

Pore scale modeling of drainage displacement patterns in association with geological sequestration of CO₂

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Key Points:

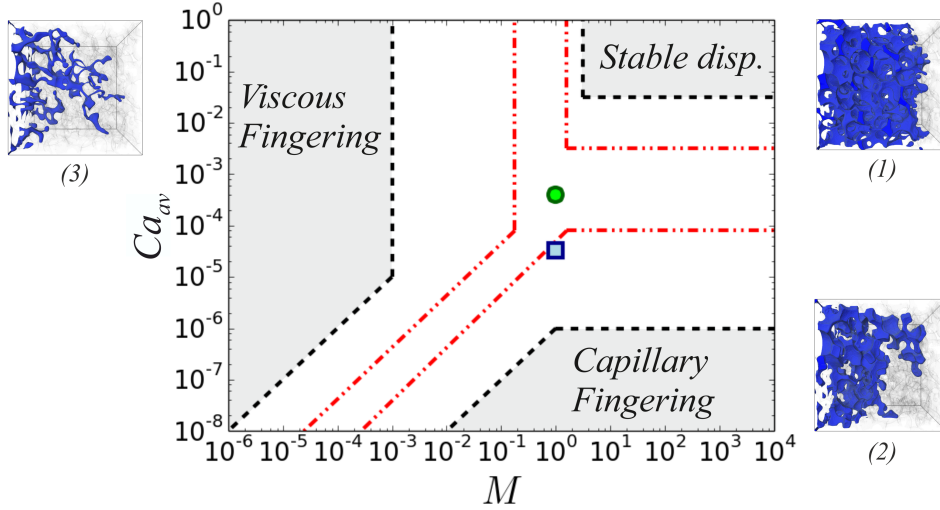
- Inertial effects cannot be neglected in the range of typical average capillary numbers (Ca_{av}) associated with multiphase flow in permeable media ($Ca_{av} \leq 10^{-3}$).
- Even as the average capillary (Ca_{av}) and Reynolds (Re_{av}) numbers decrease away from the injection point, inertial effects become important during abrupt jump events (Haines jumps).
- The local Ca_l and Re_l during jump events is orders of magnitude higher than the average flow dimensionless numbers.
- The maximum local Reynolds number Re_l during jumps is of the order 10^1 .
- The Ohnesorge number, which reflects the thermophysical properties of the system under investigation, links the capillary and Reynolds numbers and should be used to restrict the parameter selection process.
- A Navier-Stokes solver should be used to investigate pore scale displacement processes.

Abstract

We investigate the immiscible displacement (drainage) of a wetting fluid in a porous medium by a non-wetting fluid using multi-GPU lattice Boltzmann simulations with the aim of better understanding the pore scale processes involved in the geological sequestration of CO₂. Correctly resolving the dynamics involved in multiphase flow in permeable media is of paramount importance for any numerical scheme. Generally the average fluid flow is assumed to be at low Reynolds numbers Re_{av} . Hence, by neglecting inertial effects, this immiscible displacement should be characterised by just two dimensionless numbers, namely the capillary number Ca_{av} and the viscosity ratio, which quantify the ratio of the relevant forces, i.e. the viscous and capillary forces. Our investigation reveals that inertial effects cannot be neglected in the range of typical capillary numbers associated with multiphase flow in permeable media. Even as the average Ca_{av} and Re_{av} decrease away from the injection point, inertial effects remain important over a transient amount of time during abrupt Haines jumps, when the non-wetting phase passes from a narrow restriction to a wider pore space. The local Re_l at the jump sites is orders of magnitude higher than the average Re_{av} , with the local dynamics being decoupled from the externally imposed flow rate. Therefore, a full Navier-Stokes solver should be used for investigating pore scale displacement processes. Using the Ohnesorge number to restrict the parameter selection process is essential, as this dimensionless number links Ca_{av} and Re_{av} and reflects the thermophysical properties of a given system under investigation.

1 Introduction

The immiscible displacement of a fluid in a porous medium is of extreme importance in a plethora of applications, including hydrocarbon recovery, geological sequestration of CO₂ in saline aquifers etc. This has prompted experimental and numerical investigations of these phenomena, as understanding the factors affecting the displacement patterns becomes of paramount importance. Lenormand et al. (1988) investigated numerically and experimentally the immiscible displacement of a wetting (w) fluid in a micro-model by a non-wetting (nw) fluid. Completely neglecting inertial effects and considering only viscous and capillary forces, the drainage displacement process is governed by the ratio of these forces. Hence, Lenormand et al. (1988) argue that this can be characterized by: a) the ratio of viscous forces in both fluids, quantified by the ratio of the



65 **Figure 1.** Lenormand’s phase diagram for drainage and the domains of validity for the three
 66 displacement patterns: (1) stable displacement, (2) capillary and (3) viscous fingering. The gray
 67 zones denote the domains indicated by the simulations of Lenormand et al. (1988), while the
 68 boundaries given by the dashed dotted lines are from the experimental work of Zhang et al.
 69 (2011). These boundaries are shown here just for qualitative purposes, as they are strongly de-
 70 pendent upon the geometry type. The three configurations shown for illustration purposes of
 71 the three typical displacement patterns are from drainage simulations in Ketton limestone (see
 72 Fig. 2). The injected non-wetting phase is shown in blue, while rock grains and wetting phase are
 73 transparent. The two points correspond to the test cases examined in section 3.1.

54 dynamical viscosities η_i ($i = w, nw$), $M = \eta_{nw}/\eta_w$, and b) the ratio of viscous forces
 55 which act in the injected fluid over capillary forces. This is given by the average capil-
 56 lary number $Ca_{av} = \eta_{nw}\bar{u}/\gamma$, where \bar{u} and γ are the average velocity of the injected fluid
 57 and interfacial tension respectively. Of course the fluids’ affinity to the porous media (i.e.
 58 wettability) should have a profound influence on the displacement patterns (Zhao et al.,
 59 2016; Singh et al., 2017; Rabbani et al., 2017). Depending on the dimensionless num-
 60 bers above, either viscous or capillary forces dominate and the displacement pattern takes
 61 one of the basic forms: (a) viscous fingering, (b) capillary fingering or (c) stable displace-
 62 ment. The domains of validity of the different basic mechanisms can be mapped onto
 63 the (Ca_{av}, M) phase plane, which has been called the “phase-diagram” for immiscible
 64 displacements (Lenormand et al., 1988), see Fig. 1.

74 At high values of the capillary number and viscosity ratios the stable displacement
75 regime is observed, where the principal force is due to the viscous forces in the injected
76 fluid. Capillary effects and pressure drop in the displaced fluid are negligible. At values
77 of the viscosity ratio below 1, i.e. when the injected fluid has lower viscosity than the
78 viscosity of the wetting fluid, and at high values of the capillary number, then the vis-
79 cuous fingering instability dictates the displacement pattern. Here the principal force is
80 due to the viscous forces in the displaced fluid, with capillary effects and pressure drop
81 in the displacing fluid being negligible. Viscous fingers formed grow towards the mean
82 direction of the flow. Finally the third regime, observed at low values of the capillary
83 number, is called the capillary fingering regime. Viscous forces are negligible and the prin-
84 cipal force is due to capillary forces. The fingering process is distinctively different to
85 the one observed in the viscous fingering regime as now fingers form loops and grow in
86 all directions. The boundaries between the different displacement pattern regimes de-
87 pend on the details of the geometry used and the degree of fluid wettability on solid sur-
88 faces.

89 Yortsos et al. (1997) also provide a phase diagram of fully developed drainage in
90 porous media with axes the viscosity ratio and capillary number and postulate a descrip-
91 tion in terms of invasion percolation in a gradient. They recognize the existence of two
92 different global patterns, depending on whether invasion is in a stabilizing gradient or
93 a destabilizing gradient, respectively, and propose that the various properties of the flow
94 regimes can be determined by the spatial variation of the percolation probability and the
95 sign of its gradient.

96 The important question to be addressed is whether we should indeed neglect com-
97 pletely inertial effects, when it comes to modeling two phase flow at the pore scale, for
98 example for CO₂ geological sequestration or hydrocarbon recovery. Furthermore, what
99 conditions, flow regimes, or systems of fluids would justify this choice and render Ca_{av}
100 and M sufficient to fully capture the physics of fluid displacement. Lenormand's phase
101 diagram (Lenormand et al., 1988) provides a qualitative picture about the overall drainage
102 displacement patterns. However, when it comes to quantifying important aspects of two-
103 phase flow, for example: a. the fluids' spatial distribution/connectedness, b. the displace-
104 ment efficiency (limiting value for the injected phase saturation) in the different flow regimes
105 or c. the flow regimes' domain boundaries, we need to question whether the above treat-
106 ment is sufficient. In terms of direct numerical modeling this is of paramount importance,

107 as it affects the setup of the simulations and the parameter selection process; this con-
108 sequently determines the reliability of the numerical results and their predictive capa-
109 bilities.

110 When considering processes like hydrocarbon recovery (brine displacing oil) or CO₂
111 geological sequestration (CO₂ displacing brine), generally, the average fluid flow at the
112 pore scale is assumed to be at low Reynolds numbers; especially, as the distance from
113 the injection point in the well bore increases. However, considerable work in the liter-
114 ature demonstrates the relevance of inertial effects in multiphase flow during both im-
115 bibition (Ferrari & Lunati, 2014; Zacharoudiou et al., 2017) and drainage (Moebius &
116 Or, 2014; Kazemifar et al., 2016; Li et al., 2017). During imbibition, interfacial oscilla-
117 tions and inertia can affect the displacement sequence/pathways (Ferrari & Lunati, 2014)
118 and hence the displacement efficiency. Capillary filling dynamics is also not fully cap-
119 tured by just the capillary number, as the choice of parameters affects the dissipation
120 of energy in the system, leading to different dynamics, e.g. wetting film propagation (Zacharoudiou
121 et al., 2017).

122 In a drainage situation, inertial effects can be also important over a transient amount
123 of time, during Haines jump events (Berg et al., 2013; Armstrong & Berg, 2013; Arm-
124 strong et al., 2015; Zacharoudiou et al., 2018). During these sharp interfacial jumps, cap-
125 illary forces accelerate the fluid interface, as it passes from a narrow restriction to a wider
126 pore body, while initially inertial and viscous forces oppose the motion. This leads lo-
127 cally to fluid velocities that can be orders of magnitude higher than the average fluid ve-
128 locity. Kazemifar et al. (2016) and Li et al. (2017) report maximum local Reynolds num-
129 bers in the range of 10^2 and 10^1 in supercritical CO₂ - water experiments in 2D homo-
130 geneous and heterogeneous porous micromodels respectively, exceeding the range of va-
131 lidity of Darcy's law. At a second stage, as the interface decelerates, inertial and cap-
132 illary forces become the driving forces, opposed by viscous forces; the time scales of this
133 stage constitute the majority of the pore draining times (Zacharoudiou & Boek, 2016).
134 As inertia becomes a driving force, it can affect the invasion of subsequent throats. The
135 numerical work of Moebius and Or (2014) on drainage in a 2D pore-throat network, re-
136 vealed the role of inertia in shifting the distribution of invaded throats by including smaller
137 invaded throats compared to invasion patterns without inertia. This confirmed previ-
138 ous observations (Moebius & Or, 2012) that inertia affects local displacement patterns.

139

140 Another important feature of these abrupt events is the extensive fluid rearrange-
 141 ment taking place in the region of the jump, which contributes significantly to the drain-
 142 ing of the pore body (Berg et al., 2013; Zacharoudiou et al., 2018). Experimentally, it
 143 has been observed that the draining of pore bodies occurs at time scales smaller than
 144 the time scales expected considering the externally imposed injection flow rate, indicat-
 145 ing that cooperative filling dynamics takes place (Berg et al., 2013). We have confirmed
 146 this observation numerically using free energy lattice Boltzmann simulations on Ketton
 147 limestone (Zacharoudiou et al., 2018). Hence, we expect that completely neglecting in-
 148 ertial effects, which is a key element of the jumps, will not enable capturing the fluid flow
 149 dynamics and the associated features of these abrupt events, e.g. pressure drop, fluid re-
 150 arrangement etc.

151 Therefore, it would be reasonable to question whether the description in terms of
 152 just the average capillary number Ca_{av} and the viscosity ratio M is sufficient to describe
 153 pore-scale two-phase flow behaviour correctly. In particular we question the usual ap-
 154 proach in numerical simulations of keeping both Ca_{av} and Re_{av} low. The questions that
 155 need to be addressed, mainly from the modeling point of view, are the following:

- 156 • How small is the average flow Reynolds number $Re_{av} = \rho \bar{u} L_s / \eta_{nw}$ (L_s being the
 157 characteristic length scale of the system)?
- 158 • When does the transition from smooth flow to burst flow (Haines jumps) take place?
- 159 • What is the local Reynolds number Re_l due to interfacial instabilities, e.g. Haines
 160 jumps?

161 Although a significant amount of research deals with the numerical modeling of two
 162 phase flow at the pore scale (Raeni et al., 2014; Tsuji et al., 2016; Yamabe et al., 2015),
 163 a discussion on the parameter selection process is in most cases lacking. In our previ-
 164 ous work in micromodels (Zacharoudiou & Boek, 2016; Zacharoudiou et al., 2017) we high-
 165 lighted the need to numerically match not only the capillary number and the viscosity
 166 ratio, but also the Ohnesorge number for a given experimental system. Y. Chen et al.
 167 (2018) follow our approach for the parameter selection process in their lattice Boltzmann
 168 simulations in 2D heterogeneous micromodels and present a rigorous discussion on the
 169 matter. Here we extend our previous investigations to three dimensional permeable me-
 170 dia. The paper is organized as follows. In the next section we provide the details of the
 171 numerical scheme we shall use. We present and discuss our results in section 3. We fo-

172 cus on demonstrating that Ca_{av} , M are not sufficient to fully describe two phase flow
 173 at the pore scale by examining simulations at fixed Ca_{av} , M and wettability conditions.
 174 We examine two cases at intermediate and low Ca_{av} (indicated in Fig. 1), with the lat-
 175 ter being characterised by Haines jumps. The flow regimes examined (intermediate $Ca_{av} \sim$
 176 3×10^{-4} , low $Ca_{av} \sim 4 \times 10^{-5}$) can be considered sufficient for the purposes of the
 177 current investigation, given that we encounter the transition to the capillary fingering
 178 regime characterised by Haines jumps, expected at the low Ca_{av} flow. Typical values
 179 for the ratio of viscous to capillary forces at the pore scale, i.e. Ca_{av} , are in the range
 180 of $10^{-10} - 10^{-3}$, depending on the distance from the injection point in the well bore (Blunt
 181 & Scher, 1995). Then we examine individual jump events at varying Ca_{av} , demonstrat-
 182 ing that the local Ca_l , Re_l are orders of magnitude higher than the corresponding av-
 183 erage values. The local dynamics become decoupled from the externally imposed flow
 184 rate, as energy stored in the system during slow drainage as surface energy is released,
 185 driving these abrupt events and the corresponding fluid redistribution. Finally conclu-
 186 sions drawn from this work are discussed in section 4.

187 2 Numerical Model

188 In this section we describe the numerical method we shall use, starting with the
 189 thermodynamics in section 2.1, the dynamical equations of motion in section 2.2 and the
 190 lattice Boltzmann implementation in section 2.3.

191 2.1 Thermodynamics of the fluid

192 We consider a binary fluid mixture containing two types of molecules, A and B.
 193 The equilibrium properties of the binary fluid can be described by a Landau free energy
 194 functional (Briant & Yeomans, 2004), in terms of an order parameter ϕ ,

$$\mathcal{F} = \int_V \left(f_b + \frac{\kappa_\phi}{2} (\partial_\alpha \phi)^2 \right) dV + \int_S f_s dS . \quad (1)$$

195 The order parameter is a scalar field, which describes the local molecular composition
 196 of the binary fluid mixture, defined as $\phi = (n_A - n_B)/(n_A + n_B)$. $n_{A,B}$ denotes the
 197 number density of the A, B molecules locally. For temperatures below a critical temper-
 198 ature T_c ($T < T_c$), the mixture undergoes phase separation into A-rich and B-rich do-
 199 mains, with the equilibrium values for the order parameter being $\phi_{eq} = \pm 1$. The first

200 term in the integrand is the bulk free energy density given by

$$f_b = \frac{a}{2}\phi^2 + \frac{b}{2}\phi^4 + \frac{c^2}{3}\rho \ln \rho, \quad (2)$$

201 where ρ is the fluid mass density and c is a lattice velocity parameter. This choice of f_b
 202 allows binary phase separation into two phases, if $a < 0$ and $b > 0$, with bulk equilib-
 203 rium solutions $\phi_{eq} = \pm(-a/b)^{1/2}$. Making the choice $a = -b$, leads to $\phi_{eq} = \pm 1$ for
 204 the bulk of the two phases. The position of the fluid-fluid interface is chosen to be the
 205 locus $\phi = 0$. The term in ρ does not affect the phase behavior and controls the com-
 206 pressibility of the fluid (Kendon et al., 2001).

207 The energetic cost of fluid-fluid interfaces is accounted for by the gradient term $(\kappa_\phi/2)(\partial_\alpha\phi)^2$,
 208 which penalises spatial variations of the order parameter ϕ by a factor κ_ϕ , for example
 209 across the fluid-fluid interface. Taking the functional derivative of the free energy with
 210 respect to ϕ gives the exchange chemical potential

$$\mu \equiv \frac{\delta\mathcal{F}}{\delta\phi} = a\phi + b\phi^3 - \kappa_\phi\partial_\gamma\gamma\phi, \quad (3)$$

211 which is constant at equilibrium (set $\mu = 0$), otherwise it would give rise to a thermo-
 212 dynamic force density $-\phi\partial_\alpha\mu$. Assuming for simplicity that the fluid-fluid interface is
 213 flat and located at $x = 0$, eq. 3 allows an interface solution of the form $\phi(x) = \tanh(x/(\sqrt{2}\xi))$,
 214 with $\xi = \sqrt{-\kappa_\phi/a}$ being the interface width (Briant & Yeomans, 2004). Throughout
 215 this work we fix $\xi = 0.81$; this was previously shown to give accurate results for the vari-
 216 ation of ϕ across the fluid-fluid interface (Kendon et al., 2001) and allows a smooth tran-
 217 sition between the two phases over a length scale of $\sim 5\xi$. The corresponding interfa-
 218 cial tension (excess free energy per unit area) is given by $\gamma = \sqrt{-8\kappa_\phi a^3/(9b^2)}$ (Briant
 219 & Yeomans, 2004).

220 The choice of free energy functional leads to a pressure tensor (Anderson et al., 1998)

$$\begin{aligned} P_{\alpha\beta} &= \left[\phi \frac{\delta\mathcal{F}}{\delta\phi} + \rho \frac{\delta\mathcal{F}}{\delta\rho} - \mathcal{F} \right] \delta_{\alpha\beta} + (\partial_\alpha\phi) \frac{\delta\mathcal{F}}{\delta(\partial_\beta\phi)} \\ &= \left[p_b - \kappa_\phi\phi\partial_\gamma\gamma\phi - \frac{\kappa_\phi}{2}(\partial_\gamma\phi)^2 \right] \delta_{\alpha\beta} + \kappa_\phi(\partial_\alpha\phi)(\partial_\beta\phi) \\ &= P^{iso}\delta_{\alpha\beta} + P_{\alpha\beta}^{chem}, \quad (4) \end{aligned}$$

221 where $p_b = \frac{c^2}{3}\rho + \frac{1}{2}a\phi^2 + \frac{3}{4}b\phi^4$ is the bulk pressure. $P_{\alpha\beta}$ consists of two terms: an isotropic
 222 contribution $P^{iso} = c^2 3\rho$ to ensure constant density and a ‘chemical’ pressure tensor
 223 contribution $P_{\alpha\beta}^{chem}$ (Kendon et al., 2001). The thermodynamics of the fluid is contained

224 in the later. It reflects the fact that in the presence of chemical potential gradients, a
 225 force density proportional to this will drive the system to equilibrium. This thermody-
 226 namic force density, $-\phi(\partial_\alpha\mu)$, can be expressed as the divergence of a ‘chemical’ pres-
 227 sure tensor $\phi(\partial_\alpha\mu) = \partial_\beta P_{\alpha\beta}^{chem}$.

228 The final term in the free energy functional, eq. 1, describes the interactions be-
 229 tween the fluid and the solid surface. Following (Cahn, 1977), the surface energy den-
 230 sity is taken to be of the form $f_s = -h\phi_s$, where ϕ_s is the value of the order parame-
 231 ter at the surface. Minimisation of the free energy gives an equilibrium wetting bound-
 232 ary condition (Briant & Yeomans, 2004)

$$\kappa_\phi \partial_\perp \phi = \frac{df_s}{d\phi_s} = -h. \quad (5)$$

233 The value of the parameter h (the surface excess chemical potential) is related to the equi-
 234 librium contact angle θ^{eq} via (Briant & Yeomans, 2004)

$$h = \sqrt{2\kappa_\phi b} \operatorname{sign} \left[\frac{\pi}{2} - \theta^{eq} \right] \sqrt{\cos \left(\frac{\alpha}{3} \right) \left\{ 1 - \cos \left(\frac{\alpha}{3} \right) \right\}}, \quad (6)$$

235 where $\alpha = \arccos(\sin^2 \theta^{eq})$ and the function sign returns the sign of its argument.

236 2.2 Equations of motion

237 The hydrodynamic equations for the system are the continuity, eq. 7, and the Navier-
 238 Stokes, eq. 8, equations for a nonideal fluid

$$\partial_t \rho + \partial_\alpha(\rho u_\alpha) = 0, \quad (7)$$

239

$$\partial_t(\rho u_\alpha) + \partial_\beta(\rho u_\alpha u_\beta) = -\partial_\beta P_{\alpha\beta} + \partial_\beta [\eta(\partial_\beta u_\alpha + \partial_\alpha u_\beta)], \quad (8)$$

240 where \mathbf{u} , \mathbf{P} , η are the fluid velocity, pressure tensor and dynamic viscosity respectively.

241 For a binary fluid the equations of motion are coupled with a convection-diffusion equa-
 242 tion,

$$\partial_t \phi + \partial_\alpha(\phi u_\alpha) = M_\phi \nabla^2 \mu, \quad (9)$$

243 that describes the dynamics of the order parameter ϕ . M_ϕ is a mobility coefficient (see
 244 Appendix A, eq. Appendix A).

245 2.3 Lattice Boltzmann method

246 The numerical algorithm we use to solve the equations of motion is based on the
 247 lattice Boltzmann method (Doolen, 1990; Benzi et al., 1992; S. Chen et al., 1992; Succi,

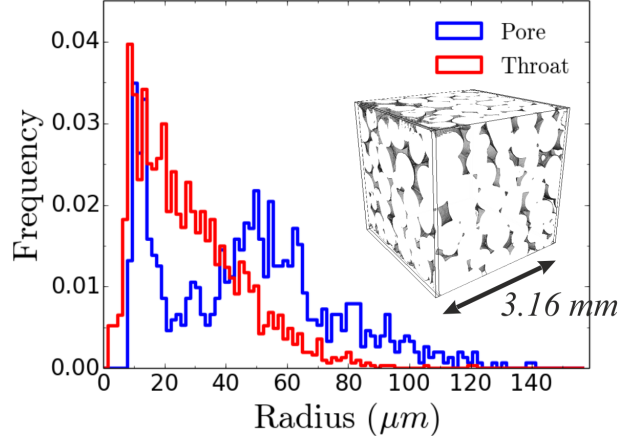
248 2001). In particular, we use an extension of the method, the free energy LB method, first
 249 introduced by Swift et al. (Swift et al., 1995, 1996), to simulate two-phase flow. This
 250 algorithm belongs to a class of hydrodynamic models, called diffuse interface models (Jacqmin,
 251 2000; Pierre & Seppecher, 1996; Briant et al., 2004; Briant & Yeomans, 2004), where the
 252 fluid-fluid interface has a finite size. Far away from a contact line, the method solves the
 253 hydrodynamic equations of motion of the fluid, i.e. the Navier-Stokes equations and the
 254 continuity equation. In the vicinity of the contact line, however, due to the finite size
 255 of the interface, the method introduces a diffusive mechanism, which regularizes the vis-
 256 cous dissipation singularity (de Gennes, 1985) and allows the contact line to slip on a
 257 solid substrate.

258 The above equations are solved using a Multiple Relaxation Time (MRT) lattice
 259 Boltzmann algorithm (D’Humières et al., 2002). Details of the implementation of the
 260 lattice Boltzmann algorithm are given in Appendix A. We note here that the numeri-
 261 cal code implementation is performed in CUDA C++ to take advantage of accelerated
 262 computing on multiple general-purpose graphics processing units (GPGPUs).

263 3 Results

264 3.1 Drainage simulations at fixed Ca_{av} and M

265 We examine here whether the classical treatment of characterizing two-phase flow
 266 (drainage) at the pore scale in terms of just the capillary number $Ca_{av} = \eta_{nw}\bar{u}/\gamma$, vis-
 267 cosity ratio $M = \eta_{nw}/\eta_w$ and wetting boundary conditions (θ^{eq}) (Lenormand et al., 1988)
 268 is sufficient to describe two phase flow. To this end we directly solve the hydrodynamic
 269 equations of motion in a three dimensional geometry reconstructed from micro-CT im-
 270 ages of Ketton limestone (Shah et al., 2016) and consider fluid flows at fixed average cap-
 271 illary number Ca_{av} and viscosity ratio ($\log M = 0$). We note here that \bar{u} is the aver-
 272 age nw phase velocity, obtained from the simulations and not the one obtained from the
 273 injection flow rate. The simulation system size is 700^3 lattice units (l.u) at a resolution
 274 of $4.52 \mu m$ per l.u. (physical system size $(3.16 \text{ mm})^3$). The fluid flow is driven by ap-
 275 plying a constant injection flow rate, which is achieved by applying velocity boundary
 276 conditions at the inlet/outlet of the simulation domain (Hecht & Harting, 2010). Small
 277 reservoirs are placed at the inlet/outlet of the simulation domain. The equilibrium con-



281 **Figure 2.** The pore and throat size distribution (Dong & Blunt, 2009) of the geometry
 282 used in the simulations (see inset). Inset: Ketton rock sample reconstructed from micro-CT
 283 images (Shah et al., 2017). The simulation system size is 700^3 lattice units (l.u) at a resolution of
 284 $4.52 \mu\text{m}$ per l.u, which corresponds to a physical system size of $(3.16 \text{ mm})^3$ with porosity 0.159.
 285 A small reservoir (16 l.u) is added at the inlet/outlet of the simulation domain.

278 tact angle is set to $\theta^{eq} = 40^\circ$, consistent with contact angle measurements in Ketton
 279 at reservoir conditions for a supercritical CO_2 - brine system (Andrew et al., 2014).

280 3.1.1 Intermediate Ca_{av}

286 For the case of intermediate capillary number (indicated in Fig. 1 with the green
 287 circle), we consider three simulations with different volumetric injection flow rates, $Q_{inj} =$
 288 $\int_A \mathbf{u}_{inj} \cdot d\mathbf{A}$, and fluid viscosities, while maintaining the same average capillary num-
 289 ber ($Ca_{av} = 3.3 \times 10^{-4}$) and viscosity ratio ($\log M = 0$). \mathbf{A} is the cross sectional area
 290 at the inlet/outlet. Fixing Ca_{av} is achieved by keeping the product $Q_{inj}\eta_{nw}$ and all other
 291 parameters (θ^{eq}, γ) constant. Simulation parameters (in lattice units - l.u) are listed in
 292 Table 1. The above choice essentially varies the Reynolds number ($Re_{av}^{(1)} = 1.0 \times 10^0$,
 293 $Re_{av}^{(2)} = 1.0 \times 10^{-2}$, $Re_{av}^{(3)} = 1.7 \times 10^{-3}$). Here we use the average invaded throat di-
 294 ameter as the characteristic length scale L_s in the definition of $Re_{av} = \rho \bar{u} L_s / \eta_{nw}$, as
 295 this controls the pressure at which pores drain, $L_s \sim 10$ l.u ($45 \mu\text{m}$). A useful dimen-
 296 sionless number that relates the capillary to Reynolds numbers is the Ohnesorge num-
 297 ber

$$Oh^2 = Ca_{av} / Re_{av} = \eta_{nw}^2 / \rho \gamma L_s, \quad (10)$$

301

Table 1. Simulation parameters in lattice units - Intermediate Ca_{av} test case*

Run	η_{nw}	u_{inj}
1	1.67×10^{-1} ($\tau_{f,nw}^{(1)} = 1.0$)	2×10^{-6}
2	6.67×10^{-2} ($\tau_{f,nw}^{(2)} = 0.7$)	5×10^{-6}
3	6.67×10^{-3} ($\tau_{f,nw}^{(3)} = 0.52$)	5×10^{-5}

*Simulations with fixed interfacial tension $\gamma = 1.17 \times 10^{-2}$ ($\kappa_\phi = 10^{-2}$, $a = -1.54 \times 10^{-2}$) and $\theta^{eq} = 40^\circ$.

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which describes the relative importance of viscous forces to inertial and interfacial tension forces. As we will show by the end of this section, Oh will prove useful in restricting the parameter selection process in terms of numerical modeling.

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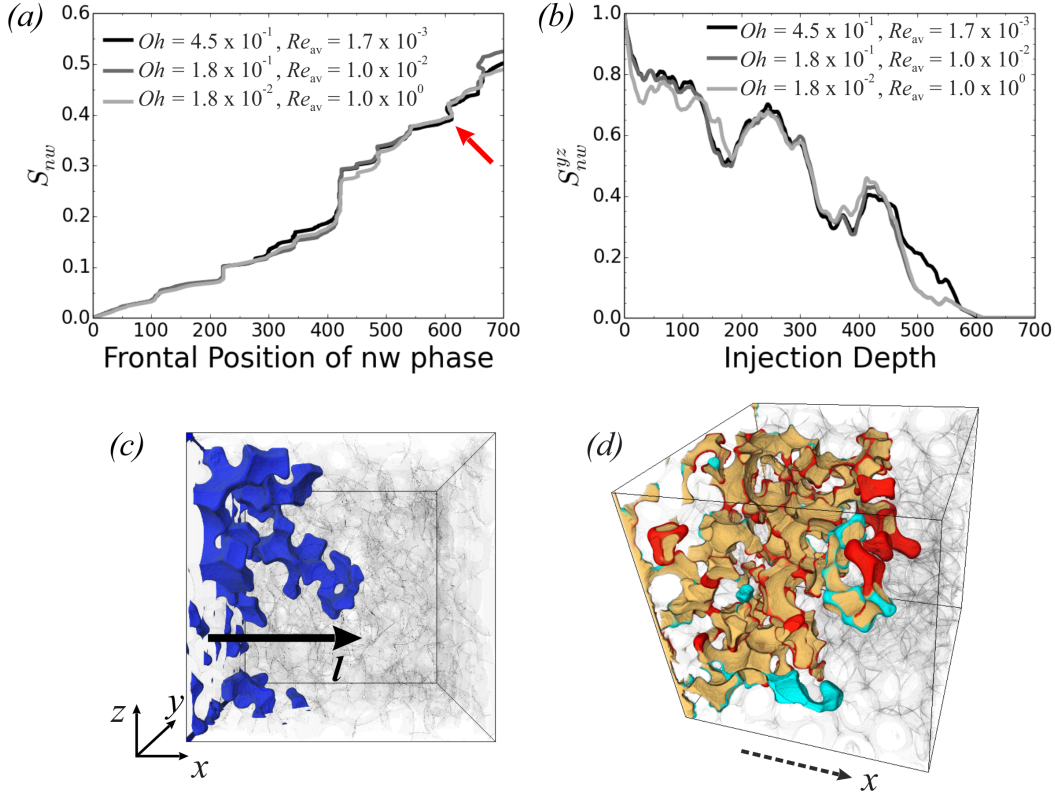
We start by examining how the injected non-wetting fluid saturates/fills the porous rock. Fig. 3(a) shows the non-wetting phase saturation S_{nw} as a function of the frontal position, which is defined as the distance of the most deeply penetrated non-wetting phase from the inlet, see Fig. 3(c). Although there is a significant degree of overlap, it is clear that results are not completely the same. More importantly, different non-wetting phase saturation is obtained at breakthrough $S_{nw}^{(br)}$ (frontal position $l = 700$). Given the overlap of the results, we further examine configurations with the same S_{nw} and frontal position, aiming at inspecting the fluids' topology in the pore space. Fig. 3(b) shows the non-wetting phase saturation at planes perpendicular to the mean direction of the flow S_{nw}^{yz} as a function of the injection depth (distance of the yz-plane from the inlet), when the total saturation $S_{nw} = 0.3975$ and the frontal position is $l = 612$ (indicated with the red arrow in Fig. 3(a)). This clearly reveals that the above simulations actually exhibit a different spatial distribution of the fluids, which means that the displacement pathways/sequence might be different. A comparison of the fluids' spatial distribution for the simulations with the smallest $Re_{av} = 1.7 \times 10^{-3}$, 1.0×10^{-2} is shown in Fig. 3(d); regions in light blue and red denote the regions occupied by the non-wetting phase only for the case $Re_{av} = 1.0 \times 10^{-2}$ and $Re_{av} = 1.7 \times 10^{-3}$ respectively, while regions in yellow are the overlapping regions in both simulations.

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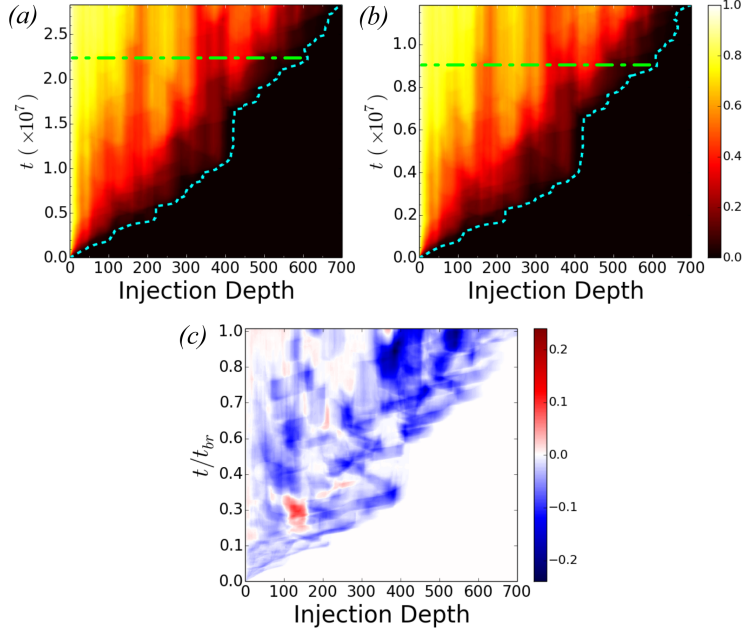
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Fig. 4 provides a more comprehensive picture of the fluids' distribution as it demonstrates the spatiotemporal plot of the non-wetting phase saturation at planes perpendicular to the mean direction of the flow S_{nw}^{yz} for (a) $Re_{av} = 1.7 \times 10^{-3}$ and (b) $Re_{av} =$



302 **Figure 3.** Drainage simulation results at fixed $Ca_{av} = 3.3 \times 10^{-4}$ ($\log M = 0$) and varying
 303 Re_{av}, Oh : (a) The non-wetting phase saturation S_{nw} as a function of the frontal position of the
 304 non-wetting phase. (b) The non-wetting phase saturation S_{nw}^{yz} at planes perpendicular to the
 305 mean direction of the flow for situations with the same total saturation $S_{nw} = 0.3975$ and frontal
 306 position $l = 612$ (indicated with the red arrow in (a)). (c) Definition of the frontal position.
 307 Non-wetting phase shown in blue, while rock grains and wetting phase are transparent. (d) Com-
 308 parison of the fluids' configuration for the simulations with $Re_{av} = 1.7 \times 10^{-3}$ and 1.0×10^{-2} .
 309 Regions in yellow are occupied by the non-wetting phase in both simulations; regions in light blue
 310 are occupied by the non-wetting phase only for the case $Re_{av} = 1.0 \times 10^{-2}$ and regions in red are
 311 occupied by the non-wetting phase only for $Re_{av} = 1.7 \times 10^{-3}$.



330 **Figure 4.** Spatiotemporal plot of the non-wetting phase saturation S_{nw}^{yz} for drainage sim-
 331 ulations at $Ca_{av} \sim 3 \times 10^{-4}$ ($\log M = 0$) and varying Re_{av} and Oh : (a) $Oh = 4.5 \times 10^{-1}$,
 332 $Re_{av} = 1.7 \times 10^{-3}$, (b) $Oh = 1.8 \times 10^{-1}$, $Re_{av} = 1.0 \times 10^{-2}$. The injection depth refers to the
 333 distance of each yz-plane, at which S_{nw}^{yz} is measured, from the inlet reservoir. The dashed line
 334 denotes the frontal position as a function of time, while the dash-dotted line corresponds to the
 335 data for S_{nw}^{yz} in Fig. 3(b). (c) The difference between (a) and (b) with red/blue denoting regions
 336 with higher S_{nw}^{yz} in (a)/(b). Time is scaled by the time it takes the non-wetting phase to reach
 337 the outlet, t_{br} .

341 1.0×10^{-2} . Although at a first glance these look similar, the differences between the two
 342 cases are shown in Fig. 4(c), where red (blue) denotes the regions with higher S_{nw}^{yz} for
 343 the case $Re_{av} = 1.7 \times 10^{-3}$ (a) ($Re_{av} = 1.0 \times 10^{-2}$ (b)). This suggests that the dis-
 344 placement sequence is different. The step-like structure that starts to emerge here for
 345 the frontal position (dashed line), as well as in Fig. 3(a) for S_{nw} as a function of time,
 346 is a characteristic feature of the capillary fingering regime (Tsuji et al., 2016).

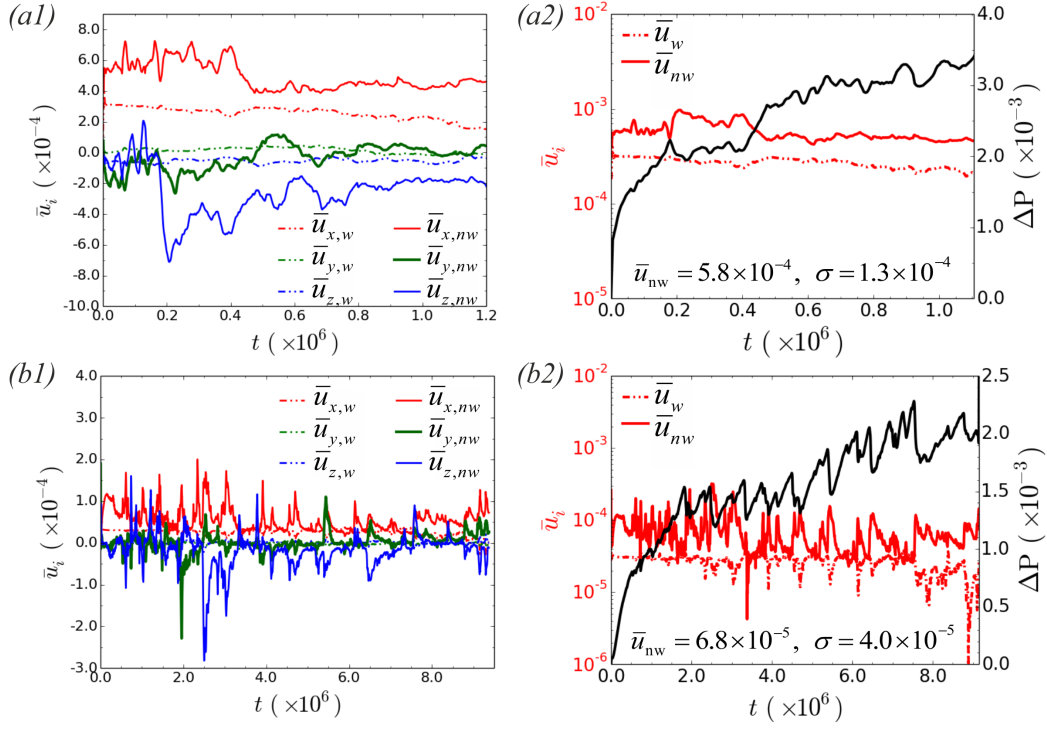
347 A reasonable question is then “Which one of these simulations is the one that best
 348 describes a given drainage situation?”. In terms of numerical modeling the question should
 349 be rephrased as “What are the relevant dimensionless numbers for the process and the
 350 flow regime under investigation?”. Matching these dimensionless numbers is essential in
 351 capturing the fluid flow behavior. In addressing these questions, it is useful to consider
 352 that the ratio of Ca_{av} to Re_{av} is given by the Ohnesorge number, $Oh^2 = Ca_{av}/Re_{av} =$
 353 $\eta_{nw}^2/(\rho\gamma L_s)$. Oh is a dimensionless number that quantifies the ratio of viscous forces to
 354 inertia and interfacial tension. Of particular importance and usefulness in terms of nu-
 355 merical modeling is the fact that the Ohnesorge number is independent of the flow rate
 356 (and the externally imposed way to drive the fluid flow). It reflects purely the thermo-
 357 physical properties of the fluids and the geometry/rock type (McKinley & Renardy, 2011).
 358 Therefore, experiments with given fluids and given geometry correspond to constant val-
 359 ues of Oh . For example considering the process of geological sequestration of CO_2 (sys-
 360 tem of CO_2 - brine), Oh varies in the range of $10^{-3} - 10^{-2}$, depending on the choice
 361 of the characteristic length-scale L_s of the system. This can be for example the mean
 362 radius of the pore throats in a drainage displacement, as the pore throats control the pres-
 363 sure at which pores drain.

364 Focusing on the case of geological sequestration of CO_2 , the important consequence
 365 of the above, with regards to numerical modeling, is that inertial effects cannot be ne-
 366 glected for two phase flow at the pore scale as the average Re_{av} is 4 to 6 orders of mag-
 367 nitude higher than the capillary number Ca_{av} ($Oh = \sqrt{Ca_{av}/Re_{av}} \sim 10^{-3} - 10^{-2}$).
 368 This is particularly important if we consider flow regimes with Ca_{av} in the range $10^{-7} -$
 369 10^{-3} and/or flow regimes characterized by sharp interfacial jumps (e.g. Haines jumps),
 370 when inertial effects become more profound. Here, for the simulations reported at in-
 371 termediate $Ca_{av} \sim 3 \times 10^{-4}$, the simulation that would best describe a CO_2 - brine
 372 system is the one with $Oh = 1.8 \times 10^{-2}$ ($Re_{av} = 1.0 \times 10^0$).

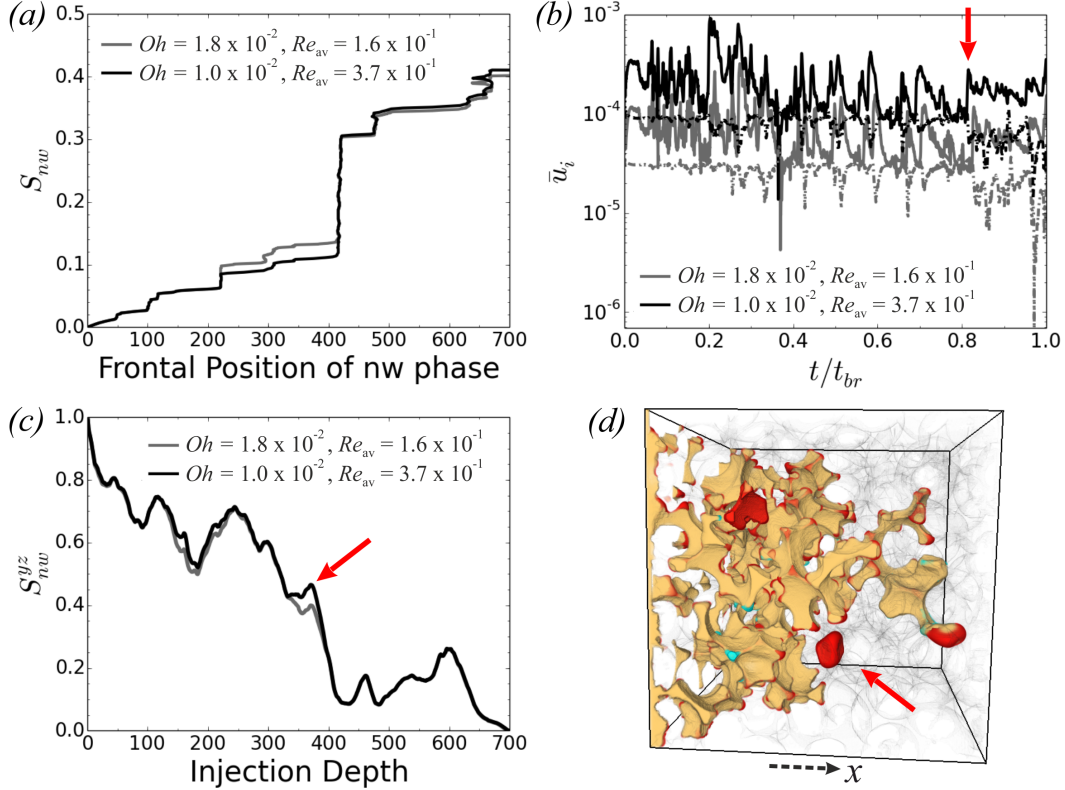
3.1.2 Low Ca_{av} - Haines jumps regime

Extending our investigation to the low Ca_{av} regime, we encounter a transition to a flow regime characterized by Haines jumps. This is evident from the flow field and the inlet-outlet pressure difference, shown in Fig. 5, where we compare simulations at $Oh = 1.8 \times 10^{-2}$ and varying Ca_{av} (intermediate versus low Ca_{av}). On the left panel we plot the average components of the velocity, while on the right panel the average magnitude of the velocity and the inlet-outlet pressure difference, $\Delta P = P_{inlet} - P_{outlet}$, is shown. A distinct change in both the flow field and the pressure difference is clearly seen: sharp interfacial jumps, indicative of Haines jumps, lead to significant increase of the non-wetting phase velocity and abrupt pressure drops, which coincide with the jumps. The transition to the capillary fingering regime with Haines jumps is also profound by examining the saturation of the injected phase, S_{nw} , as a function of the frontal position, see Fig. 3(a) versus Fig. 6(a). At $Ca_{av} = 3.9 \times 10^{-5}$ the step-like structure of S_{nw} versus frontal position is a consequence of consecutive forward and backward Haines jump events (Tsuji et al., 2016).

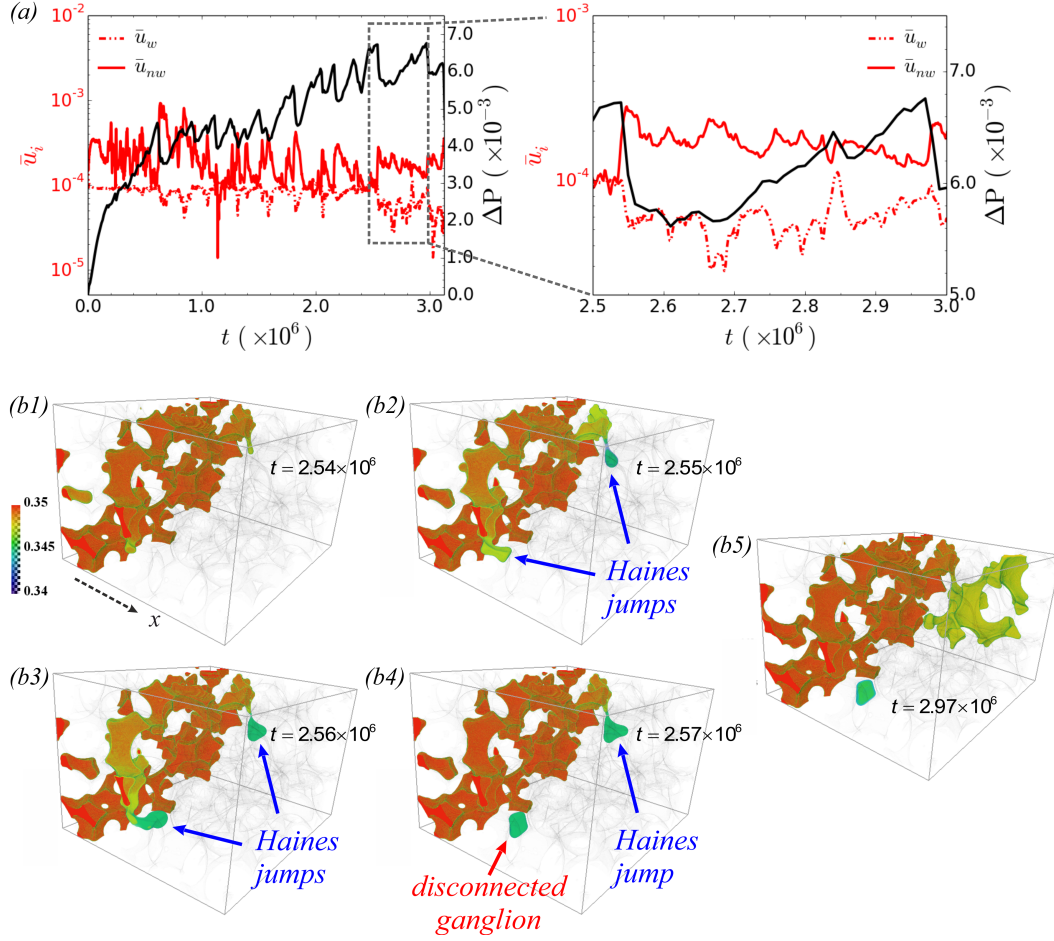
For the low Ca_{av} case (fixed average $Ca_{av} = 3.9 \times 10^{-5}$) we examine two situations with $Oh^{(1)} = 1.8 \times 10^{-2}$ ($Re_{av}^{(1)} = 1.6 \times 10^{-1}$) and $Oh^{(2)} = 1.0 \times 10^{-2}$ ($Re_{av}^{(2)} = 3.7 \times 10^{-1}$) by varying the interfacial tension and the injection flow rate, while the fluids' viscosity remains fixed ($\eta_{nw} = 6.67 \times 10^{-3}$). Simulation parameters are as follows: (1) $\gamma^{(1)} = 1.17 \times 10^{-2}$ ($\kappa_{\phi} = 1 \times 10^{-2}$, $a = -1.54 \times 10^{-2}$), $u_{inj}^{(1)} = 5 \times 10^{-6}$, (2) $\gamma^{(2)} = 3.51 \times 10^{-2}$ ($\kappa_{\phi} = 3 \times 10^{-2}$, $a = -4.62 \times 10^{-2}$) and $u_{inj}^{(2)} = 1.5 \times 10^{-5}$. Numerically it was not possible to examine situations with a much higher variation in Oh , like the intermediate Ca_{av} case (section 3.1.1) by varying the fluids' viscosity, as this would be computationally very expensive. Given that there is not a significant change in the Ohnesorge number (it could be argued that this corresponds to similar sets of experiments), it would be reasonable to expect that the results reported in Fig. 6 should look similar. The similarity observed in the flow field (Fig. 6(b)) is also an indication that the displacement pathways and the sequence of Haines jumps are the same in both cases. Comparing the fluids' configurations at breakthrough, see Fig. 6(c),(d), reveals that indeed the drainage displacement process is almost similar in both situations, except for mainly a disconnected ganglion (indicated with a red arrow), observed in the case $Oh^{(2)} = 1.0 \times 10^{-2}$ ($Re_{av}^{(2)} = 3.7 \times 10^{-1}$). Interestingly, despite the small difference in Oh , this is the result of distal (non-local) snap-off, that actually persists throughout the simulation time



388 **Figure 5.** Results from simulations with $Oh = 1.8 \times 10^{-2}$ and: (a) $Ca_{av} = 3.3 \times 10^{-4}$,
 389 (b) $Ca_{av} = 3.9 \times 10^{-5}$. Left panel: Average value for the components of the velocity for the
 390 wetting/non-wetting phases (dashed/solid lines). Right panel: The corresponding average magni-
 391 tude of the velocity and inlet-outlet pressure difference. The abrupt pressure drop and the inter-
 392 facial jumps are indicative of Haines jumps. The occurrence of Haines jumps, as Ca_{av} decreases,
 393 is reflected in the distribution of the velocities observed (higher ratio of standard deviation σ to
 394 mean velocity \bar{u}_{nw}).



395 **Figure 6.** Results from simulations at fixed $Ca_{av} = 3.9 \times 10^{-5}$ and varying Re_{av}, Oh . (a) The
 396 non-wetting phase saturation S_{nw} as a function of the frontal position of the non-wetting phase.
 397 (b) The average magnitude of the velocity versus time in scaled units $t^* = t/t_{br}$. Solid/dashed
 398 lines correspond to the average velocity of the non-wetting/wetting phase. (c) The distribution of
 399 the non-wetting phase per injection depth, quantified by S_{nw}^{yz} , at $t = t_{br}$. (d) Visualization of the
 400 non-wetting phase configurations at $t = t_{br}$. Regions in yellow are occupied by the non-wetting
 401 phase in both simulations; regions in light blue are occupied by the non-wetting phase only for
 402 the case $Re_{av} = 1.6 \times 10^{-1}$ ($Oh = 1.8 \times 10^{-2}$) and regions in red are occupied by the non-wetting
 403 phase only for $Re_{av} = 3.7 \times 10^{-1}$ ($Oh = 1.0 \times 10^{-2}$). The red arrows indicate the time of distal
 404 snap-off (b) and the corresponding disconnected non-wetting phase.



427 **Figure 7.** Results from simulations at $Ca_{av} = 3.9 \times 10^{-5}$ and $Oh = 1.0 \times 10^{-2}$. (a) The av-
 428 erage magnitude of the velocity versus time and the inlet/outlet pressure difference. Right panel:
 429 Zoom in at the times for a jump event that develops distal snap-off. (b) Visualizations of the
 430 non-wetting phase displacement that develops distal snap-off event (indicated with the red arrow
 431 in (b4)). The non-wetting phase is colored by the bulk pressure to demonstrate the pressure drop
 432 during Haines jump events.

423 and produces a long-lasting fluid configuration (as it remains disconnected) in agreement
 424 with the experimental observations of Andrew et al. (2015). Contrarily, local snap-off
 425 events generate disconnected fluid configurations that rapidly reconnect with the con-
 426 nected non-wetting phase region (Andrew et al., 2015).

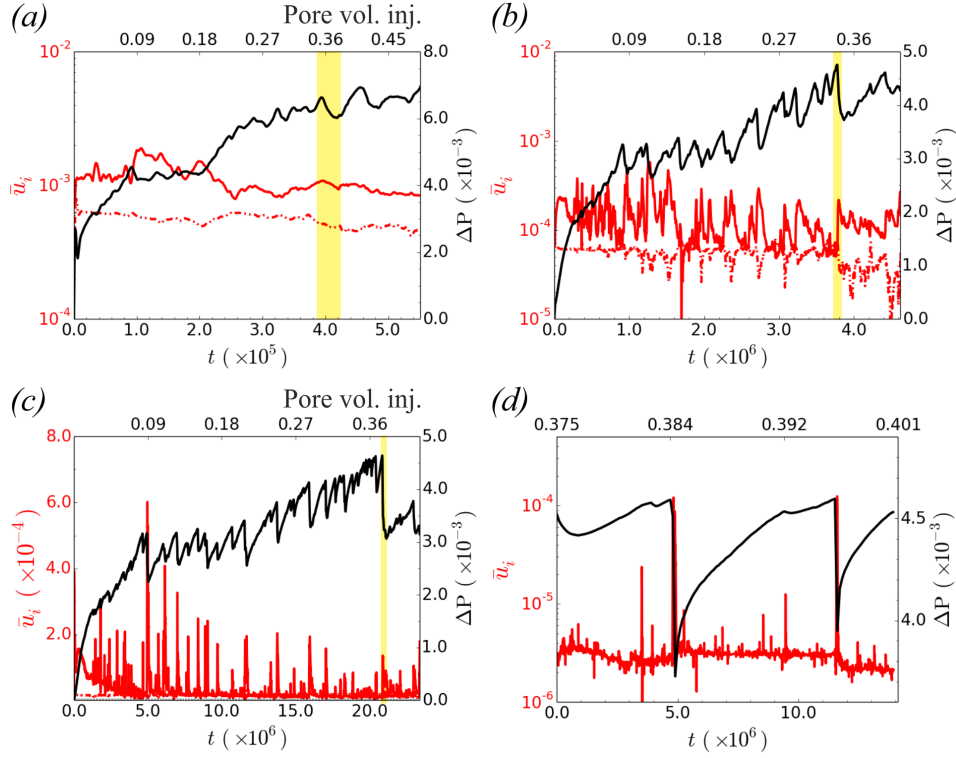
433 Fig. 7 shows that the distal snap-off event is caused by the transient low dynamic
 434 capillary pressure during the abrupt interface movement (high local Re_l), which is then
 435 preserved in the capillary pressure of the disconnected ganglion. For the case of $Oh^{(2)} =$

436 1.0×10^{-2} ($Re_{av}^{(2)} = 3.7 \times 10^{-1}$), the higher interfacial tension affects the dynamics of
 437 the Haines jump events, as the higher driving capillary forces result in larger interfacial
 438 velocities and pressure drop during the events. Consequently, the bigger reduction in cap-
 439 illary pressure observed, even in throats significantly away from the jump event, can cause
 440 the disconnection of the non-wetting phase (distal snap-off). Hence, Ca_{av} and M can-
 441 not fully describe a drainage displacement process, even in the low Ca_{av} regime (lower
 442 Re_{av}), since the choice of parameters can affect the dynamics of Haines jumps, in which
 443 case inertial effects become locally more profound. This results in not capturing the dy-
 444 namics of Haines jumps and the associated fluids' redistribution (Zacharoudiou et al.,
 445 2018). Especially in the case of jumps accompanied with distal snap-off, which can af-
 446 fect the drainage displacement process and the fluids' configuration, since distal snap-
 447 off has a persistent impact on the wetting phase flow field and the sequence of subsequent
 448 drainage events (Andrew et al., 2015). Moreover, distal snap-off can potentially have
 449 a negative impact on the displacement efficiency, which is an extremely significant fea-
 450 ture for CO₂ storage and hydrocarbon recovery, as it blocks the access to regions of the
 451 pore space that were accessible to the non-wetting phase prior the Haines jump event
 452 (Zacharoudiou et al., 2018).

460

461 **3.2 Local dynamics - Energy conversion during drainage in the low Ca_{av}** 462 **regime**

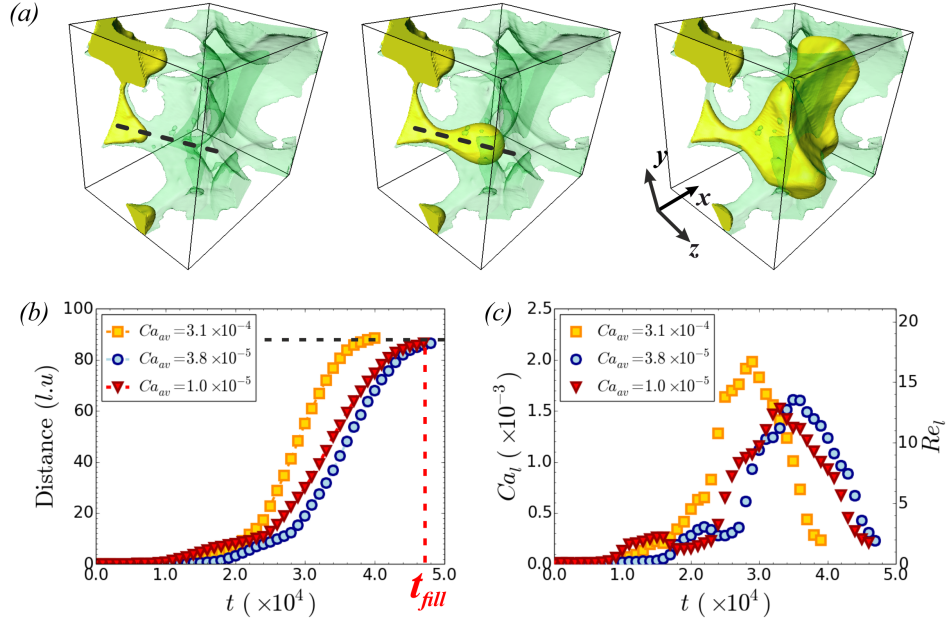
463 There is often the argument that two phase flow at the pore scale is at low Reynolds
 464 numbers and therefore inertial effects are not important. How small is the average flow
 465 Re_{av} has been addressed in the previous section by considering the system's Oh , which
 466 links Ca_{av} and Re_{av} . Another question to be addressed is what is the actual local Reynolds
 467 number Re_l , especially during jump events and as the average Ca_{av} (Re_{av}) decreases
 468 even further, e.g. far away from the injection point. Here we examine the local dynam-
 469 ics and the associated dimensionless numbers (Re_l , Ca_l), during low Ca_{av} flow. To this
 470 end we consider simulations with $\log M = 0$, $Oh = 1.3 \times 10^{-2}$ and varying Ca_{av} , see
 471 Fig. 8. This is achieved by keeping all parameters fixed ($\eta_i = 6.67 \times 10^{-3}$ ($i=w, nw$),
 472 $\gamma = 2.34 \times 10^{-2}$ ($\kappa_\phi = 2 \times 10^{-2}$, $a = -3.08 \times 10^{-2}$), $\theta^{eq} = 40^\circ$) and varying the injec-
 473 tion flow rate Q_{inj} ($u_{inj}^{(a)} = 10^{-4}$, $u_{inj}^{(b)} = 10^{-5}$, $u_{inj}^{(c)} = 2 \times 10^{-6}$, $u_{inj}^{(d)} = 2 \times 10^{-7}$). The
 474 sharp increase in non-wetting fluid velocity and the decrease in pressure signal observed



453 **Figure 8.** The average magnitude of the velocity for the wetting/non-wetting phase
 454 (dashed/solid red lines) and the inlet/outlet pressure difference for simulations with $\log M = 0$,
 455 $Oh = 1.3 \times 10^{-2}$ and varying Ca_{av} : (a) $Ca_{av} = 3.1 \times 10^{-4}$ ($u_{inj}^{(a)} = 10^{-4}$), (b) $Ca_{av} = 3.8 \times 10^{-5}$
 456 ($u_{inj}^{(b)} = 10^{-5}$), (c) $Ca_{av} = 1.0 \times 10^{-5}$ ($u_{inj}^{(c)} = 2 \times 10^{-6}$) and (d) $Ca_{av} = 1.0 \times 10^{-6}$ ($u_{inj}^{(d)} = 2 \times 10^{-7}$).
 457 The yellow shaded regions indicate the time scales of a jump event analysed in Fig.9. Simulation
 458 (d) results were obtained from restarting simulation (c) at time $t^{(c)} = 20.08 \times 10^6$ and decreasing
 459 the injection flow rate from $u_{inj}^{(c)} = 2 \times 10^{-6}$ to $u_{inj}^{(d)} = 2 \times 10^{-7}$.

475 for Ca_{av} smaller than $\sim 10^{-5}$, Fig. 8(b)-(d), indicates flow regime with Haines jumps.

476



477 **Figure 9.** (a) Pore drainage event over multiple geometrically defined pore spaces. The non-
 478 wetting invading fluid is shown in yellow. The rock and wetting phase are shown transparent
 479 and semi-transparent respectively for effective visualization. (b) The distance traveled by the
 480 fluid-fluid interface in the pore space versus time for the simulations reported in Fig. 8(a)-(c).
 481 The pore filling time (t_{fill}) can be approximated by the time the interface reached the distance
 482 plotted with the horizontal dashed line. (c) The corresponding local dimensionless numbers Ca_l
 483 and Re_l using the interfacial velocity obtained from the data in (b). Time $t = 0$ denotes the start
 484 of the event.

485 We focus on a particular drainage event, see Fig. 8(a)-(c) for the time frames con-
 486 sidered and Fig. 9 for the draining site. Results presented in Fig. 9(b)-(c) reveal that the
 487 jump dynamics (interfacial velocity, pore filling time) in the capillary fingering regime
 488 ($Ca_{av} \sim 10^{-5}$) are similar despite a decrease by a factor of 5 in the externally imposed
 489 injection flow rate ($u_{inj} = 10^{-5} - Ca_{av} = 3.8 \times 10^{-5}$, $u_{inj} = 2 \times 10^{-6} - Ca_{av} = 1.0 \times$
 490 10^{-5}), in agreement with previous observations (Armstrong & Berg, 2013; Zacharoudiou
 491 & Boek, 2016; Li et al., 2019). Moreover, the local maximum $Ca_l \sim 10^{-3}$ and $Re_l \sim$
 492 1.5×10^1 are independent of and orders of magnitude higher than the corresponding aver-
 493 average flow values. This holds in general for the jump events (not shown here). Decreas-

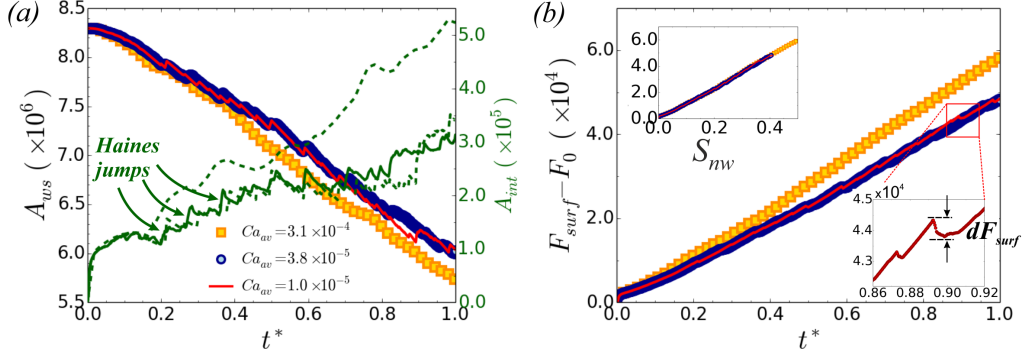
494 ing Ca_{av} to 10^{-6} , see simulation reported in Fig. 8(d), reveals local maximum Reynolds
 495 number of the order $Re_l \sim 10^1$. The above indicate that a further decrease in Ca_{av} will
 496 not affect the local dynamics (Ca_l, Re_l) significantly, but rather this becomes decoupled
 497 from the externally imposed flow rate. More importantly, the local value of Re_l indicates
 498 that inertial effects are important and the numerical scheme used should solve the full
 499 Navier-Stokes equations, honoring momentum balance, and all relevant forces, i.e., cap-
 500 illary, viscous, and inertial forces acting simultaneously in order to capture the dynam-
 501 ics and the fluid rearrangement during these abrupt events.

502 The decoupling of local dynamics from the externally imposed flow rate can be un-
 503 derstood by considering the energy conversion during drainage. The externally performed
 504 work of pressure, $W_p = \int \Delta P Q_{inj} dt$, drives the fluid flow, hence converted to kinetic
 505 energy E_k , and stores energy in the system as surface energy F_{surf} . The energy balance
 506 states (Ferrari & Lunati, 2014)

$$\frac{dW_p}{dt} - \Phi = \frac{dE_k}{dt} + \frac{dF_{surf}}{dt}, \quad (11)$$

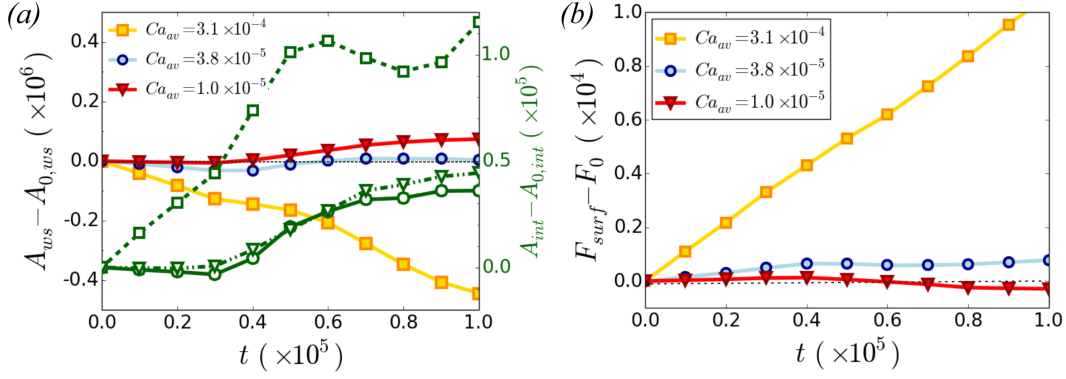
507 where the energy losses are accounted for by the viscous dissipation rate Φ . The change
 508 in surface energy is given by $dF_{surf} = \gamma dA_{int} + \gamma_{ws} dA_{ws} + \gamma_{ns} dA_{ns}$, where dA_{int} , dA_{ws}
 509 and dA_{ns} are the increments of the areas of the fluid-fluid, solid-wetting fluid and solid
 510 - non wetting fluid interfaces respectively and γ , γ_{ws} , γ_{ns} the corresponding surface ten-
 511 sions. Since $dA_{ns} = -dA_{ws}$ and the total solid surface area $A_{tot}^s = A_{ns} + A_{ws}$ is con-
 512 stant, the change in surface energy can be expressed as $dF_{surf} = \gamma(dA_{int} - \cos \theta^{eq} dA_{ws})$.
 513 In a drainage situation $dA_{ws} < 0$ and $dA_{int} > 0$; hence, F_{surf} is expected to increase
 514 monotonically, except during Haines jumps. In this case, the energetic cost of newly cre-
 515 ated interfaces ($dA_{int} > 0$) can be balanced due to fluid rearrangement and the released
 516 energy due to wetting at the imbibition sites ($dA_{ws} > 0$), leading to a decrease in the
 517 overall F_{surf} , see Fig. 10.

526 Fig. 10 presents results for the time evolution of the interfacial areas and the cor-
 527 responding change in surface energy for the simulations reported in Fig. 8(a)-(c). The
 528 transition to the capillary fingering regime with Haines jumps ($Ca_{av} \sim 10^{-5}$), as the
 529 externally imposed flow rate decreases, becomes evident as results demonstrate a higher
 530 overlapping degree, but also by examining the monotonicity of the A_{ws} curve in Fig. 10(a).
 531 High Ca_{av} flow (smooth flow) leads to a monotonic decrease of A_{ws} (increase of F_{surf});
 532 non-monotonicity emerges with the onset of Haines jumps (burst flow). During the jump



518 **Figure 10.** (a) The time evolution of interfacial areas (wetting phase - solid, A_{ws} , fluid-fluid,
 519 A_{int}) for simulations with $\log M = 0$, $Oh = 1.3 \times 10^{-2}$ and varying Ca_{av} . Time is normalised
 520 by the breakthrough time, $t^* = t/t_{br}$. The occurrence of Haines jumps is marked with a sharp
 521 increase in A_{ws} (imbibition in surrounding pore throats) and A_{int} . Results for A_{int} (green lines):
 522 $Ca_{av} = 3.1 \times 10^{-4}$ (dashed line), $Ca_{av} = 3.8 \times 10^{-5}$ (solid line), $Ca_{av} = 1.0 \times 10^{-5}$ (dashed-
 523 dotted line). (b) The corresponding surface energy $F_{surf} - F_0$, where $F_0 = \gamma_{ws} A_{tot}^s$ is constant.
 524 Inset: (lower right) Surface energy released during Haines jumps; (upper left) surface energy data
 525 collapse versus non-wetting phase saturation.

533 events, the wetting phase-solid interfacial area A_{ws} increases, releasing energy, previously
 534 stored in the system, ($dF_{surf}^{rel} = \gamma \cos \theta^{eq} dA_{ws} > 0$) necessary for accelerating the flu-
 535 ids, see sharp peaks in Fig. 8(b)-(c), and creating new interfaces ($dF_{surf}^{int} = \gamma dA_{int} >$
 536 0). The overlap of the results $F_{surf}(t^*)$ for $Ca_{av} \sim 10^{-5}$ indicates that the distribu-
 537 tion of fluids and the displacement sequence is similar despite the decrease in Ca_{av} from
 538 3.8×10^{-5} to 1.0×10^{-5} . The change in injection flow rate affects the rate at which
 539 menisci are “charged”, but the “discharging” (Haines jumps) is controlled by the amount
 540 of energy available (dF_{surf}^{rel}), the viscous dissipation rate, the connectedness of the flu-
 541 ids and the structure of the porous rock. As the externally imposed flow rate decreases,
 542 the rate at which energy is provided to the system decreases (dW_p/dt) and, therefore,
 543 during the time scales of jump events, the energy balance (eq. 3.1.2) essentially becomes
 544 $-\frac{dF_{surf}}{dt} - \Phi = \frac{dE_k}{dt}$. Given that most of the nw phase needed for the draining of the
 545 pore body comes from fluid redistribution (Berg et al., 2013; Zacharoudiou et al., 2018),
 546 it explains why the local dynamics during jump events, at the same site, becomes de-
 547 coupled from the externally imposed flow rate, as Ca_{av} decreases beyond a certain limit.
 548 See also Fig. 11, where results for the time evolution of the interfacial areas and the sur-



557 **Figure 11.** (a) The time evolution of interfacial areas (wetting phase - solid A_{ws} , fluid-fluid
 558 A_{int}) during the event shown in Fig. 9(a) and simulations with $\log M = 0$, $Oh = 1.3 \times 10^{-2}$
 559 and varying Ca_{av} . $A_{0,ws}$ and $A_{0,int}$ denote the interfacial areas at the start of the event. Re-
 560 sults for A_{int} (green lines): $Ca_{av} = 3.1 \times 10^{-4}$ (dashed line), $Ca_{av} = 3.8 \times 10^{-5}$ (solid line),
 561 $Ca_{av} = 1.0 \times 10^{-5}$ (dashed-dotted line). (b) The corresponding surface energy $F_{surf} - F_0$, where
 562 F_0 is the reference surface energy at the start of the event.

549 face energy, during the event in Fig. 9(a), start to converge as Ca_{av} decreases from $Ca_{av} =$
 550 3.1×10^{-4} to 1.0×10^{-5} . Further decrease of the injection flow rate will make the above
 551 more profound. Here examining flow regimes with Ca_{av} smaller than 10^{-6} was not possi-
 552 ble, as this would have been computationally very costly, given the resources available
 553 for this research (8 NVIDIA Tesla P100 GPUs). We must note though, that, to the best
 554 of our knowledge, the capillary number flow regimes examined numerically in this manuscript
 555 are the lowest reported so far in the literature employing a direct numerical simulation
 556 approach.

563 4 Conclusions

564 We examine here whether the description of two-phase flow (drainage) in terms of
 565 just the capillary number and the viscosity ratio and neglecting inertial effects, based
 566 on the pioneering work of Lenormand et al. (1988) and its extension to fully developed
 567 drainage by Yortsos et al. (1997), is sufficient to describe the flow at the pore scale cor-
 568 rectly. This investigation has obvious implications for large scale numerical modeling of
 569 CO_2 geological sequestration and enhanced oil recovery. Especially, considering the fact
 570 that the acceleration of numerical algorithms using GPGPUs can enable the study of two
 571 phase flow at the pore-scale at smaller capillary numbers Ca_{av} than what was possible

572 so far. We essentially examine whether direct numerical simulation approaches need to
 573 solve the full Navier Stokes equations and consider all relevant forces, i.e. capillary, vis-
 574 cuous and inertial forces, acting simultaneously, in order to capture the physics of fluid-
 575 fluid displacement and interfacial phenomena at the pore scale.

576 Given the range of the relevant dimensionless numbers in porous media flows, and
 577 especially the ratio of viscous to capillary forces, $10^{-10} < Ca_{av} < 10^{-3}$, depending
 578 on the distance from the injection point in the well bore (Blunt & Scher, 1995), we ex-
 579 amined two test cases at intermediate ($Ca_{av} \sim 10^{-4}$) and low capillary numbers ($Ca_{av} \sim$
 580 10^{-5}). A distinct change in the flow regime is observed between the above two situations,
 581 as the later is characterised by abrupt jumps in the location of the fluid-fluid interface
 582 (Haines jumps), expected for low Ca_{av} flow. Our investigation clearly shows that iner-
 583 tial effects cannot be neglected in neither of the two flow regimes, i.e. we cannot exam-
 584 ine the fluid flow as being in the limit of zero Reynolds number. Generally two-phase
 585 flow at the pore scale is assumed to be at low Re_{av} . How small, though, is Re_{av} , can be
 586 answered by examining the Ohnesorge number ($Oh^2 = Ca_{av}/Re_{av}$). This dimension-
 587 less number is fixed for a given system and reflects the thermophysical properties of the
 588 specific system under investigation, i.e. the fluids' properties and the length scale depen-
 589 dencies originating from the porous medium geometry. Considering that Oh is typically
 590 in the range of 10^{-3} - 10^{-2} for a system of brine- CO_2 at the pore scale, i.e. Re_{av} is 4 to
 591 6 orders of magnitude higher than Ca_{av} , it becomes clear that the usual approach in nu-
 592 merical simulations of keeping both Ca_{av} and Re_{av} low, without respecting the ratio of
 593 the two, is fundamentally wrong, especially for the intermediate Ca_{av} flow regime.

594 As Ca_{av} and Re_{av} decrease further, inertial effects are still important over a tran-
 595 sient amount of time during abrupt jump events (Haines jumps), when the non-wetting
 596 phase passes from a narrow restriction to a wider pore body. Our results demonstrate
 597 that local jump dynamics become decoupled from the externally imposed flow rate, and
 598 locally Ca_l and Re_l are orders of magnitude higher than the average corresponding val-
 599 ues. Moreover, the displacement sequence as well as the fluids' distribution in the porous
 600 rock can be affected significantly by the choice of the simulation parameters. Therefore,
 601 matching the system's Oh , as well as using a Navier-Stokes solver are essential in resolv-
 602 ing the fluid dynamics during these abrupt events and capturing the fluids' spatial dis-
 603 tribution/connectedness, given that significant fluid rearrangement takes place. Other-
 604 wise the predicting capabilities of the numerical scheme should be questioned. In terms

605 of numerical modeling, fixing Oh based on fluid properties and characteristic length scales,
 606 provides a convenient way of restricting the parameter selection process.

607 Finally, it would be interesting to examine how the consideration of: (a) Haines jumps
 608 and inertial effects occurring on a single and/or multiple pore levels and (b) snap-off events,
 609 both distally and locally, alter in a macroscopic sense the invasion percolation aspects
 610 of a drainage process, e.g. sign and gradient of percolation probability, or the transition
 611 between the macroscopically observed flow regimes, as captured by the Lenormand phase
 612 diagram (Lenormand et al., 1988) and in its extension to fully-developed drainage by Yortsos
 613 et al. (1997). This remains a challenge left for future studies.

614 Appendix A Lattice Boltzmann Method

Here we provide details on the implementation of the free energy lattice Boltzmann
 (Swift et al., 1995, 1996; Briant & Yeomans, 2004; Krüger et al., 2017) algorithm we use
 to solve the hydrodynamic equations of motion, the continuity and Navier Stokes equa-
 tions coupled to an advection diffusion equation for the order parameter (composition).
 The dynamics are introduced by discretised Boltzmann equations for two sets of distri-
 bution functions, $f_i(\mathbf{r}, t)$, $g_i(\mathbf{r}, t)$, which are related to the local fluid density $\sum_i f_i(\mathbf{r}, t) =$
 ρ and order parameter $\sum_i g_i(\mathbf{r}, t) = \phi$ Collision step: $f'_i(\mathbf{r}, t) = f_i(\mathbf{r}, t) + \Omega_{f,i}(\mathbf{r}, t)$,

$$g'_i(\mathbf{r}, t) = g_i(\mathbf{r}, t) + \Omega_{g,i}(\mathbf{r}, t) ,$$

$$\text{Propagation step: } f_i(\mathbf{r} + \mathbf{e}_i \Delta t, t + \Delta t) = f'_i(\mathbf{r}, t) ,$$

615 $g_i(\mathbf{r} + \mathbf{e}_i \Delta t, t + \Delta t) = g'_i(\mathbf{r}, t)$, The distribution functions are discrete in
 616 time and space, with the time step Δt and lattice spacing Δx being set to unity. The
 617 subscript i refers to the discrete set of velocity directions $\{\mathbf{e}_i\}$. Here we use a three di-
 618 mensional model with 19 discrete velocity vectors (D3Q19). Eq. Appendix A states that
 619 the time evolution of the distribution functions proceeds as follows: a) a collision step
 620 described by the collision operator $\Omega_{f/g,i}$ and b) a propagation step with velocity \mathbf{e}_i to
 621 the neighbouring lattice point $\mathbf{r} + \mathbf{e}_i \Delta t$ at the next time step $t + \Delta t$. In a concise form,
 622 the lattice Boltzmann equation, for f_i for example, becomes $f_i(\mathbf{r} + \mathbf{e}_i \Delta t, t + \Delta t) = f_i(\mathbf{r}, t) +$
 623 $\Omega_{f,i}(\mathbf{r}, t)$.

$$\text{The collision operators } \Omega_{f,i} = -M^{-1}SM[f_i(\mathbf{r}, t) - f_i^{eq}(\mathbf{r}, t)] ,$$

$$\Omega_{g,i} = -\frac{\Delta t}{\tau_g}[g_i(\mathbf{r}, t) - g_i^{eq}(\mathbf{r}, t)] , \text{ relax the distribution functions to their equilib-}$$

rium values (f_i^{eq}, g_i^{eq}) with relaxation time scales $\tau_{f,j}$ ($j = w, nw$) and τ_g . These re-

laxation times are related to the transport coefficients, dynamic viscosity η_j ($j = w, nw$) and mobility M_ϕ in the hydrodynamic equations via $\eta_j = \rho c^2 \Delta t (\tau_{f,j} - 1/2)/3$,

624 $M_\phi = \Delta t \Gamma (\tau_g - 1/2)$, where $c = \Delta x / \Delta t$ and Γ is a tunable parameter in the
 625 equilibrium distribution functions g_i^{eq} , see eq. A4. By expanding the chemical potential
 626 μ in powers of $\phi - \phi_{eq}$, for small deviations from equilibrium,

$$\mu = (a + 3b\phi_{eq}^2)(\phi - \phi_{eq}) + O((\phi - \phi_{eq})^2), \quad (A1)$$

627 the diffusive term $M_\phi \nabla^2 \mu$ in the convection-diffusion eq. 9 can be written as $D \nabla^2 \phi$ with
 628 $D = M_\phi (a + 3b\phi_{eq}^2)$ the diffusion coefficient. We note here that it is also possible to
 629 define the mobility coefficient M_ϕ to be a function of the order parameter ϕ in such a
 630 way that is restricting diffusion to the vicinity of fluid-fluid interfaces (F. O. Alpak et
 631 al., 2019).

632 In order to improve accuracy and stability, we adopt a Multiple Relaxation Time
 633 (MRT) (D’Humières et al., 2002) approach for the evolution of the distribution functions
 634 f_i^s , associated with the fluid density ρ , meaning that different relaxation rates are adopted
 635 for different linear combinations of the distribution functions (moments of f_i^s). The ma-
 636 trix M performs a transformation, so that collisions are performed in moment space. First
 637 the f_i^s are mapped to moment space, then the moments are relaxed towards equilibrium,
 638 and finally the relaxed moments are mapped back to population space. The information
 639 for the relaxation rates, $\omega = 1/\tau$, for the different linear combinations of the distribu-
 640 tion functions (moments of f_i^s) is contained in the relaxation matrix S . Following (C. Poo-
 641 ley et al., 2009), the relaxation times responsible for generating the viscous terms in the
 642 Navier-Stokes equation are set to τ_f , those related to conserved quantities to infinity and
 643 all the others, which correspond to non-hydrodynamic modes, to unity. On the other hand,
 644 a single relaxation time approximation is sufficient for the g_i^s , given that the mobility
 645 coefficient M_ϕ in Eq. eq:advection_diffusion_eqcanbetunedbytheindependentparameter Γ
 646 through eq. eq:Mobility. This allows us to fix the relaxation time $\tau_g = 1$. As shown by
 647 C. Pooley et al. (2009), this approach suppresses spurious currents at the contact line,
 648 while improving significantly the numerical stability of the algorithm as well (Lallemand
 649 & Luo, 2000). Throughout this work we choose a relative small value for $\Gamma = 1.0$ and
 650 refer the reader to the work by Ledesma-Aguilar et al. (2007) on how this (Γ and con-
 651 sequently the diffusion coefficient D) affects the contact line motion. Given the low Ca_{av}
 652 flow examined in this work, any wetting films adhered to the side walls are very small
 653 in size (less than 1% of the channel width) (Ledesma-Aguilar et al., 2007). Hence, in-

654 vaded pore spaces are fully occupied by the injected non-wetting phase and the value
 655 for Γ is not affecting the results.

656 The moments of the distribution functions are related to the physical quantities,
 657 mass density ρ , momentum density $\rho \mathbf{u}$ and composition

$$\sum_{i=0}^{18} f_i = \rho, \quad \sum_{i=0}^{18} f_i e_{i\alpha} = \rho u_\alpha, \quad \sum_{i=0}^{18} g_i = \phi. \quad (\text{A2})$$

658

659 The equilibrium distributions functions are defined as a power series in the veloc-
 660 ity (C. M. Pooley & Furtado, 2008)

$$\begin{aligned} f_i^{eq} = & w_i c^2 \left(p_b - \kappa_\phi \nabla^2 \phi + e_{i\alpha} \rho u_\alpha + 32c^2 [e_{i\alpha} e_{i\beta} - c^2 3\delta_{\alpha\beta}] \rho u_\alpha u_\beta \right) \\ & + \kappa_\phi c^2 \left(w_i^{xx} \partial_x \phi \partial_x \phi + w_i^{yy} \partial_y \phi \partial_y \phi + w_i^{zz} \partial_z \phi \partial_z \phi \right. \\ & \left. + w_i^{xy} \partial_x \phi \partial_y \phi + w_i^{xz} \partial_x \phi \partial_z \phi + w_i^{yz} \partial_y \phi \partial_z \phi \right), \quad (\text{A3}) \end{aligned}$$

661

$$g_i^{eq} = w_i c^2 \left(\Gamma \mu + e_{i\alpha} \phi u_\alpha + 32c^2 [e_{i\alpha} e_{i\beta} - c^2 3\delta_{\alpha\beta}] \phi u_\alpha u_\beta \right), \quad (\text{A4})$$

with the coefficients (C. M. Pooley & Furtado, 2008) $w_{1-6} = 16$, $w_{7-18} = 112$,

$$w_{1,2}^{xx} = w_{3,4}^{yy} = w_{5,6}^{zz} = 512,$$

$$w_{3-6}^{xx} = w_{1,2,5,6}^{yy} = w_{1-4}^{zz} = -13,$$

$$w_{7-10}^{xx} = w_{15-18}^{xx} = w_{7-14}^{yy} = w_{11-18}^{zz} = -124,$$

$$w_{11-14}^{xx} = w_{15-18}^{yy} = w_{7-10}^{zz} = 112,$$

$$w_{1-6}^{xy} = w_{1-6}^{yz} = w_{1-6}^{xz} = 0,$$

$$w_{7,10}^{xy} = w_{11,14}^{yz} = w_{15,18}^{xz} = 14,$$

$$w_{8,9}^{xy} = w_{12,13}^{yz} = w_{16,17}^{xz} = -14,$$

$$w_{11-18}^{xy} = w_{7-10}^{yz} = w_{15-18}^{yz} = w_{7-14}^{xz} = 0. \quad \text{The values for } i = 0 \text{ are chosen to}$$

conserve the local mass and composition $f_0^{eq} = \rho - \sum_{i \neq 0} f_i$,

662

$$g_0^{eq} = \phi - \sum_{i \neq 0} g_i, \quad \text{The above choice for the coefficients in eqs. eq:f}_{e}q_{i}l_{i}d_{i}str -$$

663

eq : def_{e}q_{d}istr_g is not unique; it was shown thought to reduce the unphysical currents, called spurious velocities, that appear

664

and the Laplacian ($\nabla^2 \phi$) that appear in the equilibrium distribution functions and the

665

chemical potential, using stencils (discrete operators).

The hydrodynamic equations of motion, continuity eq:continuity, Navier-Stokes eq:Navier-Stokes and convection diffusion equation eq:advection_{d}if_{f}usion_{e}q, can be obtained by performing a Chapman -

Enskog expansion (Luo, 2000) on the discretised Boltzmann equations (eq. eq : evolution_distr_funct), while the following

$$\rho, \quad \sum_{i=0}^{18} f_i^{eq} e_{i\alpha} = \rho u_\alpha, \quad \sum_{i=0}^{18} f_i^{eq} e_{i\alpha} e_{i\beta} = P_{\alpha\beta} + \rho u_\alpha u_\beta,$$

$$\sum_{i=0}^{18} f_i^{eq} e_{i\alpha} e_{i\beta} e_{i\gamma} = \frac{\rho c^2}{3} (u_\alpha \delta_{\beta\gamma} + u_\beta \delta_{\alpha\gamma} + u_\gamma \delta_{\alpha\beta}),$$

$$\sum_{i=0}^{18} g_i^{eq} = \phi, \quad \sum_{i=0}^{18} g_i^{eq} e_{i\alpha} = \phi u_\alpha, \quad \sum_{i=0}^{18} g_i^{eq} e_{i\alpha} e_{i\beta} = \Gamma \mu \delta_{\alpha\beta} + \phi u_\alpha u_\beta.$$

Finally we would like to point out that the free energy lattice Boltzmann method is capable of handling high viscosity ratios up to 10^3 . For validation of the numerical method we refer the reader to the work reported in (Zacharoudiou & Boek, 2016; Zacharoudiou et al., 2017; F. O. Alpak et al., 2019). This covers the dynamics of capillary filling, demonstrating that the method can capture the correct dynamics of imbibition in the limits of short and long time scales (different regimes for the imbibition length Vs time), as well as for varying viscosity ratio (we considered viscosity ratios $M = \eta_{nw}/\eta_w$ in the range $10^{-3} \leq M \leq 1$) (Zacharoudiou & Boek, 2016). The scaling of the dynamic contact angle with the interface velocity (capillary number) was found to be in excellent agreement with Cox theoretical prediction (Cox, 1986) for both 2D simulations (Zacharoudiou & Boek, 2016) and 3D channels with a rectangular cross-section (Zacharoudiou et al., 2017). The method was also shown to correctly predict fluid connectivity in imbibition in Gildehauser sandstone and simulate relative permeability data in close agreement with results from Darcy-scale core flooding experiments (F. Alpak et al., 2018). Further validation of the method investigating snap-off in constricted capillary tubes, Haines jumps and capillary desaturation on real-rock systems is reported in (F. O. Alpak et al., 2019).

A1 Boundary Conditions

In the lattice Boltzmann method we encounter three types of boundary conditions: a) the no-slip boundary condition on the velocity field b) the wetting boundary condition, eq. 5, on the composition and c) the boundary conditions at the inlet-outlet of the simulation domain that is related to how the fluid flow is driven.

The no-slip boundary condition, referring to the condition at solid boundaries where the fluid has zero velocity relative to the boundary, is implemented using the midlink

692 bounce-back method proposed by Ladd and Verberg (2001). Incoming populations are
 693 reflected back towards the lattice nodes they came from, resulting in recovering the wall
 694 location (zero velocity) half-way between the fluid and solid node.

Regarding the implementation of the wetting boundary condition, eq. $\phi = \phi_w - c \cos \theta$ at the substrate in equilibrium. Using this condition and the values for the composition ϕ at the fluid nodes neighbouring the boundary we assign the appropriate values for ϕ at the solid boundary nodes following Niu et al. (2007). The main advantage of this method is that the terms $\partial_\alpha \phi$ and $\nabla^2 \phi$, needed for the evaluation of f_i^{eq} and g_i^{eq} in eqs. $f_i^{eq} = \phi_i - \phi$, $g_i^{eq} = \phi_i - \phi$ can be calculated.

695 Finally, we choose to drive the fluid flow by applying a constant injection flow rate,
 696 i.e. applying velocity boundary conditions at the inlet/outlet of the simulation domain.
 697 In particular we adopt the approach proposed by Hecht and Harting (2010) to two-phase
 698 flow to estimate the missing populations (f_i , g_i) at the inlet/outlet domain boundaries.
 699 The above choice of driving the fluid flow enables the study of the low Ca_{av} flow regime
 700 characterised by Haines jumps, which wouldn't be possible in an alternative scenario of
 701 using a body force (pressure gradient) to drive the fluid flow (Zacharoudiou et al., 2018).
 702

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