

## **Archaeometrical results related to Neolithic amphibolite stone implements from Northeast Hungary**

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**Abstract:** 28 amphibolite Neolithic polished stone implements deriving from different archaeological localities and cultures in Northeast Hungary (Borsod-Abaúj-Zemplén County) were archaeometrically analysed by mainly non-destructive methods (MS, EDS/SEM, PGAA). Bulk chemistry of the samples showing subalkali characteristics. The amphibolite polished stone tools were divided into two groups based on their mineral components and metamorphic evolution. A single Ca-amphibole approach was used to calculate peak P-T conditions to determine a thermobarometric model for the amphibolite implements. Data of the studied samples were compared to those of the nearest amphibolite outcrops in Gemericum, Veporicum, Tatricum and Zemplinikum (Slovakia). The Variscan P-T loop covered the thermobarometric data of the analysed stone implements and the amphibolite outcrops. The source areas are assumed to be these fields and/or the crossing riverbeds flowing through them to Borsod-Abaúj-Zemplén County, the archaeological collecting territory of the amphibolite stone axes.

**Keywords:** amphibolite, polished stone tools, provenance field, Neolithic, Hungary

**Mineral abbreviations:** *Amp*: amphibole; *Mg-hb*: magnesio-hornblende; *Mg-f-hb*: magnesio-ferri-hornblende; *Fhb*: Ferro-hornblende; *Act*: actinolite; *Fact*: ferro-actinolite; *Sdg*: sadanagaite; *Fsdg*: ferro-sadanagaite; *Tsch*: tschermakite; *Fe2-fts*: Ferro-ferri-tschermakite; *fts*: Ferro-tschermakite; *f3-ts*: Ferri-Tschermakite; *Prg*: pargasite; *Fprg*: ferro-pargasite; *aug*: Augite; *Di*: diopside; *Cpx*: clinopyroxene; *Ab*: albite; *Olg*: oligoclase; *And*: andesine; *Lab*: labradorite; *Byt*: bytownite; *Pl*: plagioclase; *Kfs*: K-feldspar; *Ep*: epidote; *Czo*: clinozoisite; *Chl*: chlorite; *Bi*: biotite; *Ms*: muscovite; *Ph*: phengite; *Qtz*: quartz; *Ttn*: titanite; *Ilm*: ilmenite; *Mgt*: magnetite; *Ap*: hydroxylapatite; *Py*: pyrite; *Cal*: calcite. *Jrs*: jarosite.

### **1. Introduction**

In this paper we present the results of an archaeometrical research program on the polished stone implement collection of the Herman Ottó Museum in Hungary which started in 2014. The collecting area covers Borsod-Abaúj-Zemplén County in Northeast Hungary. The museum owns approximately 500 Neolithic polished stone implements, most of which are metabasites or more specifically in some cases blueschists (Kereskényi et al., 2018), but a significant number of volcanic and sedimentary stone implements are possessed as well according to our observations. Until now 28 tools have been shown to be made of amphibolite. Stone tools which are reached and characterized by the amphibolite facies assemblage; hornblende and plagioclase feldspar, were classified as amphibolite. Contact metamorphosed metabasites are not included among these lithotypes.

The archaeological localities where these amphibolite polished stone tools were found are shown in Fig. 1. Most of the findings have an archaeological context, but 10 of the axes were stray finds (Table 1). Borsod/Derékegyháza (Edelény) and Szerencs/Taktaföldvár are the sites with the two largest sets of finds. The Borsod/Derékegyháza (Edelény) locality belongs to the Middle Neolithic Bükk culture (Kalicz & Makkay, 1977). The Szerencs/Taktaföldvár locality represents the late Neolithic Tisza culture (Selján, 2005). Tiszadorogma (Kalicz and Makkay, 1977) and Tiszavalk/Kenderföldek are related to the Alföld Linear Pottery culture (Csengeri, 2013). Findings from Szirmabesenyő and its vicinity (Kalicz and Makkay, 1977) and Miskolc/Aldi2 are linked to the Bükk or Alföld Linear Pottery cultures (Csengeri, 2011). The Hejőkürt/Lidl locality is related to the early Tiszadob culture (Csengeri, 2015). The Emőd locality belongs to one or more of the Alföld Linear Pottery, Tiszadob, Tiszadob-Bükk or Bükk cultures according to the related ceramics (Csengeri, 2013) (Table 1).

As regards the raw materials, amphibolite and amphibolite-greenschist stone tools are abundant implements in the Neolithic and Aeneolithic periods (Hovorka et al., 2001, Méres et al., 2001, Přichystal, 2013), being described as common and often used raw materials due to flexibility and elasticity of the ready-made items because of the needle-like of amphiboles (Hovorka et al., 2001).

The most probable amphibolite source localities are not too far from the archaeological sites in the Western Carpathians. The Western Carpathians are defined by three major geological units: Veporicum, Gemicum and Tatricum (Plašienka et al., 1997). In the Veporicum, amphibolites are not dominant, but some large bodies are exposed in the Čierna Hora and the Branisko Mountains (Faryad, 1999, Faryad et al., 1999, 2005, Faryad & Jacko, 2002). Furthermore, the LAC (Leptynit-Amphibolite Complex) outcrops in Veľký Zelený stream and in Čierny Balog (Putiš et al., 1997). The Gneiss-Amphibolite Complex (GAC) in the Gemicum has two main outcrops in Rudňany (Radvanec et al., 2017) and in Klátov (Faryad, 1990, 1999, Faryad and Spišiak, 1999). The Ochtiná group is an exposure of amphibolite rocks at the contact zone of Veporicum and Gemicum (Novotná et al., 2015). There are only distant, minor occurrences in the Tatricum (Ivan et al., 2001). The most extensive outcrops are in the Tribeč Mountains (Faryad, 1999) and Little Carpathians (Ivan et al., 2001). Zemplinicum is the closest geological unit of the

Carpathians to the archaeological localities; but here the amount of amphibolites is much less than in the above regions (Faryad, 1995, Faryad and Vozárová, 1997). The vicinities of major riverbeds of Slovakia which cross Hungary cannot be excluded as possible provenance fields (Fig. 2).

In the southwestern part of Slovakia, at the Bajč archaeological site, a large number of polished stone tools were excavated with various types of amphibolite and greenschist stone axes. According to Hovorka and Cheben (1997), the possible provenance field of these amphibolites can be the Little Carpathians, Slovak Ore Mountains, Tribeč Mountains, the eastern part of the Bohemian Massif, Eastern/Northern Alps and also pebbles of the Hron and Danube riverbeds.

In this research we provide petrological and mineralogical data of the Neolithic amphibolite implements and compare them with the former published data of nearby Slovakian amphibolite source localities and specification of the most probable provenance of the raw material.

## **2. Analytical techniques and materials**

In addition to macroscopic and stereomicroscopic descriptions, the magnetic susceptibilities of all the axes were measured by a KT-5 Kappameter, using thickness correction (Bradák et al., 2009, Szakmány et al., 2011b, Kereskényi et al., 2015 a, b.).

Non-destructive methods were applied since most of the implements are intact. While destructive analyses were performed only on some broken stone implements (Table 1).

The chemical elements in the bulk material of 24 polished archaeological stone objects (Table 1) have been measured by the prompt-gamma activation analysis (PGAA) system of the Budapest Neutron Centre. The present state of the PGAA facility in Budapest was described by Szentmiklósi et al. (2010) in detail. During the analysis, the entire objects were placed in a horizontal beam of cold neutrons guided from the Reactor. The flux of the cold neutron beam equivalent to the thermal beam was  $1.2 \times 10^8$  neutrons  $\text{cm}^{-2} \text{s}^{-1}$ . During the experiments, the beam was setup to 24  $\text{mm}^2$  and 44  $\text{mm}^2$  size cross-section. With such beam collimation, the obtained count rate proved to be optimal from spectroscopic points. Since neutrons can travel into the deepest parts of the investigated object, the obtained composition data represents the bulk. Promptly after the neutron absorption, characteristic gamma radiation is emitted, which was detected with a HPGe-BGO detector system. The spectra were collected with a 36,000-channel analyser. To identify characteristic gamma lines, our PGAA library has been used. The quantitative elemental compositions have been calculated based on the  $k_0$ -method (Révay, 2009). In silicates, the PGAA method allows to determine all the major elements and some geochemically important traces (such as B, Cl, Sc, V, Cr, Co, Nd, Sm and Gd), non-destructively, see Kasztovszky et al. (2008), Szakmány and Kasztovszky (2004), and Szakmány et al. (2011a). In these experiments, the spectrum acquisition has been performed for 1200 s -3000 s, to gain enough counts in the relevant spectrum peaks to produce statistically reliable results.

Quantitative chemical spot analyses on the mineral phases were performed applying a JEOL JXA 8600 Superprobe electron-microprobe in energy-dispersive mode (EDS/SEM) at the Institute of Mineralogy and Geology, University of Miskolc. The operating conditions were as follows: accelerating voltage 15 kV, probe current 20 nA, and a 1-5  $\mu\text{m}$  beam diameter. The intact tools were analysed on their original surfaces using the method of Bendő et al. (2013), while polished surface sections were made from the broken implements for their analyses (Table 1).

### 3. Results

#### 3.1. Macroscopic description and magnetic susceptibility

Most of the stone implements are intact and having traces of use. They are 4.0-11.5 cm long, 1.0-6.5 cm wide and 0.5-4.0 cm thick. Regarding their archaeological typology, most of them are flat axes and chisels, but axes D21, D32, D39 have shoe-last form, D38 and D40 are shaft drilled stone axes, and D48 is a tongue-shaped axe.

Macroscopically the tools are fine-grained, compact, dominantly foliated, having grey, greyish black, greyish green and dark brown colour. Green and brown patches and bands can be recognized by naked eye.

Magnetic susceptibility values show a homogenous distribution from  $0.31\text{-}0.85 \cdot 10^{-3}$  SI. 4 samples appear from  $1.44\text{-}1.77 \cdot 10^{-3}$  SI and sporadically a continuous upward from 6.47 to  $46.7 \cdot 10^{-3}$  SI (Table 1).

#### 3.2. Bulk rock chemistry

Bulk compositional data analysed by PGAA for the amphibolite implements are listed in Table 2. The tools are characterized by  $\text{SiO}_2$  content ranging from 46.55 to 59.00 wt%. Sample D12 has an elevated  $\text{TiO}_2$  content= 3.70 wt% and remarkably differs in some of the trace elements from the other analysed samples. Representing the data in a TAS diagram, the analysed samples plot in the basalt, trachybasalt, basaltic andesite and andesite fields, having subalkaline characteristics since the  $\text{Na}_2\text{O}+\text{K}_2\text{O}$  content varies between 0.92-5.17 wt%. (Fig. 3). In the AFM diagram most of the samples show tholeiitic affinity (Fig. 4).

#### 3.3. Petrography and mineral chemistry

A total of 22 polished stone tools were analysed by EDS/SEM. Eight of them are intact stone axes and were analysed by the "original surface" method (Bendő et al., 2013). In these cases, the zoned mineralogical features (e.g. the existence of tiny or scarce phases) the zonation of the amphiboles could not always be perceived.

Most of the stone implements are fine-grained, but some stone tools (D08, D12, D13, D15, D18 and D30) are coarse-grained (Figs. 5 A—D, 6 A—C). Most of the axes are foliated (Figs. 7 A—D) except B04, B05, B12, D17, D18, D33, D45 and D47 samples.

The amphibolites consist mainly of Ca-amphibole (Mg-hb, Act, Tsch, Prg, Sdg), plagioclase, epidote, clinozoisite, chlorite and quartz. Some samples contain clinopyroxene, K-feldspar, biotite in minor quantity. Titanite and ilmenite are abundant accessory minerals (Table 3).

According to the textural relations, chemical compositions and the assemblage of coexisting phases, two varieties of the detected stone implements can be distinguished. The first variety (*Group1*) includes six samples and it consists of Ca-amphiboles with increasing Al(tot) content from cores to rims. Actinolite and magnesio-hornblende are preserved in the cores with Al(tot) 0.14-2.84 apfu (atoms per formula unit), while at the rims magnesio-hornblende, pargasite and tschermakite have crystallized with Al(tot) 0.73-2.83 apfu (Fig. 8).

The second group (*Group2*) includes 16 pieces of stone tools and their amphibole content corresponds to hornblende, pargasite, tschermakite and sadanagaite in the core with actinolite at the rims, just the opposite of the order in *Group1* tools. *Group2* was divided into four subgroups based on mineral associations:

- *Group2a* includes most of the implements in which no K-feldspar or biotite were detected (B12, D02, D04, D08, D18, D21, D32, D47),
- K-feldspar±biotite-bearing amphibolites (*Group2b*: B04, B05, C02, D48),
- Clinopyroxene±biotite-bearing amphibolites (*Group2c*: D15, D34, D45),
- Biotite-bearing amphibolite without K-feldspar and clinopyroxene (*Group2d*: D17).

### 3.3.1. Amphibole

The ACES Excel spreadsheet was used to calculate cation numbers of amphiboles. Since the  $\text{Fe}^{3+}/\text{Fe}(\text{tot})$  ratio cannot be measured by electron-microprobe, the smallest maximum deviations criteria were applied for adjusting the valences of Fe by automatically selecting one or more of four cation normalization schemes: sum Si to Ca (+Li) = 15; sum Si to Mg (+Li) = 13; sum Si to Na = 15; and sum Si to K = 16 apfu (Locock, 2014). Applied amphibole nomenclature corresponds to the recent IMA rules (Hawthorne et al., 2012).

#### *Group1*

In each detected commonly zoned Ca-amphibole of the *Group1* variety, the  $^{\text{B}}\text{Na}/(\text{Ca}+\text{Na})$  varies in a narrow range of 0.00-0.19 both in the cores and rims. Overlapping  $^{\text{A}}\text{Na}$  and  $X_{\text{Mg}}$  values characterize amphiboles in the cores and at the rims.  $^{\text{A}}\text{Na}$  ranges over 0.00-0.65 apfu and  $X_{\text{Mg}}$  is fairly homogenous with values of 0.49-0.79.

Actinolite with  $^{IV}Al$  0.12-0.44,  $^{VI}Al$  0.01-0.65 and  $Al(tot)$  0.14-1.03 apfu was detected in the cores of all samples except D30. Magnesian-hornblende occurs in the cores of the D05, D30 and D31 samples with  $^{IV}Al$  0.33-1.13,  $^{VI}Al$  0.50-1.71,  $Al(tot)$  0.44-2.84 apfu. In the D05, D12, D13 and D33 samples, magnesian-hornblende–magnesian-ferri-hornblende with  $^{IV}Al$  0.28-1.12,  $^{VI}Al$  0.60-1.64 and  $Al(tot)$  0.73-2.75 apfu makes a rim of actinolite-magnesian-hornblende (Fig. 5 B, D). Tschermakite was detected at the rims of earlier crystallized amphiboles in D05, D30 and D31 samples with  $^{IV}Al$  1.10-1.48,  $^{VI}Al$  1.15-1.55 and  $Al(tot)$  2.50-2.63. Pargasite with  $^{IV}Al$  0.72-1.13,  $^{VI}Al$  1.52-1.70,  $Al(tot)$  2.31-2.83 apfu was described from the rims on the actinolite–magnesian-hornblende in the D30 and D31 samples.

### *Group2*

In *Group2* the  $^{A}Na$  values are 0.00-0.60 apfu,  $^{B}Na/(Ca+Na)$  ranges from 0.00 to 0.19 and  $X_{Mg}$  is 0.30-0.79, values which overlap in the cores and rims.

Generally, in *Group2a* (Figs. 6 G–I, Fig. 7 D) most of the samples have a magnesian-hornblende core with  $^{IV}Al$  0.01-0.93,  $^{VI}Al$  0.07-1.79,  $Al(tot)$  0.84-2.29 apfu, and actinolite–hornblende was observed with  $^{IV}Al$  0.18-0.54,  $^{VI}Al$  0.00-0.50,  $Al(tot)$  0.23-0.89 apfu at the rims. Ferri-tschermakite and ferro-ferri-tschermakite were determined for the core of sample D08 (Fig. 7 C–D) with  $Al(tot)$  2.13 apfu. In the core of the D47 stone implement pargasite–ferropargasite prevail with  $Al(tot)$  2.64-3.31 apfu in addition to magnesian-hornblende. In the core of the B12 tool pargasite, ferropargasite, sadanagaite and ferrosadanagaite (Fig. 6–I) are present with an increased  $Al(tot)$  from 2.97 to 3.45 apfu.

In *Group2b* most of the samples have magnesian-hornblende–ferro-hornblende cores, although in D48 pargasite and ferro-pargasite were detected also. The amphibole cores have ranges of  $^{IV}Al$  0.41-0.75,  $^{VI}Al$  0.39-1.40,  $Al(tot)$  0.86-2.00 apfu, while at the actinolite rims a significant decrease shows up, with ranges of  $^{IV}Al$  0.23-0.48,  $^{VI}Al$  0-0.42,  $Al(tot)$  0.36-0.66 apfu.

In *Group2c* magnesian-hornblende–ferro-hornblende cores were recorded having  $^{IV}Al$  0.30-0.92,  $^{VI}Al$  0.60-1.30,  $Al(tot)$  0.83-2.21 apfu and at the rims magnesian-hornblende–actinolite were observed with  $^{IV}Al$  0.07-0.35,  $^{VI}Al$  0.27-0.64,  $Al(tot)$  0.52-1.03 apfu.

*Group2d* consists of only the D17 implement, in which magnesian-hornblende–ferropargasite exists in the core with  $^{IV}Al$  0.35-1.03,  $^{VI}Al$  0.39-1.68  $Al(tot)$  0.86-2.73 apfu and with magnesian-hornblende–actinolite rimming the earlier crystallized amphiboles with  $^{IV}Al$  0.13-0.42,  $^{VI}Al$  0.05-0.43,  $Al(tot)$  0.18-0.76 apfu.

Chemical compositions of the calcic amphiboles are presented in Tables 4, 5.

The detected amphiboles of the amphibolites of *Group1* and *Group2* lie in the actinolite, magnesian-hornblende, tschermakite, pargasite, and sadanagaite fields (Figs. 8, 9) in the  $^{A}(Na+K+2Ca)$  vs  $^{C}(Al+Fe^{3+}+2Ti)$  diagrams.

### 3.3.2 Pyroxenes

The pyroxene formula was calculated on the basis of 6 oxygens. The FeO/Fe<sub>2</sub>O<sub>3</sub> ratio was estimated assuming a total number of cations of 4. Normalized pyroxene end members found in amphibolites are plotted in the En-Fs-Wo diagram (Fig. 10) of Morimoto (1989). *Augite* forms hypidioblasts and xenoblasts as a relict phase next to actinolite and hornblende in small amounts in the D34 (En<sub>0.37</sub> Fs<sub>0.28</sub> Wo<sub>0.34</sub>) and D45 samples (En<sub>0.28-0.39</sub> Fs<sub>0.25-0.50</sub> Wo<sub>0.23-0.42</sub>). *Diopside* is observed as a relict phase in the D15 sample. Its chemical composition varies in a narrow range En<sub>0.38-0.39</sub> Fs<sub>0.15</sub> Wo<sub>0.47</sub>.

### 3.3.3 Feldspars

Plagioclase shows various chemical compositions from albite to bytownite in the amphibolites (Fig. 11).

In the D31 sample of *Group1*, the textural relations or core to rim zoning of the plagioclases was not well distinguished due to the limitations of the “original surface” method. Its composition varies from oligoclase to andesine with An<sub>23.0-30.0</sub>. In the D05 and D13 stone axes, a labradorite composition with a narrow range of An<sub>54.3-61.9</sub> is present in the core (PI1), presumably as a relict magmatic phase, while at the rim (PI2) it is andesine with An<sub>47.1-49.9</sub> (Table 3). In the D12, D30 and D33 stone implements, albite (An<sub>0.0-6.7</sub>) is the pre-kinematic plagioclase (PI1) (Table 3). In samples D30 and D33, oligoclase (An<sub>15.9-29.6</sub>) (Fig. 11), and in the D12 sample andesine (An<sub>37.9-44.5</sub>) are the post-kinematic phases (PI2) (Table 3).

In samples D04, D08, D45 and D48 of *Group2*, it was not possible to define the textural relations of plagioclases with andesine (An<sub>40.3-49.5</sub>) and labradorite (An<sub>50.9-64.8</sub>) compositions. In the stone implements C02 and D15, plagioclase of labradorite and bytownite composition (An<sub>63.1-89.0</sub>) is a relict phase (PI1) while oligoclase-andesine (An<sub>13.4-46.5</sub>) is the typical composition for the rim (PI2). For the other samples of *Group2*, the relict phases have values of An<sub>24.4-65.7</sub> for PI1 and the post-kinematic minerals were albite, described with An<sub>0.0-9.1</sub> content as PI2 (Table 3, Figs. 6—E, G, I). Potassic feldspar, as a relict phase is detected in *Group2b* (Figs. 6—E, G).

### 3.3.4 Other minerals

*Epidote/clinozoisite* is present in 9 samples (Table 6). *Epidote/clinozoisite* is mainly zoned, where Fe-content increases from the core to rim. In the D02 and D04 tools the amount of epidote and clinozoisite reaches 30 vol% (Fig. 7—B). *Biotite* was identified in tools C02, D15, and D17 (*Group2*). In samples C02 and D17, biotite is partially replaced by chlorite. Biotite occurs up to 50 µm in diameter. Its chemical compositions are listed in Table 7. *Ilmenite* forms idioblastic and hypidioblastic grains. In the samples B12 (Fig. 6—I), D15 and D18 (Fig. 6—C) ilmenite is partly replaced by titanite and contains magnetite and Ti-bearing magnetite as inclusions. MnO content in ilmenite from samples B12, D05 and D18 reaches 3.00-3.97 wt%. In sample D18 the size of ilmenite tablets reaches 600 µm in diameter. *Titanite* is a common accessory mineral and was found in nearly all the tools except B05, D13, D31 and D33. Its Al<sub>2</sub>O<sub>3</sub> content is

typically high, varying in the range of 1.26-5.79 wt%. The highest value was recorded in the D05 sample but shows elevated values (2.28-4.01 wt%) in the B04, D15, D18 and D45 samples as well.

*Magnetite* is observed in the D04, D18, D21 and D33 samples. In D04 it is distributed along foliations with grain size up to 200  $\mu\text{m}$  (Fig. 7—B). In tool D18, magnetite and Ti-bearing magnetite ( $\text{TiO}_2$ : 4.29 wt%) is recognized as inclusions in Mn-bearing ilmenite (Fig. 6—C). *Chlorite* was observed in five samples. In sample D13, chlorite of clinoclone composition is present at the contact of actinolite and epidote. In tool D02 it was observed at the rim of magnesio-hornblende. In sample D04, chlorite developed inside hematite and both in D02 and D04 samples chlorite was identified as chamosite. The chemical compositions of the chlorites are presented in Table 8.

## 4. Discussion

### 4.1 Thermobarometric estimations

As no garnet was found in the studied samples, the application of thermobarometric methods was limited. Since amphibole is a common mineral present in all amphibolite polished stone tools, the single Ca-amphibole approach of Gerya et al. (1997) was applied to calculate peak pressure and temperature ( $P_{(T_{\text{max}})}-T_{\text{max}}$ ) conditions. The method was modified for  $\text{Fe}^{3+}$  by Zenk & Schulz (2004). Based on their suggestion, the 13CNK method (Locock, 2014) was used to determine the amphibole formulae. In this method, the absolute errors of the calculation are generally reduced to  $\pm 1.2$  kbar and  $\pm 37$   $^{\circ}\text{C}$ .

#### Group1

In sample D33, the magnesio-hornblende in the rim implies  $T_{\text{max}}$  460  $^{\circ}\text{C}$  and  $P_{(T_{\text{max}})}$  2.8 kbar. In the D05 sample, tschermakite which rims magnesio-hornblende crystallized at 540  $^{\circ}\text{C}$  and 5.3 kbar. The D12 and D13 samples have very similar mineral assemblages and their P-T data calculated from magnesio-hornblende in the rim yield  $T_{\text{max}}$  of 560–570  $^{\circ}\text{C}$  and  $P_{(T_{\text{max}})}$  of 5 kbar. Tschermakite in the rim (Table 3) in D30 and D31 samples implies  $T_{\text{max}}$  615–635  $^{\circ}\text{C}$  and 5–6 kbar for  $P_{(T_{\text{max}})}$ . In D30, pargasite which rims magnesio-hornblende yielded  $T_{\text{max}}$  of 615  $^{\circ}\text{C}$  and 5.6 kbar for  $P_{(T_{\text{max}})}$  (Table 9). In the D31 sample, the pargasite rim (Table 3) crystallized at 635  $^{\circ}\text{C}$  and 6.2 kbar and the magnesio-hornblende core (Table 3) yields the same P-T data (Table 9, Fig. 12).

#### Group2

In sample D08 of *Group2a*, tschermakite occurring along with magnesio-hornblende yielded 600  $^{\circ}\text{C}$  and 4.8 kbar in Table 9. In sample B12, the core pargasite recorded 670  $^{\circ}\text{C}$  and 6.5 kbar, the core sadanagaite yields similar data of 680  $^{\circ}\text{C}$  and 7.3 kbar. In the D47 sample, core pargasite crystallization recorded 690  $^{\circ}\text{C}$  and 7.0 kbar. In the D48 sample (*Group2b*), ferro-pargasite crystallization yielded  $T_{\text{max}}$ : 590  $^{\circ}\text{C}$  and  $P_{(T_{\text{max}})}$  4.6 kbar. In sample D17 of *Group2d*, ferro-pargasite rimmed by actinolite reached 610



°C and 5.7 kbar. In the other samples of *Group2*, the peak metamorphism was determined from magnesio-hornblende with  $T_{\max}$  of 450–610 °C and  $P_{(T_{\max})}$  of 2.8–5.7 kbar (Table 9).

Plotting the  $T_{\max}$  versus the  $P_{(T_{\max})}$  values, a continuous transition of simultaneously increasing values in the medium-temperature- and high-temperature fields can be observed (Table 9, Fig. 12). The medium-temperature group presents  $T_{\max} \approx 550$  °C and  $P_{(T_{\max})} \approx 5.3$  kbar. For the high-temperature group, in which tschermakite-, pargasite-bearing amphibolite implements are present, the thermal maximum,  $T_{\max} \approx 600$  °C and  $P_{(T_{\max})}$  5–6 kbar are characteristic. The sadanagaite-bearing sample (B12) and a pargasite-bearing sample (D47) reached the highest peak temperatures. The B12 sample suggests  $T_{\max}$  of 680 °C and  $P_{(T_{\max})}$  of 7.3 kbar and the D47 sample yields  $T_{\max}$  of 690 °C and  $P_{(T_{\max})}$  of 7.0 kbar (Table 9, Fig. 12).

The tools in *Group2* were overprinted by greenschist facies metamorphism as magnesio-hornblende and actinolite surrounding earlier crystallized Ca-amphiboles.

The results of the thermobarometric calculations are listed in Table 9 and plotted in Fig. 12.

Based on the mineral assemblage and textural relationships, two metamorphic events can be recognized in both groups. In *Group1*, the M1 event is featured by the greenschist-amphibolite facies assemblage involving zoned Ca-amphibole with lower Al(tot). The M2 event is represented by higher Al(tot) values, indicating a progressive increase of pressure and temperature from the greenschist-amphibolite boundary to the amphibolite facies. In *Group2* the amphibolite facies M1 event was followed by the M2 event which was a retrograde stage of greenschist metamorphism recorded by the appearance of actinolite rimmed magnesio-hornblende, tschermakite, pargasite or sadanagaite and by the formation of albite, chlorite, epidote and titanite.

#### 4.2. Possible provenance fields

Our petrological and mineralogical data were compared to the amphibolite bodies of Gemericum, Veporicum, Tatricum and Zemplinicum in Slovakia (Fig. 2) (Král' et al., 1997). Since amphibolite is a common and abundant raw material, the largest and the nearest amphibolite bodies are assumed to be the material sources for the investigated amphibolite stone implements. According to Přichystal (2013), stone raw material of polished implements could also have originated from riverbeds, furthermore Török (1996) declares that amphibolite is one of the most common pebble in Slaná- and Hornád-riverbeds; so the major rivers of Slovakia are also plotted on the map (Fig. 2).

The bulk chemistries of Gemericum, Veporicum and Tatricum amphibolites are very similar to each other and also to that of our samples, having subalkaline, tholeiitic character (Fig. 3) (Bajanik & Hovorka, 1981, Hovorka et al., 1993, Ivan et al., 2001, Faryad et al., 2005, Ivan & Méres, 2015). On the AFM diagram, the distribution of the data of the samples and the data of the possible provenance fields overlap each other (Fig. 4) (Bajanik & Hovorka, 1981, Hovorka et al., 1993, Ivan et al., 2001, Faryad et al., 2005, Ivan & Méres, 2015). Using only the major element composition of bulk chemistry, the provenance

field cannot be specified. Magnesio-hornblende and actinolite are very common minerals in amphibolites. Based only on their presence, selecting the provenance field is nearly impossible. Regarding the fact that systematic archaeometrical work on amphibolite implements has not been done which we could compare to, we have used various diagnostic minerals as well as the typical thermobarometric P-T data of the amphiboles for comparison to the data of the possible source regions.

All the above lithotectonic units where amphibolite occurs show multiphase (Variscan and Alpine) metamorphic evolutions with various intensities. Some localities do not show traces of the Alpine overprint at all, while others did not preserve relict textures of the Variscan metamorphic event (Bezák et al., 1993, Hovorka et al., 1993). Prograde textured amphibolites were preserved in some outcrops in the Ochtiná group (Novotná et al., 2015) and some places in the Veporicum and Tatricum (Krist et al., 1992). Most of the amphibolite occurrences show a retrograde overprint as well. The two subsequent metamorphic events can be characterized by significantly different metamorphic gradients. For the Variscan event 40 °C/km is typical, while the Alpine event suggests 10 °C/km (Bezák et al., 1993). The results of the thermobarometric calculations (according to Gerya et al., 1997, Zenk and Schulz 2004) (Fig. 12) show that all the studied samples fit well to the Variscan metamorphic gradient (Bezák et al., 1993); no trace of a high pressure, low-medium temperature overprint is observable. The Variscan metamorphic gradient (Bezák et al., 1993) (Fig. 12) and the Variscan P-T loop (Putiš et al., 1997) (Fig. 13) cover all the available fields that are considered here as possible source regions for the present group of amphibolite polished implements. The thermobarometric data of some possible provenance fields and the samples fit each other well as shown in Fig. 13. Regarding Tatricum, the Little Carpathians can be considered as a possible source region, while the thermobarometric data of Tribeč Mountains exclude these fields as source regions (Fig. 13).

In *Group1* the tschermakite- and pargasite-bearing prograde textured samples reached the highest metamorphic grade. In *Group1* the estimated thermobarometric data mostly range 540-635 °C and 5-6.2 kbar. But the D33 sample has a lower  $T_{max}$  460 °C and  $P_{(Tmax)}$  2.8 kbar. Comparing to the prograde textured amphibolites from the Ochtiná group (Novotná et al., 2015) and sporadically from Veporicum and Tatricum (Krist et al., 1992) the majority of these *Group1* samples very likely originated from one of these areas (Table 10). According to the P-T estimation of D30 and D31, they correlate obviously to GAC-Klátov. D05, D12, D13 samples fit well to the Ochtiná group, but the P-T data of the Branisko, Čierna Hora, Čierny Balog, Little Carpathians and GAC-Klátov territories also match well. The D33 implement can be originated not only from Ochtiná group but Čierna Hora too (Fig. 13) (Table 10).

In the B12 sample (*Group2a*), besides the ordinary amphiboles, the uncommon sadanagaite and ferro-sadanagaite, as well as pargasite and ferro-pargasite also appear. The compositional range of these rare amphiboles is equivalent to pargasite and tschermakite, respectively, according to the previous nomenclature of amphiboles (Leake et al., 1997). Tschermakite, pargasite and their ferroan counterparts

were found in several samples: D05, D30, D31 (*Group1*), B12, D08, D47 (*Group2a*), D45 (*Group2c*), D48 (*Group2b*), D17 (*Group2d*). These samples have reached the highest metamorphic temperature (Fig. 12, Table 9). Results on pargasite-bearing Variscan amphibolites were published by Radvanec et al. (2017) in the GAC (Gneiss-Amphibolite Complex), Klátov area. Tschermakite-bearing amphibolites occur in several localities: the Ochtiná group (Novotná et al., 2015), Čierny Balog (Putiš et al. 1997), Čierna Hora (Faryad & Jacko, 2002) and also the Little Carpathians (Ivan et al., 2001). The calculated thermobarometric data of B12 and D47 samples would agree with the material originated from Zemplinicum or GAC-Klátov regions. The estimated P-T data and the mineral assemblage of D17 and D48 pargasite-bearing implements fit well to GAC-Klátov, but the Ochtiná group, Branisko Mountains and even the distant Little Carpathians cannot be excluded. The D08 (*Group2*) sample matches well to the Ochtiná group, Branisko or the Little Carpathians, but Stražovské Mountains can be considered also (Fig. 13) (Table 10).

Two epidote-rich amphibolite implements, D02 and D04 samples (*Group2a*) can be originated from Ochtiná group or GAC-Klátov. The calculated  $T_{\max} \approx 500$  °C and  $P_{(T_{\max})} \approx 3.7$  kbar are similar for the two stone implements (Fig. 13).

The D18, D21 and D32 samples from *Group2a*, have a very common amphibolite mineral assemblage without any diagnostic minerals and their estimated P-T data overlap with earlier published thermobarometric data of numerous possible provenance fields: Branisko Mountains, Čierna Hora, Ochtiná group, GAC-Klátov or even with the Little Carpathians. (Fig. 13).

Potassic feldspar is present as a relict phase in *Group2b*. Hence this variety may be correlated to Čierna Hora and Branisko Mountains (Faryad & Jacko, 2002) amphibolites as these rocks contain potassic feldspar and Mn-bearing ilmenite as accessory phases. Furthermore, plagioclase with a maximum  $An_{43}$  content was mentioned from here. The C02 sample contains bytownite as a relict phase also. The plagioclase compositions of these three stone tools and their mineral assemblages are rather similar to the amphibolites of Čierna Hora. Amphibolites from Čierna Hora have a tholeiitic basalt and basaltic andesite composition (Faryad et al., 2005), similar to the above mentioned three implements. The well-matching thermobarometry data of *Group2b* and Čierna Hora confirm the correlation (Figs. 6—G, E, 14). In the case of the D48 sample the estimated P-T values indicate Branisko Mountain as a possible provenance field, but the Ochtiná group and GAC-Klátov cannot be excluded either (Table 10).

*Group2c* corresponds to clinopyroxene- and biotite-bearing amphibolites described from Gemericum (Faryad, 1997, Hovorka & Spišiak, 1997). The mineral assemblage and the estimated thermobarometry data of *Group2c* indicate GAC-Klátov as a possible provenance field, but Branisko Mountain, Čierna Hora and Little Carpathians cannot be excluded because of the matching P-T data. Biotite-bearing amphibolites were earlier investigated from the above-mentioned territories (Vozárová & Faryad, 1997, Ivan et al., 2001) (Fig. 13).

Higher magnetic susceptibility values were measured from some samples of *Group1* (D33) and *Group2* (B12, C02, D04, D08, D18), although the presence of magnetite was not always supported by the EDS/SEM observations. Greenstones and amphibolites with high magnetite content were described from the Little Carpathians (Ivan et al., 2001) and the Slovak Ore Mountains (Faryad & Peterec, 1987). Therefore, the raw material of these polished stone tools might originate from these territories (Table 10).

The schematic map of amphibolite outcrops and the major rivers shows that the Hornád- and Slaná rivers cross the Gemericum, Veporicum and Zemplinicum regions and flow to Borsod-Abaúj-Zemplén county, the archaeological collecting territory of the investigated amphibolite implements (Fig. 2). This confirms these rivers as deliverers of raw materials from the source areas.

## Conclusions

Based on the mineralogical assemblages and textural relations of the amphibolite stone implements, two groups have been distinguished: *Group1* and *Group2*. Six samples preserving the features of prograde metamorphism were assigned to *Group1*. *Group2* contains most of the samples which are overprinted by retrograde metamorphism. The results of thermobarometric calculations of the investigated polished stone tools were compared to the earlier published thermobarometric data for the amphibolite field samples, as expressed in the Variscan metamorphic gradient (Fig. 12) and in Variscan P-T loop (Fig. 13).

Metamorphic evolution, thermobarometric data, petrographic-, and mineralogical observations have been taken into account in the classification of the possible provenance fields of the amphibolite polished stone implements. Even the distance of the source area and the archaeological localities is not a negligible factor because amphibolite is a diverse, common and a widespread rock in Slovakia. According to this fact, we suppose that the nearby amphibolite outcrops may be considered particularly (Table 10) For the *Group1* and for the D02, D04 samples of *Group2a* Ochtina group and GAC Klátov is the most likely provenance fields, but for the *Group1* Veporicum and Tatricum cannot be excluded either. In the case of D04 Little Carpathians is a presumptive area. The B12 and D47 samples (*Group2a*) have reached the highest metamorphic grade, they may originate from GAC-Klátov or the Zemplinicum unit. D08, D18, D21 and D32 samples of *Group2a* may have derived from several fields of Gemericum, Veporicum, Tatricum. The most possible source areas of the *Group2b* are Čierna Hora and Branisko, but in the case of C02 sample Tatricum cannot be excluded either, furthermore the D48 implement can be originated from GAC-Klátov or Ochtina group too. The samples of *Group2c* and *Group2d* most likely stem from GAC Klátov, but the possible provenance field of *Group2c* may be Branisko, Čierna Hora or the Little Carpathians as well (Table 10).

To identify the provenance field at site level in the case of amphibolite archaeological findings, cannot be solved without specific minerals or signs except for individual cases. However, because of the overlap of the P-T values of the provenance fields and the uncertainty of the estimated P-T values for the archaeological samples, it is not possible to unambiguously assign any provenance field to any single sample. Based on the mineral assemblages, textural features and thermobarometry estimations it can be concluded that the provenance fields of our samples could be the Gemicum, Veporicum, Tatricum and the Zemplanicum (Figs. 13, 14) (Table 10).

As amphibolite is a widespread and common rock in Slovakia no correlation has been detected between lithotypes and the typology of the artefacts, furthermore nor in their chronological and cultural aspects.

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Fig. 1: Schematic map showing archaeological localities of amphibolite stone tools that are deposited at the Herman Ottó Museum. Abbreviation: BAZ: Borsod-Abaúj-Zemplén.

Fig. 2: Geographical setting of the main lithotectonic units containing amphibolite bodies of the Central-Western Carpathians and depicting the principal rivers of Slovakia. The schematic map is modified after Král' et al. (1997). Abbreviations on the map: B & CH: Branisko and Čierna Hora Mountains, G: Gemericum, V: Veporicum, LT: Low Tatras, HT: High Tatras, LF: Low Fatras, HF: High Fatras, S: Strážovské Mountains, Z: Žiar Mountains, I: Inovec Mountains, T: Tribeč Mountains, LC: Little Carpathians.

Fig. 3: Amphibolite implements from the Herman Ottó Museum and their potential provenances plotted in the Total Alkali- Silica (TAS) diagram.

Fig. 4: AFM diagram of the investigated amphibolite implements and their possible fields of amphibolite localities plotted in AFM (Irvine and Baragar, 1971) diagram. A = alkali ( $\text{Na}_2\text{O}+\text{K}_2\text{O}$ ), F = FeO (total), M = MgO. IAT = island arc tholeiite, MORB = mid-ocean ridge basalt.

Fig. 5: A—B Macroscopic and BSE images of the coarse-grained sample D12 (*Group1a*). C—D Macroscopic and BSE images of the coarse-grained sample D13 (*Group1a*). E—F Macroscopic and BSE images of sample D30 (analysed by original surface) (*Group1b*).

Fig. 6: A—B Macroscopic and BSE images of the coarse-grained, ilmenite-rich sample D18 (*Group2a*). C— Ilmenite is partly replaced by titanite with magnetite inclusions in the BSE image of the sample D18. D— E Macroscopic and BSE images of the sample B04 (*Group2b*). F—G Macroscopic and BSE images of the sample B05 (*Group2b*). H—I Macroscopic and BSE images of the sample B12 (*Group2a*).

Fig. 7: A—B Macroscopic and BSE images of the foliated, epidote-rich D04 sample (*Group2a*). C—D Macroscopic and BSE images of the well-foliated D08 sample (*Group2a*).

Fig. 8: Chemical composition of calcic amphiboles of *Group1* amphibolites plotted in the  $^{\text{C}}(\text{Al}+\text{Fe}^{3+}+2\text{Ti})$  vs.  $^{\text{A}}(\text{Na}+\text{K}+2\text{Ca})$  diagram. Arrows indicate the phase changes from core to rim.

Fig. 9: Chemical composition of calcic amphiboles of *Group2* and their potential provenances plotted in the  $^{\text{C}}(\text{Al}+\text{Fe}^{3+}+2\text{Ti})$  vs.  $^{\text{A}}(\text{Na}+\text{K}+2\text{Ca})$  diagram. The arrows indicate the phase changes from core to rim. Circles signify the samples analysed by the “surface method”.

Fig. 10: Clinopyroxene compositions of D15, D34, D45 samples plotted in the En-Fs-Wo ternary diagram (Morimoto, 1989).

Fig. 11: Compositions of feldspars plotted in Or-Ab-An ternary diagram

Fig. 12: P-T diagram showing thermobarometric estimated calculation results  $T_{\text{max}}$  and  $P(T_{\text{max}})$  of the amphibolite polished stone tools plotted with the Variscan (V) and Alpine (A) metamorphic gradient (Bezák et al., 1993). Red coloured symbols signify the *Group1* samples, while the other samples are in *Group2*.

Fig. 13: Thermobarometric data of amphibolite implements and of field samples from Gemericum, Veporicum, Tatricum and Zemplinicum. The grey loop indicates the estimated Variscan P-T loop (Putiš et al., 1997). Red coloured symbols signify the *Group1* samples, while the other samples are all in *Group2*. Abbreviations: GAC-K: Gneiss Amphibolite Complex, Klátov; C-L: Čierna Hora, lower unit; B: Branisko Mountains; CB: Čieny Balog; Ze: Zemplinicum; T: Tribeč Mountains; Sv: Stražovské Mountains; LC: Little Carpathians.

Fig. 14: The schematic geological setting of the amphibolites from the surrounding area after Král' et al., (1997). The arrows show the relation between the possible provenance fields and archaeological localities of the studied stone implements. For the abbreviations on the map, see those on Fig. 13.

**Table 1**

Amphibolite implements from the collection of the Herman Ottó Museum giving the different analytical methods and magnetic susceptibility (MS) values in units of  $10^{-3}$  SI. (Y: yes; N: no; \*: "Original surface", semi-quantitative method; AVK: Alföld Linear Pottery).

Sample	Inventory number	Locality	EDS/SEM	PGAA	MS	Archaeological typology	Age	Culture / Phase
B04	70.1.23	Szerencs-Taktaföldvár	Y	Y	0.73	Flat chisel	Late Neolithic	Tisza
B05	70.1.24	Szerencs-Taktaföldvár	Y	Y	0.63	Flat chisel	Late Neolithic	Tisza
B12	74.44.15	Szerencs-Taktaföldvár	Y	Y	6.47	Flat chisel	Late Neolithic	Tisza
C02	53.206.3	Tizsadorogma	Y	Y	10.51	Flat chisel	Middle Neolithic	AVK
D02	70.1.149	Szerencs-Taktaföldvár	Y	Y	0.52	Flat chisel	Late Neolithic	Tisza
D04	58.42.2	Fancsal-farm	Y	Y	46.8	Flat chisel	Neolithic	Stray find
D05	not inventoried	Lidl Hejőkürt 2mh. S2106 2005. 06. 14.	Y	Y	1.69	Flat chisel	Middle Neolithic	AVK
D08	53.238.5	Szirmabesenyő's vicinity	Y	Y	17.17	Flat chisel	Middle Neolithic	Bükk AB, AVK
D12	53.248.1	Unknown	Y	Y	n.a.	Flat chisel	Neolithic	Stray find
D13	53.250.1	Bereg county	Y	N	0.68	Flat chisel	Neolithic	Stray find
D15	53.208.1	Emőd, Vaskó-sheer	Y	Y	0.62	Flat axe	Neolithic	Stray find
D17	74.44.11	Szerencs-Taktaföldvár	Y	Y	0.39	Flat chisel	Late Neolithic	Tisza
D18	not inventoried	Miskolc Aldi2 S153 2009. 09. 11.	Y	Y	24.33	Flat chisel	Middle Neolithic	AVK, Bükk
D21	53.160.20	Borsod/Derékegyháza (Edelény)	Y	Y	8.52	Shoe-last axe	Middle Neolithic	Bükk AB, B-C
D26	74.44.10	Szerencs-Taktaföldvár	N	Y	0.50	Flat chisel	Late Neolithic	Tisza
D30*	53.160.11	Borsod/Derékegyháza (Edelény)	Y	Y	1.69	Flat chisel	Middle Neolithic	Bükk AB, B-C
D31*	53.160.31a	Borsod/Derékegyháza (Edelény)	Y	N	1.44	Flat chisel	Middle Neolithic	Bükk AB, B-C

D32*	53.160.31b	Borsod/Derékegyháza (Edelény)	Y	Y	0.31	Shoe-last axe	Middle Neolithic	Bükk AB, B-C
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**Table 1 continued**

Sample	Inventory number	Locality	EDS/ SEM	PGAA	MS	Archaeological typology	Age	Culture/ Phase
D33*	53.160.150	Borsod/Derékegyháza (Edelény)	Y	Y	29.62	Flat chisel	Middle Neolithic	Bükk AB, B-C
D34*	53.229.1	Muhi, Bala Hill	Y	Y	n.a.	Flat chisel	Neolithic	Unknown
D35	70.1.206	Szerencs-Taktaföldvár	N	Y	0.38	Axe fragment	Late Neolithic	Tisza
D37	67.13.3	Tolcsva's vicinity	N	Y	0.72	Flat axe	Neolithic	Unknown
D38	67.12.1	Egerlövő	N	Y	0.49	Shaft drilled axe	Neolithic	Unknown
D39	67.13.2	Tolcsva's vicinity	N	Y	n.a.	Shoe-last axe	Neolithic	Stray find
D40	53.187.3b	Hangács, Ludas Balk	N	Y	1.79	Shaft drilled axe	Neolithic	Unknown
D41	77.44.14	Szerencs-Taktaföldvár	N	Y	n.a.	Flat chisel	Late Neolithic	Tisza
D45*	72.11.303	Tiszavalk, Kenderföldek	Y	N	0.85	Flat chisel	Middle Neolithic	AVK
D47*	75.25.20	Unknown	Y	N	0.61	Flat chisel	Neolithic	Stray find
D48*	53.206.2	Tiszadorogma	Y	N	0.56	Tongue-shaped axe	Middle Neolithic	AVK

**Table 2**

PGAA results. The major components are given in wt%, the trace elements are in ppm. The amount of oxides is calculated from the elemental concentration, based on the oxidation numbers. The number of digits indicates the uncertainties of concentration values. “<D.L.” stands for “less than the Detection Limit”.

Sample	B04	B05	B12	C02	D02	D04	D05	D08	D12	D15	D17	D18
SiO <sub>2</sub>	54.18	50.39	47.77	47.85	47.00	49.00	50.00	50.00	50.00	47.00	59.00	52.00
TiO <sub>2</sub>	0.46	2.28	2.29	2.63	1.61	1.53	1.92	1.48	3.70	1.29	1.07	1.70
Al <sub>2</sub> O <sub>3</sub>	13.58	16.11	14.36	15.04	13.00	13.50	15.50	13.60	13.00	15.00	17.80	13.90
Fe <sub>2</sub> O <sub>3</sub> *	10.07	12.53	12.71	13.13	11.70	12.40	9.60	12.30	13.20	10.70	5.40	11.20
MnO	0.21	0.17	0.18	0.24	0.23	0.21	0.23	0.26	0.18	0.28	0.10	0.17
MgO	5.76	5.47	7.06	4.95	7.40	3.50	6.30	7.90	5.60	7.80	3.00	6.90
CaO	7.48	4.74	8.30	9.32	13.10	11.80	10.90	12.10	9.80	14.90	7.20	8.80
Na <sub>2</sub> O	4.12	4.70	3.52	3.50	1.30	1.30	2.27	1.53	2.50	0.73	2.97	3.15
K <sub>2</sub> O	1.04	0.38	1.47	0.86	0.13	0.18	0.25	0.12	<D.L.	0.19	1.17	0.24
H <sub>2</sub> O	1.59	3.16	2.22	2.27	1.69	1.72	1.59	1.62	1.21	1.63	2.24	1.30
SO <sub>3</sub>	1.38	<D.L.	<D.L.	<D.L.	2.60	1.22	1.72	0.43	0.48	<D.L.	<D.L.	0.41
Total	99.90	99.92	99.89	99.76	99.76	96.36	100.28	101.34	99.67	99.52	99.95	99.77
B	10	1	2	1	6	3	4	3	2	12	20	3
Cl	407	305	689	798	196	200	220	290	23	270	105	180
Sc	<D.L.	39	<D.L.	65	53	40	101	51	<D.L.	29	23	12
V	134	234	185	409	480	310	670	340	390	170	160	230
Cr	<D.L.	<D.L.	<D.L.	506	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.
Co	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.
Nd	<D.L.	<D.L.	29	<D.L.	15	<D.L.	21	33	34	21	22	17
Sm	1	4	4	5	2	2	3	2	6	2	3	5
Gd	2	6	5	7	4	4	5	4	7	3	4	7

\* Total Fe as Fe<sub>2</sub>O<sub>3</sub>.

**Table 2 continued**

Sample	D21	D26	D30	D32	D33	D34	D35	D37	D38	D39	D40	D41
SiO <sub>2</sub>	47.00	53.00	50.00	49.00	52.80	48.00	46.55	56.22	50.00	52.00	47.00	51.00
TiO <sub>2</sub>	1.71	1.41	2.06	1.78	2.19	1.32	1.39	1.08	0.82	1.10	2.50	2.12
Al <sub>2</sub> O <sub>3</sub>	13.60	13.20	15.11	14.80	13.64	13.80	15.78	18.11	15.60	11.20	12.90	15.30
Fe <sub>2</sub> O <sub>3</sub> *	11.40	8.50	11.10	10.20	11.75	11.10	11.85	5.96	9.70	12.50	12.60	10.90
MnO	0.19	0.28	0.17	0.19	0.20	0.20	0.22	0.14	0.10	0.34	0.39	0.19
MgO	7.30	7.20	7.60	6.50	5.90	7.40	8.94	3.39	6.60	7.10	6.20	4.20
CaO	10.30	11.40	7.40	12.70	7.77	14.70	9.96	8.79	12.40	13.00	11.20	9.20
Na <sub>2</sub> O	1.63	2.80	4.30	1.87	3.76	1.78	2.84	3.04	2.20	0.81	2.05	3.90
K <sub>2</sub> O	0.16	0.32	0.33	0.16	0.42	0.25	0.44	1.00	0.28	0.40	0.31	0.82
H <sub>2</sub> O	1.59	1.30	1.68	1.67	0.97	1.66	1.94	2.24	1.89	1.49	2.21	2.30
SO <sub>3</sub>	5.00	0.59	0.64	1.28	0.53	<D.L.	<D.L.	<D.L.	<D.L.	0.31	2.70	<D.L.
Total	99.88	101.89	100.38	100.15	99.93	100.21	99.93	99.97	99.59	100.25	100.06	99.93
B	4	4	3	3	4	6	5	23	6	4	7	2
Cl	200	201	42	250	190	0	<D.L.	90	510	570	290	690
Sc	47	33	<D.L.	65	<D.L.	<D.L.	46	34	19	9	44	21
V	410	250	360	490	268	360	336	<D.L.	250	310	430	270
Cr	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.
Co	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.	<D.L.
Nd	14	21	32	19	<D.L.	80	<D.L.	22	230	20	16	30
Sm	3	3	4	3	7	2	3	4	1	2	4	3
Gd	4	4	7	5	10	4	5	5	2	3	5	4

\* Total Fe as Fe<sub>2</sub>O<sub>3</sub>.

**Table 3** Main diagnostic minerals of *Groups- 1* and *2* of amphibolite stone implements.

Sample	Amphibole		Plagioclase		Kfs	Bi	Ep/Czo	Cpx
	core	rim	P11 (core)	P12 (rim)				
<i>Group1</i>								
D05	Act; Mg-hb	Mg-hb; Tsch	Lab	And				
D12	Act	Mg-hb	Ab	And				
D13	Act	Mg-hb	Lab	And				
D30	Mg-hb	Tsch; Prg	Ab	Olg				
D31	Act; Mg-hb	Tsch; Prg	And, Olg	And, Olg				
D33	Act	Mg-hb; Mg-f-hb	Ab	Olg			X	
<i>Group2</i>								
<i>Group2a</i>								
B12	Sdg; Fsdg; Prg; Fprg	Mg-hb; Act	And	Ab			X	
D02	Mg-hb; Fhb	Act; Fact	Lab, And, Olg	Ab			X	
D04	Mg-hb	Mg-hb	Lab, And	Lab, And			X	
D08	Mg-hb; Mg-f-hb; Fe2-ftscht; Fe3-Ts	Act	Lab, And	Lab, And			X	
D18	Mg-hb	Act	Lab, And, Olg	Ab				
D21	Mg-f-hb	Mg-f-hb	Byt, And	Ab				
D32	Mg-hb	Act	And, Olg	Ab				
D47	Mg-hb; Prg; Fprg	Act	And	Ab			X	
<i>Group2b</i>								
B04	Mg-hb	Act	And, Olg	Ab	X			
B05	Mg-hb	Act; Fact	Lab, And	Ab	X			
C02	Fhb, Mg-hb	Fhb, Act	Byt	And, Olg	X	X		
D48	Fprg, Mg-hb	Fprg, Mg-hb	And, Lab	And, Lab	X		X	
<i>Group2c</i>								
D15	Mg-hb	Act	Byt, Lab	And		X	X	X
D34	Mg-hb	Mg-hb; Act	And	Ab				X
D45	Fhb	Act	Lab; And	Lab; And				X
<i>Group2d</i>								
D17	Fprg; Mg-hb	Mg-hb; Act	And, Olg	Ab		X	X	

**Table 4** Chemical composition of amphiboles (*Group 1*).

No.	1	2	3	4	5	6	7	8
Sample	D05-15	D30	D30	D31	D30	D30	D31	D31
SiO <sub>2</sub>	47.83	45.05	45.71	44.21	44.32	44.45	43.47	43.48
TiO <sub>2</sub>	0.16	0.35	0.30	0.26	0.30	0.40	0.45	0.50
Al <sub>2</sub> O <sub>3</sub>	15.60	15.68	15.09	16.54	13.94	13.45	16.53	14.88
MnO	0.23	0.00	0.00	0.11	0.05	0.00	0.10	0.19
FeO*	12.60	9.14	9.12	10.40	12.34	14.51	11.44	12.54
Fe <sub>2</sub> O <sub>3</sub> *	0.00	3.66	4.04	2.93	2.10	0.00	1.45	1.60
MgO	8.80	12.21	11.56	10.65	11.48	11.62	10.91	10.98
CaO	10.80	8.82	9.16	10.28	11.15	11.29	10.88	11.63
Na <sub>2</sub> O	1.64	2.76	2.63	2.34	2.06	1.97	2.58	1.88
K <sub>2</sub> O	0.32	0.28	0.34	0.25	0.24	0.29	0.17	0.31
H <sub>2</sub> O**	2.05	2.04	2.04	2.03	2.00	2.01	2.03	2.01
<b>Total</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>
Cation numbers based on 23 oxygens.								
Si	6.85	6.45	6.55	6.38	6.46	6.48	6.30	6.36
Al	1.15	1.55	1.45	1.62	1.54	1.52	1.70	1.64
<b>ΣT</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>
Ti	0.02	0.04	0.03	0.03	0.03	0.04	0.05	0.06
Al	1.48	1.10	1.10	1.19	0.86	0.79	1.13	0.92
Fe <sup>3+</sup>	0.00	0.40	0.44	0.32	0.23	0.23	0.16	0.18
Mn <sup>2+</sup>	0.03	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Fe <sup>2+</sup>	1.51	0.87	0.97	1.18	1.38	1.40	1.31	1.46
Mg	1.88	2.61	2.47	2.29	2.50	2.53	2.36	2.39
<b>ΣC</b>	<b>4.91</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>
Mn <sup>2+</sup>	0.00	0.00	0.00	0.01	0.01	0.00	0.01	0.02
Fe <sup>2+</sup>	0.00	0.23	0.13	0.08	0.13	0.13	0.08	0.08
Ca	1.66	1.35	1.41	1.59	1.74	1.76	1.69	1.82
Na	0.34	0.42	0.47	0.32	0.13	0.10	0.22	0.08



<b>ΣB</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>
Na	0.11	0.35	0.26	0.34	0.46	0.45	0.51	0.46
K	0.06	0.05	0.06	0.05	0.05	0.06	0.03	0.06
<b>ΣA</b>	<b>0.17</b>	<b>0.40</b>	<b>0.33</b>	<b>0.38</b>	<b>0.50</b>	<b>0.51</b>	<b>0.54</b>	<b>0.51</b>
<i>Species</i>	<i>Tsch</i>	<i>Tsch</i>	<i>Tsch</i>	<i>Tsch</i>	<i>Prg</i>	<i>Prg</i>	<i>Prg</i>	<i>Prg</i>

\* Total Fe was measured as FeO. FeO/Fe<sub>2</sub>O<sub>3</sub> ratio was calculated with ACES\_2 Excel spreadsheet (after Locock, 2014).

\*\* H<sub>2</sub>O was calculated from the stoichiometry: OH = 2 *apfu*.

**Table 4 continued**

No.	9	10	11	12	13	14	15
Sample	D05-6	D05	D05	D12	D12	D13	D13
SiO <sub>2</sub>	51.86	50.10	51.99	48.05	45.61	46.12	49.49
TiO <sub>2</sub>	0.22	0.14	0.23	0.33	0.27	0.23	0.25
Al <sub>2</sub> O <sub>3</sub>	6.31	8.67	5.47	9.79	13.39	13.04	9.36
MnO	0.37	0.22	0.31	0.19	0.22	0.16	0.14
FeO*	13.24	11.18	13.67	15.09	15.44	13.92	13.04
Fe <sub>2</sub> O <sub>3</sub> *	0.24	1.40	0.59	2.16	1.93	2.60	1.33
MgO	13.89	13.68	13.50	10.18	8.60	9.76	12.42
CaO	11.50	12.05	11.79	11.30	11.32	11.11	11.22
Na <sub>2</sub> O	0.25	0.40	0.33	0.81	1.04	0.87	0.60
K <sub>2</sub> O	0.10	0.12	0.08	0.10	0.20	0.18	0.12
H <sub>2</sub> O**	2.05	2.04	2.04	2.00	2.00	2.01	2.04
<b>Total</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>
Cation numbers based on 23 oxygens.							
Si	7.45	7.17	7.50	7.03	6.70	6.72	7.13
Al	0.55	0.83	0.50	0.97	1.30	1.28	0.87
<b>ΣT</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>
Ti	0.02	0.02	0.02	0.04	0.03	0.03	0.03
Al	0.51	0.64	0.43	0.72	1.01	0.96	0.72
Fe <sup>3+</sup>	0.03	0.15	0.06	0.24	0.21	0.29	0.14



Cation numbers based on 23 oxygens.										
Si	6.01	6.00	5.94	6.01	6.11	6.23	6.31	6.03	5.98	6.00
Al	2.00	2.01	2.06	2.00	1.89	1.78	1.69	1.97	2.02	2.00
<b>ΣT</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>
Ti	0.01	0.05	0.00	0.01	0.02	0.01	0.01	0.07	0.05	0.05
Al	1.42	1.30	1.39	1.42	1.24	1.20	1.04	1.15	1.23	1.31
Fe <sup>3+</sup>	0.17	0.18	0.19	0.17	0.19	0.17	0.16	0.11	0.14	0.08
Mn <sup>2+</sup>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01
Fe <sup>2+</sup>	1.75	1.63	1.62	1.75	1.63	1.67	1.92	1.93	1.95	1.95
Mg	1.65	1.83	1.80	1.65	1.92	1.97	1.88	1.74	1.63	1.61
<b>ΣC</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>
Mn <sup>2+</sup>	0.00	0.00	0.01	0.00	0.01	0.00	0.03	0.05	0.06	0.05
Fe <sup>2+</sup>	0.10	0.11	0.10	0.10	0.10	0.10	0.06	0.01	0.02	0.00
Ca	1.76	1.72	1.74	1.76	1.76	1.72	1.78	1.64	1.62	1.60
Na	0.14	0.18	0.16	0.14	0.13	0.18	0.13	0.30	0.30	0.36
<b>ΣB</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>
Na	0.51	0.57	0.61	0.51	0.52	0.56	0.54	0.67	0.64	0.66
K	0.02	0.03	0.03	0.02	0.04	0.03	0.07	0.21	0.21	0.21
<b>ΣA</b>	<b>0.53</b>	<b>0.60</b>	<b>0.64</b>	<b>0.53</b>	<b>0.57</b>	<b>0.58</b>	<b>0.60</b>	<b>0.88</b>	<b>0.85</b>	<b>0.87</b>
<i>Species</i>	<i>F-sdg</i>	<i>Sdg</i>	<i>Sdg</i>	<i>F-sdg</i>	<i>Prg</i>	<i>Prg</i>	<i>Fprg</i>	<i>Fprg</i>	<i>Fprg</i>	<i>Fprg</i>

\* Total Fe was measured as FeO. FeO/Fe<sub>2</sub>O<sub>3</sub> ratio was calculated with ACES\_2 Excel spreadsheet (after Locock, 2014).

\*\* H<sub>2</sub>O was calculated from the stoichiometry: OH = 2 *apfu*.

**Table 5 continued**

No.	11	12	13	14	15	16	17	18	19	20
Sample	D08-5	C02-25	B05-42	D08-29	D15-4	D17-44	C02-11	B05-27	D08-30	D15-2
Core/rim										
SiO <sub>2</sub>	40.24	49.72	49.65	47.48	47.19	52.80	50.94	52.89	53.51	52.82
TiO <sub>2</sub>	0.00	0.00	0.13	0.08	0.03	0.08	0.03	0.07	0.11	0.00
Al <sub>2</sub> O <sub>3</sub>	14.28	5.99	8.83	10.88	12.12	5.07	4.87	4.12	4.66	5.06
MnO	0.20	0.14	0.00	0.28	0.09	0.32	0.26	0.00	0.27	0.16
FeO*	18.48	22.82	12.01	10.46	12.51	14.10	20.58	17.83	8.99	9.53
Fe <sub>2</sub> O <sub>3</sub> *	10.07	0.03	2.50	3.82	1.19	1.00	0.12	0.00	1.09	1.42
MgO	6.73	7.45	12.81	12.65	11.47	13.18	9.23	11.78	17.05	16.03
CaO	7.52	11.64	11.03	11.34	11.64	10.14	11.65	10.83	11.79	12.34
Na <sub>2</sub> O	0.25	0.13	0.97	0.85	1.27	1.11	0.21	0.48	0.41	0.54
K <sub>2</sub> O	0.23	0.14	0.04	0.15	0.47	0.17	0.12	0.00	0.06	0.05
H <sub>2</sub> O**	1.91	1.95	2.03	2.03	2.03	2.04	1.98	2.01	2.07	2.06
<b>Total</b>	<b>99.91</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>
Cation numbers based on 23 oxygens.										
Si	6.12	7.47	7.14	6.85	6.83	7.61	7.58	7.71	7.56	7.50
Al	1.88	0.53	0.86	1.15	1.17	0.39	0.42	0.29	0.44	0.50
<b>ΣT</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>	<b>8.00</b>
Ti	0.00	0.00	0.01	0.01	0.00	0.01	0.00	0.01	0.01	0.00
Al	0.68	0.53	0.64	0.69	0.90	0.47	0.43	0.42	0.33	0.35
Fe <sup>3+</sup>	1.15	0.00	0.27	0.41	0.13	0.11	0.01	0.00	0.12	0.15
Mn <sup>2+</sup>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Fe <sup>2+</sup>	1.64	2.79	1.33	1.17	1.49	1.58	2.51	2.01	0.95	1.11
Mg	1.53	1.67	2.75	2.72	2.48	2.83	2.05	2.56	3.59	3.39
<b>ΣC</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>	<b>5.00</b>
Mn <sup>2+</sup>	0.03	0.02	0.00	0.04	0.01	0.04	0.03	0.00	0.03	0.02
Fe <sup>2+</sup>	0.71	0.07	0.12	0.10	0.03	0.12	0.05	0.17	0.11	0.02
Ca	1.23	1.87	1.70	1.75	1.81	1.57	1.86	1.69	1.78	1.88

Na	0.04	0.04	0.18	0.12	0.16	0.27	0.06	0.14	0.08	0.08
<b>ΣB</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>2.00</b>	<b>1.99</b>	<b>2.00</b>	<b>2.00</b>
Na	0.04	0.00	0.09	0.12	0.20	0.04	0.00	0.00	0.04	0.07
K	0.05	0.03	0.01	0.03	0.09	0.03	0.02	0.00	0.01	0.01
<b>ΣA</b>	<b>0.08</b>	<b>0.03</b>	<b>0.10</b>	<b>0.15</b>	<b>0.29</b>	<b>0.07</b>	<b>0.03</b>	<b>0.00</b>	<b>0.05</b>	<b>0.08</b>
<i>Species</i>	<i>F-f-tsch</i>	<i>Fhb</i>	<i>Mg-hb</i>	<i>Mg-hb</i>	<i>Mg-hb</i>	<i>Mg-hb</i>	<i>Fact</i>	<i>Act</i>	<i>Act</i>	<i>Act</i>

**Table 6** Chemical composition of epidotes.

No.	1	2	3	4	5	6	7	8	9	10
Sample	B12	D02	D02	D04	D04	D08	D08	D15	D15	D17
SiO <sub>2</sub>	39.52	39.36	39.32	39.13	39.15	38.44	38.83	38.66	39.26	38.83
TiO <sub>2</sub>	0.00	0.07	0.46	0.04	0.00	0.00	0.03	0.04	0.08	0.16
Al <sub>2</sub> O <sub>3</sub>	25.43	26.82	28.95	27.05	29.63	23.82	29.75	27.69	29.52	23.69
Fe <sub>2</sub> O <sub>3</sub> *	11.04	10.10	7.22	9.52	6.54	13.53	6.72	8.81	6.40	13.24
MnO	0.00	0.00	0.09	0.07	0.00	0.10	0.00	0.06	0.08	0.12
MgO	0.37	0.09	0.07	0.27	0.30	0.17	0.26	0.28	0.41	0.04
CaO	21.46	21.54	21.86	21.79	22.25	21.92	22.37	22.43	21.99	21.95
Na <sub>2</sub> O	0.18	0.00	0.00	0.14	0.14	0.06	0.05	0.05	0.25	0.00
K <sub>2</sub> O	0.07	0.10	0.09	0.06	0.05	0.06	0.06	0.06	0.07	0.07
H <sub>2</sub> O**	1.92	1.92	1.94	1.92	1.94	1.90	1.93	1.92	1.94	1.90
<b>Total</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>	<b>100.00</b>
Cation numbers based on 12.5 oxygens.										
Si	3.08	3.06	3.03	3.01	3.01	3.04	2.99	3.00	3.02	3.06
<b>ΣT</b>	<b>3.08</b>	<b>3.06</b>	<b>3.03</b>	<b>3.01</b>	<b>3.01</b>	<b>3.04</b>	<b>2.99</b>	<b>3.00</b>	<b>3.02</b>	<b>3.06</b>
Ti	0.00	0.00	0.03	0.00	0.00	0.00	0.00	0.00	0.00	0.01
Al	2.34	2.46	2.63	2.45	2.69	2.22	2.69	2.54	2.68	2.20
Fe <sup>3+</sup>	0.65	0.59	0.42	0.55	0.38	0.80	0.39	0.52	0.37	0.79
Mg	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
<b>ΣM</b>	<b>2.99</b>	<b>3.05</b>	<b>3.08</b>	<b>3.00</b>	<b>3.07</b>	<b>3.02</b>	<b>3.08</b>	<b>3.06</b>	<b>3.05</b>	<b>3.00</b>
Mn <sup>2+</sup>	0.00	0.00	0.01	0.00	0.00	0.01	0.00	0.00	0.01	0.01

Mg	0.04	0.01	0.01	0.03	0.03	0.02	0.03	0.03	0.05	0.00
Ca	1.79	1.79	1.80	1.80	1.83	1.86	1.85	1.87	1.81	1.86
Na	0.03	0.00	0.00	0.02	0.02	0.01	0.01	0.01	0.04	0.00
K	0.01	0.01	0.01	0.01	0.00	0.01	0.01	0.01	0.01	0.01
<b>ΣA</b>	<b>1.87</b>	<b>1.81</b>	<b>1.83</b>	<b>1.86</b>	<b>1.88</b>	<b>1.91</b>	<b>1.90</b>	<b>1.92</b>	<b>1.92</b>	<b>1.88</b>
<i>Species</i>	<i>Ep</i>	<i>Ep</i>	<i>Czo</i>	<i>Ep</i>	<i>Czo</i>	<i>Ep</i>	<i>Czo</i>	<i>Ep</i>	<i>Czo</i>	<i>Ep</i>

\* Total Fe was measured as Fe<sub>2</sub>O<sub>3</sub>.

\*\* H<sub>2</sub>O was calculated from the stoichiometry: OH = 1 *apfu*.

**Table 7** Chemical composition of biotites.

No.	1	2	3	4
Sample	D15	D17	D17	D17
SiO <sub>2</sub>	39.72	38.09	39.79	39.06
TiO <sub>2</sub>	0.17	0.22	0.15	0.15
Al <sub>2</sub> O <sub>3</sub>	16.42	18.55	16.70	17.45
FeO*	14.49	14.18	13.20	13.33
MnO	0.06	0.12	0.12	0.10
MgO	15.59	15.19	16.24	16.38
CaO	0.50	0.11	0.15	0.17
Na <sub>2</sub> O	0.20	0.06	0.00	0.00
K <sub>2</sub> O	8.75	9.39	9.52	9.25
H <sub>2</sub> O**	4.10	4.10	4.11	4.11
Total	100.00	100.00	100.00	100.00
Cation numbers based on 11 oxygens				
Si	2.90	2.79	2.90	2.85
Al	1.10	1.21	1.10	1.15
<b>ΣT</b>	<b>4.00</b>	<b>4.00</b>	<b>4.00</b>	<b>4.00</b>
Al	0.32	0.39	0.33	0.35
Ti	0.01	0.01	0.01	0.01



Al	1.41	1.49	1.41	1.42	1.65	1.41	1.61	0.36	0.56	0.87
Ti	0.01	0.01	0.00	0.01	0.00	0.00	0.00	0.00	0.00	0.00
Fe <sup>3+</sup>	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.19	1.14	0.42
Fe <sup>2+</sup>	1.76	1.68	2.36	2.30	3.44	2.56	2.45	3.93	3.34	3.77
Mn	0.01	0.01	0.03	0.03	0.01	0.02	0.01	0.01	0.00	0.00
Mg	2.67	2.32	1.80	1.90	0.56	1.82	1.56	0.45	0.91	0.81
Ca	0.02	0.18	0.13	0.09	0.03	0.00	0.00	0.02	0.01	0.03
Na	0.01	0.04	0.00	0.00	0.00	0.00	0.00	0.03	0.04	0.07
K	0.00	0.01	0.01	0.02	0.00	0.00	0.00	0.01	0.00	0.03
<b>ΣM</b>	<b>5.88</b>	<b>5.74</b>	<b>5.74</b>	<b>5.75</b>	<b>5.69</b>	<b>5.82</b>	<b>5.63</b>	<b>6.00</b>	<b>6.00</b>	<b>6.00</b>

\* Total Fe was measured as FeO. FeO/Fe<sub>2</sub>O<sub>3</sub> ratio was calculated from the stoichiometry: ΣM = 6.00 *apfu*.

\*\* H<sub>2</sub>O was calculated from the stoichiometry: OH = 8 *apfu*.

**Table 9**

Thermobarometry data of polished stone tools. MT: Medium-temperature amphibolites; HT: High-temperature amphibolites. \*: measured with "original-surface" method.

Sample	Number of measurements	T <sub>avg</sub> (°C)	P <sub>avg</sub> (kbar)	Standard deviation T (°C)	Standard deviation P (kbar)	T <sub>max</sub> (°C)	P <sub>(Tmax)</sub> (kbar)
MT							
B04	2	441	2.7	8.49	0.12	450	2.8
B05	3	510	3.9	7.58	0.17	516	4.0
C02	4	442	2.7	12.88	0.18	450	2.8
D02	2	480	3.4	22.45	0.41	500	3.7
D04	1	510	3.7	0.00	0.00	510	3.7
D05	1	540	5.3	0.00	0.00	540	5.3
D09	3	480	3.0	28.36	0.38	510	3.5
D18	3	530	4.0	14.32	0.18	545	4.1
D21	8	540	2.8	18.66	0.18	600	3.4
D32*	1	548	4.0	0.00	0.00	548	4.0
D33*	3	455	2.7	9.73	0.21	460	2.8
HT							
B12	6	670	7.1	12.53	0.30	680	7.3
D08	6	580	4.7	38.66	0.55	600	4.8
D12	3	540	4.7	25.75	0.59	570	5.0
D13	5	520	4.0	28.14	0.56	560	5.0





Little Carpathians	*	*				*		*	*	*	
Tribeč											
Stražovské Mountains											
Zemplinicum			**								

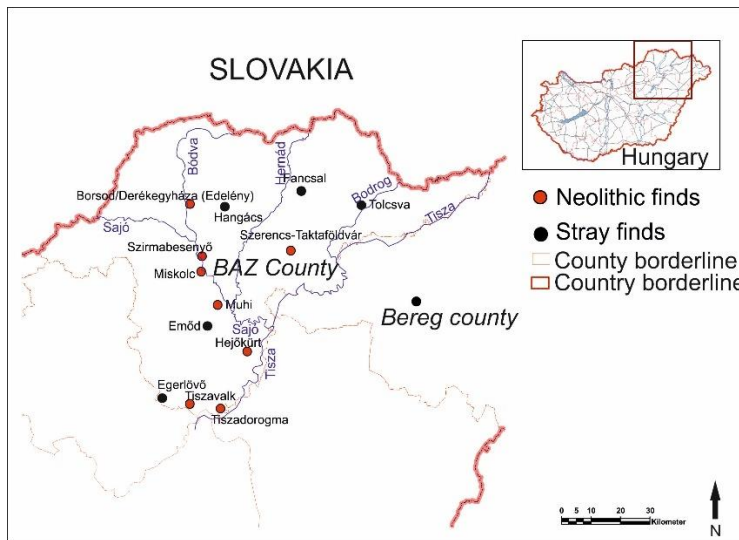


Figure 1

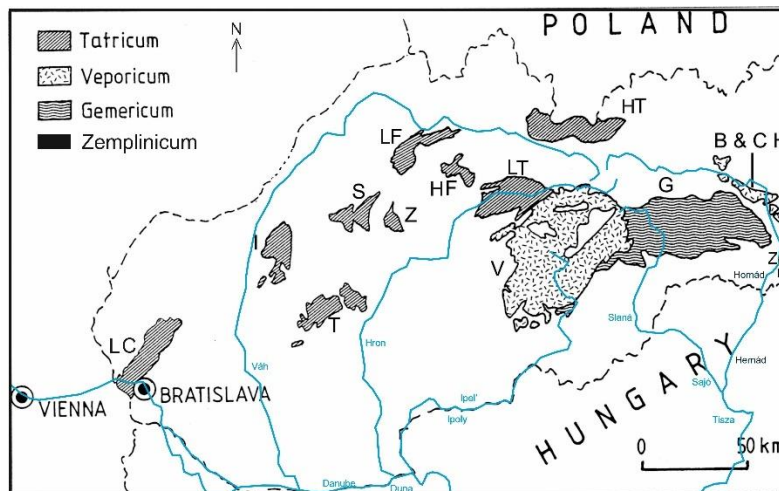


Figure 2

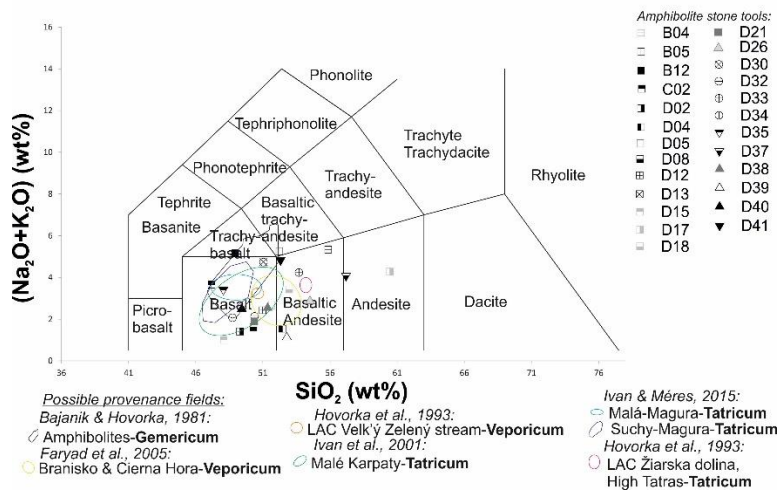


Figure 3

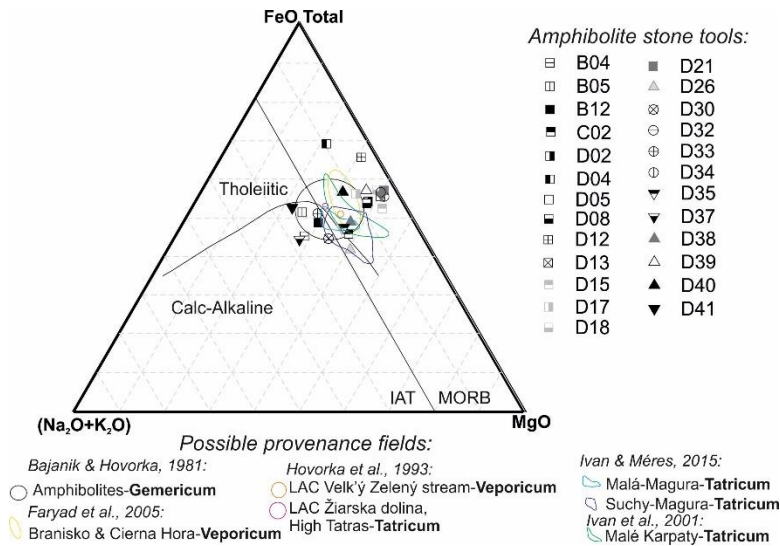


Figure 4

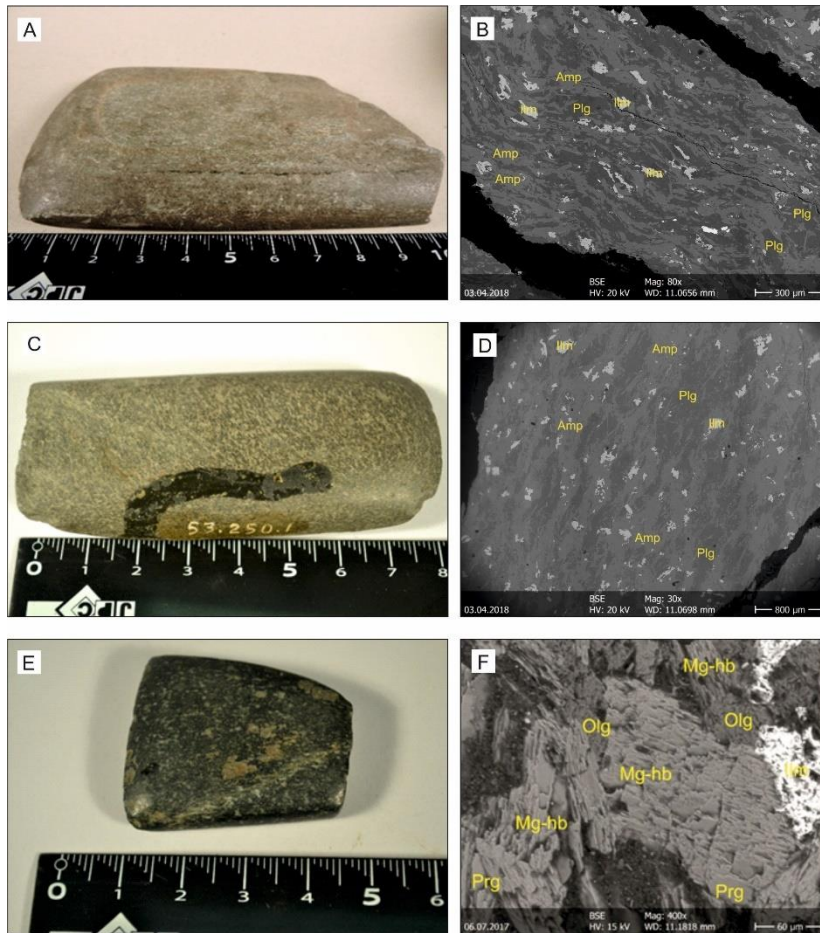


Figure 5

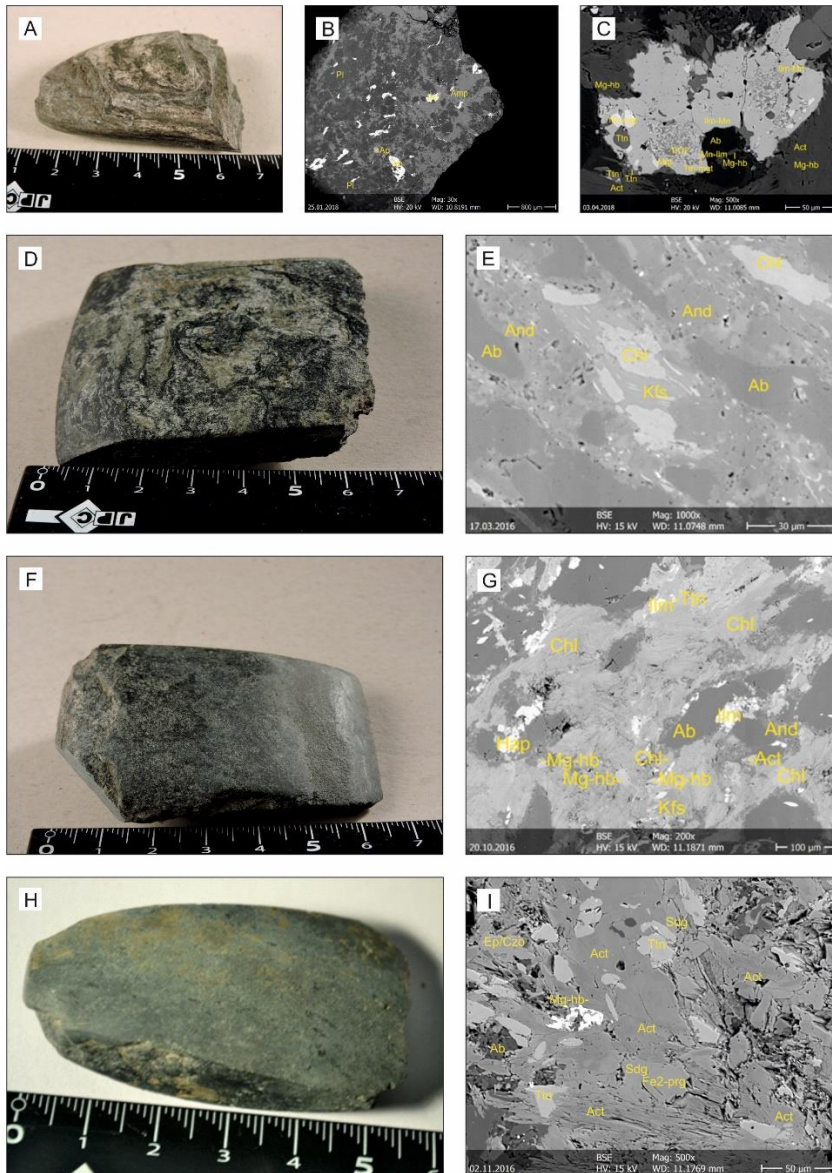


Figure 6

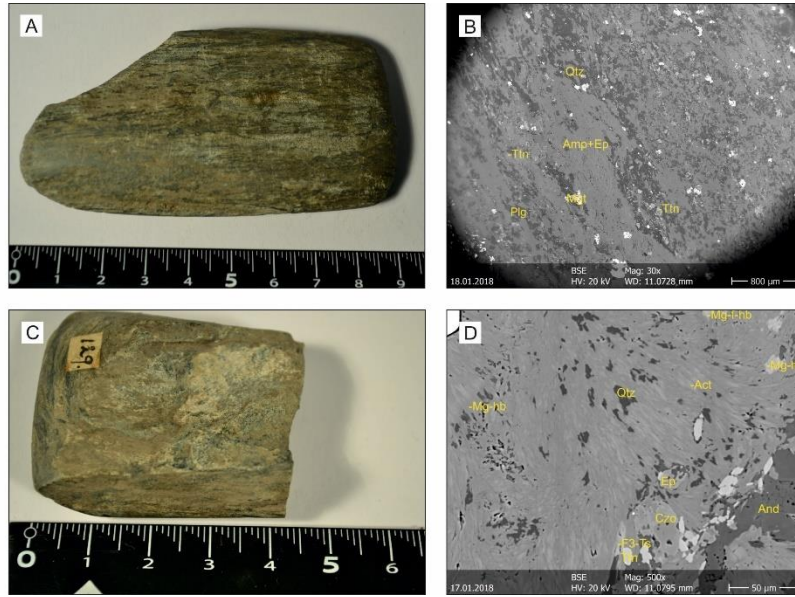


Figure 7

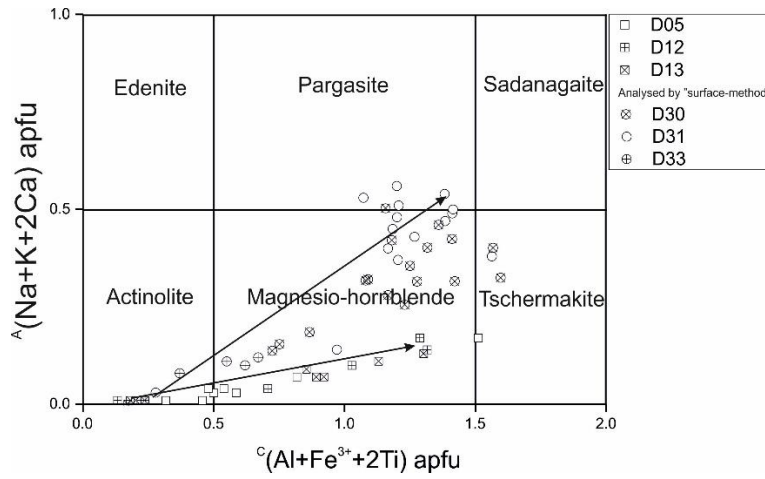


Figure 8

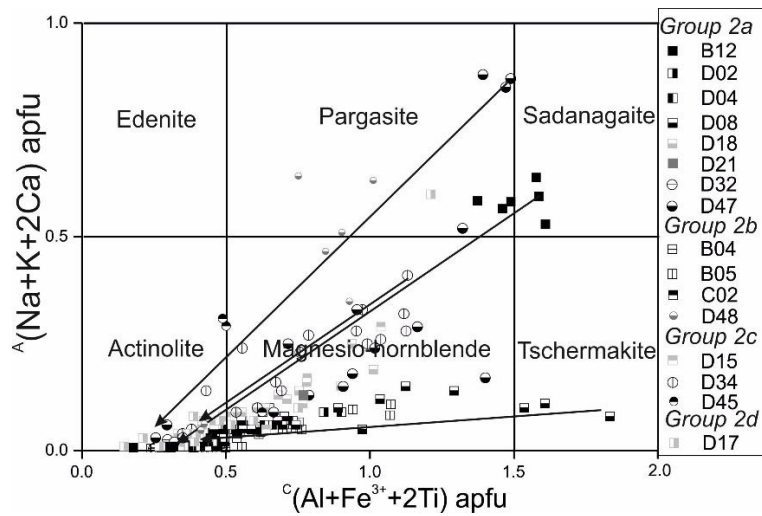


Figure 9