang *et al.*, 2011). Reference guided assembly is used for single genomes, when a 'gold standard' assembled genome is available to assist the assembly of the raw sequence data (Cattonaro *et al.*, 2010). Typically, the raw sequencing reads are aligned against the reference genome, to help guide the assembly (Vezzi *et al.*, 2011). De novo sequence assembly is used in the absence of a reference genome and can be used for both single and metagenome datasets (Zhang *et al.*, 2011). De novo assemblers aim to merge the short sequencing reads based on common overlapping fragments of a specified length (k-mers) (Miller *et al.*, 2010). These assemblers therefore aim to simplify the raw dataset and reproduce the metagenome in the environmental sample (Nagarjan and Pop, 2013). There are three key approaches to de novo assembly, which include; Greedy algorithms, Overlap Layout Consensus algorithms (OLC) and de Bruijn graph algorithms (Miller *et al.*, 2010).

Greedy sequence assembly algorithms work by selecting the highest scoring read overlaps, before merging raw reads into longer contigs (Miller *et al.*, 2010; Zhang *et al.*, 2000). This is carried out until all the available merges have been exhausted (Narzi and Misha, 2011). Heuristic corrections are carried out on each merge, with assembly gaps left in regions where corrections are not possible (Zhang *et al.*, 2000). Greedy algorithms consequently provide a set of assembled contigs, for a single genome or metagenome (Pop, 2009). Examples of assemblers based on this algorithm include PHRAP, CAP3, PCAP, TIGR, SSAKE and Phusion (Sutton *et al.*, 1995; Huang and Madan, 1999; Mullikin and Ning, 2003; Bastide and McCombie, 2007; Warren *et al.*, 2007; Simpson *et al.*, 2009; Pignatelli and Moya, 2011).

OLC algorithms are optimally used on raw reads exceeding 200bp in length (Zhang *et al.*, 2011; Deng *et al.*, 2015). This algorithm works by graphically representing the sequencing reads and the base pair overlaps between them (Miller *et al.*, 2010; Nagarjan and Pop, 2013). The overlaps are generated through pairwise comparisons of the raw reads, by in built aligners (Narzi and Misha, 2011). Each graph node comprises a metagenome read, with the edges

connecting nodes representing the overlaps between reads (Miller *et al.*, 2010). Contigs are graphically shown by the paths running through the graph, connecting nodes (Deng *et al.*, 2015). The aim of OLC algorithms is to simplify the raw assembly graph by connecting each node (raw read) to a single 'Hamiltonian' path (Miller *et al.*, 2010). However, the presence of inter and intra genome repeats and sequencing errors result in multiple diverging paths (Miller *et al.*, 2010; Deng *et al.*, 2015). Whilst OLC algorithms are computationally expensive, they can incorporate both forward and reverse DNA strands, alongside distinguishing the 5' and 3' read ends (Miller *et al.*, 2010; Narzi and Misha, 2011). Examples of OLC based sequence assemblers include Edena, MIRA and Celera (Denisov *et al.*, 2008; Hernandez *et al.*, 2008).

Finally, de Bruijn graph-based assemblers are optimal for short reads around 100bp, such as those obtained from Illumina sequencing (Miller *et al.*, 2010). Similar to OLC, the algorithm represents the assembly graphically, through a series of edges and nodes (Nagarjan and Pop, 2013). Sequencing reads are split up into fragments of length  $k$  ( $k$ -mers) and represented by edges on the graph (Narzi and Misha, 2011). The  $k$ -mer length is specified by the user and must be below the read length (Narzi and Misha, 2011). The defined  $k$ -mer length is significant, as reads significantly longer than length  $k$  will be fragmented, prior to being re-assembled (Miller *et al.*, 2010). Nodes on the graph represent the read overlaps, of length  $k-1$  (Nagarjan and Pop, 2013). Consequently, nodes connect two edges if they share a common sequence of  $k-1$  length (Nagarjan and Pop, 2013). As with OLC algorithms, de Bruijn graphs aim to simplify the assembly by merging reads into contigs and forming a single sequence (an Eulerian path) (Pevzner *et al.*, 2001; Conway *et al.*, 2012). Repeats and errors prevent the assembly of a single large sequence, so contigs are merged and extended before the paths branch, resulting in numerous Eulerian paths (Pevzner *et al.*, 2001; Narzi and Misha, 2011; Pignatelli and Moya, 2011). Miller *et al.*, (2010) identifies three key types of errors that may be represented in a fragmented de Bruijn graph (Miller *et al.*, 2010). 'Bubbles' are created when the Eulerian path splits and subsequently recombines due to errors in the centre of a read (Miller *et al.*, 2010). 'Spurs' are created when a second Eulerian path is generated, due to an error at the 3' end of a sequencing read (Pevzner *et al.*, 2001; Miller *et al.*, 2010). Finally, 'cycles' form when two paths converge due to repeat sequences, present in multiple reads, which ideally need to be separated (Miller *et al.*, 2010). Whilst these graphical representations are computationally expensive, both forward and reverse DNA reads can be represented as  $k$ -mers (Narzi and Misha, 2011). Assemblers that utilise the de Bruijn graph algorithm include Velvet, IDBA, SOAPdenovo, CLC, AllPaths, metaSPAdes and AbySS (Butler *et al.*, 2008; Zerbino and Birney, 2008; Simpson *et al.*, 2009; Li *et al.*, 2010; Miller *et al.*, 2010; Peng *et al.*, 2010; Nurk *et al.*, 2017).



## Sequence annotation

Following DNA sequencing and assembly, contigs can be annotated with their taxonomic and functional identity, using a database (Thomas *et al.*, 2012). Annotation software and pipelines exist, such as MG-RAST, which allow short unassembled reads to be annotated (Wooley *et al.*, 2010; Vincent *et al.*, 2013). However, assembled contigs with longer lengths pose fewer challenging alignments to gene databases and can provide more accurate results, given a high-quality assembly (Thomas *et al.*, 2012). Taxonomic and functional annotation of assembled contigs can be carried out using a range of algorithms and software platforms, such as MEGAN and IMG/M (Huson *et al.*, 2007; Chen *et al.*, 2017). These programs provide homology and sequence similarity alignments to databases such as NCBI GenBank, KEGG and eggNOG, to reveal the taxonomic distribution and functional potential of metagenomes (Kaneisha and Goto *et al.*, 2000; Benson *et al.*, 2005; Huerta-Cepas *et al.*, 2015).

## The choice of assembler

When conducting a genome or metagenome assembly, the user is faced with numerous choices. In the absence of a reference genome, a *de novo* assembly must be carried out, however the assembly algorithm and specific assembler need to be selected (Zhang *et al.*, 2011). There is currently an absence of formalised guidance on assembler selection and evaluation (Arumugam *et al.*, 2010). The decision-making is therefore left to the user, with the dataset size, complexity and availability of computational power all key considerations (Finotello *et al.*, 2011). Consequently, it is important for the user to evaluate and compare assemblers, to make an informed decision (Kurtz *et al.*, 2004; Peng *et al.*, 2012; Howe *et al.*, 2014).

Greedy algorithms are simple, less computationally expensive, and apply well to single genomes, with minimal repeats (Zhang *et al.*, 2011). On the other hand, the graph-based algorithms, OLC and de Bruijn, are more computationally expensive, however function better for complex genomes or metagenome samples, with repeat structures (Zhang *et al.*, 2011). Furthermore, as all NGS platforms vary in terms of the error rate, read lengths, coverage depth and evenness, different assemblers will be better aligned to one platform than another (Earl *et al.*, 2011). Whilst the quality of the assembly does inherit issues from the sequencing platform (base call errors, short read length), the assembler selected will introduce variances through contig length, miss-joined contigs and single nucleotide polymorphisms (SNPs) (Finotello *et al.*, 2011; Salzberg *et al.*, 2012). For example, some assemblers, such as SOAPdenovo prioritise extending the contig lengths, at the expense of accuracy (thereby

increasing the number of erroneous contigs) (Salzberg *et al.*, 2012). It is therefore important to compare different assembly algorithms and assemblers for the sequenced dataset in question, however this is seldom done (Finotello *et al.*, 2011; Lin *et al.*, 2011).

As metagenomics has been newly introduced into the field of microbial ecology, it is still not widely utilised. This largely relates to the cost, computational intensity and popularity of other community analysis approaches such as 16s amplicon sequencing (Handelsman, 2004). However, the use of this approach is rising, with more papers incorporating metagenomes to study functional ecology (Eloe-Fadrosh *et al.*, 2016a). Consequently, there is a growing need for guidance on assembler use and choice for environmental data analysis. Additional guidance may aid the incorporation of this technique into the research community, highlighting the benefit and need for metagenome assembly and evaluation.

#### Selecting an appropriate assembler

To choose an appropriate assembler, several platforms can be compared using the same dataset, and evaluated using selected metrics. However, a range of evaluation metrics exist, none of which provide a complete or standardised evaluation of assembly quality (Arumugam *et al.*, 2010). Traditionally, studies have looked at measures of the assembly size, including the number of contigs produced, the mean contig lengths and the contig N50 length (Earl *et al.*, 2011). However, a long contig length does not fully indicate the assembly quality, as long contigs may be a result of misjoins, and the N50 value cannot be directly compared across assemblies (Finotello *et al.*, 2011). Consequently, an additional measure can look at the accuracy of the assembly, by mapping the assembled contigs to a reference genome if available, identifying any mismatched regions (Miller *et al.*, 2010; Narzisi and Misha, 2011). Quantification of the metagenome coverage may also be utilised, which involves mapping the raw sequence reads to the assembled contigs and identifying the number of which can be correctly aligned (Earl *et al.*, 2011; Huson *et al.*, 2001). The optimal assembly will utilise all raw sequencing reads supplied, and any errors will be random (Kumar and Blaxter, 2010; Bräutigam *et al.*, 2011). Typically, as assemblies tend to trade-off contig length and quality, it is important to combine a series of measures that look at the size, coverage and accuracy of the assembly (Nagarjan and Pop, 2013).

Traditional studies may only choose a single assembler, with selection based on previous findings. This may be due to the complex nature of the raw data and the computational intensity required to carry out an assembler comparison (Najaran and Pop, 2013). Evaluation metrics are often presented to validate the choice of assembler, for example by showing N50

values in comparison to those of other studies (Miller *et al.*, 2010). Some studies have carried out comparisons between assembly algorithms to resolve a specified dataset or have used simulated datasets to test assembler quality (Narzisi and Misha, 2011; Lin *et al.*, 2011). However, these have been restricted to single genome assemblies or using contigs from human samples, or combined human and environmental datasets (Huson *et al.*, 2001; Mavromatias *et al.*, 2007; Earl *et al.*, 2011; Pignatelli and Moya, 2011; Saltzberg *et al.*, 2012). More recently, a study by Vollmers *et al.*, (2017) compared assemblers using real Illumina sequencing data for forest soil and algal biofilms. This study highlighted the importance of comparing assemblers on multiple metrics, however did not investigate parameter settings, all assembly algorithms or use simulated data to test the accuracy of assemblers (Vollmers *et al.*, 2017). Additionally, Sczyrba *et al.*, (2017) compared a range of software for metagenomic data analysis, including assemblers, using artificial metagenomes. However, as this study utilised newly sequenced genomes, the analysis was not specific to bacteria, and in particular, those from heterogeneous environmental samples, such as soils. Thus far, there is no assembler comparison available specifically for studies investigating bacterial communities obtained from sediment or soil.

## **1.7 Summary and research gaps**

Diverse and unique microbial life can be found in proglacial regions, including forefield soils and proglacial fjords. These microbes can be subject to distinct environmental gradients and biogeochemical perturbations created by ice melt, which may modify community structure and function. In proglacial soils, ice melt and glacier retreat expose undeveloped soils, providing an opportunity to study microbial colonisation and development. These microbial communities are likely to modify in line with soil development and may have key roles in soil biogeochemical cycling, such as the nitrogen cycle. Understanding community composition and development in terms of structure and function will aid our understanding of microbial colonisation in extreme environments, and the role of these organisms in nutrient cycling. In proglacial fjords, microbial communities are subject to environmental perturbations from glacial freshwater fluxes. These fluxes create distinct changes to salinity, sediment, light attenuation and nutrients in fjords, which again may influence the microbial community taxonomic and functional diversity. As meltwater fluxes continue to increase with climate change, meltwater may pose challenges to the wider biogeochemical functioning of fjords. This is particularly important as fjords in regions such as Chilean Patagonia have been shown to be hotspots of primary productivity, and support commercially significant fisheries. Understanding current community composition and function will help us understand how microbial communities may contribute to biogeochemical cycling and how this may change in the future.

Metagenomics is a technique which can be used to profile both the taxonomic and functional diversity of microbial communities. It can therefore be applied to the understanding of microbial communities in proglacial regions. The assembly of metagenome sequencing data can improve downstream functional and taxonomic annotation, and thus, can raise the accuracy of ecological conclusions. However, the lack of guidance and complex nature of DNA assembly means there is a minimal uptake within the field of microbial ecology. Selecting an appropriate assembler for the data type in question is paramount, to obtain an improved outcome and minimise erroneous results. Consequently, additional guidance for the selection of metagenome assemblers is needed, to improve the outcome of assembly-based studies and encourage the use of DNA assembly. This is particularly needed in the field of soil microbial ecology, as limited focus has been made to this research group.

## **1.8 Aims and objectives**

**Objective 1:** To compare the performance of five publicly-available metagenome assemblers for soil bacterial communities

Metagenomics involves the study of all genes present in a microbial community sample. The analysis of the community gene pool provides information on both the taxonomic and functional composition. This provides an insight into the potential role of the microbial community in local biogeochemical cycles. Assembling short metagenome DNA sequencing reads is beneficial to obtain longer fragments (contigs) for alignment to functional gene databases. However, there is currently no formalised guidance for the assembly of complex bacterial communities isolated from soil samples. In particular, the choice of metagenome assemblers for this research community requires further investigation. This study investigated the performance five publicly available metagenome assemblers, spanning the three dominant assembly algorithms (OLC, de Bruijn graph and Greedy). The assemblers were evaluated based on a series of artificial soil metagenomes of varying complexity, with evaluation based on assembly size, completeness and contiguity. Overall, the choice of assembler was shown to influence assembly quality and thus downstream functional or taxonomic annotation. The more complex de Bruijn graph assemblers, metaSPAdes and CLC provided the highest quality assemblies. This is because these assemblers have more tuneable parameter settings to fit the data type in question and are better suited for short reads returned from Illumina sequencing. It is hoped that the results from this work will help guide assembler selection for the soils community, alongside highlighting the importance of informed choice in the use of metagenome assembly tools.

**Objective 2:** To investigate the similarities and differences in taxonomic composition of diazotrophic bacteria in metagenomes sampled from four Arctic glacier forefields

Published in FEMS Microbiology (Nash *et al.*, 2018).

Microbial nitrogen fixation is crucial for building labile nitrogen stocks and facilitating higher plant colonization in oligotrophic glacier forefield soils. Here, the diazotrophic bacterial community structure across four Arctic glacier forefields was investigated using metagenomic analysis. In total, 70 soil metagenomes were used for taxonomic interpretation based on 185 nitrogenase (*nif*) sequences, extracted from assembled contigs. The low number of recovered genes highlights the need for deeper sequencing in some diverse samples, to uncover the complete microbial populations. A key group of forefield diazotrophs, found throughout the forefields, was identified using a *nifH* phylogeny, associated with *nifH* Cluster I and III. Sequences related most closely to groups including Alphaproteobacteria, Betaproteobacteria, Cyanobacteria and Firmicutes. Using multiple *nif* genes in a Last Common Ancestor analysis revealed a diverse range of diazotrophs across the forefields. Key organisms identified across the forefields included *Nostoc*, *Geobacter*, *Polaromonas* and *Frankia*. Nitrogen fixers which are symbiotic with plants were also identified, through the presence of root associated diazotrophs, which fix nitrogen in return for reduced carbon. Additional nitrogen fixers identified in forefield soils were metabolically diverse, including fermentative and sulfur cycling bacteria, halophiles and anaerobes.

**Objective 3:** To investigate the bulk microbial community composition along a chronosequence of soil succession in the Midtre Lovénbreen forefield, Svalbard, using metagenomics

Arctic glaciers are currently undergoing retreat with global climate change, revealing undeveloped soils at the glacier terminus. These soils can be used to study succession, utilising a chronosequence based approach along a transect of soil age and development. Understanding microbial communities in the Arctic is important for our understanding of microbial function, diversity and importance in harsh oligotrophic, cold conditions. Additionally, microbial succession in Arctic forefields may provide insights into how extra-terrestrial life may colonise on cold planets such as Mars. Whilst some research has been carried out using 16s rRNA sequencing on forefield communities, metagenomics has yet to be applied to understanding succession. This study used metagenomes spanning the forefield of Midtre Lovénbreen glacier, Svalbard, to investigate microbial community development. A combination of metagenome DNA reads, assemblies and genome binning were used to gain

an insight into microbial taxonomy and function during soil development. A diverse range of microbes were recovered, including those with carbon, sulfur and nitrogen cycling metabolisms. Cyanobacteria were detected in early soils, attributed to their ability to fix carbon and nitrogen in oligotrophic conditions, alongside the production of protective EPS. The microbial community was shown to modify along the chronosequence, in line with the establishment of labile carbon and nitrogen pools. Overall, this study provided an insight into the diversity and metabolic potential of microbial communities during forefield succession and highlighted the potential of metagenomic analysis for those studying Arctic microbial ecology.

**Objective 4:** To investigate the composition and potential function of microbial communities sourced from benthic metagenomes in a Chilean fjord.

The fjord systems of Chilean Patagonia contain three UNSECO bio-reserves, support commercially important Salmon fisheries and host high rates of primary productivity. These fjords are particularly of interest due to the interaction of glacial meltwater from Patagonian ice fields with marine waters, harboring a range of physico-chemical conditions for microbial life. However, the microbiology of the fjord sediments, crucial for understanding the wider ecosystem functioning, has received limited research attention.

This study applied metagenomics to understand the taxonomic diversity of uncultured benthic sediment metagenomes from a Patagonian fjord. The uncovered taxonomic diversity was used to drive genome binning and functional analysis, providing insights into the novelty, metabolic potential and ecological diversity of these microbial communities. In particular, extremophiles associated with anoxia and oligotrophy were detected. Despite the harsh nature of the environment, organisms with carbon, nitrogen and sulfur metabolisms were detected, indicating a range of potential biogeochemical cycles. Additionally, the results highlight the novelty of the microbial community, with the potential for new genomes within the samples. The findings provide an insight into this unique environment, whilst highlighting areas where targeted single cell genome sequencing and culture-based studies may be beneficial.

## Chapter 2: Comparison of publicly available metagenome assemblers for soil bacterial communities

Maisie V. Nash<sup>1</sup>; Alexandre M. Anesio<sup>2</sup>; Gary Barker<sup>3</sup> and Patricia Sánchez-Baracaldo<sup>1</sup>

<sup>1</sup>School of Geographical Sciences, University of Bristol, BS8 1SS, UK

<sup>2</sup>Department of Environmental Science, Aarhus University, PO box 358, Denmark

<sup>3</sup>School of Life Sciences, University of Bristol, BS8 1TQ, UK

### Contributions and acknowledgements

All work presented in this chapter was carried out by MN. Co-authors AA, GB and PSB provided methodological advice and guidance on editing. This work was presented as part of a conference talk during the Polar and Alpine Microbiology Symposium (2017).

### Funding

This work was carried out as part of the NERC GW4+ Doctoral Training Partnership, awarded to MN. Funding support was provided through NERC grants NE/J02399X/1, awarded to AA, and a Royal Society Research Fellowship awarded to PSB.

### 2.1 Introduction

Natural soils incorporate complex and dynamic bacterial communities, which play important roles in local biogeochemical cycles, such as nutrient cycling and carbon fixation (Handelsman *et al.*, 1998; Van Der Heijden *et al.*, 2008). Investigating these communities from a genomic perspective aids understanding of their taxonomic composition and functional potential, and how this changes between locations or over environmental gradients (Cong *et al.*, 2015). However, due to the complex structure of soils, small spatial scale of biogeochemical processes and heterogeneity of bacterial life, it is difficult to extract and culture cells for analysis (Riesenfeld *et al.*, 2004; Teeling and Glockner, 2012). Whilst some studies have used 16s rRNA amplicon sequencing and microarrays to investigate soil bacterial communities, these provide limited information on novel uncultured organisms and the functional genes they contain (Fierer *et al.*, 2012; Cong *et al.*, 2015). A more comprehensive understanding of microbial communities may be obtained through metagenomics, as ideally the complete microbial DNA is sequenced (Daniel, 2005; Handelsman, 2004). This provides the user with information on the community composition and functional genes, therefore highlighting the potential contribution of the community to local biogeochemical processes (Tringe *et al.*, 2005; Tringe and Rubin, 2005).

There is currently a limited (but growing) uptake of metagenomic sequencing in the microbial ecology community, largely related to the considerable use and lower cost of 16s amplicon

sequencing (Risenfeld *et al.*, 2004). However, metagenomics has been explored more extensively in other fields, for example in studying the human gut microbiome (Guill *et al.*, 2006; Qin *et al.*, 2010; Qin *et al.*, 2012). The approach is gaining interest from the microbial ecology community, as more research has been published exemplifying the possibilities with metagenomic data from environmental samples (Piganeau and Moreau, 2007; Makelprang *et al.*, 2011; Pearce *et al.*, 2012; Eloë-Fadrosh *et al.*, 2016a). As there is a lack of guidance on metagenome data analysis, the majority of studies use short unassembled sequencing reads, in programs such as MG-RAST (Meyer *et al.*, 2008; Glass *et al.*, 2010). This may not be optimal, as the short DNA fragments may not provide good alignments to taxonomic and functional databases during metagenome annotation, limiting the potential of the data analysis (Van der Walt *et al.*, 2017). Consequently, assembling the DNA fragments into longer contigs, using a metagenome assembler, can improve the dataset annotation quality (Nagarjan and Pop, 2013). Assembled metagenomes have enabled scientists to explore community functional genes in greater detail, such as through comparing environments or even assembling complete genomes from metagenomic datasets (Sharon and Banfield, 2013; Eloë-Fadrosh *et al.*, 2016b). However, it must be acknowledged that assembly may not always be the optimal approach, for example when investigating highly diverse, novel and fragmented metagenomes, which can be harder to assemble.

A limited number of ecological studies have started assembling metagenome DNA, however these studies often select a single assembler without justification for that choice (Najaran and Pop, 2013). This is widely accepted, as there is limited formalised assembler comparisons or guidance specifically for the analysis of bacterial microbial communities sourced from soils. However, as assemblers are based on multiple algorithms, they will work differently depending on the sequencing platform, community structure and complexity of the input dataset (Zhang *et al.*, 2011). Testing several assemblers for the type of dataset in question, selecting the best performing algorithm, can enhance the quality of the metagenome assembly, and therefore annotation (Kurtz *et al.*, 2004; Peng *et al.*, 2012; Howe *et al.*, 2014). More accurate functional annotations will not only increase the reliability of the ecological conclusions drawn, but also enable more detailed analysis of the microbial community, such as extracting single draft genomes from community datasets (Eloë-Fadrosh *et al.*, 2016b). Several studies have attempted assembler comparison for animal, human and mixed environmental samples (Mavromatis *et al.*, 2007; Lin *et al.*, 2011; Zhang *et al.*, 2011; Peng *et al.*, 2012; Deng *et al.*, 2015; Sczyrba *et al.*, 2017; Vollmers *et al.*, 2017). However, many of these comparisons have been for single genome assemblies, not metagenomes, or focused on mixed environmental metagenomes from numerous habitats (Earl *et al.*, 2011; Salzberg *et al.*, 2012; van der Walt *et al.*, 2017; Vollmers *et al.*, 2017). The recently completed CAMI inter-comparison aimed to



evaluate software for metagenomic data analysis, including assemblers, based on simulated metagenomes (Sczyrba *et al.*, 2017). However, this study used newly sequenced genomes from a range of sources, spanning Bacteria, Archaea and Fungi (Sczyrba *et al.*, 2017). Thus far, there has been no assembler comparison focused on complex soil bacterial communities, comparing all assembly algorithms.

When conducting an assembler comparison, simulating artificial microbial communities to use as the test dataset(s) is often useful (Mavromatis *et al.*, 2007; Sczyrba *et al.*, 2017; van der Walt *et al.*, 2017). This allows the user to construct the input dataset, and therefore evaluate how well the assemblers do in recreating it (Mavromatis *et al.*, 2007). These simulated datasets can contain DNA sequences of choice from publicly available databases, alongside artificially introduced sequencing errors, to reflect typical metagenome data preparation (Mavromatis *et al.*, 2007). It is important to conduct assembler comparisons using multiple evaluation measures, including both contiguity and completeness (Arumugam *et al.*, 2010; Vollmers *et al.*, 2017). This will highlight any assemblers that are producing long, but erroneous contigs (Kumar and Blaxter, 2010; Bräutigam *et al.*, 2011).

Here, five publicly available metagenome assemblers, spanning the three main assembly algorithms, were compared for four artificially curated soil metagenomes of different complexity. Assembler evaluation was carried out using both the contiguity and completeness of metagenome assemblies. This study aims provide a more standardized method for the assembly of complex soil metagenome data, aiding assembler choice within the soil bacterial ecology community.

## **2.2 Methodology**

For this analysis, five metagenome assemblers were selected for comparison, spanning the three key assembly algorithms (Arumugam *et al.*, 2010; Zhang *et al.*, 2011). A summary of these assemblers can be identified in Table 2.1.

**Table 2.1:** The five assemblers used in the evaluation, detailing the assembler version and algorithm, alongside assembly specific details.

Assembler	Algorithm	Assembly details
MetaSPAdes 3.7.0 (Nurk <i>et al.</i> , 2013)	de Bruijn graph	Assembly run in paired end mode, with error correction switched off (--only-assembler flag) as base quality values are required for this, however are not supplied in the MetaSim datasets. Read coverage cut-off was not set as this cannot be used in paired end mode.  Key parameter to modify was the kmer length, which ranged between 21 and 71.
SSAKE 3.8.4 (Warren <i>et al.</i> , 2007)	Greedy	Assembly run in paired end mode.  Key parameters to modify were the minimum contig coverage depth, the minimum number of overlapping bases required for a consensus sequence join, and to trim bases when all other extension prospects have been trialled.
ABYSS 1.9.0 (Simpson <i>et al.</i> , 2009)	de Bruijn graph	Key parameters to modify were the kmer length, the minimum contig coverage and to pop assembly bubbles below a threshold value.
MIRA 4.0 (Chavreux, B., 2014)	Overlap layout consensus	Assembly run without base quality values, using the --no_qualities flag, as these are not supplied by Metasim. Assemblies were run using the flag for Illumina input data (SOLEXA_SETTINGS).  Key parameters to modify were the minimum read length used in the assembly and the minimum contig length.
CLC 4.4.1 (Qiagen Bioinformatics, 2016)	de Bruijn graph	Assembly run in paired end mode without scaffolding.  The key parameters to modify were the minimum output contig length, the minimum word size for the de Bruijn graph, and the maximum assembly bubble size allowed.

To test these assemblers, an artificial metagenome was created in MetaSim, incorporating 123 complete soil bacterial genomes downloaded from NCBI GenBank, identified in Table 2.2 (Richter *et al.*, 2008). Simulated Illumina sequencing errors were introduced to the metagenome to replicate typical sequencing error profiles, such as declining sequencing quality with run time (Richter *et al.*, 2008). Four variants of the artificial metagenome were created (A-D), with increasing levels of complexity, each containing 10,000,000 reads, 80 base pairs (bp) in length (Table 2.3). The metagenome size was selected as a trade-off between accurately representing soil metagenome sizes, whilst minimising computational requirements. Metagenome complexity modifications were carried out by amending the abundance profile of bacterial species in the MetaSim simulations, in line with the methodology of Mavromatis *et al.*, (2007). The low complexity community (Metagenome A)

contained a small number of high abundance organisms, with the remainder of the bacterial population at a lower abundance. This theoretically would allow a good assembly of the high abundance bacteria, due to an increase in read coverage for those organisms (Mavromatis *et al.*, 2007). In contrast, the high complexity community (Metagenome D) contains organisms at a range of abundance levels, with varying levels of coverage. As the population abundance is more variable, the assemblers may create a smaller and more erroneous output, for example by increasing the number of chimeric contigs, whereby reads from different organisms are merged (Mende *et al.*, 2012). This is because there would be reduced sequencing support for each read from low abundance organisms, thereby causing difficulties in assembling regions with long repeats or INDELS. Furthermore, a default dataset with even coverage (or abundance) across all bacteria (Metagenome A) was also included. This is to remove the effect of high abundance organisms, which, due to the associated increase in read coverage, tend to pull the average assembly quality up (Mavromatis *et al.*, 2007). In total 100 assemblies were carried out, with the full list available in Appendix 1 Table 1.

**Table 2.2:** Complete bacterial genome sequences used for artificial metagenomes created in MetaSim, sourced from NCBI GenBank. GenBank ID numbers are given, alongside genome description.

No.	NCBI accession number	NCBI sequence name
1	CP007128.1	<i>Gemmatimonadetes</i> bacterium KBS708, complete genome
2	NC_000964.3	<i>Bacillus subtilis</i> subsp. <i>subtilis</i> str. 168 chromosome, complete genome
3	NC_002516.2	<i>Pseudomonas aeruginosa</i> PAO1 chromosome, complete genome
4	NC_002942.5	<i>Legionella pneumophila</i> subsp. <i>pneumophila</i> str. Philadelphia 1
5	NC_002947.3	<i>Pseudomonas putida</i> KT2440 chromosome, complete genome
6	NC_003030.1	<i>Clostridium acetobutylicum</i> ATCC 824 chromosome, complete genome
7	NC_003155.4	<i>Streptomyces avermitilis</i> MA-4680 = NBRC 14893, complete genome
8	NC_003210.1	<i>Listeria monocytogenes</i> EGD-e chromosome, complete genome
9	NC_003212.1	<i>Listeria innocua</i> Clip11262 complete genome
10	NC_003366.1	<i>Clostridium perfringens</i> str. 13 DNA, complete genome
11	NC_003450.3	<i>Corynebacterium glutamicum</i> ATCC 13032 chromosome, complete genome
12	NC_003888.3	<i>Streptomyces coelicolor</i> A3(2) chromosome, complete genome
13	NC_003902.1	<i>Xanthomonas campestris</i> pv. <i>campestris</i> str. ATCC 33913 chromosome
14	NC_004668.1	<i>Enterococcus faecalis</i> V583 chromosome, complete genome
15	NC_004722.1	<i>Bacillus cereus</i> ATCC 14579 chromosome, complete genome
16	NC_005085.1	<i>Chromobacterium violaceum</i> ATCC 12472, complete genome
17	NC_005296.1	<i>Rhodopseudomonas palustris</i> CGA009 complete genome
18	NC_005957.1	<i>Bacillus thuringiensis</i> serovar <i>konkukian</i> str. 97-27 chromosome
19	NC_006177.1	<i>Symbiobacterium thermophilum</i> IAM 14863 DNA, complete genome
20	NC_006270.3	<i>Bacillus licheniformis</i> ATCC 14580, complete genome"
21	NC_006361.1	<i>Nocardia farcinica</i> IFM 10152 DNA, complete genome
22	NC_006582.1	<i>Bacillus clausii</i> KSM-K16 DNA, complete genome
23	NC_006834.1	<i>Xanthomonas oryzae</i> pv. <i>oryzae</i> KACC 10331, complete genome
24	NC_007005.1	<i>Pseudomonas syringae</i> pv. <i>syringae</i> B728a chromosome, complete genome
25	NC_007404.1	<i>Thiobacillus denitrificans</i> ATCC 25259, complete genome
26	NC_007406.1	<i>Nitrobacter winogradskyi</i> Nb-255, complete genome
27	NC_007761.1	<i>Rhizobium etli</i> CFN 42, complete genome
28	NC_008009.1	<i>Candidatus Koribacter versatilis</i> Ellin345, complete genome
29	NC_008255.1	<i>Cytophaga hutchinsonii</i> ATCC 33406, complete genome
30	NC_008536.1	<i>Solibacter usitatus</i> Ellin6076, complete genome
31	NC_008555.1	<i>Listeria welshimeri</i> serovar 6b str. SLCC5334 complete genome
32	NC_008593.1	<i>Clostridium novyi</i> NT, complete genome
33	NC_008711.1	<i>Arthrobacter aurescens</i> TC1, complete genome

34	NC_008726.1	<i>Mycobacterium vanbaalenii</i> PYR-1, complete genome
35	NC_009328.1	<i>Geobacillus thermodenitrificans</i> NG80-2, complete genome
36	NC_009434.1	<i>Pseudomonas stutzeri</i> A1501, complete genome
37	NC_009439.1	<i>Pseudomonas mendocina</i> ymp, complete genome
38	NC_009441.1	<i>Flavobacterium johnsoniae</i> UW101, complete genome
39	NC_009515.1	<i>Methanobrevibacter smithii</i> ATCC 35061, complete genome
40	NC_009636.1	<i>Sinorhizobium medicae</i> WSM419 chromosome, complete genome
41	NC_009674.1	<i>Bacillus cytotoxicus</i> NVH 391-98, complete genome
42	NC_009720.1	<i>Xanthobacter autotrophicus</i> Py2, complete genome
43	NC_009792.1	<i>Citrobacter koseri</i> ATCC BAA-895, complete genome
44	NC_010001.1	<i>Clostridium phytofermentans</i> ISDg, complete genome
45	NC_010002.1	<i>Delftia acidovorans</i> SPH-1, complete genome
46	NC_010337.2	<i>Heliobacterium modesticaldum</i> lce1, complete genome
47	NC_010571.1	<i>Opitutus terrae</i> PB90-1, complete genome
48	NC_010572.1	<i>Streptomyces griseus</i> subsp. <i>griseus</i> NBRC 13350, complete genome
49	NC_010617.1	<i>Kocuria rhizophila</i> DC2201 DNA, complete genome
50	NC_010655.1	<i>Akkermansia muciniphila</i> ATCC BAA-835, complete genome
51	NC_010725.1	<i>Methylobacterium populi</i> BJ001, complete genome
52	NC_010995.1	<i>Cellvibrio japonicus</i> Ueda107, complete genome
53	NC_011000.1	<i>Burkholderia cenocepacia</i> J2315 chromosome 1, complete genome
54	NC_011001.1	<i>Burkholderia cenocepacia</i> J2315 chromosome 2, complete genome
55	NC_011002.1	<i>Burkholderia cenocepacia</i> J2315 chromosome 3, complete genome
56	NC_011666.1	<i>Methylocella silvestris</i> BL2, complete genome
57	NC_011886.1	<i>Arthrobacter chlorophenolicus</i> A6, complete genome
58	NC_012483.1	<i>Acidobacterium capsulatum</i> ATCC 51196, complete genome
59	NC_012490.1	<i>Rhodococcus erythropolis</i> PR4 DNA, complete genome
60	NC_012560.1	<i>Azotobacter vinelandii</i> DJ, complete genome
61	NC_012660.1	<i>Pseudomonas fluorescens</i> SBW25 complete genome
62	NC_012669.1	<i>Beutenbergia cavernae</i> DSM 12333, complete genome
63	NC_012778.1	<i>Eubacterium eligens</i> ATCC 27750, complete genome
65	NC_012781.1	<i>Eubacterium rectale</i> ATCC 33656, complete genome
65	NC_012881.1	<i>Desulfovibrio salexigens</i> DSM 2638, complete genome
66	NC_012969.1	<i>Methylovorus glucosetrophus</i> SIP3-4, complete genome
67	NC_013061.1	<i>Pedobacter heparinus</i> DSM 2366, complete genome
68	NC_013131.1	<i>Catenulispora acidiphila</i> DSM 44928, complete genome
69	NC_013132.1	<i>Chitinophaga pinensis</i> DSM 2588, complete genome
70	NC_013595.1	<i>Streptosporangium roseum</i> DSM 43021, complete genome
71	NC_013739.1	<i>Conexibacter woesei</i> DSM 14684, complete genome
72	NC_013743.1	<i>Haloterrigena turkmenica</i> DSM 5511, complete genome
73	NC_013757.1	<i>Geodermatophilus obscurus</i> DSM 43160, complete genome
74	NC_013861.1	<i>Legionella longbeachae</i> NSW150, complete genome
75	NC_013892.1	<i>Xenorhabdus bovienii</i> SS-2004 chromosome, complete genome
76	NC_013929.1	<i>Streptomyces scabiei</i> 87.22 complete genome
77	NC_013947.1	<i>Stackebrandtia nassauensis</i> DSM 44728, complete genome
78	NC_014103.1	<i>Bacillus megaterium</i> DSM319, complete genome
79	NC_014158.1	<i>Tsukamurella paurometabola</i> DSM 20162, complete genome
80	NC_014170.1	<i>Xenorhabdus nematophila</i> ATCC 19061 plasmid XNC1_p, complete genome
81	NC_014228.1	<i>Xenorhabdus nematophila</i> ATCC 19061 chromosome, complete genome
82	NC_014259.1	<i>Acinetobacter oleivorans</i> DR1, complete genome
83	NC_014323.1	<i>Herbaspirillum seropedicae</i> SmR1, complete genome
84	NC_014391.1	<i>Micromonospora aurantiaca</i> ATCC 27029, complete genome
85	NC_014551.1	<i>Bacillus amyloliquefaciens</i> DSM7 complete genome
86	NC_014622.2	<i>Paenibacillus polymyxa</i> SC2, complete genome
87	NC_014623.1	<i>Stigmatella aurantiaca</i> DW4/3-1, complete genome
88	NC_014734.1	<i>Paludibacter propionicipigenes</i> WB4, complete genome
89	NC_014814.1	<i>Mycobacterium gilvum</i> Spyr1, complete genome
90	NC_014958.1	<i>Deinococcus maricopensis</i> DSM 21211, complete genome
91	NC_015177.1	<i>Pedobacter saltans</i> DSM 12145, complete genome
92	NC_015514.1	<i>Cellulomonas fimi</i> ATCC 484, complete genome
93	NC_015677.1	<i>Ramlibacter tataouinensis</i> TTB310, complete genome
94	NC_016109.1	<i>Kitasatospora setae</i> KM-6054, complete genome
95	NC_016629.1	<i>Desulfovibrio africanus</i> str. Walvis Bay, complete genome
95	NC_016803.1	<i>Desulfovibrio desulfuricans</i> ND132, complete genome
97	NC_016845.1	<i>Klebsiella pneumoniae</i> subsp. <i>pneumoniae</i> HS11286 chromosome
98	NC_017384.1	<i>Ketogulonigenium vulgare</i> WSH-001 chromosome, complete genome
99	NC_017770.1	<i>Solitalea canadensis</i> DSM 3403, complete genome
100	NC_017960.1	<i>Enterococcus faecium</i> DO chromosome, complete genome

101	NC_018750.1	<i>Streptomyces venezuelae</i> ATCC 10712 complete genome
102	NC_020800.1	<i>Xanthomonas axonopodis</i> Xac29-1, complete genome
103	NC_020815.1	<i>Xanthomonas citri</i> subsp. citri Aw12879, complete genome
104	NC_020816.1	<i>Xanthomonas citri</i> subsp. citri Aw12879 plasmid pXcaw19, complete genome
105	NC_020817.1	<i>Xanthomonas citri</i> subsp. citri Aw12879 plasmid pXcaw58, complete genome
106	NC_020990.1	<i>Streptomyces albus</i> J1074, complete genome
107	NC_020995.1	<i>Enterococcus casseliflavus</i> EC20, complete genome
108	NZ_AP014683.1	<i>Burkholderiales bacterium</i> GJ-E10 DNA, complete genome
109	NZ_CP002190.1	<i>Bdellovibrio bacteriovorus</i> W, complete genome
110	NZ_CP007215.1	<i>Enterobacter sacchari</i> SP1, complete genome
111	NZ_CP007557.1	<i>Citrobacter freundii</i> CFNIH1, complete genome
112	NZ_CP009124.1	<i>Streptomyces lividans</i> TK24, complete genome
113	NZ_CP009576.1	<i>Listeria ivanovii</i> subsp. londoniensis strain WSLC 30151, complete genome
114	NZ_CP009962.1	<i>Collimonas arenae</i> strain Cal35, complete genome
115	NZ_CP009963.1	<i>Collimonas arenae</i> strain Cal35 plasmid, complete sequence
116	NZ_CP010028.1	<i>Deinococcus swuensis</i> strain DY59, complete genome"
117	NZ_CP010946.1	<i>Mycobacterium chelonae</i> genome
118	NZ_CP011253.2	<i>Pandoraea oxalativorans</i> strain DSM 23570, complete genome
119	NZ_CP011451.1	<i>Nitrosomonas communis</i> strain Nm2, complete genome
120	NZ_CP012329.1	<i>Bacillus pumilus</i> strain NJ-M2, complete genome
121	NZ_CP012382.1	<i>Streptomyces ambofaciens</i> ATCC 23877, complete genome
122	NZ_CP013106.1	<i>Halomonas huangheensis</i> strain BJGMM-B45, complete genome
123	NZ_HG916826.1	<i>Pseudomonas pseudoalcaligenes</i> CECT 5344 complete genome

**Table 2.3:** Four artificial bacterial metagenomes created in MetaSim. The community complexity description is given, alongside the number of organisms at each abundance level. The relative abundance values used in MetaSim are given in parentheses.

Metagenome	Community complexity	Metagenome composition
A	Default test community (all organisms at same abundance)	123 even abundance (100)
B	Low complexity	4 high abundance (200) 119 standard abundance (100)
C	Medium complexity	20 high abundance (200) 93 standard abundance (100) 10 low abundance (50)
D	High complexity	20 high abundance (200) 20 medium-high abundance (150) 63 standard abundance (100) 20 low abundance (50)

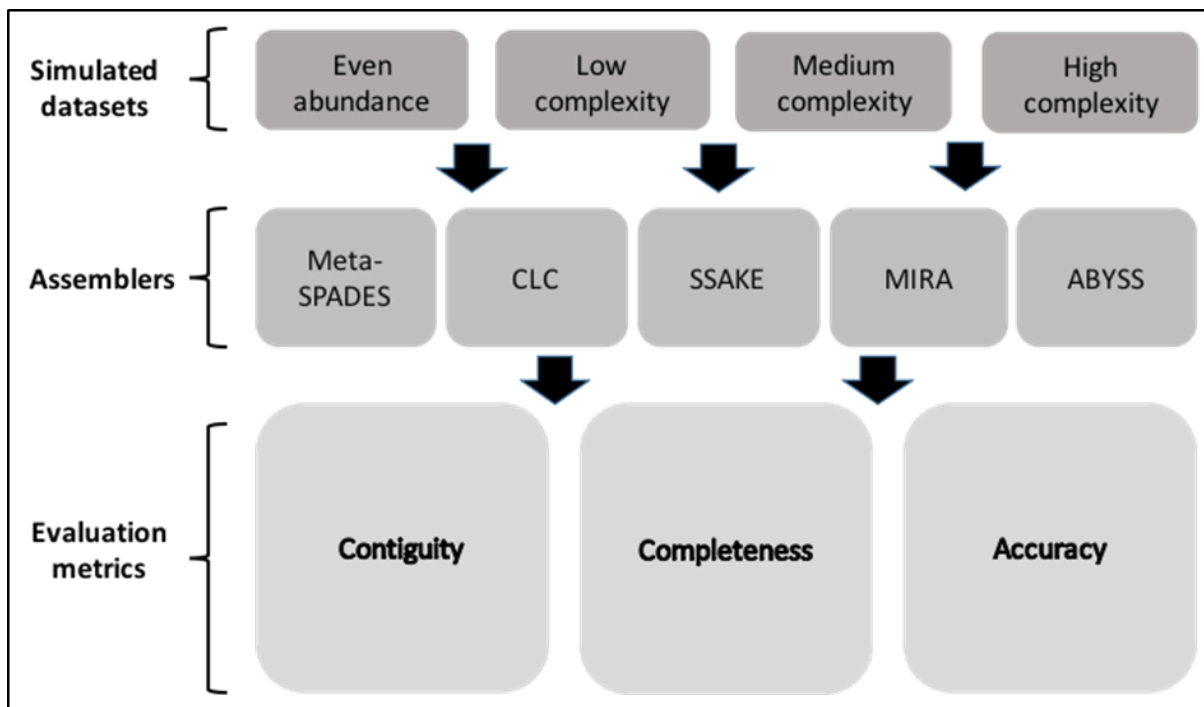
Each of the assemblers were subsequently run on the four metagenome variants (A-D) in turn. However, whilst many studies currently run assemblies using default parameter settings, it must be acknowledged that modifying the parameter values to fit the input metagenome size, structure and complexity can produce a better assembly (Howe *et al.*, 2014). Consequently, whilst a full parameter space optimisation was not the aim of this analysis, for each assembler, five parameter sets were tested on each metagenome, to obtain an improved result. The

importance of this is to generate longer and more accurate contigs, which may provide enhancements for subsequent functional and phylogenetic annotations (Nielsen *et al.*, 2014). A schematic for this methodology can be identified in Figure 2.1, and the full list of parameter sets used is available in Appendix 1 Table 2. Key parameters to modify were the k-mer length (or word size) used to build assemblies, alongside the minimum read length and coverage. Parameter values were selected systematically, to test periodic intervals across the parameter space.

The final assemblies were evaluated based on both contiguity and completeness metrics. This allows the comparison to incorporate both the assembly length and coverage of the input dataset, to avoid large erroneous contigs masking the quality of the assembly (Earl *et al.*, 2011; Finotello *et al.*, 2011). The assembly contiguity was evaluated in Quast, using the measures of maximum contig length, contig N50 and number of contigs over 1000bp (Gurevich *et al.*, 2013). Assembly size was also evaluated, based on number of contigs in the assembly, and overall assembly size (bp). To analyse assembly completeness, the input metagenome reads were mapped to the assembled datasets using BWA (Li and Durbin, 2009). This identifies the amount of the input metagenome that has been used in the assembly, shown through the percentage coverage of the input dataset. Assembly chimerism was explored by searching for chimeric contigs using VSEARCH, enabling reference-based chimera detection (Rognes *et al.*, 2016). Using VSEARCH, the output metagenome assemblies were searched against the input artificial metagenomes using a global alignment to detect for chimeric sequences.

Additionally, the accuracy of the best metagenome assemblies in terms of the taxonomic distribution of organisms was evaluated using MetaPhlan2 (Segata *et al.*, 2012). This method uses clade-specific markers to profile the assembled metagenome taxonomy and provide an output taxonomic distribution of the dataset (Segata *et al.*, 2012). This allows an insight into the presence of chimeric contigs, which may create erroneous organisms which were not present in the input dataset, due to joining sequences from unrelated bacteria (Lai *et al.*, 2012). This may also cause assemblers to eliminate highly abundant organisms, particularly when containing DNA sequences with numerous repeat regions or INDELS, as the assemblers find these difficult to resolve. This can be identified by viewing the top most abundant genus' between assemblies of the same metagenome, in comparison with the input dataset. Overall, the best performing assembly will most closely match the taxonomic profile of the input metagenome. Phylogenetic heat maps were created displaying the 25 most abundant clades in each metagenome in a logarithmic scale. Clustering was performed by average linkage and Euclidean distances for both clades and samples (Segata *et al.*, 2012). Finally, the best

performing parameter set for each assembler (for each metagenome) was selected based on the above metrics.



**Figure 2.1:** Overview schematic of the methodology used in assembler comparison.

## 2.3 Results

### 2.3.1 Assembly coverage and contiguity

The best overall assemblies from the five tested assemblers are summarised in Table 2.4, with the full results available in Appendix 1 Table 3. As the study is primarily focused on assembler comparison (rather than parameter optimisation), the parameter set shown in Table 2.4 is the best performing out of the five tested, and not a product of testing the whole parameter space. For each assembly, a range of contiguity and completeness metrics are presented, including assembly N50, commonly used to describe contiguity. However, as the N50 is not a standardised metric, and relies on the dataset size, it should not be used to directly compare assemblies of different magnitudes.

**Table 2.4:** Summary table of metagenome assemblies, split by metagenome complexity (A-D). The single overall best performing parameter set for each assembler is shown, with results for all parameter sets available in Appendix 1. The parameter set is given in parentheses. Assembly size (number of contigs) and contiguity (N50, maximum contig length, contigs over 1000bp) statistics are given. The assembly completeness is identified by % coverage, which is the percentage of raw input metagenome reads that could be mapped to the final assembly.

Assembler	CLC (2)	ABYSS (2)	MIRA (1)	MetaSPAdes (1)	SSAKE (4)
Assembly	A	A	A	A	A
Number of contigs	1334018	1227327	186625	666045	351881
Contigs over 1000bp	1369	143	373	939	21
Maximum contig length	5060	4641	5387	5299	1725
N50	232	109	203	597	122
% coverage of BWA mapping	59.35	22.13	10.30	42.08	8.89

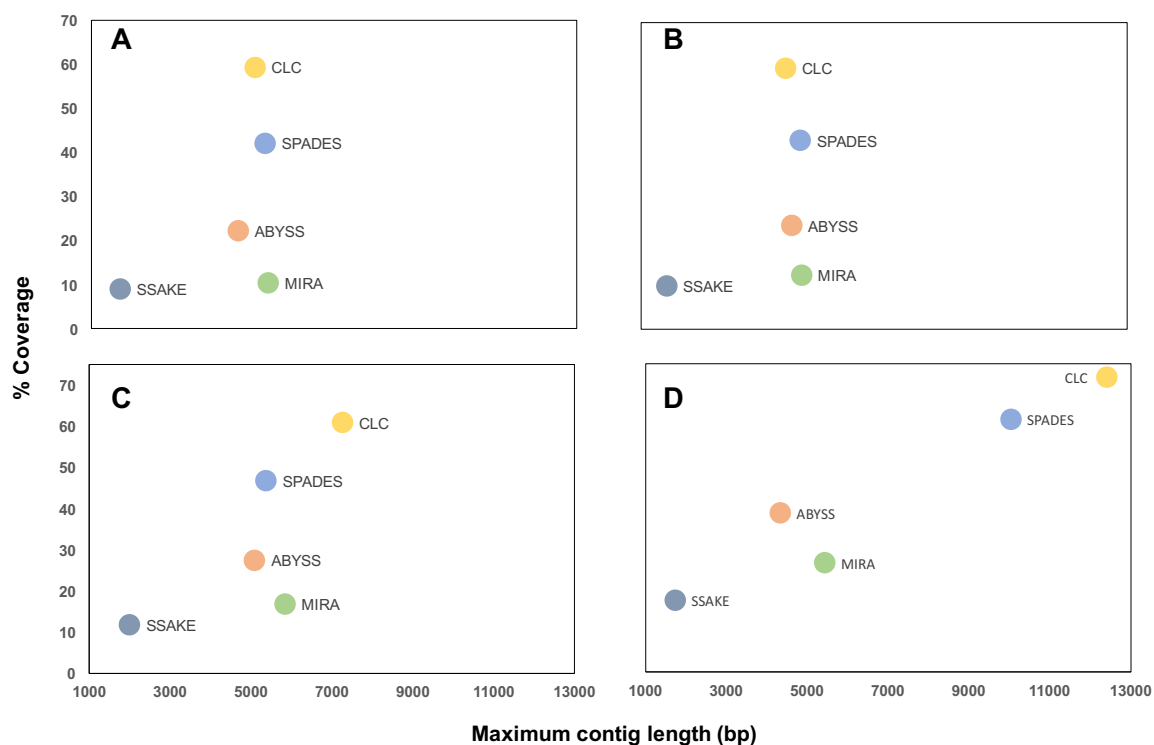
Assembler	CLC (2)	ABYSS (2)	MIRA (1)	MetaSPAdes (1)	SSAKE (4)
Assembly	B	B	B	B	B
Number of contigs	1289769	1204906	205976	645872	366959
Contigs over 1000bp	3234	146	389	2285	29
Maximum contig length	4585	4730	4773	4940	1643
N50	234	112	210	624	123
% coverage of BWA mapping	59.56	23.56	12.24	43.03	9.75

Assembler	CLC (2)	ABYSS (2)	MIRA (1)	MetaSPAdes (1)	SSAKE (4)
Assembly	C	C	C	C	C
Number of contigs	1179804	1189639	272144	613234	417318
Contigs over 1000bp	6144	147	418	4622	28
Maximum contig length	7265	5090	5552	5386	1999
N50	243	121	217	662	124
% coverage of BWA mapping	60.80	27.47	16.81	46.69	11.82

Assembler	CLC (2)	ABYSS (2)	MIRA (1)	MetaSPAdes (1)	SSAKE (4)
Assembly	D	D	D	D	D
Number of contigs	1038353	1253133	400485	613787	536263
Contigs over 1000bp	10876	189	456	8162	32
Maximum contig length	12412	4345	5931	10045	1743
N50	281	138	218	682	126
% coverage of BWA mapping	67.00	36.04	24.67	57.31	16.13

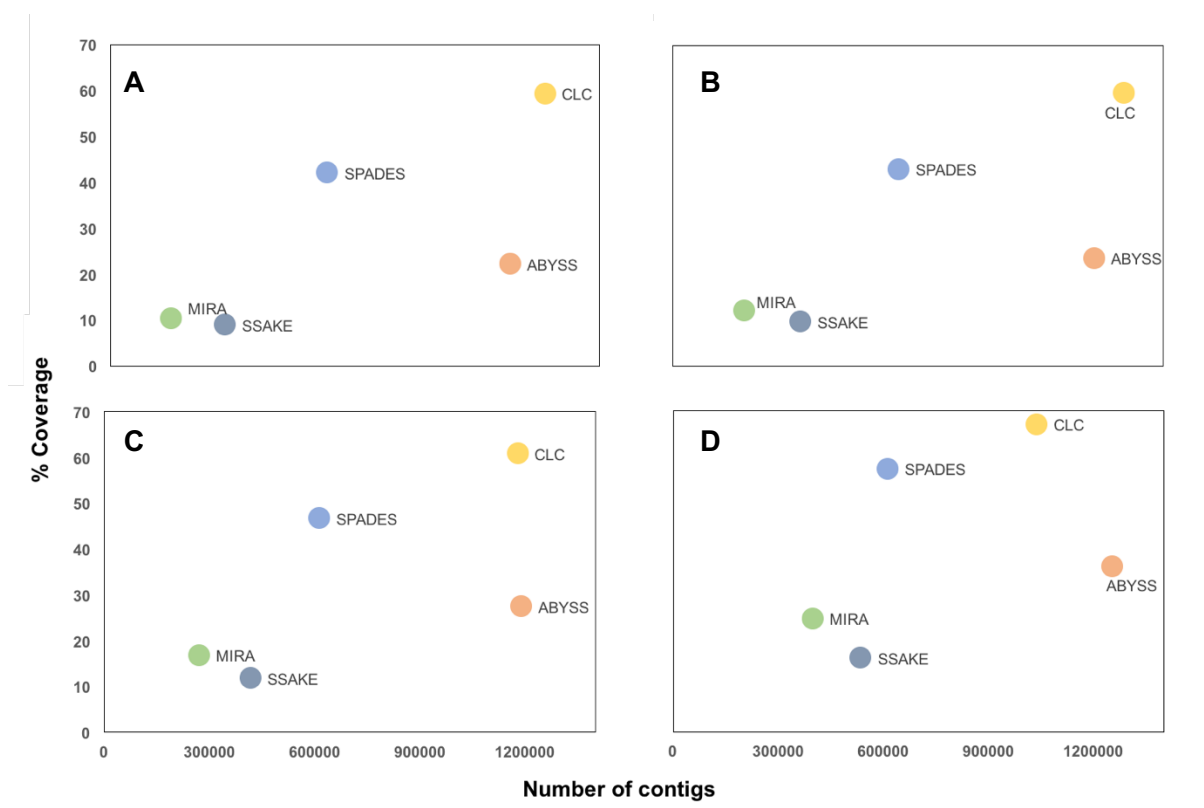


As the best performing assembly will have both the longest contigs, and include the greatest portion of the input dataset, it is important to evaluate based on both contiguity and completeness, as opposed to a single metric. Figure 2.2 exemplifies the percentage coverage against a commonly used contiguity metric, maximum contig length. The results are provided for each metagenome (A-D), with increasing complexity (Table 2.4; Figure 2.2). For all metagenomes, the de Bruijn based assemblers, CLC, metaSPAdes and ABYSS output the assemblies with the highest coverage and maximum contig length (Table 2.4; Figure 2.2). For the best performing assembler, CLC, the coverage and contiguity scores range between 59 to 67% and 4585 to 12412 bp, between metagenomes, respectively (Table 2.4; Figure 2.2). In contrast, the two worst scoring assemblers are MIRA and SSAKE, based on OLC and Greedy assembly algorithms (Table 2.4; Figure 2.2). The lowest scoring assembler for all metagenomes, SSAKE, output coverage and contiguity scores ranging between 8 to 16% and 1643 to 1999bp, respectively.



**Figure 2.2:** Maximum contig length (bp) and percentage coverage for assembled artificial metagenomes (A-D). For each assembler, the results for the best performing parameter set from the five tested are shown. Metagenome A: even organism abundance; Metagenome B: low organism complexity; Metagenome C: medium organism complexity; Metagenome D: high organism complexity.

A similar pattern to that observed in contiguity is also identified when analysing assembly size (as number of contigs), with assembly coverage (Figure 2.3). The largest assemblies with the most coverage are produced by de Bruijn based assemblers CLC, metaSPAdes and ABYSS (Figure 2.3). Again, CLC is the best performing assembler across all metagenomes tested, with assembly sizes ranging between 1038353 to 1334018 contigs in metagenomes D and A respectively (Table 2.4). These assemblies are in contrast to those obtained from SSAKE, which range between 351881 and 536263 contigs for metagenomes A and D. Results from a one-way analysis of variance (ANOVA) on the best performing parameter sets, identifies significant differences between assemblers for each maximum contig length, number of contigs and percentage coverage ( $P = <0.05$ , Appendix 1 Table 4). This identifies that significant differences in assembly coverage, contiguity and size can be obtained by using different assemblers.



**Figure 2.3:** Number of contigs and percentage coverage for assembled artificial metagenomes (A-D). For each assembler, the results for the best performing parameter set from the five tested are shown.

**Table 2.5:** Number of contigs, maximum contig length and percentage BWA coverage for each assembler. Values are reported as averages across all assemblies carried out for each assembler, over the four metagenomes (A-D), and the five parameter sets tested for each metagenome assembly.

Assembler	Number of contigs	Maximum contig length	% coverage
CLC	455,954	5,728	30
metaSPAdes	188,893	3,699	16
ABYSS	246,006	2,206	8
MIRA	266,292	3,967	13
SSAKE	241,520	1,363	6

On average, de Bruijn based assembler ABYSS produces a maximum contig length of 2,206, with a mean maximum contig number of 246,006 (Table 2.5). This is comparable to both metaSPAdes and CLC, which obtained a maximum number of contigs, on average of 188,893 and 455,954, respectively (Table 2.5). However, ABYSS obtains lower assembly coverage values than the other de Bruijn graph assemblers, CLC and metaSPAdes (Table 2.5).

When looking at the best performing parameter sets (Table 2.4) and overall assembly averages (Table 2.5) SSAKE and MIRA, based on simple Greedy and OLC algorithms, performed worst in terms of assembly coverage (Figures 1-2). SSAKE produced the least complete metagenome assemblies, with coverage values at 6% of the input data on average, and MIRA outputting a slightly higher average of 13% (Table 2.5). SSAKE assembler also produced the lowest number of contigs, and contigs over 1000bp, with maximum values of 536,263 and 32 contigs, for assembly D (Table 2.4). This compares to the best performing CLC assembler, which output at total of 1,038,353 contigs for assembly D, with 10,876 contigs over 1000bp (Table 2.4). This represents a percentage difference of 64% and 199% for total number of contigs and contigs over 1000bp between CLC and SSAKE for assembly D (high complexity) (Table 2.4). This pattern is reflected in the lower complexity datasets, with percentage differences of 111% and 196% for these variables between CLC and SSAKE for assembly B (low complexity) (Table 2.4). This identifies SSAKE as producing less complete assemblies, with smaller more fragmented contigs, than the de Bruijn based assembler CLC.

For OLC based assembler, MIRA, the number of contigs output is comparable to de Bruijn based assemblers ABYSS and metaSPAdes on average (Table 2.5). Average maximum contig numbers are 226,292 for MIRA, compared to 188,893 and 246,006 contigs for metaSPAdes and ABYSS respectively, however, these are still lower than CLC averages at 455,954 contigs (Table 2.5). When looking at contiguity statistics, maximum contig length is

again comparable (if not higher than) metaSPAdes and ABYSS for both the best performing parameter sets (Table 2.4, Figures 1-2) and the assemblies on average (Table 2.5). For example, for high complexity metagenome (D) MIRA has a maximum contig length of 5,931bp compared to 4,345bp, 10,045bp and 12,412bp for ABYSS, metaSPAdes and CLC respectively (Table 2.4). This indicates that both assembly size and contig length are comparable to the more complex de Bruijn based assemblers, whilst coverage is significantly lower (Table 2.4). This indicates that metaSPAdes and CLC outperform ABYSS in terms of assembly quality, as they are able to incorporate a greater amount of the input data in the resulting assemblies, thereby increasing assembly coverage.

### 2.3.2 Modifying community complexity

For each assembler, it is useful to compare the results across the different artificial metagenomes tested, as this will highlight how metagenome complexity influences the assembly outcome. Interestingly, in terms of the assembly coverage, N50 and number of contigs over 1000bp, overall assembler performance increases with greater metagenome complexity (Table 2.4; Table 2.5). For example, on average, assembly coverage (for best performing assemblers) rises from 36% to 50% between assembly A and assembly D (Table 2.4). This trend can be seen across all assemblers tested, with number of contigs over 1000bp increasing between 13% (CLC) and 770% (metaSPAdes), with an average increase of 2,473 (56%) in maximum contig length between metagenome A and D (Table 2.4). This indicates that with an increase in metagenome complexity, both the assembly contiguity and coverage improve, creating longer contiguous sequences, compared to shorter, fragmented sequences. However, whilst there are improvements in the assembly coverage and contiguity, the overall assembly size does not show a significant difference between low and high complexity metagenomes. On average, the assembly size only increases by 15,225 contigs (2%) between metagenomes A and D (Table 2.4). For de Bruijn graph assemblers, CLC and metaSPAdes, the number of contigs decreased by 22% and 8%, respectively (Table 2.4). Therefore, whilst contig length and coverage increases with greater complexity, the number of contigs assembled does not show a substantial improvement.

To test the significance of the metagenome complexity on the output assembly quality. A one-way ANOVA was carried out for each assembler independently. The contiguity, complexity and assembly size scores outlined in Table 2.4 were compared for each assembler across the four metagenomes tested. Overall, no significant difference could be observed in the assembly outcomes, when analysing metagenomes of different complexity (Appendix 1 Table 5). Consequently, whilst the metagenome complexity may play a role in modifying contiguity and

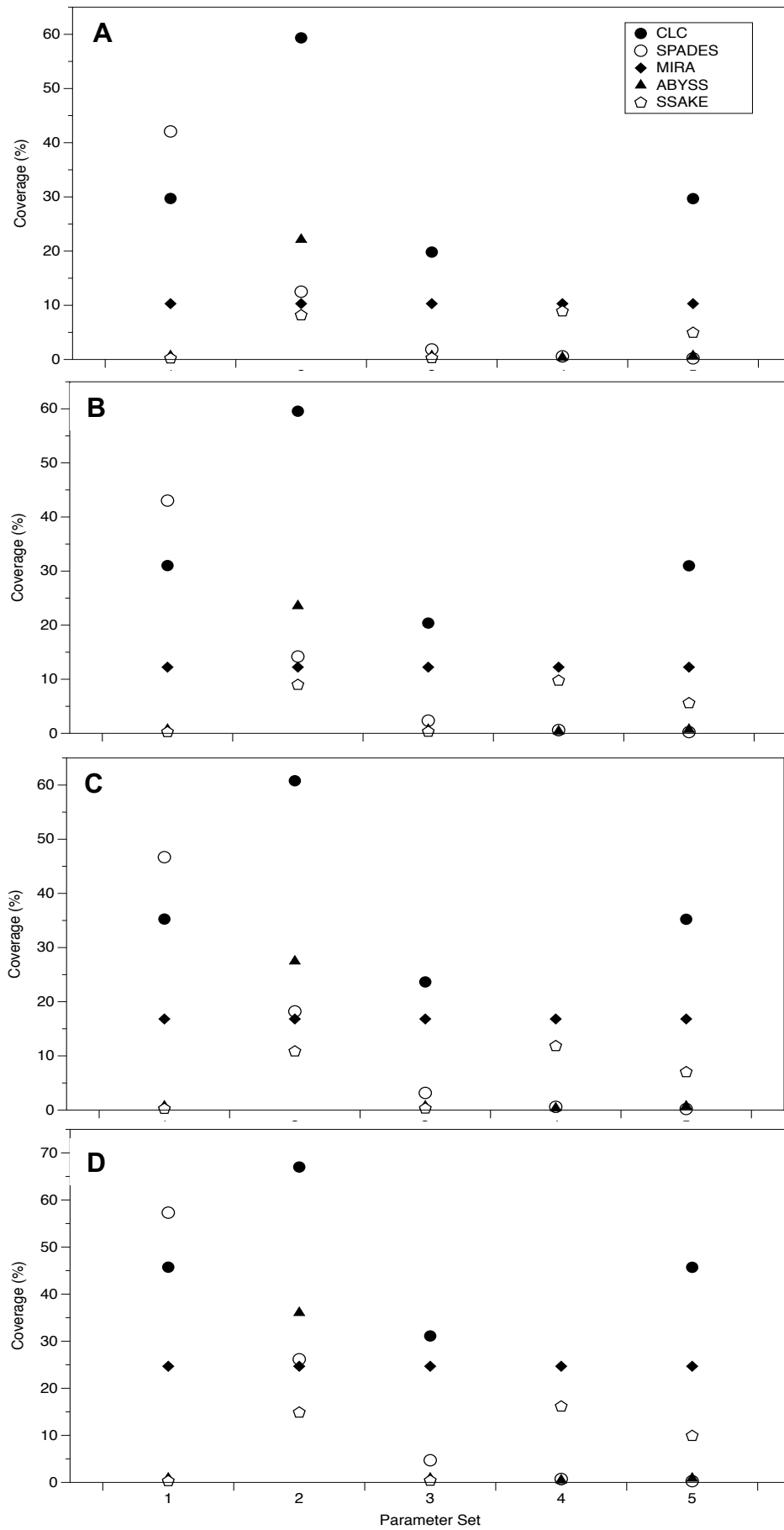
completeness of assemblies, overall this change was not substantial. Therefore, it can be argued that the optimal assembler selected for the data type (e.g. soil or sediment microbial communities) could be applied to numerous datasets of a similar type, with different community compositions.

### 2.3.3 Assembler parameterisation

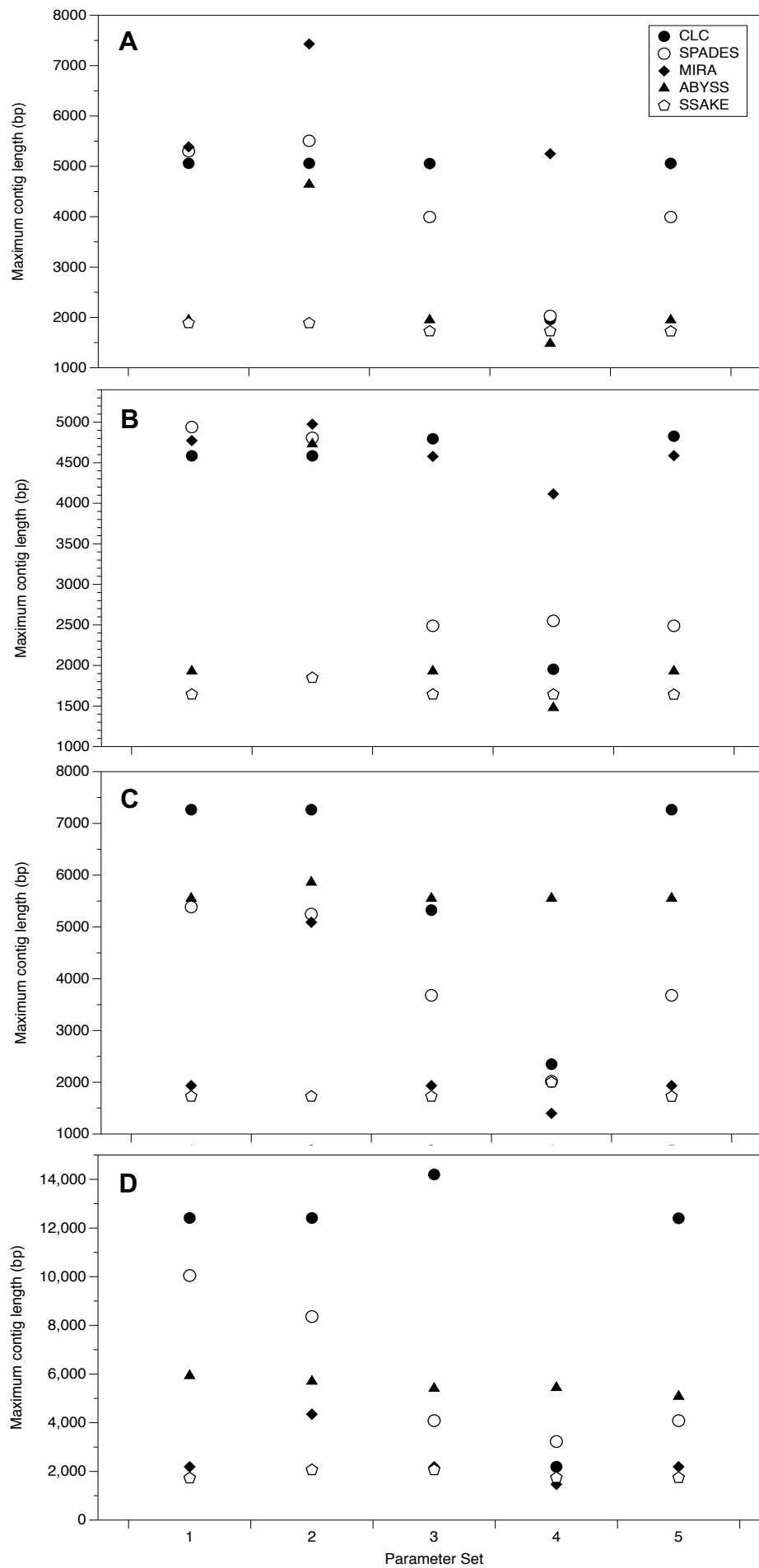
Whilst it may be possible to identify a suitable generic assembler for soil microbial communities, each assembler has a range of parameter values that can be modified (Appendix 1 Table 2). Whilst assembler optimisation is not the fundamental aim of this study, it is important to acknowledge the influence these have on assembly outcomes. Hence, for each assembler, 5 different parameter sets were tested, with the full results available in Appendix 1 Table 3. This gives an indication of the sensitivity of assembly outcomes to parameter values and helps highlight assemblers where more thorough parameter optimisation may be required. The results of parameter testing for each assembler are given in Figures 4 and 5, showing percentage coverage and maximum contig length, respectively. This gives an indication of the possible spread of assembly quality that can be obtained by modifying parameters, such as the Kmer length. De Bruijn graph assemblers CLC, MetaSPAdes and ABYSS scored the highest values for both contiguity and coverage (Figure 2.4; Figure 2.5). CLC recorded the maximum values in both metrics, across all metagenomes tested, followed by metaSPAdes (Figure 2.4; Figure 2.5). For example, for CLC, maximum contig length reached 14,158 bp and a coverage of 67% for parameter sets 2 and 3 respectively (Metagenome D; Figure 2.4; Figure 2.5). This is compared to SSAKE, the worst performing assembler, which reached maximum values of 2064 bp (contig length) and 16% (coverage) for parameter sets 3 and 4 respectively (Metagenome D) (Figure 2.4). MetaSPAdes and ABYSS recorded maximum coverage values of 57% and 36% for Metagenome D, with the highest contig lengths reaching 10,045 bp and 50,90 bp respectively (Figure 2.4; Figure 2.5).

Despite de Bruijn based assemblers showing the most promising assembly outcomes, the spread of contiguity, completeness and size values across parameter sets far exceeds the simpler OLC and Greedy algorithm-based assemblers (Figure 2.4). This indicates simpler assemblers have a smaller range of outcomes, in comparison to more complex, highly modifiable assemblers. For example, for Metagenome D, CLC and metaSPAdes have a range of maximum contig lengths between 11,964 bp and 6,818 bp, respectively (Figure 2.4). This is compared to SSAKE, which had a range of 278 bp. This pattern is also reflected in the range of coverage scores, with CLC showing a spread of 66%, compared to 15% for SSAKE, in

assemblies of Metagenome D (Figure 2.4). Consequently, whilst de Bruijn based assemblers can provide the highest scoring coverage and contiguity values, they can also provide the largest range in scores. This is compared to the assemblers based on simpler OLC and Greedy algorithms, that express a narrower range of outcomes, despite modifying parameter values. This is indicative of pronounced sensitivity to parameterisation in more complex assemblers, and thus a greater need for parameter set optimisation or testing.



**Figure 2.4:** Coverage (%) of metagenome assemblies against the input dataset, for artificial metagenomes A-D. Results for each parameter set (1-5) are shown, with values for each assembler distinguished. Assemblers tested include: CLC, metaSPAdes, MIRA, ABYSS, SSAKE.

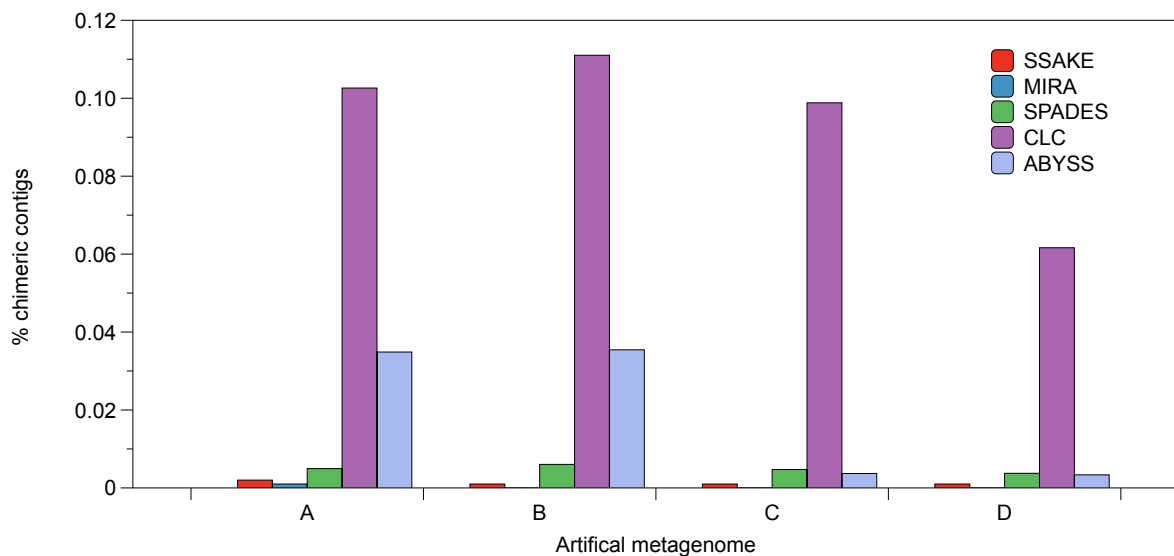


**Figure 2.5:** Maximum contig length (bp) of metagenome assemblies against the input dataset, for artificial metagenomes A-D. Results for each parameter set (1-5) are shown, with values for each assembler distinguished. Assemblers tested include: CLC, metaSPAdes, MIRA, ABYSS, SSAKE.



### 2.3.4 Assembly chimeras

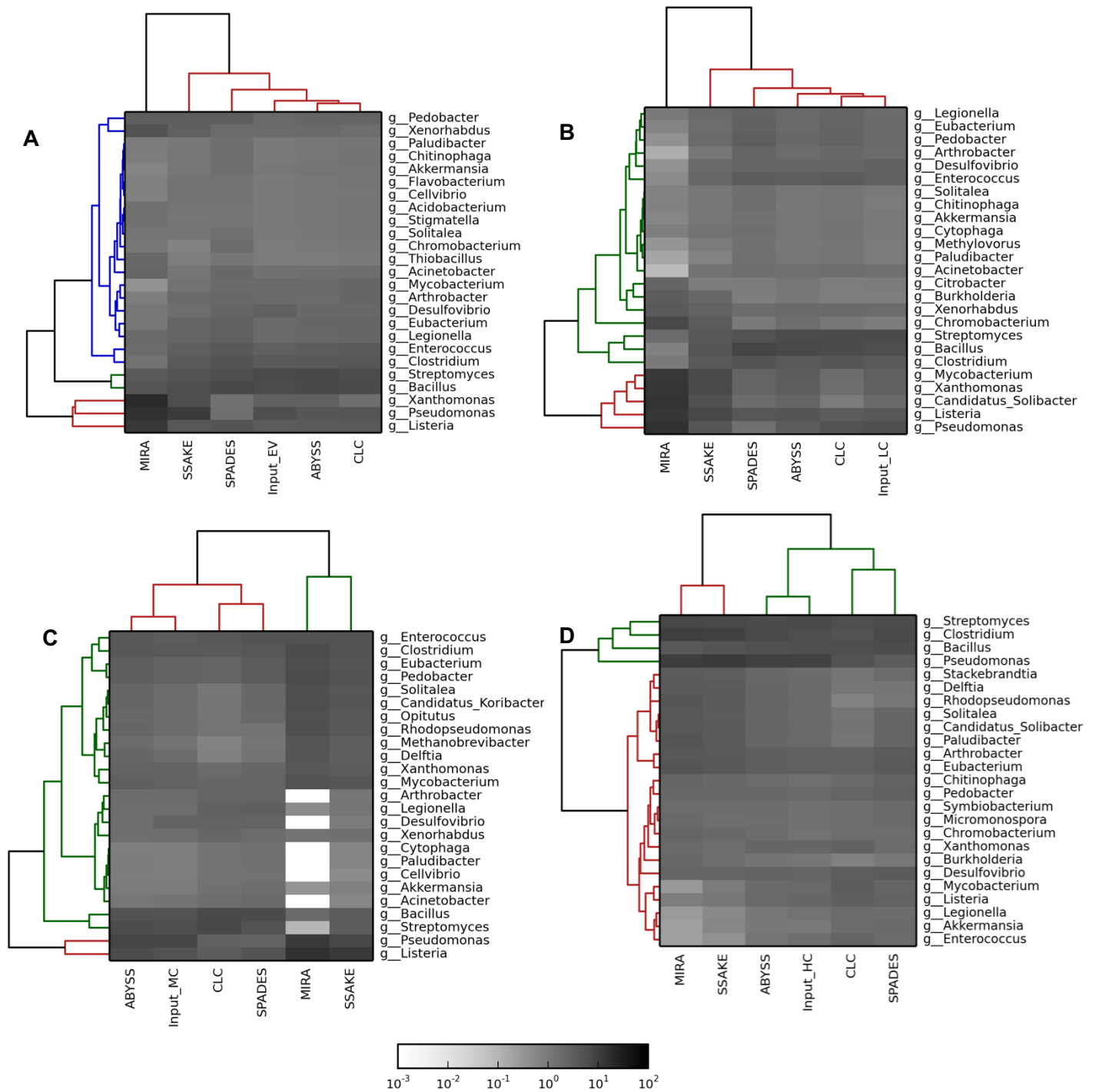
For all assemblies, chimeric (incorrectly assembled) contigs were identified using VSEARCH, which utilises an optimal global aligner to identify misassembled contigs in the final assemblies (Rognes *et al.*, 2016). This highlights which assemblers produce more accurate contigs, in comparison to those which produce longer, but incorrect sequences, which would adversely affect downstream functional analysis (Wooley *et al.*, 2010). As identified in Table 2.4, both CLC and metaSPAdes produced the longest contig lengths and metagenome coverage. However, CLC produced the largest percentage of chimeric contigs, in relation to the total number of contigs assembled, with a maximum number of chimeras at 0.1% for metagenome B, compared to 0.006% for metaSPAdes (Figure 2.6). This indicates that whilst CLC produces long contigs, they are slightly more prone to errors. The lowest number chimeric contigs were found in MIRA and SSAKE assemblers, accounting for a maximum of 0.002% of total assembled contigs (Figure 2.6). However, as these assemblies are generally smaller and more fragmented, they may be less useful for downstream functional analysis. For all assemblers tested, assemblies of metagenome D (high complexity) produced the lowest percentage of chimeric contigs, ranging between 0 and 0.06%, between MIRA and CLC assemblers respectively (Figure 2.6). This is in agreement with the coverage and contiguity statistics presented in Table 2.4, identifying metagenome D as having the best overall assembly statistics across all the artificial metagenomes tested (Table 2.4).



**Figure 2.6:** Chimeras as a percentage of total contigs for assemblies of metagenomes A-D, for each assembler. The results for the best performing parameter set for each assembler are shown (identified in Table 2.4).

### 2.3.5 Taxonomic distribution of assemblies

Alongside using chimera searches to highlight assembly accuracy, identifying the taxonomic distribution of outputs can also provide an insight into assembly quality. MetaPhlan2 was used to evaluate the taxonomic distribution of assemblies of each metagenome across all five assemblers tested, in comparison to the input unassembled metagenome (Figure 2.7; Segata *et al.*, 2012). This helps to highlight erroneous assemblies by isolating unusual community profiles, or over/under abundant organisms in assembled data (Segata *et al.*, 2012). Greedy based assembler, MIRA, was shown to perform poorly in representing the input metagenome in resulting assemblies. At the genus level, this assembler over represented genus' including *Listeria*, *Pseudomonas* and *Xanthomonas*, in comparison to the input dataset (Figure 2.7). This was shown to vary significantly between input metagenomes, with substantial over representation of *Gammaproteobacteria* and *Deltaproteobacteria* for high complexity Metagenome C, in comparison to the other assemblers (Figure 2.7). This pattern was not reflected in Metagenomes A, B and D. However, SSAKE, the second greedy-based assembler tested, appeared to more accurately represent the input metagenome in resulting assemblies (Figure 2.7). Additionally, whilst outputs from ABYSS, CLC and metaSPAdes (de Bruijn graph based assemblers), did not completely recreate the input dataset, they did not show a consistent bias in the taxonomic distribution of organisms (Figure 2.7). Consistent bias in assemblies is more problematic for downstream analysis, as inaccurate conclusions about the community composition may be drawn, as opposed to random assembly error, which is unlikely to skew the overall community distribution.



**Figure 2.7:** Taxonomic heat maps for each artificial metagenome assembled (A-D), showing the top 25 most abundant genus', clustered based on Euclidean distances. Relative genus abundance for the best parameter set for each assembler are shown, alongside the input microbial community for comparison (flagged by Input\_MC).

## 2.4 Discussion

### 2.4.1 Premise of the study

This study tested the performance of five metagenome assembly tools for the assembly of complex soil bacterial metagenomes. The artificial metagenomes were simulated using an identical community composition, but were modified based on organism abundance profiles, to produce metagenomes at a range of complexities.

Simulated metagenome datasets are generally simplifications of true environmental communities (Vazquez-Castellanos *et al.*, 2014). However, as the aim of this study was to test metagenome assemblers, rather than to study the taxonomy of the dataset, this does not undermine the conclusions drawn (Vazquez-Castellanos *et al.*, 2014). Additionally, using test datasets where the species composition is known *a priori*, allows a direct comparison of the assemblers (Mavromatis *et al.*, 2007). This would not be possible using true bacterial metagenomes, as the species composition is unknown prior to assembly (Mavromatis *et al.*, 2007). Consequently, using simulated datasets allows a comparison of how well the resulting assemblies represent the 'true' input bacterial sequence composition. The best performing assemblers will most accurately represent the composition of the test dataset, have minimal chimeric contigs, and have a large contig length to maximise the quality of downstream functional and taxonomic annotation (Howe and Chain, 2015).

### 2.4.2 Assembly evaluation metrics

This study has demonstrated the importance of using more than one evaluation metric for assembly comparison, and that measures incorporating contiguity, completeness and accuracy should be used. This is important as no single metric is comprehensive enough to summarise overall assembly quality (Kurtz *et al.*, 2004; Peng *et al.*, 2012; Vollmers *et al.*, 2017). When looking at assemblies based solely on contiguity, CLC shows the overall highest number of contigs and contig length, in comparison to SSAKE which produces the shortest contigs, and therefore most fragmented assemblies. This is important, as long contigs, such as those produced by CLC (and metaSPAdes) can provide better alignments to gene databases for functional annotation. However, when also considering the number of chimeric contigs, CLC exemplified the highest number of chimeras across all tested datasets (Figure 2.6). This indicates that whilst CLC is producing long contigs, these are sometimes miss-assembled, incorporating read fragments from different organisms. This could subsequently produce some incorrect taxonomic and functional annotations in downstream analysis

(Mavromatis *et al.*, 2007). Whilst this is not a substantial proportion of the overall number of contigs produced during CLC assemblies (0.1%) it is important to select an assembler that can produce both long, and correctly assembled contigs. As metaSPAdes was able to assemble contigs with substantial lengths (Table 2.4) and with a very minimal number of chimeras, at 0.006% of total contigs (Figure 2.6), this assembler also provides a reasonable trade-off between these metrics. Interestingly, an inter-comparison of assemblers by Vollmers *et al.*, (2017) based on real Illumina sequencing data (forest soil and algal biofilm) selected metaSPAdes as having the best trade-off between assembly size and coverage of diversity, in comparison to the assemblers tested. Whilst this study was primarily focused on de Bruijn graph-based assemblers, and did not explore parameter values, it does indicate the performance of this assembler is upheld in real sequence data (Vollmers *et al.*, 2017).

#### 2.4.3 Simulated bacterial communities

This study also demonstrated the use of simulated datasets in conducting metagenome assembly evaluation. Using artificial metagenome datasets allowed the assemblies to be compared based on percentage coverage and taxonomic composition, in comparison to the 'true' community structure (Sczyrba *et al.*, 2017). This consequently enabled the accuracy of the assemblers to be evaluated, alongside using more standard measures of assembly size and contiguity. Out of the two simpler 'Greedy' based assemblers tested (MIRA and SSAKE), MIRA showed the highest number of contigs, contig length and coverage of the input dataset (Table 2.4). However, when investigating the taxonomic composition of MIRA assemblies in comparison to the input dataset, MIRA is shown to both under and over represent the 25 most abundant organisms (Figure 2.7). This assembler is therefore not completely representing the true species composition and therefore may produce incorrect downstream annotations and ecological conclusions. Consequently, this demonstrated the value of using artificial datasets, and the importance of selecting a good assembler, in order to produce un-fragmented accurately assembled contigs (Risenfeld *et al.*, 2004).

Here, metagenome assemblers were also tested on several simulated communities, at a range of complexities, as opposed to a single dataset. This is because each assembler is formatted based on a different algorithm, with varying parameter options and will therefore suit datasets with different sizes and community structures (Miller *et al.*, 2010; Sczyrba *et al.*, 2017). Consequently, assemblers showed some differences in performance across the simulated communities, in relation to the algorithm on which they are based. Overall, high complexity metagenome 'D' was shown to be assembled with the highest coverage, contig lengths and the minimal number of chimeras (Figure 2.2; Figure 2.6). This contrasts with low

complexity metagenome 'B', and even abundance metagenome 'A', which demonstrated highest numbers of chimeric contigs and lowest contig lengths (Table 2.4; Figure 2.6). These differences in outcomes between metagenomes can be attributed to the community abundance structure, as the taxonomic composition is identical. These differences in assembly outcome may be due to raised abundance of some organisms, and subsequent increases in its read coverage in the dataset. Organisms with high abundance are likely to be well assembled, pulling the quality of the assembly up, as there is more data support for each DNA sequence. Metagenome B was classified as the low complexity community as it was comprised of 4 highly abundant organisms, with the remainder as low abundance flanking organisms (Table 2.3). This is in line with Mavromatis *et al.*, (2007) whom created test communities in MetaSim, to reflect simpler bacterial populations with a select group of organisms. This is similar to Metagenome A, where all the organisms have the same level of abundance. It is likely that these metagenomes were assembled poorly as there were limited organisms at a high abundance, and therefore most organisms had a low read coverage. This low coverage could have increased the difficulty in resolving artificial sequencing errors, gaps and repeats (Sims *et al.*, 2014). On the other hand, Metagenome D had more organisms at a higher abundance, and therefore the increase in read coverage may have improved the assembly outcome, despite having a more complex species distribution (Table 2.4; Sims *et al.*, 2014). However, the overall differences between assemblies of the different metagenomes were found to be insignificant for all assemblers (Appendix 1 Table 5). In the instance of low coverage datasets, read based analysis, as opposed to assemblies may therefore be beneficial for taxonomic investigations. This is because using reads will be able to profile more of the community, as opposed to assemblies which may not utilise all of the sequencing data.

#### 2.4.4 Choice of assembler

In contrast to community complexity, overall assembler selection was found to have a significant influence on the quality of the assembly (Appendix 1 Table 4). This indicates that overall assembler selection for the type of data in question (e.g. soil or sediments) is important, however the complexity of that community does not have a substantial influence on the assembly outcome (Table 2.4; Figure 2.2; Figure 2.3). For metaSPAdes assemblies, the percentage coverage increased by 15% between Metagenomes A to D, up to a maximum of 57% (Table 2.4). This can be compared to the increase of 7% up to a maximum coverage of 16% for SSAKE. Thus, the difference between SSAKE and metaSPAdes coverage scores for Assembly D is 41%, exceeding the 15% and 7% increase in coverage obtained by changing metagenome complexity for metaSPAdes and SSAKE, respectively (Table 2.4).

#### 2.4.5 The importance of parameterisation

The analysis highlighted that modifying key parameter values, such as the kmer length, can influence the assembly outcome (Figure 2.4; Figure 2.5). This was an important aspect to test, as many studies will run assemblies using default parameter values, which may not obtain the most optimal results (Howe *et al.*, 2014; Vollmers *et al.*, 2017). This was shown to be particularly evident for more complex de Bruijn Graph assemblers such as CLC and metaSPAdes (Figure 2.4; Figure 2.5). This is likely to relate to the wider range of parameter settings available in these assemblers or their sensitivity to changing key variables such as k-mer values. Whilst simpler assemblers often produce lower quality results, these are generally more consistent between parameterisations. Consequently, in order to obtain an improved assembly outcome, it is beneficial to iterate through several sets of parameter values, before making the final selection. This is in agreement with the CAMI inter-comparison, which identified parameters such as the kmer length, to be highly influential on assembly outcome (Sczyrba *et al.*, 2017). For studies which are highly focused on obtaining the most optimal assembly quality, parameter space optimisation can be used to select the most favourable parameter values (Chikhi & Medvedev, 2013). Some assemblers, such as metaSPAdes already incorporate optimisation of kmer length selection, by iterating through several values during the assembly process, before selecting the most appropriate to the dataset in question (Nurk *et al.*, 2017).

#### 2.4.6 Best performing assemblers

Overall, CLC and metaSPAdes, de Bruijn graph-based assemblers, were shown to perform best in terms of contiguity, size and coverage across all datasets (Figure 2.2; Figure 2.3). Greedy algorithm based assembler, SSAKE was shown to perform worst over these metrics. Furthermore, OLC based assembler, MIRA, did not show a substantial improvement in assembly outcome from the worst performing assembler, SSAKE (Figure 2.2; Figure 2.3). However, when considering the number of incorrectly assembled contigs, CLC shows the highest proportion of chimeras relative to its dataset size (Figure 2.6). Whilst this is not a substantial proportion, it is higher than the other assemblers tested. In contrast, simpler assemblers, MIRA and SSAKE show a smaller proportion of incorrectly assembled contigs in their resulting assemblies (Figure 2.6). However, as these assemblers have created much shorter contig lengths, and therefore smaller, more fragmented assemblies, the number of long, incorrectly joined reads is lower overall. However, as the contig lengths are much shorter for these assemblers than the de Bruijn graph-based assemblers, CLC and metaSPAdes, the accuracy of downstream functional annotation could be limited (Narzisi and Misha, 2011).

Consequently, it can be suggested that metaSPAdes and CLC produce suitable assemblies for the investigation of complex bacterial metagenomes formed from environmental samples. This is because the assemblies produced are accurate, large and provide long contigs for downstream annotations (Table 2.4; Figure 2.6; Figure 2.7). Whilst CLC does produce slightly more chimeric contigs, this is arguably a small proportion of contigs and therefore should not impact the inferences from the bulk of the data.

## **2.5 Conclusion**

This study has used artificially simulated metagenomic data to test publicly available metagenome assemblers, to provide better guidance on assembler selection for those investigating soil and sediment bacterial communities. The assemblers were selected to cover the range of assembly algorithms available, including de Bruijn graph assemblers (metaSPAdes, CLC, ABYSS), OLC assemblers (MIRA) and simple Greedy assemblers (SSAKE).

The study has shown the importance of using more than one metric when evaluating assemblies. Using one type of metric, such as contiguity, may provide misleading results, as long contigs may not necessarily be accurate. Incorrectly assembled contigs may provide inaccurate functional and taxonomic annotation in downstream analysis and may subsequently influence ecological conclusions. Evaluating based on assembly size may also be misleading, as large assemblies of short contigs will not provide optimal alignments to databases for functional annotation. Fragmented contigs may therefore inaccurately represent key groups of the dataset, as shown by MIRA assemblies, which missed out several highly abundant genera's. Consequently, testing the assemblers on artificial communities proved useful, as the assemblies could be evaluated based on their accuracy and coverage of the 'true' dataset. A combination of contiguity, assembly size, coverage and accuracy metrics are therefore needed to adequately summarise the quality of the metagenome assembly, or to compare between assemblies.

Assembler selection was shown to be important for the output quality, with de Bruijn based assemblers, CLC and metaSPAdes producing assemblies with the highest coverage, contiguity and assembly size. Assembly outcomes were shown to be significantly different across the assemblers tested, and therefore identifying a suitable assembler for the data type is important for obtaining high quality results. Whilst the assemblers were tested across communities at a range of complexities, no significant difference could be found between these, when looking at a single assembler. It is arguably more important to select an appropriate assembler for the data type, which can then be applied to similar data without



significant disadvantages. Some differences were obtained when modifying the community complexity, mainly in terms of assembly coverage, N50 and number of contigs over 1000bp. Therefore, if computational power and time permits, it is recommended to test the assemblers for the dataset in question, rather than just the data type. However, we would not expect substantial reductions in assembly quality when using an assembler tested on a similar data type.

Modifying parameter values was also shown to influence output quality. Whilst the assembly itself can improve the dataset quality for functional and taxonomic annotation, by testing several parameter sets, the assembly outcome can be improved. Whilst providing a full parameter optimisation was not the aim of this study, we have shown that the more complex de Bruijn graph assemblers are more sensitive to parameter settings, such as the kmer length. It would therefore be beneficial to test a few parameter sets, or carry out a parameter optimisation, if assembly quality was a key concern.

Overall, we have identified the metaSPAdes and CLC assemblers as providing the highest quality assemblies for soil and sediment bacterial datasets. These assemblers are recommended for use on datasets of similar types. However, the more the user attempts to improve the assembly outcome (i.e. testing for the dataset in question and modifying parameter settings), the greater the outcome quality can be. Improving the quality of the assembly will not only increase the reliability of the functional and taxonomic annotations, but also allow more detailed investigation of the microbial community. For example, improved annotation may aid extraction of draft genomes from metagenome datasets.

## **2.6 Observations**

The results of this chapter have highlighted a series of observations, when considering metagenome assembly within an analysis pipeline.

- Assembling metagenome reads can increase their length, which may help downstream annotation.
- Assemblers do not all perform the same, the most appropriate choice will depend on the data type and complexity.
- If possible, testing a selection of assemblers may help determine the most appropriate choice.
- Testing of assemblers can be carried out on real environmental data, or a simulated community (which will allow the evaluation of accuracy).

- Assemblers should be evaluated on multiple metrics: contiguity, completeness, size and accuracy.
- Parameter values can influence assembly outcomes, especially for de Bruijn graph assemblers. Testing several parameter sets may help improve assembly quality.
- Once testing has been carried out, the same assembler can be used on multiple samples from a similar environment, without significant cost to assembly quality.
- Assembly may not always be the most appropriate choice, for example, given high fragmented datasets with low sequencing coverage.

## **2.7 Recommendations for best practice**

The results of this analysis have highlighted some recommendations for best practice when using metagenome assembly as part of an analysis pipeline. Here it has been shown that de Bruijn graph assemblers, such as CLC and metaSPAdes are most appropriate for complex soil metagenomes. This is likely related to the use of a graph-based algorithm which identifies read overlaps and read layout, alongside multiple tuneable parameters. However, we recommend that if the resources are available, several de Bruijn graph-based assemblers are tested on the dataset under analysis, to obtain an improved outcome. Comparing and selecting an assembler can increase read coverage, assembly contiguity and completeness, which will improve downstream annotation opportunities. This analysis has shown that the simpler Greedy assemblers and those designed for single genomes should not be applied to metagenomic assembly. This is because these assemblers produce small localised assemblies through read overlaps, often creating fragmented datasets which hinder annotation.

Once an assembler has been selected, multiple parameter sets should be tested to obtain an improved result. This is especially the case for de Bruijn graph assemblers, which are sensitive to key parameter values such as the kmer length. A consideration should also be made to the expected composition and complexity of the dataset. This analysis has shown substantial differences in assembly outcome based on the complexity of the community, and that assembler choice should take this into account. For example, a simple community monopolised by a few microbes may assemble reasonably well with a Greedy based assembler, given the increased read coverage for these organisms. However, with a complex community structure, such as found in soils, a variability in microbial abundance and therefore

read coverage requires the use of a more tuneable assembler. Given limited computational resources and time, it is recommended that those assembling soil metagenomes use a de Bruijn graph assembler, and test several values of kmer length as a minimum. Some assemblers, such as metaSPAdes incorporate kmer length optimisation as part of the assembly process.

## **2.8 Limitations and additional work**

As the aim of this study was to test overall assembler function on simulated bacterial communities, parameter space optimisation was not carried out. Whilst the study could have been conducted by running each assembler on default values, we also wanted to show how modifying some parameter settings could enhance the results available. In this study we tested a total of five parameter sets for each assembler, including the default values. Whilst this was sufficient to highlight the sensitivity of more complex de Bruijn graph assemblers to parameters such as the kmer length, greater analysis of the parameter space could be carried out. This would include running optimisation of key parameters such as the kmer length, minimum contig length and bubble size (Chikhi and Mendeleev, 2013). Identifying the optimal parameter set is likely to improve the quality of the assembly outcome. This would be particularly important for studies that are interested in more detailed aspects of the microbial community structure, such as gene arrangements or assembling single draft genomes (Eloe-Fadrosh *et al.*, 2016b).

A subset of publicly available assemblers were selected for this study, in order to cover the three main assembly algorithms available. To obtain the most optimal assembler for soil microbial data, all publicly available assemblers could be included. This would require collaboration between several research centres, as carried out by Assemblethon 2 for the assembly of single vertebrate genomes (Bradman *et al.*, 2013). However, we have shown here that by selecting a de Bruijn graph assembler, such as CLC or metaSPAdes, assemblies of improved quality can be obtained compared to using simpler OLC or Greedy assemblers, which are better fitted to longer sequencing reads. In future, the soils community may wish to conduct a wider inter-comparison for metagenome data, in order to highlight the best performing assembler. This in itself has limitations, as the assemblers may only be tested on one data type, and as the field is constantly evolving with the release of new assemblers, this will also only be a snapshot of the assembly platforms available.

This study tested the selected assemblers on a series of artificial bacterial metagenomes. This was carried out so that the 'true' bacterial community was known, and therefore the assembly accuracies could be evaluated. The assemblers could also be tested on real bacterial data,

however, this would limit the evaluation metrics to contiguity, coverage and size criteria. The assemblies could also be repeated in order to test the reliability of each assembly, and how variable the assembly outcomes can be. As far as the authors are aware, there is no current knowledge on the reliability of assemblers when repeating assemblies of the same dataset.

# Chapter 3: Metagenomic insights into diazotrophic communities across Arctic glacier forefields

Maisie V. Nash<sup>1</sup>; Alexandre M. Anesio<sup>2</sup>; Gary Barker<sup>3</sup>, Martyn Tranter<sup>1</sup>, Gilda Varliero<sup>3</sup>, Emiley A. Eloë-Fadrosh<sup>4</sup>, Torben Nielsen<sup>4</sup>, Thomas Turpin-Jelfs<sup>1</sup>, Liane G. Benning<sup>567</sup> and Patricia Sánchez-Baracaldo<sup>1</sup>

<sup>1</sup>School of Geographical Sciences, University of Bristol, BS8 1SS, UK

<sup>2</sup>Department of Environmental Science, Aarhus University, PO box 358, Denmark

<sup>3</sup>School of Life Sciences, University of Bristol, BS8 1TQ, UK

<sup>4</sup>DOE Joint Genome Institute, 2800 Mitchell Drive, Walnut Creek, CA 94598, US

<sup>5</sup>GFZ German Research Centre for Geosciences, Telegrafenberg, 14473 Potsdam, Germany

<sup>6</sup>School of Earth and Environment, University of Leeds, LS2 9JT, Leeds, UK

<sup>7</sup>Department of Earth Sciences, Free University of Berlin, Malteserstr, 74-100, Building A, 12249, Berlin, Germany

## Contributions and acknowledgements

This chapter has been published in *FEMS Microbiology Ecology*, volume 94, issue 9 (12 June 2018). The full reference is detailed below:

Nash, M. V., Anesio, A. M., Barker, G., Tranter, M., Varliero, G., Eloë-Fadrosh, E. A., ... & Sánchez-Baracaldo, P. (2018). Metagenomic insights into diazotrophic communities across Arctic glacier forefields. *FEMS microbiology ecology*, 94(9), fiy114

MN carried out the data analysis, interpretation and manuscript formulation. AA, GB and MT carried out field sampling. TTJ supplied the raw carbon and nitrogen data for the Midtre Lovénbreen forefield. GV carried out the read mapping. All co-authors were involved in the refinement and corrections of the manuscript. Additional support came from Katherine Wright, Joshua Blacker and James Bradley during field sampling. Jane Coghill and Christy Waterfall at the Bristol Genomics Facility carried out the DNA sequencing and Nathan Christmas provided data analysis support.

## Funding

This work was carried out as part of the NERC GW4+ Doctoral Training Partnership, awarded to MN. Funding support was provided through NERC grants NE/J02399X/1, awarded to AA and NE/J022365/1 awarded to LGB. Funding support for PS-B came from a Royal Society University Research Fellowship and LGB also acknowledges the financial support through the Helmholtz Recruiting Initiative. Funding support for GV came from the European Union's Horizon 2020 research and innovation programme under the Marie Skłodowska-Curie grant agreement No 675546 – MicroArctic. The work conducted by the U.S. Department of Energy Joint Genome Institute, a DOE Office of Science User Facility, is supported by the Office of Science of the U.S. Department of Energy under Contract No. DE-AC02-05CH11231.

### 3.1 Introduction

Arctic glaciers are undergoing fast retreat, exposing soils that have been locked under ice for thousands of years (Bradley, Singarayer and Anesio, 2014). Microbial communities have been identified as the primary colonisers of these newly exposed soils (Schmidt *et al.*, 2008) and are important for building up initial carbon and nitrogen pools, enhancing soil stability through the release of exopolymeric substances, and mediating forefield soil pH (Sattin *et al.*, 2009; Schulz *et al.*, 2013). However, there is a lack of coherent understanding on the diversity and biogeochemical importance of these bacterial communities in relation to nitrogen fixation (Brankatschk *et al.*, 2011). Bacterial nitrogen fixation uses the enzyme nitrogenase to convert atmospheric nitrogen ( $N_2$ ) into fixed ammonia ( $NH_3$ ) for biological uptake by non-diazotrophic organisms (Brill 1975). As nitrogen is a key nutrient for microbe and plant growth, nitrogen limited forefield soils may place restrictions on heterotroph colonisation, productivity and succession (Duc *et al.*, 2009). Subsequently, diazotrophic organisms have been proposed as crucial facilitators of succession in newly exposed forefield soils (Knelman *et al.*, 2012). Nitrogen-fixing cyanobacteria have been identified as key in building these initial nitrogen stocks, and therefore expediting the establishment of heterotrophic organisms (Kaštovská *et al.*, 2005; Schmidt *et al.*, 2008; Duc *et al.*, 2009).

Whilst the importance of early diazotrophs is evident, similarities and variations in the nitrogen-fixing communities across forefields, in terms of both diversity and phylogeny, have received limited attention. The majority of research to date has focused on understanding changes in nitrogen fixation within individual forefields, along transects or chronosequences of soil development (Duc *et al.*, 2009; Brankatschk *et al.*, 2011). Thus far, the taxonomic diversity and abundance of the *nifH* gene, encoding nitrogenase for nitrogen fixation, has been shown to decrease with soil age and distance from the glacier terminus, in line with increasing fixed nitrogen in soils, and a reduced need for diazotrophy (Duc *et al.*, 2009; Brankatschk *et al.*, 2011). The dominant diazotrophic community composition in forefields is likely to be influenced by factors such as soil physicochemical status, climate, topography, the establishment of plants and any disturbances, such as water flow pathways, which may elicit both similarities and differences in diazotrophy between sites (Hodkinson, *et al.*, 2002; Nicol *et al.*, 2005; Schütte *et al.*, 2010; Liu *et al.*, 2012). Furthermore, the current body of evidence surrounding microbial succession in forefields has a limited geographical range, with most studies conducted in the Damma Glacier forefield in Switzerland (Duc *et al.*, 2009; Frey *et al.*, 2010;

Bernasconi *et al.*, 2011; Brankatschk *et al.*, 2011; Brunner *et al.*, 2011; Zumsteg *et al.*, 2013, 2013; Bradley *et al.*, 2015). Investigation across multiple glacier forefields is needed to fully explore similarities and differences between forefields in terms of diazotrophic community composition and their phylogenetic relations (Schütte *et al.*, 2010). This will help highlight the microbial community diversity involved in nitrogen fixation among glacier forefields.

Bacterial nitrogen fixation is encoded by clustered nitrogenase (*nif*) genes, typically through an enzyme containing an iron (Fe) cofactor and a molybdenum-iron (Mo-Fe) cofactor (Dixon and Kahn, 2004). Overall, the abundance of bioavailable nitrogen controls the transcription of nitrogenase genes, whilst the variant of nitrogenase transcribed is regulated by the presence of molybdenum (Oda *et al.*, 2005; Teixeira *et al.*, 2008). In the absence of Mo, nitrogenase is transcribed with vanadium (Fe-V co-factor), or exclusively with iron (Fe-Fe cofactor) in the absence of both Mo and V (Raymond *et al.*, 2004; Teixeira *et al.*, 2008). These nitrogenases are in turn encoded by the *nifHDK*, *vnfH-vnfDGK* and *anfHDGK* operons (Dixon and Kahn 2004; Teixeira *et al.*, 2008). The phylogenetically conserved *nifH* gene can be used to classify bacterial diazotrophs into Clusters I-IV based on the nitrogenase (Chien and Zinder 1996). Cluster I covers the typical Mo *nifH*, whilst Cluster II covers the alternative *vnfH* and Cluster III generally includes a diverse range of anaerobic bacteria (Zehr *et al.*, 2003). Furthermore, Cluster IV contains organisms with 'nif-like' sequences, as opposed to conventional *nif* genes (Zehr *et al.*, 2003).

Previous research conducted on microbial succession in glacial forefields, including those on functional genes, has mostly focused on marker gene data, such as the universal bacterial marker 16s rRNA and amplified *nifH* (Schmidt *et al.*, 2008; Brankatschk *et al.*, 2011; Rime *et al.*, 2015). However, studies are now applying alternative methods, such as metagenomics, to study microbial communities (Wooley, Godzik and Friedberg, 2010). This is because metagenomics can provide gene sequences for the entire microbial community gene pool, rather than target sequences (Handelsman 2004; Daniel 2005). Thus, both microbial diversity and functional potential can be inferred using one approach (Wooley, Godzik and Friedberg 2010; Thomas *et al.*, 2012). To maximise the quality of the output metagenome, the short DNA fragments from next generation sequencing can be assembled (Vázquez-Castellanos *et al.*, 2014). This generates longer continuous DNA reads (contigs), which can provide enhanced functional and taxonomic annotations (Howe *et al.*, 2014; Vázquez-Castellanos *et al.*, 2014).

In this study, we investigated 70 soil metagenomes spanning transects and chronosequences across four Arctic forefields in N-Sweden, Greenland and Svalbard. The datasets have been assembled separately and subsequently annotated for use in a comparative metagenomics analysis. This study leads on from Chapter 2, by utilising metagenome assembly with metaSPAdes in an environmental setting. Here, we use metagenomics to present an investigation into the taxonomy and phylogenetic relationships of the functional genes recovered relating to bacterial nitrogen fixation in the four forefields. This analysis aims to contribute to the existing knowledge on pioneer microbial communities, helping to identify key genera of diazotrophic bacteria, which may have a key role building labile nitrogen stocks and soil development in oligotrophic forefield soils.

## **3.2 Materials and methods**

### **3.2.1 Field sampling**

Four Arctic glacier forefields were selected for sampling and analysis, in front of Rabots glacier (Rb), N-Sweden (67° 54' 25.6284" N, 18° 26' 51.0792" E); Storglaciären (St), N-Sweden (67° 52' 21.1116" N, 18° 34' 2.676" E); Midtre Lovénbreen (MI), Svalbard (79° 6' 1.8" N, 12° 9' 21.996" E) and Russell Glacier (RI), Greenland (67° 9' 23.4324" N, 50° 3' 50.342" W). Samples were obtained in July 2013 (Midtre Lovénbreen) and July 2014 (Russell, Rabots and Storglaciären). Surface soil from each site was sampled using a chronosequence/transect-based approach, constructing three parallel transects along the forefield moving away from the terminus (Bradley, Singarayer and Anesio 2014). Chronosequence-based sampling was used to capture the diversity in nutrient concentration and microbial taxonomy of each forefield, to make more holistic comparisons between glacial forefields. Bulk surface samples were collected into sterile Whirlpak bags, and frozen at -20°C. Observationally, the sites comprised soils at very different development stages. A 'typical' smooth successional chronosequence from bare ground, to more developed, plant colonised soil was observed in the MI forefield. However, the other sites sampled had a more heterogeneous chronosequence, with earlier and often more patchy plant colonisation.

### **3.2.2 Soil organic carbon and total nitrogen content**

Soil total nitrogen (TN) and total organic carbon (TOC) were determined using mass spectrometry on a FlastEA 1112 nitrogen and carbon elemental analyser. The protocol described in Hedges and Stern (1984) was used for sample preparation. In brief, for TN



analysis soil samples were weighed and dried at 50°C overnight, before subsamples were transferred into tin capsules. For TOC analysis, 2 ml of 1 M HCL was incrementally added to 0.1 g of sample ( $W_o$ ) until effervescence stopped. Subsequently samples were again dried overnight at 50°C, left to equilibrate with hygroscopic salts, and re-weighed ( $W_f$ ). Finally, subsamples were transferred into tin vials for analysis. The percentage of TOC in each sample was calculated using a correction for acidification induced weight change (Equation 2.1). Where possible, three environmental replicates were analysed for each TN and TOC per sampling site.

**Equation 2.1:** Correction for weight change during acidification of samples for organic carbon elemental analysis.

$$\%OC = \left[ \frac{100 \times \text{mgOC}}{\text{mg sample}} \right] \times \left[ \frac{W_f}{W_o} \right]$$

Where  $W_o$  is sediment dry weight prior to acidification, and  $W_f$  is dry weight after acidification.

### 3.2.3 DNA extraction, library preparation and sequencing

As this study was focused on the microbial diversity in bulk surface soil, DNA was extracted using a Mo-Bio DNAEasy PowerSoil DNA extraction kit (QIAGEN, UK), with DNA yield quantified using a Qubit 2.0 fluorometer. Samples that yielded less than 50 ng of DNA during extractions were pooled with their field replicates prior to sequencing. This method has been previously shown to obtain high DNA yields from soils and has been used for soil microbial diversity analysis in a number of studies, including root microbiomes (Fierer *et al.*, 2007; Allison *et al.*, 2008; Inceoğlu *et al.*, 2010; Carvalhais *et al.*, 2013; Vishnivetskaya *et al.*, 2014). However, as this approach is not directly targeting the soil rhizosphere communities, there may be limitations to DNA extraction from this subset of the microbial community. Metagenomes were sequenced using an Illumina Next-Seq 500 (Rb, St and RI) and an Illumina-Mi Seq (MI), with a TruSeq library prep kit at the University of Bristol Genomics facility. A total of 70 metagenomes were sequenced across the four sites using 2x 150bp (Rb, ST, RI) and 2x 100bp (MI) paired-end reads (Appendix 2 Table 1). Sequencing read output for each site can be identified in Appendix 2 Table 2, ranging between 3 817 852 and 10 510 0186 reads per metagenome.

### 3.2.4 Metagenome assembly and annotation

The 70 sequenced datasets were quality trimmed and subsequently assembled individually using the metaSPAdes 3.10.0-dev assembler, a development release of metaSPAdes, tested in Chapter 2 (Bankevich *et al.*, 2012). These assemblies were carried out in collaboration with the DOE Joint Genome Institute (Walnut Creek, CA), using the BFC algorithm for read error correction (Li 2015), and the --meta and --only-assembler flags. Furthermore, incremental Kmer lengths were used (22, 33, 55 and 77) to identify the most appropriate parameter value for assembly. Assembly size for each metagenome ranged between 241,660 and 429,543,524 bases (Appendix 2 Table 2). Functional annotation of the 70 metagenomes was subsequently carried out using the Integrated Microbial Genomes with Microbiome Samples (IMG/M) system (Chen *et al.*, 2017). Rarefaction curves were created in MG-RAST 4.0.3 for each metagenome (Appendix 2 Figures 1–4; Meyer *et al.*, 2008). Each metagenome was evaluated based on the number of contigs assembled and species obtained, to highlight metagenomes that may be under sampled through sequencing. Under sampling can occur in highly diverse metagenomes, where the sequencing is not adequate to reveal all taxa present in the sample (Torsvik and Øvreås 2002). Consequently, in under sampled datasets, some organisms, particularly those which were less abundant, may not be included in the output metagenome (Rodriguez and Konstantinidis 2014).

For each metagenome, the *nifH* gene for nitrogen fixation was searched using the Basic Local Alignment Search Tool for Proteins (BLAST-p) with an e-value of  $1e^{-5}$  and extracted. As *nifH* genes are generally found in a phylogenetically conserved nitrogenase cluster (with *nifD*, *K*, *N* and *E*), these genes were also searched for and extracted (Howard and Rees 1996). *Nif* genes were dereplicated, removing duplicate copies, using VSEARCH 2.6.0, leaving a total of 185 assembled *nif* genes for subsequent analysis (Rognes *et al.*, 2016). The *nif* genes used for the analysis have been deposited in GenBank, under accession numbers MH551286 - MH551470. Gene abundance was calculated as a combined value of *nifHDKNE*, normalised in relation to the abundance of the bacterial single copy housekeeping gene, *rpoB*, for each site (Vos *et al.*, 2012; Ishii *et al.*, 2015). As this method relies on sequencing unamplified genes, the *nif* gene counts are limited and may not be exhaustive for individual samples. This is particularly the case for unamplified sequencing of complex microbiome datasets, such as soil samples (Rodriguez and Konstantinidis 2014). Additionally, diazotrophs can contain multiple different *nif* genes, and several copies of a single variant, so should not be used as a measure to enumerate the explicit number of diazotrophs in each sample (Zehr *et al.*, 2003).

Finally, the raw sequencing reads were mapped to the extracted nif contigs for each metagenome using the BWA-MEM algorithm (Li and Durbin 2009). The alignment score (AS) of each read/contig is reported, which numerically indicates the quality of the alignments.

### 3.2.5 Nif taxonomy

The taxonomic distribution of all nif sequences (HDKNE) was carried out using a Last Common Ancestor (LCA) analysis in MEGAN 6.9.0 (Huson *et al.*, 2016). For each forefield, nifHDKNE sequences were nucleotide BLAST (BLASTn) searched against an NCBI GenBank database of complete bacterial genomes. The sequences were subsequently binned based on the NCBI taxonomy, using an LCA algorithm, and visualised at the genus level for each forefield (Huson *et al.*, 2016).

### 3.2.6 Gene phylogeny

A phylogeny for nifH, based on clusters identified in Zehr *et al.*, (2003), was carried out, as this gene is supported by the largest body of research. Sample nifH sequences were aligned to sequences of cultured isolates, largely derived from the phylogeny by Deslippe and Egger (2006). GenBank and UniProtKB accession numbers for cultured isolates are available in Table 3.1. DNA sequence alignments were generated in SATé 2.2.7, using MAFT, MUSCLE and FASTTREE (Liu *et al.*, 2011). The GTR+CAT model was implemented, with the decomposition set to longest (to minimise long branch attraction) and a maximum number of iterations set to 8. Alignments were manually edited in Mesquite, alongside generating Nexus and Phylip format files (Maddison and Maddison 2017). Maximum likelihood phylogenies were carried out using the CIPRES implementation of RAXML-HPC2 8.2.10 on XSEDE<sup>1</sup>, with 1000 bootstrap iterations (Stamatakis, 2014). The GTR+G model of nucleotide substitutions was implemented, as identified with j model test (Guindon and Gascuel 2003; Darriba *et al.*, 2012). Trees were evaluated using Figtree 1.4.3<sup>2</sup>, before annotation with EvoView v2<sup>3</sup> (He *et al.*, 2016). Graphical enhancements were made using Inkscape 0.92.2<sup>4</sup>. Comparisons between nifH sample sequences and cultured isolates were made using NCBI BLASTn<sup>5</sup>, to identify nearest cultured relatives.

**Table 3.1:** GenBank and UniProtKB accession numbers for *nifH* sequences derived from Deslippe and Egger (2006), for use in the *nifH* phylogeny.

Database	Accession number	Species	Gene
GenBank	X13519.1	<i>Azotobacter vinelandii</i>	<i>vnfH</i>
GenBank	AY367395.1	<i>Kiebsiella variicola</i>	<i>nifH</i>
GenBank	AF484674.1	<i>Methylomonas rubra</i>	<i>nifH</i>
GenBank	AF216883.1	<i>Azomonas agilis</i>	<i>nifH</i>
GenBank	V01215.1	<i>Rhizobium meliloti</i>	<i>nifH</i>
GenBank	U97122.1	<i>Azoarcus tolulyticus</i>	<i>nifH</i>
UniProtKB	P26251	<i>Azorhizobium caulinodans</i>	<i>nifH</i>
GenBank	AJ515294.1	<i>Paenibacillus azotofixans</i>	<i>nifH</i>
GenBank	Z31716.1	<i>Nostoc</i> sp.	<i>nifH</i>
UniProtKB	P33178	<i>Anabaena</i> sp.	<i>nifH</i>
UniProtKB	P08925	<i>Frankia alni</i>	<i>nifH</i>
GenBank	X57006.1	<i>Frankia</i> sp.	<i>nifH</i>
GenBank	ABQ25379.1	<i>Geobacter uraniireducens</i>	<i>nifH</i>
GenBank	M23528.1	<i>Azotobacter vinelandii</i>	<i>anfH</i>
UniProtKB	P16269	<i>Azotobacter vinelandii</i>	<i>anfH</i>
UniProtKB	Q07942	<i>Azotobacter capsulatus</i>	<i>anfH</i>
GenBank	AF065617.1	<i>Chlorobium tepidum</i>	<i>anfH</i>
GenBank	AY221832.1	<i>Pelodictyon lutolum</i> CC11OA0	<i>anfH</i>
GenBank	AF227926.1	<i>Desulfovibrio salexigens</i>	<i>nifH</i>
GenBank	AY040513.1	<i>Desulfomicrobium baculatum</i>	<i>anfH</i>
UniProtKB	P25767	<i>Methanococcus thermolithotrophicus</i>	<i>nifH</i>
GenBank	AY029234.1	<i>Methanosarcina mazei</i>	<i>nifH</i>
UniProtKB	P00456	<i>Clostridium pasteurianum</i>	<i>nifH</i>
GenBank	AF065618.1	<i>Desulfonema limicola</i>	<i>anfH</i>
GenBank	AF216881.1	<i>Acetobacterium woodii</i>	<i>anfH</i>

### 3.3 Results and Discussion

#### 3.3.1 Soil carbon and nitrogen

**Table 3.2:** Summary statistics for total nitrogen (TN) and total organic carbon (TOC) across the four forefields (Midtre Lovénbreen MI, Russell RI, Storglaciären St and Rabots Rb). The average, minimum, maximum and standard deviation (SD) across each forefield is given. The detection limit for both TN and TOC was  $1 \text{ mg g}^{-1}$ . Sites recording values below detection (b.d) are shown.

<b>TN (<math>\text{mg g}^{-1}</math>)</b>	<b>Average</b>	<b>Minimum</b>	<b>Maximum</b>	<b>SD</b>
<b>MI</b>	b.d.	b.d	4.90	1.56
<b>RI</b>	1.95	b.d	6.94	2.15
<b>St</b>	b.d.	b.d	4.19	0.93
<b>Rb</b>	1.04	b.d	3.35	1.33

<b>TOC (<math>\text{mg g}^{-1}</math>)</b>	<b>Average</b>	<b>Minimum</b>	<b>Maximum</b>	<b>SD</b>
<b>MI</b>	10.56	b.d.	72.36	21.14
<b>RI</b>	26.36	b.d.	82.70	26.35
<b>St</b>	2.78	b.d.	27.89	6.25
<b>Rb</b>	6.81	b.d.	22.90	9.66

The range of values obtained within and between forefields for TOC and TN for samples from each forefield is listed in Table 3.2. These values include TOC and TN from both microbial and plant sources. Looking at average nutrient contents, comparing across the forefields, TN content ranges from averages below detection to  $1.95 \text{ mg g}^{-1}$ , between St and RI, respectively (Table 3.2). TOC content follows the same trend, increasing from the two Swedish glaciers (St and Rb), to MI and RI. Results from a one-way ANOVA analysis for each nutrient did not show any statistically significant differences in the TN measured between forefields ( $P > 0.05$ ). However, concentrations of TOC were found to vary significantly ( $P = 0.002$ ) (Table 3.3). Additional analysis of the TOC variance between forefields using a post-hoc Tukey analysis revealed the significant difference was between the St and RI forefields, with RI containing almost 10 times the TOC content of St on average ( $P < 0.01$ , Table 3.2; Table 3.4).

**Table 3.3** ANOVA comparing differences between the four forefields (Midtre Lovénbreen, Russell, Rabots and Storglaciären) based on total nitrogen (TN) and total organic carbon (TOC). Significant differences observed between the forefields are noted at the 0.01 or 0.05 level.

	<b>TN</b>	<b>TOC</b>
<b>f-ratio value</b>	2.46	5.375
<b>p-value</b>	0.071	0.002357
<b>Significance level</b>	Not significant	0.05

**Table 3.4:** Results of a post-Hoc Tukey analysis, comparing differences between forefields (Midtre Lovénbreen MI, Russell RI, Storglaciären St and Rabots Rb), based on total nitrogen (TN) and total organic carbon (TOC). Significant differences between the forefields are noted at the 0.01 significance level.

<b>TN</b>	<b>Treatments</b>	<b>Q statistic</b>	<b>p-value</b>	<b>Inference</b>
	<b>MI vs RI</b>	2.65	0.25	insignificant
	<b>MI vs Rb</b>	0.14	0.89	insignificant
	<b>MI vs St</b>	0.89	0.89	insignificant
	<b>RI vs Rb</b>	1.54	0.68	insignificant
	<b>RI vs St</b>	3.64	0.05	insignificant
	<b>Rb vs St</b>	0.72	0.89	insignificant
<b>TOC</b>	<b>Treatments</b>	<b>Q statistic</b>	<b>p-value</b>	<b>Inference</b>
	<b>MI vs RI</b>	3.56	0.07	insignificant
	<b>MI vs Rb</b>	0.52	0.89	insignificant
	<b>MI vs St</b>	1.66	0.62	insignificant
	<b>RI vs Rb</b>	2.79	0.21	insignificant
	<b>RI vs St</b>	5.39	0.00	** p<0.01
	<b>Rb vs St</b>	0.56	0.89	insignificant

Samples from the RI forefield revealed the widest range in both TOC (below detection—82.70mg g<sup>-1</sup>) and TN (below detection—6.94mg g<sup>-1</sup>), respectively (Table 3.2). This contrasts with the Rb forefield, where TOC and TN values expressed a smaller range, from below detection to 22.90 mg g<sup>-1</sup> and below detection up to 3.35 mg g<sup>-1</sup>, respectively. A range of values is expected across sites within each forefield, due to soil development which takes place over successional chronosequences and given variations in sources of autochthonous

and allochthonous material (Bradley, Singarayer and Anesio 2014), for example, in the deposition of aeolian material (such as soot), or the presence of ancient *in situ* organic pools, exposed by glacier retreat (Schulz *et al.*, 2013; Bradley *et al.*, 2015). For example, across the MI chronosequence TN and TOC increase from below detection and 2.85 mg g<sup>-1</sup>, to 4.4 mg g<sup>-1</sup> and 14.5 mg g<sup>-1</sup>, in line with expected soil development (Table 3.5; Bradley *et al.*, 2016). However, whilst differences in soil nutrient content do occur between sites, the values fall into the general range observed from other forefields (1–2 mg g<sup>-1</sup> nitrogen, and 0.1–40 mg g<sup>-1</sup> carbon) (Bradley, Singarayer and Anesio 2014) and are indicative of a generally oligotrophic environment.

**Table 3.5:** Initial and final concentrations of total nitrogen (TN) and total organic carbon (TOC) in forefield soils. Values are shown at the start of the transect/ chronosequence (by the glacier terminus) and at the end of the transect.

	Start TN (mg g <sup>-1</sup> )	End TN (mg g <sup>-1</sup> )	Start TOC (mg g <sup>-1</sup> )	End TN (mg g <sup>-1</sup> )
<b>MI</b>	b.d	4.40	2.85	14.47
<b>St</b>	b.d	b.d	b.d	1.24
<b>Rb</b>	b.d	1.74	b.d	b.d
<b>RI</b>	b.d	1.33	b.d	17.46

### 3.3.2 Rarefaction analysis

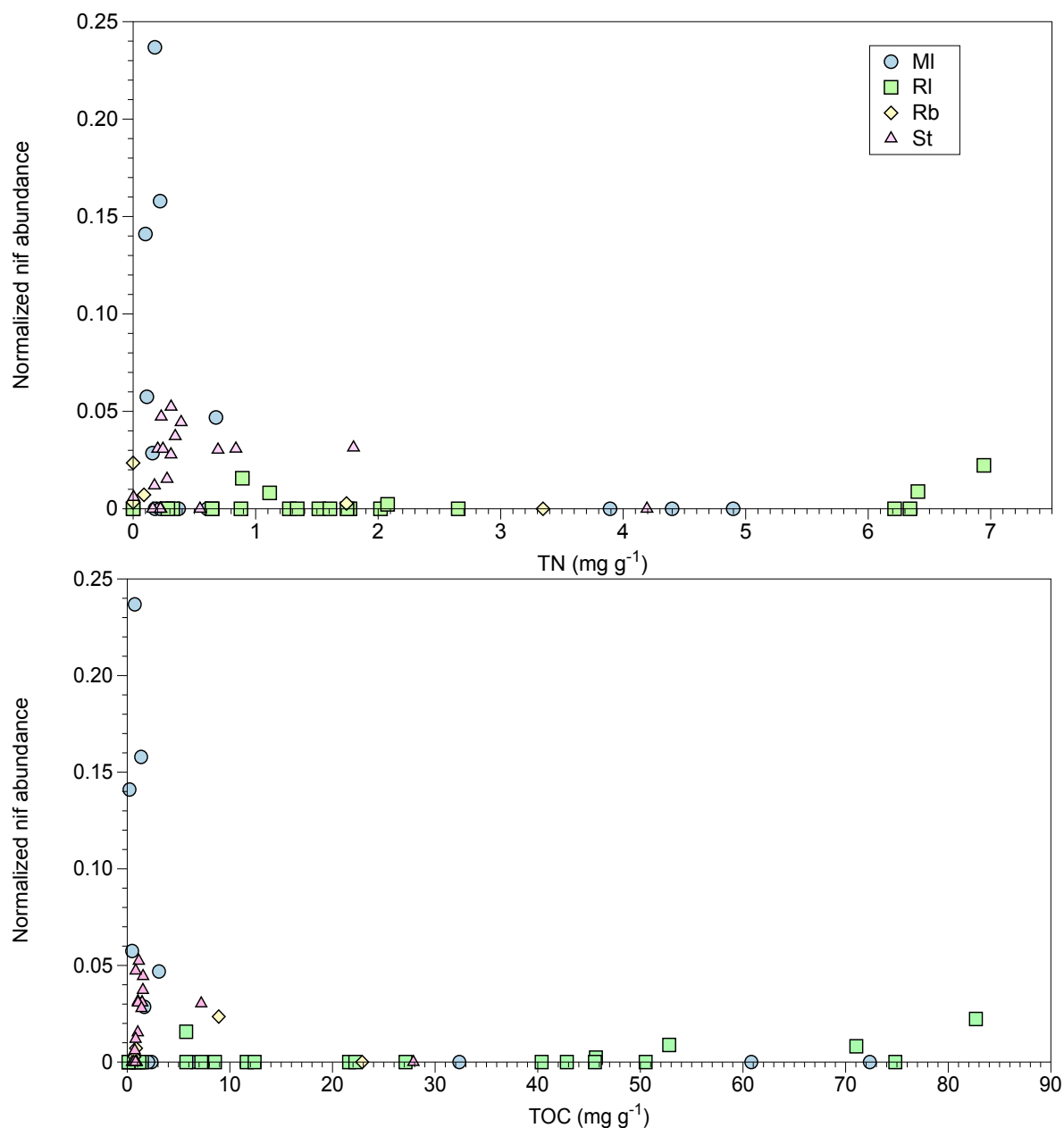
Rarefaction analysis was used to investigate the coverage of diversity in each metagenome, identifying any datasets where species content may be under sampled (Appendix 2 Figures 1-4). For each forefield, an assortment of both adequately sequenced and under sampled metagenomes were obtained (Figure 1-4, Appendix 2). Metagenomes that show rarefaction curves to reach saturation are likely to adequately profile the microbial diversity in the samples, for example metagenomes MI 7, RI 15, St 16 and St 17 (Figures 1, 2 and 3, Appendix 2). However, those metagenomes in which species number does not reach saturation are most likely to exclude taxa, for example ML1, ML 20, RI 14 and RI 20 (Figure 1 and 2, Appendix 2). In these metagenomes, the least abundant taxa are most probably excluded from the dataset, due to the reduced abundance of DNA for sequencing from these organisms (Rodriguez and Konstantinidis 2014). Whilst this does not detract from conclusions drawn on the organisms present in the samples, the full depth of diversity in under-sampled metagenomes cannot be highlighted. This issue is often prevalent in highly complex datasets such as soil and can only

be resolved through continued deeper sequencing of those metagenomes (Rodriguez and Konstantinidis 2014).

### 3.3.3 Nif genes recovered

The total abundance of dereplicated *rpoB* normalised contigs containing nif genes (nifHDKNE), in relation to the variation of TN and TOC, spanning all sampling sites is shown in Figure 3.1. A total of 185 nif genes contained on assembled contigs were recovered from the datasets. In 75% of samples where nif genes were detected, the TN and TOC concentrations fell below 1 and 5 mg g<sup>-1</sup>, respectively (Figure 3.1). Conversely, in samples where nif genes were not detected, 61% and 49% measured below 1 and 5 mg g<sup>-1</sup>, of TN and TOC, respectively (Figure 3.1). As sequencing output varied substantially between metagenomes, further sequencing may reveal additional genes due to the complex nature of soil microbiome samples (Table 2, Appendix 2; Rodriguez and Konstantinos 2014). However, this may indicate that samples with limited TN/TOC could have a larger relative abundance of genes for diazotrophy, as these were recovered through the sequencing effort undertaken. Interestingly, a similar trend between nitrogen fixation and TN has been reflected by the assays carried out by Telling *et al.*, (2011), whereby fixation rates on Arctic glaciers were negatively correlated with total inorganic nitrogen content. Additionally, a link between nif gene abundance and activity is supported theoretically, as fixation becomes less metabolically beneficial when labile nitrogen stocks increase (Gutschink *et al.*, 1978). When applied to forefield soils, both TN and TOC have been shown to increase over successional chronosequences, indicating nitrogen fixation may become less profitable with soil development (Duc *et al.*, 2009; Brankatschk *et al.*, 2011; Bradley, Singarayer and Anesio 2014). Furthermore, research by Brankatschk *et al.*, (2011) identified a link between nif gene abundance and enzyme activity, indicating sites with high numbers of nif genes, such as Storglaciären, would have enhanced nitrogen fixation activities. However, the relationship between gene abundance and nitrogen fixation activity is not always fully defined, as areas with low nitrogenase activity have previously been linked to high gene abundance in the Damma Glacier (Swiss Alps) (Duc *et al.*, 2009).





**Figure 3.1:** Relationship between normalized *nif* gene abundance (*nifHDKNE*) and concentration of total organic carbon (TOC) and total nitrogen (TN) per gram of soil, across all sampling sites. *Nif* gene abundance values are normalized against the bacterial single copy housekeeping gene, *rpoB*, for each metagenome. Values across the different forefields are noted, including: Midtre Lovénbreen (MI), Russell (RI), Rabots (Rb) and Storglaciären (St).

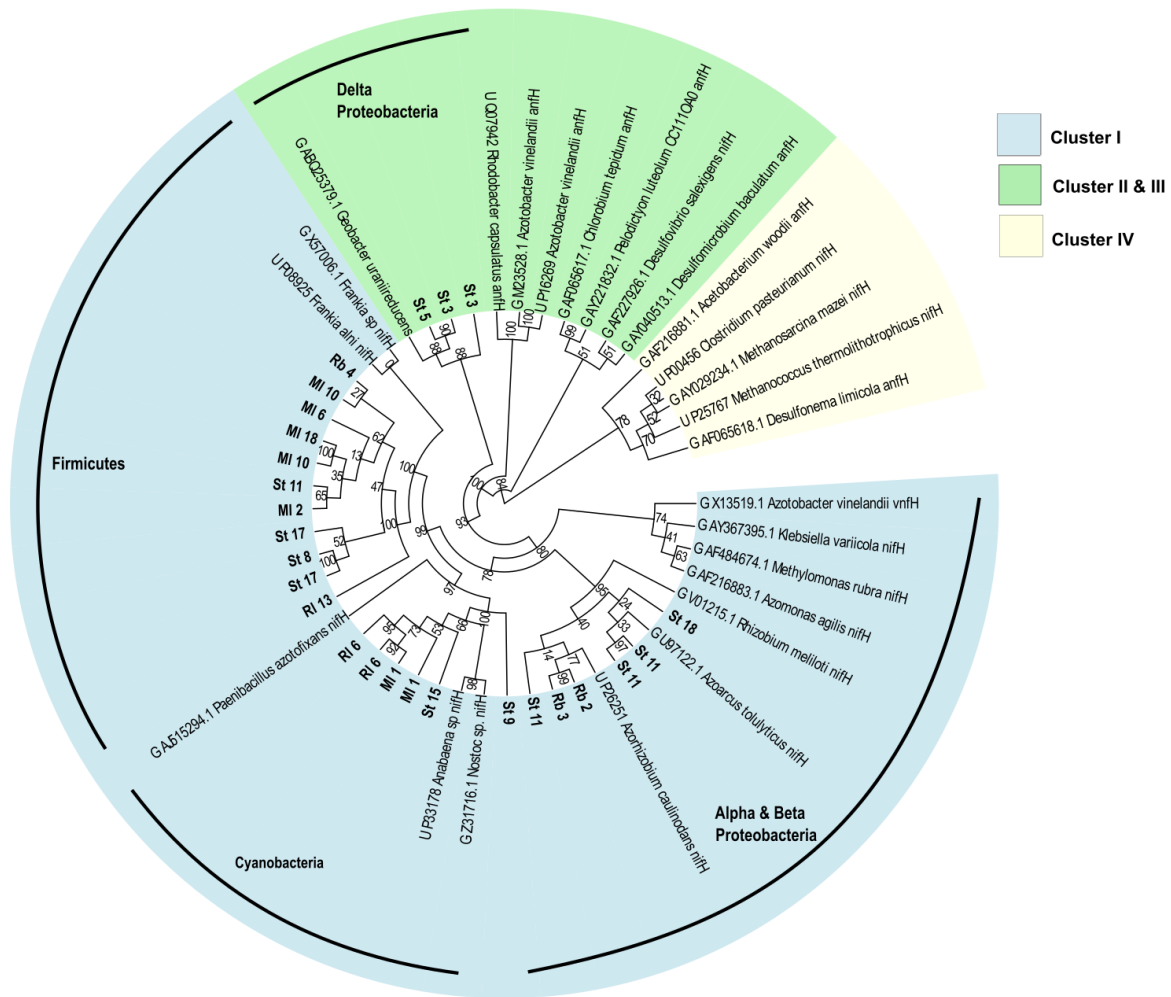
The results of mapping sequencing reads to the *nif* genes is provided in Table 3 (Appendix 2). This highlights the Alignment Score (AS), which indicates the alignment quality between reads and contigs (Table 3, Appendix 2). The number of *nif* genes for each score threshold is provided, alongside the percentage of reads with AS over 60. The Alignment Score ranges between 0 and the maximum length of the reads (0–100 for MI dataset and 0–150 for Rb, St and RI datasets). For each forefield, the percentage of alignments with an AS greater than 60

was  $1.06 \times 10^{-3}$  (MI),  $4.23 \times 10^{-5}$  (RI),  $2.38 \times 10^{-4}$  (Rb) and  $9.56 \times 10^{-4}$  (St). Plots of the normalised nif genes recovered and the number of reads aligning to genes with an AS over 60, for each metagenome, are available in Figure 5–8 (Appendix 2).

#### 3.3.4 Nitrogenase clusters

Our newly sampled bacteria were analysed and grouped with previously published relatives, as shown in Zehr *et al.*, (2003). Forefield sequences were distributed across Cluster I (23 sample sequences) and III (3 sample sequences), with no representatives in Cluster II or IV (Figure 3.2). Thus, 88.5% of sample sequences were attributed to Cluster I, which contains the typical Mo nifH, indicating the presence of plentiful molybdenum in soils for the nitrogenase cofactor (Zehr *et al.*, 2003).

Environmental samples in Cluster I included the groups Alphaproteobacteria, Betaproteobacteria, Cyanobacteria and Firmicutes (Figure 3.2). The first group, associated with Alphaproteobacteria and Betaproteobacteria, incorporated five environmental samples that clustered most closely with *Azorhizobium caulinodans* and *Azoarcus tolulyticus*. These are plant-associated diazotrophs, important for establishing stocks of fixed nitrogen for legume uptake, supporting plant growth (Hurek and Hurek 1995; Dreyfus, Garcia and Gillis 1988). The second group was comprised of six sample sequences, clustering with the Cyanobacteria, *Nostoc* and *Anabaena*, which are free living nitrogen fixers (Zehr *et al.*, 2003). Cyanobacteria have been proposed as crucial for building labile nitrogen pools in newly exposed soils, important for facilitating heterotroph colonisation, and have been identified in other forefields using 16s rRNA amplicon sequencing (Schmidt *et al.*, 2008; Duc *et al.*, 2009; Frey *et al.*, 2013). Group 3 contained 11 highly related sample nifH sequences, grouping closely to *Frankia*. This genus is composed of nitrogen-fixing bacteria that are symbionts of actinorhizal plant roots, and again provides evidence for bacterial support of plant growth and establishment, through supplies of fixed nitrogen (Benson and Silvester 1993). Whilst the forefields may have a low diversity of root symbiotic diazotrophs, this may also relate to sub-optimal cell lysis and separation of root-associated cells during the DNA extraction process, or that these organisms were at a low abundance and thus not captured through sequencing.



**Figure 3.2:** *nifH* maximum likelihood phylogeny of sample sequences (bold) and sequenced samples derived from NCBI GenBank and UniProtKB. Most sample sequences were obtained from the *nifH* phylogeny of Deslippe and Egger, (2006). For study samples, the Sample ID is given, corresponding to Appendix 2 Table 1. For sequenced samples, the database, organism name and gene are given. Bootstrap support values are given, based on 1000 tree iterations. The *nifH* clusters (derived from Zehr *et al.*, 2003) are denoted by leaf colours (Cluster I-IV). The tree is rooted on Cluster IV, as this group contains divergent ‘*nif*-like’ sequences (Zehr *et al.*, 2003). Key groups containing sample sequences are noted, including Firmicutes, Cyanobacteria, Alphaproteobacteria, Betaproteobacteria and Deltaproteobacteria.

Environmental samples were also present in Cluster III, which is attributed to a group of anaerobic bacteria (Zehr *et al.*, 2003). The three sample sequences clustered most closely to *Geobacter uraniireducens*, an anaerobe common in sediments under metal reducing conditions, capable of dissimilatory Fe(III) reduction (Shelobolina *et al.*, 2008). However, no sample sequences were linked to Cluster II, which is associated with organisms containing the alternative *anfH*, containing an Fe–Fe cofactor, used in the absence of molybdenum and Vanadium (Zehr *et al.*, 2003).

These results reflect those of Duc *et al.*, (2009), who used clone libraries to evaluate the phylogeny of diazotrophs across the Damma Glacier, Switzerland. Interestingly, *nifH* sequences from their analysis also grouped with nitrogenase Clusters I and III (Duc *et al.*, 2009). Additionally, genera identified by Duc *et al.*, (2009) included the key genera identified in this analysis, such as *Geobacter*, *Nostoc* and *Anabaena*, suggesting that these organisms are common across forefields (Duc *et al.*, 2009). The prevalence of these organisms may be due to adaptations or attributes to cold environments, such as cold or UV tolerance, and the release of protective exudates (Tamaru *et al.*, 2005; Chattopadhyay 2006; Pattanaik *et al.*, 2007). Cyanobacteria such as *Nostoc* have been shown to produce extracellular polysaccharides (EPS) which are important for desiccation and freeze-thaw tolerance in Arctic environments (Tamaru *et al.*, 2005). *Geobacter* are commonly found in anaerobic environments, and therefore may tolerate any anoxia in forefield soils created by frequent meltwater flooding and the formation of melt pools (Duc *et al.*, 2009). The consistent identification of *Geobacter*, *Nostoc* and *Frankia* in forefield soils using *nifH* analysis indicates that a core group of diazotrophs may be present across Arctic forefields. These diazotrophs may be important for facilitating plant colonisation and establishment, either by building labile pools in newly exposed soils (Cyanobacteria) or through symbiosis (*Frankia*, *Azorhizobium*).

Results from BLASTn searching each *nifH* sequence against cultured isolates revealed forefield sample sequences were divergent, with sequence identity ranging between 80%–95% (Table 3.6). This indicates that the diazotrophs present in the samples are novel compared to those which have been previously identified and may be unique or contain adaptations to cold oligotrophic forefield conditions. However, as less abundant organisms will contribute to a minor proportion of the unamplified sequenced DNA and *nifH* gene pool, using additional *nif* genes may help highlight the presence of rare organisms in samples (Cowan *et al.*, 2005). This may be especially helpful for metagenomes where sequencing coverage was not sufficient to profile the complete community structure, and thereby some low abundance organisms may not have been represented in the final dataset (Figure 1–4, Appendix 2).

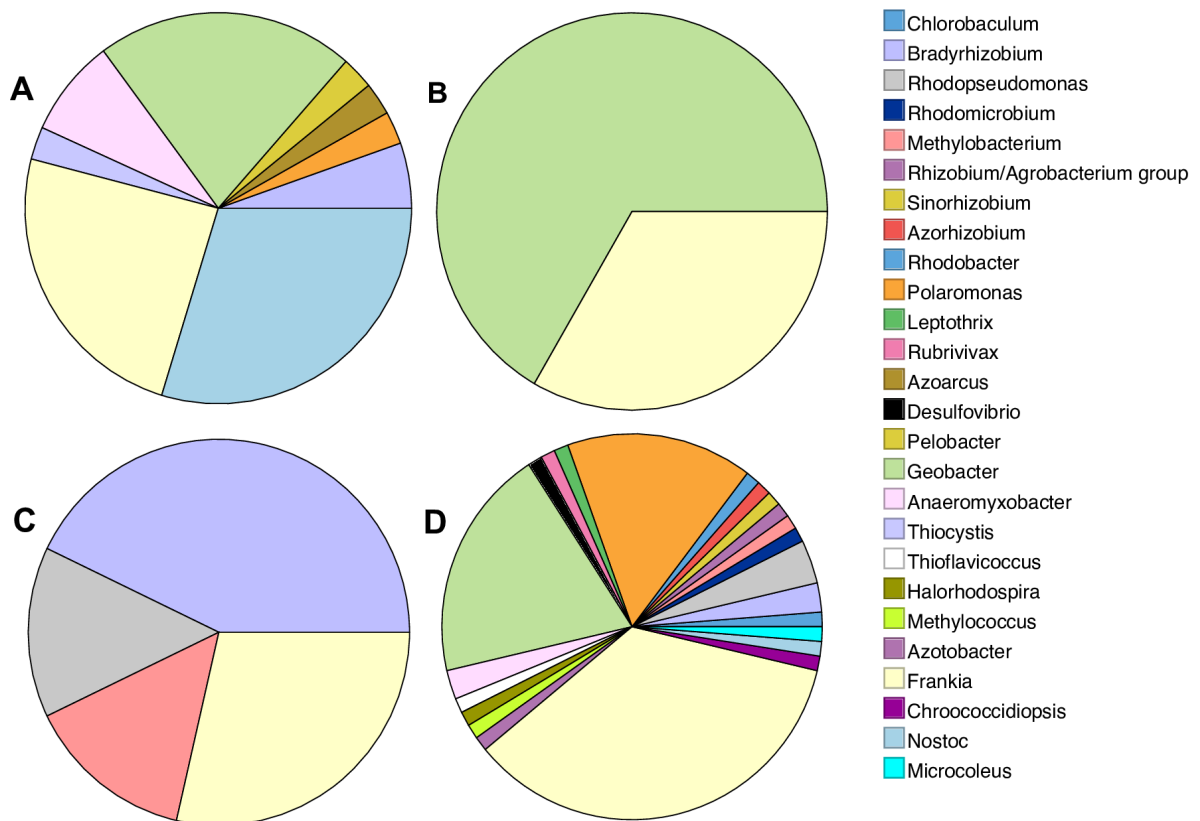
**Table 3.6:** NCBI blastn matches for sample sequences, against cultured isolates. The best match accession number and % identity is given. Sequences with no significant matches have been left blank.

Sample	Cluster	Blast match	Accession number	% similarity
Rb10	I	<i>Frankia</i> sp.	X57006.1	83
MI 10	I	<i>Frankia</i> sp.	CP000820.1	84
MI 6	I	<i>Frankia alni</i> str.	CT573213.2	82
MI 18	I	-	-	-
MI 10	I	<i>Frankia casuarinae</i> strain	CP000249.1	82
St 11	I	<i>Frankia</i> sp.	AY115490.2	82
MI 2	I	<i>Frankia alni</i> str.	CT573213.2	82
St 17	I	<i>Frankia</i> HRN18a	X17522.1	81
St 8	I	<i>Frankia</i> sp.	X73983.1	81
St 17	I	<i>Frankia</i> sp.	HM026362.1	81
RI13	I	<i>Leptosprillum ferriphilum</i>	JN390678.1	85
St3	III	<i>Geobacter uraniireducens</i>	CP000698.1	88
St3	III	<i>Geobacter lovleyi</i>	CP001089.1	90
St5	III	<i>Geobacter uraniireducens</i>	CP000698.1	89
RI 6	I	<i>Nostoc flagelliforme</i>	AP018269.1	95
RI 6	I	<i>Scytonema</i> sp.	AP018268.1	91
MI1	I	<i>Nostoc punctiforme</i>	CP001037.1	94
MI1	I	<i>Scytonema</i> sp.	AP018268.1	90
St15	I	<i>Scytonema</i> sp.	AP018268.1	84
St9	I	<i>Anabaena variabilis</i>	AP018216.1	84
Rb2	I	<i>Bradyrhizobium oligotrophicum</i>	AP012603.1	94
Rb3	I	<i>Bradyrhizobium oligotrophicum</i>	AP012603.1	91
St 11	I	<i>Polaromonas naphthalenivorans</i> CJ2	CP000529.1	87
St18	I	<i>Bradyrhizobium oligotrophicum</i> S58	AP012603.1	80
St11	I	<i>Polaromonas naphthalenivorans</i> CJ2	CP000529.1	94
St11	I	<i>Polaromonas naphthalenivorans</i> CJ2	CP000529.1	90

### 3.3.5 Diazotroph community structure

LCA analysis with multiple *nif* genes (HDKNE) identified the key organisms consistent between two or more forefields, including *Geobacter*, *Frankia* and *Nostoc*, which were also highlighted in the *nifH* analysis. Additional genera, for example *Polaromonas*, *Pelobacter* and *Microcoleus* were also identified here through the inclusion of additional *nif* genes (*nifDKNE*)

(Figure 3.3). This suggests including multiple nitrogenase genes provides a more holistic view of the diazotroph community structure in each forefield, due to the low copy number of these genes in unamplified samples. This is a particular issue of highly diverse metagenome samples, such as those from soils, as sequencing depth may not profile the complete community structure (Rodriguez and Konstantinidis 2014).



**Figure 3.3:** Taxonomic distribution of *nif* (HDKNE) genes for each forefield at the genus level: Midtre Lovénbreen MI (A), Russell RI (B), Rabots Rb (C), Storglaciären St (D). The total *nif* gene sequence count for each site was 42, 15, 13 and 91, respectively.

The assignment of *nif* genes in the RI forefield covers two key genera, *Geobacter* and *Frankia*. Limited research has been conducted into the presence of *Frankia* in Greenland; however, these organisms are typically associated with common actinorhizal plants (Benson and Silvester 1993; Chaia *et al.*, 2010). This group forms nitrogen-fixing root nodules with *Frankia* in exchange for reduced carbon and therefore are commonly found as early colonisers of undeveloped, oligotrophic soils (Wall 2000; Schwinter 2012). This is in agreement with the limited nitrogen content detected in this forefield, at 2.04 TN g<sup>-1</sup> (Figure

3.1 and Table 3.2). Additionally, the presence of plants has been identified as a key control on microbial community structure over the Damma Glacier forefield, Switzerland (Miniaci *et al.*, 2007). Furthermore, the identification of the anaerobic *Geobacter* indicates the presence of periodically saturated and anoxic conditions along the forefield, possibly attributed to meltwater flooding (Duc *et al.*, 2009). *Geobacter* are dissimilatory metal and sulfur reducing bacteria and have been proposed as key players in sediment nutrient cycles, oxidation of organic matter, bioremediation and soil gleying (Lovley *et al.*, 1993; Childers *et al.*, 2002; Methe *et al.*, 2003). *Geobacter* have been consistently identified across glacier forefield soils, which may relate to their metabolic diversity, thereby making these organisms well suited to fluctuating environmental conditions in forefield soils (Duc *et al.*, 2009; Edwards and Cook 2015; Rime *et al.*, 2015). This group has been shown to use chemotaxis to access Fe(III) oxides as an electron acceptor, which may explain their prevalence over other non-motile Fe(III) reducers (Hartmann and Brunner 2015). Whilst deeper sequencing in some metagenomes may highlight additional rare diazotrophic bacteria in RI samples, it is likely that *Geobacter* and *Frankia* were the most dominant nitrogen fixers present, as these were identified through direct sequencing of unamplified DNA (Cowan *et al.*, 2005; Figure 2, Appendix 2).

Similarly, to RI, the taxonomic diversity detected in the N-Swedish Rb forefield was largely comprised of root associated diazotrophs, including the genera *Bradyrhizobium*, *Frankia*, *Methylobacterium* and *Rhodopseudomonas* (Figure 3.3). This may relate to the lack of bare soil observed at this forefield, and therefore limited requirement for free living diazotrophs (Miniaci *et al.*, 2007). This site also had a low average soil nitrogen content, at 1.04 mg g<sup>-1</sup> (Figure 3.1 and Table 3.2), which, alongside the detection of *Rhizobia*, Fabaceae root-nodule symbionts, indicates that nitrogen limitation for plant growth may have been occurring in soils (Mylona *et al.*, 1995). Actinorhizal and legume plants, which directly benefit from biological nitrogen fixation through symbiosis, such as Clover, are likely to prevail in developing forefield soils (Fagerli and Svenning 2005; Chaia *et al.*, 2010). This is because they maintain a competitive advantage over other plants in nitrogen limited conditions, typical of newly exposed soils (Menge and Hedin 2009; Bradley, Singarayer and Anesio 2014). Additionally, Rb had a lower average soil TOC content than other forefields, at 6.8 mg g<sup>-1</sup> (Figure 3.1 and Table 3.2). Thus, *Rhizobia* are likely to benefit from symbiosis with plants through the supply of reduced carbon (Denison and Kiers 2004). Plants may therefore be acting as a control on the forefield microbial community structure, endorsing the presence of

root-associated diazotrophs (Miniaci *et al.*, 2007). Rarefaction curves for Rb sites were shown to be nearing saturation, indicating much of the microbial community structure was profiled (Figure 4, Appendix 2). Additional sequencing for these samples may reveal further low abundance taxa; however, it is likely that the most dominant fraction of diazotrophs have been identified adequately through our analysis.

The *nif* genes recovered from the MI forefield showed a wider taxonomic diversity of diazotrophs and contained sequences linked to the genera *Nostoc*, *Polaromonas*, *Bradyrhizobium*, *Pelobacter*, *Azoarcus* and *Anaeromyxobacter*. The presence of the Cyanobacteria, *Nostoc*, was expected due to the greater extent of bare soil observed in this forefield, enhancing the need for early colonisers (Frey *et al.*, 2013). Additionally, EPS production enables this group to resist harsh freeze-thaw cycles, common in Arctic environments (Tamaru *et al.*, 2005). Given the high latitude of this forefield, it is also not surprising to find *Polaromonas*, which are known psychrophiles (Irgens *et al.*, 1996). The presence of *Bradyrhizobium* and *Frankia* indicate plants may require additional fixed nitrogen through symbiosis, corresponding with the low nitrogen stocks detected (Benson and Silvester 1993; Mylona *et al.*, 1995; Chaia *et al.*, 2010; Figure 3.1 and Table 3.2). Additionally, the presence of legume symbiotic diazotrophs is interesting, as Fabaceae are non-native to Svalbard, having been introduced over the 20th Century (Fagerli and Svenning 2005). The absence of early plant colonisation in the forefield may also have been a control on overall microbial community structure, endorsing a range of non-symbiotic diazotrophs (Knelman *et al.*, 2012). Alongside *Geobacter*, the identification of *Pelobacter*, *Thiocystis* and *Anaeromyxobacter*, again indicates permanent or periodic anaerobic conditions in the glacier forefield, similarly to RI (Schink and Stieb, 1983; Sanford *et al.*, 2002). *Pelobacter* are anaerobic organisms containing diverse fermentative metabolisms, which may make this group well suited to the rapidly changing conditions in forefield soils (Schink, 2006). For example, *Pelobacter* have been shown to ferment acetylene using acetylene hydratase to acetate for cell growth or using nitrogenase to ethylene through nitrogen fixation (Akob *et al.*, 2017). The genomic results for the MI forefield falls in line with 16s amplicon data presented by Bradley *et al.*, (2016). This study also found *Frankia*, *Rhizobium*, *Nostoc* and *Geobacter* in the MI forefield (Bradley *et al.*, 2016). The identification of additional organisms such as *Devosia*, *Sphingomonas* and *Rhodoplanes* may relate to the use of amplification in their methodology, thereby aiding the discovery of low abundance organisms (Bradley *et al.*, 2016). Additionally, some metagenomes from this forefield would



have benefitted from greater sequencing depth in order to completely profile the microbial community composition (Figure 1, Appendix 2). Therefore, deep sequencing of these samples may reveal additional low abundance diazotrophs, unidentified in this analysis.

Finally, the St forefield contained sequences relating to *Nostoc*, *Geobacter*, *Rhizobium*, *Polaromonas* and *Frankia*, in line with the other forefields sampled (Figure 3.3). This supports the identification of a core group of diazotrophs present across Arctic glacier forefields. However, several diazotrophs detected at this site may also have importance in sulfur cycling, alongside nitrogen fixation (Figure 3.3). The detection of the anaerobic diazotrophs *Geobacter* and *Desulfovibrio* indicates the potential for sulfur reduction, whereby energy is gained through reducing sulfur (S) or sulfate ( $\text{SO}_4^{2-}$ ) to hydrogen sulfide ( $\text{H}_2\text{S}$ ), with the oxidation of organic carbon (Boopathy and Kulpa 1993; Caccavo *et al.*, 1994). However, inorganic S and  $\text{SO}_4$  have been found to be limiting for both plants and microbes in newly exposed glacier forefield soils (Allison *et al.*, 2007; Prietzel *et al.*, 2013). Nevertheless, desulfonating bacteria, whom metabolise organically bound sulfur to labile sulfates, have been found in forefield soils, and may therefore help overcome S limitation (Schmalenberger and Noll, 2009; Prietzel *et al.*, 2013). Additionally, suitable anaerobic growth conditions for sulfur reducing bacteria may occur frequently in stagnated proglacial meltwater pools and during periods of meltwater flushing (Duc *et al.*, 2009). Furthermore, the detection of organisms such as *Chlorobaculum*, *Thioflaviccoccus*, *Halorhodospira* and *Thiocystis* indicates the potential for St forefield bacteria to carry out both nitrogen fixation and sulfur oxidation (Figure 3.3). These organisms have the potential to oxidise  $\text{H}_2\text{S}$  to S and  $\text{SO}_4$ , alongside gaining fixed nitrogen through diazotrophy (Imhoff and Pfenning 2001; Chan, Morgan-Kiss and Hanson 2008; Peduzzi *et al.*, 2011; Challacombe *et al.*, 2013). The ability of these organisms to overcome nitrogen limitation through fixation, and to respire anaerobically in anoxic soils, may make this group well suited to harsh forefield environments. Additionally, as *Halorhodospira* is also halophilic, this may indicate resistance to high salinity environments, such as ice brine channels, or evaporation ponds in the St forefield (DasSarma and DasSarma 2006).

The diazotroph community composition observed using LCA nifHDKNE analysis was again largely consistent with those found at the Damma Glacier, Switzerland (Duc *et al.*, 2009; Frey *et al.*, 2013). This includes genera such as *Methylobacterium*, *Bradyrhizobium*, *Azotobacter*, *Anabaena*, *Nostoc* and *Geobacter* (Duc *et al.*, 2009). This supports the results from the nifH phylogeny, indicating the presence of

consistent genera across forefields, which may be well adapted to the cold, oligotrophic and high UV conditions. Plant colonisation has also been identified as an influence on the diazotrophic community composition, in agreement with studies on the Damma Glacier, Switzerland (Miniaci *et al.*, 2007; Duc *et al.*, 2009; Zumsteg *et al.*, 2013). However, it is important to acknowledge that additional factors, such as latitude, bedrock mineralogy, organic matter and aeolian nitrogen deposition, may also have an influence on diazotroph community structure and abundance (Duc *et al.*, 2009; Zumsteg *et al.*, 2013). Some genera found by Duc *et al.*, (2009), such as *Oscillatoria*, *Ideonella* and *Paenibacillus* were not identified in this study (Figure 3.3). This may relate to the absence of these organisms in the four forefields in this analysis, but also may relate to the alternate approach used. As this analysis uses unamplified *nifH* sequences, some low abundance organisms may not be sequenced due to incomplete sequencing depth in highly complex samples (Rodriguez and Konstantinos 2014; Figures 1–4, Appendix 2). Thus, it cannot be ruled out that these organisms were also not present in the forefields, but at a lower abundance than those captured by the sequencing effort (Prakash and Taylor, 2012). In order to profile the complete community of some metagenomes, including low abundance organisms, deeper sequencing would be required, due to the diverse nature of soil samples (Rodriguez and Konstantinos 2014). Despite this, this analysis has been able to capture a diverse group of diazotrophs that appear to be common across glacier forefields and are likely the most abundant fraction of the nitrogen-fixing community, as these were captured by unamplified DNA sequencing (Rodriguez and Konstantinos 2014).

### **3.4 Conclusions**

Overall, this study has used metagenomics to understand the diversity of diazotrophs across four Arctic glacier forefields. The results of Chapter 2 were applied to this analysis, to assemble metagenome contigs for interpretation. The subsequent analysis used a *nifH* phylogeny to identify a key group of diazotrophs across four Arctic forefields, associated with both Cluster I and III nitrogenase, linked to aerobic and anaerobic organisms containing the typical Mo *nifH* (Zehr *et al.*, 2003). Incorporating multiple *nif* genes (HDKNE) revealed additional organisms from unamplified metagenome samples, compared to using the *nifH* gene exclusively. This may relate to the complex nature of soil metagenome samples, whereby sequencing depth is not always adequate to profile the complete microbial community diversity. Thus, to reveal all low abundance diazotrophs, some metagenomes

would require additional deep sequencing. Key diazotrophs were found to be metabolically diverse, including genera such as *Geobacter*, *Frankia*, *Nostoc*, *Polaromonas* and *Bradyrhizobium*. A range of diazotrophic organisms outside the key group were also highlighted, including halophiles, psychrophiles and bacteria associated with fermentative metabolisms and sulfur cycling. Therefore, this analysis has shown a diverse group of diazotrophic bacteria present in Arctic forefield soils, including a consistent core subset. These diazotrophs have the potential to build labile nitrogen stocks in forefield soils, which may support further colonisation and soil development.

### **3.5 Limitations and Future work**

Metagenomics was applied in this study to investigate the diversity of diazotrophic bacteria across four Arctic glacier forefields. However, metagenomics is unable to provide evidence for the activity of microbial nitrogen fixation, as it focuses on providing genomic verification for the potential of the pathway (Wooley *et al.*, 2010). To clarify if the diazotrophs recovered were active, transcriptomics could be used. Transcriptomics involves the sequencing of microbial community RNA, highlighting the function of actively expressed genes (Wang *et al.*, 2009). However, this technique only provides a snapshot of the expression at the time of sampling, so cannot be extrapolated to year-round activity (Lowe *et al.*, 2017). It would also be interesting to investigate the rate of nitrogen fixation in each forefield and if there are any differences between sites. To do this, gasometric incubations of N<sub>2</sub> uptake could be implemented using forefield soil samples, for example utilising acetylene reduction assays, or alternatively utilizing a radio tracer for N<sub>2</sub> incorporation, such as <sup>15</sup>N<sub>2</sub> (Burriss *et al.*, 1972; Hardy *et al.*, 1973). This would highlight how active community nitrogen fixation was and how this varies over the chronosequences or between forefields.

Furthermore, as shown by the rarefaction analysis (Figures 1-4, Appendix 2) the sequencing coverage was not sufficient to profile the complete microbial community in all forefield samples. The recovered diazotrophs are likely to be those which were most abundant in the samples, as they would account for a greater portion of the sequenced DNA (Rodriguez and Konstantinidis 2014). To reveal less abundant diazotrophic organisms, additional deep sequencing could be carried out. Not only would this provide a more complete picture of the microbial community, but the enhanced sequencing coverage may enable more detailed analyses, for example extracting complete genomes from the metagenomes (Sharon and Banfield, 2013; Eloë-Fadrosch *et al.*, 2016b). This would allow a more refined investigation of

the microbial genomes, for example, by providing genomic evidence for cold, desiccation or UV tolerance adaptations which enable these microbes to survive in Arctic environments (Christmas *et al.*, 2016).

## Footnotes

- 1 <https://www.phylo.org/>
- 2 <http://tree.bio.ed.ac.uk/software/figtree/>
- 3 <http://www.evolgenius.info/evolview/>
- 4 <https://inkscape.org/en/>
- 5 <https://blast.ncbi.nlm.nih.gov/>

# Chapter 4: Microbial community development over a gradient of soil succession along the Midtre Lovénbreen glacier forefield, Svalbard

Maisie V. Nash<sup>1</sup>; Alexandre M. Anesio<sup>2</sup>; Gary Barker<sup>3</sup>; Martyn Tranter<sup>1</sup>, Thomas Turpin-Jelfs<sup>1</sup>, Liane G. Benning<sup>4,5,6</sup> and Patricia Sánchez-Baracaldo<sup>1</sup>

<sup>1</sup>School of Geographical Sciences, University of Bristol, BS8 1SS, UK

<sup>2</sup>Department of Environmental Science, Aarhus University, 4000-Roskilde, Denmark

<sup>3</sup>School of Life Sciences, University of Bristol, BS8 1TQ, UK

<sup>4</sup>GFZ German Research Centre for Geosciences, Telegrafenberg, 14473 Potsdam, Germany

<sup>5</sup>School of Earth and Environment, University of Leeds, LS2 9JT, Leeds, UK

<sup>6</sup>Department of Earth Sciences, Free University of Berlin, Malteserstr, 74-100, Building A, 12249, Berlin, Germany

## Contributions and acknowledgements

MN carried out the analysis, interpretation and formulation of the manuscript. AA, MT and LB assisted with fieldwork logistics and sampling. TTJ supplied the raw TN and TOC data. PSB and AA helped refine drafts of the manuscript. Additional support came from James Bradley and Katherine Wright during fieldwork sampling.

## Funding

This work was carried out as part of the NERC GW4+ Doctoral Training Partnership, awarded to MN. Funding support was provided through NERC grants NE/J02399X/1, awarded to AA and NE/J022365/1 awarded to LGB. Funding support for PS-B came from a Royal Society University Research Fellowship and LGB also acknowledges the financial support through the Helmholtz Recruiting Initiative.

## 4.1 Introduction

Arctic glaciers face continued ice retreat with warming global temperatures (Zemp *et al.*, 2015). As the terminus of a glacier retreats, undeveloped soils are exposed which were previously shielded by ice cover (Bradley *et al.*, 2014). A transect of soil succession can often be identified in glacier forefields, with newly exposed bare soil close to the glacier terminus and more developed plant colonized soils moving outwards (as reviewed by Bradley *et al.*, 2014). These soils pose an interesting opportunity to understand how land is first colonized, the changes to microbe and plant communities during soil succession and the impact these communities have on soil structure and physicochemical characteristics during development (Edwards and Cook, 2015). Investigating microbial communities during forefield soil succession is important to understand Arctic microbial diversity and the role of microbial

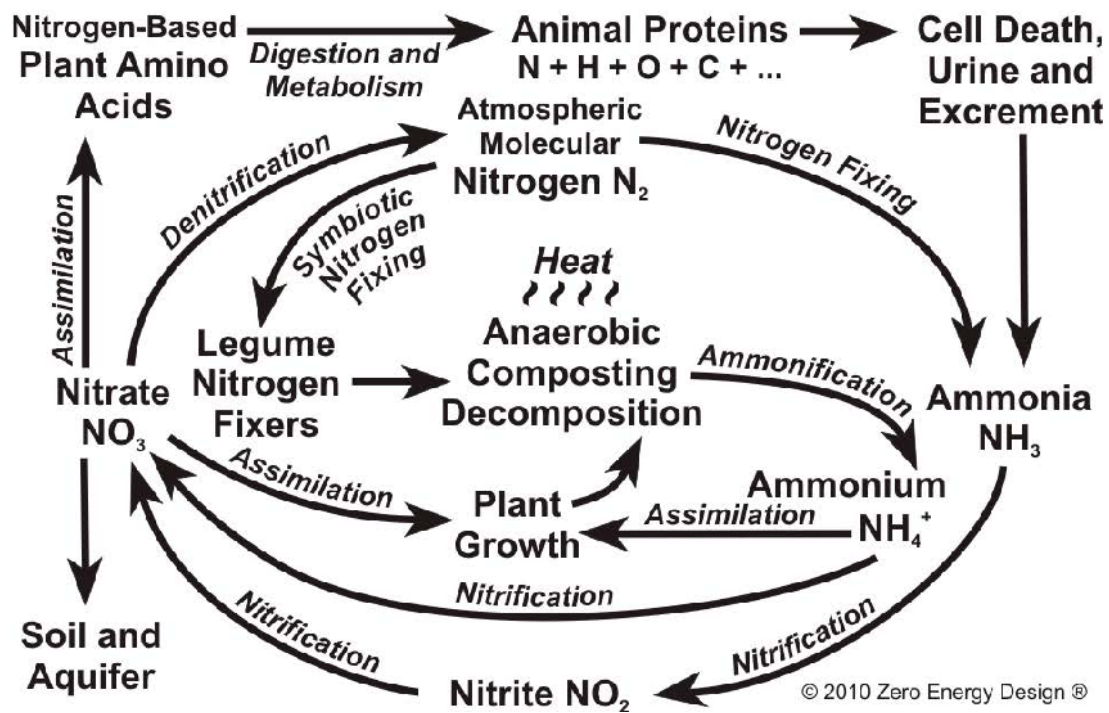
communities in soil biogeochemical cycles (Bradley *et al.*, 2014). Additionally, this information will help improve our understanding of how Arctic microbial ecology and nutrient cycles may modify in the future, given the continued retreat of glaciers with global warming (Edwards and Cook, 2015).

A chronosequence based approach can be applied to study soil succession in glacier forefields, employing a space for time substitution, with older soils at a greater distance from the glacier terminus (Tscherko *et al.*, 2003; Bradley *et al.*, 2014). Thus far, the majority of forefield chronosequence studies have focused on plant communities and soil structure during succession (Ohtonen *et al.*, 1999; Strauss *et al.*, 2009; Knelman *et al.*, 2012). A general increase in organic carbon and nitrogen has been shown during succession, alongside a reduction in pH, as soils become colonised by plants (Ohtonen *et al.*, 1999; Strauss *et al.*, 2009; Knelman *et al.*, 2012; Turpin-Jelfs *et al.*, 2019). Microbial community colonisation and development in forefields is important for influencing soil nutrient cycles, physicochemical status and structure (Tscherko *et al.*, 2003; Hahn and Quideau, 2013). Microbial communities are typically the initial colonisers of newly exposed soils and thus facilitate soil development (Chapin *et al.*, 1994; Hahn and Quideau, 2013). In particular, phototrophic and diazotrophic microbes are important for building labile carbon and nitrogen stocks, facilitating the establishment of heterotrophic microbial populations and plants (Chapin *et al.*, 1994; Hahn and Quideau, 2013). Microbial carbon and nitrogen fixation is particularly important following glacier retreat, as newly exposed soils are often oligotrophic (Chapin *et al.*, 1994). Additionally, microbes help to stabilise soils for plant colonisation through the secretion of extracellular polymeric substances (EPS) during growth (Magner and Thomas, 2011).

The initial coloniser microbial communities in glacial forefields have attracted substantial research, as these microbes are key facilitators for further soil development. However, the composition of these initial communities remains unclear (Kastovska *et al.*, 2005). On one hand, diazotrophic and phototrophic bacteria may constitute early colonisers, as these microbes can fix carbon and nitrogen in nutrient depleted conditions, subsequently facilitating labile nutrient pools for heterotroph colonisation (Hodkinson *et al.*, 2002). However, heterotrophic microbes may also be present in initial soils given the availability of organic carbon, which may be derived from ancient organic matter in overridden soils or from aeolian deposition (Kastovska *et al.*, 2005; Bardgett *et al.*, 2007; Bradley *et al.*, 2014). The pioneer microbial community may be derived from ice surface microbes, aeolian deposition, or legacy subglacial microbial communities (Anesio *et al.*, 2009; Boyd *et al.*, 2010). Furthermore, the composition of forefield microbial communities has been shown to modify with soil succession, related to changes in environmental factors such as nutrients, organic matter, water flow

pathways and pH (Tscherko *et al.*, 2003; Sakata Bekku *et al.*, 2004; Frey *et al.*, 2013). Additionally, research by Knelman *et al.*, (2012) identified that plant colonisation (alongside the type of plant), as having a key influence on microbial community structure during succession, as opposed to soil pH and nutrient status.

Nutrient and organic matter inputs to forefields may be from both allochthonous (glacial runoff, animal droppings and aeolian deposition) and autochthonous sources (*in situ* microbial activity) (Bradley *et al.*, 2014). Bioavailable nitrogen is needed by both plants and microbes for protein synthesis (Tresder, 2008). Nitrogen is supplied to forefields by diazotrophic nitrogen fixation, aeolian deposition, degradation of organic matter and in washing from snowmelt (Bradley *et al.*, 2014; Bradley *et al.*, 2015; Duc *et al.*, 2009). Nitrogen fixation is important in newly exposed soils, especially when the total nitrogen content is initially low (Duc *et al.*, 2009; Turpin-Jelfs *et al.*, 2019). Denitrification has been shown to increase in line with soil succession, attributed to the increase in water logged soils following plant colonization (Kandeler *et al.*, 2006; Schulz *et al.*, 2011). Additionally, nitrification has been shown to increase with soil age, due to a greater availability of bioavailable nitrogen in older soils (Brankatschk *et al.*, 2011). Nitrogen mineralization from organic matter is an important mechanism for recycling bioavailable nitrogen in forefield soils (Brankatschk *et al.*, 2011). This process converts organic nitrogen from cell death or excrement to labile ammonium (Figure 4.1). A schematic identifying the key processes involved in the soil nitrogen cycle is available in Figure 4.1



**Figure 4.1:** The soil nitrogen cycle. Sourced from Abatenh *et al.*, (2018)

Furthermore, organic carbon is a crucial nutrient for microbial growth and forms the backbone of molecules. Organic carbon can be supplied through aeolian deposition, carbon fixation and may also be present in the form of ancient organic matter in overridden soils (Hodkinson *et al.*, 2002; Guelland *et al.*, 2013). During the course of succession, organic carbon stocks increase in soils, particularly following plant colonization, whereby plant litter can be recycled by microbial communities back into the soil carbon pool (Knelman *et al.*, 2012; Zumsteg *et al.*, 2013). In early soils, phototrophic bacteria such as cyanobacteria can be crucial, as these microbes fix carbon dioxide into labile organic carbon (Strauss *et al.*, 2012). These microbes may therefore be key facilitators of heterotrophic colonization, as heterotrophs rely on the availability of labile fixed carbon (Freeman *et al.*, 2009; Zumsteg *et al.*, 2013). Finally, sulfur may be cycled in forefield soils through microbial redox reactions. Energy can be produced by chemolithotrophic bacteria through the oxidation of sulphur, coupled to the reduction of carbon or nitrate (Koltz *et al.*, 2011, Wainwright 1978). Sulfate can also be reduced when used as a terminal electron acceptor for the oxidation of organic matter, in the absence of oxygen, forming hydrogen sulfide (Widdel and Hansen 1992). Isotopic evidence for microbial sulfate reduction has been found in Arctic subglacial systems (Wadham *et al.*, 2004) and therefore may continue in soils following ice retreat.



The Damma Glacier, Switzerland, has often been used as a study site to investigate soil, microbial and plant succession (Bernasconi *et al.*, 2008; Duc *et al.*, 2009; Frey *et al.*, 2010; Bernasconi *et al.*, 2011; Brunner *et al.*, 2011). Comparatively, less research is available for the Midtre Lovénbreen forefield, Svalbard (Kastovska *et al.*, 2005; Schutte *et al.*, 2009; Bradley *et al.*, 2016; Nash *et al.*, 2018). Bradley *et al.*, (2016) combined field and laboratory data from the Midtre Lovénbreen glacier with the forefield microbial model SHIMMER 1.0, to help inform parameter values (Bradley *et al.*, 2016). The model results show that microbial biomass developed over the chronosequence and microbial activity was important for nutrient fixation and recycling in soils (Bradley *et al.*, 2016). Consequently, microbial communities may be important facilitators for plant colonization in these oligotrophic forefield soils (Bradley *et al.*, 2016). Furthermore, the microbial community of five Svalbard Glaciers, including Midtre Lovénbreen was investigated by Kastovska *et al.*, (2005). Both cryoconite and barren soils were found to contain the highest abundance of cyanobacteria, however cyanobacteria were also found in vegetated sites. Schutte *et al.*, (2009) found different microbial community compositions in surface and mineral depth soils, alongside a difference between newly exposed and more developed soils in the Midtre Lovénbreen forefield. However, this study also exemplified how forefield disturbances, such as water flow pathways may disrupt the pattern of succession (Schutte *et al.*, 2009). Finally, the Midtre Lovénbreen forefield was used in Chapter 3 of this thesis to investigate the diversity of diazotrophs between four Arctic glacier forefields using functional genes for nitrogen fixation (*nif* genes) (Nash *et al.*, 2018). Diazotrophs were identified in the Midtre Lovénbreen forefield and included cyanobacteria (*Nostoc*), root associated bacteria (*Rhizobia*, *Frankia*) and anaerobes (*Geobacter*) (Nash *et al.*, 2018).

Thus far, the majority of research on microbial communities in glacial forefields has used 16s rRNA sequencing or targeted functional gene amplification to investigate microbial community taxonomy (Sigler and Zeyer, 2002; Nemergut *et al.*, 2007; Knelman *et al.*, 2012; Zumsteg *et al.*, 2013). Metagenomics involves unamplified sequencing of the microbial community gene pool, providing information on both the taxonomy and functional genes present (Wooley *et al.*, 2010). Metagenomics has yet to be applied to understanding microbial community development during soil succession. Metagenomic sequencing would be a useful contribution to the current body of literature, as it would provide more information on microbial community functional potential and their role in soil biogeochemical cycles during succession. Furthermore, genome binning is a method which can be used to extract single draft genomes from community metagenome data (Albertsen *et al.*, 2013; Nielsen *et al.*, 2014). This technique is beneficial as it allows draft genomes to be extracted without the need for culturing (Kunin *et al.*, 2008; Albertsen *et al.*, 2013). The completeness of the draft genome can vary

depending on the depth of sequencing and the complexity of the microbial community under investigation (Kunin *et al.*, 2008; Albertsen *et al.*, 2013). This means that the gene content, functional pathways and taxonomy of genomes of interest can be investigated further. This technique has yet to be applied to forefield samples and would be beneficial to explore the novelty of genomes in this harsh environment, which is not possible with 16s rRNA analysis.

This study aims to use metagenomics to explore the microbial community composition along a chronosequence of soil succession in the Midtre Lovénbreen forefield, Svalbard. In particular, we hypothesize that the microbial community composition, function and metabolic pathways modify during succession, in line with soil development, the establishment of labile nutrient pools and plants. We hypothesize that autotrophic carbon and nitrogen-fixing bacteria will be prevalent in low nutrient early soils, with heterotrophic microbes dominant in older soils with established nutrient pools. We expect carbon and nitrogen fixation to be detected in recently deglaciated soils, with pathways such as denitrification, carbon remineralization and sulfate reduction to occur in line with nutrient pool development. Additionally, we hypothesize that extremophilic, well adapted or unique genomes may be found in the metagenomes, related to the harsh environmental conditions found in forefield soils. Whilst Chapter 3 looked at the broad diversity of diazotrophs between forefields, this analysis investigates whole community changes within a forefield during succession. In addition, this work follows on from Chapter 2, as the metagenome assembler, metaSPAdes, is applied for functional analysis. The results of this analysis aim to contribute to the current body of literature on forefield microbial communities, providing additional evidence for community functional potential, which has not previously been explored in the context of succession.

## **4.2 Methodology**

### **4.2.1 Field Sampling**

Surface sediment was sampled along the forefield of Midtre Lovénbreen glacier, Svalbard, in July 2013. Additional sampling of cryoconite sediment and dark ice from the glacier surface was carried out for comparison. Chronosequence based sampling of the forefield was implemented, whereby sediment was obtained from three parallel transects, moving out from the glacier terminus (Bradley *et al.*, 2014; Nash *et al.*, 2018). This chronosequence sampling technique was applied to capture the changes in microbial community composition and nutrient content with soil development along the forefield. Additionally, this sampling approach enabled the sites to be dated from their year of exposure following glacier retreat (Brankatschk *et al.*, 2011; Bradley *et al.*, 2014; Bradley *et al.*, 2015). Bulk sediment was sampled in triplicate

from each site into Whirlpak™ bags, and frozen at -20° C prior to analysis. The sampling site locations, transect numbers (T1-T3) and associated soil age are provided in Table 4.1, with sites ranging from 0 – 2000 years since glacier retreat (Brankatschk *et al.*, 2011; Bradley *et al.*, 2015).

**Table 4.1:** Sampling site locations and associated soil exposure age (measured in years since glacier retreat), in line with Brankatschk *et al.*, (2011) and Bradley *et al.*, (2015). Samples obtained from the glacier surface are marked with “\*”.

Sample Number	Latitude / Longitude	Transect number (T1 - T3)	Estimated soil age (Years)
1	79.101 / 12.156	T2	0
2	79.112 / 12.175	T2	3
3	79.112 / 12.258	T3	3
4	79.118 / 12.094	T1	5
5	79.114 / 12.196	T2	5
6	79.104 / 12.279	T3	5
7	79.152 / 12.216	T1	29
8	79.151 / 12.254	T2	29
9	79.141 / 12.090	T3	29
10	78.928 / 12.254	T1	50
11	78.927 / 12.077	T2	50
12	78.908 / 12.164	T3	50
13	78.901 / 12.076	T2	50 – 113
14	78.901 / 12.076	T2	50 – 113
15	78.901 / 12.076	T2	50 - 113
16	78.990 / 12.083	T2	113
17	78.992 / 12.230	T2	113
18	78.979 / 12.332	T2	113
19	79.768 / 12.144	T2	2000
20	79.768 / 12.144	T2	2000
21	79.768 / 12.144	T2	2000
22	79.484 / 12.092	N/A	*Cryoconite hole
23	79.484 / 12.092	N/A	*Dark Ice

#### 4.2.2 DNA extraction, library preparation and sequencing

In total, 23 samples were selected for DNA extraction and metagenome sequencing (Table 4.1), spanning the range of soil ages. DNA from bulk soil samples was extracted in line with

the methodology detailed in Chapter 3, section 3.2.3 (Nash *et al.*, 2018). Metagenomes were sequenced using an Illumina-Mi Seq, with a TruSeq library prep kit at the University of Bristol Genomics facility. The 23 selected sites were sequenced for metagenomics, providing 2x 100bp paired-end reads. Sequencing read output for each site ranged between 4013376 – 66567072 reads per metagenome (Table 4.2).

**Table 4.2:** Statistics for unassembled and assembled metagenomes, including contiguity and completeness metrics. The % Coverage is calculated using the percentage of raw reads mapped back to the assembly.

Sample	Unassembled Reads	Assembled Contigs	Contigs > 1000bp	Largest Contig	% Coverage
1	18800440	58428	445	52064	23.2
2	23227446	15534	18	43675	3.7
3	17611153	34535	154	314188	16.0
4	20076741	14007	7	46511	2.9
5	22256997	26820	31	32120	6.4
6	21156969	33757	91	51529	7.6
7	22729904	19399	199	106118	4.6
8	20076740	14007	7	46511	2.9
9	22977962	15533	19	43675	2.7
10	57959962	60667	221	150624	5.1
11	38013530	22180	3	67008	1.9
12	38014196	22214	3	67008	1.9
13	4013376	851	0	8102	1.5
14	22329785	13861	38	1072806	3.7
15	22371213	11373	16	19032	2.4
16	22371175	11374	16	19032	2.4
17	38409045	19556	296	256358	3.5
18	66567072	75376	423	768884	6.9
19	23966913	5931	0	4646	0.9
20	20702648	4905	0	5800	0.9
21	29107763	14288	0	8822	1.6
22	19259216	89619	878	371579	26.6
23	22255419	26820	31	32120	9.5

#### 4.2.3 Rarefaction curves

Rarefaction curves for quality trimmed sequencing reads were created in MG RAST 4.0.3 (Glass *et al.*, 2010; Figure 4.2). Rarefaction curves display the number of species obtained with increasing sequencing effort (number of reads). Curves which reach saturation (i.e. more sequencing effort does not increase the species count) highlight metagenomes which profile the microbial diversity sufficiently (Rodriguez and Konstantinidis, 2014). Metagenomes which

are under saturated profile the more abundant fractions of the microbial community, however may not sample those which are less abundant. Under saturated rarefaction curves are common when investigating complex microbial communities, such as those in soil, as the sequencing depth may not be enough to recover the entire microbial community composition (Rodriguez and Konstantinidis, 2014). In these cases, additional deep re-sequencing would be useful if the less abundant microbial fractions were of particular interest.

#### 4.2.4 Metagenome assembly and mapping

The 23 metagenomes were quality trimmed and subsequently assembled with metaSPAdes v3.11.1 (Nurk *et al.*, 2017). The size and contiguity of the assembled metagenomes were evaluated using QUAST on the KBASE platform (Arkin *et al.*, 2016). Raw reads were mapped back to the assemblies using Bowtie 2 v2.3.2, to identify the proportion of the input reads which were incorporated into the final assemblies (Langmead and Salzberg, 2012). Assembly size for each metagenome ranged between 851 – 89619 contigs and read coverage ranged between 0.9 – 23.2 % (Table 4.2). These read coverage scores mean that over 75% of reads in each sample were not incorporated in to the assembled metagenome. The low coverage scores may be attributed to the complex community structure, combined with a sequencing depth which was insufficient to fully profile all the organisms. Combined, these can result in short fragmented assemblies, as the assembler cannot resolve repeat regions or insertion/deletions (INDELs).

#### 4.2.5 Taxonomic annotation

Read-based taxonomic annotation for each metagenome was carried out in Kaiju v1.5.0 (Menzel *et al.*, 2016). Read based annotation was used, as opposed to assembly annotation, due to the low read recruitment in assemblies (Table 4.2). Consequently, using trimmed reads enabled more of the dataset to be annotated, which is particularly important when aiming to gain an overview of the complete community composition. Kaiju is a kmer based taxonomic classifier, based on protein sequences (Menzel *et al.*, 2016). This classifier is suited to novel metagenome samples and those with sequencing errors, as sequence conservation is generally greater in protein sequences than the corresponding DNA (Menzel *et al.*, 2016).

#### 4.2.6 Metagenome binning

Genome binning was applied to recover draft microbial genomes from the metagenomes, using MaxBin2 v2.2.4 (Bankevich *et al.*, 2012; Wu *et al.*, 2015). This binning algorithm uses

an expectation- maximum approach, based on contig tetra nucleotide frequencies, to recover draft genomes (Wu *et al.*, 2015). MaxBin2 employs the assembled metagenomes and coverage information from raw reads to group contigs into discrete genome bins. In this instance, using assembled metagenomes was beneficial to provide longer sequences for tetra nucleotide frequencies to be assessed during binning and subsequent gene annotation. BinUtil v1.0.1 was used to extract the genomes from the metagenomes using the KBASE platform (Arkin *et al.*, 2016). Each genome bin was functionally annotated using RAST v0.0.12 on KBASE (Aziz *et al.*, 2008; Overbeek *et al.*, 2013; Arkin *et al.*, 2016). Features were predicted using glimmer3 and prodigal in RAST, before matching to the SEED ontology (Aziz *et al.*, 2008; Overbeek *et al.*, 2013). For each sampling site, extracted genomes were pooled and a function profile was created in KBASE, based on the SEED Ontology (Overbeek *et al.*, 2013; Arkin *et al.*, 2016). The SEED categories relating to nitrogen cycling (nitrogen fixation, nitrification and denitrification) were selected for analysis, given the importance of the nutrient for microbial growth and the limited nitrogen stocks recorded (Table 4.3).

Each genome bin was assigned a taxonomy using BLASTn against NCBI RefSeq ( $e 10^{-5}$ ). For each sampling site, the genome bins were aligned to closely related genomes on the KBASE platform, using Species Tree Builder, based on a subset of COG groups (Appendix 3 Table 1). The alignments were used to generate a phylogenetic tree for each site, using Fast Tree 2.2.10 (Aziz *et al.*, 2008). Given the incomplete sequencing depth and low assembly read recruitment, we do not aim to gain any near complete draft genomes or sample all genomes present. However, genome binning can provide an insight into the functional potential of some of the most abundant genomes in the forefield. Additionally, this exemplified the potential of metagenomics to reveal genome specific functions, given the availability of deep sequencing.

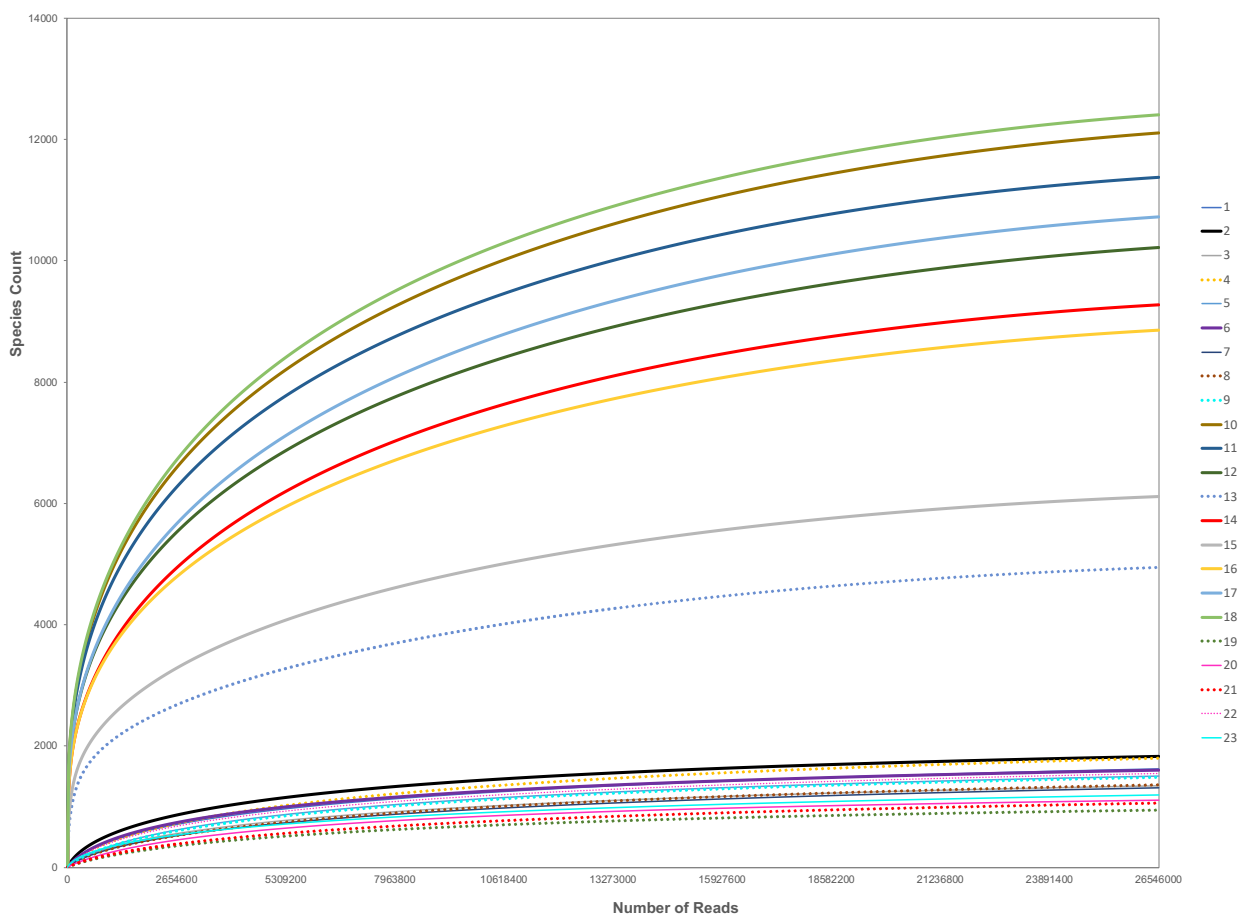
#### 4.2.7 Soil metadata

Soil pH and temperature for each of the sampling sites was obtained using a Hanna pH meter. Soil total nitrogen (TN) and total organic carbon (TOC) for each of the 23 sediment samples was evaluated using the protocol detailed in Chapter 3, section 3.2.2 (Nash *et al.*, 2018).

## 4.3 Results

### 4.3.1 Rarefaction curves

Rarefaction curves for each of the 23 metagenomes are displayed in Figure 4.2. Mid transect sites (50-113 years) display more under saturation than earlier sites, which could be related to an increase in species count and diversity but no increase in sequencing effort (Figure 4.2). Additional sequencing of these samples would be beneficial to recover the less abundant microorganisms. However, no samples were shown to be severely under sequenced and therefore we can be confident in profiling the most abundant constituents of the microbial communities for each site.



**Figure 4.2:** Rarefaction curves for the 23 metagenomes sampled from Midtre Lovénbreen, Svalbard. The curves display the number of species with increasing sampling (or number of reads). Curves that reach saturation display an adequate sequencing depth to profile the microbial community.

### 4.3.2 Soil metadata

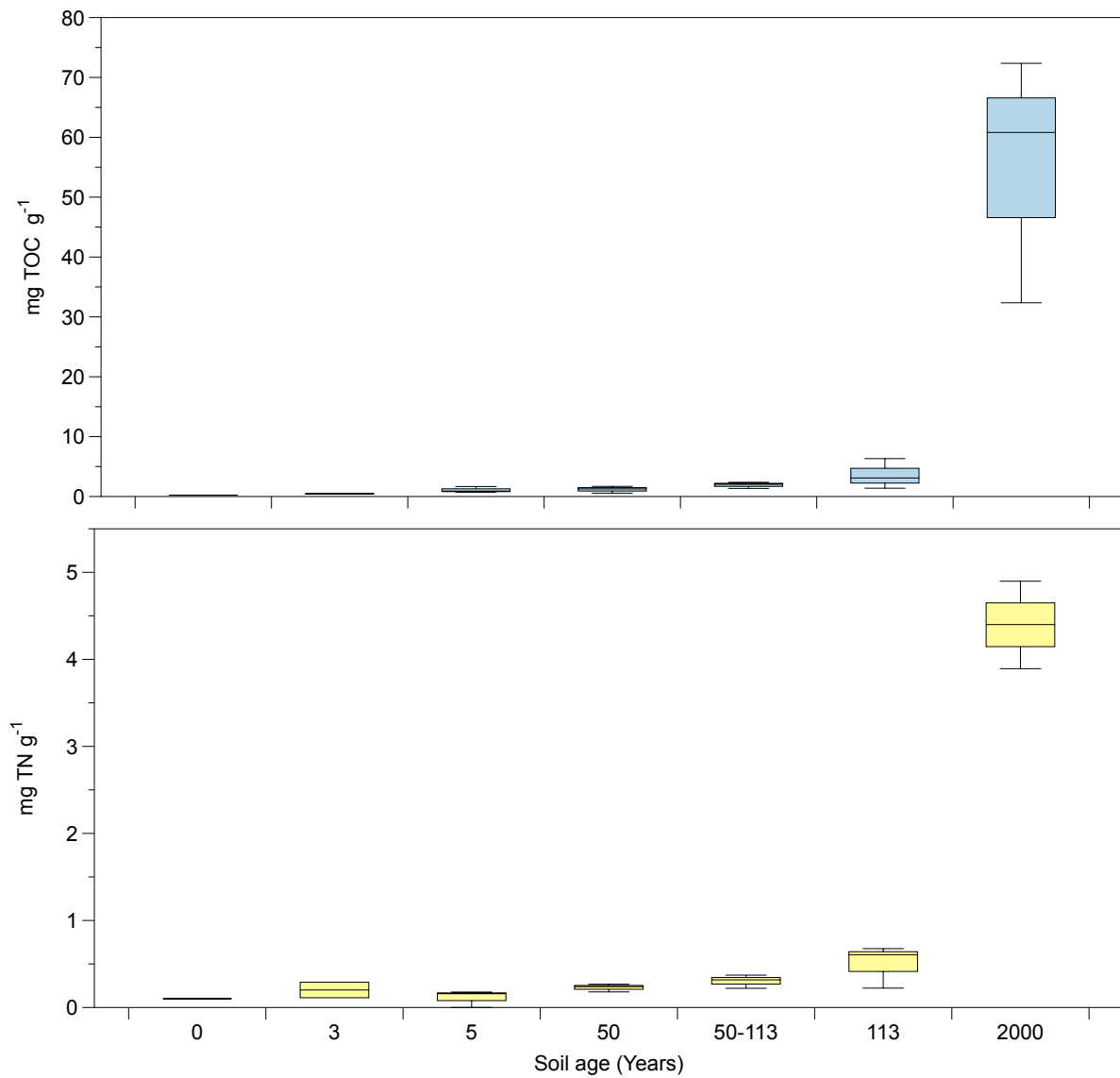
Soil temperature, pH, TN and TOC measurements were obtained from each metagenome sampling site and are displayed in Table 4.3. Soil temperature ranged between 4.8 – 8.5 °C,

between sites 1- 15 respectively. Soil pH ranged around neutral, between 6.2 – 8.5, and does not show a distinct variance with soil age (Table 4.3). The TOC values fell in to the general range observed in Arctic soils, with most soils ranging between 0 – 40 mg g<sup>-1</sup> TOC (Chu *et al.*, 2010). The results for TOC and TN are displayed in Table 4.3 and shown graphically in Figure 4.3. TOC ranged between 2.47 mg g<sup>-1</sup> – 86.97 mg g<sup>-1</sup>, between sites 8 – 21 (ages 29 – 2000) (Figure 4.3). The highest TOC values were found in the oldest soils (2000 years), the furthest sampled from the glacier terminus (Figure 4.3). Furthermore, in total, 86% of forefield sites had a TN content which fell below the instrumental limit of detection of 1 mg g<sup>-1</sup> (Table 4.3). However, older sites 19-21 (2000 years) had detectable TN, ranging between 3.89 - 4.89 mg g<sup>-1</sup> (Table 4.3). Consequently, both TN and TOC values showed an increase with rising soil age (Figure 4.3). Results of a one way ANOVA between all sites showed a significant difference in both TN and TOC, at the 0.01 significance level (Table 4.4). A post-hoc Tukey identified this significant difference to occur between earlier soils (ages 0-113 years) and the 2000 year soils (Table 4.4) This identifies a statistically significant increase in TN and TOC along the forefield, as expected with soil development.



**Table 4.3:** Total organic carbon (TOC), total nitrogen (TN), temperature and pH for metagenome sampling sites 1 – 21, soil ages 0- 2000 years. For TN and TOC analysis the instrumental limit of detection was 1 mg g<sup>-1</sup>, with samples falling below this limit marked as b.d. TN and TOC vales are not available for ice sampling sites 22 and 23.

Sample Number	Soil age (Years)	TOC (mg g <sup>-1</sup> )	TN (mg g <sup>-1</sup> )	Temperature (°C)	pH
1	0	6.68	b.d.	4.8	7.5
2	3	4.30	b.d.	7.6	7.6
3	3	3.89	b.d.	5.9	7.0
4	5	3.72	b.d.	6.6	7.5
5	5	3.68	b.d.	7.5	6.2
6	5	5.07	b.d.	7.6	7.3
7	29	4.07	b.d.	7.0	7.5
8	29	2.47	b.d.	8.3	7.5
9	29	2.96	b.d.	8.0	7.5
10	50	3.48	b.d.	7.4	7.5
11	50	5.47	b.d.	8.4	7.5
12	50	7.17	b.d.	8.5	7.3
13	50-113	3.09	b.d.	8.5	7.4
14	50-113	4.95	b.d.	8.4	7.5
15	50-113	2.72	b.d.	8.5	7.3
16	113	4.67	b.d.	7.8	7.5
17	113	11.53	b.d.	7.8	8.5
18	113	9.97	b.d.	7.7	7.5
19	2000	54.88	3.89	7.4	7.5
20	2000	67.08	4.39	7.4	7.5
21	2000	86.97	4.89	8.1	7.5



**Figure 4.3:** Box plots for total organic carbon (TOC) and total nitrogen (TN) for Midtre Lovénbreen forefield sampling sites, spanning soil ages 0 – 2000 years. The boxplots show the mean, first quartile, third quartile and range of the replicate field samples (where available).

**Table 4.4:** One-way ANOVA comparing TN and TOC differences between all soils. A post-hoc Tukey identifies which sites display the significant difference. Significant differences observed between the sites are noted at the 0.01 or 0.05 level.

	Total organic carbon (TOC)	Total Nitrogen (TN)
<b>p value</b>	1.5e-08	7.3e-15
<b>significance level</b>	0.01	0.01
<b>Pot hoc Tukey</b> 0-113 vs. 2000 years	All (P<0.01)	All (P<0.01)

### 4.3.3 Microbial community composition

The read-based microbial community composition of Midtre Lovénbreen metagenomes is displayed in Figure 4.4 and in Appendix 3 Table 2 (Genus Level) and in Table 4.5 (Phylum level). Metagenomes spanning the forefield chronosequence (0 years – 2000 years) are displayed alongside cryoconite and dark ice samples for comparison. The most abundant phyla recovered included Proteobacteria and Actinobacteria, accounting for 35% and 21% of annotated protein coding genes, respectively (Table 4.5). Other phyla assigned included Cyanobacteria, Chloroflexi, Verrucomicrobia, Firmicutes and Bacteroidetes, amongst others (Table 4.5). The percentage of protein coding genes attributed to Acidobacteria, Verrucomicrobia and Nitrospirae increased along the forefield, from 1.0%, 0.6% and 0% at 0 years to 11.7%, 4.9% and 0.9% at 2000 years, respectively (Table 4.5). This does not show an absolute change in organism abundance, due to differing depths of sequencing (Figure 4.2). However, as the sequencing profiles the most abundant fraction of the microbial community, it shows these phyla become a greater fraction of the microbial community recovered at older soil ages. Furthermore, in site 0 and dark ice, cyanobacteria account for 13.8% of the protein coding genes (Table 4.5). However, this reduces to 1.59% in soils aged 2000 years (Table 4.5). Whilst the cryoconite sample contains relatively few protein coding genes associated with cyanobacteria (0.7%), 13.8% of genes are attributed to the phototroph Chloroflexi, compared to the range of 1.4 - 3.9% in forefield samples (Table 4.5). Interestingly, no clear pattern can be seen in the distribution of Archaea and Fungi (Table 4.5). The Fungi Ascomyota is present in soils aged 0,3 and 50-113 years, with the Archaea, Thaumarchaeota and Eurayarchaeota present in 113 years and 3 years, 113 years and basal ice, respectively (Table 4.5).

Between 34.6 – 57.9 % of metagenome reads were annotated using Kaiju at the genus level (Table 4.6). The reads were assigned to a range of taxa, associated with diverse metabolisms including nitrogen, carbon and sulfur cycling (Figure 4.4). In particular, genera associated with nitrogen fixation (*Rhizobium*, *Nostoc*, *Oscillatoria*), denitrification (*Conexibacter*), nitrification (*Nitrospira*), sulfur oxidation (*Chromatiales*, *Thiobacillus*), sulfur reduction (*Shewanella*) and carbon fixation (*Phormedesmis*, *Nostoc*, *Oscillatoria*, *Synechococcales*, *Rhodofera*, *Rhodoplanes*, *Crococcales*) were identified (Figure 4.4; Appendix 3 Table 2). Genera associated with both aerobic (*Niabella*, *Variovorax*, *Rhodococcus*) and anaerobic metabolisms (*Rhodofera*, *Geobacter*, *Bacillus*) were also identified (Figure 4.4). Whilst we cannot determine the change in abundance of microbes along the chronosequence, due to differing metagenome sequencing depths, we can identify changes to the most abundant fraction of the microbial community. Soils aged 0 years contained the diazotrophs *Rhizobium*

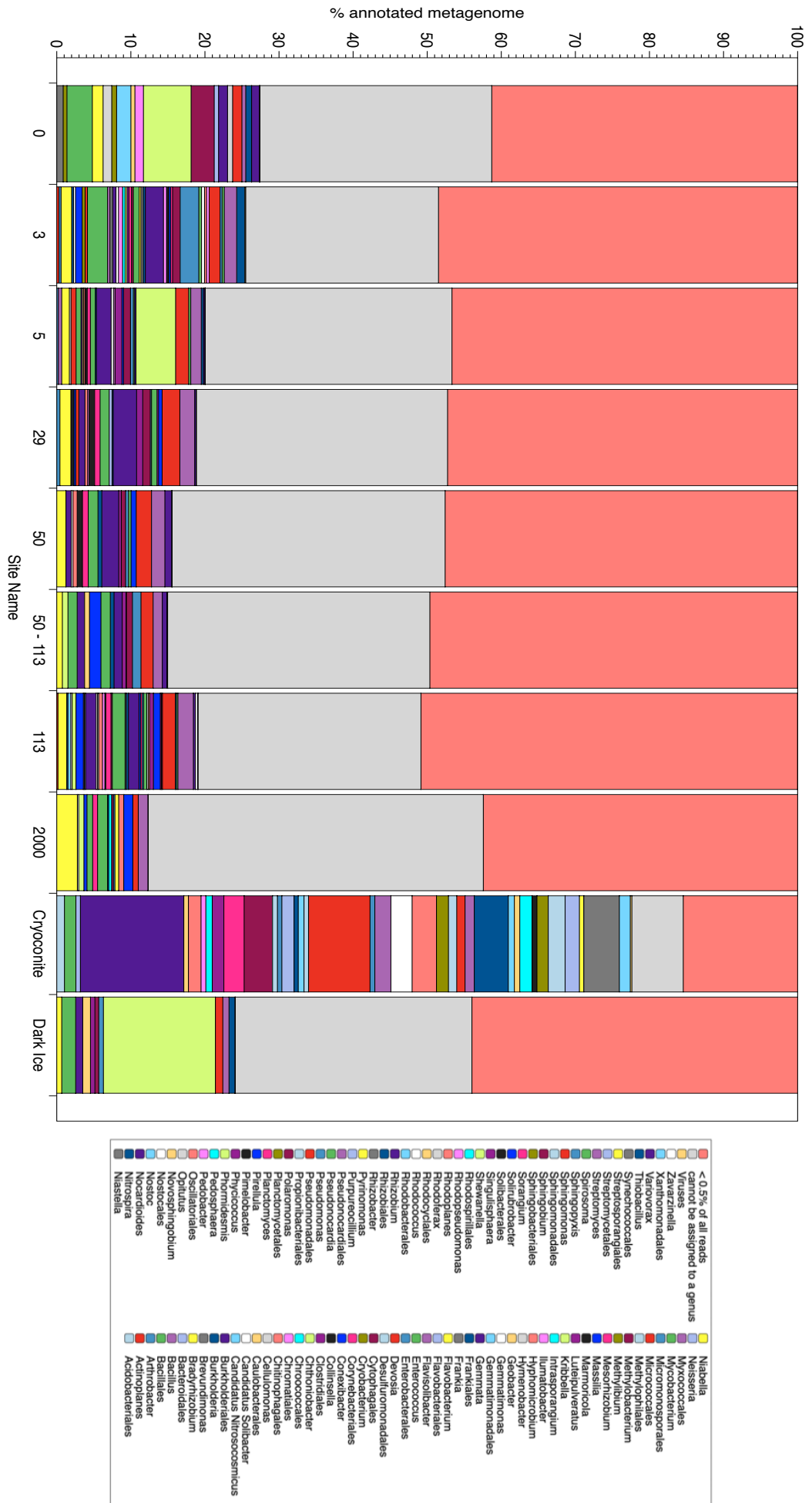
and *Nostoc*, alongside the sulfur oxidizer *Thiobacillus* (Figure 4.4). In soils aged 3-5 years, the recovered microbial community included nitrifying, denitrifying and sulfur reducing microbes (Figure 4.4). Carbon fixing (and often diazotrophic) cyanobacteria detected in the metagenomes included *Synechococcales*, *Phormidesmis*, *Oscillatoriales* and *Nostocales* (Figure 4.3). Cyanobacteria were only recovered from soils aged 0 years and cryoconite (8.4 and 11% of annotated reads) and not found in more developed soil sites (Figure 4.4). Root associated microbes, including symbiotic diazotrophs (*Frankia*, *Rhizobium*, *Massilia*, *Burkholderia*) were detected throughout the forefield and accounted for 1.2% - 3.4% of annotated reads, between soils ages 0 years and 2000 years, respectively (Figure 4.4).

**Table 4.5:** Read based microbial community composition of Midtre Lovénbreen sample sites, presented as the percentage of annotated sequences at the phylum level. Replicate metagenomes for each soil age/ sampling site have been merged.

	0	3	5	29	50	50 - 113	113	2000	Cryoconite	Dark ice
<i>Acidobacteria</i>	1.01	1.62	3.31	4.65	6.20	4.75	4.96	11.73	1.79	1.79
<i>Actinobacteria</i>	8.77	22.67	24.39	29.50	24.28	17.61	25.29	15.53	20.64	20.64
<i>Ascomycota</i>	1.33	4.01	0.00	0.00	0.00	0.66	0.00	0.00	0.00	0.00
<i>Bacteroidetes</i>	13.08	4.14	7.26	7.55	7.59	8.00	7.44	6.24	1.21	10.19
<i>Candidatus Rokubacteria</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.90	0.00	0.00
<i>Chloroflexi</i>	1.44	2.30	2.92	3.22	3.90	2.98	2.74	2.74	13.81	1.21
<i>Cyanobacteria</i>	13.76	1.96	1.50	1.83	1.48	3.04	2.37	1.59	0.71	13.81
<i>Deinococcus-Thermus</i>	0.00	0.30	0.00	0.00	0.00	0.00	0.21	0.00	0.69	0.71
<i>Euryarchaeota</i>	0.00	0.27	0.00	0.00	0.00	0.00	0.23	0.00	0.00	0.69
<i>Firmicutes</i>	5.07	4.96	3.74	2.49	2.37	3.85	3.15	3.32	4.30	4.30
<i>Gemmatimonadetes</i>	0.00	0.76	1.08	1.13	1.61	1.05	1.16	1.35	0.77	0.77
<i>Nitrospirae</i>	0.00	0.34	0.53	0.73	0.83	0.79	0.86	0.89	0.00	0.00
<i>Planctomycetes</i>	1.16	2.75	2.93	4.10	4.47	6.10	6.21	3.59	1.71	4.30
<i>Proteobacteria</i>	42.04	45.09	40.84	35.44	36.96	38.55	35.12	36.21	38.98	0.77
<i>Thaumarchaeota</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.24	0.00	0.00	0.00
<i>Verrucomicrobia</i>	0.58	1.21	1.74	1.66	1.98	3.52	2.36	4.95	0.53	0.53

**Table 4.6:** Percentage classified and unclassified reads for each metagenome at the genus level, following read based taxonomic annotation in Kaiju.

Sample	% Classified	% Unclassified
1	57.95	42.05
2	40.31	59.69
3	56.89	43.11
4	52.44	47.56
5	53.17	46.83
6	53.13	46.87
7	50.99	49.01
8	52.44	47.56
9	52.5	47.5
10	49.44	50.56
11	49.65	50.35
12	49.7	50.3
13	34.82	65.18
14	47.73	52.27
15	47.8	52.2
16	45.86	54.14
17	34.56	65.44
18	46.04	53.96
19	47.02	52.98
20	46.75	53.25
21	46.34	53.66
22	36.71	63.29
23	53.17	46.83



**Figure 4.4:** Read based community composition, expressed as the percentage of annotated reads for Midtre Lovénbreen sample sites at the genus level. Replicate samples for each soil age have been combined. The legend is ordered top to bottom, in line with the graph bars. A tabular format is available in Appendix 3 Table 2.

#### 4.3.4 Genome bins

In total, 83 genome bins were extracted from the metagenomes (Table 4.7). Genome bins were extracted from all metagenomes, except for sample 13. The percentage identity of the extracted bins ranged between 70.5 – 100 % identity to NCBI GenBank (Table 4.7). In total, 21% of bins fell below 80% identity to cultured representatives on NCBI GenBank, with only 8.5% with over 95% identity (Table 4.7). The bins were assigned to organisms with nitrogen metabolisms (16% of bins), sulfur metabolisms (11%) and also cyanobacteria (5%), anaerobes (12%) and plant associated organisms (4%) (Table 4.7). Some BLAST assignments were more prevalent than others, such as *Polaromonas* (5% of bins), *Sphingomonas* (5%), *Conexibacter woesei* (6%), *Sulfuricaulis limicola* (4%) and *Thiobacillus denitrificans* (3%). However, due to the low percentage identity of BLAST matches to publicly available sequences, we cannot be confident in the specificity of the assignments. The extracted bins were annotated with functional genes in RAST, with each bin containing between 144 - 9902 genes (Table 4.7). The genes were matched to known functions using the SEED database, with each bin containing between 126 – 3005 distinct functions (Table 4.7). Binned genomes were grouped with their field replicate samples and a function profile created, based on the SEED ontology. A subset of SEED functions was extracted for investigation, focused on nitrogen cycling pathways.

For each sample (1-23), extracted bins were used in maximum likelihood phylogenies based on a subset of COG groups (Appendix 3, Figure 1 – 22). This was used to audit the BLAST results and to investigate how well bins aligned to KBASE genomes. The trees show 77% of sample bins to be uniquely branching, clustering independently from publicly available KBASE genomes, in agreement with the low percentage identity of BLAST match results (Appendix 3, Figure 1-22).

**Table 4.7:** Genome binning results for the 23 assembled metagenomes. For each metagenome, the extracted bins are listed with the top BLASTn hit to NCBI GenBank, along with the RAST annotation results.

Blastn Results						RAST annotation results	
Sample	Acession No	% identity	E value	Max Score	Species	Number of genes	SEED Functions
<b>1</b>							
Bin 1	CP002355.1	94.81	0	5243	<i>Sulfuricurvum kujiense</i>	2849	1070
Bin 2	CP024785.1	92.15	0	13771	<i>Nostoc flagelliforme</i>	3440	1157
Bin 3	CP003178.1	76.26	0	1629	<i>Niastella koreensis</i>	1421	648
Bin 4	CP011131.1	79.79	0	1321	<i>Lysobacter gummosus</i>	1362	884
Bin 5	CP026692.1	88.52	0	3860	<i>Nostoc</i> sp.	806	373
Bin 6	CP000116.1	86.93	0	3068	<i>Thiobacillus denitrificans</i>	1992	867
Bin 7	CP019508.1	91.72	0	2747	<i>Brevundimonas</i>	1194	867
Bin 8	CP003614.1	89.43	0	2645	<i>Oscillatoria nigro-viridis</i>	1177	515
Bin 9	CP027482.1	81.53	0	4942	<i>Aeromicrobium</i>	2224	960
Bin 10	CP000116.1	83.22	0	3788	<i>Thiobacillus denitrificans</i>	3553	1161
<b>2</b>							
Bin 1	CP025581.1	80.17	0	3472	<i>Nocardioides</i>	3005	1050
Bin 2	CP001854.1	92.76	0	3864	<i>Conexibacter woesei</i>	162	73
<b>3</b>							
Bin 1	AP014879.1	84.39	0	4165	<i>Sulfuricaulis limicola</i>	2625	1140
Bin 2	CP012573.1	84.04	0	6135	<i>Clavibacter capsici</i>	1329	751
Bin 3	AP012057.1	80.85	0	2261	<i>Ilumatobacter coccineus</i>	2022	815
Bin 4	CP010554.1	87.37	0	2030	<i>Rugosibacter aromaticivorans</i>	1060	563
Bin 5	CP012371.1	80.27	0	2462	<i>Nitrosospira briensis</i>	1535	798
Bin 6	CP015079.1	83.56	0	2970	<i>Nocardioides dokdonensis</i>	2748	992
Bin 7	AP014879.1	86.33	0	3921	<i>Sulfuricaulis limicola</i>	1194	593
Bin 8	CP000116.1	84.61	0	1947	<i>Thiobacillus denitrificans</i>	841	468
<b>4</b>							
Bin 1	CP015732.1	94.33	0	3085	<i>Arthrobacter</i>	3212	1057
Bin 2	CP000529.1	90.13	0	3720	<i>Polaromonas naphthalenivorans</i>	1144	612
<b>5</b>							
Bin 1	AP014879.1	76.92	0	2534	<i>Sulfuricaulis limicola</i>	2609	996
Bin 2	CP031145.1	92.32	0	4575	<i>Intrasporangium calvum</i>	1574	581
Bin 3	CP009111.1	85.94	0	1930	<i>Rhodococcus opacus</i>	1574	1073
Bin 4	CP001854.1	78.26	0	2145	<i>Conexibacter woesei</i>	2031	787
<b>6</b>							
Bin 1	CP002399.1	83.06	0	3013	<i>Micromonospora</i>	5386	11259

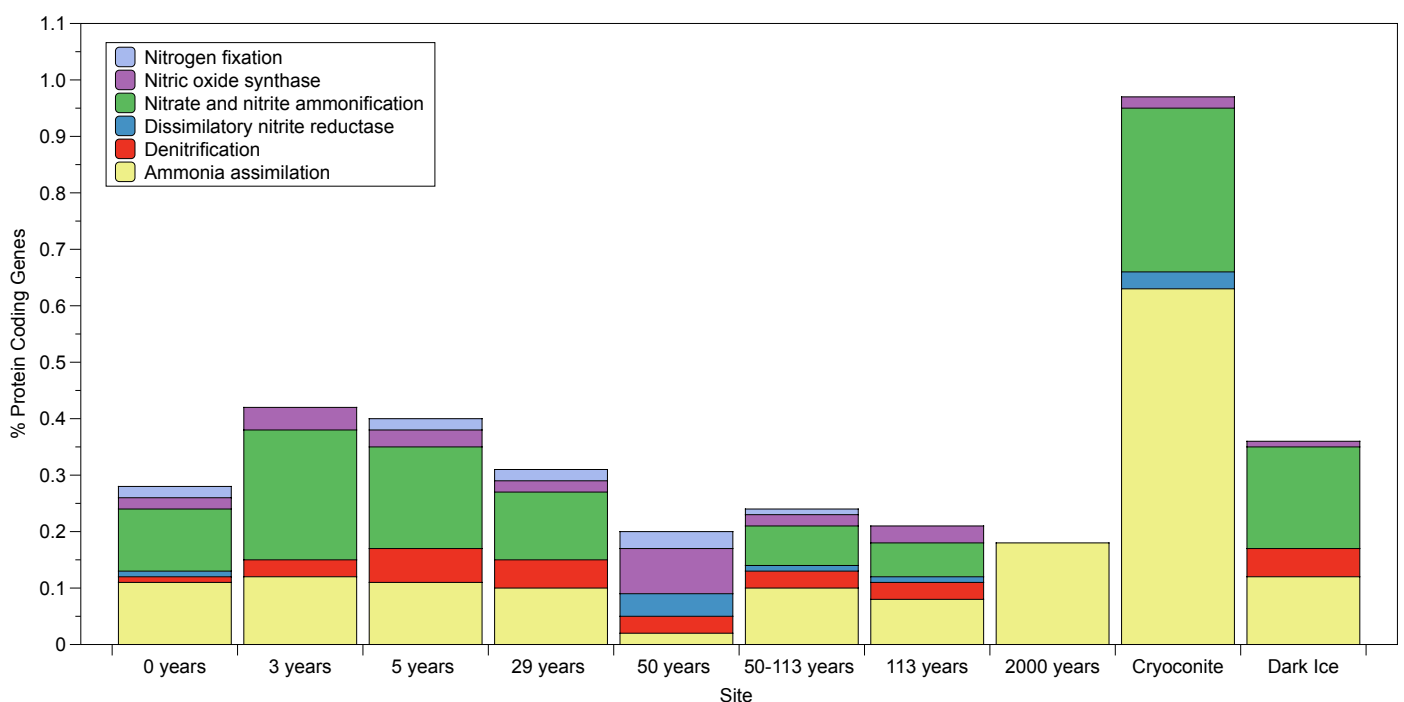
Bin 2	CP004036.1	70.05	2.00E-45	196	<i>Sphingomonas</i>	4866	1200
Bin 3	CP002199.1	90.2	2.00E-06	67.6	<i>Cyanothece</i>	3388	940
Bin 4	CP006005.1	93.33	9.00E-07	67.6	<i>Vibrio parahaemolyticus</i>	2224	272
<b>7</b>							
Bin 1	CP002994.1	82.83	2.00E-13	89.8	<i>Streptomyces violaceusniger</i>	2678	861
Bin 2	CP021181.1	74.87	1.00E-78	307	<i>Sphingomonas wittichii</i>	3426	1086
Bin 3	CP012184.1	77.19	7.00E-157	566	<i>Pseudonocardia</i>	3089	905
Bin 4	CP011339.1	81.37	2.00E-11	84.2	<i>Microcystis panniformis</i>	2315	726
<b>8</b>							
Bin 1	CP015732.1	94.33	0	3085	<i>Arthrobacter</i>	3212	1057
Bin 2	CP000529.1	90.13	0	3720	<i>Polaromonas naphthalenivorans</i>	1130	611
<b>9</b>							
Bin 1	LT827010.1	95.24	0	2918	<i>Actinoplanes</i>	2996	1041
Bin 2	CP001854.1	92.76	0	3864	<i>Conexibacter woesei</i>	153	67
<b>10</b>							
Bin 1	CP014989.1	83.94	0	3241	<i>Serinicoccus</i>	1303	742
Bin 2	CP010954.1	79.31	0	1493	<i>Sphingobium</i>	1966	924
Bin 3	CP000316.1	83.2	0	2499	<i>Polaromonas</i>	4632	791
Bin 4	CP030865.1	85.71	5.00E-46	198	<i>Micromonospora</i>	1900	760
Bin 5	CP001124.1	81.21	0	1123	<i>Geobacter bemidjiensis Bem</i>	2919	1098
Bin 6	CP000533.1	91.78	0	3690	<i>Streptomyces lunaelactis</i>	1619	552
Bin 7	CP026304.1	94.06	0	3241	<i>Polaromonas naphth.</i>	4633	1049
Bin 8	CP000531.1	96.7	0	1358	<i>Polaromonas naphth.</i>	852	381
<b>11</b>							
Bin 1	FO117623.1	91.54	0	2510	<i>Blastococcus saxobsidens</i>	2136	791
Bin 2	CP006644.1	92.08	0	2643	<i>Sphingomonas sanxanigenens</i>	1291	516
<b>12</b>							
Bin 1	FO117623.1	91.65	0	2521	<i>Blastococcus saxobsidens</i> DD2	2142	772
Bin 2	CP006644.1	92.08	0	2643	<i>Sphingomonas sanxanigenens</i>	1285	515
<b>14</b>							
Bin 1	CP029343.1	91.03	0	4739	<i>Massilia oculi</i>	3661	1109
Bin 2	CP029343.1	93.05	0	5033	<i>Massilia oculi</i>	2665	1124
<b>15</b>							
Bin 1	CP027775.1	95.55	0	2837	<i>Clostridium botulinum</i>	427	136
Bin 2	LT827010.1	95.41	0	2942	<i>Actinoplanes</i>	2138	798
<b>16</b>							
Bin 1	CP027775.1	95.55	0	2837	<i>Clostridium botulinum</i>	429	135
Bin 2	CP002479.1	75.86	2.00E-21	117	<i>Geobacter</i>	2137	795
<b>17</b>							
Bin 1	CP009241.1	91.3	0.00004	62.1	<i>Paenibacillus</i>	2394	1006
Bin 2	CP009241.1	91.3	0.00004	62.1	<i>Paenibacillus</i>	3534	891
Bin 3	CP009571.1	81.75	0	734	<i>Sphingomonas taxi</i>	6927	1083



Bin 4	CP031968.1	80	9E-10	78.7	<i>Chromobacterium rhizoryzae</i>	1495	600
<b>18</b>							
Bin 1	CP031146.1	92.86	5E-11	82.4	<i>Pseudomonas plecoglossicida</i>	2094	830
Bin 2	AP014946.1	76.36	0	2464	<i>Variibacter gotjawalensis</i>	2839	1258
Bin 3	CP001854.1	76.88	9E-143	520	<i>Conexibacter woesei</i>	5058	958
Bin 4	CP021904.1	100	0.000003	65.8	<i>Alkalitalea saponilacus</i>	1170	499
Bin 5	CP020569.1	85.71	0	2311	<i>Streptomyces gilvosporeus</i>	5924	1015
Bin 6	CP026952.1	81.25	5E-11	82.4	<i>Aeromicrobium</i>	5313	1191
Bin 7	CP011271.1	78.03	0	983	<i>Gemmata</i>	5397	1016
Bin 8	CP011132.1	82.42	3E-10	80.5	<i>Citrobacter amalonaticus</i>	3734	910
<b>19</b>							
Bin 1	CP001854.1	87.68	0	1101	<i>Conexibacter woesei</i>	331	138
Bin 2	CP001389.1	76.65	1E-27	137	<i>Sinorhizobium fredii</i>	188	93
<b>20</b>							
Bin 1	CP022522.1	97.56	6E-08	71.3	<i>Pseudoalteromonas</i>	144	56
Bin 2	CP018171.1	89.74	7E-32	150	<i>Mesorhizobium oceanicum</i>	181	83
<b>21</b>							
Bin 1	CP011491.1	78.61	3E-55	228	<i>Mycolicibacterium vaccae</i>	673	243
Bin 2	CP001197.1	80.31	1E-44	193	<i>Desulfovibrio vulgaris</i>	432	172
<b>22</b>							
Bin 1	AP017308.1	80.64	0	2817	<i>Leptolyngbya</i> sp.	2239	1026
Bin 2	CP016768.2	90.99	0	4482	<i>Candidatus Nanopelagicus limnes</i>	9902	636
Bin 3	CP000494.1	82.08	0	1205	<i>Bradyrhizobium</i>	2663	894
Bin 4	CP016768.2	79.64	0.00E+00	2198	<i>Candidatus Nanopelagicus limnes</i>	1948	707
Bin 5	CP016282.1	73.68	3.00E-127	468	<i>Cryobacterium arcticum</i>	1266	534
<b>23</b>							
Bin 1	AP014879.1	76.92	0	2534	<i>Sulfuricaulis limicola</i>	2608	995
Bin 2	CP031145.1	92.32	0.00E+00	4575	<i>Intrasporangium calvum</i>	1584	583
Bin 3	CP009111.1	85.94	0	1930	<i>Rhodococcus opacus</i>	3320	1072
Bin 4	CP001854.1	78.26	0.00E+00	2145	<i>Conexibacter woesei</i>	2029	784

For the binned genomes at each sampling site, a SEED function profile was created for nitrogen cycling pathways. This profile highlights the percentage of protein coding genes attributed to each functional classification. Whilst the functional profiles were not exhaustive of all microbes (due to inadequate sequencing depth in some samples and incomplete metagenome assembly), it does show the most abundant nitrogen cycling pathways in the genomes that were recovered from the binning process. Additionally, whilst the binning did not sample all genes in each binned genome due to inadequate coverage, it exemplifies the potential of this approach for deeply sequenced samples, highlighting genome functionality.

The SEED classifications for nitrogen cycling pathways are available in Figure 4.5. The pathways include nitrogen fixation, nitrification, denitrification and ammonia assimilation (Figure 4.5). Nitrogen fixation was present in extracted genomes up to 50 years and was not detected in binned genomes from cryoconite and dark ice samples (Figure 4.5). However, ammonia assimilation was found in all samples, reaching 0.63% of recovered protein coding genes in the cryoconite sample (Figure 4.5). Denitrification, nitrite and nitrate ammonification and dissimilatory nitrite reductase were largely found in genome bins throughout the forefield, however were not detected in bins from 2000 year soils (Table 4.3; Figure 4.5). Ammonia assimilation and ammonification were the most prevalent nitrogen cycling pathways, accounting for up to 0.63% and 0.29% of protein coding genes, respectively (Figure 4.5). Whilst SEED genome annotation can't be used to profile the complete nitrogen cycling pathways in the metagenomes (due to inadequate sequencing and assembly), it does provide an insight into the functional potential of abundant genomes.



**Figure 4.5:** SEED function profile for nitrogen metabolism in genome bins across sampling sites. The percentage of protein coding genes attributed to each pathway is shown.

## 4.4 Discussion

### 4.4.1 Total nitrogen and total organic carbon

Soil carbon and nitrogen content were investigated to understand changes to soil nutrient content along the chronosequence. This provides important contextual information from which hypotheses regarding community composition and functional shifts can be investigated. The soil total nitrogen and total organic carbon content revealed changes in soil nutrient content with succession. As shown by Figure 4.3, the highest TN and TOC values were found in the soils which were furthest from the glacier terminus and therefore had been exposed for the longest duration. This is in agreement with research by Turpin-Jelfs *et al.*, (2019) who found both carbon and nitrogen content to increase in soils along the Midtre Lovénbreen forefield. This may be because these soils have been exposed by glacier retreat and have been subject to soil development, in line with our hypothesis (Bradley *et al.*, 2014). TOC accumulates in forefield soils from allochthonous and autochthonous sources, such as through deposition of soot and organic matter, alongside the action of autotrophic microbial communities (Guelland *et al.*, 2013; Bradley *et al.*, 2014). The increased TOC content observed here in older soils is consistent with studies from other forefields, such as the Damma Glacier (Switzerland) and has been represented in SHIMMER, a numerical model of glacier forefield succession (Guelland *et al.*, 2013; Bradley *et al.*, 2016). The increase in TOC content with soil age is a crucial factor in soil development and facilitates colonization by plants and heterotrophic microbes (Schutte *et al.*, 2009). This largely explains why initial (newly exposed) soils typically support minimal vegetation and a greater plant density can be observed in the later more developed soils (Tscherko *et al.*, 2005). Variance within the forefield may relate to disturbances such as water flow pathways or variance in soil type (Bekku *et al.*, 2004; Frey *et al.*, 2013). Whilst older soils contained the highest TOC content, initial soils (0 – 5 years) were also found to contain detectable levels of TOC, between 3.89 – 6.68 mg g<sup>-1</sup> (Table 4.3). The presence of TOC in newly exposed soils has been previously attributed to ancient subglacial organic carbon and aeolian deposition, which may be crucial for supporting initial heterotrophic microbial communities, prior to the establishment of autotrophs (Hodkinson *et al.*, 2002; Schulz *et al.*, 2013).

The majority of sites across the forefield had TN levels below detectable limits (Table 4.3). As nitrogen is crucial for microbial and plant protein synthesis, minimal levels are indicative of low nutrient conditions and potential nitrogen limitation (Turpin-Jelfs *et al.*, 2019). Typically, nitrogen is supplied to forefield soils through autochthonous sources such as microbial nitrogen fixation or remineralization, alongside allochthonous sources such as snowmelt or aeolian deposition (Bradley *et al.*, 2014). In general, studies have shown that microbial

nitrogen fixation (diazotrophy) is the most important source in early sites, helping to facilitate the establishment of higher microbes and plants (Brankatschk *et al.*, 2011). The lack of detectable nitrogen in the early soils may relate to an inactivity of diazotrophs or a tight coupling between nitrogen fixation and consumption in soils. Low TN values are often observed in recently deglaciated forefields due to a lack of soil development, with studies across forefields from Canada, Antarctic, Svalbard, Austria, Italy and Switzerland finding values ranging between 0.1 – 2 mg g<sup>-1</sup> (Bradley *et al.*, 2014). However, at a soil age of 2000 years a TN content of 3.89 – 4.89 mg g<sup>-1</sup> was detected (Figure 4.3). This is consistent with the expected increase in TN with soil development, in line with the TOC results. Whilst the source of the additional TN cannot be isolated, it may come from a combination of diazotrophy, nitrogen remineralization or increased aeolian deposition (Bradley *et al.*, 2014).

#### 4.4.2 Microbial community composition

A chronosequence of metagenomes were evaluated along the forefield, to test our hypothesis that microbial community composition and function would change alongside soil succession and nutrient pool development. We hypothesized autotrophic organisms would be dominant in early soils, with more heterotrophic nutrient cycling appearing later in the chronosequence, relating to nutrient availability.

Throughout all metagenomes, the read based taxonomic annotation shows both aerobic and anaerobic organisms can be identified, with no sites containing exclusively anaerobes or microaerophilic organisms (Figure 4.4). Aerobes identified include organisms such as *Niabella*, *Variovorax*, *Rhodococcus* and *Pseudomonas* (Figure 4.4). Anaerobic organisms recovered from the metagenomes include *Rhodoferax*, *Opitutus*, *Geobacter* and *Bacillus* (Figure 4.4). The presence of strict and facultative anaerobes in forefield soils has been previously attributed to the periodic flushing by subglacial meltwater (Duc *et al.*, 2009). However, as anaerobic strains have been identified in each site, this may indicate anoxic or microaerophilic micro environments, alongside the aerobic soil surface layer. This could be attributed to poor drainage of glacial meltwater. Anoxic environments are often found in poorly drained subglacial systems, and therefore these anaerobes may be a legacy from the subglacial environment (Wadham *et al.*, 2008; Boyd *et al.*, 2010).

Extremophilic and psychrophilic organisms were recovered from the metagenomes, alongside those with traits to aid colonization and survival in the cold, high UV, oligotrophic conditions typical of Arctic forefields (Figure 4.4). Sequences attributed to *Bacillus* were identified in soils at 5 years following glacier retreat (Figure 4.4). This extremophile is endospore forming and

is resistant to both cold and desiccation, with psychrophilic, acidophilic, alkaliphilic, halotolerant, and halophilic properties (Joan *et al.*, 2011). The production of endospores allows *Bacillus* to remain in a dormant state under stressful conditions (Nicholson *et al.*, 2000). *Bacillus* are typically found in extreme environments, such as deserts and Arctic soils, due to their high resistance to environmental stressors (Rüger *et al.*, 2000). In addition, *Candidatus Solibacter* and *Pseudomonas* were recovered from sites ranging from 3 – 2000 years since exposure (Figure 4.4). These organisms secrete extracellular polymeric substances (EPS), consisting of polysaccharides, proteins, lipids and DNA (Yang *et al.*, 2011). This EPS creates a biofilm surrounding the microbes, helping to reduce temperature and nutrient fluctuations (Flemming and Wingender, 2010). This is a crucial mechanism for environmental tolerance, particularly to cold and desiccation stressors in the Arctic. Bacteria with cold-tolerance mechanisms were also identified in the forefield metagenome samples and dark ice, such as *Cryobacterium* and *Polaromonas* (Figure 4.4). Both species are psychrophiles and found in cold environments such as the Arctic, which may be due to cold tolerance mechanisms such as EPS production and cell membrane fluidity (Irgens *et al.*, 1996; Suzuki *et al.*, 1997).

The cyanobacteria, *Nostoc*, *Oscillatoria*, *Phormidesmis* and *Synechococcales* were found in recently exposed soils (0 years) and cryoconite samples, in support for our hypothesis (Figure 4.4). Cyanobacteria have previously been proposed as early colonizers of forefield soils and are often found in cryoconite (Zumsteg *et al.*, 2013; Christner *et al.*, 2003; Edwards *et al.*, 2011). Cyanobacteria are photosynthetic and many are diazotrophic, and therefore they do not rely on fixed sources of carbon (and nitrogen) for growth (Mitusi *et al.*, 1986). This is crucial in newly exposed soils and cryoconite on the glacier surface, as labile carbon and nitrogen stocks are typically limited (Christner *et al.*, 2003). Furthermore, cyanobacteria are resilient to environmental stressors such as sub-freezing temperatures, attributed to the production of protective EPS, which buffers temperature, desiccation and pH (de los Rios *et al.*, 2015). The recovery of the mat forming cyanobacteria, *Oscillatoria*, from the cryoconite metagenome is in agreement with the surrounding literature, which highlights it as one of the most abundant cryoconite cyanobacteria (Müller *et al.*, 2015; Edwards *et al.*, 2011). Cyanobacterial sequences were not recovered from more developed soil sites, however this may relate to inadequate sampling, as including more sample sites and DNA extraction from rock biofilms may highlight the presence of cyanobacteria. Additional deep sequencing of samples would help to highlight less abundant microbes in the metagenome samples, however the results show that cyanobacteria are not a dominant fraction of microbial communities in older soil samples. This reflects the results of Zumsteg *et al.*, (2013), who found cyanobacteria to decline with increasing soil age across the Damma Glacier forefield, Switzerland.

The microbial community composition of newly exposed soils has previously been the subject of scientific debate. The presence of heterotrophic microbial populations in early colonizer communities is not consistent between forefields. Some studies have identified heterotrophic colonizers in newly exposed soils, whilst others suggest autotrophic bacteria comprise the bulk of the initial community composition (Hodkinson *et al.*, 2002; Kastovska *et al.*, 2005). The composition of colonizer communities in this study (Site 0 years) contained both autotrophs (*Nostoc*, *Phormedesmis*, *Rhodofera*, *Thiobacillus*) and heterotrophs (Gemmata, Rhodospirillales, Sphingopyxis) (Figure 4.4). This may relate to the presence of low, but detectable TOC at this site, of 6.68 mg g<sup>-1</sup> (Table 4.3). This TOC may have been sourced from aeolian deposition or from the subglacial environment (Kastovska *et al.*, 2005; Bardgett *et al.*, 2007). The presence of TOC can provide a labile carbon pool for heterotrophic colonization (Kastovska *et al.*, 2005). However, as the TOC content of newly exposed soils will vary between glaciers, based on the content of overridden material and deposited material, the presence of heterotrophs is likely to be inconsistent between glaciers.

Furthermore, runoff from the ice surface and cryoconite may be a key fertilizing mechanism for newly exposed soils (Kastovska *et al.*, 2005; Stibal *et al.*, 2006). This dispersal may be due to the drainage of supraglacial meltwater through moulins to the glacier terminus (Stibal *et al.*, 2006). In our study site, autotrophic and heterotrophic microbial populations found in the cryoconite or dark ice sample were also found in newly exposed soils (0 years), including *Nostoc*, *Thiobacillus*, *Sphingomonas*, *Gemmata*, *Frankia*, *Corynebacteriales*, *Bacteriodiales* and *Rhodospirillales* (Figure 4.4). Consequently, ice surface microbial populations may be important for facilitating the colonization and composition of pioneer microbial communities. Again, this is largely a function of aeolian deposition, as glacier ice is most often fertilized by microbial populations in wind driven material (Marshall and Chalmers, 1997). Additionally, the composition of subglacial microbial communities may mediate the composition of initial colonizer communities, for example supplying anaerobic or sulfur cycling organisms, however subglacial samples were not available for validation in this analysis. Consequently, the composition of pioneer microbial communities is likely to vary between glaciers, influenced by the content of depositional material, ice microbial surface populations, subglacial communities and the nutrient content of initial soils.

#### 4.4.3 Microbial metabolisms – Carbon

A wide range of microbial metabolic potential was identified in the forefield and glacier metagenomes, including chemolithoautotrophs, organoheterotrophs, chemoheterotrophs and chemolithotrophs, involved in carbon, nitrogen and sulfur cycling (Figure 4.4). This identifies

microbial communities to have a potential influence on a multitude of biogeochemical cycles in forefield soils. Heterotrophic organisms potentially involved in microbial carbon cycling were found in soils 3 years following glacier retreat (Figure 4.4). This includes organisms such as *Micromonospora*, which grows off decaying organic matter in soils (White *et al.*, 1996). Additionally, bacteria capable of utilizing multiple carbon sources, including aromatic compounds, were recovered, such as *Rhodococcus opacus*, *Sphingomonas* and *Rugosibacter aromaticivorans* (Figure 4.4). The presence of carbon cycling microbes in sites above 3 years may relate to the minimal carbon content in newly exposed soils (Figure 4.3). Although TOC does not increase significantly from 0 years to 3 years, deposition of aromatic hydrocarbons from fossil fuel combustion and carbon fixation by phototrophic cyanobacteria, may provide sufficient labile carbon for heterotrophic carbon cycling bacteria to colonize (Margesin *et al.*, 2003; Frey *et al.*, 2013). The importance of phototrophs in facilitating the colonization of heterotrophic microbes and plants in forefield soils is supported by the surrounding literature, highlighting the important source of carbon phototrophs provide (Kastovska *et al.*, 2005; Frey *et al.*, 2013). Consequently, this provides support for our hypothesis that heterotrophic cycling would appear in later sites, due to nutrient availability.

#### 4.4.4 Microbial metabolisms – Nitrogen

Microbes with soil nitrogen cycling potential were also highlighted using the read based taxonomic annotation. Nitrogen-fixing cyanobacteria were recovered from 0 year soils and cryoconite metagenomes using read annotation and genome binning, including *Nostoc*, *Oscillatoria* and *Leptolyngbya* (Figure 4.4; Table 4.7). Nitrogen-fixing cyanobacteria may be fundamental for the build-up of labile nitrogen in nutrient deplete initial soils, as they convert nitrogen gas ( $N_2$ ) to ammonium ( $NH_4^+$ ) (Duc *et al.*, 2009; Figure 4.1). These cyanobacteria may therefore be important facilitators for the colonization of heterotrophs, considering the low nitrogen stocks observed at site 0 years (Figure 4.3). This is supported by SEED function profiles, which identified the nitrogen fixation pathway to be present in soils up to 50 years (Figure 4.5). This provides support for our hypothesis that autotrophic nitrogen cycling microbes would be found in early soils, due to low nutrient availability. The findings are supported by Turpin-Jelfs *et al.*, (2019) who found biological nitrogen fixation to occur in newly exposed soils in the Midtre Lovénbreen forefield, and decreased long the chronosequence.

Our results show the presence of cyanobacteria in cryoconite samples. As these microbes can fix carbon and often nitrogen, they are able to withstand the oligotrophic conditions on the glacier surface (Stibal *et al.*, 2006; Cameron *et al.*, 2012). The flushing of meltwater from cryoconite, through moulins to the glacier terminus, may be an important dispersal

mechanism, alongside wind drift, for cyanobacteria to reach newly exposed soils (Mueller *et al.*, 2001). Interestingly, the nitrogen fixation pathway was not detected in genome bins from the cryoconite sample, however this may relate to poor recovery of diazotrophic (*nif*) genes during sequencing due to EPS interference with DNA extraction, or a reduced need for nitrogen fixation (Figure 4.5). However, data on sediment nitrogen content was not available for this site. Root associated diazotrophs were found in metagenome samples exposed for 113 years or more (Figure 4.4). These organisms are symbiotic, fixing nitrogen in return for organic carbon substrates from plant roots (Franche *et al.*, 2009). The root associated diazotrophs recovered from read annotation and genome binning included *Mesorhizobium oceanicum*, *Bradyrhizobium*, *Sinorhizobium fredii* and *Chromobacterium rhizoryzae* (Figure 4.4; Table 4.7). The identification of rhizobia in older soil sites is consistent with the colonization of plants at later stages of soil development (Knelman *et al.*, 2012). However, *Bradyrhizobium* were detected in cryoconite samples, despite a lack of vascular plants (Figure 4.4). This may be the result of wind dispersal and bird droppings, common mechanisms for fertilizing the glacier surface with microorganisms (Kastovska *et al.*, 2005).

Bacteria related to nitrification (*Nitrospira briensis*) and denitrification (*Conexibacter woesei*, *Thiobacillus denitrificans*) were also found in soils after 3 years of exposure (Figure 4.4; Table 4.7). *Nitrospira briensis* are involved in ammonia oxidation, the first step in nitrification, oxidizing ammonia (NH<sub>3</sub>) to Nitrite (NO<sub>2</sub><sup>-</sup>) (Figure 4.1; Teske *et al.*, 1994; Daims *et al.*, 2015). As nitrifying bacteria rely on labile ammonia, they may only occur after 3 years of soil exposure, when sufficient stocks have accumulated from deposition, ammonification and nitrogen fixation. This is supported by Brankatschk *et al.*, (2011) who found low potential nitrification and denitrification rates at initial forefield sites, attributed to the lack of available nitrate and ammonium (Brankatschk *et al.*, 2011). However, if this is the case, the production or deposition of labile ammonia must be tightly coupled to uptake, given the minimal TN values recorded in our sampling sites (Table 4.3). Denitrifying bacteria are largely responsible for the loss of labile nitrogen from soil environments. Denitrification involves the reduction of fixed nitrate (NO<sub>3</sub><sup>-</sup>) to nitrogen gas (N<sub>2</sub>) by anaerobic or facultatively anaerobic bacteria (Figure 4.1; Aulakh *et al.*, 1992). Through denitrification, nitrate/nitrite is used as the terminal electron acceptor in respiration, in the absence of oxygen (Simon *et al.*, 2009). Denitrifying bacteria were recovered from soils exposed for 3 years or more and from the dark ice sample (Figure 4.4). This may relate to the requirement of labile nitrate, hindering the occurrence in newly exposed soils (0 years) (Brankatschk *et al.*, 2011). This supports the research of Kandeler *et al.*, (2006), who found evidence of denitrifying communities developing under soil succession, driven by an increasing availability of organic substrates. Denitrification is common in soils and may occur along the Midtre Lovénbreen forefield in wet or waterlogged conditions, where



the oxygen supply is limited (Christensen *et al.*, 1990). The presence of *Conexibacter woesei* in the dark ice metagenome indicates the potential for denitrification on the glacier surface (Figure 4.4). Denitrifying bacteria have been found in anoxic glacier ice and basal samples where oxygen is limited, including the Midtre Lovénbreen subglacial environment (Hodson *et al.*, 2005; Simon *et al.*, 2009; Ansari *et al.*, 2013). This is also supported by the SEED function profile for genome bins, which found denitrification to occur throughout the forefield and in the dark ice sample (Figure 4.5). Interestingly, the nitrification and denitrification pathways were not found in soils aged 2000 years in the SEED profiles. This may be due to the genome sampling that was carried out (i.e. based on assembled, binned, genomes) and therefore is not as comprehensive in terms of diversity as the read based taxonomic annotation. The results presented here therefore support the hypothesis that heterotrophic nitrogen cycling would appear in later soil sites, requiring an initial buildup of labile nitrogen stocks.

#### 4.4.5 Microbial metabolisms – Sulfur

Evidence for the potential of microbial sulfur cycling was also found in the forefield metagenomes and cryoconite samples. Sulfur oxidizing microbes were found in early soils, aged 0 and 5 years, alongside the cryoconite sample by both read annotation and genome binning and included *Thiobacillus denitrificans*, *Sulfuricurvum kujinse* and *Sulfuricaulis limicola* (Figure 4.4; Table 4.7). Sulfur oxidation is the process by which elemental ( $S^0$ ) or reduced ( $H_2S$ ,  $HS^-$ ) sulfur is oxidized to sulfate ( $SO_4^{2-}$ ), coupled with the reduction of oxygen (aerobic) or nitrate (anaerobic) (Eriksen *et al.*, 1998). This mechanism is used for energy production by chemolithotrophic sulfur oxidizers and has been identified in subglacial systems (Wainright 1978; Bottrell and Tranter, 2002). In addition, evidence for microbial sulfate reduction was provided by the recovery of *Desulfovibrio vulgaris* and *Geobacter bemidjensis* *Bem* sequences from more developed soil sites (aged 50 – 2000 years) (Figure 4.4; Table 4.7). Sulfate reducers use sulfate ( $SO_4^{2-}$ ) as a terminal electron acceptor for anaerobic respiration, reducing it to hydrogen sulfide ( $H_2S$ ) and aid the degradation of organic matter (Eriksen *et al.*, 1998). The presence of anaerobic sulfur reducers is indicative of anoxic micro-environments along the chronosequence, for example waterlogged soils, where oxygen availability is limited. Again, the presence of sulfate reducers may be a legacy from the subglacial environment, however, additional sampling of subglacial sediments would be needed to validate this. The location of sulfur reducing bacteria in more developed forefield soils compared to sulfur oxidizers, is interesting. An explanation for this is that sulfate reduction requires the availability of  $SO_4^{2-}$ , produced by oxidation, and therefore cannot occur in newly exposed nutrient deplete soils. However, additional measurements of sulfur species would be needed to validate this. Whilst sulfur reducing bacteria may be present in early soil samples,

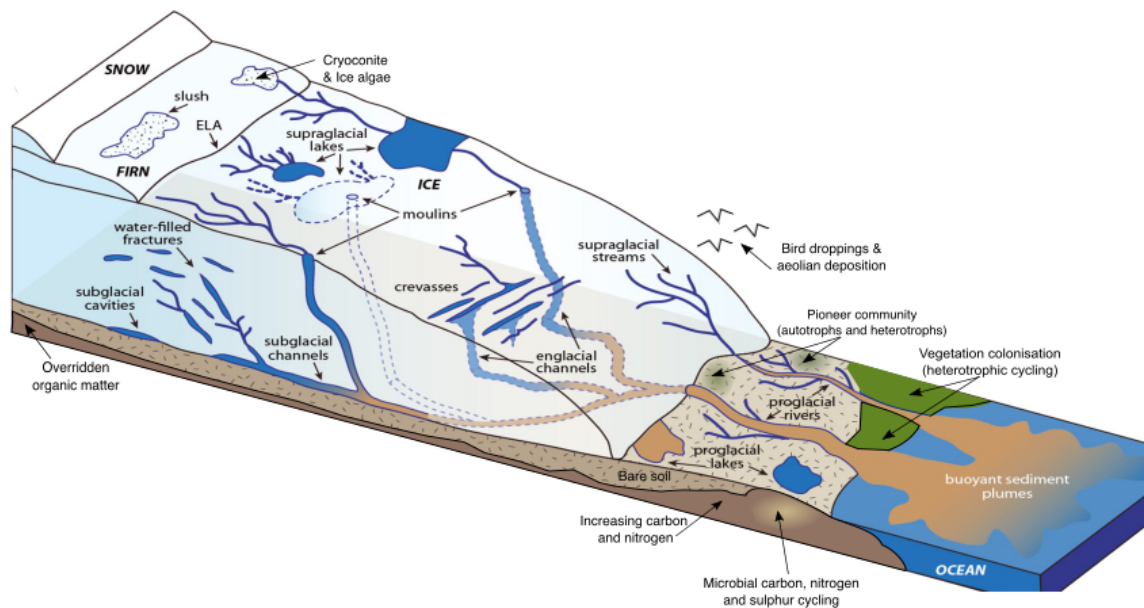
they were not recovered by the read based annotation carried out. Consequently, sulfur reducing microbes do not constitute a major fraction of the microbial community composition at early stages of soil succession in the Midtre Lovénbreen forefield. Additional research on the transcription of sulfur cycling genes (using transcriptomics) would be beneficial to highlight if the sulfur oxidation and reduction pathways were indeed active in the forefield.

#### 4.4.6 Unique genomes

Genome binning was carried out to investigate the presence of novel genomes in forefield soils, as expected by our hypothesis. The results of this analysis have indicated a degree of novelty of genomes extracted during the binning process, with 21% of bins falling below 80% identity to NCBI GenBank (Table 4.7). Of these bins, three matched most closely to the denitrifying bacteria, *Conexibacter woesei* and two were associated with the sulfur oxidizing *Sulfuricaulis limicola* (Table 4.7). This indicates the samples may contain unique species with ecological importance in local biogeochemical cycles, for example in nitrogen and sulfur cycling. The low percentage identity of these bins is related to using incomplete genome databases to perform taxonomy assignments (Albertsen *et al.*, 2013). As the complete global microbial diversity has not been captured on genome databases, it is likely that some samples may not match closely to the genomes that are available (Albertsen *et al.*, 2013). In the future, developments in NGS technologies, analysis pipelines and methodologies such as metagenomics will allow more environmental microbial diversity to be profiled in online databases. However, the results of this analysis do provide scope for further single cell sequencing and culture based studies, to fully ascertain the novelty of genomes found during the binning process.

### 4.5 Conclusions

This study has used metagenomics to explore the microbial community composition along a chronosequence of soil succession in the Midtre Lovénbreen forefield, Svalbard. Prior to this study, most research on forefields had focused on plant succession, changes to soil physicochemical properties or utilized 16s rRNA or single gene sequencing. For the first time, this study implements metagenomics to understanding the succession of microbial communities in a glacial forefield. We aimed to test several hypotheses: (1) community composition and function would modify with soil development; (2) autotrophic microbes would be present in newly exposed soils, with heterotrophic nutrient cycling occurring in later sites; and (3) extremophilic or novel microbes may be identified due to the harsh environmental conditions in Arctic forefields.



**Figure 4.6:** Schematic outlining proglacial features and processes described in this chapter. Key drainage features include: englacial, subglacial and proglacial pathways. Microbial features include: cryoconite and ice algae, the pioneer community and plant colonization. Soil nutrient content is displayed graphically, highlighting overridden organic matter and the buildup of labile organic carbon and nitrogen with succession. Nutrient sources are autochthonous (microbial fixation by autotrophs) and allochthonous (deposition). Source: The figure is an adaptation of Chu et al., (2014).

The key pathways and processes discussed in this analysis are highlighted in Figure 4.6. This study finds that along the successional chronosequence, both total organic carbon (TOC) and total nitrogen (TN) increased, in line with soil development. This is consistent with previous findings from glacier forefields, which identify TN and TOC pools to rise with soil development due to aeolian deposition, microbial fixation and the establishment of plants. This study also finds a range of extremophilic microbes to be present along the forefield, which may be adapted to the high UV, oligotrophic, cold conditions present, in line with our hypothesis. Additionally, we recovered cyanobacterial sequences from newly exposed soils and cryoconite, in line with previous forefield literature. This may be related to the ability of cyanobacteria fix carbon (and often nitrogen) in low nutrient conditions, alongside the secretion of protective exopolymeric substances (EPS) to buffer the harsh environmental conditions. The composition of the initial pioneer microbial community has previously been the subject of debate, in particular, the prevalence of heterotrophs in newly exposed soils. Our findings conclude that both autotrophic and heterotrophic bacteria were present in the initial pioneer community. This may be attributed to the availability of an initial TOC pool, from

overridden soils or aeolian deposition, which heterotrophic bacteria could utilize. Additionally, heterotrophic and autotrophic microbes identified in the pioneer community were also recovered from glacier ice samples. Consequently, we conclude that the community composition of newly exposed soils may be mediated by the microbes supplied in runoff from the glacier surface or from subglacial systems. In response to the debate on pioneer community composition, we suggest the presence of heterotrophic bacteria in newly exposed soils is likely to vary between glaciers, due to differences in the ice surface microbial communities, aeolian deposition and overridden TOC stocks.

The use of metagenome sequencing in this study enabled an investigation into the carbon, nitrogen and sulfur cycling pathways along the chronosequence, through read annotation, genome binning and SEED functional pathway annotation. The use of functional genes provides evidence for microbial biogeochemical cycling potential, which is not possible with taxonomic markers such as 16s rRNA. Our study concludes that forefield soils contain the potential for microbial carbon fixation, heterotrophic carbon cycling, nitrogen fixation, nitrification, denitrification, alongside sulfur oxidation and reduction. We show that microbes associated with heterotrophic carbon, nitrogen and sulfur cycling were recovered from older soils, as these pathways require labile nutrient stocks which may not be present in newly exposed soils, in agreement with our hypothesis. The use of metagenomics in our study also allowed the recovery of discrete genome bins from the samples. Whilst the genomes extracted were not comprehensive of the complete diversity (due to incomplete assembly and read coverage), they provide an insight into the functional potential of microbes in forefield samples. Several of these genome bins did not align closely to those on NCBI GenBank, indicating a degree of genome novelty. In particular, several genomes related to denitrifying and sulfur oxidizing microbes were recovered, highlighting these pathways as potentially prevalent in the forefield soils. Whilst extracting complete draft genomes was not the aim of this analysis, due to sequencing limitations, the results demonstrate the potential of metagenomics to explore functionality on a single genome scale, given deep sequencing.

Overall, this study has provided an insight into the diversity and metabolic potential of microbial communities along a successional chronosequence. It is hoped that this work will help to stimulate further research exploring functional activity rates, gene expression using transcriptomics and probe deeper into genome novelty using single cell sequencing and culture based studies.

#### 4.6 Limitations and Future Work

To build on the results of this study, additional deep metagenome sequencing would be beneficial in under sampled sites (Figure 4.2). This would enable less abundant microbes to be recovered and thus fully profile the microbial community composition in each site. Additionally, deep sequencing would provide more read coverage for each microbe, thus enabling a greater number (and more complete) genomes to be extracted during the binning process (Albertsen *et al.*, 2013). Consequently, this may provide the opportunity to build draft metagenome-assembled genomes, enabling an insight into genome structure, diversity and novelty (Albertsen *et al.*, 2013; Hugerth *et al.*, 2015). In particular, building near complete draft genomes would enable a comparison of the average nucleotide identity (ANI) between sample sequences and previously sequenced genomes (Konstantinidis and Tiedje, 2005). This information would be beneficial to fully ascertain how unique our genomes are in comparison to those available on public databases. Draft microbial genomes may also be extracted using culture based studies and single cell sequencing, focused on those genomes highlighted in the binning process (Table 4.7). The benefit of single cell sequencing would be an increased coverage for the target genome, increasing the likelihood of a fully closed genome assembly (Blainey 2013; Shapiro *et al.*, 2013). Additionally, culture based studies may be beneficial, to understand the life cycle and growth conditions of any newly isolated genomes (Marx 2017).

Furthermore, (meta) transcriptomics could also be applied to each sampling site, to highlight which metabolic pathways are active. As transcriptomics involves sequencing the transcribed mRNA, it goes one step further from metagenomics, as it identifies which processes are active at a snapshot in time (Moran *et al.*, 2013). This would be beneficial to understand which aspects of biogeochemical cycles are functioning and any changes along the forefield. Finally, metagenomics supplemented with 16s rRNA analysis may also be useful. Not only would this enable a comparison between the taxonomy conclusions of each method, but the addition of 16s analysis would enable the less abundant fractions of the microbial community to be profiled (Franzosa *et al.*, 2015). The reduced cost and ability of this method to fully profile the community composition enables further diversity metrics and gene abundance calculations to be applied. Finally, a subglacial sample from Midtre Lovénbreen would have been beneficial for this analysis. This would allow a comparison between newly exposed soils, to ice surface and subglacial systems, to identify how influential supraglacial or subglacial microbes are on the composition of the pioneer microbial community.

# Chapter 5: Sequencing the deep - microbial communities recovered from benthic fjord sediments, Chilean Patagonia

Nash, M.V<sup>1</sup>., Anesio, A.M<sup>2</sup>., Barker, G<sup>3</sup>., Wadham, J<sup>1</sup>., Tranter, M<sup>1</sup>., Hawkings, J<sup>4</sup>., Beaton, A<sup>5</sup>., Chin Ng, H<sup>6</sup>., and Sánchez-Baracaldo, P<sup>1</sup>.

<sup>1</sup>School of Geographical Sciences, University of Bristol, BS8 1SS, UK

<sup>2</sup>Department of Environmental Science, Aarhus University, PO box 358, Denmark

<sup>3</sup>School of Life Sciences, University of Bristol, BS8 1TQ, UK

<sup>4</sup>Department of Earth, Ocean and Atmospheric Science, Florida State University, Tallahassee, FL 32304, USA

<sup>5</sup>National oceanography Centre, European Way, Southampton, SO14 3ZH, UK

<sup>6</sup>Department of Earth Sciences, University of Bristol, BS8 1RL, UK

## Contributions and acknowledgements

MN carried out the sampling, analysis and interpretation of the sediments data, alongside the writing of the manuscript. AB supplied the underway sampling data. JH, JW and HCN assisted with funding, fieldwork logistics and sampling. PSB and AA assisted with the planning and refinement of the manuscript.

## Funding

This work was carried out as part of the NERC GW4+ Doctoral Training Partnership, awarded to MN. Funding support was provided through a NERC grant NE/J02399X/1, awarded to AA and a Royal Society University Research Fellowship, awarded to PSB. Additional funding support came from the PISCES project and a PISCES Knowledge Exchange Scholarship, awarded to MN.

## 5.1 Introduction

The fjord systems of Chilean Patagonia present a unique habitat for microbial life, draining glacial melt water into marine fjord systems. These fjords have been hypothesized as a hotspot of primary productivity at the land-ocean interface and are crucial for understanding coastal ecosystem functioning and diversity (Iriarte *et al.*, 2007). The fjord systems of Chilean Patagonia host three UNESCO bio-reserves and support commercially important Salmon fisheries (Haussemann and Forsterra, 2009; Niklitscheck *et al.*, 2013). High rates of primary productivity have been observed in Chilean fjords, alongside harboring diverse and unique ecosystems, such as the Patagonian cold-water corals (Iriarte *et al.*, 2007). However, the microbiology of the fjord sediments, crucial for understanding the wider ecosystem functioning, has received limited research attention. These fjords are particularly of interest due to the

interaction of glacial meltwater from Patagonia ice fields with marine waters, harboring a range of physico-chemical conditions for microbial life.

The Patagonian ice fields drain into downstream fjords, influencing the fjord's physical characteristics, with implications on primary productivity (PP) (Iriarte *et al.*, 2014). The fjords are also supplied by marine Sub-Antarctic water, which mixes in the inner fjords with freshwater runoff from terrestrial and glacial rivers (Palma and Silva, 2004; Iriarte *et al.*, 2014). The greater density of the marine waters means that a vertical stratification of the water column is often observed (Palma and Silva, 2004). Consequently, Patagonia fjords often contain a surface freshwater lens, over more dense marine waters, promoting a range of dynamic conditions for endemic and novel organisms (Iriarte *et al.*, 2010). Corresponding vertical gradients in oxygen, salinity, nutrients and light have been observed, all of which have implications on PP (Iriarte *et al.*, 2014). Oxygen can reduce to near a hypoxic state with depth, caused by the oxidation of organic matter in the inner fjords (Gonzalez *et al.*, 2013; Iriarte *et al.*, 2014). Horizontal gradients in oxygen have also been observed, reducing towards the inner fjords and glacier outflows (Davila *et al.*, 2002; Silva and Vargas, 2014). This relates to the supply of oxygen from the Sub-Antarctic water, which is depleted by microbial consumption towards the fjord head (Silva and Vargas, 2014). Consequently, deep waters in the inner fjords are more likely to experience anoxic conditions.

Research on the hydrochemistry of Patagonia fjords indicates that the glacial meltwater is low in nutrients, and that mixing plumes from the Sub-Antarctic water may help to stimulate phytoplankton blooms (Aracena *et al.*, 2011; Gonzalez *et al.*, 2013; Montero *et al.*, 2017). Additionally, the inner fjord regions, dominated by glacial runoff, typically have lower nitrate and phosphate than the distal marine waters (Aracena *et al.*, 2011). The glacial runoff is sediment dense, largely attributed to inorganic matter from glacial weathering (Aracena *et al.*, 2011). Alongside the low nutrient content of glacial runoff, the high sediment yield may block out light for photosynthetic microorganisms (Aracena *et al.*, 2011; Landaeta *et al.*, 2012; Meerhoff *et al.*, 2013). Consequently, the biological export from inner fjord regions can be reduced (Silva, 2008). However, the fjords are also fed by terrestrial rivers which contain increased loads of organic matter, silicon, nitrate and phosphate to fjord surface waters (Mayer *et al.*, 1998; Tréguer *et al.*, 2013; Gonzalez *et al.*, 2013).

On average, Patagonian glaciers discharge 70 km<sup>3</sup> of freshwater a year into downstream fjords and this flux is increasing with glacier thinning (Lenaerts *et al.*, 2014; Schaefer *et al.*, 2017). The implications of a greater meltwater flux on the microbial communities is hard to ascertain, particularly because fjord PP is a result of complex interactions between light,

nutrients and temperature, all of which may be influenced by glacial runoff (Landaeta *et al.*, 2012). For example, increased sediment flux may reduce water column and benthic microbial activity due to increased light attenuation, with implications propagating up the food chain to fish populations (Landaeta *et al.*, 2012; Gonzalez *et al.*, 2013). Additionally, a study by Gutiérrez *et al.*, (2015) suggests that a change in hydrographic conditions from increased glacial meltwater may influence the microbial community structure, favoring freshwater dominant species such as nano and pico plankton. Consequently, a greater understanding of the current microbial community composition and function in these fjords is needed, to help understand how the ecology and ecosystem services may change in the future.

Biological analysis of the fjords has focused on the water column, and in particular, on plankton activity and diversity (Iriarte *et al.*, 2007; Gonzalez *et al.*, 2013). Ecological indicators have classified benthic microbial communities at a 'good' status, however communities may be unbalanced due to the high outflow of glacial meltwater into the fjords (Quiorga *et al.*, 2013). The overall water column net primary productivity (NPP) and export production has been shown to increase when moving out from the inner fjords, largely related to the high sediment flux and freshwater in inner fjords (Aracena *et al.*, 2011; Gonzalez *et al.*, 2013). Research using 16s rRNA amplicon sequencing has been carried out on the water column microbial communities (Gutiérrez *et al.*, 2015). This research indicates that the glacial meltwater influences the microbial community structure in inner fjords, increasing the dominance of freshwater and cold adapted communities (Gutiérrez *et al.*, 2015). Consequently, meltwater drainage has been proposed as a key influence on biological activity and diversity in Patagonian fjords.

Chilean fjords present an interesting opportunity to study the influence of glacial meltwater on marine habitats, as they form the intersection between the land and open ocean. Currently, little is known about the diversity and function of microbial communities in these fjord environments. This information will aid our understanding of microbial diversity, fjord biogeochemical cycling and may help inform predictions to how these functions may change with increased ice melt in future years. This is especially important in Chilean Patagonia, given the presence of commercially important Salmon fisheries, which rely on fjord ecosystem functioning. Additionally, the interaction of saline and fresh waters, alongside dark and cold conditions within the sediments may harbor novel and ecologically significant microbes. Techniques such as metagenomics and genome binning, which enable both the functional and taxonomic diversity of uncultured microbial communities to be uncovered, have yet to be applied to this environment. These approaches may be highly beneficial for answering questions on the diversity, importance and uniqueness of communities in this region.

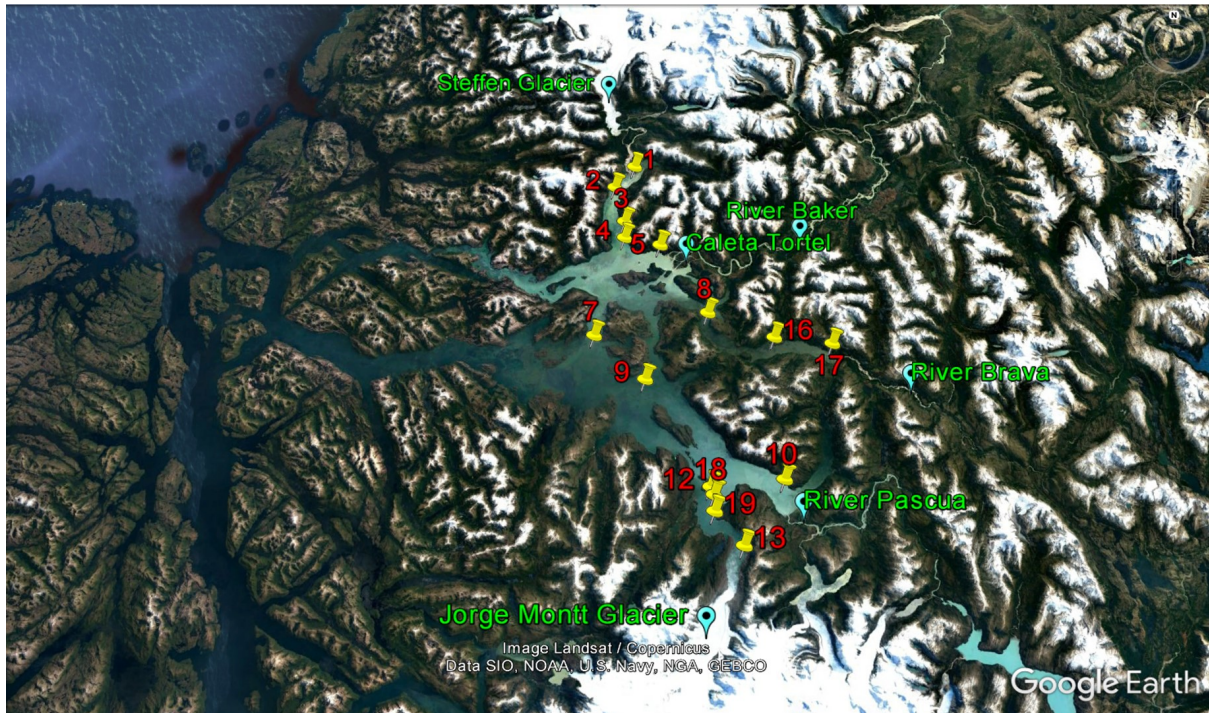


This study applied metagenomics to understand the taxonomic diversity of uncultured benthic sediment metagenomes from Patagonian fjords. The uncovered taxonomic diversity was used to drive genome binning and functional analysis, providing insights into the novelty, metabolic potential and ecological diversity of these microbial communities. The findings of Chapter 2 are used to inform metagenome assembly for functional annotation, however the limitations of assembly highlighted by Chapter 4 are also considered. The findings aim to provide an insight into this unique environment, whilst highlighting areas where targeted single cell genome sequencing and culture-based studies may be beneficial.

## **5.2 Methods**

### **5.2.1 Field Sampling**

Sampling was undertaken to investigate the microbial community composition of the benthic fjord sediments along the Steffen fjord and the outflow to the Baker channel, Chilean Patagonia. A total of 17 sites were investigated by the PISCES project for temperature, oxygen, salinity, pH and turbidity with depth using a CTD with a microfluidic colorimetric analyzer (Figure 5.1). From these sites, 5 were selected for analysis of benthic sediment microbial community composition (sites 1,2,4,5 and 7, Table 5.1). For these sites, integrated sediment grab samples were taken in a Van Veer grab sampler and stored in sterile Eppendorf tubes. Samples were frozen at -20°C for analysis at the University of Bristol. Temperature, oxygen, salinity, pH and turbidity of the fjord network surface waters were measured using underway sampling. This data was recorded using an EXO 2 sonde during the boat cruise track, which provided continuous measurements of the fjords. This data is used here to supplement the microbial community composition data obtained from the sediment samples from Sites 1,2,4,5 and 7.



**Figure 5.1:** Sampling sites for aqueous nutrients, salinity, pH and dissolved oxygen. Sampling sites for sediment microbial community composition were 1,2,4,5 and 7. Source: Google Earth

**Table 5.1:** Benthic sediment sampling site locations

Sample	Coordinates South	Coordinates West	Date sampled
1	S47 38.373	W73 40.046	18/02/2017
2	S47 40.748	W73 42.897	20/02/2017
4	S47 46.396	W73 41.770	20/02/2017
5	S47 47.263	W73 36.512	16/02/2017
7	S47 56.756	W73 45.839	17/02/2017

### 5.2.2 Metagenome DNA extraction and sequencing

DNA extraction and sequencing was carried out in line with the protocols detailed in Chapters 3 and 4 (Nash *et al.*, 2018). DNA for metagenomics was extracted using the same protocol applied in Chapter 3 and 4 (section 3.2.3; Nash *et al.*, 2018). Metagenomes were sequenced using an Illumina Next-Seq 500, with a TruSeq library prep kit at the University of Bristol Genomics facility (Nash *et al.*, 2018). A total of 5 metagenomes were sequenced (one metagenome per site) using 2x 150bp reads.

### 5.2.3 Metagenome read annotation

The 5 sequenced metagenomes were quality trimmed using Trimmomatic V0.38, with quality checks carried out using FASTQC v0.11.7 (Bolger *et al.*, 2014; Babraham Bioinformatics). Read based taxonomic annotation was carried out in Kaiju v1.5.0 (Menzel *et al.*, 2016). Kaiju is a kmer based approach which utilises protein sequences, which are more highly conserved when compared to nucleotide sequences (Menzel *et al.*, 2016). This method has been shown to be well suited to novel and divergent metagenome samples (Menzel *et al.*, 2016). Interactive hierarchical microbial community structures of the annotated samples were obtained in Krona, using the KBASE platform (Ondov *et al.*, 2015; Arkin *et al.*, 2016).

### 5.2.4 Metagenome assembly, genome binning and annotation

The sequenced reads for each metagenome were assembled using metaSPAdes V3.11.1 with kmer length optimisation and reads mapped back to the metagenomes using Bowtie 2 v2.3.2 (Langmead and Salzberg, 2012; Nurk *et al.*, 2017). The assembled metagenomes were imported in to JGI IMG/ MER for annotation and gene calling (Chen *et al.*, 2017). Genome binning was used to extract novel genomes from the assembled metagenomes, using MaxBin2 v2.2.4, which clusters genomes based on sequence coverage and tetra-nucleotide frequencies of assembled contigs (Bankevich *et al.*, 2012; Wu *et al.*, 2015). The genomes were extracted using BinUtil v1.0.1 on KBASE (Arkin *et al.*, 2016). The taxonomic identity of the extracted bins were assigned using a BLAST-n search against all complete NCBI RefSeq genomes (e-value  $10^{-5}$ ). The top BLAST hit for each bin was used to assign the taxonomy, based on the BLAST max score value, with an E-value below  $1 \times 10^{-5}$ . Thresholds of 95 and 85 % identity or better to classify bins to species and genus level were used, in line with analysis carried out by Camparano *et al.*, (2016). Bins classified at below 80% identity were proposed as potentially unique, as they cannot be accurately placed given the current available DNA sequences.

### 5.2.5 16s rRNA and dsrAB phylogenies

To demonstrate the potential capabilities of metagenome analysis, 16s rRNA and dissimilatory sulfite reductase (dsrAB) phylogenies were carried out, based on the results from genome binning. Whilst other explorative analyses and phylogenies could be carried out, these were selected to provide an example for how metagenome sequencing can be used.

The archaeon genus *Nitrosopumilus* was selected for phylogenetic analysis based on the 16s SSU rRNA marker gene. This is because the genus was abundant, accounting for 10-26% of genome bins across the samples and may have biogeochemical significance in the sediment nitrogen cycle. 16s rRNA sequences from reference *Nitrosopumilus* genomes were obtained from NCBI GenBank, based on the phylogeny of Qin *et al.*, (2017). *Nitrosopumilus* 16s sequences were extracted from the assembled metagenome samples using JGI IMG/MER, using BLASTn, with an e-value of  $10^{-5}$  (Chen *et al.*, 2017). As the assembled metagenomes incorporated unamplified DNA from complex sediment microbial communities, the 16s sequences obtained may not cover the complete set of *Nitrosopumilus* genomes in the samples (Rodriguez and Konstantinidis, 2014). However, those obtained through BLAST searching are likely to be those most prevalent in the metagenomes, as these have been recovered by sequencing (Rodriguez and Konstantinidis, 2014). Metagenome and reference sequence alignments were created in SATé 2.2.7, using MAFT, MUSCLE and FASTTREE with the GTR+CAT model, in line with Chapter 3 (Liu *et al.*, 2011; Nash *et al.*, 2018). Manual inspection of the 16s rRNA sequence alignment and production of Phylip files was carried out using Mesquite 3.2 (Maddison and Maddison, 2017). A 16s maximum likelihood phylogeny was generated using RAXML-HPC2 8.2.10 on XSEDE through the CIPRES Science Gateway, with 1000 bootstrap iterations, implementing the GTR+G nucleotide substitution model (Stamatakis, 2014; Nash *et al.*, 2018). Final trees were produced using Figtree 1.4.3 and visual modifications were made in Inkscape 0.91. Similarity evaluation between *Nitrosopumilus* 16s sample sequences and NCBI GenBank relatives were made using NCBI BLASTn.

Additionally, genome binning and taxonomic analysis of the metagenomes highlighted the presence of sulfur cycling microbes. As metagenomes contain functional genes (alongside taxonomic markers such as 16s rRNA), the functional gene *dsrAB* (dissimilatory sulfite reductase) for sulfite reduction was explored phylogenetically to demonstrate the potential of using metagenomes for functional exploration. This gene encodes the reduction of sulfite ( $\text{SO}_3^{2-}$ ) to sulfide ( $\text{S}^{2-}$ ) in anaerobic respiration by both bacteria and archaea (Müller *et al.*, 2015). Reference *dsrAB* sequences were obtained from NCBI GenBank based on the phylogeny of Moreau *et al.*, (2010). Sequences were obtained from assembled metagenomes with IMG/MER using BLASTn with an e-value of  $10^{-5}$ , with the alignment and phylogeny generated as outlined above. The metagenome *dsrAB* sequences were compared to nearest relatives using BLASTn against NCBI GenBank.

#### 5.2.6 Sediment organic carbon and total nitrogen content

Sediment total nitrogen (TN) and total organic carbon (TOC) were determined in triplicate for each sample site, using mass spectrometry. The protocol demonstrated in Chapters 3 and 4 was applied (section 3.2.2; Nash *et al.*, 2018).

### 5.3 Results and Discussion

#### 5.3.1 Fjord metadata and influence of glacial meltwater

The results of underway sampling the surface waters of the Steffen fjord and Baker channel for temperature, oxygen, salinity, pH and turbidity are shown in Table 5.2 and Figure 5.2. The results of depth profiles for each site are shown in Table 5.2 and Figure 5.3.

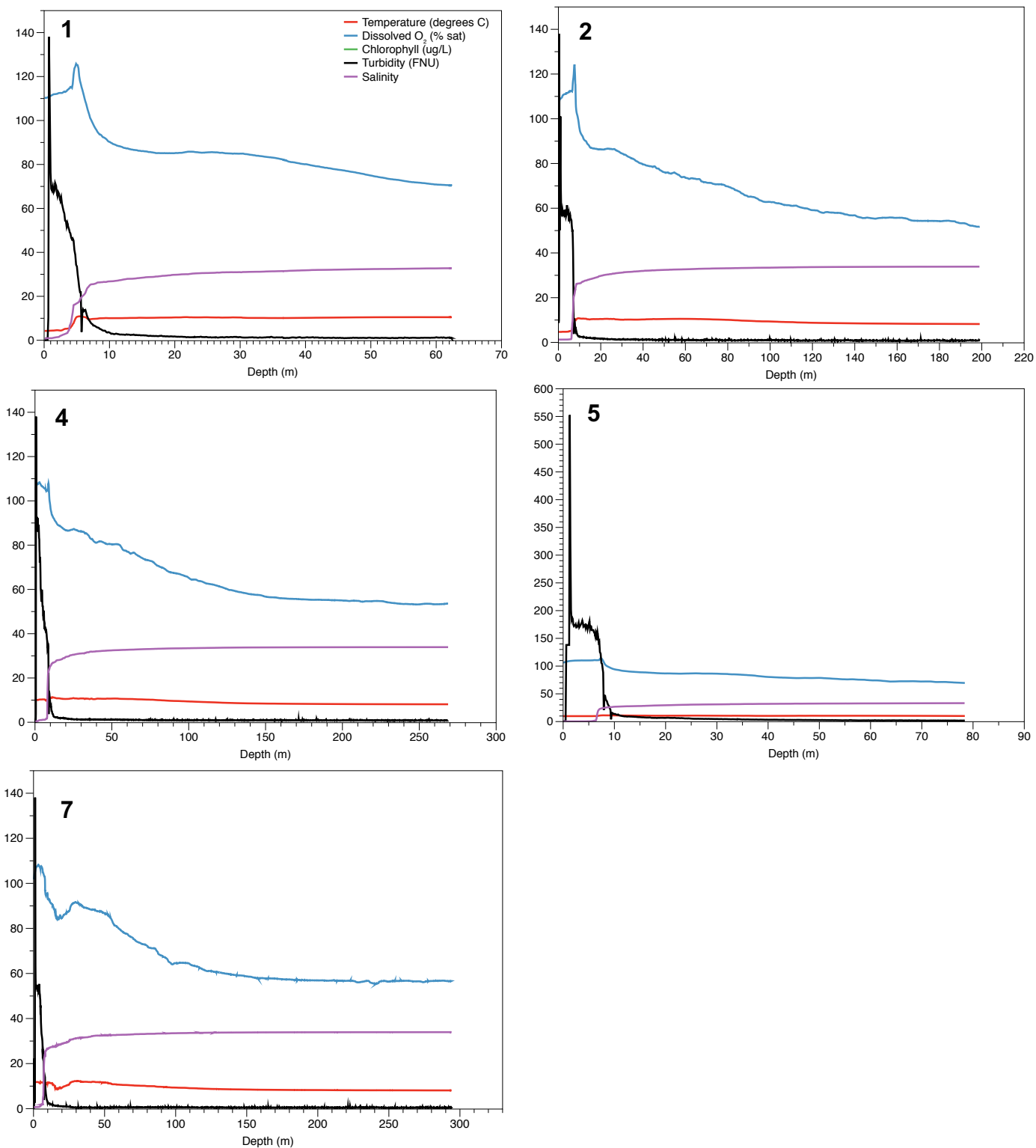
**Table 5.2:** Metadata for benthic sediment sampling sites, courtesy of the PISCES project.

Site	Water depth (m)	Water temperature (°C)	Surface water dissolved oxygen (% sat)	Bottom water salinity	Surface water pH	Turbidity (FNU)
1	61	10	104	32	6.76	40
2	199	8	104	34	7.65	38
4	268	8	102	34	7.98	50
5	78	8	104	34	7.65	50
7	360	9	104	34	7.71	30

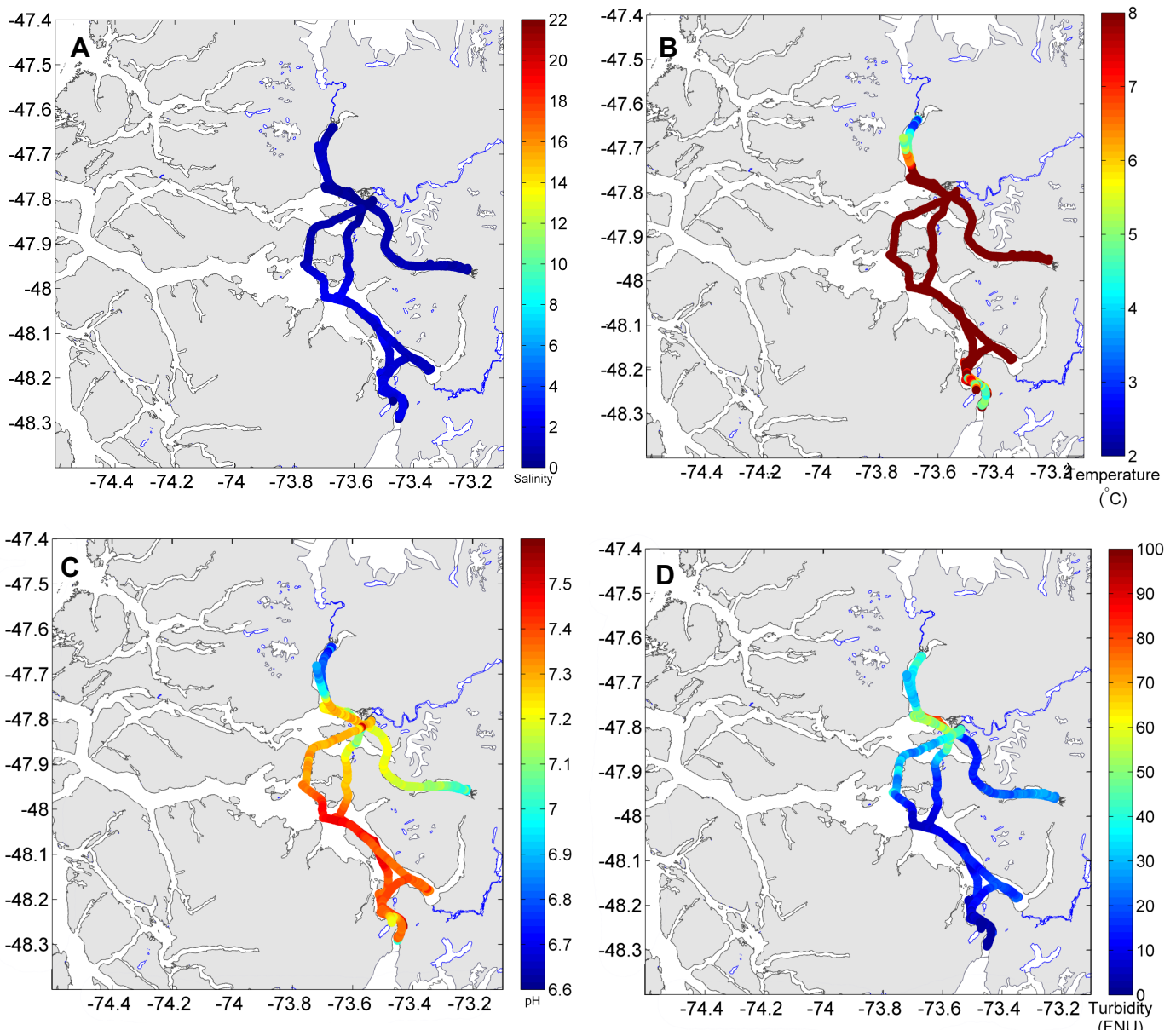
For sediment sampling sites, water depth ranged between 61 - 360m, between sites 1 and 7. This reflects the movement from shallower inner waters of the Steffen Fjord, to more distal, marine dominated sites (Iriarte *et al.*, 2014). Surface water temperature ranged between 8 - 10°C and displayed limited variation with depth using a CTD sensor (Table 5.2; Figure 5.2). Surface water temperatures reduced at sites 1 and 13 due to the outflow of the Steffen and Jorge Montt glaciers (Figure 5.1; Figure 5.3). Surface water pH ranged between 6.8 – 8 at sediment sampling sites, with the lowest pH found at sample site 1, at the outflow of the Steffen glacier into the fjord (Table 5.2). Surface water salinity ranged between 0.4 – 1.3 and increased to 32 – 34 in bottom waters (Table 5.2; Figure 5.3; Figure 5.4). This reflects the distinct salinity stratification in the Steffen fjord, whereby fresh terrestrial and glacial runoff largely remained at the fjord surface, over a layer of dense saline marine water (Pickard, 1971; Iriarte *et al.*, 2014). This indicates a lack of mixing between the two water masses (Iriarte *et al.*, 2014).

Surface water turbidity ranged between 30 – 50 FNU, however reduced to 0.4 – 1.2 FNU at the sediment bed (Table 5.2; Figure 5.2). This again reflects the distinct fjord stratification, whereby surface waters contain more suspended sediment from terrestrial runoff than marine bottom waters (Iriarte *et al.*, 2014). The turbidity of surface waters showed distinct variation in the fjord network, increasing from 0 FNU in the Baker channel to 40 FNU in the Steffen fjord (Figure 5.3). This may relate to sediment flux from the terrestrially fed River Baker and glacial flour from the Steffen glacier, which is largely retained in the fresh surface waters of the Steffen fjord (Figure 5.1). Surface waters were saturated with oxygen, however this reduced with depth to 51 – 71% saturation at the sediment bed (Table 5.2; Figure 5.2). This is due to air-sea gas exchange with oxygen in surface waters maintaining atmospheric equilibrium, influenced by wind velocity (Broecker and Peng, 1974). This is not maintained with depth due to inadequate mixing and microbial consumption of oxygen (Leon-Munoz *et al.*, 2013).

The freshwater glacial runoff in the Steffen fjord and wider Baker channel was therefore shown to have distinct influences on the fjord physicochemical characteristics. Glacial and terrestrial runoff provided a less dense freshwater lens over more saline bottom waters throughout the fjord (Iriarte *et al.*, 2014; Figure 5.2). This freshwater lens had greater turbidity, oxygen content and lower salinity than benthic waters (Table 5.2; Figure 5.2). This information provides the physicochemical background for understanding the microbial community structure in the benthic sediment metagenomes sampled.



**Figure 5.2:** CTD depth profiles for sediment sampling cites 1,2,4,5,7. Temperature, dissolved oxygen, chlorophyll, turbidity and salinity are plotted with depth for each sampling site.



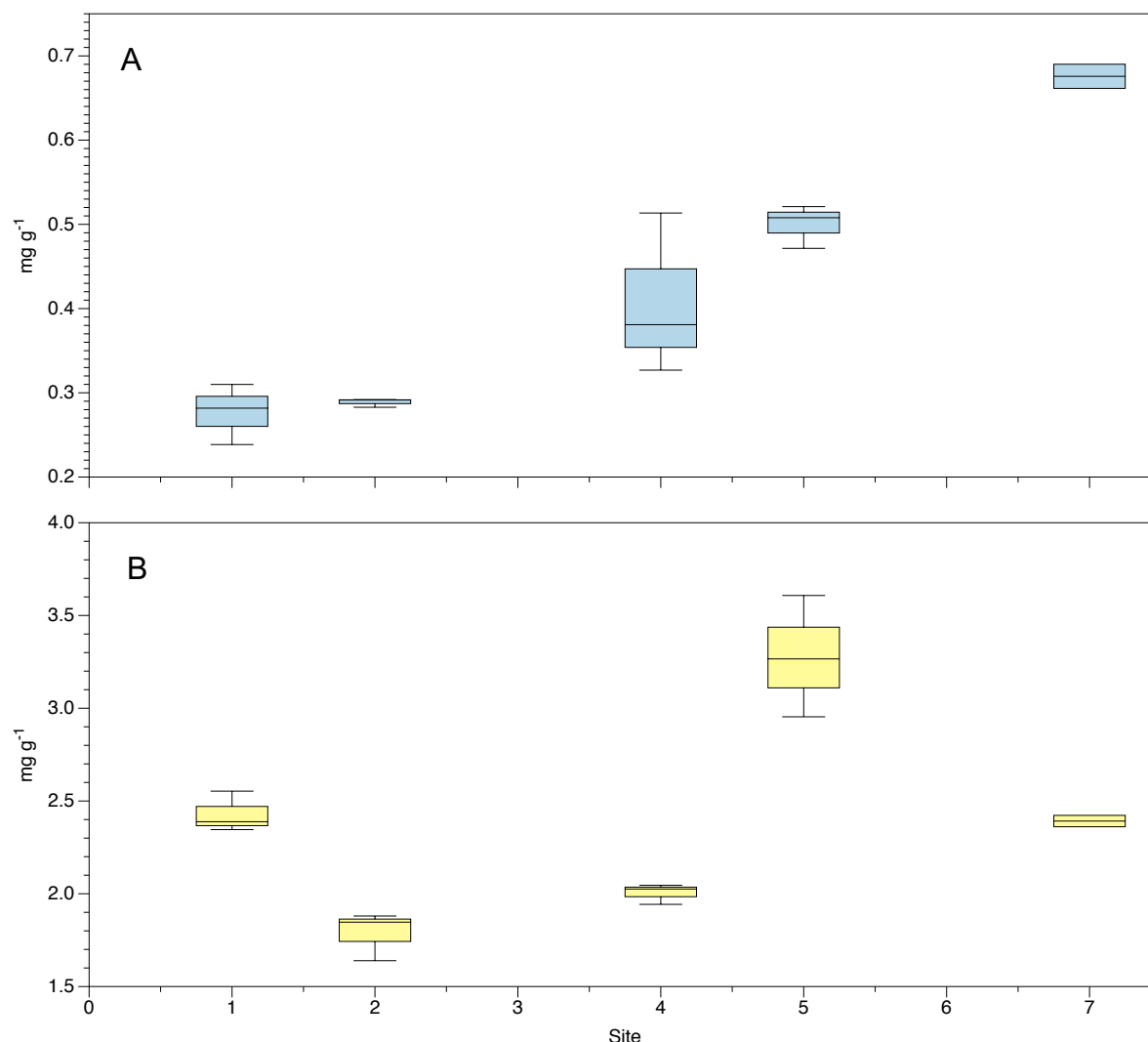
**Figure 5.3:** Underway sampling data of surface waters for (A) Salinity, (B) Temperature, (C) pH and (D) Turbidity. Data courtesy of Alex Beaton, PISCES project.

### 5.3.2 Sediment TN and TOC

The results for sediment total organic carbon (TOC) and total nitrogen (TN) are displayed in Figure 5.4. TOC ranged between  $3 \text{ mg g}^{-1}$  to  $13.5 \text{ mg g}^{-1}$ , from Site 2 to Site 17 respectively. The sites sampled for metagenomics along the Steffen Fjord (Sites 1-7) display a smaller range of values, between  $3 - 5.8 \text{ mg g}^{-1}$  TOC (Figure 5.4). The range of TN values from sediments was more constrained, ranging between  $0.2 - 1 \text{ mg g}^{-1}$  (Figure 5.4). The limit of detection was  $1 \text{ mg g}^{-1}$  for both TOC and TN. These values are indicative of low nutrient



conditions in the benthic fjord sediments, which may mediate the microbial communities which they can sustain. In particular, the minimal values of TN indicate nitrogen may be limiting, and thus may influence the diversity and structure of microbial life in the sediments.



**Figure 5.4:** A) total nitrogen (TN) and B) total organic carbon (TOC) per gram of sediment for metagenome sampling sites (1,2,4,5,7). The limit of detection was  $1 \text{ mg g}^{-1}$  for both TOC and TN. The relative standard deviation of measurement was 1.44% and 4.86% for TOC and TN respectively, calculated by measuring a series of 9 standards.

### 5.3.3 Metagenome assembly and annotation

Following sequencing, the five benthic sediment metagenomes were assembled to create longer contigs for functional annotation. Metagenome assembly has been shown to improve the annotation quality of the datasets, due to longer DNA sequence lengths for alignment-based interpretation (Nagarajan and Pop, 2013, Chapter 2). The output quality for metagenome assembly and read mapping are displayed in Table 5.3. The number of contigs ranged between 25,112 – 64,579, with contig size between 10,000 – 1,000,000 bp (Table 5.3).

The read coverage of the metagenomes ranged between 87.1- 95.4%, highlighting the majority of the raw read data was used in the assemblies (Table 5.3). However, following taxonomic and functional annotation in JGI IMG/MER 98.76 – 99.33% of reads were not assigned to a taxonomy at 90% identity (Table 5.4). This may highlight a substantial novelty of genomes in the samples which cannot be identified using current databases, as these do not fully profile current global microbial diversity (Ferrer *et al.*, 2005).

**Table 5.3:** Metagenome assembly statistics for microbial community metagenomes, highlighting assembly size and contig length, produced in Quast (Gurevich *et al.*, 2013). Output quality metrics for raw read mapping against the metagenomes are also provided.

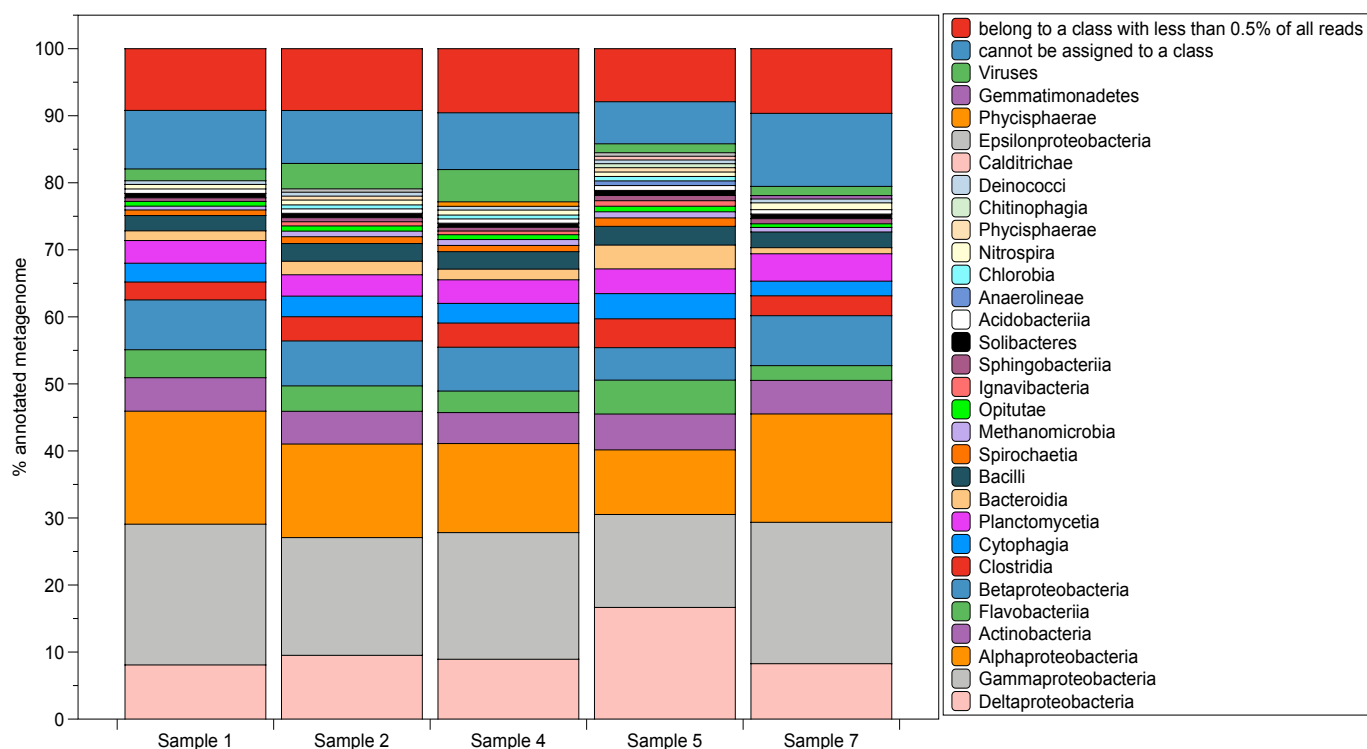
	Site 1	Site 2	Site 4	Site 5	Site 7
<b>Assembly parameters</b>					
# contigs	47703	44713	42395	64579	25112
# contigs (>= 0 bp)	47703	44713	42395	64579	25112
# contigs (>= 1000 bp)	47703	44713	42395	64579	25112
# contigs (>= 10000 bp)	2374	1581	2050	3227	946
# contigs (>= 100000 bp)	6	14	14	10	0
# contigs (>= 1000000 bp)	0	0	0	0	0
Largest contig	169641	297725	360451	227878	82647
Total length	197488635	174470052	175632561	268241056	95362614
Total length (>= 0 bp)	197488635	174470052	175632561	268241056	95362614
Total length (>= 1000 bp)	197488635	174470052	175632561	268241056	95362614
Total length (>= 10000 bp)	42735865	33490431	41852693	62069788	15473582
Total length (>= 100000 bp)	762550	1960800	2074273	1341724	0
Total length (>= 1000000 bp)	0	0	0	0	0
N50	4300	3842	4189	4236	3774
N75	2731	2602	2691	2725	2598
L50	11553	11316	9829	15263	6821
L75	26303	25409	23230	35496	14572
GC (%)	51.32	47.16	45.67	46.93	50.36
<b>Mapping parameters</b>					
Total Reads	174621110	161768224	164078036	155907364	164920816
Unmapped Reads	160588571	148029200	149463098	136710343	157328781
Mapped Reads	14032539	13739024	14614938	19197021	7592035
Singletons	958195	981434	853002	1273877	798535
% coverage	91.96	91.5	91.1	87.1	95.4

**Table 5.4:** Percentage of taxonomically unassigned assembled sequences at 60 and 90% identity, using JGI IMG/MER

Sample	% unassigned sequences (90% identity)	% unassigned sequences (60% identity)
1	98.97	68.35
2	99.15	80.8
4	99.33	83.46
5	99.24	75.94
7	98.76	65.19

Taxonomic annotation of quality trimmed (unassembled) sequencing reads was carried out following the limited scope of assembled metagenome annotation (Table 5.4). The read classifier Kaiju was selected, as this classifier has been shown to have a higher sensitivity to underrepresented genera than other kmer based classifiers (Menzel *et al.*, 2016). Furthermore, the read classification is based upon protein sequences, which are more conserved and resilient to sequencing errors than DNA based analysis (Menzel *et al.*, 2016).

The overall microbial community composition of the five metagenomes, at the class level, can be identified in Figure 5.5. The full classification from class to species level is available in Appendix 4, Tables 1-5. In total, only 25% of the sequenced reads could be annotated with a taxonomy at the class level (Figure 5.5). Of these reads, 20% could not be annotated at the class level, or were at a very low abundance (Figure 5.5). This indicates a presence of either novel, and/or highly fragmented DNA in the datasets. As current annotation techniques rely on the use of known sequences, any novel or unique organisms will not be assigned a taxonomy (Menzel *et al.*, 2016). This is particularly the case for environmental samples, as known sequenced genomes are often biased towards those which have a medical application (Menzel *et al.*, 2016). The use of binning techniques may help to delineate the presence of novel genomes in the samples. However, as the samples were derived from low nutrient sediments, it is possible that a large fraction of the DNA is degraded (i.e. not living) due to the harsh environmental conditions and exposure to water (Taberlet *et al.*, 2012; Bohmann *et al.*, 2014). As DNA is generally fragmented during the degradation process, the shorter length hinders the ability to annotate proteins, either by sequence alignment, sequence composition or tetranucleotide frequencies (Taberlet *et al.*, 2012; Menzel *et al.*, 2016). To reveal the composition of this more degraded DNA, amplification using short primers would be needed, for example using metabarcoding (Taberlet *et al.*, 2012; Leray and Knowlton, 2015) however this is beyond the scope of the current analysis.



**Figure 5.5:** Kaiju taxonomic classification of reads for the five sequenced samples, at the Class level. Only the annotated sequences are shown, which account for 25% of the total metagenome sequenced for each site. The percentage of the classified metagenome attributed to each Class is shown. Sequences that have been classified but cannot be assigned to a class and very low abundance reads, are also provided. The legend is ordered in line with the sequence of bars in the main plot. The full taxonomic classification from class – species level is available in Appendix 4, Tables 1-5.

From the classified reads, the five samples show a similar overall bacterial community composition, with Deltaproteobacteria, Gammaproteobacteria, Alphaproteobacteria and Actinobacteria dominant in all samples (Figure 5.5). The presence of organisms related to *Methanomicrobia*, *Chlorobia*, *Clostridia*, *Nitrospira* and *Anaerolineae* is indicative of anoxic conditions in the surface sediments, thus eliciting the detection of anaerobic organisms and the potential for methanogenesis (Eisen *et al.*, 2002; Luker *et al.*, 2010; Mackelprang *et al.*, 2011; Matsuura *et al.*, 2015). This aligns with the reduction in oxygen saturation with depth, which may continue below the sediment surface due to microbial activity in sediments and lack of oxygen replenishment (Figure 5.2; Oschmann, 2001). The microbial communities in the benthic fjord sediments were subjected to low nutrient, saline and dark conditions (Figure 5.4; Figure 5.5). Consequently, there may be evidence for adaptations within the bacterial community, as the environment may be highly selective for the organisms which can survive (Figure 5.5). For example, *Acidobacteria* (1.4 - 1.6% of metagenomes) have been shown to produce large amounts of exopolysaccharide (EPS), which is both protective and adhesive (Ward *et al.*, 2009). The adhesive properties of EPS can increase the nutrient uptake, which is beneficial in a low TN environment such as the Steffen fjord (Weiner *et al.*, 1995; Ward *et al.*, 2009). Additionally, *Acidobacteria* are facultative anaerobes and use ferric iron reduction

in the absence of oxygen for respiration (Blothe *et al.*, 2008). Facultative anaerobic respiration may be beneficial in the Steffen fjord given the potential for low oxygen conditions. Furthermore, taxa related to the aerobic chemoorganotroph, *Deinococci*, were identified in sediment samples (Copeland *et al.*, 2012). This organism is a known extremophile and can survive ionizing radiation, cold and oligotrophy (Copeland *et al.*, 2012). Additionally, evidence for the selection of resilient microbes is supported by the identification of the extremophile, *Epsilonproteobacteria* (Figure 5.5). These bacteria are commonly found in extreme environments such as hydrothermal vents and cold seeps and obtains energy through chemolithotrophy, thus can survive without organic compounds and light (Takai *et al.*, 2005). This may be significant given the low TOC in the sediment samples (Figure 5.4). This consequently provides additional evidence for the selection of well adapted resilient organisms to the environmental conditions in the benthic sediment of the Steffen fjord, if active.

The hierarchical structure of the Kaiju microbial community composition was visualized in interactive Krona plots for each site, to derive more detail from the classifications (Appendix 4 Tables 1-5; Table 5.5). Overall, Proteobacteria accounted for 46 - 55% of bacteria in the metagenomes and provided evidence for sulfur cycling potential within the fjord sediments (Table 5.5). The presence of *Desulfovibrionales* and *Desulfobacterales* indicates the potential for anaerobic sulfate reduction, producing sulfide ( $\text{H}_2\text{S}$ ) from sulfate ( $\text{SO}_4^{2-}$ ) (Kuever, 2014, Table 5.5). Sulfate reducing bacteria are commonly found in seawater due to the abundance of sulfate, alongside in anaerobic sediments, as they contribute to the degradation of organic matter (Goldharber and Kaplan, 1974). Additional evidence for sulfur cycling potential is provided through the identification of the purple sulfur bacteria *Chromatiales*, accounting for 3-6% of the metagenomes (Table 5.5). *Chromatiales* are typically anaerobic, and utilise the waste  $\text{H}_2\text{S}$  from sulfate reduction, and oxidise it to elemental sulfur (S) (Imhoff, 2005). Consequently, the sediment metagenomes provide evidence for microbial sulfur cycling potential within the sediments. In order to validate the activity of sulfur cycling, further work using transcriptomics could be implemented to identify the transcription of key marker genes, such as *dsrAB* and *apsA* for sulfate reduction (Wagner *et al.*, 2005).

Archaea accounted for 3-5% of cellular organisms within the metagenomes (Table 5.5). The depth of sequencing used in shotgun metagenomics is often suitable to profile bacterial communities, however may not have the coverage to isolate the archaea and eukaryotes present (Hugenholtz and Tyson, 2008). Deep sequencing is often required to reveal the full breath of these communities (Narasimgarao *et al.*, 2012). However, the archaea identified do provide support for the presence of anaerobic conditions at the sediment bed. The presence of the methanogens *Methanomicrobia*, *Methanobacterales*, *Methanococcales* and

*Thermoplasmata* indicate the potential production of methane (CH<sub>4</sub>), using carbon dioxide (CO<sub>2</sub>) as the terminal electron acceptor in anaerobic respiration (Table 5.5; Valentine, 2002; Bonin and Boone, 2006). Methanogenesis is significant for carbon cycling, acting as the final stage in the degradation of organic material following the preferential use of other electron acceptors such as oxygen, sulfate and nitrate (Zeikus, 1977; Valentine, 2002). Methane has been shown to be a substantial contributor to global climate change, having a warming potential greater than that of CO<sub>2</sub> (Valentine, 2002). The presence of methanogenesis may therefore have implications on both the local and global carbon cycle over long timescales, for both the degradation and recycling of organic matter, and production of CH<sub>4</sub>.

**Table 5.5:** Extract of *Kaiju* read based classification for the 5 metagenome samples, based on those discussed in this analysis. The classification name and percentage of metagenome assigned is provided. Full classification for each metagenome is available in Appendix 4 Tables 1-5.

	Classification rank	Sample 1 (%)	Sample 2 (%)	Sample 4 (%)	Sample 5 (%)	Sample 7 (%)
Bacteria	Domain	93	92	90	96	91
Archaea	Domain	5	4	5	3	7
Viruses	Domain	2	4	5	1	2
Proteobacteria	Phylum	55	49	49	46	54
Terrabacteria	Phylum	15	17	17	19	16
FCB group/ Sphingobacteria	Phylum	12	13	11	17	8
PVC group	Phylum	5	5	6	6	6
Thaumarchaeota	Phylum	3	2	3	3	5
Other	-	10	14	14	9	11
Methanomicrobia	Class	0.6	0.8	0.9	0.9	0.7
Epsilonproteobacteria	Class	0.4	0.5	0.5	0.5	0.4
Clostridia	Class	3	4	4	4	3
Bacilli	Class	2	3	3	3	2
Thermoplasmata	Class	0.08	0.1	0.2	0.1	0.1
Other	-	93.92	91.6	91.4	91.5	93.8
Desulfovibrionales	Order	0.7	0.9	0.9	1	0.8
Desulfobacterales	Order	1	2	2	7	2
Chromatiales	Order	5	4	4	3	6
Micrococcales	Order	0.08	0.8	0.7	0.8	0.7
Oscillatoriales	Order	0.7	0.5	0.5	0.5	0.6
Methanobacteriales	Order	0.1	0.2	0.3	0.2	0.2
Methanococcales	Order	0.09	0.2	2	0.1	0.1
Other	-	92.33	91.4	89.6	87.4	10.4
Nitrobacter	Genus	0.09	0.08	0.07	0.09	0.08
Rhizobium	Genus	0.7	0.3	0.3	0.3	0.4
Pseudomonas	Genus	1	1	1	1	1
Other	-	98.21	98.62	98.63	98.61	98.52

Finally, the metagenomes in this study provide evidence for nitrogen cycling potential within the sediments. The presence of organisms relating to *Thaumarchaeota* (3 – 5 % of the metagenomes), chemolithotrophic ammonia oxidisers, suggests that microbial nitrification

may be taking place within the sediments (Table 5.5; Park *et al.*, 2012). Ammonia oxidising archaea convert ammonia ( $\text{NH}_3$ ) to nitrite ( $\text{NO}_2^-$ ), which can be utilised by nitrite oxidisers in such as *Nitrobacter*, to convert  $\text{NO}_2^-$  to nitrate ( $\text{NO}_3^-$ ) for assimilation (Jetten, 2008). In addition to this, the presence of *Epsilonproteobacteria* provides evidence for denitrification potential within the sediments, converting  $\text{NO}_3^-$  to nitrogen gas ( $\text{N}_2$ ) and is a key pathway for fixed nitrogen removal from sediments (Murdock and Juniper, 2017). Nitrogen-fixing bacteria such as *Clostridia*, *Rhizobium* and  $\text{N}_2$  fixing cyanobacteria such as *Micrococcales*, and *Oscillatoriales* were identified in the sediment metagenomes (Table 5.5). These organisms may be instrumental in providing labile fixed nitrogen into the sediments, to be utilized by heterotrophic bacteria for protein assimilation (Bergman *et al.*, 1997). If the nitrogen fixers were active in the sediments, they may be crucial for facilitating heterotrophic microbial communities, given the oligotrophic conditions identified (Figure 5.4). Finally, ammonification is performed by bacteria to convert organic nitrogen to more bioavailable ammonium ( $\text{NH}_4^+$ ) (Gruber, 2008). Evidence for the potential of this pathway is shown through the presence of organisms relating to *Pseudomonas* and *Bacilli* in the sediments (Table 5.5). Consequently, the metagenomes sampled contain organisms capable of the complete nitrogen cycle, despite the low nitrogen concentrations sampled (Figure 5.4). If active, the nitrification and denitrification pathways may be tightly coupled to microbial nitrogen fixation to obtain sufficient nitrogen stocks.

#### 5.3.4 Genome binning

Genome binning was carried out using assembled metagenomes, to isolate discrete genomes which may be unique compared to those currently available in reference datasets. The results for each metagenome are shown in Tables 5.6 – 5.10, identifying the genome bins (discrete genomes) alongside the BLASTn match of each genome to NCBI GenBank. Metagenome 1 contained 36 genome bins, with 67% of bins containing less than 80% identity to cultured relatives on NCBI GenBank (Table 5.6). This indicates that the metagenomes may contain several novel species and/or strains, which have yet to be cultured or fully sequenced. This pattern was continued within the remainder of the metagenomes, with 58%, 55%, 71% and 63% of genome bins below 80% identity to GenBank for samples 2,4,5 and 7 respectively (Tables 5.6 – 5.10). This highlights a need for further culture-based studies or single cell sequencing, to isolate complete genome sequences for these potentially novel organisms. The presence of novel sequences may relate to the dark, cold and low nutrient conditions within the fjord sediments, lack of previous exploration of these sediments and the incomplete nature of current DNA reference databases.

In line with the taxonomy, the genome bins obtained were broadly similar between the metagenomes, however Sample 5 contained a wider range of genomes than Sample 7 (Tables 5.6 – 5.10). The genome bins predominately spanned the bacterial phyla Proteobacteria, Cyanobacteria, Bacteroidetes, Chloroflexi, Firmicutes and Verrucomicrobia, alongside the archeal phyla Thaumarchaeota and Euryarchaeota (Tables 5.6 – 5.10). The Thaumarchaeota *Nitrosopumilus* was common between metagenomes, accounting for between 10 – 26% of genome bins for each site. The binned *Nitrosopumilus* sequences contained between 75 – 98% identity to NCBI GenBank reference genomes, highlighting the presence of both unique and previously isolated genomes in the samples (Tables 5.6 – 5.10). *Nitrosopumilus* are ammonia-oxidising archaea, which use CO<sub>2</sub> as a carbon source to convert ammonia (NH<sub>3</sub>) to nitrite (NO<sub>2</sub><sup>-</sup>) during nitrification (Banning *et al.*, 2015). These organisms may therefore be significant in sediment nitrogen cycling, helping to provide nitrite (and subsequently nitrate) for biological uptake by heterotrophic organisms (Francis *et al.*, 2005). This may be substantial for facilitating the activity of heterotrophic microbes, given the potential for nitrogen limitation in the sediments (Figure 5.4).

Interestingly, sequences relating to the Proteobacterium *Magnetospirillum* were found in samples 1, 4 and 7 (Tables 5.6 – 5.10). This organism is microaerophilic, magneto tactic and capable of producing high quality magnetite from low iron aquatic environments (Matsunaga *et al.*, 2005). However, sequences ranged between 74 – 78% identity to *Magnetosprillum*, highlighting the potential for a novel species in the metagenomes (Tables 5.6 – 5.10). As *Magnetosprillum* has a wide range of commercial applications, such as bioremediation, manufacture and pharmaceuticals, the isolation of a new novel species may provide additional commercial applications to those already identified (Safarik and Safarikova, 2004). This again highlights the need for further culture-based analysis and single cell sequencing of environmental genomes from unique environments, such as those sampled here.

The binned genomes also displayed evidence of sulfur cycling, in line with the results of the read based taxonomic annotation (Figure 5.5). The presence of the chemolithotrophic *Thioalkalivibrio*, in metagenomes 1, 2 and 7 indicates the potential of sulfur oxidation in the sediments, whereby elemental sulfur (S) is converted to sulfate (SO<sub>4</sub><sup>2-</sup>) (Sorokin *et al.*, 2001). In turn, the presence of the anaerobic *Desulfococcus*, *Desulfuromonas*, *Desulfobacter* and *Desulfobacterium* highlight the presence of sulfur and sulfate reduction (Taylor and Parkes, 1983; Kleindienst *et al.*, 2014). These organisms have been shown to oxidise organic compounds such as acetate and pyruvate using sulfate and sulfur, to reduced forms of sulfur (Brysch *et al.*, 1987).



Consequently, the genome binning has identified the potential for novel species with ecological importance in the sediment metagenomes. Deep re-sequencing of the metagenomes may help to recover longer more complete metagenome assembled genomes. This may be beneficial to gain a better understating of the strain level diversity and metabolisms of unique species. This, alongside genome sequencing of cultured stains would help provide new draft genomes to publicly available databases.

**Table 5.6:** Metagenome Sample 1 assembled bins and top BLAST match for each bin, with % NCBI GenBank identity and bit score. In total, 36 assembled bins representing distinct genomes within Sample 1 metagenome were identified. The genomes with an % identity less than 80% to GenBank sequences are highlighted.

Bin number	Match GenBank accession number	% identity	Bit score	Match definition
1	CP010868.1	81.207	4525	<i>Candidatus Nitrosopumilus piranensis</i> strain D3C
2	CP007451.1	92.906	2769	<i>Draconibacterium orientale</i> strain FH5T
3	CP001339.1	75.937	3073	<i>Thioalkalivibrio sulfidiphilus</i>
4	CP003842.1	79.143	13367	<i>Candidatus Nitrosopumilus koreensis</i> AR1
5	CP016268.1	79.051	3114	<i>Woeseia oceani</i> strain XK5
6	AP007255.1	75.303	1166	<i>Magnetospirillum magneticum</i> AMB-1
7	AP012978.1	75.372	859	Endosymbiont of unidentified scaly snail isolate Monju DNA
8	CP002453.1	76.899	1035	<i>Cellulophaga algicola</i> DSM 14237
9	CP003560.1	82.543	857	<i>Flammeovirga</i> sp. MY04 chromosome 1
10	AP018052.1	80.625	1574	<i>Shewanella halifaxensis</i> HAW-EB4
11	AP018052.1	77.399	1448	<i>Thiohalobacter thiocyanaticus</i> DNA
12	CP002568.1	77.305	1777	<i>Polymorphum gilvum</i> SL003B-26A1
13	CP010869.1	81.564	2023	<i>Confluentimicrobium</i> sp. EMB200-NS6
14	CP003843.1	98.205	3698	<i>Candidatus Nitrosopumilus</i> sp. AR2
15	LN614827.1	72.869	460	<i>Legionella fallonii</i> LLAP-10
16	CP000393.1	77.075	1085	<i>Trichodesmium erythraeum</i> IMS101
17	CP014646.1	74.338	1081	<i>Thauera humireducens</i> strain SgZ-1
18	CP009416.1	81.239	436	<i>Jeotgalibacillus malaysiensis</i> strain D5
19	CP013355.1	78.817	2726	<i>Lutibacter profundus</i> strain LP1 chromosome
20	CP023451.1	83.537	2422	<i>Rhizorhabdus dicambivorans</i> strain Ndbn-20
21	CP015848.1	78.456	2015	<i>Magnetospirillum</i> sp. ME-1
22	CP000866.1	83.057	3886	<i>Nitrosopumilus maritimus</i> SCM1
23	CP013099.1	74.967	652	<i>Candidatus Tenderia electrophaga</i> isolate NRL1
24	CP002026.1	75.507	771	<i>Starkeya novella</i> DSM 506
25	AP014879.1	75.341	1037	<i>Sulfuricoccus limicola</i> DNA
26	AP013066.1	81.522	1352	<i>Sulfuricella denitrificans</i> skB26 DNA
27	CP000390.1	77.996	926	<i>Chelativorans</i> sp. BNC1
28	CP007451.1	85.331	4113	<i>Draconibacterium orientale</i> strain FH5T
29	CP011392.1	75.39	789	<i>Dehalogenimonas</i> sp. WBC-2
30	CP013411.1	75.745	1227	<i>Burkholderia thailandensis</i> strain 2002721643
31	CP003230.1	81.084	1557	<i>Cycloclasticus</i> sp. P1
32	CP002056.1	76.696	1140	<i>Methylothermus versatilis</i> 301
33	CP016268.1	78.819	1533	<i>Woeseia oceani</i> strain XK5
34	CP011110.1	78.28	881	<i>Pseudomonas chlororaphis</i> strain PCL1606
35	CP013355.1	77.904	3009	<i>Lutibacter profundus</i> strain LP1 chromosome
36	AP012305.1	76.957	1461	<i>Azoarcus</i> sp. KH32C plasmid pAZKH DNA

**Table 5.7:** Metagenome Sample 2 assembled bins and top BLAST match for each bin, with % identity and bit score. In total, 36 assembled bins representing distinct genomes within Sample 2 metagenome were identified. The genomes with an % identity less than 80% to GenBank sequences are highlighted.

Bin number	Match	GenBank accession number	% identity	Bit score	Match definition
1	CP021431.1		100	150	<i>Loktanella vestfoldensis</i> strain SMR4r
2	CP003843.1		97.522	3856	<i>Candidatus Nitrosopumilus</i> sp. AR2
3	CP011412.1		81.859	5535	<i>Sedimenticola thiotaurini</i> strain SIP-G1
4	CR522870.1		78.462	457	<i>Desulfotalea psychrophila</i> LSV54
5	CP001147.1		82.116	1735	<i>Thermodesulfobacterium yellowstonii</i> DSM 11347
6	CP000478.1		77.095	1048	<i>Syntrophobacter fumaroxidans</i> MPOB
7	CP001339.1		75.921	3068	<i>Thioalkalivibrio sulfidiphilus</i> HL-EbGr7
8	CP018632.1		72.188	893	<i>Granulosicoccus antarcticus</i> IMCC3135
9	CP000252.1		80.675	737	<i>Thioalkalivibrio sulfidiphilus</i> HL-EbGr7
10	CP011070.1		78.046	8626	<i>Thioalkalivibrio sulfidiphilus</i> HL-EbGr7
11	CP016268.1		80.768	4023	<i>Woeseia oceani</i> strain XK5
12	CP017478.1		71.649	292	<i>Lutibacter</i> sp. LPB0138
13	CP003843.1		81.137	3921	<i>Candidatus Nitrosopumilus</i> sp. AR2
14	CP011070.1		85.923	2769	<i>Candidatus Nitrosopumilus adriaticus</i> strain NF5
15	CP011036.1		86.065	2265	<i>Pseudoalteromonas nigrifaciens</i> strain KMM 661
16	CP013355.1		77.878	2198	<i>Lutibacter profundus</i> strain LP1 chromosome
17	CP003843.1		85.51	2846	<i>Candidatus Nitrosopumilus</i> sp. AR2
18	CP011412.1		84.806	1216	<i>Sedimenticola thiotaurini</i> strain SIP-G1
19	CP001131.1		75.762	1206	<i>Anaeromyxobacter</i> sp. K
20	CP001032.1		84.372	1886	<i>Opitutus terrae</i> PB90-1
21	CP011070.1		75.438	983	<i>Candidatus Nitrosopumilus adriaticus</i> strain NF5
22	CP010869.1		81.564	2023	<i>Confluentimicrobium</i> sp. EMB200-NS6
23	CP000142.2		79.661	1079	<i>Pelobacter carbinolicus</i> DSM 2380
24	CP011070.1		84.783	2320	<i>Candidatus Nitrosopumilus adriaticus</i> strain NF5
25	CP020892.1		75.18	761	<i>Pseudomonas</i> sp. M30-35 chromosome
26	CP011412.1		74.583	1262	<i>Sedimenticola thiotaurini</i> strain SIP-G1 chromosome
27	CP000083.1		74.989	1014	<i>Colwellia psychrethraea</i> 34H
28	CP018889.1		77.537	712	<i>Beggiatoa leptomitiformis</i> strain D-401 chromosome
29	AP018042.1		76.407	1153	<i>Marinifilaceae bacterium</i> SPP2 DNA
30	CP012398.1		76.8	987	<i>Chelatococcus</i> sp. CO-6
31	CP016268.1		77.68	2095	<i>Woeseia oceani</i> strain XK5
32	CP020555.1		79.577	1653	<i>Streptomyces</i> sp. Sge12
33	CP013355.1		78.791	2067	<i>Lutibacter profundus</i> strain LP1
34	CP000859.1		78.682	1262	<i>Desulfococcus oleovorans</i> Hxd3
35	LN890655.2		90.747	1195	<i>Ardenticatena</i> sp. Cfx-K strain Cfx-K
36	CP023439.1		79.824	1461	<i>Thauera</i> sp. K11 chromosome

**Table 5.8:** Metagenome Sample 4 assembled bins and top BLAST match for each bin, with % identity and bit score. In total, 31 assembled bins representing distinct genomes within Sample 4 metagenome were identified. The genomes with an % identity less than 80% to GenBank sequences are highlighted.

Bin number	Match GenBank accession number	% identity	Bit score	Match definition
1	CP011412.1	81.844	5529	<i>Sedimenticola thiotaurini</i> strain SIP-G1
2	CP003843.1	79.766	4977	<i>Candidatus Nitrosopumilus sediminis</i>
3	LT934425.1	82.249	2551	<i>Candidatus Kuenenia stuttgartiensis</i> isolate
4	CP003787.1	71.384	355	<i>Riemerella anatipestifer</i> RA-CH-1
5	CP019936.1	76.165	2981	Chromatiaceae bacterium 2141T.STBD.0c.01a chromosome
6	CP011070.1	79.409	10133	<i>Candidatus Nitrosopumilus adriaticus</i> strain NF5
7	CP012358.1	82.108	1635	<i>Oblitimonas alkaliphila</i> strain B4199
8	FO203512.1	83.839	2992	<i>Oleispira antarctica</i> strain RB-8
9	CP000282.1	75.114	1862	<i>Saccharophagus degradans</i> 2-40
10	CP000866.1	81.693	1371	<i>Nitrosopumilus maritimus</i> SCM1
11	CP016268.1	80.094	3238	<i>Woeseia oceani</i> strain XK5
12	CP003843.1	87.927	965	<i>Candidatus Nitrosopumilus</i> sp. AR2
13	CP018889.1	74.03	220	<i>Beggiatoa leptomitiformis</i> strain D-401
14	CP003843.1	86.121	941	<i>Candidatus Nitrosopumilus</i> sp. AR2
15	CP003843.1	84.052	1910	<i>Candidatus Nitrosopumilus</i> sp. AR2
16	CP003842.1	77.178	6458	<i>Candidatus Nitrosopumilus koreensis</i> AR1
17	CP015136.1	80.136	1354	<i>Luteitalea pratensis</i> strain DSM 100886
18	CP016268.1	77.352	965	<i>Woeseia oceani</i> strain XK5
19	CP003843.1	81.619	1943	<i>Candidatus Nitrosopumilus</i> sp. AR2
20	AP007255.1	75.343	1170	<i>Magnetospirillum magneticum</i> AMB-1
21	CP023741.1	75.465	1107	<i>Sphingobium yanoikuyae</i> strain S72
22	CP001720.1	74.05	939	<i>Desulfotomaculum acetoxidans</i> DSM 771
23	CP005934.1	90.769	1125	<i>Candidatus Methanomassiliicoccus intestinalis</i> Issoire-Mx1
24	AP018052.1	89.758	2455	<i>Thiohalobacter thiocyanaticus</i> DNA
25	CP019343.1	73.404	857	<i>Oceanicoccus sagamiensis</i> strain NBRC
26	AP018052.1	75.158	1524	<i>Thiohalobacter thiocyanaticus</i>
27	CP012154.1	77.689	1871	<i>Wenzhouxiangella marina</i> strain KCTC 42284
28	CP012040.1	77.632	592	<i>Cyclobacterium amurskyense</i> strain KCTC 12363
29	CP001707.1	75.194	1173	<i>Kangiella koreensis</i> DSM 16069
30	CP002271.1	74.648	638	<i>Cyclobacterium amurskyense</i> strain KCTC 12363
31	CP000316.1	82.536	2983	<i>Polaromonas</i> sp. JS666

**Table 5.9:** Metagenome Sample 5 assembled bins and top BLAST match for each bin, with % identity and bit score. In total, 51 assembled bins representing distinct genomes within Sample 5 metagenome were identified. The genomes with an % identity less than 80% to GenBank sequences are highlighted.

Bin number	Match GenBank accession number	% identity	Bit score	Match definition
1	CP000724.1	83.259	2180	<i>Alkaliphilus metalliredigens</i> QYMF
2	CP016268.1	79.953	3103	<i>Woeseia oceani</i> strain XK5
3	CP012851.1	76.274	1037	<i>Persicobacter</i> sp. JZB09
4	LT981265.1	77.833	1290	<i>Candidatus Nitrosocaldus cavascurensis</i> strain SCU2
5	CP013118.1	73.115	878	<i>Salinivirga cyanobacteriivorans</i> strain L21-Spi-D4
6	CP002031.1	76.244	1387	<i>Geobacter sulfurreducens</i> KN400
7	CP001032.1	76.652	1676	<i>Opitutus terrae</i> PB90-1
8	CP009788.1	73.92	682	<i>Geobacter pickeringii</i> strain G13
9	CP010802.1	75.07	2220	<i>Desulfuromonas soudanensis</i> strain WTL chromosome
10	CP002271.1	78.881	737	<i>Stigmatella aurantiaca</i> DW4/3-1
11	CP003843.1	91.176	3602	<i>Candidatus Nitrosopumilus</i> sp. AR2
12	CP003360.1	78.68	4414	<i>Desulfomonile tiedjei</i> DSM 6799
13	CP003350.1	74.774	1461	<i>Frateuria aurantia</i> DSM 6220
14	CP013355.1	75.697	590	<i>Lutibacter profundus</i> strain LP1
15	CP009505.1	83.006	1927	<i>Methanosarcina</i> sp. MTP4
16	FO203503.1	73.2	1469	<i>Desulfobacula toluolica</i> Tol2
17	LT934425.1	75.995	815	<i>Candidatus Kuenenia stuttgartiensis</i>
18	CP012358.1	82.159	1640	<i>Oblitimonas alkaliphila</i> strain B4199 chromosome
19	CP013355.1	74.638	1397	<i>Lutibacter profundus</i> strain LP1 chromosome
20	CP003985.1	73.498	669	<i>Desulfocapsa sulfexigens</i> DSM 10523
21	CP003843.1	86.747	15400	<i>Candidatus Nitrosopumilus</i> sp. AR2
22	CP006900.2	73.958	484	<i>Pandoraea pnomenusa</i> 3kgm
23	CP003273.1	76.135	1574	<i>Desulfotomaculum gibsoniae</i> DSM 7213
24	CP019913.2	75	928	<i>Desulfococcus multivorans</i> strain DSM 2059
25	CP016268.1	79.47	1724	<i>Woeseia oceani</i> strain XK5
26	CP000859.1	83.796	1672	<i>Desulfococcus oleovorans</i> Hxd3
27	CP000934.1	74.105	1664	<i>Cellvibrio japonicus</i> Ueda107
28	CP025791.1	75.197	994	<i>Flavivirga eckloniae</i> strain ECD14 chromosome
29	CU207366.1	73.988	130	<i>Gramella forsetii</i> KT0803
30	CP013457.1	80.632	1099	<i>Burkholderia</i> sp. MSMB617WGS
31	CP013118.1	74.824	734	<i>Salinivirga cyanobacteriivorans</i> strain L21-Spi-D4
32	CP003843.1	86.724	2320	<i>Candidatus Nitrosopumilus</i> sp. AR2
33	CP003843.1	94.042	9022	<i>Candidatus Nitrosopumilus</i> sp. AR2
34	CP001087.1	76.225	1709	<i>Desulfobacterium autotrophicum</i> HRM2
35	CP003985.1	74.536	881	<i>Desulfocapsa sulfexigens</i> DSM 10523
36	CP003843.1	88.372	1705	<i>Desulfobacterium autotrophicum</i> HRM2
37	CP003843.1	88.209	4833	<i>Desulfobacterium autotrophicum</i> HRM2
38	CP016268.1	79.01	2320	<i>Woeseia oceani</i> strain XK5
39	CP013355.1	81.96	1448	<i>Lutibacter profundus</i> strain LP1 chromosome
40	CP000473.1	77.081	1903	<i>Candidatus Solibacter usitatus</i> Ellin6076
41	CP003380.1	89.854	1404	<i>Methylophaga frappieri</i> strain JAM7
42	CP011454.1	78.822	931	<i>Gemmatimonas phototrophica</i> strain AP64
43	CP010904.1	75.319	669	<i>Kiritimatiella glycovorans</i> strain L21-Fru-AB
44	CP013355.1	78.896	2442	<i>Lutibacter profundus</i> strain LP1 chromosome
45	CP011125.1	73.835	1144	<i>Sandaracinus amylolyticus</i> strain DSM 53668
46	CP013355.1	77.89	5387	<i>Lutibacter profundus</i> strain LP1 chromosome
47	CP003389.1	77.644	992	<i>Corallococcus coralloides</i> DSM 2259
48	CP003843.1	87.817	3410	<i>Candidatus Nitrosopumilus</i> sp. AR2
49	CP013355.1	84.56	1770	<i>Lutibacter profundus</i> strain LP1 chromosome
50	CP015080.1	78.171	1742	<i>Desulfuromonas</i> sp. DDH964
51	CP006587.1	78.762	1112	<i>Hymenobacter</i> sp. APR13

**Table 5.10:** Metagenome Sample 7 assembled bins and top BLAST match for each bin, with % identity and bit score. In total, 19 assembled bins representing distinct genomes within Sample 7 metagenome were identified. The genomes with an % identity less than 80% to GenBank sequences are highlighted.

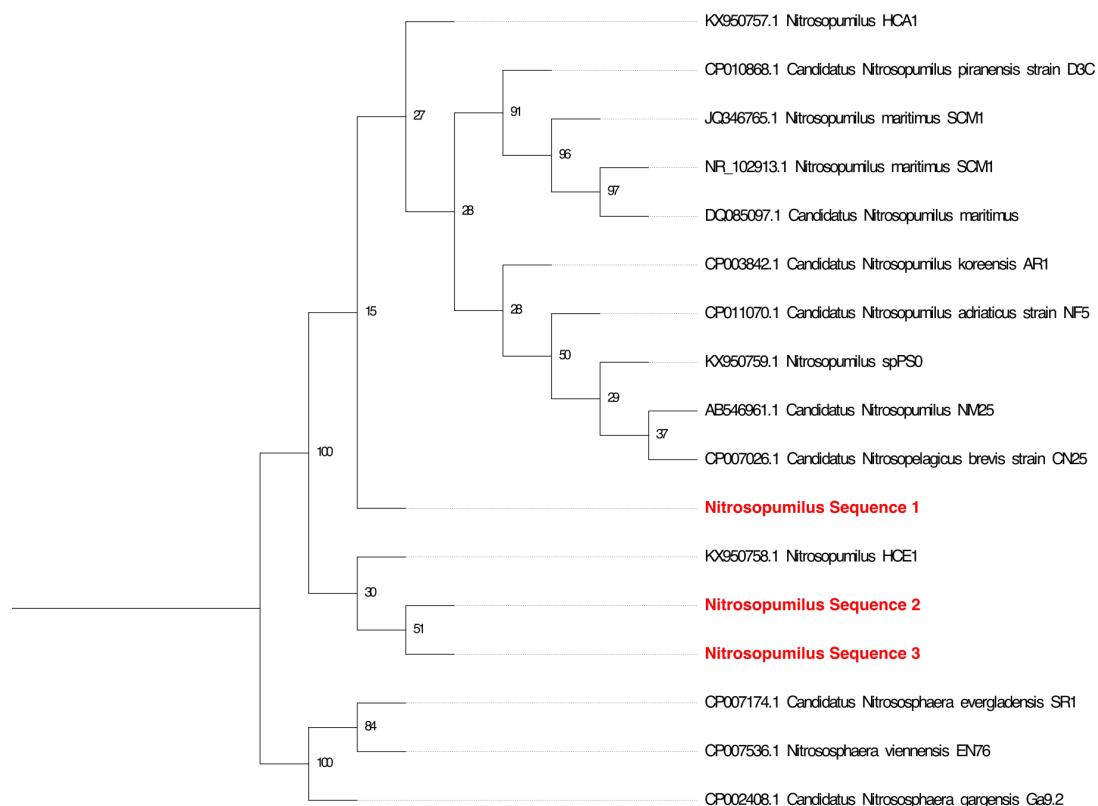
Bin number	Match GenBank accession number	% identity	Bit score	Match definition
1	CP021324.1	81.804	3290	<i>Candidatus Nitrosomarinus catalina</i> strain SPOT01
2	CP000866.1	79.394	4019	<i>Nitrosopumilus maritimus</i> SCM1
3	CP001339.1	78.306	4396	<i>Thioalkalivibrio sulfidophilus</i> HL-EbGr7
4	CP003843.1	82.51	3799	<i>Candidatus Nitrosopumilus</i> sp. AR2
5	AP014936.1	76.285	1186	<i>Sulfurifustis variabilis</i> DNA
6	CP003843.1	78.281	4008	<i>Candidatus Nitrosopumilus</i> sp. AR2
7	CP001339.1	81.686	1576	<i>Thioalkalivibrio sulfidophilus</i> HL-EbGr7
8	CP019630.1	78.577	1794	<i>Labrenzia aggregata</i> strain RMAR6-6
9	CP017689.1	79.23	680	<i>Thalassotalea crassostreae</i> strain LPB0090
10	CP003360.1	90.76	1700	<i>Desulfomonile tiedjei</i> DSM 6799
11	CP010868.1	78.458	1273	<i>Candidatus Nitrosopumilus piranensis</i> strain D3C
12	CP014944.1	86.356	2593	<i>Colwellia</i> sp. PAMC 20917 chromosome
13	CP016268.1	79.204	2979	<i>Woeseia oceani</i> strain XK5
14	CP012358.1	81.239	1400	<i>Oblitimonas alkaliphila</i> strain B4199
15	CP017689.1	78.912	1186	<i>Thalassotalea crassostreae</i> strain LPB0090
16	CP010868.1	78.585	2423	<i>Candidatus Nitrosopumilus piranensis</i> strain D3C
17	CP000083.1	83.764	1925	<i>Colwellia psychrerythraea</i> 34H
18	CP021106.3	79.99	1526	<i>Nitrosospira lacus</i> strain APG3
19	LN997848.1	74.536	1681	<i>Magnetospirillum</i> sp. XM-1

### 5.3.5 Phylogenetic analysis – 16s rRNA

To explore the findings from genome binning further and to highlight the potential of metagenome data analysis, the genus *Nitrosopumilus* was selected for 16s rRNA analysis, as this archaeon was identified in multiple genome bins and may be ecologically important in the sediment nitrogen cycle, if active. The 16s sequences were obtained from the metagenomes, however as amplification was not used, some less abundant strains may not have been recovered through sequencing. However, it is likely that the *Nitrosopumilus* sequences obtained were those which were the most dominant in the samples (Rodriguez and Konstantinos, 2014). The metagenome 16s sequences relating to *Nitrosopumilus* were aligned with and subsequently evaluated in a maximum likelihood phylogeny with GenBank cultured relatives (Figure 5.6). The results show that the sample sequences align most closely with *Nitrosopumilus HCE1* (*Nitrosopumilus oxclinae*) and *Nitrosopumilus HCA1* (*Nitrosopumilus cobalaminigenes*) (Figure 5.6). When the 16s sample sequences were BLAST searched, *Nitrosopumilus* sequence 1, 2 and 3 aligned at 99% sequence similarity to *Nitrosopumilus oxclinae*, *Nitrosopumilus cobalaminigenes* and *Candidatus Nitrosopumilus sediminis* (Park *et al.*, 2012; Qin *et al.*, 2017; Figure 5.6). This supports the genome bins and phylogeny in identifying these sequences as belonging to *Nitrosopumilus* genomes. However, whilst 99% similarity was shown using the 16s sequences, work by Park *et al.*, (2012) has shown that when looking at overall genome nucleotide identity, the nucleotide similarity can reduce below 80%, providing evidence for new species or strains, in line with the results from

genome binning. Consequently, this highlights the need for culturing and full genome sequencing of these organisms, as these may constitute new genomes to the *Nitrosopumilus* genus, as found by Park *et al.*, (2012).

As highlighted above, the presence of *Nitrosopumilus* genomes provided evidence for ammonia oxidising archaea (AOA) in the sediments, which if active, would contribute to nitrogen cycling (Banning *et al.*, 2015). These AOA are well suited to low nitrogen environments such as the fjord sediments sampled, as they have a high affinity to ammonia (Qin *et al.*, 2017; Figure 5.4). Additionally, they are suited to the low organic carbon contents of the sediments as they efficiently use CO<sub>2</sub> as a carbon source (Qin *et al.*, 2017). This could explain why these genomes were consistently identified using genome binning, as these genomes may be tolerant to low nutrient conditions (Tables 5.6 – 5.10). Furthermore, the 99% similarity with *Nitrosopumilus cobalaminigenes* indicates the genomes may have additional adaptive mechanisms to the extreme conditions in the fjord sediments. *Nitrosopumilus cobalaminigenes* is psychrotolerant and can survive in a wide salinity range (10 – 40 PSU) (Qin *et al.*, 2017). This suggests these genomes may be able to survive preferentially in the cold fjord conditions and withstand fluctuating salinities during mixing of stratified waterbodies (Qin *et al.*, 2017).



**Figure 5.6:** 16s rRNA Maximum likelihood phylogeny of *Nitrosopumilus* sample sequences with cultured relatives, obtained from NCBI GenBank. The accession numbers for cultured relatives are shown and metagenome sample sequences highlighted in bold. Bootstrap

support values are provided for each branch node, resulting from 1000 tree iterations. The tree is rooted on *Nitrososphaera*, in line with the phylogeny of Quin *et al.*, (2017).

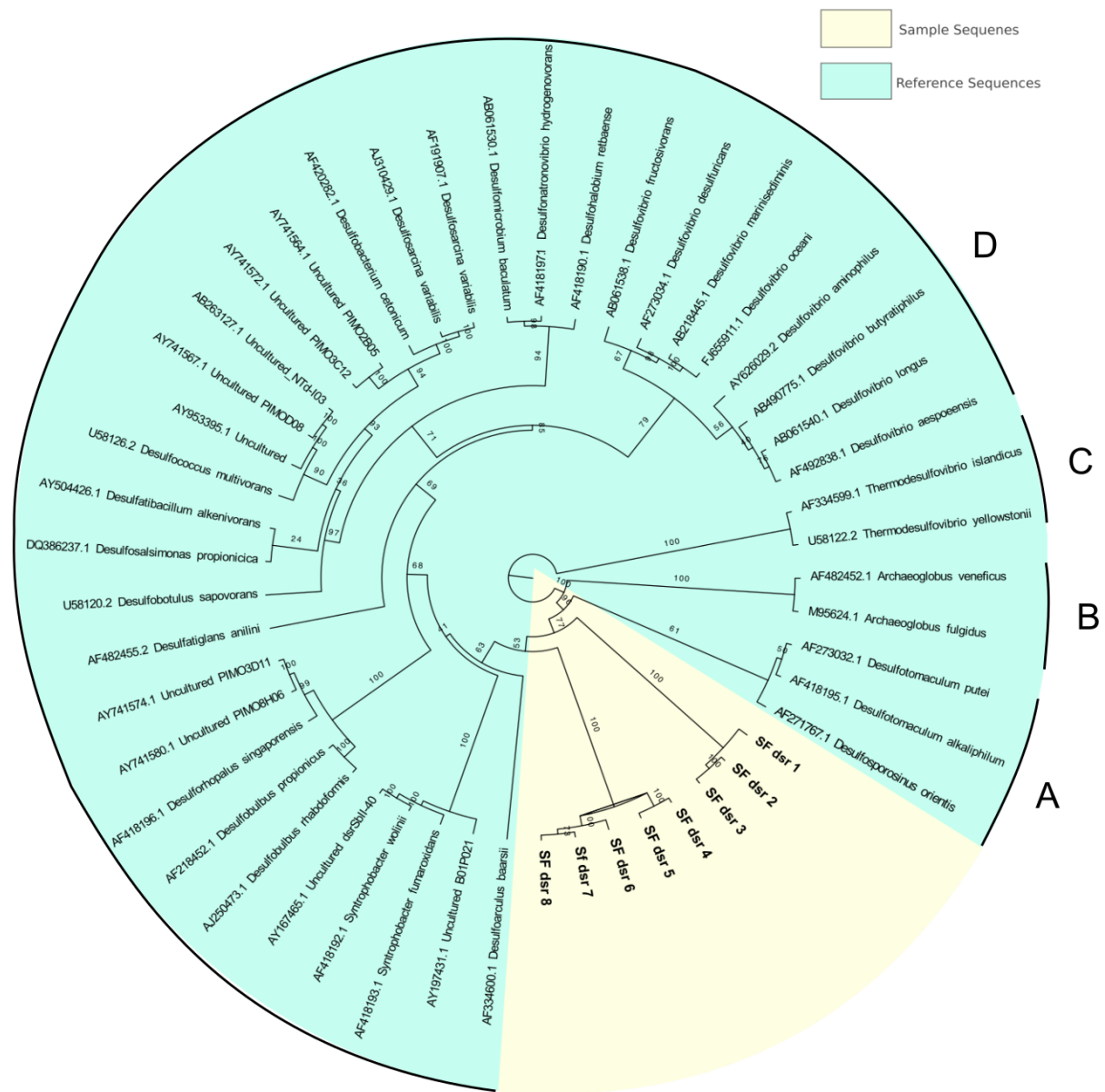
### 5.3.6 Phylogenetic analysis – dissimilatory sulfite reductase

A key advantage of metagenomics, in comparison to 16s rRNA sequencing, is that functional genes can be investigated (Wooley *et al.*, 2010). This means that functional pathways can be searched for, and the taxonomy subsequently assigned (Handelsman, 2004). The benefit of this is that it provides more accurate evidence for biogeochemical cycling in the sediments, as the genes encoding the functional pathways are present (Handelsman, 2004). Whilst this does not show activity of the pathways, it highlights the potential for these to occur (Wooley *et al.*, 2010).

The findings of both the read based taxonomy and genome bins (Figure 5.5, Tables 5.6 – 5.10) identified the presence of sulfur cycling microbes in the sediments. To investigate this further, and demonstrate the potential of metagenome sequencing, the *dsrAB* gene (dissimilatory sulfite reductase) was investigated. This gene encodes the reduction of sulfite ( $\text{SO}_3^{2-}$ ) to sulfide ( $\text{H}_2\text{S}$ ) (Müller *et al.*, 2015). This is a key reduction step in the sulfur cycle, whereby prokaryotes reduce sulfite for energy in combination with the oxidation of organic matter, during anaerobic respiration (Müller *et al.*, 2015). A maximum likelihood phylogeny of metagenome sample *dsrAB* genes and those from GenBank relatives is shown in Figure 5.7. Here, cultured relatives have been grouped in to the classes of Clostridia, Archaeoglobi, Nitrospira and Deltaproteobacteria (Figure 5.7), based on the phylogeny of Moreau *et al.*, (2010). The sample sequences (Sf *dsr* 1 – 7) are grouped together, branching between Clostridia and Deltaproteobacteria (Figure 5.7). This may highlight a unique set of sulfite reducers in the samples. To investigate this further, the sample *dsrAB* sequences were BLASTn searched against cultured isolates in NCBI GenBank, with results shown in Table 11. The sequences were matched with several sulfite reducing microbes, with 75% of samples identified under the class of Gammaproteobacteria (Table 8). Samples SF *dsr* 1 – 2 were classified as Deltaproteobacteria and Chlorobia, respectively. However, as the matches received a low score, and identity of 82%, there is limited confidence in these assignments. Overall, the identity of matches ranged between 76 – 93%, suggesting that these sequences may be unique (Table 11). Additional single genome sequencing would help isolate the full genomes of these prokaryotes, thereby validating the suggestions made here. Furthermore, future work with meta-transcriptomics would help highlight if these sulfite reducers were active, as this technique investigates the transcribed mRNA (Carvalhais *et al.*, 2013).

The identification of *dsrAB* sequences in the samples supports the potential of sulfur cycling highlighted by the read based taxonomy and genome bins (Figure 5.5; Table 5.5; Tables 5.6

– 5.10). This contributes to the evidence indicating the potential for biogeochemical cycling within these oligotrophic sediments. The sulfite may be lost by volatilisation or mineral formation, however may also be oxidised back to sulfite and sulfate by chemo-lithotrophic prokaryotes such as *Chromatiales*, identified in the read based taxonomic annotation (Holmer and Stockholm, 2001; Imhoff, 2005).



**Figure 5.7:** Maximum likelihood phylogeny of the Dissimilatory Sulfite Reductase (*dsrAB*) gene, including metagenome sample *dsr* sequences (bold and yellow, SF *dsr* 1 - 7) and reference genes (blue) obtained from NCBI GenBank and the phylogeny of Moreau et al., (2010). Reference sequences display the GenBank accession number, followed by the species name. Reference sequences are grouped in to classes A) Clostridia; B) Archaeoglobi; C) Nitrospira and D) Deltaproteobacteria.



**Table 5.11:** BLASTn matches for metagenome *dsrAB* sequences 1-7 against NCBI GenBank. The BLAST match name, GenBank accession number, % identity and score is provided.

Sample Name	Blast Match	Accession Number	% identity	Total Score
sf dsr 1	<i>Sulfurifustis variabilis</i>	AP014936.1	93	686
sf dsr 2	<i>Desulfobulbus</i> sp. ORNL	CP021255.1	82	95.3
sf dsr 3	<i>Chlorobaculum limnaeum</i> strain DSM 1677	CP017305.1	82	149
sf dsr 4	<i>Thioalkalivibrio paradoxus</i> ARh 1	CP007029.1	78	431
sf dsr 5	<i>Sulfuricaulis limicola</i> DNA, complete genome	AP014879.1	76	490
sf dsr 6	<i>Thiohalobacter thiocyanaticus</i>	AP018052.1	91	854
sf dsr 7	<i>Thiohalobacter thiocyanaticus</i>	AP018052.1	82	848
sf dsr 8	<i>Thioflaviccoccus mobilis</i> 8321	CP003051.1	92	81

### 5.3.7 Implications for biogeochemical cycling

This analysis has used metagenomics to highlight the microbial community composition and functional potential of benthic fjord sediments. Whilst productivity assays would be needed to measure rates of activity, the presence of functional genes can be used to indicate potential biogeochemical implications of fjord microbial communities.

Interestingly, the fjord waters of Chilean Patagonia had been proposed as a hotspot of primary productivity, hosting three UNESCO bio reserves and support commercially important fisheries (Iriate *et al.*, 2007; Haussemann and Forsterra 2009; Niklitscheck *et al.*, 2013). However, the results from our analysis indicated a low nutrient environment, one which would not typically be classed as productive. In line with previous literature, we confirmed the presence of a surface freshwater lens, supplied by glacial runoff (Arancena *et al.*, 2011). This surface lens was low nutrient and high turbidity, attributed to the outflow of inorganic sediments from glacial runoff (Arancena *et al.*, 2011). We propose that the turbid surface waters may block light, limiting surface productivity and thus the export of organic matter to deep sediments. This is supported by previous research which has shown productivity to increase moving away from inner fjord waters, attributed to light and nutrient availability (Silva 2008; Arancena *et al.*, 2011; Gonzalez *et al.*, 2013). However, nutrient export from terrestrial surface waters may provide some limiting nutrients, helping to stimulate biological productivity to inner fjord waters (Gonzalez *et al.*, 2013).

With rising global temperatures, it is likely that glacial meltwater export will increase in this region. Increased meltwater export may reduce benthic and surface water biological productivity, through limiting light penetration and nutrient availability. Decreased surface productivity will not only limit nutrient export to benthic sediments, but may also have

implications on wider ecosystem functioning and fisheries (Landaeta *et al.*, 2012; Gonzalez *et al.*, 2013).

Despite limitations on light and nutrient availability, we were able to detect genomic potential for microbial carbon, sulfur and nitrogen cycling in the sediments investigated. Potentially attributed to the extreme conditions, we found a combination of anaerobes, extremophiles and chemolithotrophs, who are resilient to the low light, oligotrophic and cold conditions. These results lay a foundation for the identification of potential novel genomes within the samples, which could relate to environmental selection for a unique subset of traits for survival. However, more exploration on the genomic level is needed to characterize these genomes and their biogeochemical functionality.

#### **5.4 Conclusions**

This study aimed to investigate the microbial community composition of the Steffen Fjord, Chilean Patagonia. This fjord sparked interest as it is fed by melt water from the Steffen glacier, and drains into the marine Baker fjord system. The lack of investigation in this region means that new microbial genomes, that have yet to be sequenced, may be in the sediments, with ecological, biotechnological or industrial uses.

The metadata showed the Steffen glacier to have a clear effect on fjord salinity, turbidity and temperature, with glacial freshwater runoff separated at the surface from the more dense marine bottom waters. The benthic sediments displayed low TN and TOC, indicating that nutrient limitation may be taking place. This indicates that increased freshwater runoff in future years may increase the fjord stratification, which could influence biological productivity. Overall, the sediment microbial community composition was largely dominated by bacteria, with archaea accounting for 3-5% of the metagenomes. The bacteria and archaea recovered related to anaerobic (and often methanogenic) taxa, indicating the presence of anoxia, or low oxygen zones within the surface sediments. Additionally, bacteria with adaptations or extremophilic qualities were detected, attributed to the oligotrophic, low oxygen, cold and dark conditions present in the sediments. Despite this, microbes relating to nitrogen and sulfur cycling organisms were found in the sediments, indicating the potential for microbial nutrient cycling in this low nutrient environment. Genome binning was used to help identify the presence of potentially novel genomes within the samples. Over half the genome bins scored below 80% identity to NCBI GenBank. This indicates that some genomes found during binning could represent new strains or species. Additionally, microbes relating to sulfur cycling organisms were found frequently among the metagenomes, and bins relating to the ammonia oxidizing archaea, *Nitrosopumilus*, were common.

Sequences relating to *Nitrosopumilus* were selected for a 16s rRNA phylogeny with sequenced relatives from the phylogeny of Park *et al.*, (2012). The sample sequences were shown to have 99% similarity to the 16s sequences of sequenced *Nitrosopumilus* relatives *Nitrosopumilus oxclinae*, *Nitrosopumilus cobalaminigenes* and *Candidatus Nitrosopumilus sediminis* (Park *et al.*, 2012; Qin *et al.*, 2017). The presence of *Nitrosopumilus*-like genomes may be related to the high affinity to ammonia, and thereby resilience to the low TN of the sediments. Additionally, the presence of these genomes highlights the potential for nitrogen cycling in the sediments, despite the oligotrophy detected. Furthermore, the detection of sulfur cycling microbes within the sediments was used to drive a *dsrAB* phylogeny, for sulfite reducing organisms in the metagenomes. This provides an addition to 16s analysis, as it shows the presence of functional genes involved in sulfur cycling. The sulfite reducers were shown to constitute a discrete group within the phylogeny and had below a 91% identity to cultured relatives.

Overall the results highlight the presence of extremophiles and nutrient cycling microbes, despite the hostile conditions. The results demonstrate the presence of potentially unique microbes within the samples, which may have ecological or commercial implications. It is hoped that this work will stimulate further culture based analysis and single cell sequencing to fully isolate novel genomes, to facilitate understanding of environmental microbial diversity. Additionally, this analysis has shown the scope of metagenomics for investigating microbial communities, as it enables taxonomic and functional potential to be inferred.

## **5.5 Limitations and Further work**

Whilst this analysis has provided an insight into the taxonomic diversity and functional potential of Patagonian fjord benthic microbial communities, it has also highlighted areas for future work. Firstly, a large fraction of the sequenced DNA was unannotated, attributed to either novelty or the presence of degraded DNA. To reveal the composition of more degraded DNA, amplification using short primers could be used, for example with metabarcoding (Taberlet *et al.*, 2012; Leray and Knowlton, 2015). This would increase the abundance of shorter DNA fragments for sequencing and annotation. Additionally, 16s rRNA amplicon sequencing and deep metagenome sequencing could also be used to uncover less abundant microbes, which may not have been recovered from this analysis. As sediment metagenomes are diverse, the sequencing carried out in this analysis is likely to focus on the most abundant fraction and may exclude some less abundant organisms (Rodriguez and Konstantinidis, 2014).

Deep metagenome sequencing may also be beneficial to recover complete genome sequences, which could be used to compile metagenome assembled draft genomes (MAGs). This would be largely related to an increase in sequencing coverage which would help resolve repeat regions during the assembly of genomes (Albertsen *et al.*, 2013). The recovery of MAGs would provide a deeper understanding of the metabolic functioning and strain level diversity of novel genomes. Culture based analysis and subsequent complete genome sequencing would also be beneficial for the *Nitrosopumilus* genomes. This would allow comparison of the average nucleotide identity (ANI) to current sequenced strains, to clarify if the species and strains recovered in this analysis are novel, alongside the growth conditions and physiology (Park *et al.*, 2012). Finally, radio-labelled  $^{13}\text{C}$  assays or meta-transcriptomics could also be used to identify if the organisms or functional pathways were active. As metagenomics only highlights the functional potential of the microbial community (rather than the active fraction), it would be interesting to investigate which pathways were more active than others in this unique environment.

# Chapter 6: Concluding discussion

## 6.1 Summary

Through the rise of DNA sequencing, our understanding of global microbial diversity has increased. Advances through 16s rRNA amplicon sequencing has aided understanding of microbial community composition in natural environments, without the need for culturing. More recently, metagenomics has been used in medical and environmental applications to directly sequence community DNA without the need for amplification. This method allows both the functional diversity and taxonomy of communities to be investigated and is especially beneficial for unique microbial samples. The more commonly-used 16s rRNA sequencing relies on publicly-available databases to assign taxa to DNA sequences, however these databases do not capture global microbial diversity. Whilst taxonomy assignments can be made using metagenome data, the function of sequences or genomes that cannot be annotated with taxonomic databases can be investigated. Consequently, more information on the ecological and biogeochemical importance of the uncultured fraction of microbial communities can be understood.

Currently, limited guidance is available for metagenomic data analysis in the field of microbial ecology. Whilst raw metagenomic sequencing data can be directly analysed for taxonomy and function, assembling the DNA reads in to longer fragments (contigs) may be beneficial. DNA assembly can be carried out using a range of different assembly algorithms and aims to improve the alignment of DNA sequences to functional or taxonomic databases. DNA assembly is not commonly used in soil microbial ecology, and when it is, the choice of assembler and parameters often requires greater attention.

This study applied metagenomics to several areas in microbial ecology, to improve understanding of microbial function and taxonomy. Prior to this study, metagenomics had not been applied to understanding microbial succession in glacial forefields. Microbes have been proposed as the initial colonisers of newly exposed soil, but more research into how these communities change and the role they play in soil biogeochemical cycles was needed. In addition, the composition of the initial pioneer microbial community has been subject to debate, so metagenomic data may help to contribute to this discussion. This is significant given the continued retreat of glaciers, exposing more land for microbial colonisation. Furthermore, microbial nitrogen fixation (diazotrophy) has been previously identified as important for building up labile nitrogen stocks in oligotrophic forefield soils, facilitating the

colonisation of higher microbes and plants during succession. However, little research had been carried out into the diversity of these diazotrophs and how they vary between forefields. In addition to this, a significant fraction of global microbial diversity has yet to be explored, including the glacially fed fjord sediments of Chilean Patagonia. These fjords are hotspots of primary productivity and support commercially important Salmon fisheries. Using metagenomics in this region will help facilitate understanding of microbial diversity, function and biogeochemical significance in these unexplored sediments.

Broadly, this study aimed to contribute to the use of metagenomics in the field of environmental microbial ecology, in terms of providing both methodological advancements and to broaden understanding of microbial diversity. In particular, the objectives of this study were to: 1) evaluate metagenome assemblers for soil microbial ecology, and subsequently apply metagenomics to investigate; 2) microbial nitrogen fixation in Arctic glacier forefields; 3) microbial diversity during succession in an Arctic forefield and 4) the diversity and functional potential of microbial communities in benthic fjord sediments. The following section provides an overview of the key findings:

**Objective 1:** To compare the performance of five publicly available metagenome assemblers for soil bacterial communities

This study found de Bruijn graph based assemblers (CLC and metaSPAdes) to provide the highest coverage and contig lengths during metagenome assembly of artificial soil bacterial datasets. However, due to the increased complexity of this algorithm, these assemblers are more sensitive to parameterisation, in particular, to components such as the kmer length. Therefore, de Bruijn graph-based assemblers can provide high quality assemblies, but also produce the largest spread of values. Consequently, testing parameter values to suit the dataset in question may improve the output quality, over simply using default values. Additionally, this study also showed the importance of evaluating assembly quality using multiple metrics, covering assembly contiguity, size and completeness. Using a spread of metrics helps to avoid large fragmented datasets or long erroneous contigs, which would not be optimal for downstream analysis. Additionally, the use of artificial test datasets was shown to be beneficial in evaluating the accuracy of metagenome assembly in producing the correct taxonomic composition.

Overall, the study found assembler selection to have a significant influence on outcome quality. This has importance for the microbial ecology community, as it shows assembler selection is an important factor to consider during sequence analysis. Fitting the assembler

(and parameterisation) to the complexity of the dataset can improve assembly outcome and thus taxonomic and functional annotation. It is therefore recommended that assembler comparison, or justification of assembler choice is made during methodology development.

**Objective 2:** To investigate the similarities and differences in taxonomic composition of diazotrophic bacteria in metagenomes sampled from four Arctic glacier forefields

This study applied the assembler selected in Objective 1 (metaSPAdes) to investigate the diversity of nitrogen-fixing bacteria across four Arctic forefields (Storglaciären, Rabots, Russell and Midtre Lovénbreen). Metagenome assembly was applied to provide longer contigs for taxonomic annotation of *nif* genes (for nitrogen fixation). The study found a diverse range of diazotrophs across the forefields, including a core group of cyanobacteria, anaerobes and extremophiles, which were identified across sites. The composition of this core group may be related to adaptive mechanisms, including tolerance to the oligotrophic, high UV and cold conditions that are typical in forefield soils. This analysis provided a new *nifH* phylogeny, demonstrating the phylogenetic distribution of Arctic diazotrophs, in relation to sequenced relatives. The study contributes to our understanding of microbial diversity in the Arctic, including how a range of bacterial species may contribute to local biogeochemical cycling through nitrogen fixation. In addition, the study exemplified how metagenomics can be applied to functional and taxonomic analysis of microbial communities in extreme environments. The use of metagenomics allowed the taxonomy of functional (nitrogen fixation, *nif*) genes to be interpreted, which would not be possible with 16s analysis or culture-based methodologies. It is hoped this work will stimulate others in the microbial ecology community to apply metagenomics to widen understanding of microbial diversity.

**Objective 3:** To investigate the bulk microbial community composition along a chronosequence of soil succession in the Midtre Lovénbreen forefield, Svalbard, using metagenomics.

Objective 3 applied metagenomics to understand microbial community composition during soil succession along the Midtre Lovénbreen forefield, Svalbard. In this study, unassembled DNA sequencing was used for community taxonomic annotation, due to limitations of assembly resulting from the community complexity and sequencing coverage. However, sequences were assembled with metaSPAdes and used for functional interpretation and genome binning, to benefit from longer contig lengths.

During succession, forefield soil total nitrogen and total organic carbon content increased, attributed to allochthonous and autochthonous sources. This included both aeolian deposition and the fixation of carbon and nitrogen by autotrophic bacteria. Throughout the forefield,

aerobic, anaerobic and extremophilic species were identified, attributed to the low nutrient, high UV and cold conditions, in line with Objective 2. Newly-exposed soils were shown to contain both autotrophic cyanobacteria and heterotrophs, providing further information on the current debate surrounding the composition of pioneer communities. The presence of heterotrophic bacteria may relate to the identification of some organic carbon in early soils, which may be sourced from overridden material or aeolian deposition. Furthermore, from 3 years post ice retreat, organisms with carbon, nitrogen and sulfur cycling metabolisms were identified, in line with soil nutrient build up. Genome binning was used to support the taxonomy and also indicated the potential for novel strains and species in the soils.

Overall, this study has used metagenomics to aid understanding of how microbial community structure (and function) modified during soil succession. This is significant for highlighting how microbes may contribute to soil development and local biogeochemical cycles in the Arctic, a region which was once thought to be abiotic. This is especially significant given the expected increase in soil exposure in the Arctic with glacier retreat in future years, and thus, microbial colonisation. Furthermore, this study has also shown how metagenome assembly, whilst beneficial for functional analysis, may not always be the most useful approach. Given datasets that are highly complex or have low read coverage, read-based taxonomic annotation may provide a better overview, than that based on incomplete assemblies.

**Objective 4:** To investigate the composition and potential function of microbial communities sourced from benthic metagenomes in a Chilean fjord.

Objective 4 applied metagenomics to understand the microbial community composition of benthic fjord sediments, fed by glacial meltwater. The glacial outflow from the Steffen glacier has a distinct influence on fjord salinity, turbidity and temperature. However, as this freshwater was stratified as a surface lens, there was limited influence on the benthic sediments. These sediments were in more saline (marine) conditions, with very low total organic carbon and nitrogen content. Alongside metagenome assembly, read-based annotation was also carried out, due to the low read recruitment in assemblies, which may be due to limited DNA quality from this extreme sampling location. This highlights that metagenome assembly may not always be the best approach and the data type/quality in question needs to be considered. A substantial fraction of the DNA reads could not be annotated, which again may be related to DNA quality or the presence of unique genomes in the samples. The majority of the recovered community composition consisted of bacteria, including anaerobes and those with potential extremophilic adaptations, to the dark, cold, oligotrophic conditions. Results from genome binning revealed that over half the genomes did not match to NCBI GenBank accurately,



indicating a degree of species novelty may be present in the samples. This provides a focus for further single cell sequencing or culture-based studies, that may wish to isolate these genomes. Additionally, 16s rRNA and *drsAB* gene phylogenies were applied to demonstrate the potential of metagenomics for both taxonomic and functional analysis.

Overall, this study has shown how metagenomics can provide useful insights into community composition and function of uncultured, potentially unique, samples. This technology can provide insights into microbial diversity, which may help to focus further analysis, such as culturing or deep re-sequencing. However, it has also been shown how metagenome assembly may not always be the most optimal tool, given a dataset that is highly diverse or fragmented. It is therefore suggested that methodologies are constructed with consideration of the dataset at hand.

## **6.2 Limitations and future opportunities**

The studies here have demonstrated metagenomics to be a useful tool for understanding microbial diversity and functional potential. However, metagenomics is unable to provide information on the activity of microbes or functional pathways because it is based on DNA sequencing (Wooley *et al.*, 2010). Transcriptomics may be a useful addition to metagenomic analysis, which involves sequencing the microbial community mRNA (Moran *et al.*, 2013). Transcriptomics is therefore able to provide an insight into the functional pathways which are active at a snapshot in time. Whilst this cannot be extrapolated to daily (or yearly) activities, it is useful to gauge if functional pathways are indeed active. Incubation experiments may also be carried out to measure rates of productivity, for example through acetylene reduction assays for microbial nitrogen fixation (Burriss *et al.*, 1972; Hardy *et al.*, 1973).

Furthermore, sequencing coverage can also be an issue, especially for complex microbial communities, as shown in this analysis. It is difficult to determine the level of sequencing required *a priori*, and therefore all genomes in diverse community samples may not be recovered. To recover the less abundant microbial fraction, or to extract metagenome assembled genomes, deep resequencing can be carried out to improve the coverage (Narasimgarao *et al.*, 2012). However, this is not necessary if the analysis aims to only gauge an insight into the dominant microbial taxonomy and functions.

Currently, there is a lack of formalised guidance for the use of metagenomics by microbial ecologists. Multiple stages during the analysis require informed judgement, for example; to assemble the metagenome or not, the choice of assembler, to undertake genome binning or to target specific functional genes. If metagenomics is used to a greater degree in future, a

methodological consensus or guidance may be achieved. This study has contributed by evaluating the use of assembly and the choice of assemblers for soil microbial ecologists. However, the development of more formalised tutorials, guidance and analysis platforms would be beneficial to improve uptake of metagenomics by the research community. More recently, platforms such as KBASE have moved towards this, by providing interactive user interfaces for scientists without substantial bioinformatic experience.

### **6.3 Implications**

The results from this analysis have demonstrated the potential of metagenomics for understanding microbial taxonomy and functional potential within the field of environmental microbial ecology. Not only can metagenomics provide insights into the diversity and potential role of microbes in local biogeochemical cycles, it can also help guide further analysis, such as single cell sequencing, genome assembly and culturing. It is hoped that this work will help to demonstrate the benefit of metagenomics, which can be used as an alternative to, or alongside, 16s or targeted gene sequencing.

Here, we have used metagenomics to contribute to the understanding of microbial ecology in glaciated regions. In particular, the work has improved understanding of microbial function and diversity in glacial forefields, which has particular significance given continued ice retreat with climate change. This analysis has identified distinct changes to microbial diversity and functional potential during forefield succession, and supplied a new *nifH* phylogeny for forefield diazotrophs. In addition, metagenomics has provided new insights into microbial community structure and function in deep low biomass Chilean fjord sediments, a largely unexplored environment. We have also discussed the benefit of metagenome assembly and how the choice of assembler can impact the output quality and downstream analysis. It has been suggested that studies which use metagenomics consider assembler choice with respect to the dataset, and acknowledge that assembly may not always be the best approach for taxonomic analysis given highly diverse or fragmented metagenomes. This is because low abundance sequences may be excluded from sequencing in complex samples, and therefore the full diversity may not be profiled. This work has identified de Bruijn graph based assemblers as the most appropriate choice for soil metagenome assemblies, however consideration needs to be made to the data type, complexity and choice of parameters during assembler selection. Our recommendations include careful assembler selection, testing and parameter optimisation to provide an improved assembly outcome for downstream annotation.

It is hoped that the application of metagenomics in this work will encourage those in the field of soil microbial ecology to explore this technique. However, we suggest that scientists do so

with methodological care and caution. Further methodological inter-comparisons and guidance may encourage the use metagenomics in microbial ecology. This is because there is limited formalised structure available for metagenomic analysis, in contrast to the body of literature using 16s rRNA sequencing. Metagenomics is a high useful tool for understanding the uncultured microbial fraction and can help guide further studies and more detailed analysis. This not only has application and benefit for understanding global microbial diversity, but can be beneficial for applied studies, for example in biotechnology or anti-microbial resistance. Whilst metagenomics is a highly beneficial tool for understanding microbial diversity, it is best applied in a holistic approach, incorporating culture based studies, modelling and microscopy, to gain a full understanding of community structure and function.

## References

- Abatenh, E., Gizaw, B., Tsegaye, Z., & Tefera, G. (2018). Microbial function on climate change - a review. *Environmental Pollution and Climate Change*, 2(147), 10-4172.
- Akob, D. M., Baesman, S. M., Sutton, J. M., Fierst, J. L., Mumford, A. C., Shrestha, Y., ... & Oremland, R. S. (2017). Detection of diazotrophy in the acetylene-fermenting anaerobe *Pelobacter* sp. strain SFB93. *Applied and Environmental Microbiology*, 83(17).
- Albertsen, M., Hugenholtz, P., Skarshewski, A., Nielsen, K. L., Tyson, G. W., & Nielsen, P. H. (2013). Genome sequences of rare, uncultured bacteria obtained by differential coverage binning of multiple metagenomes. *Nature Biotechnology*, 31(6), 533.
- Alkan, C., Sajjadian, S., & Eichler, E. E. (2011). Limitations of next-generation genome sequence assembly. *Nature Methods*, 8(1), 61-65.
- Allison, S. D., & Treseder, K. K. (2008). Warming and drying suppress microbial activity and carbon cycling in boreal forest soils. *Global Change Biology*, 14(12), 2898-2909.
- Allison, V. J., Condon, L. M., Peltzer, D. A., Richardson, S. J., & Turner, B. L. (2007). Changes in enzyme activities and soil microbial community composition along carbon and nutrient gradients at the Franz Josef chronosequence, New Zealand. *Soil Biology and Biochemistry*, 39(7), 1770-1781
- Anesio, A. M., Hodson, A. J., Fritz, A., Psenner, R., & Sattler, B. (2009). High microbial activity on glaciers: importance to the global carbon cycle. *Global Change Biology*, 15(4), 955-960.
- Ansari, A. H., Hodson, A. J., Kaiser, J., & Marca-Bell, A. (2013). Stable isotopic evidence for nitrification and denitrification in a High Arctic glacial ecosystem. *Biogeochemistry*, 113(1-3), 341-357.
- Ansorge, W. J. (2009). Next-generation DNA sequencing techniques. *New Biotechnology*, 25(4), 195-203.
- Aracena, C., Lange, C. B., Iriarte, J. L., Rebolledo, L., & Pantoja, S. (2011). Latitudinal patterns of export production recorded in surface sediments of the Chilean Patagonian fjords (41–55 S) as a response to water column productivity. *Continental Shelf Research*, 31(3-4), 340-355.
- Arkin, A. P., Stevens, R. L., Cottingham, R. W., Maslov, S., Henry, C. S., Dehal, P., ... & Sneddon, M. W. (2016). The DOE systems biology knowledgebase (KBase). *bioRxiv*, 096354.
- Arumugam, M., Harrington, E.D., Foerstner K.U., Raes, J., Bork, P. (2010) SmashCommunity: a metagenomic annotation and analysis tool. *Bioinformatics*, 26, 2977–2978
- Aulakh, M. S., Doran, J. W., & Mosier, A. R. (1992). 'Soil denitrification—significance, measurement, and effects of management'. In Stewart, B.A. *Advances in Soil Science*. Springer: New York, 1-57.
- Aziz, R. K., Bartels, D., Best, A. A., DeJongh, M., Disz, T., Edwards, R. A., ... & Meyer, F. (2008). The RAST Server: rapid annotations using subsystems technology. *BMC Genomics*, 9(1), 75.
- Bankevich, A., Nurk, S., Antipov, D., Gurevich, A. A., Dvorkin, M., Kulikov, A. S., ... & Pyshkin, A. V. (2012). SPAdes: a new genome assembly algorithm and its applications to single-cell sequencing. *Journal of Computational Biology*, 19(5), 455-477.

- Banning, N. C., Maccarone, L. D., Fisk, L. M., & Murphy, D. V. (2015). Ammonia-oxidising bacteria not archaea dominate nitrification activity in semi-arid agricultural soil. *Scientific Reports*, 5, 11146.
- Barber, S. A. (1995) *Soil nutrient bioavailability: a mechanistic approach*. John Wiley & Sons: USA.
- Bardgett, R. D., Richter, A., Bol, R., Garnett, M. H., Bäumler, R., Xu, X., ... & Wanek, W. (2007). Heterotrophic microbial communities use ancient carbon following glacial retreat. *Biology Letters*, 3(5), 487-490.
- Bastide, M., & McCombie, W. R. (2007). Assembling genomic DNA sequences with PHRAP. *Current Protocols in Bioinformatics*, 11-4.
- Benn, D., & Evans, D. J. (2014). *Glaciers and Glaciation*. Routledge: London.
- Benson, D. A., Karsch-Mizrachi, I., Lipman, D. J., Ostell, J., & Wheeler, D. L. (2005). GenBank. *Nucleic Acids Research*, 33, 34-38.
- Benson, D. R., & Silvester, W. B. (1993). Biology of *Frankia* strains, actinomycete symbionts of actinorhizal plants. *Microbiological Reviews*, 57(2), 293-319.
- Bergman, B., Gallon, J. R., Rai, A. N., & Stal, L. J. (1997). N<sub>2</sub> fixation by non-heterocystous cyanobacteria. *FEMS Microbiology Reviews*, 19(3), 139-185.
- Bernasconi, S. M. (2008). Weathering, soil formation and initial ecosystem evolution on a glacier forefield: a case study from the Damma Glacier, Switzerland. *Mineralogical Magazine*, 72(1), 19-22.
- Bernasconi, S. M., Bauder, A., Bourdon, B., Brunner, I., Bünemann, E., Chris, I., ... & Frossard, E. (2011). Chemical and biological gradients along the Damma glacier soil chronosequence, Switzerland. *Vadose Zone Journal*, 10(3), 867-883.
- Bhatia, M., Sharp, M., & Foght, J. (2006). Distinct bacterial communities exist beneath a high Arctic polythermal glacier. *Applied and Environmental Microbiology*, 72(9), 5838-5845.
- Blainey, P. C. (2013). The future is now: single-cell genomics of bacteria and archaea. *FEMS Microbiology Reviews*, 37(3), 407-427
- Błażewicz, J., Formanowicz, P., Kasprzak, M., Schuurman, P., & Woeginger, G. J. (2002). DNA sequencing, Eulerian graphs and the exact perfect matching problem'. In Brandstadt, A., *Graph-Theoretic Concepts in Computer Science*. Springer: Berlin Heidelberg, 13-24.
- Blothe, M., Akob, D. M., Kostka, J. E., Goschel, K., Drake, H. L. & Kusel, K. (2008). pH gradient-induced heterogeneity of Fe(III)-reducing microorganisms in coal mining-associated lake sediments. *Applied Environmental Microbiology*, 74, 1019–1029.
- Bohmann, Kristine, Alice Evans, M. Thomas P. Gilbert, Gary R. Carvalho, Simon Creer, Michael Knapp, W. Yu Douglas, and Mark De Bruyn. (2014) Environmental DNA for wildlife biology and biodiversity monitoring. *Trends in Ecology & Evolution*, 29 (6), 358-367.
- Bolger, A. M., Lohse, M., & Usadel, B. (2014). Trimmomatic: a flexible trimmer for Illumina sequence data. *Bioinformatics*, 30(15), 2114-2120.

- Bonin, A. S., & Boone, D. R. (2006). 'The order methanobacteriales'. In: Dworkin, M., Falkow, S., Schleifer K.H., Stackebrandt, E. (Eds). *The Prokaryotes*. Springer: New York, 231-243.
- Boopathy, R., and Kulpa, C. F. (1993). Nitroaromatic compounds serve as nitrogen source for *Desulfovibrio* sp.(B strain). *Canadian Journal of Microbiology*, 39(4), 430-433.
- Bottrell, S. H., & Tranter, M. (2002). Sulfide oxidation under partially anoxic conditions at the bed of the Haut Glacier d'Arolla, Switzerland. *Hydrological Processes*, 16(12), 2363-2368.
- Boyd, E. S., Skidmore, M., Mitchell, A. C., Bakermans, C., & Peters, J. W. (2010). Methanogenesis in subglacial sediments. *Environmental Microbiology Reports*, 2(5), 685-692.
- Bradley J, Singarayer J, Anesio A. (2014) Microbial community dynamics in the forefield of glaciers. *Proceedings of the Royal Society B*, 281.
- Bradley, J. A., Anesio, A. M., Singarayer, J. S., Heath, M. R., and Arndt, S. (2015). SHIMMER (1.0): a novel mathematical model for microbial and biogeochemical dynamics in glacier forefield ecosystems. *Geoscientific Model Development Discussions*, 8(8), 6143-6216.
- Bradley, J. A., Arndt, S., Sabacká, M., Benning, L. G., Barker, G. L., Blacker, J. J., ... & Tranter, M. (2016). Microbial dynamics in a High Arctic glacier forefield: a combined field, laboratory, and modelling approach. *Biogeosciences*, 13(19), 5677.
- Brankatschk, R., Töwe, S., Kleineidam, K., Schloter, M., & Zeyer, J. (2011). Abundances and potential activities of nitrogen cycling microbial communities along a chronosequence of a glacier forefield. *The ISME Journal*, 5(6), 1025-1037.
- Branton, D., Deamer, D. W., Marziali, A., Bayley, H., Benner, S. A., Butler, T., ... & Jovanovich, S. B. (2008). The potential and challenges of nanopore sequencing. *Nature Biotechnology*, 26(10), 1146-1153.
- Bräutigam, A., Mullick, T., Schliesky, S., & Weber, A. P. (2011). Critical assessment of assembly strategies for non-model species mRNA-Seq data and application of next-generation sequencing to the comparison of C3 and C4 species. *Journal of Experimental Botany*, 62(9), 3093-3102.
- Breen, K., & Levesque, E. (2006). Proglacial succession of biological soil crusts and vascular plants: biotic interactions in the High Arctic. *Botany*, 84(11), 1714-1731.
- Bremner, J. M. (1965). Organic nitrogen in soils. *Soil Nitrogen*, 93-149.
- Brill, W. J. (1975). Regulation and genetics of bacterial nitrogen fixation. *Annual Reviews in Microbiology*, 29(1), 109-129.
- Broecker, W. S., & Peng, T. H. (1974). Gas exchange rates between air and sea. *Tellus*, 26(1-2), 21-35.
- Brunner, I., Plötze, M., Rieder, S., Zumsteg, A., Furrer, G., and Frey, B. (2011). Pioneering fungi from the Damma glacier forefield in the Swiss Alps can promote granite weathering. *Geobiology*, 9(3), 266-279.
- Burris, R. H. (1972). 'Nitrogen fixation—Assay methods and techniques'. In: Colowick, S.P. *Methods in enzymology*. Elsevier: Netherlands, 24, 415-431.

- Butler, J., MacCallum, I., Kleber, M., Shlyakhter, I. A., Belmonte, M. K., Lander, E. S., ... & Jaffe, D. B. (2008). ALLPATHS: de novo assembly of whole-genome shotgun microreads. *Genome Research*, 18(5), 810-820.
- Caccavo, F., Lonergan, D. J., Lovley, D. R., Davis, M., Stolz, J. F., & McInerney, M. J. (1994). *Geobacter sulfurreducens* sp. nov., a hydrogen-and acetate-oxidizing dissimilatory metal-reducing microorganism. *Applied and Environmental Microbiology*, 60(10), 3752-3759.
- Cameron, K. A., Hodson, A. J., & Osborn, A. M. (2012). Carbon and nitrogen biogeochemical cycling potentials of supraglacial cryoconite communities. *Polar Biology*, 35(9), 1375-1393.
- Carvalhais, L. C., Dennis, P. G., Badri, D. V., Tyson, G. W., Vivanco, J. M., & Schenk, P. M. (2013). Activation of the jasmonic acid plant defence pathway alters the composition of rhizosphere bacterial communities. *PLoS One*, 8(2), e56457.
- Cattonaro, F., Policriti, A., & Vezzi, F. (2010). *Enhanced reference guided assembly*. Bioinformatics and Biomedicine (BIBM). IEEE International Conference.
- Chaia, E. E., Wall, L. G., & Huss-Danell, K. (2010). Life in soil by the actinorhizal root nodule endophyte Frankia. A review. *Symbiosis*, 51(3), 201-226.
- Challacombe, J. F., Majid, S., Deole, R., Brettin, T. S., Bruce, D., Delano, S. F., ... & Reitenga, K. G. (2013). Complete genome sequence of *Halorhodospira halophila* SL1. *Standards in Genomic Sciences*, 8(2), 206.
- Chapin, F. S., Walker, L. R., Fastie, C. L., & Sharman, L. C. (1994). Mechanisms of primary succession following deglaciation at Glacier Bay, Alaska. *Ecological Monographs*, 64(2), 149-175.
- Chattopadhyay, M. K. (2006). Mechanism of bacterial adaptation to low temperature. *Journal of Biosciences*, 31(1), 157-165.
- Chen M,A., Markowitz, V.M., Chu, K., Palaniappan, K., Szeto, E., Pillay, M. (2017). IMG/M: integrated genome and metagenome comparative data analysis system. *Nucleic Acids Research*, 45.
- Chevreux (2014) MIRA [Online] Available from: <http://www.chevreux.org/index0.html> [Accessed 25-01-2018].
- Chien, Y. T., and Zinder, S. H. (1996). Cloning, functional organization, transcript studies, and phylogenetic analysis of the complete nitrogenase structural genes (nifHDK2) and associated genes in the archaeon *Methanosarcina barkeri* 227. *Journal of Bacteriology*, 178(1), 143-148.
- Chikhi, R., & Medvedev, P. (2013). Informed and automated k-mer size selection for genome assembly. *Bioinformatics*, 30(1), 31-37.
- Childers, Susan E., Stacy Ciufo, and Derek R. Lovley. (2002). *Geobacter metallireducens* accesses insoluble Fe (III) oxide by chemotaxis. *Nature*, 416 (6882), 767-769.
- Christmas, N. A., Barker, G., Anesio, A. M., & Sánchez-Baracaldo, P. (2016). Genomic mechanisms for cold tolerance and production of exopolysaccharides in the Arctic cyanobacterium *Phormidesmis priestleyi* BC1401. *BMC Genomics*, 17(1), 533.
- Christner, B. C., Kvitko, B. H., & Reeve, J. N. (2003). Molecular identification of bacteria and eukarya inhabiting an Antarctic cryoconite hole. *Extremophiles*, 7(3), 177-183.

- Christner, B. C., Skidmore, M. L., Priscu, J. C., Tranter, M., & Foreman, C. M. (2008). 'Bacteria in subglacial environments'. In: Margesin, R., Schinner, F., Marx, J.C., Gerday, C. (Eds) *Psychrophiles: from biodiversity to biotechnology*. Springer: Berlin Heidelberg, 51-71.
- Chu, H., Fierer, N., Lauber, C. L., Caporaso, J. G., Knight, R., & Grogan, P. (2010). Soil bacterial diversity in the Arctic is not fundamentally different from that found in other biomes. *Environmental Microbiology*, 12(11), 2998-3006.
- Chu, V. W. (2014). Greenland ice sheet hydrology: A review. *Progress in Physical Geography*, 38(1), 19-54.
- Collins, D. N. (1979). Hydrochemistry of meltwaters draining from an alpine glacier. *Arctic and Alpine Research*, 307-324.
- Conen, F., Yakutin, M. V., Zumbrunn, T., & Leifeld, J. (2007). Organic carbon and microbial biomass in two soil development chronosequences following glacial retreat. *European Journal of Soil Science*, 58(3), 758-762.
- Cong, J., Yang, Y., Liu, X., Lu, H., Liu, X., Zhou, J., ... & Zhang, Y. (2015). Analyses of soil microbial community compositions and functional genes reveal potential consequences of natural forest succession. *Scientific Reports*, 5.
- Conway, T., Wazny, J., Bromage, A., Zobel, J., & Beresford-Smith, B. (2012). Gossamer—a resource-efficient de novo assembler. *Bioinformatics*, 28(14), 1937-1938.
- Copeland, A., Zeytun, A., Yassawong, M., Nolan, M., Lucas, S., Hammon, N., ... & Goodwin, L. A. (2012). Complete genome sequence of the orange-red pigmented, radioresistant *Deinococcus proteolyticus* type strain (MRP T). *Standards in Genomic Sciences*, 6(2), 240.
- Cowan, D., Meyer, Q., Stafford, W., Muyanga, S., Cameron, R., & Wittwer, P. (2005). Metagenomic gene discovery: past, present and future. *Trends in Biotechnology*, 23(6), 321-329.
- Cowton, T., Nienow, P., Bartholomew, I., Sole, A., & Mair, D. (2012). Rapid erosion beneath the Greenland ice sheet. *Geology*, 40(4), 343-346.
- D'Amico, M. E., Freppaz, M., Filippa, G., & Zanini, E. (2014). Vegetation influence on soil formation rate in a proglacial chronosequence (Lys Glacier, NW Italian Alps). *Catena*, 113, 122-137.
- Daims, H., Lebedeva, E. V., Pjevac, P., Han, P., Herbold, C., Albertsen, M., ... & Kirkegaard, R. H. (2015). Complete nitrification by *Nitrospira* bacteria. *Nature*, 528(7583), 504.
- Daniel, R. (2005). The metagenomics of soil. *Nature Reviews Microbiology*, 3(6), 470-478.
- Darriba, D., Taboada, G.L., Doallo, R., Posada, D. (2012) jModelTest 2: more models, new heuristics and parallel computing. *Nature Methods*, 9(8), 772.
- DasSarma, S., & DasSarma, P. (2006). Halophiles. *Engineering in Life Sciences*, 1-13.
- Davidson, E. A., & Janssens, I. A. (2006). Temperature sensitivity of soil carbon decomposition and feedbacks to climate change. *Nature*, 440(7081), 165-173.
- Davila, P.M., D. Figueroa, & E. Muller. (2002). Freshwater input into the coastal ocean and its relation with the salinity distribution off austral Chile. *Continental Shelf Research*, 22(3): p. 521-534.



- Delmont, T. O., Robe, P., Cecillon, S., Clark, I. M., Constancias, F., Simonet, P., ... & Vogel, T. M. (2011). Accessing the soil metagenome for studies of microbial diversity. *Applied and Environmental Microbiology*, 77(4), 1315-1324.
- Deng, X., Naccache, S. N., Ng, T., Federman, S., Li, L., Chiu, C. Y., & Delwart, E. L. (2015). An ensemble strategy that significantly improves de novo assembly of microbial genomes from metagenomic next-generation sequencing data. *Nucleic Acids Research*, 43(7), e46-e46.
- Denison, R. F., & Kiers, E. T. (2004). Why are most rhizobia beneficial to their plant hosts, rather than parasitic?. *Microbes and Infection*, 6(13), 1235-1239.
- Denisov, G., Walenz, B., Halpern, A. L., Miller, J., Axelrod, N., Levy, S., & Sutton, G. (2008). Consensus generation and variant detection by Celera Assembler. *Bioinformatics*, 24(8), 1035-1040.
- Deslippe, J. R., and Egger, K. N. (2006). Molecular diversity of nifH genes from bacteria associated with high arctic dwarf shrubs. *Microbial Ecology*, 51(4), 516-525.
- Dethier, M. N., & Schoch, G. C. (2005). The consequences of scale: assessing the distribution of benthic populations in a complex estuarine fjord. *Estuarine, Coastal and Shelf Science*, 62(1), 253-270.
- Dixon, R. and Kahn, D. (2004). Genetic regulation of biological nitrogen fixation. *Nature reviews Microbiology*, 2.8.
- Dreyfus, B., Garcia, J. L., and Gillis, M. (1988). Characterization of Azorhizobium caulinodans gen. nov., sp. nov., a stem-nodulating nitrogen-fixing bacterium isolated from Sesbania rostrata. *International Journal of Systematic and Evolutionary Microbiology*, 38(1), 89-98.
- Duc, L., Noll, M., Meier, B. E., Bürgmann, H., and Zeyer, J. (2009). High diversity of diazotrophs in the forefield of a receding alpine glacier. *Microbial Ecology*, 57(1), 179-190.
- Dutton, A., Carlson, A. E., Long, A. J., Milne, G. A., Clark, P. U., DeConto, R., ... & Raymo, M. E. (2015). Sea-level rise due to polar ice-sheet mass loss during past warm periods. *Science*, 349(6244).
- Earl, D., Bradnam, K., John, J. S., Darling, A., Lin, D., Fass, J., ... & Nguyen, N. (2011). Assemblathon 1: a competitive assessment of de novo short read assembly methods. *Genome Research*, 21(12), 2224-2241.
- Edwards, A., & Cook, S. (2015). Microbial dynamics in glacier forefield soils show succession is not just skin deep. *Molecular Ecology*, 24(5), 963-966.
- Egli, M., Filip, D., Mavris, C., Fischer, B., Götze, J., Raimondi, S., & Seibert, J. (2012). Rapid transformation of inorganic to organic and plant-available phosphorous in soils of a glacier forefield. *Geoderma*, 189, 215-226.
- Eisen, J. A., Nelson, K. E., Paulsen, I. T., Heidelberg, J. F., Wu, M., Dodson, R. J., ... & Hickey, E. K. (2002). The complete genome sequence of Chlorobium tepidum TLS, a photosynthetic, anaerobic, green-sulfur bacterium. *Proceedings of the National Academy of Sciences*, 99(14), 9509-9514.
- Eloe-Fadrosh, E. A., Ivanova, N. N., Woyke, T., & Kyrpides, N. C. (2016a). Metagenomics uncovers gaps in amplicon-based detection of microbial diversity. *Nature Microbiology*, 1(4), 15032.

- Eloe-Fadrosh, E. A., Paez-Espino, D., Jarett, J., Dunfield, P. F., Hedlund, B. P., Dekas, A. E., ... & Li, W. J. (2016b). Global metagenomic survey reveals a new bacterial candidate phylum in geothermal springs. *Nature communications*, 7, 10476.
- Erlich, H. A. (1989). *PCR technology*. Stockton press: UK.
- Fagerli, I. L., & Svenning, M. M. (2005). Arctic and subarctic soil populations of *Rhizobium leguminosarum* biovar *trifolii* nodulating three different clover species: characterisation by diversity at chromosomal and symbiosis loci. *Plant and soil*, 275(1), 371-381.
- Fastie, C. L. (1995). Causes and ecosystem consequences of multiple pathways of primary succession at Glacier Bay, Alaska. *Ecology*, 76(6), 1899-1916.
- Ferrer, M., Martínez-Abarca, F., & Golyshin, P. N. (2005). Mining genomes and 'metagenomes' for novel catalysts. *Current Opinion in Biotechnology*, 16(6), 588-593.
- Fierer, N., Breitbart, M., Nulton, J., Salamon, P., Lozupone, C., Jones, R., ... and Knight, R. (2007). Metagenomic and small-subunit rRNA analyses reveal the genetic diversity of bacteria, archaea, fungi, and viruses in soil. *Applied and Environmental Microbiology*, 73(21), 7059-7066.
- Fierer, N., Jackson, J. A., Vilgalys, R., & Jackson, R. B. (2005). Assessment of soil microbial community structure by use of taxon-specific quantitative PCR assays. *Applied and Environmental Microbiology*, 71(7), 4117-4120.
- Fierer, N., Leff, J. W., Adams, B. J., Nielsen, U. N., Bates, S. T., Lauber, C. L., ... & Caporaso, J. G. (2012). Cross-biome metagenomic analyses of soil microbial communities and their functional attributes. *Proceedings of the National Academy of Sciences*, 109(52), 21390-21395.
- Finotello, F., Lavezzo, E., Fontana, P., Peruzzo, D., Albiero, A., Barzon, L., ... & Toppo, S. (2011). Comparative analysis of algorithms for whole-genome assembly of pyrosequencing data. *Briefings in Bioinformatics*, bbr063.
- Flemming, H. C., & Wingender, J. (2010). The biofilm matrix. *Nature Reviews Microbiology*, 8(9), 623.
- Franche, C., Lindström, K., & Elmerich, C. (2009). Nitrogen-fixing bacteria associated with leguminous and non-leguminous plants. *Plant and Soil*, 321(1-2), 35-59.
- Francis, C. A., Roberts, K. J., Beman, J. M., Santoro, A. E., & Oakley, B. B. (2005). Ubiquity and diversity of ammonia-oxidizing archaea in water columns and sediments of the ocean. *Proceedings of the National Academy of Sciences*, 102(41), 14683-14688.
- Franzosa, E. A., Hsu, T., Sirota-Madi, A., Shafquat, A., Abu-Ali, G., Morgan, X. C., & Huttenhower, C. (2015). Sequencing and beyond: integrating molecular'omics' for microbial community profiling. *Nature Reviews Microbiology*, 13(6), 360.
- Freeman, K.R., Pescador, M.Y., Reed, S.C., Costello, E.K., Robeson, M.S., Schmidt, S.K. (2009) Soil CO<sub>2</sub> flux and photoautotrophic community composition in high-elevation, 'barren' soil. *Environmental Microbiology*, 11, 674–686.
- Frenot, Y., Gloaguen, J. C., Cannavacciuolo, M., & Bellido, A. (1998). Primary succession on glacier forelands in the subantarctic Kerguelen Islands. *Journal of Vegetation Science*, 9(1), 75-84.

- Frey, B., Bühler, L., Schmutz, S., Zumsteg, A., & Furrer, G. (2013). Molecular characterization of phototrophic microorganisms in the forefield of a receding glacier in the Swiss Alps. *Environmental Research Letters*, 8(1), 015033.
- Frey, B., Rieder, S. R., Brunner, I., Plötze, M., Koetzsch, S., Lapanje, A., ... & Furrer, G. (2010). Weathering-associated bacteria from the Damma glacier forefield: physiological capabilities and impact on granite dissolution. *Applied and Environmental Microbiology*, 76(14), 4788-4796.
- Gaby, J. C., & Buckley, D. H. (2012). A comprehensive evaluation of PCR primers to amplify the *nifH* gene of nitrogenase. *PloS one*, 7(7), e42149.
- Glass, E. M., Wilkening, J., Wilke, A., Antonopoulos, D., & Meyer, F. (2010). Using the metagenomics RAST server (MG-RAST) for analyzing shotgun metagenomes. *Cold Spring Harbor Protocols*, (1).
- González, H. E., Castro, L. R., Daneri, G., Iriarte, J. L., Silva, N., Tapia, F., ... & Vargas, C. A. (2013). Land–ocean gradient in haline stratification and its effects on plankton dynamics and trophic carbon fluxes in Chilean Patagonian fjords (47–50 S). *Progress in Oceanography*, 119, 32-47.
- Göransson, H., Venterink, H. O., & Bååth, E. (2011). Soil bacterial growth and nutrient limitation along a chronosequence from a glacier forefield. *Soil Biology and Biochemistry*, 43(6), 1333-1340.
- Grada, A., & Weinbrecht, K. (2013). Next-generation sequencing: methodology and application. *Journal of Investigative Dermatology*, 133(8), e11.
- Gruber, N. (2008). The marine nitrogen cycle: overview and challenges. *Nitrogen in the Marine Environment*, 2, 1-50.
- Guelland, K., Hagedorn, F., Smittenberg, R. H., Göransson, H., Bernasconi, S. M., Hajdas, I., & Kretzschmar, R. (2013). Evolution of carbon fluxes during initial soil formation along the forefield of Damma glacier, Switzerland. *Biogeochemistry*, 113(1-3), 545-561.
- Guindon, S. and Gascuel, O. (2003). A simple, fast and accurate method to estimate large phylogenies by maximum-likelihood. *Systematic Biology*, 52, 696-704.
- Gurevich, A., Saveliev, V., Vyahhi, N., & Tesler, G. (2013). QUASt: quality assessment tool for genome assemblies. *Bioinformatics*, 29(8), 1072-1075.
- Gutiérrez, M. H., Galand, P. E., Moffat, C., & Pantoja, S. (2015). Melting glacier impacts community structure of Bacteria, Archaea and Fungi in a Chilean Patagonia fjord. *Environmental Microbiology*, 17(10), 3882-3897.
- Hahn, A. S., & Quideau, S. A. (2013). Shifts in soil microbial community biomass and resource utilization along a Canadian glacier chronosequence. *Canadian Journal of Soil Science*, 93(3), 305-318.
- Hajirasouliha, I., Hormozdiari, F., Alkan, C., Kidd, J. M., Birol, I., Eichler, E. E., & Sahinalp, S. C. (2010). Detection and characterization of novel sequence insertions using paired-end next-generation sequencing. *Bioinformatics*, 26(10), 1277-1283.
- Handelsman, J. (2004). Metagenomics: application of genomics to uncultured microorganisms. *Microbiology and Molecular Biology Reviews*, 68(4), 669-685.

- Handelsman, J., Rondon, M. R., Brady, S. F., Clardy, J., & Goodman, R. M. (1998). Molecular biological access to the chemistry of unknown soil microbes: a new frontier for natural products. *Chemistry & Biology*, 5(10).
- Hardy, R., Burns, R. C., & Holsten, R. D. (1973). Applications of the acetylene-ethylene assay for measurement of nitrogen fixation. *Soil Biology and Biochemistry*, 5(1), 47-81.
- Hawkings, J. R., Wadham, J. L., Tranter, M., Lawson, E., Sole, A., Cowton, T., ... & Telling, J. (2015). The effect of warming climate on nutrient and solute export from the Greenland Ice Sheet. *Geochemical Perspective Letters*, 1, 94-104.
- Hawkings, J. R., Wadham, J. L., Tranter, M., Raiswell, R., Benning, L. G., Statham, P. J., ... & Telling, J. (2014). Ice sheets as a significant source of highly reactive nanoparticulate iron to the oceans. *Nature Communications*, 5.
- He, Z., Zhang, H., Gao, S., Lercher, M. J., Chen, W. H., and Hu, S. (2016). Evolvew v2: an online visualization and management tool for customized and annotated phylogenetic trees. *Nucleic acids research*, 44(W1).
- Hedges, J. I., and Stern, J. H. (1984). Carbon and nitrogen determinations of carbonate-containing solids. *Limnology and Oceanography*, 29(3), 657-663.
- Hernandez, D., François, P., Farinelli, L., Østerås, M., & Schrenzel, J. (2008). De novo bacterial genome sequencing: millions of very short reads assembled on a desktop computer. *Genome Research*, 18(5), 802-809.
- Hodkinson, I. D., Coulson, S. J., & Webb, N. R. (2003). Community assembly along proglacial chronosequences in the high Arctic: vegetation and soil development in north-west Svalbard. *Journal of Ecology*, 91(4), 651-663.
- Hodkinson, I. D., Webb, N. R., & Coulson, S. J. (2002). Primary community assembly on land—the missing stages: why are the heterotrophic organisms always there first? *Journal of Ecology*, 90(3), 569-577.
- Hodson, A. (2006). Biogeochemistry of snowmelt in an Antarctic glacial ecosystem. *Water Resources Research*, 42.
- Hodson, A. J., Mumford, P. N., Kohler, J., & Wynn, P. M. (2005). The High Arctic glacial ecosystem: new insights from nutrient budgets. *Biogeochemistry*, 72(2), 233-256.
- Hodson, A., Anesio, A. M., Tranter, M., Fountain, A., Osborn, M., Priscu, J., ... & Sattler, B. (2008). Glacial ecosystems. *Ecological Monographs*, 78(1), 41-67.
- Holmer, M., & Storkholm, P. (2001). Sulfate reduction and sulfur cycling in lake sediments: a review. *Freshwater Biology*, 46(4), 431-451.
- Hood, E., & Scott, D. (2008). Riverine organic matter and nutrients in southeast Alaska affected by glacial coverage. *Nature Geoscience*, 1(9), 583-587.
- Hood, E., Fellman, J., Spencer, R. G., Hernes, P. J., Edwards, R., D'Amore, D., & Scott, D. (2009). Glaciers as a source of ancient and labile organic matter to the marine environment. *Nature*, 462(7276), 1044-1047.
- Hopwood, M.J., Statham, P.J., Tranter, M & Wadham, J.L. (2014) Glacial flours as a potential source of Fe(II) and Fe(III) to polar waters. *Biogeochemistry*, 443-452.

- Horner, D. S., Pavesi, G., Castrignanò, T., De Meo, P. D. O., Liuni, S., Sammeth, M., ... & Pesole, G. (2009). Bioinformatics approaches for genomics and post genomics applications of next-generation sequencing. *Briefings in Bioinformatics*, bbp046.
- Howard, J.B., and Rees, D.C. (1996) Structural basis of biological nitrogen fixation. *Chem Rev*, 96: 2965–2982.
- Howe, A. C., Jansson, J. K., Malfatti, S. A., Tringe, S. G., Tiedje, J. M., & Brown, C. T. (2014). Tackling soil diversity with the assembly of large, complex metagenomes. *Proceedings of the National Academy of Sciences*, 111(13), 4904-4909.
- Howe, A., & Chain, P. S. (2015). Challenges and opportunities in understanding microbial communities with metagenome assembly (accompanied by IPython Notebook tutorial). *Frontiers in Microbiology*, 6, 678.
- Huang, X., & Madan, A. (1999). CAP3: A DNA sequence assembly program. *Genome Research*, 9(9), 868-877.
- Huerta-Cepas, J., Szklarczyk, D., Forslund, K., Cook, H., Heller, D., Walter, M. C., ... & Jensen, L. J. (2015). eggNOG 4.5: a hierarchical orthology framework with improved functional annotations for eukaryotic, prokaryotic and viral sequences. *Nucleic Acids Research*, 44(D1), D286-D293.
- Hugenholtz, P., & Tyson, G. W. (2008). Microbiology: metagenomics. *Nature*, 455(7212), 481.
- Hugerth, L. W., Larsson, J., Alneberg, J., Lindh, M. V., Legrand, C., Pinhassi, J., & Andersson, A. F. (2015). Metagenome-assembled genomes uncover a global brackish microbiome. *Genome Biology*, 16(1), 279.
- Hurek, T., and Reinhold-Hurek, B. (1995). Identification of grass-associated and toluene-degrading diazotrophs, *Azoarcus* spp., by analyses of partial 16S ribosomal DNA sequences. *Applied and Environmental Microbiology*, 61(6), 2257-2261.
- Huson, D. H., Auch, A. F., Qi, J., & Schuster, S. C. (2007). MEGAN analysis of metagenomic data. *Genome Research*, 17(3).
- Huson, D. H., Beier, S., Flade, I., Górska, A., El-Hadidi, M., Mitra, S., ... and Tappu, R. (2016). MEGAN community edition-interactive exploration and analysis of large-scale microbiome sequencing data. *PLoS Computational Biology*, 12(6), e1004957.
- Huson, D. H., Halpern, A. L., Lai, Z., Myers, E. W., Reinert, K., & Sutton, G. G. (2001). Comparing assemblies using fragments and mate-pairs. *Algorithms in Bioinformatics*, 294-306.
- Illumina (2014). Estimating sequencing coverage [Online]. Available from: [https://www.illumina.com/documents/products/technotes/technote\\_coverage\\_calculation.pdf](https://www.illumina.com/documents/products/technotes/technote_coverage_calculation.pdf) [Accessed 13-2-2018].
- Imhoff, J. F. (2005). 'Chromatiales ord. nov'. In Holt. J., Sneath. P., *Bergey's Manual of systematic bacteriology*. Springer: Boston, 1-59.
- Imhoff, J. F., and Pfennig, N. (2001). Thioflavicoccus mobilis gen. nov., sp. nov., a novel purple sulfur bacterium with bacteriochlorophyll b. *International Journal of Systematic and Evolutionary Microbiology*, 51(1), 105-110.

- Inceoğlu, Ö., Hoogwout, E. F., Hill, P., & van Elsas, J. D. (2010). Effect of DNA extraction method on the apparent microbial diversity of soil. *Applied and Environmental Microbiology*, 76(10), 3378-3382.
- IPCC, (2013). *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*. Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.). Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 1535.
- Irgens, R. L., Gosink, J. J., and Staley, J. T. (1996). *Polaromonas vacuolata* gen. nov., sp. nov., a psychrophilic, marine, gas vacuolate bacterium from Antarctica. *International Journal of Systematic and Evolutionary Microbiology*, 46(3), 822-826.
- Iriarte, J. L., González, H. E., & Nahuelhual, L. (2010). Patagonian fjord ecosystems in southern Chile as a highly vulnerable region: problems and needs. *AMBIO: A Journal of the Human Environment*, 39(7), 463-466.
- Iriarte, J. L., González, H. E., Liu, K. K., Rivas, C., & Valenzuela, C. (2007). Spatial and temporal variability of chlorophyll and primary productivity in surface waters of southern Chile (41.5–43 S). *Estuarine, Coastal and Shelf Science*, 74(3), 471-480.
- Iriarte, J. L., Pantoja, S., & Daneri, G. (2014). Oceanographic processes in Chilean Fjords of Patagonia: from small to large-scale studies. *Progress in Oceanography*, 129, 1-7.
- Ishii, S. I., Suzuki, S., Tenney, A., Norden-Krichmar, T. M., Nealson, K. H., & Bretschger, O. (2015). Microbial metabolic networks in a complex electrogenic biofilm recovered from a stimulus-induced metatranscriptomics approach. *Scientific Reports*, 5.
- Janssen, P. H., Yates, P. S., Grinton, B. E., Taylor, P. M., & Sait, M. (2002). Improved culturability of soil bacteria and isolation in pure culture of novel members of the divisions Acidobacteria, Actinobacteria, Proteobacteria, and Verrucomicrobia. *Applied and Environmental Microbiology*, 68(5), 2391-2396.
- Jetten, M. S. (2008). The microbial nitrogen cycle. *Environmental Microbiology*, 10(11), 2903-2909.
- Joan, L., Slonczewski, & John W. F. (2011), *Microbiology: An Evolving Science (2nd Edition)*, Norton: London.
- Jumpponen, A. (2003). Soil fungal community assembly in a primary successional glacier forefront ecosystem as inferred from rDNA sequence analyses. *New Phytologist*, 158(3), 569-578.
- Kandeler, E., Deiglmayr, K., Tschirko, D., Bru, D., & Philippot, L. (2006). Abundance of narG, nirS, nirK, and nosZ genes of denitrifying bacteria during primary successions of a glacier foreland. *Applied and Environmental Microbiology*, 72(9), 5957-5962.
- Kaštovská, K., Elster, J., Stibal, M., & Šantrůčková, H. (2005). Microbial assemblages in soil microbial succession after glacial retreat in Svalbard (High Arctic). *Microbial Ecology*, 50(3), 396.
- Keck, A., Wiktor, J., Hapter, R., & Nilsen, R. (2000). Phytoplankton assemblages related to physical gradients in an arctic, glacier-fed fjord in summer. *ICES Journal of Marine Science*, 56, 203-214.

- Kirk, J. L., Beaudette, L. A., Hart, M., Moutoglis, P., Klironomos, J. N., Lee, H., & Trevors, J. T. (2004). Methods of studying soil microbial diversity. *Journal of Microbiological Methods*, 58(2), 169-188.
- Kleindienst, S., Herbst, F. A., Stagars, M., Von Netzer, F., Von Bergen, M., Seifert, J., ... & Knittel, K. (2014). Diverse sulfate-reducing bacteria of the Desulfosarcina/Desulfococcus clade are the key alkane degraders at marine seeps. *The ISME journal*, 8(10), 2029.
- Knelman, J. E., Legg, T. M., O'Neill, S. P., Washenberger, C. L., González, A., Cleveland, C. C., & Nemergut, D. R. (2012). Bacterial community structure and function change in association with colonizer plants during early primary succession in a glacier forefield. *Soil Biology and Biochemistry*, 46, 172-180.
- Knight, P.G (1999). *Galciars*. Stanley Thornes: Cheltenham UK.
- Konstantinidis, K. T., & Tiedje, J. M. (2005). Genomic insights that advance the species definition for prokaryotes. *Proceedings of the National Academy of Sciences*, 102(7), 2567-2572.
- Kuever, J. (2014). 'The family Desulfobacteraceae'. In: Dworkin, M., Falkow, S., Schleifer K.H., Stackebrandt, E. (Eds) *The Prokaryotes*. Springer: Berlin, Heidelberg, 45-73.
- Kumar, S., & Blaxter, M. L. (2010). Comparing de novo assemblers for 454 transcriptome data. *BMC genomics*, 11(1), 571.
- Kunin, V., Copeland, A., Lapidus, A., Mavromatis, K. & Hugenholtz, P. (2008) A bioinformatician's guide to metagenomics. *Molecular Biology Reviews*, 72, 557–578.
- Kurtz, S., Phillippy, A., Delcher, A. L., Smoot, M., Shumway, M., Antonescu, C., & Salzberg, S. L. (2004). Versatile and open software for comparing large genomes. *Genome Biology*, 5(2), R12.
- Lai, B., Ding, R., Li, Y., Duan, L., & Zhu, H. (2012). A de novo metagenomic assembly program for shotgun DNA reads. *Bioinformatics*, 28(11), 1455-1462.
- Landaeta, M. F., López, G., Suárez-Donoso, N., Bustos, C. A., & Balbontín, F. (2012). Larval fish distribution, growth and feeding in Patagonian fjords: potential effects of freshwater discharge. *Environmental Biology of Fishes*, 93(1), 73-87.
- Langmead, B., & Salzberg, S. L. (2012). Fast gapped-read alignment with Bowtie 2. *Nature Methods*, 9(4), 357.
- Lawson, E. C., Wadham, J. L., Tranter, M., Stibal, M., Lis, G. P., Butler, C. E., ... & Dewsbury, P. (2014). Greenland Ice Sheet exports labile organic carbon to the Arctic oceans. *Biogeosciences*, 11(14), 4015-4028.
- Lee, B., Han, Y., Huh, Y., Lundstrom, C., Siame, L. L., Lee, J. I., ... & ASTER Team. (2013). Chemical and physical weathering in south Patagonian rivers: A combined Sr–U–Be isotope approach. *Geochimica et Cosmochimica Acta*, 101, 173-190.
- Lenaerts, J. T., Van Den Broeke, M. R., van Wessem, J. M., van de Berg, W. J., van Meijgaard, E., van Ulf, L. H., & Schaefer, M. (2014). Extreme precipitation and climate gradients in Patagonia revealed by high-resolution regional atmospheric climate modeling. *Journal of Climate*, 27(12), 4607-4621.

- Leray, M., & Knowlton, N. (2015). DNA barcoding and metabarcoding of standardized samples reveal patterns of marine benthic diversity. *Proceedings of the National Academy of Sciences*, 112(7), 2076-2081.
- Li, H. and Durbin R. (2009) Fast and accurate short read alignment with Burrows-Wheeler Transform. *Bioinformatics*, 25, 1754-60.
- Li, Y., Hu, Y., Bolund, L., & Wang, J. (2010). State of the art de novo assembly of human genomes from massively parallel sequencing data. *Human Genomics*, 4(4), 1.
- Lin, Y., Li, J., Shen, H., Zhang, L., Papasian, C. J., & Deng, H. W. (2011). Comparative studies of de novo assembly tools for next-generation sequencing technologies. *Bioinformatics*, 27(15), 2031-2037.
- Liu, G. X., Hu, P., Zhang, W., Wu, X., Yang, X., Chen, T., ... and Li, S. W. (2012). Variations in soil culturable bacteria communities and biochemical characteristics in the Dongkemadi glacier forefield along a chronosequence. *Folia Microbiologica*, 57(6), 485-494.
- Liu, K., Warnow, T. J., Holder, M. T., Nelesen, S. M., Yu, J., Stamatakis, A. P., and Linder, C. R. (2011). SATE-II: very fast and accurate simultaneous estimation of multiple sequence alignments and phylogenetic trees. *Systematic Biology*, 61(1), 90-106.
- Liu, Y., Yao, T., Jiao, N., Kang, S., Zeng, Y., & Huang, S. (2006). Microbial community structure in moraine lakes and glacial meltwaters, Mount Everest. *FEMS microbiology letters*, 265(1), 98-105.
- Lovley, D. R., Giovannoni, S. J., White, D. C., Champine, J. E., Phillips, E. J. P., Gorby, Y. A., & Goodwin, S. (1993). *Geobacter metallireducens* gen. nov. sp. nov., a microorganism capable of coupling the complete oxidation of organic compounds to the reduction of iron and other metals. *Archives of Microbiology*, 159(4), 336-344.
- Lowe, R., Shirley, N., Bleackley, M., Dolan, S., & Shafee, T. (2017). Transcriptomics technologies. *PLoS Computational Biology*, 13(5), e1005457.
- Mackelprang, R., Waldrop, M. P., DeAngelis, K. M., David, M. M., Chavarria, K. L., Blazewicz, S. J., ... & Jansson, J. K. (2011). Metagenomic analysis of a permafrost microbial community reveals a rapid response to thaw. *Nature*, 480(7377), 368-371.
- Maddison, W. P. and D.R. Maddison. (2017). Mesquite: a modular system for evolutionary analysis. Version 3.2 [Online]. Available at: <http://mesquiteproject.org> [Accessed 08-10-17].
- Malard, L. A., & Pearce, D. A. (2018). Microbial diversity and biogeography in Arctic soils. *Environmental Microbiology Reports*, 10(6), 611-625.
- Mardis, E. R. (2008). The impact of next-generation sequencing technology on genetics. *Trends in Genetics*, 24(3), 133-141.
- Margesin, R., Gander, S., Zacke, G., Gounot, A. M., & Schinner, F. (2003). Hydrocarbon degradation and enzyme activities of cold-adapted bacteria and yeasts. *Extremophiles*, 7(6), 451-458.
- Marshall, W. A. & Chalmers, M. O. (1997) Airborne dispersal of Antarctic algae and cyanobacteria. *Ecography*, 20, 585-594
- Marx, V. (2017). Microbiology: the return of culture. *Nature Methods*, 14(1), 37-41.



- Matsunaga, Okamura, Fukuda, Wahyudi, Murase, and Takeyama (2005). Complete Genome Sequence of the Facultative Anaerobic Magnetotactic Bacterium *Magnetospirillum* sp. strain AMB-1. *DNA Research*, 12, 157–166.
- Matsuura, N., Tourlousse, D. M., Sun, L., Toyonaga, M., Kuroda, K., Ohashi, A., ... & Sekiguchi, Y. (2015). Draft genome sequence of *Anaerolineae* strain TC1, a novel isolate from a methanogenic wastewater treatment system. *Genome Announcements*, 3(5), e01104-15.
- Matthews, J. A., & Vater, A. E. (2015). Pioneer zone geo-ecological change: Observations from a chronosequence on the Storbreen glacier foreland, Jotunheimen, southern Norway. *CATENA*, 135, 219-230.
- Mavromatis, K., Ivanova, N., Barry, K., Shapiro, H., Goltsman, E., McHardy, A. C., ... & Lapidus, A. (2007). Use of simulated data sets to evaluate the fidelity of metagenomic processing methods. *Nature Methods*, 4(6), 495-500.
- Mayer, L. M., Keil, R. G., Macko, S. A., Joye, S. B., Ruttenger, K. C., & Aller, R. C. (1998). Importance of suspended particulates in riverine delivery of bioavailable nitrogen to coastal zones. *Global Biogeochemical Cycles*, 12(4), 573-579.
- Meerhoff, E., Castro, L. & Tapia, F. (2013). Influence of freshwater discharges and tides on the abundance and distribution of larval and juvenile *Munida gregaria* in the Baker river estuary, Chilean Patagonia. *Continental Shelf Research*, 61–62(0), 1-11.
- Mende, D. R., Waller, A. S., Sunagawa, S., Järvelin, A. I., Chan, M. M., Arumugam, M., ... & Bork, P. (2012). Assessment of metagenomic assembly using simulated next generation sequencing data. *PLoS one*, 7(2), e31386.
- Menge, D. N., & Hedin, L. O. (2009). Nitrogen fixation in different biogeochemical niches along a 120 000-year chronosequence in New Zealand. *Ecology*, 90(8), 2190-2201.
- Menzel, P., Ng K., and Krogh, A. (2016). Fast and sensitive taxonomic classification for metagenomics with Kaiju. *Nature Communications*, 7, 11257.
- Meola, M., Lazzaro, A., & Zeyer, J. (2014). Diversity, resistance and resilience of the bacterial communities at two alpine glacier forefields after a reciprocal soil transplantation. *Environmental Microbiology*, 16(6), 1918-1934.
- Methe, B. A., Nelson, K. E., Eisen, J. A., Paulsen, I. T., Nelson, W., Heidelberg, J. F., ... & Dodson, R. J. (2003). Genome of *Geobacter sulfurreducens*: metal reduction in subsurface environments. *Science*, 302(5652), 1967-1969.
- Metzker, M. L. (2010). Sequencing technologies—the next generation. *Nature Reviews Genetics*, 11(1), 31-46.
- Meyer, F., Paarmann, D., D'Souza, M., Olson, R., Glass, E. M., Kubal, M., ... & Wilkening, J. (2008). The metagenomics RAST server—a public resource for the automatic phylogenetic and functional analysis of metagenomes. *BMC Bioinformatics*, 9(1), 386.
- Miller, J., Koren, S., Sutton, G. (2010). Assembly algorithms for next-generation sequencing data. *Genomics*, 95: 315–327

- Mindl, B., Anesio, A. M., Meirer, K., Hodson, A. J., Laybourn-Parry, J., Sommaruga, R., & Sattler, B. (2007). Factors influencing bacterial dynamics along a transect from supraglacial runoff to proglacial lakes of a high Arctic glacier. *FEMS Microbiology Ecology*, 59(2), 307-317.
- Miniaci, C., Bunge, M., Duc, L., Edwards, I., Bürgmann, H., and Zeyer, J. (2007). Effects of pioneering plants on microbial structures and functions in a glacier forefield. *Biology and Fertility of Soils*, 44(2), 289-297.
- Montero, P., Pérez-Santos, I., Daneri, G., Gutiérrez, M. H., Igor, G., Seguel, R., ... & Crawford, D. W. (2017). A winter dinoflagellate bloom drives high rates of primary production in a Patagonian fjord ecosystem. *Estuarine, Coastal and Shelf Science*, 199, 105-116.
- Moran, M. A., Satinsky, B., Gifford, S. M., Luo, H., Rivers, A., Chan, L. K., ... & Smith, C. B. (2013). Sizing up metatranscriptomics. *The ISME journal*, 7(2), 237.
- Moreau, J. W., Zierenberg, R. A., & Banfield, J. F. (2010). Diversity of dissimilatory sulfite reductase genes (*dsrAB*) in a salt marsh impacted by long-term acid mine drainage. *Applied and Environmental Microbiology*, 76(14), 4819-4828.
- Moreau, M., Mercier, D., Laffly, D., & Roussel, E. (2008). Impacts of recent paraglacial dynamics on plant colonization: a case study on Midtre Lovénbreen foreland, Spitsbergen (79° N). *Geomorphology*, 95(1), 48-60.
- Mueller, D. R., Vincent, W. F., Pollard, W. H., & Fritsen, C. H. (2001). Glacial cryoconite ecosystems: a bipolar comparison of algal communities and habitats. *Nova Hedwigia Beiheft*, 123, 173-198.
- Müller, A. L., Kjeldsen, K. U., Rattei, T., Pester, M., & Loy, A. (2015). Phylogenetic and environmental diversity of *DsrAB*-type dissimilatory (bi) sulfite reductases. *The ISME journal*, 9(5), 1152.
- Mullikin, J. C., & Ning, Z. (2003). The phusion assembler. *Genome Research*, 13(1), 81-90.
- Murdock, S. A., & Juniper, S. K. (2017). Capturing compositional variation in denitrifying communities: A multiple primer approach that includes autotrophic denitrifying Epsilonproteobacteria. *Applied and Environmental Microbiology*, AEM-02753.
- Myers, E. W., Sutton, G. G., Delcher, A. L., Dew, I. M., Fasulo, D. P., Flanigan, M. J., ... & Anson, E. L. (2000). A whole-genome assembly of *Drosophila*. *Science*, 287(5461), 2196-2204.
- Mylona, P., Pawlowski, K., & Bisseling, T. (1995). Symbiotic nitrogen fixation. *The Plant Cell*, 7(7), 869.
- Nagarajan, N., & Pop, M. (2013). Sequence assembly demystified. *Nature Reviews Genetics*, 14(3), 157-167.
- Narasingarao, P., Podell, S., Ugalde, J. A., Brochier-Armanet, C., Emerson, J. B., Brocks, J. J., ... & Allen, E. E. (2012). De novo metagenomic assembly reveals abundant novel major lineage of Archaea in hypersaline microbial communities. *The ISME journal*, 6(1), 81.
- Narzisi, G and Mishra, B. (2011) Comparing De Novo genome assembly: the long and short of it. *PLoS ONE*, 6.
- Nash, M. V., Anesio, A. M., Barker, G., Tranter, M., Varliero, G., Eloe-Fadrosh, E. A., ... & Sánchez-Baracaldo, P. (2018). Metagenomic insights into diazotrophic communities across Arctic glacier forefields. *FEMS Microbiology Ecology*, 94(9), fyy114.

- Nemergut, D. R., Anderson, S. P., Cleveland, C. C., Martin, A. P., Miller, A. E., Seimon, A., & Schmidt, S. K. (2007). Microbial community succession in an un-vegetated, recently de-glaciated soil. *Microbial Ecology*, 53(1), 110-122.
- Newton, C. R., Graham, A., & Ellison, J. S. (1997). *PCR*. BIOS Scientific Publishers: Chicago
- Nicholson, W. L., Munakata, N., Horneck, G., Melosh, H. J., & Setlow, P. (2000). Resistance of *Bacillus* endospores to extreme terrestrial and extraterrestrial environments. *Microbiology and Molecular Biology Reviews*, 64(3), 548-572.
- Nicol, G. W., Tscherko, D., Embley, T. M., and Prosser, J. I. (2005). Primary succession of soil Crenarchaeota across a receding glacier foreland. *Environmental Microbiology*, 7(3), 337-347.
- Nielsen, H. B., Almeida, M., Juncker, A. S., Rasmussen, S., Li, J., Sunagawa, S., ... & Pelletier, E. (2014). Identification and assembly of genomes and genetic elements in complex metagenomic samples without using reference genomes. *Nature Biotechnology*, 32(8).
- Nübel, U., Garcia-Pichel, F., & Muyzer, G. (1997). PCR primers to amplify 16S rRNA genes from cyanobacteria. *Applied and Environmental Microbiology*, 63(8), 3327-3332.
- Nurk, S., Bankevich, A., Antipov, D., Gurevich, A. A., Korobeynikov, A., Lapidus, A., ... & Stepanauskas, R. (2013). Assembling single-cell genomes and mini-metagenomes from chimeric MDA products. *Journal of Computational Biology*, 20(10), 714-737.
- Nurk, S., Meleshko, D., Korobeynikov, A., & Pevzner, P. A. (2017). metaSPAdes: a new versatile metagenomic assembler. *Genome Research*, 27(5), 824-834.
- Oda, Y., Samanta, S. K., Rey, F. E., Wu, L., Liu, X., Yan, T., ... and Harwood, C. S. (2005). Functional genomic analysis of three nitrogenase isozymes in the photosynthetic bacterium *Rhodospseudomonas palustris*. *Journal of Bacteriology*, 187(22), 7784-7794.
- Ohtonen, R., Fritze, H., Pennanen, T., Jumpponen, A., & Trappe, J. (1999). Ecosystem properties and microbial community changes in primary succession on a glacier forefront. *Oecologia*, 119(2), 239-246.
- Ondov, B. D., Bergman, N. H., & Phillippy, A. M. (2015). 'Krona: Interactive Metagenomic Visualization in a Web Browser'. In Nelson, K.E. (Eds). *Encyclopedia of Metagenomics: Genes, Genomes and Metagenomes: Basics, Methods, Databases and Tools*, 339-346.
- Oschmann, W. (2001). 'Oxygen in the ocean'. In D. Briggs., P. Crowther. (Eds.) *Palaeobiology II*. Wiley-Blackwell: Oxford, 470-472.
- Overbeek, R., Olson, R., Pusch, G. D., Olsen, G. J., Davis, J. J., Disz, T., ... & Vonstein, V. (2013). The SEED and the Rapid Annotation of microbial genomes using Subsystems Technology (RAST). *Nucleic Acids Research*, 42(D1), D206-D214.
- Ozsolak, F. (2012). Third-generation sequencing techniques and applications to drug discovery. *Expert Opinion on Drug Discovery*, 7(3), 231-243.
- Palma, S., & Silva, N. (2004). Distribution of siphonophores, chaetognaths, euphausiids and oceanographic conditions in the fjords and channels of southern Chile. *Deep Sea Research Part II: Topical Studies in Oceanography*, 51(6), 513-535.

- Park, S. J., Kim, J. G., Jung, M. Y., Kim, S. J., Cha, I. T., Ghai, R., ... & Rhee, S. K. (2012). Draft genome sequence of an ammonia-oxidizing archaeon, "Candidatus Nitrosopumilus sediminis" AR2, from Svalbard in the Arctic Circle. *Journal of Bacteriology*, 194(24), 6948-6949.
- Pattanaik, B., Schumann, R., and Karsten, U. (2007). Effects of ultraviolet radiation on cyanobacteria and their protective mechanisms. *Algae and Cyanobacteria in Extreme Environments*, 29-45.
- Pearce, D. A., Newsham, K. K., Thorne, M. A., Calvo-Bado, L., Krsek, M., Laskaris, P., ... & Wellington, E. M. (2012). Metagenomic analysis of a southern maritime Antarctic soil. *Frontiers in Microbiology*, 3.
- Peduzzi, S., Welsh, A., Demarta, A., Decristophoris, P., Peduzzi, R., Hahn, D., & Tonolla, M. (2011). Thiocystis chemoclinalis sp. nov. and Thiocystis cadagnonensis sp. nov., motile purple sulfur bacteria isolated from the chemocline of a meromictic lake. *International Journal of Systematic and Evolutionary Microbiology*, 61(7), 1682-1687.
- Peng, Y., Leung, H. C., Yiu, S. M., & Chin, F. Y. (2012). IDBA-UD: a de novo assembler for single-cell and metagenomic sequencing data with highly uneven depth. *Bioinformatics*, 28(11), 1420-1428.
- Pevzner, P. A., Tang, H., & Waterman, M. S. (2001). An Eulerian path approach to DNA fragment assembly. *Proceedings of the National Academy of Sciences*, 98(17), 9748-9753.
- Pickard, G. L. (1971). Some physical oceanographic features of inlets of Chile. *Journal of the Fisheries Board of Canada*, 28(8), 1077-1106.
- Piganeau, G., & Moreau, H. (2007). Screening the Sargasso Sea metagenome for data to investigate genome evolution in *Ostreococcus* (Prasinophyceae, Chlorophyta). *Gene*, 406(1).
- Pignatelli, M., & Moya, A. (2011). Evaluating the fidelity of de novo short read metagenomic assembly using simulated data. *PLoS One*, 6(5), e19984.
- Pop, M. (2009). Genome assembly reborn: recent computational challenges. *Briefings in Bioinformatics*, 10(4), 354-366.
- Prado-Fiedler, R. (2009). Winter and summer distribution of dissolved oxygen, pH and nutrients at the heads of fjords in Chilean Patagonia with possible phosphorus limitation. *Revista de Biología Marina y Oceanografía*, 44(3), 783-789.
- Prakash, T., & Taylor, T. D. (2012). Functional assignment of metagenomic data: challenges and applications. *Briefings in Bioinformatics*, 13(6), 711-727
- Prietzl, J., Wu, Y., Dümig, A., Zhou, J., & Klysubun, W. (2013). Soil sulfur speciation in two glacier forefield soil chronosequences assessed by SK-edge XANES spectroscopy. *European Journal of Soil Science*, 64(2), 260-272.
- Qaigen Bioinformatics (2016) CLC Assembly Cell. [Online]. Available from: <http://resources.qaigenbioinformatics.com/manuals/clcassemblycell/current/index.php?manual=Introduction.html> [Accessed 25-01-2018].
- Qin, J., Li, R., Raes, J., Arumugam, M., Burgdorf, K. S., Manichanh, C., ... & Mende, D. R. (2010). A human gut microbial gene catalogue established by metagenomic sequencing. *Nature*, 464(7285), 59.

- Qin, J., Li, Y., Cai, Z., Li, S., Zhu, J., Zhang, F., ... & Peng, Y. (2012). A metagenome-wide association study of gut microbiota in type 2 diabetes. *Nature*, 490(7418), 55-60.
- Qin, W., Heal, K. R., Ramdasi, R., Kobelt, J. N., Martens-Habben, W., Bertagnolli, A. D., ... & Devol, A. H. (2017). *Nitrosopumilus maritimus* gen. nov., sp. nov., *Nitrosopumilus cobalaminigenes* sp. nov., *Nitrosopumilus oxycliniae* sp. nov., and *Nitrosopumilus ureiphilus* sp. nov., four marine ammonia-oxidizing archaea of the phylum Thaumarchaeota. *International Journal of Systematic and Evolutionary Microbiology*, 67(12), 5067-5079.
- Quail, M. A., Smith, M., Coupland, P., Otto, T. D., Harris, S. R., Connor, T. R., ... & Gu, Y. (2012). A tale of three next generation sequencing platforms: comparison of Ion Torrent, Pacific Biosciences and Illumina MiSeq sequencers. *BMC Genomics*, 13(1), 341.
- Quaiser, A., Ochsenreiter, T., Klenk, H. P., Kletzin, A., Treusch, A. H., Meurer, G., ... & Schleper, C. (2002). First insight into the genome of an uncultivated crenarchaeote from soil. *Environmental Microbiology*, 4(10), 603-611.
- Raiswell, R. (1984). Chemical models of solute acquisition in glacial meltwaters. *Journal of Glaciology*, 30(104), 49-57.
- Raymond, J., Siefert, J. L., Staples, C. R., and Blankenship, R. E. (2004). The natural history of nitrogen fixation. *Molecular Biology and Evolution*, 21(3), 541-554.
- Reddy, P. V. V., Rao, S. S. S. N., Pratibha, M. S., Sailaja, B., Kavya, B., Manorama, R. R., ... & Shivaji, S. (2009). Bacterial diversity and bioprospecting for cold-active enzymes from culturable bacteria associated with sediment from a melt water stream of Midtre Lovénbreen glacier, an Arctic glacier. *Research in Microbiology*, 160(8), 538-546.
- Reed, D. C., Algar, C. K., Huber, J. A., & Dick, G. J. (2014). Gene-centric approach to integrating environmental genomics and biogeochemical models. *Proceedings of the National Academy of Sciences*, 111(5), 1879-1884.
- Renner, M., Arimitsu, M. L., & Piatt, J. F. (2012). Structure of marine predator and prey communities along environmental gradients in a glaciated fjord. *Canadian Journal of Fisheries and Aquatic Sciences*, 69(12), 2029-2045.
- Riesenfeld, C. S., Schloss, P. D., & Handelsman, J. (2004). Metagenomics: genomic analysis of microbial communities. *Annual Reviews Genetics*, 38, 525-552.
- Rime, T., Hartmann, M., Brunner, I. (2015) Vertical distribution of the soil microbiota along a successional gradient in a glacier forefield. *Molecular Ecology*, 24, 1091–1108.
- Rinnan, R., Rousk, J., Yergeau, E., Kowalchuk, G. A., & Bååth, E. (2009). Temperature adaptation of soil bacterial communities along an Antarctic climate gradient: predicting responses to climate warming. *Global Change Biology*, 15(11), 2615-2625.
- Roberts, R. J., Carneiro, M. O., & Schatz, M. C. (2013). The advantages of SMRT sequencing. *Genome Biology*, 14(6), 405.
- Rodriguez, L. M., & Konstantinidis, K. T. (2014). Estimating coverage in metagenomic data sets and why it matters. *The ISME journal*, 8(11), 2349.
- Rognes, T., Flouri, T., Nichols, B., Quince, C., & Mahé, F. (2016). VSEARCH: a versatile open source tool for metagenomics. *PeerJ*, 4, e2584.

- Rondon, M. R., August, P. R., Bettermann, A. D., Brady, S. F., Grossman, T. H., Liles, M. R., ... & Tiong, C. L. (2000). Cloning the soil metagenome: a strategy for accessing the genetic and functional diversity of uncultured microorganisms. *Applied and Environmental Microbiology*, 66(6), 2541-2547.
- Rüger, H. J., Fritze, D., & Spröer, C. (2000). New psychrophilic and psychrotolerant *Bacillus marinus* strains from tropical and polar deep-sea sediments and emended description of the species. *International Journal of Systematic and Evolutionary Microbiology*, 50(3), 1305-1313.
- Sakata Bekku, Y., Nakatsubo, T., Kume, A., & Koizumi, H. (2004). Soil microbial biomass, respiration rate, and temperature dependence on a successional glacier foreland in Ny-Ålesund, Svalbard. *Arctic, Antarctic, and Alpine Research*, 36(4), 395-399.
- Salzberg, S. L., Phillippy, A. M., Zimin, A., Puiu, D., Magoc, T., Koren, S., ... & Marçais, G. (2012). GAGE: A critical evaluation of genome assemblies and assembly algorithms. *Genome Research*, 22(3), 557-567.
- Sanford, R. A., Cole, J. R., and Tiedje, J. M. (2002). Characterization and description of *Anaeromyxobacter dehalogenans* gen. nov., sp. nov., an aryl-halo-respiring facultative anaerobic myxobacterium. *Applied and Environmental Microbiology*, 68(2), 893-900.
- Sanger, F., Air, G. M., Barrell, B. G., Brown, N. L., Coulson, A. R., Fiddes, J. C., ... & Smith, M. (1977). Nucleotide sequence of bacteriophage  $\phi$ X174 DNA. *Nature*, 265(5596), 687.
- Sattin, S. R., Cleveland, C. C., Hood, E., Reed, S. C., King, A. J., Schmidt, S. K., ... & Nemergut, D. R. (2009). Functional shifts in unvegetated, perhumid, recently-deglaciated soils do not correlate with shifts in soil bacterial community composition. *The Journal of Microbiology*, 47(6), 673-681.
- Schadt, E. E., Turner, S., & Kasarskis, A. (2010). A window into third-generation sequencing. *Human Molecular Genetics*, 19(R2), R227-R240.
- Schink, B. (2006). 'The genus *Pelobacter*'. In: Dworkin, M., Falkow, S., Schleifer K.H., Stackebrandt, E. (Eds). *The Prokaryotes*. Springer: New York, 5-11.
- Schink, B., and Stieb, M. (1983). Fermentative degradation of polyethylene glycol by a strictly anaerobic, gram-negative, nonsporeforming bacterium, *Pelobacter venetianus* sp. nov. *Applied and Environmental Microbiology*, 45(6), 1905-1913.
- Schmalenberger, A., & Noll, M. (2009). Shifts in desulfonating bacterial communities along a soil chronosequence in the forefield of a receding glacier. *FEMS Microbiology Ecology*, 71(2), 208-217.
- Schmidt, S. K., Reed, S. C., Nemergut, D. R., Grandy, A. S., Cleveland, C. C., Weintraub, M. N., ... & Martin, A. M. (2008). The earliest stages of ecosystem succession in high-elevation (5000 metres above sea level), recently deglaciated soils. *Proceedings of the Royal Society of London B: Biological Sciences*, 275(1653), 2793-2802.
- Schmieder, R., & Edwards, R. (2011). Quality control and pre-processing of metagenomic datasets. *Bioinformatics*, 27(6), 863-864
- Schulz, S., Brankatschk, R., Dümig, A., Kögel-Knabner, I., Schloter, M., and Zeyer, J. (2013). The role of microorganisms at different stages of ecosystem development for soil formation. *Biogeosciences*, 10(6), 3983-3996.

- Schütte, U. M., Abdo, Z., Bent, S. J., Williams, C. J., Schneider, G. M., Solheim, B., & Forney, L. J. (2009). Bacterial succession in a glacier foreland of the High Arctic. *The ISME journal*, 3(11), 1258-1268.
- Schütte, U. M., Abdo, Z., Foster, J., Ravel, J., Bunge, J., Solheim, B., and Forney, L. J. (2010). Bacterial diversity in a glacier foreland of the high Arctic. *Molecular Ecology*, 19(s1), 54-66.
- Sczyrba, A., Hofmann, P., Belmann, P., Koslicki, D., Janssen, S., Dröge, J., ... & Bremges, A. (2017). Critical Assessment of Metagenome Interpretation—a benchmark of metagenomics software. *Nature Methods*, 14(11), 1063.
- Segata, N., Waldron, L., Ballarini, A., Narasimhan, V., Jousson, O., & Huttenhower, C. (2012). Metagenomic microbial community profiling using unique clade-specific marker genes. *Nature Methods*, 9(8), 811-814.
- Shapiro, E., Biezuner, T., & Linnarsson, S. (2013). Single-cell sequencing-based technologies will revolutionize whole-organism science. *Nature Reviews Genetics*, 14(9), 618.
- Sharon, I., & Banfield, J. F. (2013). Genomes from metagenomics. *Science*, 342(6162).
- Shelobolina, E. S., Vrionis, H. A., Findlay, R. H., and Lovley, D. R. (2008). *Geobacter uraniireducens* sp. nov., isolated from subsurface sediment undergoing uranium bioremediation. *International Journal of Systematic and Evolutionary Microbiology*, 58(5), 1075-1078.
- Shendure, J., & Ji, H. (2008). Next-generation DNA sequencing. *Nature Biotechnology*, 26(10), 1135-1145.
- Sigler, W. V., & Zeyer, J. (2002). Microbial diversity and activity along the forefields of two receding glaciers. *Microbial Ecology*, 43(4), 397-407.
- Silva, N. (2008). 'Physical and chemical characteristics of the surface sediments in the austral Chilean channels and fjords'. In Silva, N., Palma, S.(Eds.), *Progress in the oceanographic knowledge of Chilean interior waters, from Puerto Montt to Cape Horn. Comité Oceanográfico Nacional- Pontificia Universidad Católica de Valparaíso, Valparaíso*, 69-75.
- Silva, N., & Vargas, C. A. (2014). Hypoxia in Chilean patagonian fjords. *Progress in Oceanography*, 129, 62-74.
- Simon, C., Wiezer, A., Strittmatter, A. W., & Daniel, R. (2009). Phylogenetic diversity and metabolic potential revealed in a glacier ice metagenome. *Applied and Environmental Microbiology*, 75(23), 7519-7526.
- Simpson, J.T., Wong, K., Jackman, S.D., Schein, J.E., Jones, S. J., and Birol, I. (2006) AbySS: a parallel assembler for short read sequence data. *Genome Research*, 19(6), 1117-1123.
- Sims, D., Sudbery, I., Illott, N. E., Heger, A., & Ponting, C. P. (2014). Sequencing depth and coverage: key considerations in genomic analyses. *Nature Reviews Genetics*, 15(2), 121-132.
- Skidmore, M. L., Foght, J. M., & Sharp, M. J. (2000). Microbial life beneath a high Arctic glacier. *Applied and Environmental Microbiology*, 66(8), 3214-3220.
- Sorokin, D. Y., Lysenko, A. M., Mityushina, L. L., Tourova, T. P., Jones, B. E., Rainey, F. A., ... & Kuenen, G. J. (2001). *Thioalkalimicrobium aerophilum* gen. nov., sp. nov. and *Thioalkalimicrobium sibericum* sp. nov., and *Thioalkalivibrio versutus* gen. nov., sp. nov., *Thioalkalivibrio nitratiss* sp. nov., novel and *Thioalkalivibrio denitrificans* sp. nov., novel

obligately alkaliphilic and obligately chemolithoautotrophic sulfur-oxidizing bacteria from soda lakes. *International Journal of Systematic and Evolutionary Microbiology*, 51(2), 565-580.

- Stamatakis, A. (2014). RAxML version 8: a tool for phylogenetic analysis and post-analysis of large phylogenies. *Bioinformatics*, 30(9), 1312-1313.
- Statham, P. J., Skidmore, M., & Tranter, M. (2008). Inputs of glacially derived dissolved and colloidal iron to the coastal ocean and implications for primary productivity. *Global Biogeochemical Cycles*, 22(3).
- Stibal, M., Hasan, F., Wadham, J. L., Sharp, M. J., & Anesio, A. M. (2012). Prokaryotic diversity in sediments beneath two polar glaciers with contrasting organic carbon substrates. *Extremophiles*, 16(2), 255-265.
- Stibal, M., Šabacká, M., & Kaštovská, K. (2006). Microbial communities on glacier surfaces in Svalbard: impact of physical and chemical properties on abundance and structure of cyanobacteria and algae. *Microbial Ecology*, 52(4), 644-654.
- Stibal, M., Tranter, M., Benning, L. G., & Řehák, J. (2008). Microbial primary production on an Arctic glacier is insignificant in comparison with allochthonous organic carbon input. *Environmental Microbiology*, 10(8), 2172-2178.
- Strauss, S. L., Garcia-Pichel, F., & Day, T. A. (2012). Soil microbial carbon and nitrogen transformations at a glacial foreland on Anvers Island, Antarctic Peninsula. *Polar Biology*, 35(10), 1459-1471.
- Strauss, S. L., Ruhland, C. T., & Day, T. A. (2009). Trends in soil characteristics along a recently deglaciated foreland on Anvers Island, Antarctic Peninsula. *Polar Biology*, 32(12), 1779-1788.
- Streit, W. R., & Schmitz, R. A. (2004). Metagenomics—the key to the uncultured microbes. *Current Opinion in Microbiology*, 7(5), 492-498.
- Sutton, G. G., White, O., Adams, M. D., & Kerlavage, A. R. (1995). TIGR Assembler: A new tool for assembling large shotgun sequencing projects. *Genome Science and Technology*, 1(1), 9-19.
- Suzuki, K. I., Sasaki, J., Uramoto, M., Nakase, T., & Komagata, K. (1997). *Cryobacterium psychrophilum* gen. nov., sp. nov., nom. rev., comb. nov., an obligately psychrophilic actinomycete to accommodate “*Curtobacterium psychrophilum*” Inoue and Komagata 1976. *International Journal of Systematic and Evolutionary Microbiology*, 47(2), 474-478.
- Swerdlow, H., Wu, S.L., Harke, H. & Dovichi, N.J. (1990) Capillary gel electrophoresis for DNA sequencing. Laser-induced fluorescence detection with the sheath flow cuvette. *Journal of Chromatography*, 516, 61–67.
- Szynkiewicz, A., Modelska, M., Buczyński, S., Borrok, D. M., & Merrison, J. P. (2013). The polar sulfur cycle in the Werenskiöldbreen, Spitsbergen: Possible implications for understanding the deposition of sulfate minerals in the North Polar Region of Mars. *Geochimica et Cosmochimica Acta*, 106, 326-343.
- Taberlet, P., Coissac, E., Hajibabaei, M., & Rieseberg, L. H. (2012). Environmental DNA. *Molecular Ecology*, 21(8), 1789-1793.
- Takai, K., Campbell, B. J., Cary, S. C., Suzuki, M., Oida, H., Nunoura, T., ... & Horikoshi, K. (2005). Enzymatic and genetic characterization of carbon and energy metabolisms by deep-sea



hydrothermal chemolithoautotrophic isolates of Epsilonproteobacteria. *Applied and Environmental Microbiology*, 71(11), 7310-7320.

- Tamaru, Y., Takani, Y., Yoshida, T., & Sakamoto, T. (2005). Crucial role of extracellular polysaccharides in desiccation and freezing tolerance in the terrestrial cyanobacterium *Nostoc commune*. *Applied and Environmental Microbiology*, 71(11), 7327-7333.
- Taylor, J., & Parkes, R. J. (1983). The cellular fatty acids of the sulfate-reducing bacteria, *Desulfobacter* sp., *Desulfobulbus* sp. and *Desulfovibrio desulfuricans*. *Microbiology*, 129(11), 3303-3309.
- Teeling, H., & Glöckner, F. O. (2012). Current opportunities and challenges in microbial metagenome analysis—a bioinformatic perspective. *Briefings in Bioinformatics*, bbs039.
- Teixeira, R. L. F., Von Der Weid, I., Seldin, L., and Rosado, A. S. (2008). Differential expression of *nifH* and *anfH* genes in *Paenibacillus durus* analysed by reverse transcriptase-PCR and denaturing gradient gel electrophoresis. *Letters in Applied Microbiology*, 46(3), 344-349.
- Telling, J., Anesio, A. M., Hawkings, J., Tranter, M., Wadham, J. L., Hodson, A. J., ... & Yallop, M. L. (2010). Measuring rates of gross photosynthesis and net community production in cryoconite holes: a comparison of field methods. *Annals of Glaciology*, 51(56), 153-162.
- Telling, J., Anesio, A. M., Tranter, M., Irvine-Fynn, T., Hodson, A., Butler, C., and Wadham, J. (2011). Nitrogen fixation on Arctic glaciers, Svalbard. *Journal of Geophysical Research: Biogeosciences*, 116(G3).
- Telling, J., Stibal, M., Anesio, A. M., Tranter, M., Nias, I., Cook, J., ... & Nienow, P. (2012). Microbial nitrogen cycling on the Greenland Ice Sheet. *Biogeosciences*, 9(7), 2431-2442.
- Teske, A., Alm, E., Regan, J. M., Toze, S., Rittmann, B. E., & Stahl, D. A. (1994). Evolutionary relationships among ammonia- and nitrite-oxidizing bacteria. *Journal of Bacteriology*, 176(21), 6623-6630.
- Thomas, T., Gilbert, J., and Meyer, F. (2012). Metagenomics—a guide from sampling to data analysis. *Microbial Informatics and Experimentation*, 2(3), 1-12.
- Torsvik, V., & Øvreås, L. (2002). Microbial diversity and function in soil: from genes to ecosystems. *Current Opinion in Microbiology*, 5(3), 240-245.
- Tranter, M., Brown, G. H., Hodson, A., Gurnell, A. M., & Sharp, M. J. (1994). Variations in the nitrate concentration of glacial runoff in alpine and sub-polar environments. *Snow and Ice covers: Interactions with the Atmosphere and Ecosystems*, 299-311.
- Tranter, M., Sharp, M. J., Lamb, H. R., Brown, G. H., Hubbard, B. P., & Willis, I. C. (2002). Geochemical weathering at the bed of Haut Glacier d'Arolla, Switzerland—a new model. *Hydrological Processes*, 16(5), 959-993.
- Tréguer, P.J. and Rocha C.L. (2013). The World Ocean Silica Cycle. *Annual Review of Marine Science*, 5(1), 477-501.
- Treseder, K. K. (2008). Nitrogen additions and microbial biomass: A meta-analysis of ecosystem studies. *Ecology Letters*, 11(10), 1111-1120.

- Tringe, S. G., & Rubin, E. M. (2005). Metagenomics: DNA sequencing of environmental samples. *Nature Reviews Genetics*, 6(11), 805-814.
- Tringe, S. G., Von Mering, C., Kobayashi, A., Salamov, A. A., Chen, K., Chang, H. W., ... & Bork, P. (2005). Comparative metagenomics of microbial communities. *Science*, 308(5721).
- Tscherko, D., Hammesfahr, U., Zeltner, G., Kandeler, E., & Böcker, R. (2005). Plant succession and rhizosphere microbial communities in a recently deglaciated alpine terrain. *Basic and Applied Ecology*, 6(4), 367-383.
- Tscherko, D., Rustemeier, J., Richter, A., Wanek, W., & Kandeler, E. (2003). Functional diversity of the soil microflora in primary succession across two glacier forelands in the Central Alps. *European Journal of Soil Science*, 54(4), 685-696.
- Turpin-Jelfs, T., Michaelides, K., Blacker, J. J., Benning, L. G., Williams, J. M., & Anesio, A. M. (2019). Distribution of soil nitrogen and nitrogenase activity in the forefield of a High Arctic receding glacier. *Annals of Glaciology*, 1-8.
- Van Der Heijden, M. G., Bardgett, R. D., & Van Straalen, N. M. (2008). The unseen majority: soil microbes as drivers of plant diversity and productivity in terrestrial ecosystems. *Ecology Letters*, 11(3), 296-310.
- Van der Walt, A. J., Van Goethem, M. W., Ramond, J. B., Makhanyane, T. P., Reva, O., & Cowan, D. A. (2017). Assembling metagenomes, one community at a time. *bioRxiv*, 120154
- Vázquez-Castellanos, J. F., García-López, R., Pérez-Brocal, V., Pignatelli, M., & Moya, A. (2014). Comparison of different assembly and annotation tools on analysis of simulated viral metagenomic communities in the gut. *BMC Genomics*, 15(1), 37.
- Vezi, F., Cattonaro, F., & Policriti, A. (2011). e-RGA: enhanced reference guided assembly of complex genomes. *EMBnet. journal*, 17(1), pp-46.
- Vincent, J., Wilke, A., Bischof, J., Desai, N., D'Souza, M., Glass, E., ... & Meyer, F. (2013). Analysis of Shotgun Metagenomes with MG-RAST. *Journal of Biomolecular Techniques*, 24, S20.
- Vincent, W. F., J. A. E. Gibson, R. Pienitz, and V. Villeneuve (2000). Ice shelf microbial ecosystems in the high Arctic and implications for life on Snowball Earth. *Naturwissenschaften* 87: 137–141.
- Vishnivetskaya, T. A., Layton, A. C., Lau, M. C., Chauhan, A., Cheng, K. R., Meyers, A. J., ... & Pfiffner, S. M. (2014). Commercial DNA extraction kits impact observed microbial community composition in permafrost samples. *FEMS Microbiology Ecology*, 87(1), 217-230.
- Vollmers, J., Wiegand, S., & Kaster, A. K. (2017). Comparing and evaluating metagenome assembly tools from a microbiologist's perspective-Not only size matters!. *PloS one*, 12(1).
- Vos, M., Quince, C., Pijl, A. S., de Hollander, M., & Kowalchuk, G. A. (2012). A comparison of rpoB and 16S rRNA as markers in pyrosequencing studies of bacterial diversity. *PLoS One*, 7(2), e30600.
- Wadham, J. L., Bottrell, S., Tranter, M., & Raiswell, R. (2004). Stable isotope evidence for microbial sulfate reduction at the bed of a polythermal high Arctic glacier. *Earth and Planetary Science Letters*, 219(3-4), 341-355.

- Wadham, J. L., Tranter, M., Tulaczyk, S., & Sharp, M. (2008). Subglacial methanogenesis: a potential climatic amplifier?. *Global Biogeochemical Cycles*, 22(2).
- Wadham, J. L., Arndt, S., Tulaczyk, S., Stibal, M., Tranter, M., Telling, J., ... & Sharp, M. J. (2012). *Potential methane reservoirs beneath Antarctica*. *Nature*, 488(7413), 633-637.
- Wadham, J. L., De'ath, R., Monteiro, F. M., Tranter, M., Ridgwell, A., Raiswell, R., & Tulaczyk, S. (2013). The potential role of the Antarctic Ice Sheet in global biogeochemical cycles. *Earth and Environmental Science Transactions of the Royal Society of Edinburgh*, 104(01), 55-67.
- Wadham, J. L., Tranter, M., Skidmore, M., Hodson, A. J., Priscu, J., Lyons, W. B., ... & Jackson, M. (2010). Biogeochemical weathering under ice: size matters. *Global Biogeochemical Cycles*, 24(3).
- Wagner, M., Loy, A., Klein, M., Lee, N., Ramsing, N. B., Stahl, D. A., & Friedrich, M. W. (2005). Functional marker genes for identification of sulfate-reducing prokaryotes. *Methods in Enzymology*, 397, 469-489.
- Wainwright, M. (1978). Microbial sulfur oxidation in soil. *Science Progress*, 459-475.
- Wall, L. G. (2000). The actinorhizal symbiosis. *Journal of plant growth regulation*, 19(2), 167-182.
- Wang, Z., Gerstein, M., & Snyder, M. (2009). RNA-Seq: a revolutionary tool for transcriptomics. *Nature Reviews Genetics*, 10(1), 57.
- Ward, D. M., Weller, R., & Bateson, M. M. (1990). 16S rRNA sequences reveal numerous uncultured microorganisms in a natural community. *Nature*, 345(6270), 63.
- Ward, N. L., Challacombe, J. F., Janssen, P. H., Henrissat, B., Coutinho, P. M., Wu, M., ... & Barabote, R. D. (2009). Three genomes from the phylum Acidobacteria provide insight into the lifestyles of these microorganisms in soils. *Applied and Environmental Microbiology*, 75(7), 2046-2056.
- Warren, R.L., Sutton, G.G., Jones, S.J.M., Holt, R.A. (2007). Assembling millions of short DNA sequences using SSAKE. *Bioinformatics*, (23)500.
- Wash, S., & Image, C. (2008). DNA sequencing: generation next-next. *Nature Methods*, 5(3), 267.
- Weiner, R., Langille, S., and Quintero, E. (1995). Structure, function and immunochemistry of bacterial exopolysaccharides. *Journal of Industrial Microbiology*, 15, 339-346.
- Weisburg, W. G., Barns, S. M., Pelletier, D. A., & Lane, D. J. (1991). 16S ribosomal DNA amplification for phylogenetic study. *Journal of Bacteriology*, 173(2), 697-703.
- White, D. C., Sutton, S. D., & Ringelberg, D. B. (1996). The genus *Sphingomonas*: physiology and ecology. *Current Opinion in Biotechnology*, 7(3), 301-306.
- Widdel, F and Hansen, T.A. (1992) 'The dissimilatory sulfate- and sulfur-reducing bacteria'. In: Balows A., Trüper H.G., Dworkin M. Harder W., Schleifer K.H. (Eds) *The Prokaryotes. A Handbook on the Biology of Bacteria: Ecophysiology, Isolation, Identification, Applications*. Springer: New York, 583-624.
- Wojcik, R., Donhauser, J., Frey, B., Holm, S., Holland, A., Anesio, A. M., ... & Benning, L. G. (2019). Linkages between geochemistry and microbiology in a proglacial terrain in the High Arctic. *Annals of Glaciology*, 1-16.

- Wolicka, D., Zdanowski, M. K., Żmuda-Baranowska, M. J., Poszytek, A., & Grzesiak, J. (2014). Sulfate reducing activity detected in soil samples from Antarctica, Ecology Glacier forefield, King George Island. *Polish Journal of Microbiology*, 63(4), 443-450.
- Wommack, K. E., Bhavsar, J., & Ravel, J. (2008). Metagenomics: read length matters. *Applied and Environmental Microbiology*, 74(5), 1453-1463.
- Wooley, J. C., Godzik, A., & Friedberg, I. (2010). A primer on metagenomics. *PLoS Computational Biology*, 6(2).
- Wu, Y. W., Simmons, B. A., & Singer, S. W. (2015). MaxBin 2.0: an automated binning algorithm to recover genomes from multiple metagenomic datasets. *Bioinformatics*, 32(4), 605-607.
- Xu, J. (2006). Invited review: microbial ecology in the age of genomics and metagenomics: concepts, tools, and recent advances. *Molecular Ecology*, 15(7), 1713-1731.
- Yang, L., Hu, Y., Liu, Y., Zhang, J., Ulstrup, J., & Molin, S. (2011). Distinct roles of extracellular polymeric substances in *Pseudomonas aeruginosa* biofilm development. *Environmental Microbiology*, 13(7), 1705-1717.
- Yde, J. C., Bárcena, T. G., & Finster, K. W. (2011). Subglacial and proglacial ecosystem responses to climate change. In *Climate Change-Geophysical Foundations and Ecological Effects*. IntechOpen.
- Yilmaz, S., Allgaier, M., & Hugenholtz, P. (2010). Multiple displacement amplification compromises quantitative analysis of metagenomes. *Nature Methods*, 7(12), 943.
- Zehr, J. P., Jenkins, B. D., Short, S. M., and Steward, G. F. (2003). Nitrogenase gene diversity and microbial community structure: a cross-system comparison. *Environmental Microbiology*, 5(7), 539-554.
- Zeikus, J. G. (1977). The biology of methanogenic bacteria. *Bacteriological Reviews*, 41(2), 514.
- Zemp, M., Frey, H., Gärtner-Roer, I., Nussbaumer, S. U., Hoelzle, M., Paul, F., ... & Bajracharya, S. (2015). Historically unprecedented global glacier decline in the early 21st century. *Journal of Glaciology*, 61(228), 745-762.
- Zerbino, D. R., & Birney, E. (2008). Velvet: algorithms for de novo short read assembly using de Bruijn graphs. *Genome Research*, 18(5), 821-829.
- Zhang, W., Chen, J., Yang, Y., Tang, Y., Shang, J., & Shen, B. (2011). A practical comparison of de novo genome assembly software tools for next-generation sequencing technologies. *PLoS one*, 6(3), e17915.
- Zhang, Z., Schwartz, S., Wagner, L., & Miller, W. (2000). A greedy algorithm for aligning DNA sequences. *Journal of Computational Biology*, 7(1-2), 203-214.
- Zumsteg, A., Bååth, E., Stierli, B., Zeyer, J., and Frey, B. (2013). Bacterial and fungal community responses to reciprocal soil transfer along a temperature and soil moisture gradient in a glacier forefield. *Soil Biology and Biochemistry*, 61, 121-132.

# Appendix 1

**A1 Table 1:** Full list of assemblies carried out in the assembler evaluation. The assembler, metagenome (A-D) and parameter set are identified.

Assembly number	Assembler	Metagenome	Parameter set
1	MIRA	A	Set 1
2	MIRA	A	Set 2
3	MIRA	A	Set 3
4	MIRA	A	Set 4
5	MIRA	A	Set 5
6	MIRA	B	Set 1
7	MIRA	B	Set 2
8	MIRA	B	Set 3
9	MIRA	B	Set 4
10	MIRA	B	Set 5
11	MIRA	C	Set 1
12	MIRA	C	Set 2
13	MIRA	C	Set 3
14	MIRA	C	Set 4
15	MIRA	C	Set 5
16	MIRA	D	Set 1
17	MIRA	D	Set 2
18	MIRA	D	Set 3
19	MIRA	D	Set 4
20	MIRA	D	Set 5
21	SSAKE	A	Set 1
22	SSAKE	A	Set 2
23	SSAKE	A	Set 3
24	SSAKE	A	Set 4
25	SSAKE	A	Set 5
26	SSAKE	B	Set 1
27	SSAKE	B	Set 2
28	SSAKE	B	Set 3
29	SSAKE	B	Set 4
30	SSAKE	B	Set 5
31	SSAKE	C	Set 1
32	SSAKE	C	Set 2
33	SSAKE	C	Set 3
34	SSAKE	C	Set 4
35	SSAKE	C	Set 5
36	SSAKE	D	Set 1
37	SSAKE	D	Set 2
38	SSAKE	D	Set 3
39	SSAKE	D	Set 4
40	SSAKE	D	Set 5
41	ABYSS	A	Set 1
42	ABYSS	A	Set 2
43	ABYSS	A	Set 3
44	ABYSS	A	Set 4
45	ABYSS	A	Set 5
46	ABYSS	B	Set 1
47	ABYSS	B	Set 2
48	ABYSS	B	Set 3
49	ABYSS	B	Set 4
50	ABYSS	B	Set 5
51	ABYSS	C	Set 1
52	ABYSS	C	Set 2
53	ABYSS	C	Set 3
54	ABYSS	C	Set 4
55	ABYSS	C	Set 5
56	ABYSS	D	Set 1

57	ABYSS	D	Set 2
58	ABYSS	D	Set 3
59	ABYSS	D	Set 4
60	ABYSS	D	Set 5
61	MetaSPAdes	A	Set 1
62	MetaSPAdes	A	Set 2
63	MetaSPAdes	A	Set 3
64	MetaSPAdes	A	Set 4
65	MetaSPAdes	A	Set 5
66	MetaSPAdes	B	Set 1
67	MetaSPAdes	B	Set 2
68	MetaSPAdes	B	Set 3
69	MetaSPAdes	B	Set 4
70	MetaSPAdes	B	Set 5
71	MetaSPAdes	C	Set 1
72	MetaSPAdes	C	Set 2
73	MetaSPAdes	C	Set 3
74	MetaSPAdes	C	Set 4
75	MetaSPAdes	C	Set 5
76	MetaSPAdes	D	Set 1
77	MetaSPAdes	D	Set 2
78	MetaSPAdes	D	Set 3
79	MetaSPAdes	D	Set 4
80	MetaSPAdes	D	Set 5
81	CLC	A	Set 1
82	CLC	A	Set 2
83	CLC	A	Set 3
84	CLC	A	Set 4
85	CLC	A	Set 5
86	CLC	B	Set 1
87	CLC	B	Set 2
88	CLC	B	Set 3
89	CLC	B	Set 4
90	CLC	B	Set 5
91	CLC	C	Set 1
92	CLC	C	Set 2
93	CLC	C	Set 3
94	CLC	C	Set 4
95	CLC	C	Set 5
96	CLC	D	Set 1
97	CLC	D	Set 2
98	CLC	D	Set 3
99	CLC	D	Set 4
100	CLC	D	Set 5

**A1 Table 2:** Parameter settings modified during assemblies, with description of the parameter function. All other settings (unlisted) were kept at the default values, which can be found in the individual assembler manuals and publications (references listed).

Assembler	Parameter details	Set 1	Set 2	Set 3	Set 4	Set 5
MetaSPAdes 3.7.0 (Nurk <i>et al.</i> , 2013)	-k : kmer length	-k : 21,33,55	-k : 41	-k : 61	-k : 71	-k : 51
SSAKE 3.8.4 (Warren <i>et al.</i> , 2007)	-w : minimum contig coverage depth	-w : 5	-w : 1	-w : 3	-w : 1	-w : 1
	-m Minimum number of overlapping bases required during contig assembly	-m : default (20)	-m : default (20)	-m : default (20)	-m : 16	-m : default (20)
	-t Number of contig bases to trim when other extension possibilities are depleted	-t : 0	-t : 0	-t : 0	-t : 0	-t : 1
ABYSS 1.9.0 (Simpson <i>et al.</i> , 2009)	-k : kmer length	-k : 64	-k : 40	-k : 64	-k : 70	-k : 64
	-c : remove contigs below N coverage threshold	-c : N/A	-c : N/A	-c : 2	-c : N/A	-c : N/A
	-b : pop bubbles less than N base pairs	-b : N/A	-b : N/A	-b : N/A	-b : N/A	-b : 192
MIRA 4.0 (Chavreux, B., 2014)	SOLEXA_SETTINGS : flag used to indicate illumine specific settings	SOLEXA_SETTINGS : off	SOLEXA_SETTINGS : on	SOLEXA_SETTINGS : on	SOLEXA_SETTINGS : on	SOLEXA_SETTINGS : on
	-AS:mrl : minimum read length	-AS:mrl= N/A	-AS:mrl= N/A	-AS:mrl= N/A	-AS:mrl= N/A	-AS:mrl= 50
	-AS:ardml : minimum contig length	-AS:ardml = N/A	-AS:ardml = N/A	-AS:ardml = 100	-AS:ardml = 150	-AS:ardml = 200
CLC 4.4.1 (Qaigen Bioinformatics, 2016)	-m : minimum output contig length	-m : default (200)	-m: 100	-m : default (200)	-m : default (200)	-m : default (200)
	-w : wordsize for the de Bruijn graph (12 – 64). Default sets the value based on the input dataset size.	-w: default	-w: default	-w: 20	-w: 60	-w: default
	-b : maximum bubble size for the de Bruijn graph	-b: default (50)	-b: default (50)	-b: default (50)	-b: default (50)	-b: 40

**A1 Table 3:** Summary results table for metagenome assemblies, covering assembly size, contiguity and completeness. The assembler results are split by parameter sets (1-5), the assembler used, and the metagenome assembled (A-D). Assembly contiguity is identified by maximum contig length, contig N50 and the number of contigs over 1000bp. The completeness of the assembly is shown through the percentage coverage (of raw reads mapped to the input metagenome). Finally the number of contigs details the assembly size.

Parameter Set 1

Assembler	CLC				ABYSS				MIRA				MetaSPAdes				SSAKE				
	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	
Number of contigs	402385	388383	374676	406221	2975	3260	3405	3405	4374	186625	205976	272144	400485	666045	645872	613234	613787	210	223	238	263
Number of contigs over 1000 bp	1366	3269	6115	10853	12	23	20	20	24	373	389	418	456	939	2285	4622	8162	10	14	16	17
Maximum contig length	5060	4585	7265	12412	1948	1930	1935	1935	2186	5387	4773	5552	5931	5299	4940	5386	10045	1887	1643	1726	1726
N50	330	337	363	410	166	163	158	158	151	203	210	217	218	597	624	662	682	374	498	428	570
% coverage of BWA mapping	29.71	31.02	35.26	45.72	0.65	0.70	0.69	0.69	0.85	10.30	12.24	16.81	24.67	42.08	43.03	46.69	57.31	0.20	0.24	0.23	0.28

Parameter Set 2

Assembler	CLC				ABYSS				MIRA				MetaSPAdes				SSAKE				
	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	
Number of contigs	1334018	1289769	1179804	1038353	1227327	1204906	1189639	1253133	186528	206026	272130	400483	202146	212900	251115	337496	326884	343188	391249	507457	
Number of contigs over 1000 bp	1369	3234	6144	10876	143	146	147	147	189	389	391	425	455	193	191	222	223	17	23	23	26
Maximum contig length	5060	4585	7265	12412	4641	4730	5090	4345	7430	4975	5861	5698	5507	4807	5248	8361	1887	1849	1726	2064	
N50	232	234	243	281	109	112	121	138	203	210	210	631	618	697	629	602	586	119	119	121	123
% coverage of BWA mapping	59.35	59.56	60.80	67.00	22.13	23.56	27.47	36.04	10.30	12.24	16.81	24.67	12.51	14.20	18.21	26.16	8.16	8.97	10.83	14.85	

Parameter Set 3

Assembler	CLC				ABYSS				MIRA				MetaSPAdes				SSAKE				
	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	
Number of contigs	287900	277612	270300	295654	2975	3260	3405	3405	4374	186600	205957	272087	400498	15810	20951	30709	48068	515	611	721	867
Number of contigs over 1000 bp	1335	2467	4943	8418	12	23	20	20	24	377	394	414	457	82	95	92	104	12	18	18	22
Maximum contig length	5056	4795	5325	14158	1948	1930	1935	1935	2186	5047	4578	5552	5412	3994	2489	3679	4085	1725	1643	1726	2064
N50	332	337	364	408	166	163	158	151	723	723	665	630	617	1030	1065	994	1098	268	292	234	238
% coverage of BWA mapping	19.82	20.38	23.65	31.11	0.65	0.70	0.69	0.69	0.85	10.30	12.23	16.81	24.67	1.86	2.35	3.16	4.70	0.26	0.31	0.29	0.37

Parameter Set 4

Assembler	CLC				ABYSS				MIRA				MetaSPAdes				SSAKE				
	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	
Number of contigs	594	607	628	658	731	696	722	722	920	186556	205960	272159	400478	1097	1115	1136	1276	351881	366959	417318	536263
Number of contigs over 1000 bp	20	31	22	29	5	8	2	2	6	373	393	416	471	26	33	28	38	21	29	28	32
Maximum contig length	1959	1953	2349	2194	1480	1478	1399	1472	1722	5250	4115	5552	5440	2030	2550	2018	3227	1725	1643	1999	1743
N50	329	337	344	370	182	194	188	188	188	721	668	630	617	945	1086	895	1227	122	123	124	126
% coverage of BWA mapping	0.49	0.52	0.50	0.61	0.36	0.39	0.35	0.47	0.47	10.30	12.24	16.81	24.67	0.58	0.61	0.60	0.73	8.89	9.75	11.82	16.13

Parameter Set 5

Assembler	CLC				ABYSS				MIRA				MetaSPAdes				SSAKE				
	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	A	B	C	D	
Number of contigs	402253	388322	374685	406257	2975	3260	3405	3405	4374	186553	206015	272083	400489	15697	20864	30636	47902	331417	346054	395637	512446
Number of contigs over 1000 bp	1353	3211	6081	10842	12	23	20	20	24	378	395	421	456	86	94	92	101	16	23	20	28
Maximum contig length	5060	4827	7265	12412	1948	1930	1935	1935	2186	5047	4587	5552	5076	3994	2489	3679	4085	1723	1641	1724	1741
N50	330	337	363	410	166	163	158	158	151	723	669	630	617	226	227	228	228	117	118	119	121
% coverage of BWA mapping	29.68	30.98	35.22	45.69	0.65	0.70	0.69	0.69	0.85	10.30	12.24	16.81	24.67	0.21	0.23	0.19	0.27	4.94	5.59	7.01	9.89



**A1 Table 4:** One way analysis of variance (ANOVA) results for best performing parameter sets. *F* value and *P* value significance for dependent variables, percentage coverage, number of contigs and maximum contig length are shown. For each dependent variable, the ANOVA was carried out between the five tested assemblers for the best performing parameter sets.

	% Coverage	Number of contigs	Maximum contig length
<b>F value</b>	59.32	92.67	4.72
<b>P value</b>	<0.0001	<0.0001	0.112

**A1 Table 5:** One way analysis of variance (ANOVA) results for each assembler, comparing the contiguity, completeness and assembly size across the four metagenomes tested. For each metagenome, the summary statistics used included: number of contigs; contigs over 1000bp; maximum contig length and % coverage. These were compared across the assembly results for metagenomes A-D, to identify significant differences stemming from metagenome complexity, rather than the assembler.

	CLC	ABYSS	MIRA	MetaSPAdes	SSAKE
<b>F value</b>	0.86	1.10	1.00	0.66	1.00
<b>P value</b>	0.49	0.41	0.42	0.59	0.43

## Appendix 2

**A2 Table 1:** Samples used in metagenomic sequencing across the four Arctic forefields, and the corresponding metadata. Samples were obtained in a transect across each forefield, moving away from the glacier terminus. For the Midtre Lovénbreen (MI), Russell (RI) and Storglaciären (St) sites, three parallel transects were conducted to obtain field replicates. This was not possible for the Rabots (Rb) site.

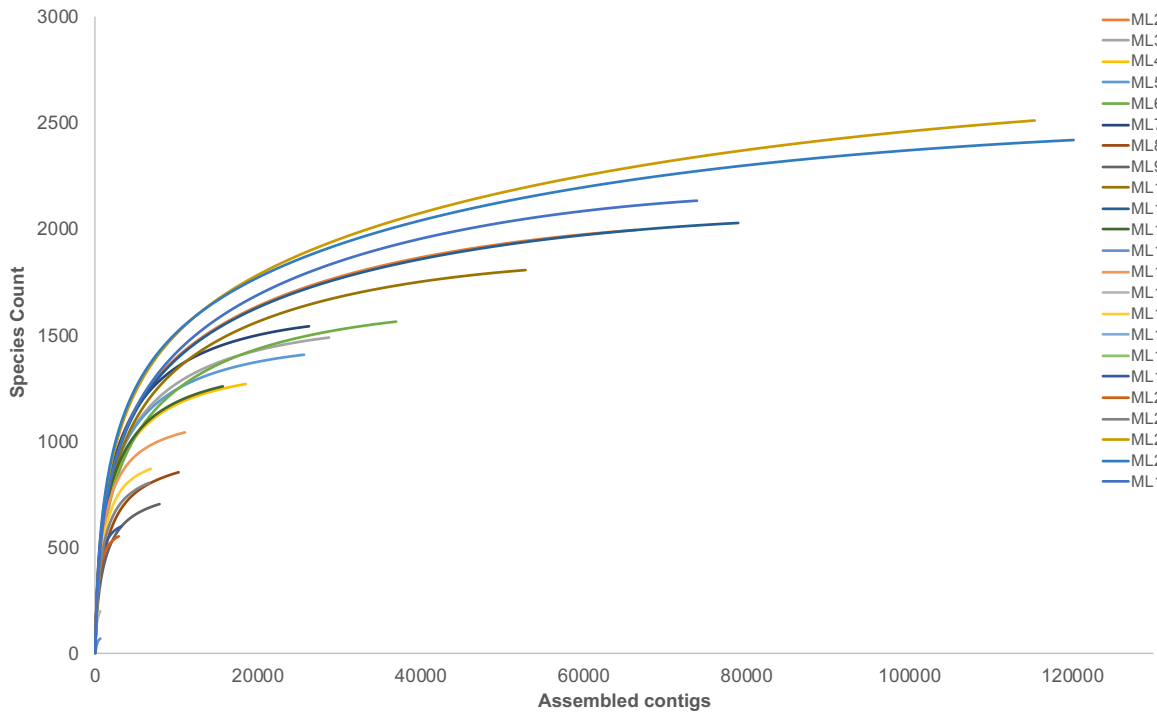
Site ID	Forefield	Latitude	Longitude	Altitude (m)	Date obtained
MI 1	MI	79.100555	12.156111	54	29/07/2013
MI 2	MI	79.112223	12.175555	44	29/07/2013
MI 3	MI	79.112222	12.258333	44	29/07/2013
MI 4	MI	79.118333	12.093611	54	29/07/2013
MI 5	MI	79.113611	12.195833	52	29/07/2013
MI 6	MI	79.104444	12.278888	52	29/07/2013
MI 7	MI	79.152555	12.215555	50	29/07/2013
MI 8	MI	79.151388	12.253611	43	29/07/2013
MI 9	MI	79.140833	12.092222	43	29/07/2013
MI 10	MI	78.927777	12.254166	35	29/07/2013
MI 11	MI	78.921666	12.076666	40	29/07/2013
MI 12	MI	78.907777	12.164444	48	29/07/2013
MI 13	MI	78.900555	12.076111	30	29/07/2013
MI 14	MI	78.900555	12.076111	40	29/07/2013
MI 15	MI	78.900555	12.076111	105	29/07/2013
MI 16	MI	78.900555	12.082777	29	29/07/2013
MI 17	MI	78.991666	12.233333	30	29/07/2013
MI 18	MI	78.978888	12.332222	30	29/07/2013
MI 19	MI	79.768333	12.143611	19	29/07/2013
MI 20	MI	79.768333	12.143611	19	29/07/2013
MI 21	MI	79.768333	12.143611	19	29/07/2013
MI 22	MI	79.484166	12.092222	105	29/07/2013
MI 23	MI	79.484166	12.092222	105	29/07/2013
RI 1	RI	67.15650902	-50.06398397	439	24/07/2014
RI 2	RI	67.15651598	-50.06386997	440	24/07/2014
RI 3	RI	67.15655998	-50.06389101	439	24/07/2014
RI 4	RI	67.16295303	-50.01826898	589	25/07/2014
RI 5	RI	67.16301103	-50.01844500	589	25/07/2014
RI 6	RI	67.16306903	-50.01828399	589	25/07/2014
RI 7	RI	67.15211598	-50.04869697	515	26/07/2014
RI 8	RI	67.15208103	-50.04859303	516	26/07/2014
RI 9	RI	67.15210701	-50.04851701	516	26/07/2014
RI 10	RI	67.15685402	-50.08261903	404	26/07/2014
RI 11	RI	67.15680499	-50.08249900	403	26/07/2014
RI 12	RI	67.15684304	-50.08236104	404	26/07/2014
RI 13	RI	67.15642001	-50.08365101	403	26/07/2014

RI 14	RI	67.15646099	-50.08366501	403	26/07/2014
RI 15	RI	67.15638698	-50.08376702	403	26/07/2014
RI 16	RI	67.15559204	-50.08486102	411	26/07/2014
RI 17	RI	67.15558601	-50.08499002	411	26/07/2014
RI 18	RI	67.15567200	-50.08488097	411	26/07/2014
RI 19	RI	67.08225802	-50.32251497	251	27/07/2014
RI 20	RI	67.08222499	-50.32243500	238	27/07/2014
RI 21	RI	67.08218602	-50.32229704	237	27/07/2014
RI 22	RI	67.05702002	-50.45979604	147	27/07/2014
RI 23	RI	67.05700200	-50.45969403	148	27/07/2014
RI 24	RI	67.05692296	-50.45960804	148	27/07/2014
Rb 1	Rb	67.910855	18.470863	1250	07/02/2014
Rb 2	Rb	67.907119	18.447522	1105	07/02/2014
Rb 3	Rb	67.907119	18.447522	1105	07/02/2014
Rb 4	Rb	67.906846	18.445550	1110	07/02/2014
Rb 5	Rb	67.872223	16.713705	1054	07/02/2014
St 1	St	67.904568	18.607115	1131	07/01/2014
St 2	St	67.904687	18.610965	1103	07/01/2014
St 3	St	67.904687	18.610965	1103	07/01/2014
St 4	St	67.904687	18.610965	1103	07/02/2014
St 5	St	67.899243	18.344347	1147	07/01/2014
St 6	St	67.899244	18.344371	1147	07/01/2014
St 7	St	67.900853	18.441750	1146	07/01/2014
St 8	St	67.900853	18.441750	1146	07/01/2014
St 9	St	67.900853	18.441750	1146	07/01/2014
St 10	St	67.900879	18.434740	1130	07/01/2014
St 11	St	67.900879	18.434740	1130	07/01/2014
St 12	St	67.900879	18.434740	1130	07/01/2014
St 13	St	67.901082	18.428257	1113	07/01/2014
St 14	St	67.901082	18.428257	1113	07/01/2014
St 15	St	67.865505	16.714941	1103	07/01/2014
St 16	St	67.865505	16.714941	1103	07/01/2014
St 17	St	67.865505	16.714941	1103	07/01/2014
St 18	St	67.903128	18.604355	1182	07/01/2014

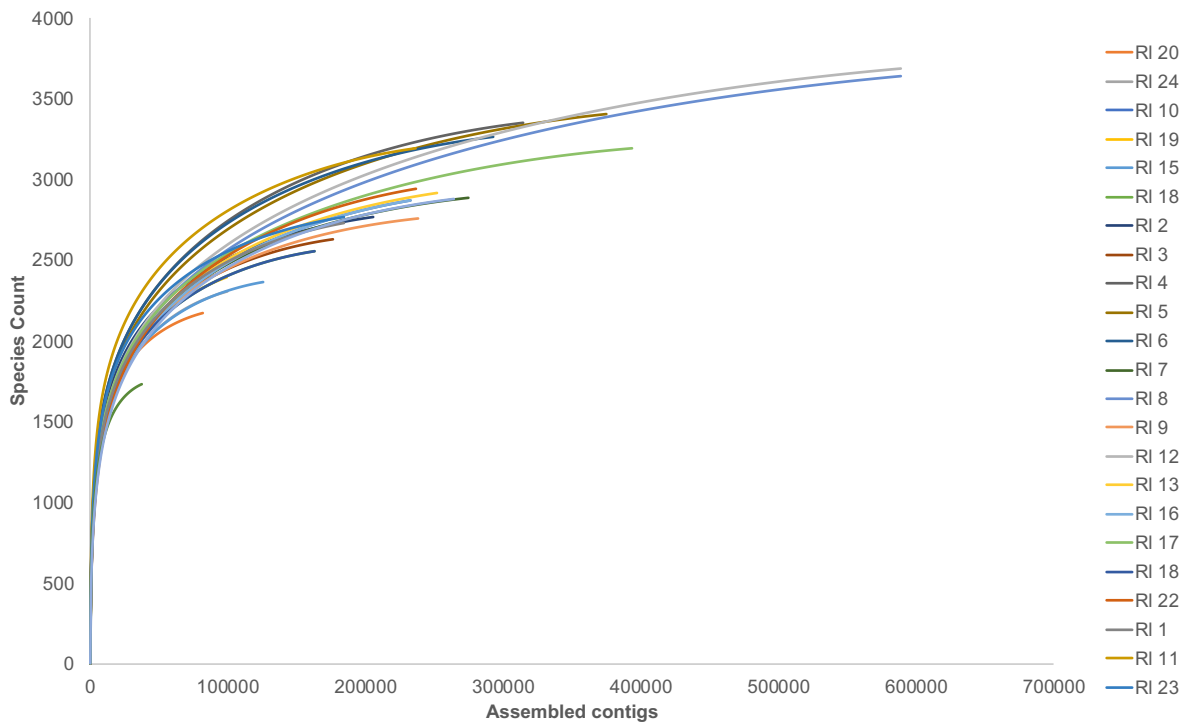
**A2 Table 2:** Output statistics for metagenome sequencing and assembly, for each site. The number of raw reads returned from sequencing is given, alongside the subsequent assembly sizes, in both sequences and bases.

	<b>Sequencing reads</b>	<b>Assembly size (sequences)</b>	<b>Assembly size (bases)</b>
MI 1	17465080	73894	33376241
MI 2	28186722	66450	32695412
MI 3	16801828	28727	17783002
MI 4	25805412	18484	6405001
MI 5	20630402	25643	10899365
MI 6	21421268	36952	16375348
MI 7	22655970	26238	13376228
MI 8	20003790	10239	3947663
MI 9	22650764	7883	3544627
MI 10	58387594	52836	25394805
MI 11	56518916	78995	29749304
MI 12	37925220	15664	4751816
MI 13	3952496	612	272071
MI 14	22179120	10999	8951989
MI 15	6433798	608	241660
MI 16	22231192	6823	2827176
MI 17	38684672	18250	16687388
MI 18	68392102	78312	52131008
MI 19	23405212	3192	998864
MI 20	20152212	2871	859271
MI 21	28777734	6620	2019562
MI 22	22182332	115393	61826702
MI 23	28234260	120239	81857371
RI 1	67748804	204281	181720264
RI 2	77474456	205426	192219830
RI 3	61199512	176255	115011779
RI 4	74992750	314429	325818814
RI 5	74220220	374929	341656927
RI 6	68161082	292741	302935645
RI 7	75975046	274719	186606224
RI 8	91844214	588635	395690349
RI 9	77023868	238198	157627061
RI 10	139381778	344195	180851046
RI 11	91962344	203511	112190691
RI 12	81681122	86956	82284058
RI 13	105431412	251779	123396624
RI 14	63240678	37335	22184640

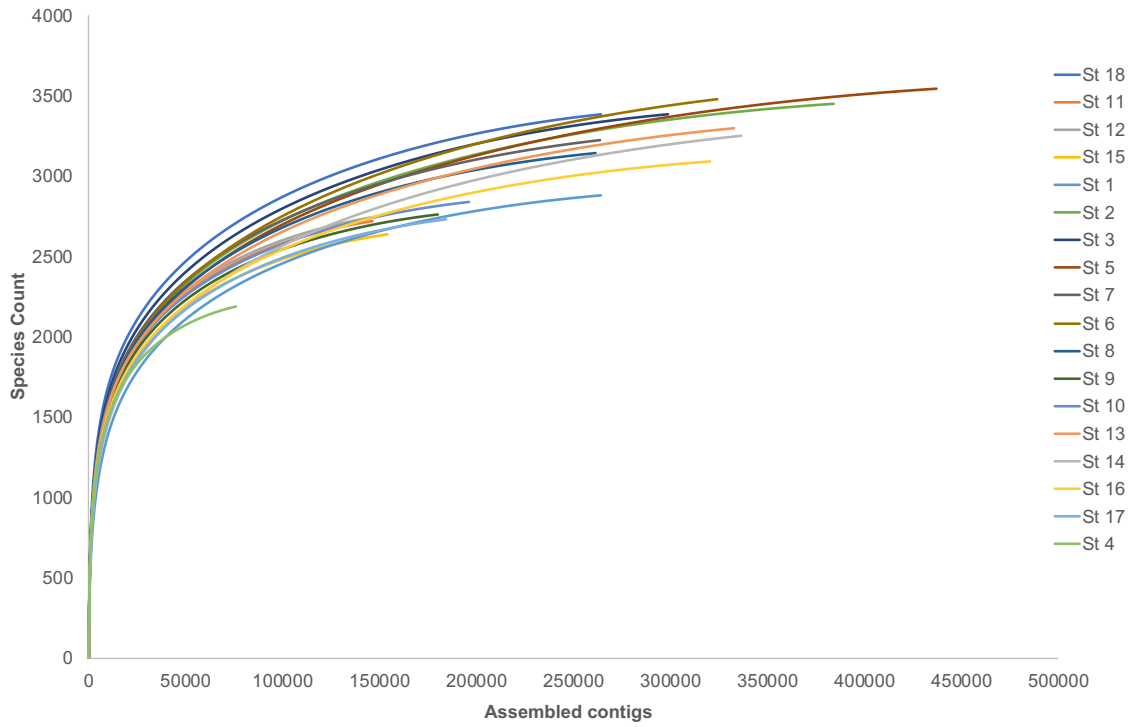
RI 15	72243790	125592	88672076
RI 16	82137112	232607	118301737
RI 17	107570706	393735	228570150
RI 18	78274268	162862	98312835
RI 19	75392190	99149	52942968
RI 20	65582044	81726	58507862
RI 21	91522052	264245	149668891
RI 22	96384726	236649	128282042
RI 23	96207238	406505	410018759
RI 24	80981820	99519	60773858
Rb 1	76498938	350951	429543524
Rb 2	72691820	275858	181694382
Rb 3	70126632	334545	249845086
Rb 4	61828928	303591	231705319
Rb 5	74667258	375682	213265070
St 1	68111048	336424	219250082
St 2	71828498	154126	83781202
St 3	85214054	320463	188269776
St 4	71411294	184331	103140952
St 5	66910678	323872	236040965
St 6	72683122	384288	290006175
St 7	60148730	298633	216231874
St 8	67844804	75868	48601571
St 9	63953088	437238	275297528
St 10	83010234	324248	214964286
St 11	74901072	263784	180223385
St 12	83405572	261266	175258662
St 13	64225756	179953	121520140
St 14	64764076	196033	140418783
St 15	66675200	146437	83396680
St 16	66848090	149134	96714442
St 17	72051286	332892	268415228
St 18	78397352	478826	377436650



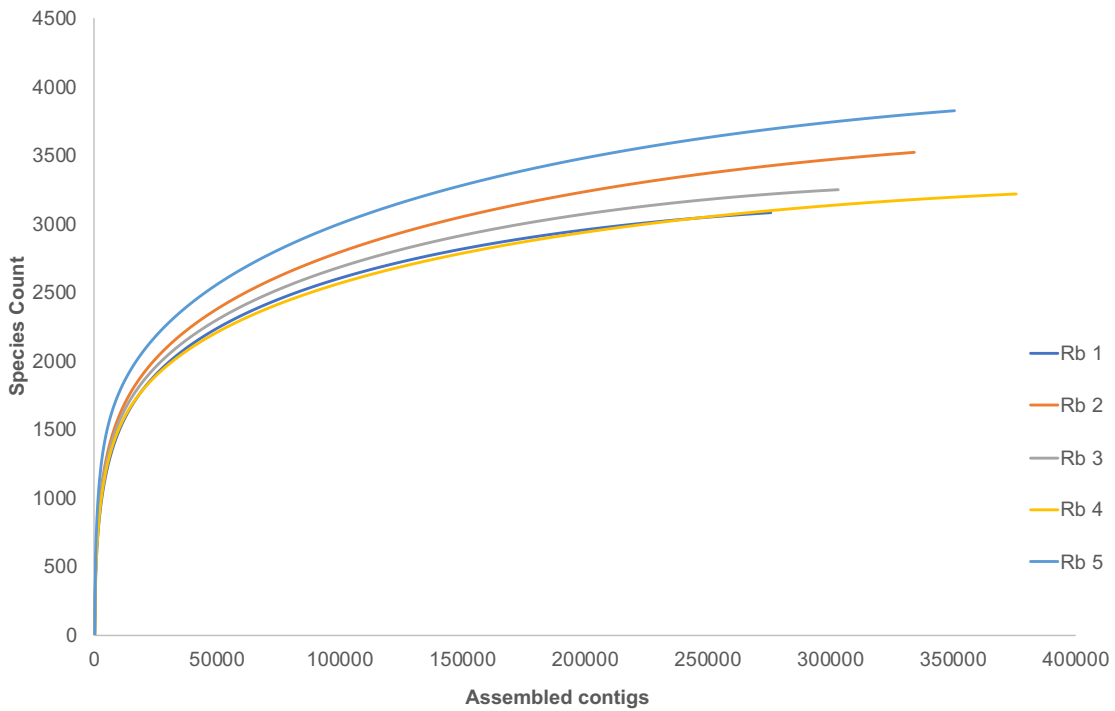
**A2 Figure 1:** Rarefaction curves for metagenomes sampled from Midtre Lovénbreen, Svalbard. The total assembled contigs in each metagenome is shown, against the total species count obtained from these contigs.



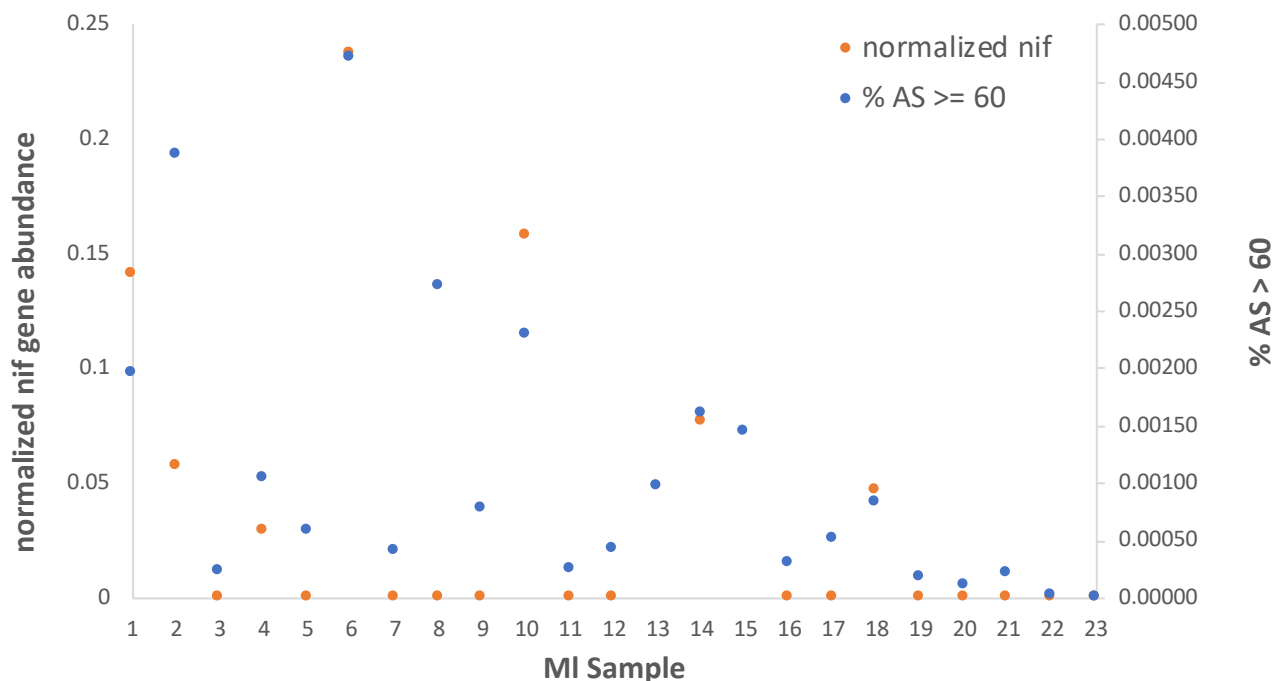
**A2 Figure 2:** Rarefaction curves for metagenomes sampled from Russell Glacier, Greenland. The total assembled contigs in each metagenome is shown, against the total species count obtained from these contigs.



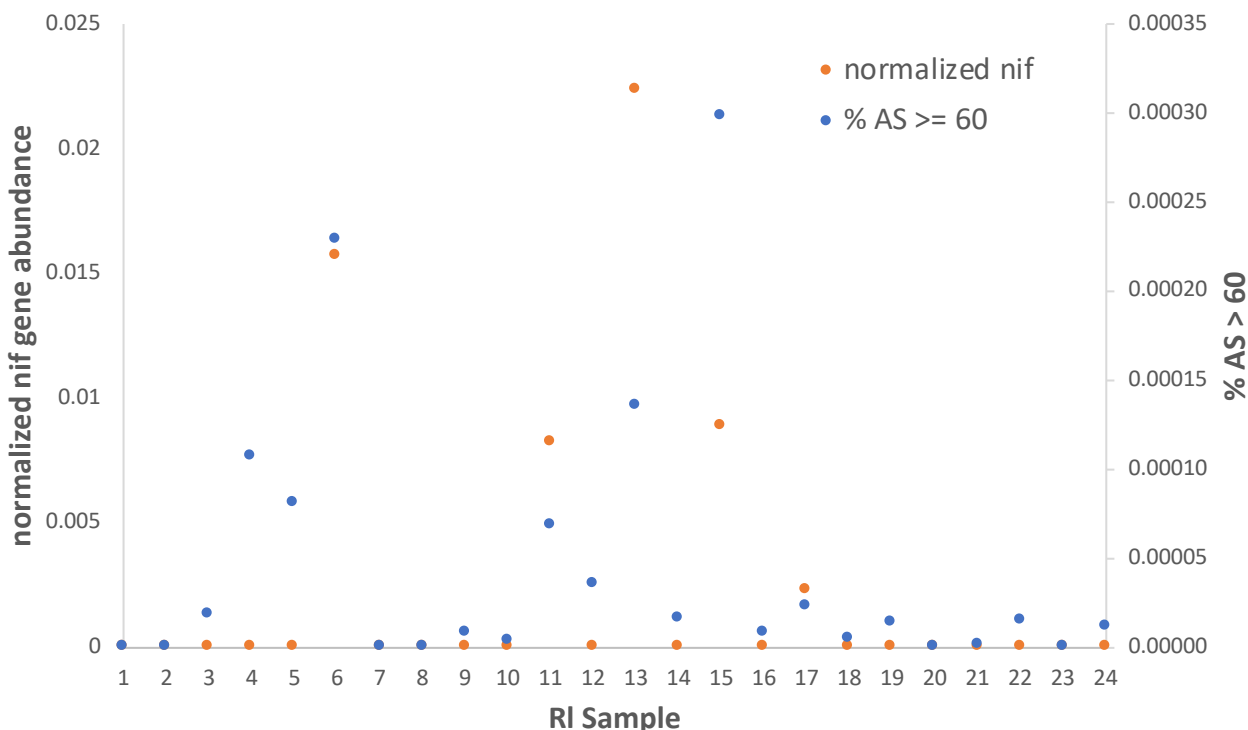
**A2 Figure 3:** Rarefaction curves for metagenomes sampled from Storglaciären, N-Sweden. The total assembled contigs in each metagenome is shown, against the total species count obtained from these contigs.



**A2 Figure 4:** Rarefaction curves for metagenomes sampled from Rabots glacier, N-Sweden. The total assembled contigs in each metagenome is shown, against the total species count obtained.

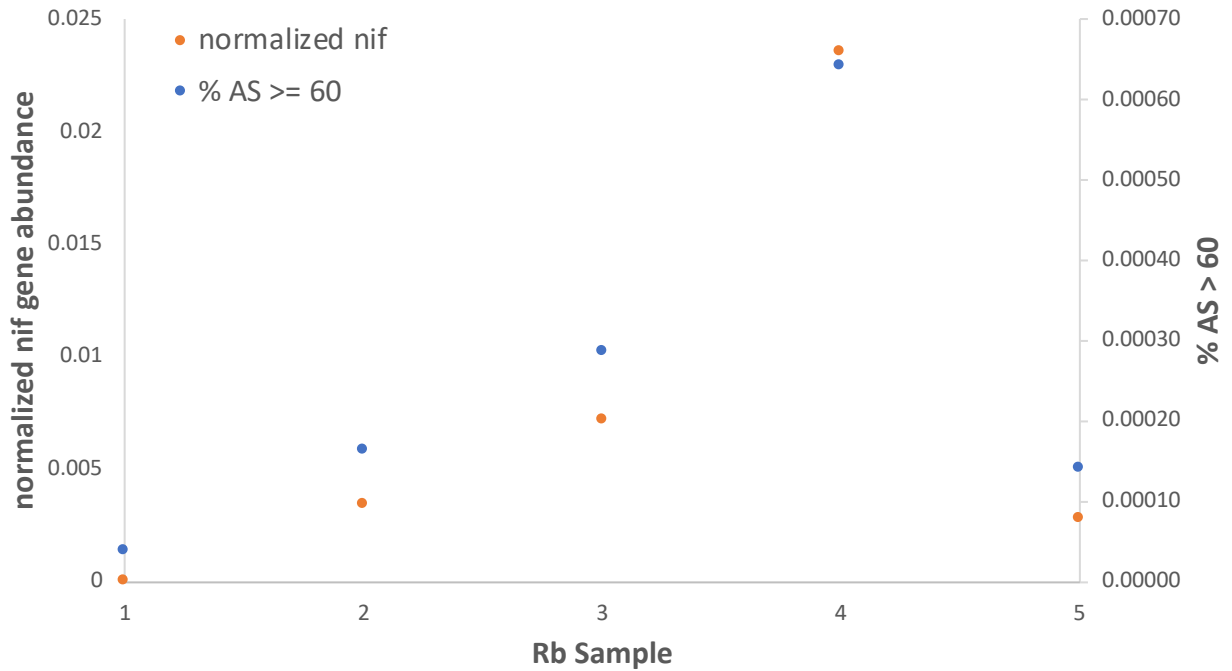


**A2 Figure 5:** *rpoB* normalized *nif* gene abundance and percentage of reads with an Alignment Score (AS) over 60, for samples obtained from Midtre Lovénbreen (MI), Svalbard. The Alignment Score ranges between 0 and the maximum length of the reads (0-100) and indicates the quality of the alignment between reads and contigs.

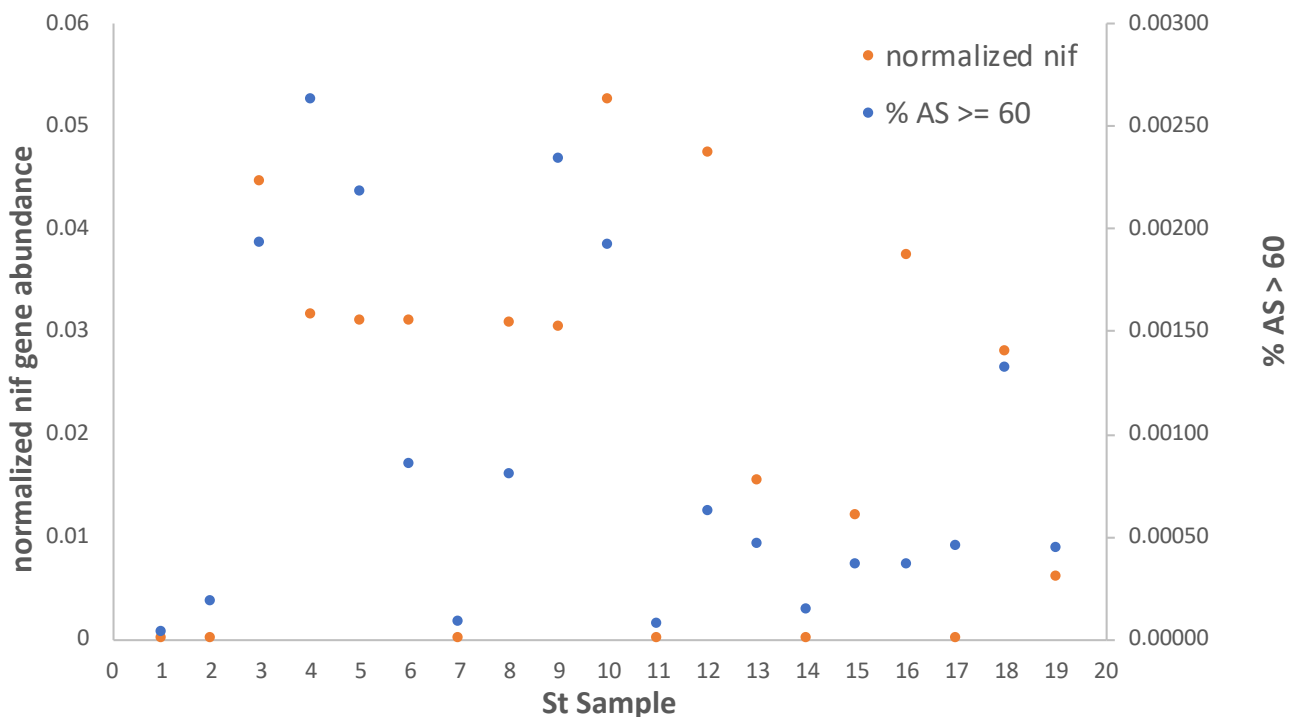


**A2 Figure 6:** *rpoB* normalized *nif* gene abundance and percentage of reads with an Alignment Score (AS) over 60, for samples obtained from Russell Glacier (RI), Greenland. The Alignment Score ranges between 0 and the maximum length of the reads (0-150) and indicates the quality of the alignment between reads and contigs.





**A2 Figure 7:** *rpoB* normalized *nif* gene abundance and percentage of reads with an Alignment Score (AS) over 60, for samples obtained from Rabots Glacier (Rb), N-Sweden. The Alignment Score ranges between 0 and the maximum length of the reads (0-150) and indicates the quality of the alignment between reads and contigs.



**A2 Figure 8:** *rpoB* normalized *nif* gene abundance and percentage of reads with an Alignment Score (AS) over 60, for samples obtained from Storglaciären Glacier (St), N-Sweden. The Alignment Score ranges between 0 and the maximum length of the reads (0-150) and indicates the quality of the alignment between reads and contigs.

**A2 Table 3:** Mapping alignments between *nif* genes and raw sequencing reads. The number of read alignments in each metagenome is shown, grouped by the alignment score (AS)  $\geq 30$ ,  $\geq 60$ ,  $\geq 90$ ,  $\geq 120$  and  $\geq 140$ . A total for each forefield is also provided. The alignment score represents the quality of the alignment and ranges between 0-100 for the MI dataset and 0-150 for the Rb, St and RI datasets. The number of sequencing reads is provided, shown as the total number of forward and reverse reads for each sample. The percentage of alignments with an AS equal or higher than 60 is listed for each metagenome sample (% AS  $\geq 60$ ).

Sample	Number of reads	AS $\geq$ 30	AS $\geq$ 60	AS $\geq$ 90	AS $\geq$ 120	AS $\geq$ 140	% AS $\geq$ 60
MI 1	17465080	407	342	280	-	-	0.0019582
MI 2	28186722	1175	1089	947	-	-	0.0038635
MI 3	16801828	62	39	20	-	-	0.0002321
MI 4	25805412	466	268	79	-	-	0.0010385
MI 5	20630402	242	118	16	-	-	0.0005720
MI 6	21421268	1653	1008	449	-	-	0.0047056
MI 7	22655970	172	91	8	-	-	0.0004017
MI 8	20003790	1040	542	69	-	-	0.0027095
MI 9	22650764	344	175	17	-	-	0.0007726
MI 10	58387594	2260	1336	230	-	-	0.0022882
MI 11	56518916	244	135	30	-	-	0.0002389
MI 12	37925220	314	161	22	-	-	0.0004245
MI 13	3952496	72	38	3	-	-	0.0009614
MI 14	22179120	660	355	46	-	-	0.0016006
MI 15	6433798	164	93	11	-	-	0.0014455
MI 16	22231192	125	68	8	-	-	0.0003059
MI 17	38684672	333	197	31	-	-	0.0005092
MI 18	68392102	932	560	93	-	-	0.0008188
MI 19	23405212	79	40	6	-	-	0.0001709
MI 20	20152212	37	20	8	-	-	0.0000992
MI 21	28777734	98	61	12	-	-	0.0002120
MI 22	22182332	24	4	0	-	-	0.0000180
MI 23	28234260	6	0	0	-	-	0.0000000
MI forefield	633078096	10909	6740	2385	0	0	0.0010646
RI 1	67748804	28	0	0	0	0	0.0000000
RI 2	77474456	20	0	0	0	0	0.0000000
RI 3	61199512	39	11	7	0	4	0.0000180
RI 4	74992750	130	80	64	23	49	0.0001067
RI 5	74220220	109	60	43	16	34	0.0000808
RI 6	68161082	240	156	121	46	90	0.0002289
RI 7	75975046	28	0	0	0	0	0.0000000
RI 8	91844214	47	0	0	0	0	0.0000000
RI 9	77023868	34	6	4	2	3	0.0000078
RI 10	139381778	32	5	1	0	0	0.0000036
RI 11	91962344	131	63	32	16	27	0.0000685

<b>RI 12</b>	81681122	61	29	18	1	9	0.0000355
<b>RI 13</b>	105431412	190	143	92	18	53	0.0001356
<b>RI 14</b>	63240678	26	10	4	0	3	0.0000158
<b>RI 15</b>	72243790	428	215	76	5	29	0.0002976
<b>RI 16</b>	82137112	23	6	1	0	0	0.0000073
<b>RI 17</b>	107570706	74	24	10	4	6	0.0000223
<b>RI 18</b>	78274268	33	3	1	0	1	0.0000038
<b>RI 19</b>	75392190	30	10	2	0	0	0.0000133
<b>RI 20</b>	65582044	16	0	0	0	0	0.0000000
<b>RI 21</b>	91522052	24	1	0	0	0	0.0000011
<b>RI 22</b>	96384726	43	14	3	0	2	0.0000145
<b>RI 23</b>	96207238	5	0	0	0	0	0.0000000
<b>RI 24</b>	80981820	30	9	3	0	1	0.0000111
<b>RI forefield</b>	1996633232	1821	845	482	131	311	0.0000423
<b>Rb 1</b>	76498938	95	29	23	2	11	0.0000379
<b>Rb 2</b>	72691820	211	117	76	21	49	0.0001610
<b>Rb 3</b>	70126632	379	199	132	55	104	0.0002838
<b>Rb 4</b>	61828928	584	396	307	99	226	0.0006405
<b>Rb 5</b>	74667258	215	105	52	10	32	0.0001406
<b>Rb Forefield</b>	355813576	1484	846	590	187	422	0.0002378
<b>St 1</b>	68111048	247	120	67	8	37	0.0001762
<b>St 2</b>	71828498	1817	1385	1031	480	749	0.0019282
<b>St 3</b>	85214054	2871	2235	1571	734	1143	0.0026228
<b>St 4</b>	71411294	1967	1552	1150	484	804	0.0021733
<b>St 5</b>	66910678	726	563	457	285	389	0.0008414
<b>St 6</b>	72683122	142	59	39	17	28	0.0000812
<b>St 7</b>	60148730	645	481	349	190	274	0.0007997
<b>St 8</b>	67844804	1956	1586	1213	784	1026	0.0023377
<b>St 9</b>	63953088	1447	1222	1003	616	851	0.0019108
<b>St 10</b>	83010234	102	60	44	25	38	0.0000723
<b>St 11</b>	74901072	577	459	369	182	297	0.0006128
<b>St 12</b>	83405572	586	384	259	100	182	0.0004604
<b>St 13</b>	64225756	119	86	71	19	44	0.0001339
<b>St 14</b>	64764076	295	229	182	83	140	0.0003536
<b>St 15</b>	66675200	350	238	185	101	150	0.0003570
<b>St 16</b>	66848090	363	302	250	127	203	0.0004518
<b>St 17</b>	72051286	1064	947	760	447	611	0.0013143
<b>St 18</b>	78397352	485	346	260	118	209	0.0004413
<b>St Forefield</b>	1282383954	15759	12254	9260	4800	7175	0.0009556

## Appendix 3

**A3 Table 1:** COG domains used in Tree construction using KBASE Species Tree Builder for binned genomes, sourced from <https://kbase.us>

COG	Gene	Description
COG0013	AlaS	Alanyl-tRNA synthetase
COG0016	PheS	Phenylalanyl-tRNA synthetase alpha subunit
COG0018	ArgS	Arginyl-tRNA synthetase
COG0030	KsgA	Dimethyladenosine transferase (rRNA methylation)
COG0041	PurE	Phosphoribosylcarboxyaminoimidazole (NCAIR) mutase
COG0046	PurL	Phosphoribosylformylglycinamide (FGAM) synthase, synthetase domain
COG0048	RpsL	Ribosomal protein S12
COG0049	RpsG	Ribosomal protein S7
COG0051	RpsJ	Ribosomal protein S10
COG0052	RpsB	Ribosomal protein S2
COG0072	PheT	Phenylalanyl-tRNA synthetase beta subunit
COG0080	RplK	Ribosomal protein L11
COG0081	RplA	Ribosomal protein L1
COG0082	AroC	Chorismate synthase
COG0086	RpoC	DNA-directed RNA polymerase, beta' subunit/160 kD subunit
COG0087	RplC	Ribosomal protein L3
COG0088	RplD	Ribosomal protein L4
COG0089	RplW	Ribosomal protein L23
COG0090	RplB	Ribosomal protein L2
COG0091	RplV	Ribosomal protein L22
COG0092	RpsC	Ribosomal protein S3
COG0093	RplN	Ribosomal protein L14
COG0094	RplE	Ribosomal protein L5
COG0096	RpsH	Ribosomal protein S8
COG0097	RplF	Ribosomal protein L6P/L9E
COG0098	RpsE	Ribosomal protein S5
COG0099	RpsM	Ribosomal protein S13
COG0100	RpsK	Ribosomal protein S11
COG0102	RplM	Ribosomal protein L13
COG0103	RpsI	Ribosomal protein S9
COG0105	Ndk	Nucleoside diphosphate kinase
COG0126	Pgk	3-phosphoglycerate kinase
COG0127	COG0127	Xanthosine triphosphate pyrophosphatase

COG0130	TruB	Pseudouridine synthase
COG0150	PurM	Phosphoribosylaminoimidazole (AIR) synthetase
COG0151	PurD	Phosphoribosylamine-glycine ligase
COG0164	RnhB	Ribonuclease HII
COG0172	SerS	Seryl-tRNA synthetase
COG0185	RpsS	Ribosomal protein S19
COG0186	RpsQ	Ribosomal protein S17
COG0215	CysS	Cysteinyl-tRNA synthetase
COG0244	RplJ	Ribosomal protein L10
COG0256	RplR	Ribosomal protein L18
COG0343	Tgt	Queuine/archaeosine tRNA-ribosyltransferase
COG0504	PyrG	CTP synthase (UTP-ammonia lyase)
COG0519	GuaA	GMP synthase, PP-ATPase domain/subunit
COG0532	InfB	Translation initiation factor 2 (IF-2; GTPase)
COG0533	QRI7	Metal-dependent proteases with possible chaperone activity

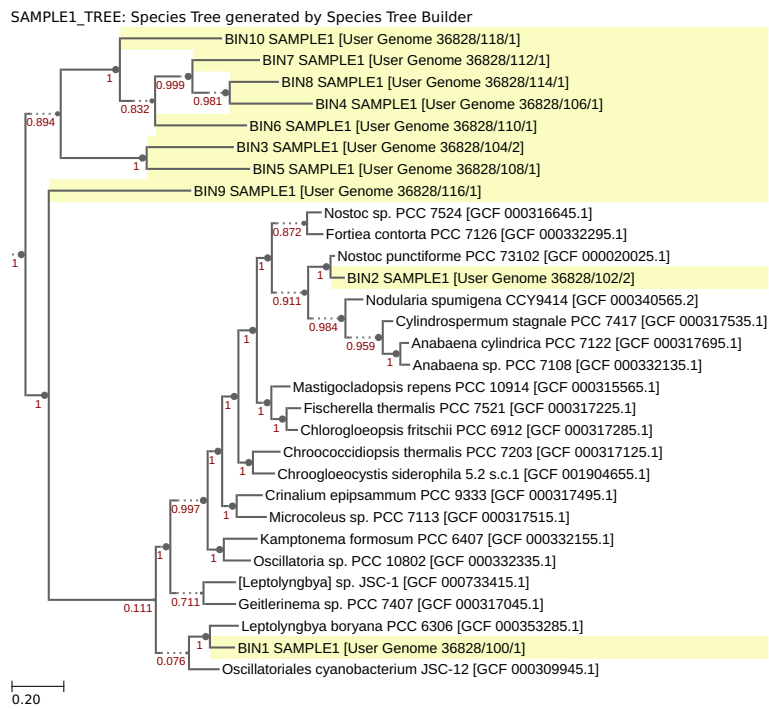
**A3 Table 2:** Read based community composition, expressed as the percentage of annotated reads for Midtre Lovénbreen sample sites at the genus level. Replicate samples for each soil age have been combined. Green indicates the recovery of the genus from each sample.

Genus	0	3	5	29	50	50 - 113	113	2000	Cryoconite	Basal Ice
<i>Acidobacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.04	0.00
<i>Actinoplanes</i>	0.00	0.30	0.00	0.00	0.00	0.00	0.16	0.00	0.00	0.00
<i>Arthrobacter</i>	0.00	0.29	0.21	0.43	0.00	0.00	0.00	0.00	0.00	0.00
<i>Bacillales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.55	0.00
<i>Bacillus</i>	0.00	0.00	0.46	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Bacteroidales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.61	0.00
<i>Bradyrhizobium</i>	0.00	1.42	1.03	1.48	1.23	0.75	1.18	2.78	0.00	0.67
<i>Brevundimonas</i>	0.87	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Burkholderia</i>	0.00	0.26	0.00	0.00	0.00	0.00	0.17	0.00	0.00	0.00
<i>Burkholderiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	13.96	0.00
<i>Cand. Nitrosocosmicus</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.37	0.00	0.00	0.00
<i>Candidatus Solibacter</i>	0.00	0.26	0.00	0.00	0.00	0.00	0.19	0.17	0.00	0.00
<i>Caulobacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.62	0.00
<i>Cellulomonas</i>	0.00	0.00	0.23	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Chitinophagales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.70	0.00
<i>Chromatiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.65	0.00
<i>Chroococcales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.86	0.00
<i>Chthoniobacter</i>	0.00	0.00	0.00	0.00	0.00	0.76	0.52	0.71	0.00	0.00
<i>Clostridiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.57	0.00
<i>Collinsella</i>	0.00	0.00	0.00	0.42	0.00	0.00	0.00	0.00	0.00	0.00
<i>Conexibacter</i>	0.00	0.89	0.00	0.23	0.00	0.00	1.00	0.41	0.00	0.00
<i>Corynebacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.74	0.00
<i>Cryobacterium</i>	0.51	0.41	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Cytophagales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	3.83	0.00
<i>Desulfuromonadales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.66	0.00
<i>Devosia</i>	0.00	0.31	0.68	0.41	0.00	0.00	0.00	0.00	0.00	0.00
<i>Enterobacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.60	0.00
<i>Enterococcus</i>	3.42	2.72	0.67	0.00	0.00	1.25	0.00	0.77	0.00	1.88
<i>Flavisolibacter</i>	0.00	0.32	0.00	0.00	0.00	0.00	0.14	0.00	0.00	0.00
<i>Flavobacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.65	0.00
<i>Flavobacterium</i>	1.45	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Frankia</i>	0.00	0.37	0.00	0.00	0.00	0.00	0.16	0.00	0.00	0.00
<i>Frankiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.57	0.00
<i>Gemmata</i>	0.00	0.45	0.19	0.83	0.72	1.00	1.43	0.00	0.00	0.00
<i>Gemmatimonadales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.77	0.00
<i>Gemmatimonas</i>	0.00	0.34	0.00	0.00	0.23	0.00	0.19	0.00	0.00	0.00
<i>Geobacter</i>	0.00	0.00	0.18	0.00	0.00	0.66	0.00	0.00	0.00	0.00

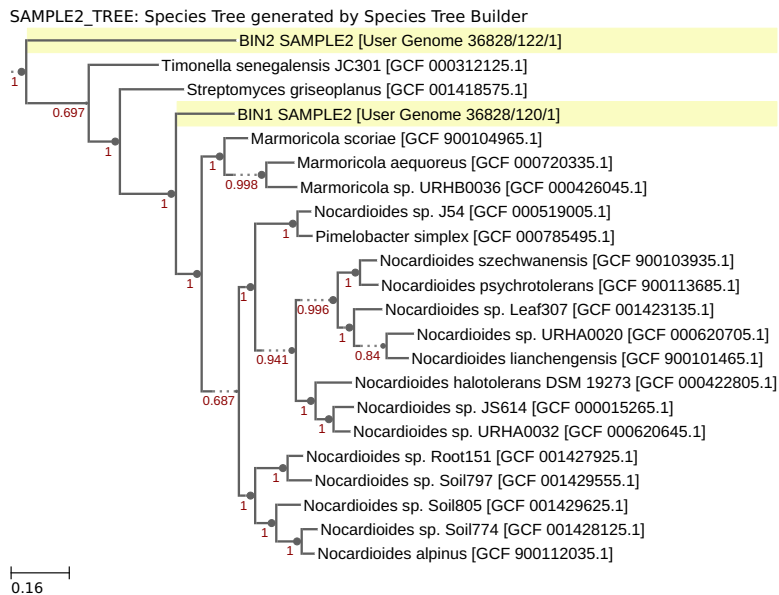
<i>Hymenobacter</i>	1.17	0.00	0.00	0.00	0.00	0.00	0.16	0.00	0.00	0.00
<i>Hyphomicrobium</i>	0.00	0.00	0.19	0.35	0.57	0.00	0.51	0.00	0.00	0.00
<i>Ilumatobacter</i>	0.00	0.55	0.00	0.18	0.00	0.00	0.32	0.00	0.00	0.00
<i>Intrasporangium</i>	0.00	0.34	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Kribbella</i>	0.00	0.25	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Luteipulveratus</i>	0.00	0.27	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Marmoricola</i>	0.00	0.00	0.29	0.77	0.71	0.00	0.13	0.00	0.00	0.00
<i>Massilia</i>	0.00	0.00	0.00	0.00	0.00	1.55	0.00	0.00	0.00	0.00
<i>Mesorhizobium</i>	0.00	0.35	0.42	0.73	0.78	0.00	0.70	0.66	0.00	0.00
<i>Methylibium</i>	0.66	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Methylobacterium</i>	0.00	0.25	0.00	0.00	0.00	0.00	0.17	0.00	0.00	0.00
<i>Methylophilales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.61	0.00
<i>Micrococcales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	8.30	0.00
<i>Micromonosporales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.65	0.00
<i>Mycobacterium</i>	0.00	0.79	0.66	1.23	1.32	1.26	1.73	1.35	0.00	0.00
<i>Myxococcales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.17	0.00
<i>Neisseria</i>	0.00	0.00	0.00	0.39	0.00	0.00	0.00	0.00	0.00	0.00
<i>Niabella</i>	0.00	0.23	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Niastella</i>	0.00	0.31	0.00	0.00	0.00	0.00	0.16	0.00	0.00	0.00
<i>Nitrospira</i>	0.00	0.29	0.16	0.16	0.52	0.50	0.29	0.16	0.00	0.00
<i>Nocardioides</i>	0.00	2.45	1.97	3.15	2.28	1.10	1.42	0.00	0.00	0.96
<i>Nostoc</i>	1.93	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Nostocales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.86	0.00
<i>Novosphingobium</i>	0.54	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.04
<i>Opitutus</i>	0.00	0.00	0.37	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Oscillatoriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	3.29	0.00
<i>Pedobacter</i>	1.15	0.41	0.18	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Pedosphaera</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.39	0.00	0.00
<i>Phormidesmis</i>	6.43	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Phycococcus</i>	0.00	0.00	0.92	0.83	0.35	0.52	0.00	0.00	0.00	0.60
<i>Pimelobacter</i>	0.00	0.32	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Pirellula</i>	0.00	0.24	0.22	0.00	0.00	0.04	0.21	0.00	0.00	0.00
<i>Planctomyces</i>	0.00	0.25	0.00	0.00	0.00	0.00	0.21	0.00	0.00	0.00
<i>Planctomycetales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.62	0.00
<i>Polaromonas</i>	3.12	1.02	0.94	1.02	0.58	0.82	0.00	0.25	0.00	0.52
<i>Propionibacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.12	0.00
<i>Pseudomonadales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.12	0.00
<i>Pseudomonas</i>	0.00	2.51	0.39	0.17	0.39	1.15	0.21	0.20	0.00	0.63
<i>Pseudonocardia</i>	0.00	0.37	0.16	0.74	0.35	0.00	0.43	0.00	0.00	0.00
<i>Pseudonocardiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.24	0.00
<i>Purpureocillium</i>	0.58	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Pyrinomonas</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.51	0.00	0.00
<i>Rhizobacter</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00

<i>Rhizobiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	4.59	0.00
<i>Rhizobium</i>	1.22	0.00	0.19	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Rhodobacterales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.82	0.00
<i>Rhodococcus</i>	0.00	0.40	0.00	0.00	0.00	0.00	0.19	0.00	0.00	0.00
<i>Rhodocyclales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.73	0.00
<i>Rhodoferax</i>	0.68	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Rhodoplanes</i>	0.00	0.27	0.00	0.00	0.00	0.00	0.19	0.68	0.00	0.00
<i>Rhodopseudomonas</i>	0.00	0.36	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<i>Rhodospirillales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.66	0.00
<i>Shewanella</i>	0.00	0.00	5.35	0.00	0.00	0.00	0.00	0.00	0.00	15.11
<i>Singulisphaera</i>	0.00	0.00	0.00	0.21	0.00	0.00	0.52	0.00	0.00	0.00
<i>Solibacterales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.65	0.00
<i>Solirubrobacter</i>	0.00	0.00	0.00	0.48	0.69	0.00	0.93	1.23	0.00	0.00
<i>Sorangium</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.14	0.00	0.00	0.00
<i>Sphingobacteriales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.53	0.00
<i>Sphingobium</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.15	0.00	0.00	0.00
<i>Sphingomonadales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.28	0.00
<i>Sphingomonas</i>	1.25	1.47	1.76	2.39	2.07	1.64	1.74	0.73	0.00	1.01
<i>Sphingopyxis</i>	0.00	0.27	0.00	0.00	0.00	0.00	0.15	0.00	0.00	0.00
<i>Spirosoma</i>	0.00	0.30	0.28	0.00	0.00	0.00	0.20	0.00	0.00	0.00
<i>Streptomyces</i>	0.52	1.67	1.42	1.99	1.81	1.24	2.09	1.32	0.00	0.83
<i>Streptomycetales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.93	0.00
<i>Streptosporangiales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.59	0.00
<i>Synechococcales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	4.80	0.00
<i>Thiobacillus</i>	0.79	1.10	0.28	0.00	0.00	0.00	0.00	0.00	0.00	0.80
<i>Variovorax</i>	1.05	0.00	0.20	0.20	0.93	0.64	0.28	0.00	0.00	0.00
<i>Xanthomonadales</i>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.47	0.00
<i>Zavarzinella</i>	0.00	0.00	0.00	0.00	0.00	0.03	0.29	0.00	0.00	0.00
Viruses	0.08	0.13	0.05	0.04	0.03	0.05	0.07	0.02	0.22	0.05
cannot be assigned genus > 0.5% of all reads	31.29	26.02	33.30	33.91	36.86	35.40	30.08	45.24	6.94	31.93
	41.29	48.47	46.66	47.25	47.58	49.64	50.84	42.42	15.44	43.96

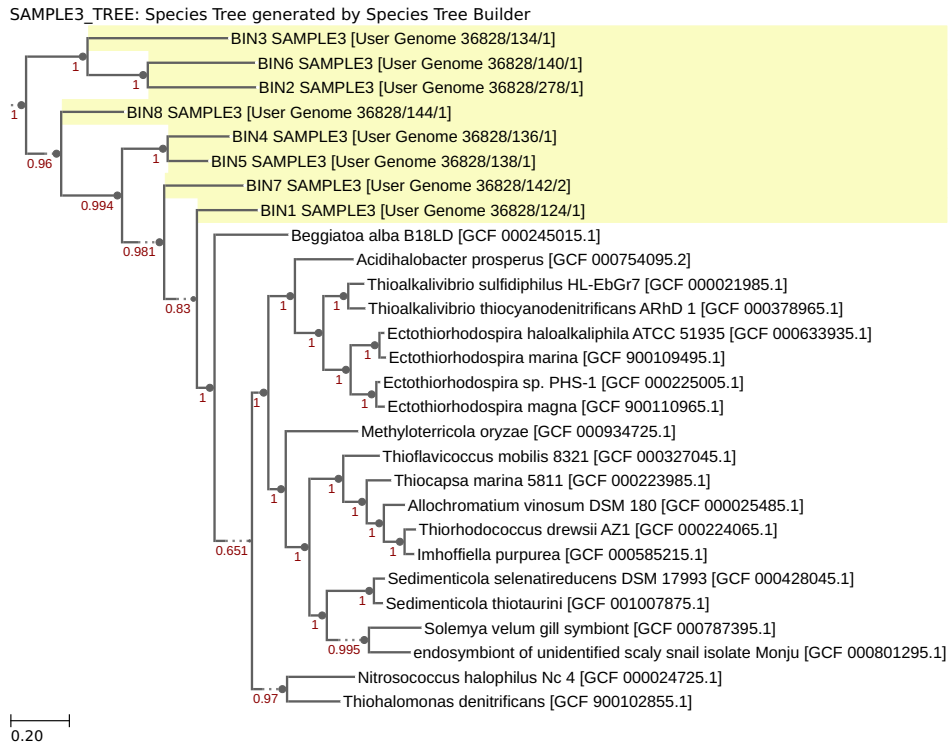




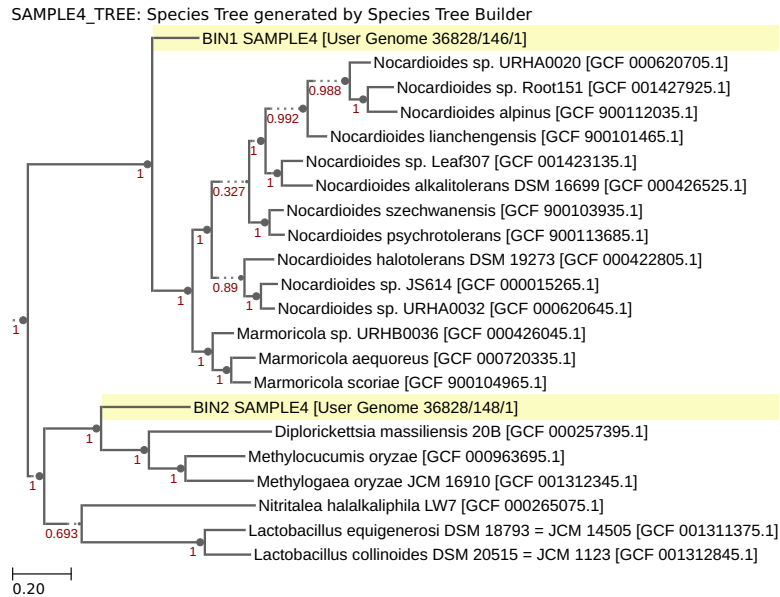
**A3 Figure 1:** Maximum likelihood Species tree for Sample 1, bins 1-10. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



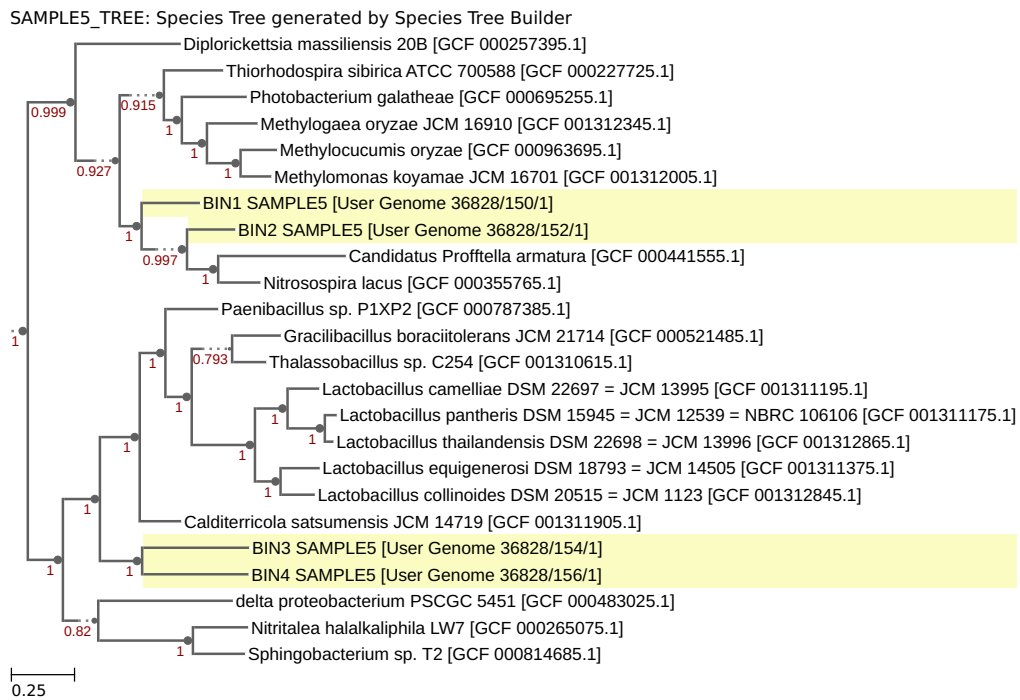
**A3 Figure 2:** Maximum likelihood Species tree for Sample 2, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



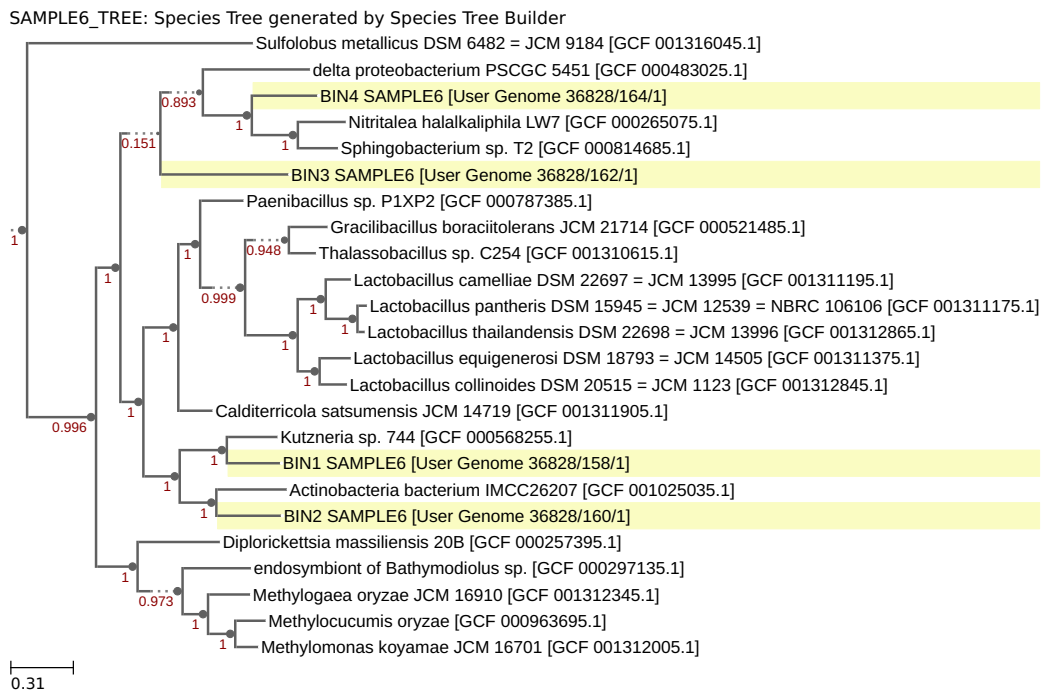
**A3 Figure 3:** Maximum likelihood Species tree for Sample 3, bins 1-8. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



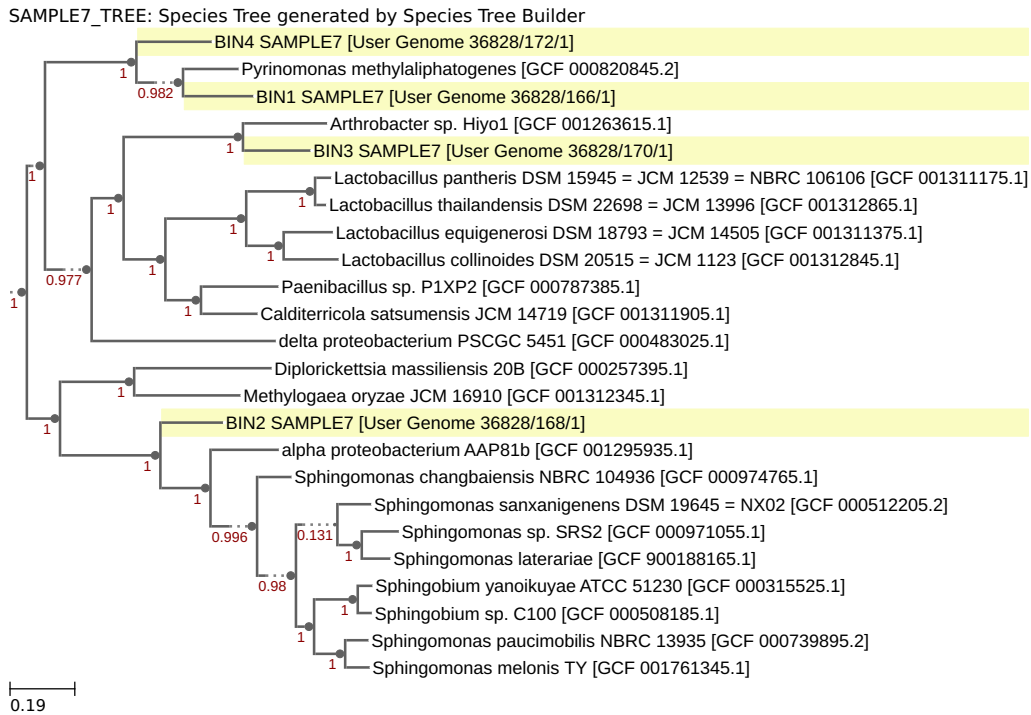
**A3 Figure 4:** Maximum likelihood Species tree for Sample 4, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



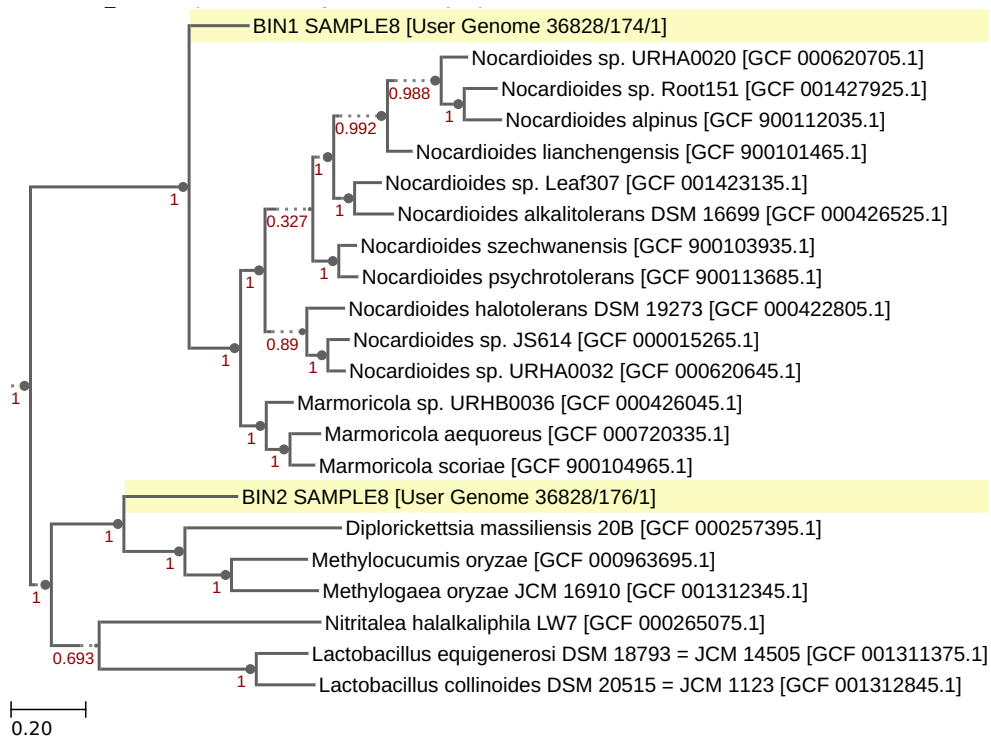
**A3 Figure 5:** Maximum likelihood Species tree for Sample 1, bins 1-4. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



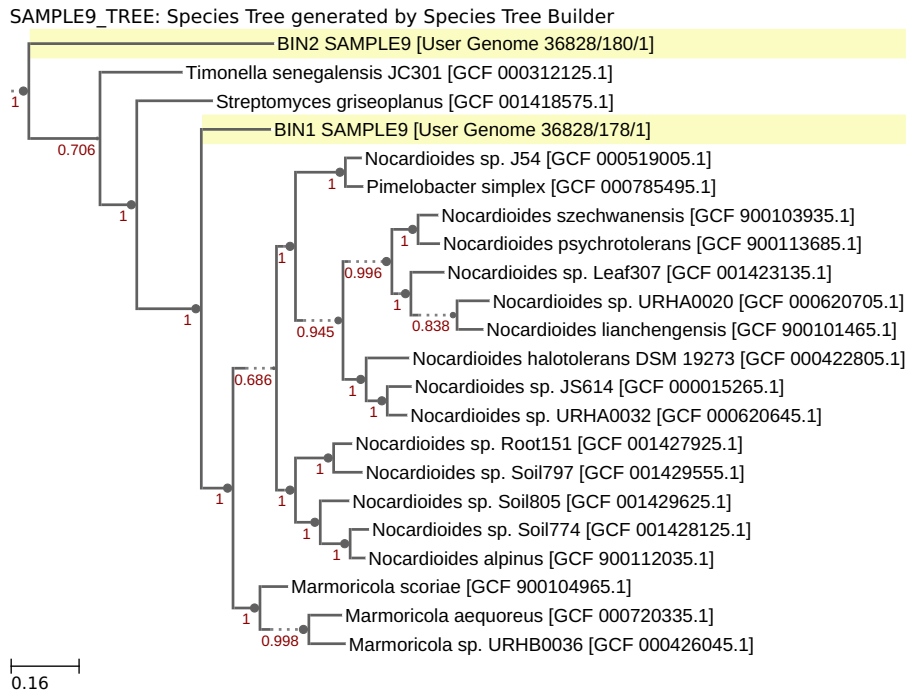
**A3 Figure 6:** Maximum likelihood Species tree for Sample 6, bins 1 - 4. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



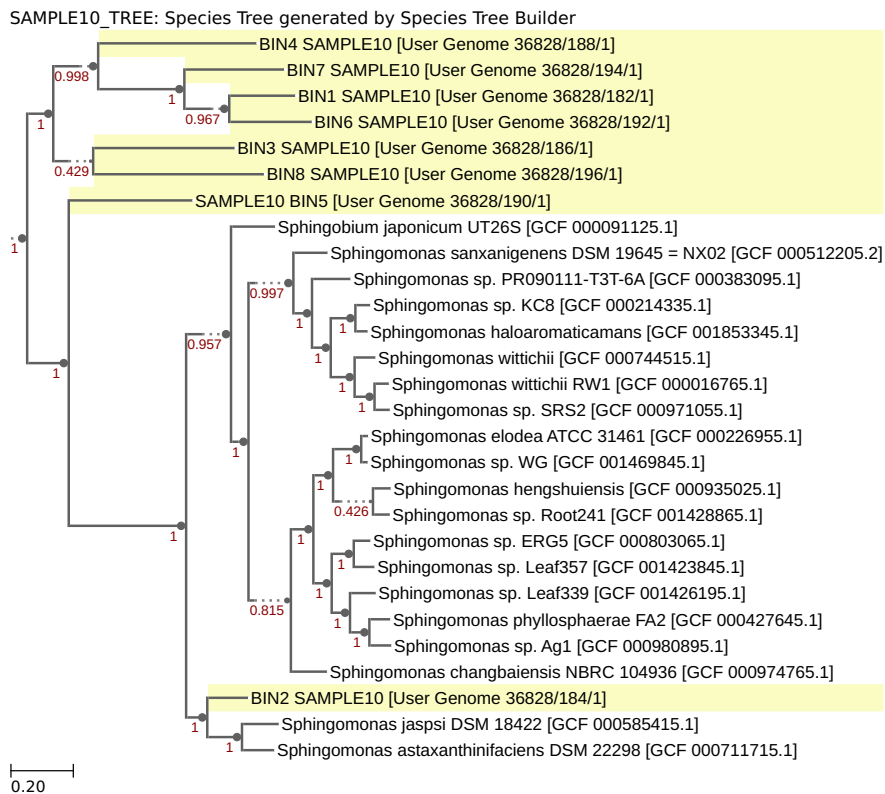
**A3 Figure 7:** Maximum likelihood Species tree for Sample 1, bins 1-4. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



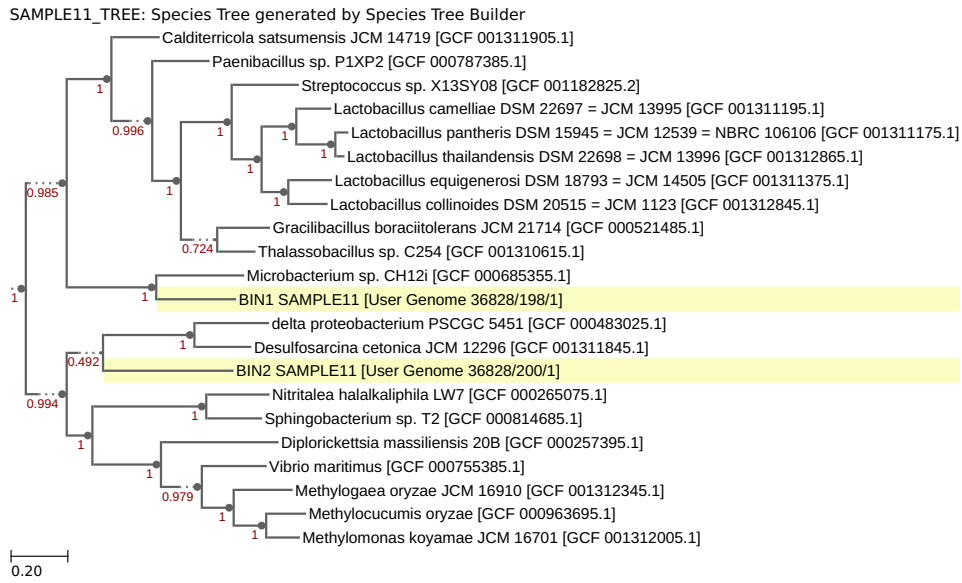
**A3 Figure 8:** Maximum likelihood Species tree for Sample 8, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



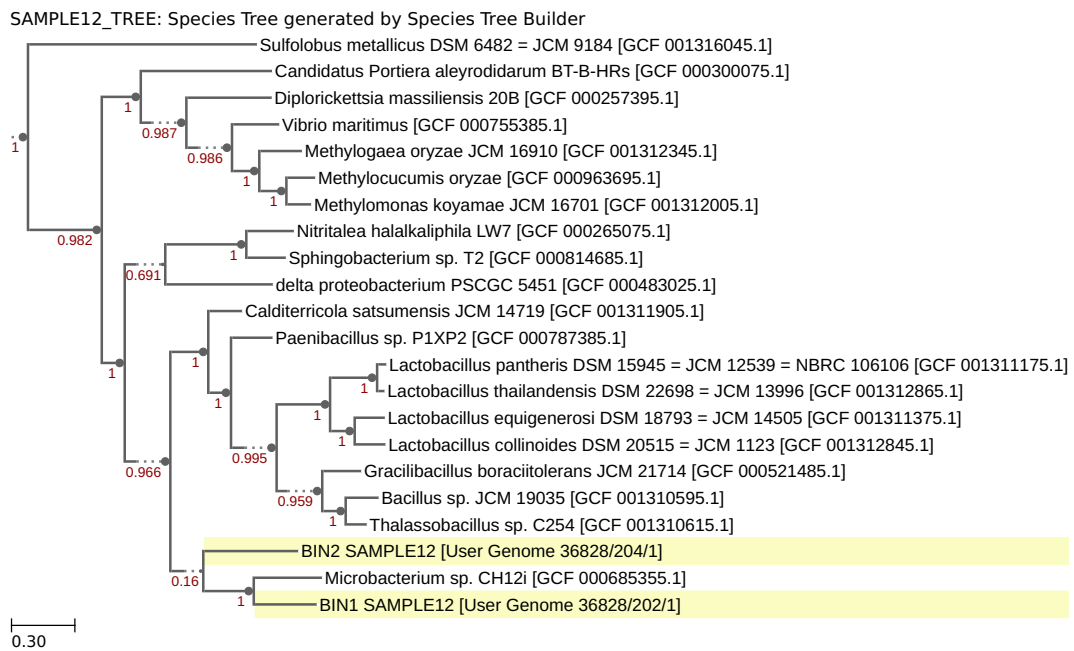
**A3 Figure 9:** Maximum likelihood Species tree for Sample 9, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



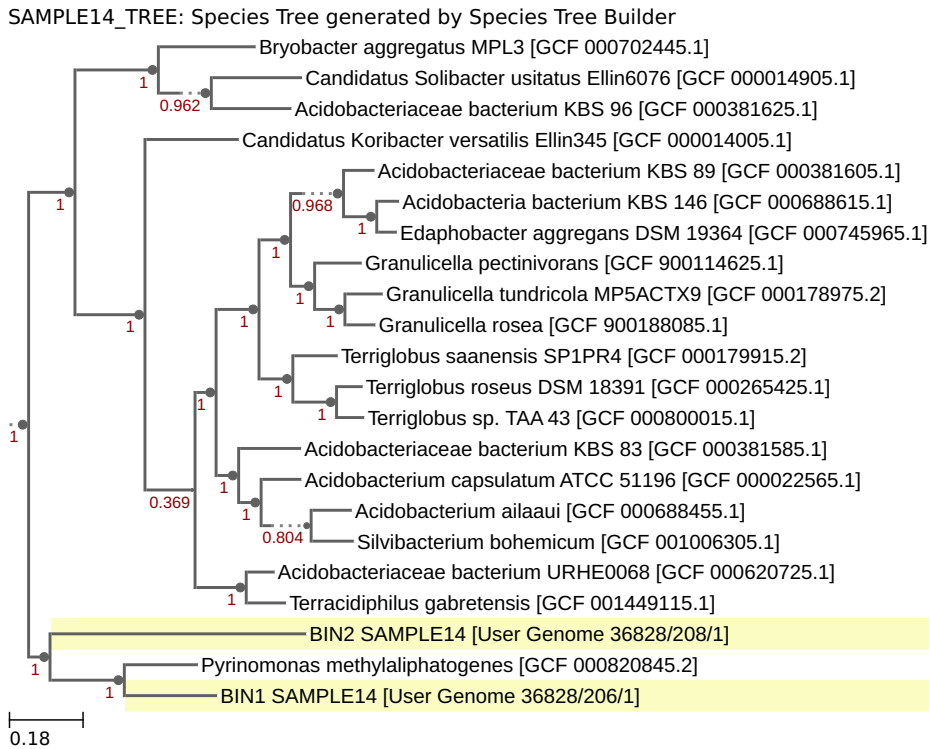
**A3 Figure 10:** Maximum likelihood Species tree for Sample 1, bins 1-8. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



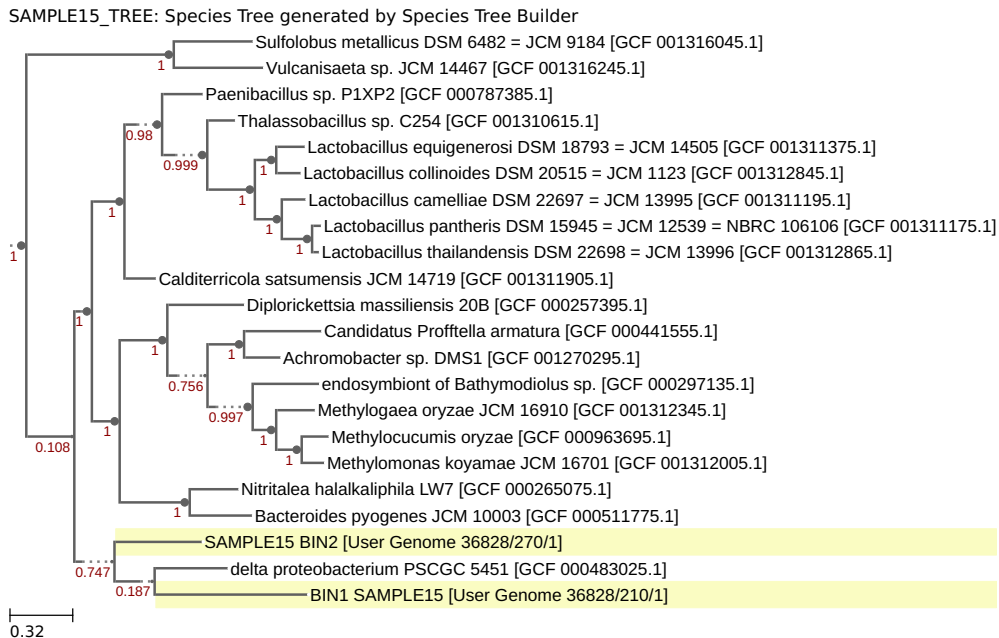
**A3 Figure 11:** Maximum likelihood Species tree for Sample 11, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



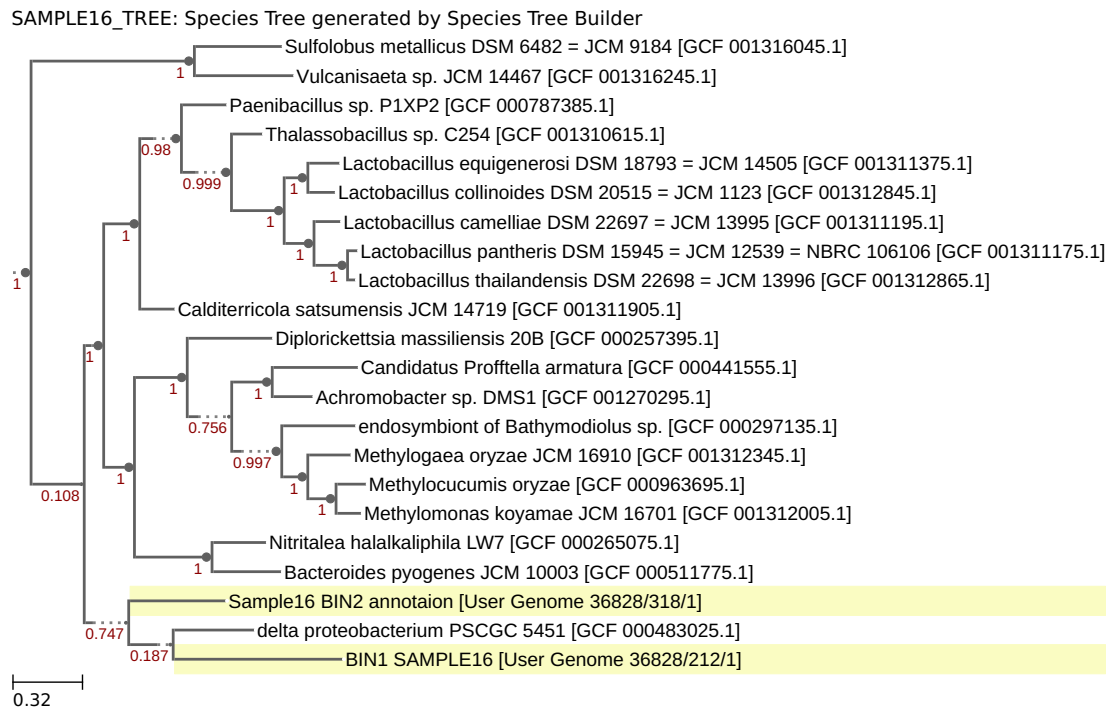
**A3 Figure 12:** Maximum likelihood Species tree for Sample 12, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



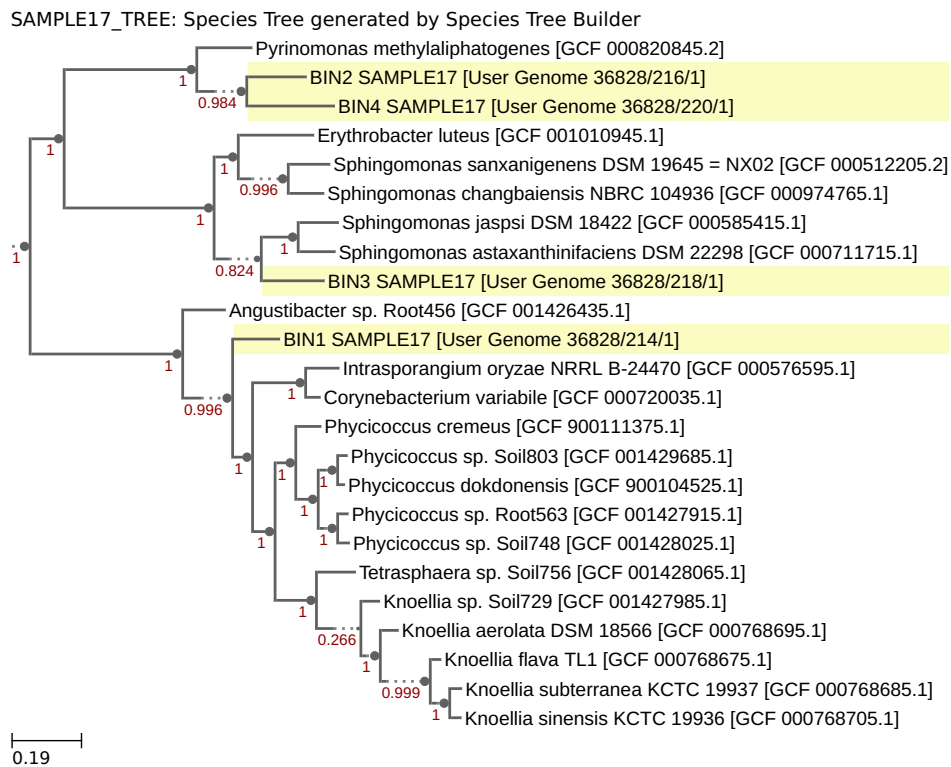
**A3 Figure 13:** Maximum likelihood Species tree for Sample 14, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



**A3 Figure 14:** Maximum likelihood Species tree for Sample 15, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



**A3 Figure 15:** Maximum likelihood Species tree for Sample 16, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.

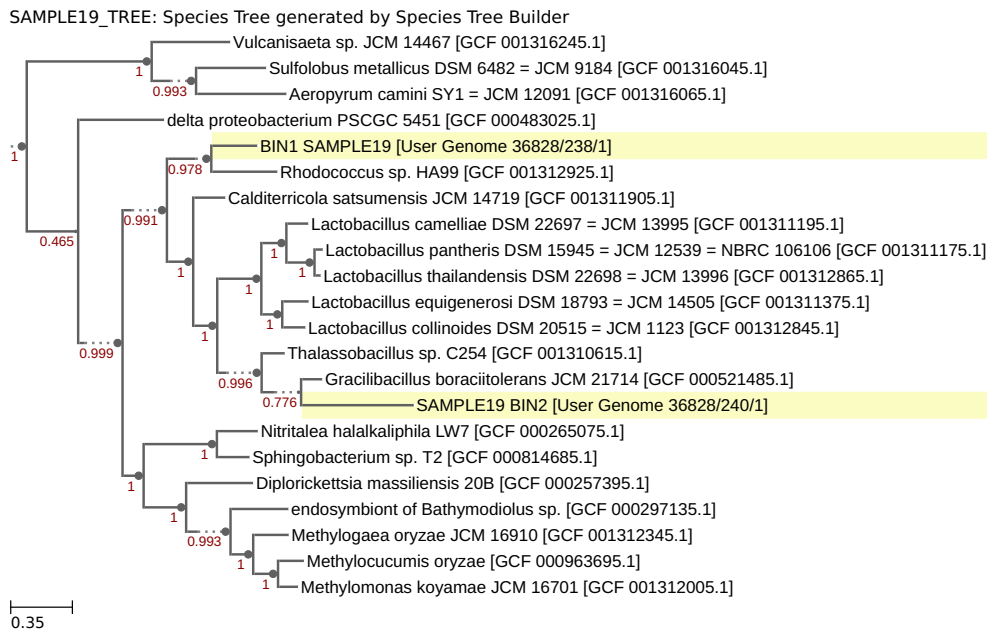


**A3 Figure 16:** Maximum likelihood Species tree for Sample 17, bins 1-4. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.

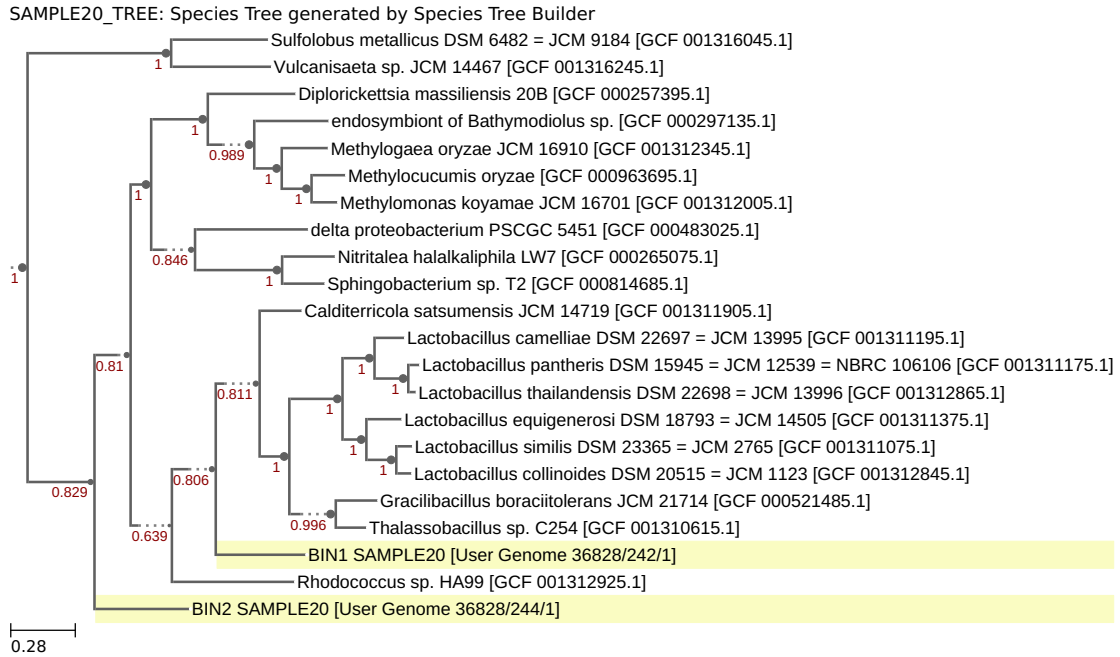




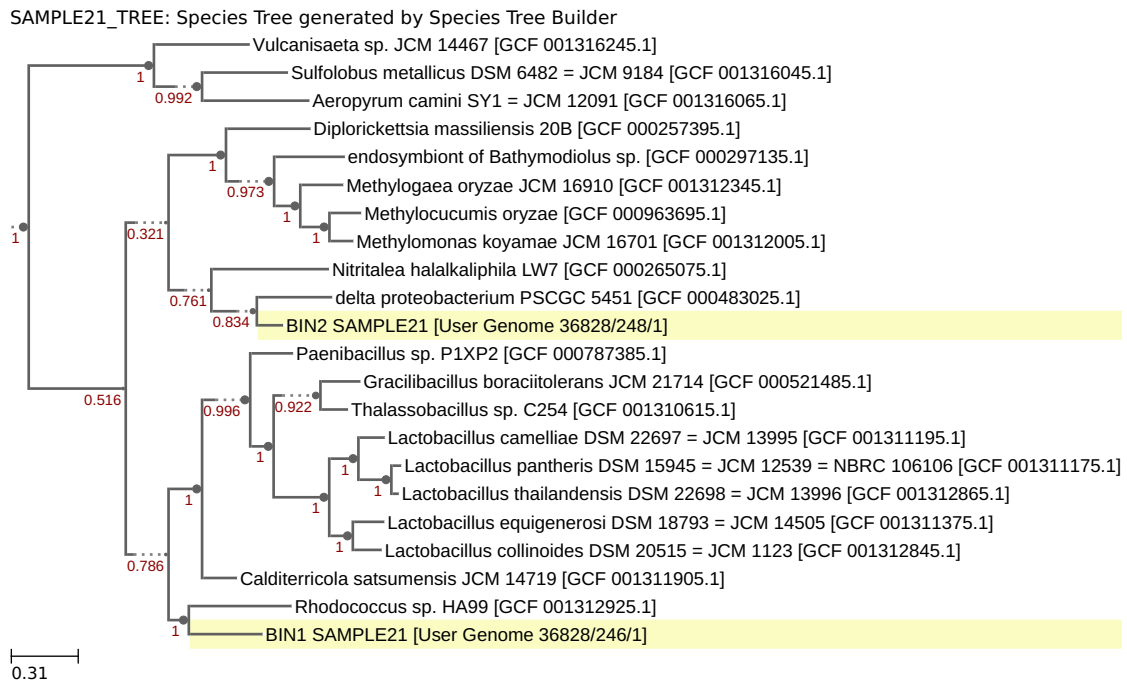
**A3 Figure 17:** Maximum likelihood Species tree for Sample 18, bins 1-8. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



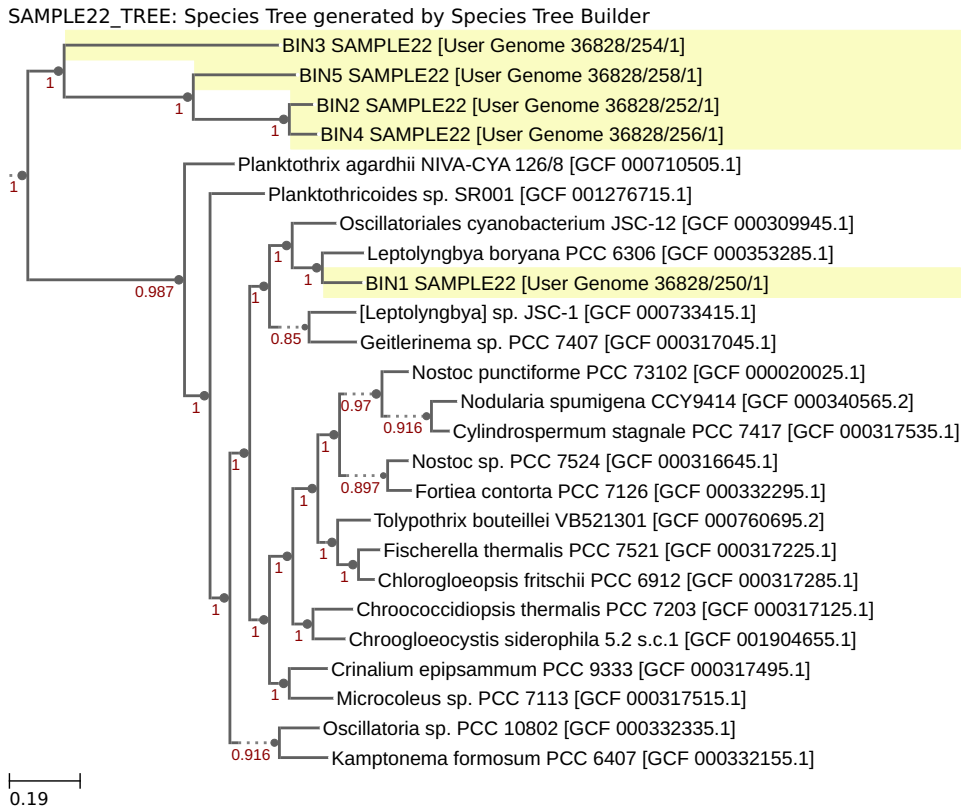
**A3 Figure 18:** Maximum likelihood Species tree for Sample 19, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



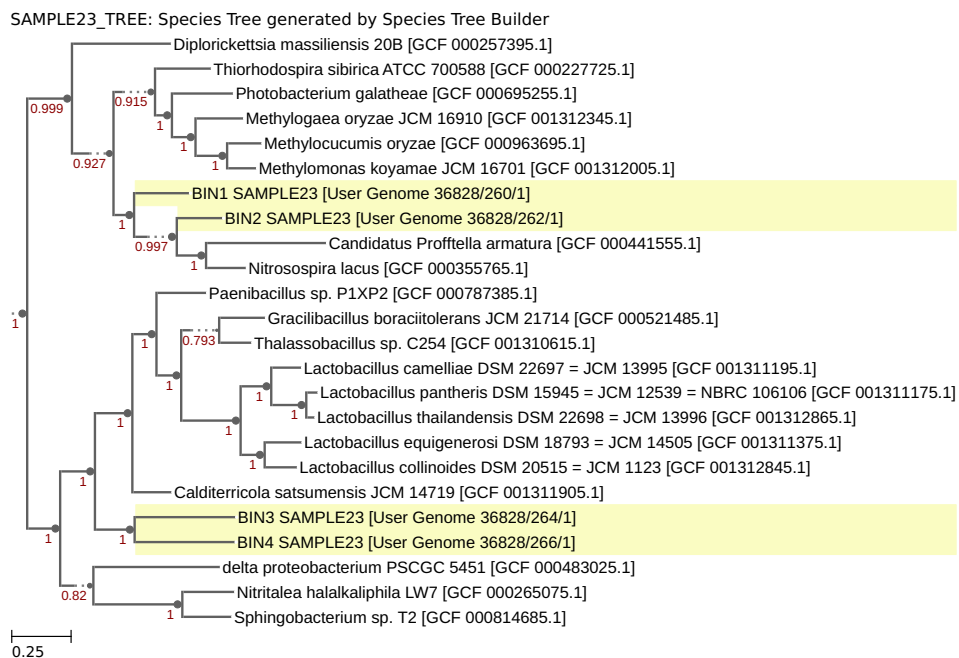
**A3 Figure 19:** Maximum likelihood Species tree for Sample 20, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



**A3 Figure 20:** Maximum likelihood Species tree for Sample 21, bins 1-2. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



**A3 Figure 21:** Maximum likelihood Species tree for Sample 22, bins 1-5. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.



**A3 Figure 22:** Maximum likelihood Species tree for Sample 23, bins 1-4. The species tree is based off alignments of COG groups to publicly available genomes on the KBASE platform.

# Appendix 4

**A4 Table 1: Kaiju taxonomic classification of metagenome reads for Sample 1. The percentage of the metagenome classified and the taxonomic assignment is given, from the class to species level. The percentage of the reads that were unclassified is also provided.**

%	class	%	family	%	genus	%	order	%	phylum	%	species
20.994911	Gamma proteobacteria	3.931496	Flavobacteriaceae	2.959363	Nitrosopumilus	6.542552	Rhizobiales	54.47688	Proteobacteria	2.706551	Woeseia oceani
16.85726	Alphaproteobacteria	3.080831	Rhodobacteraceae	2.706551	Woeseia	5.263172	Chromatiales	10.4805	Bacteroidales	1.679244	Candidatus Nitrosopumilus sediminis
8.074667	Deltaproteobacteria	2.959363	Nitrosopumilaceae	1.472738	Pseudomonas	4.648056	Burkholderiales	5.654521	Actinobacteria	1.042004	Candidatus Nitrosopumilus adriaticus
7.43486	Betaproteobacteria	2.706551	Woeseiaceae	0.867306	Halioglobus	4.169882	Flavobacteriales	5.388607	Firmicutes	0.65466	Candidatus Solibacter usitatus
4.995546	Actinobacteria	2.381379	Planctomycetaceae	0.814504	Geobacter	3.443874	Rhodobacteriales	3.769552	Planctomycetes	0.583466	Fuerstia marisgermanicae
4.169882	Flavobacteria	2.323833	Rhodobacteriaceae	0.773717	Bradyrhizobium	3.386523	Planctomycetales	3.233559	Thaumarchaeota	-----	-----
3.386523	Planctomycetia	1.711983	Burkholderiaceae	0.708511	Streptomyces	3.072975	Myxococcales	2.002861	Cyanobacteria	1.760236	Viruses
2.806206	Cytophagia	1.597047	Pseudomonadaceae	0.677328	Burkholderia	2.959363	Rhodospirillales	1.480751	Acidobacteria	16.12795	cannot be assigned to a species
2.66413	Clostridia	1.545995	Sphingomonadaceae	0.654666	Candidatus Solibacter	2.901428	Rhodospirillales	1.462165	Euryarchaeota	75.44589	species < 0.5% of all reads
2.281646	Bacilli	1.423672	Bradyrhizobiaceae	0.65081	Mycobacterium	2.806206	Cytophagales	1.365232	Chloroflexi	-----	-----
1.444824	Bacteroidia	1.192517	Hyphomicrobiaceae	0.583466	Fuerstia	2.685127	Aleromonadales	1.048709	Verucomicrobia	70.54341	unclassified
0.832079	Spirochaetia	1.190262	Comamonadaceae	0.550183	Azospirillum	2.107571	Sphingomonadales	0.832079	Spirochaetes	-----	-----
0.709328	Opitutae	1.147958	Rhizobiaceae	0.528953	Desulfovibrio	2.095362	Cellulimonales	0.675267	Nitrospirae	-----	-----
0.675267	Nitrospira	1.126993	Ecobithorhodospiraceae	0.51725	Nitrospira	1.984431	Clostridiales	0.549639	Deinococcus-Thermus	-----	-----
0.665197	Acidobacteria	1.048925	Chromatiaceae	0.513984	Sphingomonas	1.908261	Pseudomonadales	-----	-----	-----	-----
0.65466	Solibacteris	1.02649	Aleromonadaceae	0.511768	Paenibacillus	1.857753	Bacillales	1.760236	Viruses	-----	-----
0.567369	Methanomicrobia	0.993518	Desulfobacteraceae	-----	-----	1.856936	Desulfuromonadales	1.61208	cannot be assigned to a phylum	-----	-----
0.549639	Deinococci	0.969528	Geobacteraceae	1.760236	Viruses	1.721043	Oceanospirillales	3.860186	phylum < 0.5% of all reads	-----	-----
0.536263	Sphingobacteria	0.867306	Halleaeae	9.951588	cannot be assigned to a genus	1.47764	Desulfobacteriales	-----	-----	-----	-----
-----	-----	0.863535	Phyllobacteriaceae	72.79704	genus < 0.5% of all reads	1.352556	Corynebacteriales	70.54341	unclassified	-----	-----
-----	-----	0.849304	Desulfuromonadaceae	-----	-----	1.150097	Xanthomonadales	-----	-----	-----	-----
-----	-----	0.840789	Cytophagaceae	-----	-----	0.99414	Enterobacteriales	-----	-----	-----	-----
-----	-----	0.817693	Rhodocyclaceae	-----	-----	0.918553	Thiotrichales	-----	-----	-----	-----
-----	-----	0.806222	Polyangiaceae	-----	-----	0.898664	Bacteroidales	-----	-----	-----	-----
-----	-----	0.754081	Bacillaceae	-----	-----	0.817693	Rhodocyclales	-----	-----	-----	-----
-----	-----	0.73258	Streptomycetaceae	-----	-----	0.757347	Micrococcales	-----	-----	-----	-----
-----	-----	0.675267	Nitrospiraceae	-----	-----	0.73258	Streptomycetales	-----	-----	-----	-----
-----	-----	0.672895	Mycobacteriaceae	-----	-----	0.715394	Desulfuromonadales	-----	-----	-----	-----
-----	-----	0.670446	Isophaeraceae	-----	-----	0.675267	Nitrospirales	-----	-----	-----	-----
-----	-----	0.665197	Acidobacteriaceae	-----	-----	0.668696	Synechococcales	-----	-----	-----	-----
-----	-----	0.65466	Solibacteraceae	-----	-----	0.665197	Acidobacteriales	-----	-----	-----	-----
-----	-----	0.653493	Xanthomonadaceae	-----	-----	0.65466	Solibacterales	-----	-----	-----	-----
-----	-----	0.643539	Alcaligenaceae	-----	-----	0.641012	Vibrionales	-----	-----	-----	-----
-----	-----	0.641012	Vibrionaceae	-----	-----	0.614028	Bacteroidetes Order II. Incertae sedis	-----	-----	-----	-----
-----	-----	0.626198	Paenibacillaceae	-----	-----	0.580472	Methylococcales	-----	-----	-----	-----
-----	-----	0.614028	Rhodothermaceae	-----	-----	0.56943	Spirochaetales	-----	-----	-----	-----
-----	-----	0.610528	Hymenobacteraceae	-----	-----	0.560876	Syntrophobacteriales	-----	-----	-----	-----
-----	-----	0.588249	Cellvibrionaceae	-----	-----	0.560798	Caulobacteriales	-----	-----	-----	-----
-----	-----	0.580472	Methylococcaceae	-----	-----	0.548433	Marinilibiales	-----	-----	-----	-----
-----	-----	0.573279	Peptococcaceae	-----	-----	0.536263	Sphingobacteriales	-----	-----	-----	-----
-----	-----	0.567913	Acetobacteraceae	-----	-----	0.515811	Opitutales	-----	-----	-----	-----
-----	-----	0.561148	Clostridiaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.560798	Caulobacteraceae	-----	-----	1.760236	Viruses	-----	-----	-----	-----
-----	-----	0.558621	Cyclobacteriaceae	-----	-----	7.556406	cannot be assigned to a order	-----	-----	-----	-----
-----	-----	0.556015	Oxalobacteraceae	-----	-----	17.72192	order < 0.5% of all reads	-----	-----	-----	-----
-----	-----	0.549289	Spirochaetaceae	-----	-----	70.54341	unclassified	-----	-----	-----	-----
-----	-----	0.546723	Desulfuromonadaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.536263	Sphingobacteriaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.523082	Piscirickettsiaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.515811	Opitutaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.510773	Erythrobacteraceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.503058	Oceanospirillaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	0.503019	Colwelliaceae	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	-----	-----	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	1.760236	Viruses	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	9.324418	cannot be assigned to a family	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	34.16439	family < 0.5% of all reads	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	-----	-----	-----	-----	-----	-----	-----	-----	-----	-----
-----	-----	70.543408	unclassified	-----	-----	-----	-----	-----	-----	-----	-----

**A4 Table 2: Kaiju taxonomic classification of metagenome reads for Sample 2. The percentage of the metagenome classified and the taxonomic assignment is given, from the class to species level. The percentage of the reads that were unclassified is also provided.**

% class	% family	% genus	% order	% phylum	% species
17.56399 Gammaproteobacteria	3.544211 Flavobacteriaceae	1.794189 Woeseia	5.42386 Rhizobiales	49.03289 Proteobacteria	1.794189 Woeseia oceanii
13.95572 Alphaproteobacteria	2.242425 Rhodobacteraceae	1.617759 Nitrosopumilus	4.131728 Burkholderiales	11.09193 Bacteroidetes	0.88622 Candidatus Nitrosopumilus sediminis
9.508138 Deltaproteobacteria	2.161755 Planctomycetaceae	1.289043 Pseudomonas	4.097165 Chromatiales	6.818052 Firmicutes	0.650965 Candidatus Solibacter usitatus
6.7027 Betaproteobacteria	2.157875 Rhodospirillaceae	0.997309 Geobacter	3.795227 Flavobacteriales	5.515596 Actinobacteria	0.588354 Candidatus Nitrosopumilus adriaticus
4.878619 Actinobacteria	1.794189 Woeseia	0.704138 Desulfococcus	3.189243 Planctomycetales	3.779418 Planctomycetes	0.503038 Draconibacterium orientale
3.795227 Flavobacteriales	1.668777 Desulfobacteraceae	0.698294 Streptomyces	3.069819 Cyphobacteriales	2.16597 Cyanobacteria	-----
3.612905 Clostridia	1.617759 Nitrosopumilaceae	0.68397 Desulfovibrio	2.847546 Myxococcales	2.136989 Euryarchaeota	3.792065 Viruses
3.189243 Planctomycoetia	1.514 Burkholderiaceae	0.650965 Candidatus Solibacter	2.677966 Rhodospirillales	1.837398 Thaumarchaeota	15.79 cannot be assigned to a species
3.069819 Cyphobacteria	1.398312 Pseudomonadaceae	0.638749 Bradyrhizobium	2.645679 Bradyrhizobiales	1.707291 Chloroflexi	75.99517 species < 0.5% of all reads
2.644529 Bacilli	1.276588 Sphingomonadaceae	0.618438 Paenibacillus	2.528986 Rhodobacteriales	1.48612 Acidobacteria	-----
2.013828 Bacteroidia	1.202146 Bradyrhizobiaceae	0.616474 Burkholderia	2.295358 Alteromonadales	1.131536 Verrucomicrobia	74.19126 unclassified
1.018722 Spirochaetia	1.156111 Geobacteraceae	0.603923 Mycobacterium	2.252868 Desulfobacteriales	1.018722 Spirochaetes	-----
0.831035 Methanomicrobia	1.058147 Comamonadaceae	0.602821 Clostridium	2.153515 Bacillales	0.721048 Nitrospirae	-----
0.777862 Opitutae	0.991178 Ectothiorhodospiraceae	0.571013 Halloglobus	2.081037 Desulfuromonadales	0.624857 Ignivibacteriales	-----
0.721049 Nitrospira	0.975513 Chromatiaceae	0.529433 Nitrospira	1.890813 Pseudomonadales	0.594103 Chlorobi	-----
0.659587 Acidobacteria	0.960998 Rhizobiaceae	0.524259 Bacillus	1.680369 Sphingomonadales	0.585672 Deinococcus-Thermus	-----
0.650965 Solibacteres	0.960854 Hyphomicrobiaceae	0.503038 Draconibacterium	1.617759 Nitrosopumilales	-----	-----
0.624857 Ignivibacteria	0.914771 Cyphobacteraceae	-----	1.537999 Cellvibrionales	3.792065 Viruses	-----
0.609883 Sphingobacteria	0.880184 Desulfuromonadaceae	3.792065 Viruses	1.440707 Oceanospirillales	2.086211 cannot be assigned to a phylum	-----
0.594103 Chlorobia	0.866484 Bacillaceae	10.13678 cannot be assigned to a genus	1.273044 Conynebacteriales	3.874124 phylum < 0.5% of all reads	-----
0.586299 Phycisphaerae	0.856999 Alteromonadaceae	72.42734 genus < 0.5% of all reads	1.246457 Bacteroidales	-----	-----
0.585672 Deinococci	0.821311 Rhodocyclaceae	-----	1.012207 Xanthomonadales	74.19126 unclassified	-----
0.500355 Epsilonproteobacteria	0.799467 Peptococcaceae	74.19126 unclassified	0.945429 Enterobacteriales	-----	-----
-----	0.75391 Clostridiaceae	-----	0.926795 Desulfuromonadales	-----	-----
3.792065 Viruses	0.752569 Paenibacillaceae	-----	0.860065 Syntrichobacteriales	-----	-----
7.8979 cannot be assigned to a class	0.740497 Polyangiaceae	-----	0.821311 Rhodocyclales	-----	-----
9.218847 class < 0.5% of all reads	0.723012 Streptomycetaceae	-----	0.815993 Thiothricales	-----	-----
-----	0.721048 Nitrospiraceae	-----	0.777766 Micrococcales	-----	-----
74.19126 unclassified	0.706198 Desulfuovibrionaceae	-----	0.757647 Marinilibiales	-----	-----
-----	0.704569 Spirochaetaceae	-----	0.730677 Spirochaetales	-----	-----
-----	0.682869 Isosphaeraceae	-----	0.723012 Streptomycetales	-----	-----
-----	0.682581 Phyllobacteriaceae	-----	0.721048 Nitrospirales	-----	-----
-----	0.660066 Rhodothermaceae	-----	0.711515 Nitrococcales	-----	-----
-----	0.659587 Acidobacteriaceae	-----	0.683683 Thermoanaerobacteriales	-----	-----
-----	0.650965 Solibacteraceae	-----	0.660066 Bacteroidetes Order II. Incertae sedis	-----	-----
-----	0.649192 Hymenobacteraceae	-----	0.659587 Acidobacteriales	-----	-----
-----	0.626007 Mycobacteriaceae	-----	0.650965 Solibacterales	-----	-----
-----	0.616857 Cyclobacteriaceae	-----	0.624857 Ignivibacteriales	-----	-----
-----	0.611013 Syntrophaceae	-----	0.605983 Sphingobacteriales	-----	-----
-----	0.605983 Sphingobacteriaceae	-----	0.594103 Chlorobiales	-----	-----
-----	0.594103 Chlorobiaceae	-----	0.589887 Opitutales	-----	-----
-----	0.589887 Opituliaceae	-----	0.574702 Vibrionales	-----	-----
-----	0.583708 Xanthomonadaceae	-----	0.545241 Oscillatoriales	-----	-----
-----	0.57954 Desulfobulbaceae	-----	0.537624 Methanosarcinales	-----	-----
-----	0.574702 Vibrionaceae	-----	0.506295 Methylcoccales	-----	-----
-----	0.571013 Halieaceae	-----	-----	-----	-----
-----	0.567564 Alcaligenaceae	-----	3.792065 Viruses	-----	-----
-----	0.543277 Porphyromonadaceae	-----	7.826906 cannot be assigned to a order	-----	-----
-----	0.538535 Flammeovirgaceae	-----	17.66002 order < 0.5% of all reads	-----	-----
-----	0.51099 Acetobacteraceae	-----	-----	-----	-----
-----	0.506295 Methylcocccaceae	-----	-----	-----	-----
-----	0.503038 Prolixibacteraceae	-----	74.19126 unclassified	-----	-----
-----	-----	-----	-----	-----	-----
3.792065 Viruses	-----	-----	-----	-----	-----
9.568976 cannot be assigned to a family	-----	-----	-----	-----	-----
36.60933 family < 0.5% of all reads	-----	-----	-----	-----	-----
-----	-----	-----	-----	-----	-----
74.19126 unclassified	-----	-----	-----	-----	-----





**A4 Table 5: Kaiju taxonomic classification of metagenome reads for Sample 7. The percentage of the metagenome classified and the taxonomic assignment is given, from the class to species level. The percentage of the reads that were unclassified is also provided.**

% class	% family	% genus	% order	% phylum	% species
21.08804	Gamma proteobacteria	4.331581 Nitrosopumilaceae	6.432116 Rhizobiales	54.229175 Proteobacteria	2.815322 Woeseia oceanii
16.16458	Alphaproteobacteria	2.842116 Planctomycetaceae	5.585598 Chromatiales	6.939337 Bacteroidetes	2.038815 Candidatus Nitrosopumilus sedimin
8.26172	Deltaproteobacteria	2.815322 Woeseiaceae	4.767477 Burkholderiales	5.784874 Firmicutes	1.882659 Candidatus Nitrosopumilus adriatic
7.449125	Betaproteobacteria	2.603863 Rhodospirillaceae	4.331581 Nitrosopumilales	5.645465 Actinobacteria	0.689345 Candidatus Solibacter usitatus
5.005143	Acinobacteria	2.603863 Rhodospirillaceae	4.089393 Planctomycetales	5.145432 Thaumarchaeota	0.686457 Fuerstia marisgermanicae
4.089393	Planctomyceta	2.037391 Flavobacteriaceae	3.209103 Rhodospirillales	4.566359 Planctomycetes	0.528627 Pirellula slaleyi
2.966957	Clostridia	1.854442 Burkholderiaceae	2.897294 Rhodobacterales	2.292808 Cyanobacteria	-----
2.341288	Bacilli	1.652026 Pseudomonadaceae	2.876612 Myxococcales	1.776448 Euryarchaeota	1.368434 Viruses
2.202966	Flavobacteria	1.456685 Bradyrhizobiaceae	2.639155 Alteromonadales	1.520026 Acidobacteria	16.14101 cannot be assigned to a species
2.182787	Cytophaga	1.389031 Sphingomonadaceae	2.202966 Flavobacteriales	1.436422 Chloroflexi	73.84934 species < 0.5% of all reads
1.018151	Nitrospira	1.275327 Desulfobacteraceae	2.182787 Cytophagales	1.018151 Nitrospirae	-----
0.906707	Bacteroidia	1.240369 Rhizobiaceae	2.161981 Clostridiales	0.812219 Vermucomicrobia	71.03285 unclassified
0.760934	Spirochaetia	1.214455 Chromatiaceae	1.947005 Pseudomonadales	0.760934 Spirochaetes	-----
0.689345	Solibacteres	1.205119 Ectothiorhodospiraceae	1.945917 Cellvibrionales	0.571705 Deinococcus-Thermus	-----
0.677288	Methanomicrobia	1.155342 Comamonadaceae	1.934236 Bacillales	0.513137 Gemmatimonadetes	-----
0.654639	Acidobacteria	1.084591 Hyphomicrobiaceae	1.800839 Sphingomonadales	-----	-----
0.571705	Deinococci	1.018151 Nitrospiraceae	1.728652 Desulfobacteriales	1.368434 Viruses	-----
0.545582	Opilatae	0.996507 Alteromonadaceae	1.706659 Oceanospirillales	1.849 cannot be assigned to a phylum	-----
0.513137	Gemmatimonadetes	0.948907 Geobacteraceae	1.660315 Desulfuromonadales	3.707502 phylum < 0.5% of all reads	-----
-----	-----	0.884561 Rhodocyclaceae	1.380617 Corynebacteriales	-----	-----
1.368434	Viruses	0.836004 Isophaeraceae	1.072701 Xanthomonadales	71.03285 unclassified	-----
10.89422	cannot be assigned to a class	0.832732 Phyllobacteriaceae	1.018151 Nitrospirales	-----	-----
9.647863	class < 0.5% of all reads	0.798947 Haliaeaceae	0.976035 Enterobacteriales	-----	-----
-----	-----	0.796268 Bacillaceae	0.952465 Thiotrichales	-----	-----
71.03285	unclassified	0.793337 Streptomycetaceae	0.884561 Rhodocyclales	-----	-----
-----	-----	0.718818 Polyangiaceae	0.794468 Desulfuromonadales	-----	-----
-----	-----	0.689345 Solibacteraceae	0.793337 Streptomycetales	-----	-----
-----	-----	0.679214 Alcaligenaceae	0.734392 Synechococcales	-----	-----
-----	-----	0.677581 Desulfuromonadaceae	0.732675 Syntrophobacteriales	-----	-----
-----	-----	0.671193 Mycobacteriaceae	0.702282 Micrococcales	-----	-----
-----	-----	0.671385 Cytophagaceae	0.689345 Solibacterales	-----	-----
-----	-----	0.654639 Acidobacteriaceae	0.654639 Acidobacteriales	-----	-----
-----	-----	0.653376 Paenibacillaceae	0.64882 Vibrionales	-----	-----
-----	-----	0.64882 Vibrionaceae	0.640112 Bacteroidales	-----	-----
-----	-----	0.610012 Methylococcales	0.610012 Methylococcales	-----	-----
-----	-----	0.606369 Xanthomonadaceae	0.583093 Bacteroidetes Order II. Incertae sedis	-----	-----
-----	-----	0.603858 Desulfuromonadaceae	0.580664 Oscillatoriaceae	-----	-----
-----	-----	0.595945 Acetabacteraceae	0.570408 Thermoanaerobacteriales	-----	-----
-----	-----	0.585939 Clostridiaceae	0.520463 Pseudonocardiales	-----	-----
-----	-----	0.583093 Rhodothermaceae	0.51816 Nitrosomonadales	-----	-----
-----	-----	0.553327 Oxalobacteraceae	0.513137 Gemmatimonadales	-----	-----
-----	-----	0.521426 Colwelliaceae	0.511504 Spirochaetales	-----	-----
-----	-----	0.520463 Pseudonocardiaceae	0.50715 Caulobacteriales	-----	-----
-----	-----	0.51816 Nitrosomonadaceae	-----	-----	-----
-----	-----	0.513137 Gemmatimonadaceae	1.368434 Viruses	-----	-----
-----	-----	0.507317 Cellvibrionaceae	8.128674 cannot be assigned to a order	-----	-----
-----	-----	0.50715 Caulobacteraceae	16.76881 order < 0.5% of all reads	-----	-----
-----	-----	0.50514 Myxococcaleae	-----	-----	-----
-----	-----	0.501289 Syntrophaceae	71.03285 unclassified	-----	-----
-----	-----	0.500619 Piscirickettsiaceae	-----	-----	-----
-----	-----	-----	-----	-----	-----
1.368434	Viruses	-----	-----	-----	-----
9.866397	cannot be assigned to a family	-----	-----	-----	-----
34.65817	family < 0.5% of all reads	-----	-----	-----	-----
-----	-----	-----	-----	-----	-----
71.03285	unclassified	-----	-----	-----	-----