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# Laboratory Experiments on Convective Entrainment using a Saline Water Tank

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<sup>6</sup> Abstract Entrainment fluxes in a shear-free convective boundary layer have been measured

7 with a saline water tank set-up. The experiments were targeted towards measuring the en-

<sup>8</sup> trainment behaviour for medium to high Richardson numbers and use a two-layer design, i.e.

<sup>9</sup> two stacked non-stratified (neutral) layers with different densities. With laser induced fluo-

<sup>10</sup> rescence (LIF), the entrainment flux of a fluorescent dye is measured for bulk Richardson

numbers in the range 30-260. It is proposed that a carefully chosen combination of top-down and bottom-up processes improves the accuracy of LIF-based entrainment observations. The

<sup>13</sup> observed entrainment fluxes are about an order of magnitude lower than reported for ther-

mal water tanks: the derived buoyancy entrainment ratio, A, is found to be  $A \approx 0.02$ , which

is to be compared with  $A \approx 0.25$  for a thermal convection tank (Deardorff et al., J. Fluid

<sup>16</sup> Mechanics, 1980, Vol.100, 41–64). An extensive discussion is devoted to the influence of

17 the Reynolds and Prandtl numbers in laboratory experiments on entrainment.

<sup>18</sup> Keywords Convective boundary layer · Entrainment · Laboratory experiments

# 19 **1 Introduction**

20 For weather, climate and air quality purposes it is of key importance to quantify the deepen-

<sup>21</sup> ing rate of a turbulent atmospheric boundary layer into the overlaying free troposphere. The

22 turbulence can originate from a variety of sources, most notably shear and buoyancy. His-

<sup>23</sup> torically, laboratory experiments targeted to mimic such situations in a downscaled setting,

<sup>24</sup> have played an important role in understanding and quantifying the process of deepening

<sup>25</sup> mixed layers, in particular the process of entrainment. Two classical examples are the water

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 (1974), Deardorff et al. (1980) (deepening by surface buoyancy).

The advantage of studying entrainment characteristics in the laboratory, rather than out-28 doors, is that 1) one has control over the initial and boundary conditions, the homogeneity 29 of the surface and the external forcings; 2) the experiments (if conducted well) are repro-30 ducible; 3) one can perform parameter studies keeping everything else constant. These as-31 pects, after all, form the cornerstone of scientific research and are not easily realized in 32 atmospheric measurement studies. The major downside of laboratory experiments is their 33 limited scale, or more precise, the fact that the attainable Reynolds and Péclet numbers are 34 orders of magnitude lower than in the geophysical case (the Péclet number is the ratio be-35 tween the convective contribution and the diffusive contribution to the flux). We come back 36 to this important aspect in Sect. 2 and in the Discussion (Sect. 5). This issue notwithstand-37 ing, in particular the thermal convection tank results on penetrative entrainment by Deardorff 38 et al. (1980) have had a big impact in boundary-layer meteorology. As the non-dimensional 39 entrainment rate was found to follow an inverse Richardson law, viz.  $ARi^{-1}$  with  $A \approx 0.25$ , 40 41 the result appeared to comply with the early views on convective entrainment (Ball 1960; Tennekes 1973) and with atmospheric observations (e.g. Stull 1976, and references therein). 42 This explains why laboratory experiments have formed an important benchmark case for 43 large-eddy simulation (LES) studies (e.g. Nieuwstadt et al. 1993; Sullivan et al. 1998; Fe-44 dorovich et al. 2004). 45 Still there are a few unresolved issues on convective entrainment, in particular the be-46 haviour for large Richardson numbers. Experiments by Turner (1968) showed that, apart 47 from a  $Ri^{-1}$  law, also a  $Ri^{-3/2}$  law can be observed, which would be even more likely for 48 high Péclet numbers such as in the atmosphere. In this context it is interesting that Deardorff 49 et al. (1980) presented their results in a subtle way: "It [the  $Ri^{-1}$  law] appears to represent 50 our data slightly better than the  $Ri^{-\frac{3}{2}}$  law". On the other hand, experiments on convective en-51 trainment by Kantha (1980) in a different set-up revealed a 'regime change' at a Richardson 52 number of order  $10^2$ . 53 Our study is therefore aimed at exploring the entrainment characteristics for relatively 54 high Richardson numbers. To this end we use a saline water tank and conduct experiments in 55 a so-called two-layer arrangement - two stacked non-stratified layers with different densities 56 (Kantha et al. 1977; Kantha 1980). Such an arrangement enables one to create a well-defined 57 buoyancy jump (Richardson number), and to choose settings such that the Richardson num-58 ber remains constant during an experiment. There are several advantages of a saline water 59 tank over a thermal tank (Hibberd and Sawford 1994a,b): there is, for example, no temper-60 ature leakage at the boundaries due to conduction or radiation, and the diffusivity of salt is 61 more than two orders smaller than the diffusivity of heat, which enables one to retain station-62 ary density jumps (see also Sect. 2). A drawback is that the surface buoyancy flux is realized 63 through applying a surface mass flux; as a result, the mixed layer deepens not only due to 64 entrainment but also due to the added mass. On the other hand the advantage of a surface 65 mass flux is that one can readily conduct 'bottom-up' diffusion experiments (Wyngaard and 66 Brost 1984) by adding a dye to the surface inflow - a feature that will be exploited herein 67

together with 'top-down' diffusion.

#### 69 2 Governing equations

<sup>70</sup> In order to be able to readily compare convection in the saline tank with its geophysical

<sup>71</sup> counterpart, as well as with convection in a thermal tank, it is most convenient to formulate

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 $\rho_0$  is a reference density and g is the acceleration due to gravity. Within the Boussinesq approximation and in the absence of the Coriolis force, the governing equations can be written as

$$\frac{\partial u_i}{\partial t} = -u_j \frac{\partial u_i}{\partial x_j} - \frac{1}{\rho_0} \frac{\partial p}{\partial x_i} + v \frac{\partial^2 u_i}{\partial x_j^2} + b \delta_{i3}, \tag{1}$$

$$\frac{\partial u_i}{\partial x_i} = 0,\tag{2}$$

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$$\frac{\partial b}{\partial t} = -u_j \frac{\partial b}{\partial x_j} + D \frac{\partial^2 b}{\partial x_j^2},\tag{3}$$

where  $u_i$  {i = 1, 2, 3} denotes the velocity in the x, y, z directions, respectively, with the z-78 direction pointing upwards, p denotes the pressure,  $\delta$  is the Kronecker symbol, v is the 79 kinematic viscosity of the carrier fluid (water/air) in  $m^2 s^{-1}$  and D is the molecular diffu-80 sivity of the quantity related to buoyancy (i.e. temperature/salinity) in  $m^2 s^{-1}$ . In the case of 81 convection in the atmospheric boundary layer, b is usually expressed in terms of the virtual 82 potential temperature  $\theta_v$  through  $b = g(\theta_v - \theta_{v0})/\theta_{v0}$  (e.g. Stull 1988). In a thermal con-83 vection tank one has  $b = g\alpha(T - T_0)$ , with T the temperature and  $\alpha$  the thermal expansion 84 coefficient of water. In a saline convection tank, on the other hand, density differences arise 85 from the salinity fraction S, i.e.  $\rho(S) = \rho_0(1+S)$ , with  $\rho_0$  is the density of pure water. The 86 buoyancy is thus related to salinity via 87

$$b = -gS. \tag{4}$$

<sup>88</sup> In the shear-free convective boundary layer, turbulence is driven by a surface buoyancy flux <sup>89</sup> *B*. The other relevant aspect defining the case is the initial condition. Instead of the com-<sup>90</sup> monly applied linear stratification (constant density gradient), we consider here a so-called <sup>91</sup> *two-layer set-up* (Kantha 1980; Deardorff et al. 1980), i.e. two stacked neutrally stratified <sup>92</sup> layers separated by an initial buoyancy jump  $\Delta b$  at initial height  $h(0) = h_0$ , see Fig. 1. Due <sup>14</sup> to entroice the neutral back to will increase

 $_{93}$  to entrainment the mixed-layer height *h* will increase

$$w_{\rm e} = \frac{{\rm d}h}{{\rm d}t}$$

where  $w_e$  is the entrainment rate. The key question is how  $w_e$  can be expressed in terms of the strength of the buoyancy jump (inversion), and the surface buoyancy flux. Equations 1-3

show that the full problem is characterized by 1) the boundary condition B; 2) the initial

conditions  $h_0$ ,  $\Delta b$ ; and 3) the fluid properties v, D. In addition the tank geometry (heightwidth) may play a role.

<sup>99</sup> Combining these parameters one can define the convective velocity scale (Deardorff 100 1970)

$$v_* = (Bh)^{1/3} \tag{5}$$

and the large eddy turnover time  $t_* = h/w_*$ . Rescaling the equations using  $w_*$  for velocity,  $t_*$  for time, *h* for length,  $\rho_0 w_*^2$  for pressure, and  $b_* = B/w_*$  for buoyancy, one obtains

$$\frac{\partial u_i}{\partial t} = -u_j \frac{\partial u_i}{\partial x_j} - \frac{\partial p}{\partial x_i} + \frac{1}{Re} \frac{\partial^2 u_i}{\partial x_j^2} + \delta_{i3}b, \qquad (6)$$

$$\frac{\partial b}{\partial t} = -u_j \frac{\partial b}{\partial x_j} + \frac{1}{Pe} \frac{\partial^2 b}{\partial x_i^2}.$$
(7)

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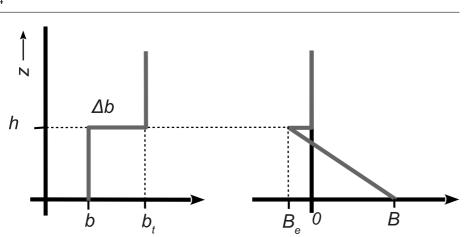


Fig. 1 Zero-order representation of the profiles of buoyancy and the buoyancy flux.

The initial condition, characterized by the buoyancy jump  $\Delta b$ , gives, after rescaling by  $b_*$ ,

the bulk Richardson number  $Ri = \Delta b/b_* = \Delta b h/w_*^2$ . If one includes the tank geometry, one

<sup>106</sup> can conclude that in essence the problem is characterized by four dimensionless numbers,

viz. the Reynolds number, the Péclet number, the Richardson number, and the tank aspect
 ratio

$$Re = \frac{w_*h}{v}, \qquad Pe = \frac{w_*h}{D}, \qquad Ri = \frac{\Delta bh}{w_*^2}, \qquad \Gamma = \frac{h}{L},$$
 (8)

where *L* is the lateral size. The non-dimensional entrainment rate  $E = w_e/w_*$  is some function of these dimensionless numbers

$$\frac{w_{\rm e}}{w_{\star}} = E(Re, Pe, Ri, \Gamma). \tag{9}$$

Instead of the  $\{Re, Pe\}$  combination, one could also use  $\{Re, Pr\}$  to define the situation, where Pr = v/D is the Prandtl number. Actually, in the case of salinity, it would be more appropriate to refer to v/D as the Schmidt number, but following e.g. Wolanski and Brush (1975) or Sayler and Breidenthal (1998) hereafter we refer to v/D as the Prandtl number both for thermal and saline situations. In laboratory experiments aimed to mimic geophysical phenomena it is usually pre-

<sup>116</sup> sumed that the Reynolds number and the Péclet number are *large enough* not to have an <sup>117</sup> sumed that the Reynolds number and the Péclet number are *large enough* not to have an <sup>118</sup> important effect on the results (e.g. Townsend 1980; Wyngaard 2010). Molecular properties <sup>119</sup> of the fluid and geometric details of the set-up then play a minor role. In Sect. 5 we come <sup>120</sup> back to the important issue of Reynolds-number independence, because it is not easy to <sup>121</sup> know *a priori* how large the Reynolds number should be to reach this regime. If the regime <sup>122</sup> is reached, the non-dimensional entrainment rate is only dependent on the Richardson num-<sup>123</sup> ber, generally assumed via a power-law relation

$$E(Ri) = aRi^{-b} \tag{10}$$

with a and b constants. As mentioned in the Introduction, the accepted values for atmo-

spheric convection are a = 0.2...0.25 and b = 1. The -1 Richardson-number law is consistent with the common assumption in mixed-layer models (Tennekes 1973) that the entrainment buoyancy flux  $B_e$  is a fixed fraction A of the surface buoyancy flux B (Ball 1960).

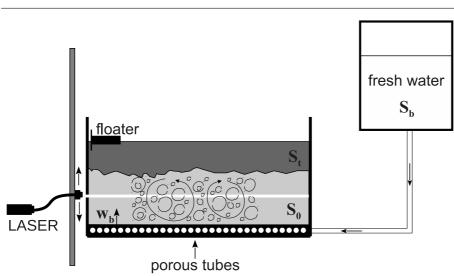


Fig. 2 Sketch of the saline convection tank set-up.

Indeed, invoking a zero-order model (Lilly 1968) for the entrainment flux  $B_e = -w_e \Delta b$ , using Eq. 10 with the Richardson number defined by Eq. 8, and Eq. 5, gives  $A = -B_e/B = a$ ; see also the sketch in Fig. 1.

In this study we focus on determining *A* rather than  $w_e$  because we are particularly interested in the behaviour at high Richardson numbers. The entrainment rate  $w_e$  might become very small for high *Ri* but, for a  $Ri^{-1}$  law, the value of *A* should be independent of the Richardson number. We emphasize here that, if the entrainment rates in the saline convection tank set-up display an entrainment law with an exponent *b* different from 1, then the measured values of *A* will show a Richardson dependence

$$A(Ri) = -B_{\ell}/B = E(Ri) Ri = aRi^{1-b}.$$
(11)

For example, a  $Ri^{-3/2}$  law would lead to A vanishing according to  $\sim 1/\sqrt{Ri}$ . In addition, if the present set-up introduces a Reynolds/Péclet number dependence in the entrainment rate

Eq. 9, then this will also show up in the flux ratio, i.e.  $A = A(Re, Pe, Ri, \Gamma)$ .

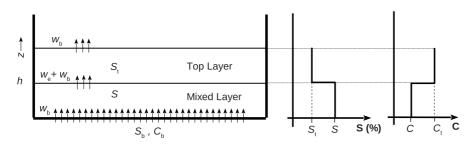
# 140 **3** Experimental set-up and measurement methodology

<sup>141</sup> 3.1 Description of the saline tank set-up

The saline convection tank set-up that we used is been sketched in Fig. 2. It consists of a 142 1m×1m glass container with height 0.5m, a scanning rail, a 2.5W Argon laser connected 143 via an optical fibre to the scanning rail, a digital 10-bit camera located in front of the tank, 144 an elevated water reservoir, and at the bottom of the convection tank a tray consisting of 49 145 parallel porous tubes (Gardena, external diameter 17 mm). The porous tubes were covered 146 with a carefully levelled bed ( $\approx$  10 mm) of small aquarium quartz pebbles (small reddish 147 stones a few mm in size) to even the surface. Another important experimental issue concerns 148 the spatial homogeneity of the surface fluxes. Tests of the homogeneity of the surface fluxes 149 are described later in this section. 150

Quantity	Symbol	Typical value	Unit
mixed-layer height (initial)	$h_0$	0.1	m
bottom inflow	wb	$2 \cdot 10^{-5}$	${ m m~s^{-1}}$
salinity inflow	Sb	0.0	%
salinity mixed layer (initial)	$S_0$	0.5	%
salinity top layer	St	0.0 - 0.4	%
dye concentration inflow	$C_{\rm b}$	0.0	$\mu \mathrm{g}\mathrm{l}^{-1}$
dye concentration mixed layer (initial)	$C_0$	10	$\mu$ g l $^{-1}$
dye concentration top layer	$C_{\rm t}$	50 - 200	$\mu$ g l $^{-1}$
convective velocity	$W_*$	$5 \cdot 10^{-3}$	m s <sup>-1</sup>
eddy turnover time	$t_*$	20	8
inversion strength (initial)	Ri	50 - 250	-
viscosity	ν	$1 \cdot 10^{-6}$	$m^{2} s^{-1}$
saline diffusivity	D	$1 \cdot 10^{-9}$	$m^{2} s^{-1}$

Table 1 Typical values of parameters used in the set-up.



**Fig. 3** Schematic view of the two layers separated by a buoyancy jump (left) and corresponding salinity and dye concentration profiles (right). The top water level increases by the inflow rate  $w_b$  whereas the mixed layer also grows by entrainment, i.e. at a rate  $w_e + w_b$ .

All entrainment experiments have been conducted within the context of a so-called two-151 layer set-up (e.g. Kantha et al. 1977; Kantha 1980) such as schematically depicted in Fig. 152 3. This initial state was created by first filling the tank through the porous tubes with well-153 mixed water with salinity  $S_0$  up to a height  $h_0 \approx 0.1$  m. Next a layer with lower salinity  $S_t$  was 154 added on top of the mixed layer resulting in a buoyancy jump of  $\Delta b = g[S_0 - S_t]$ . Placing 155 this layer on top was done by a careful (slow) filling procedure that consisted of dripping 156 water on top of four floating devices (not depicted in the sketch) that were submerged just 157 below the water surface. Due to the design of these floaters, only a flow in a lateral direction 158 occurs. In this way very sharp buoyancy jumps could be created, which by virtue of the 159 small diffusivity of salinity  $(D \approx 10^{-9} \text{m}^2 \text{ s}^{-1})$  remained sharp until the experiment was 160 started. For example, the time for the diffusion region to become  $\approx 5\%$  of the initial mixed-161 layer height,  $\delta = 0.05h_0$ , can be estimated by  $\delta^2/D \approx 7$ hr, i.e. much longer than the typical 162 duration of an experiment (15min). This highlights one of the advantages of the saline set-up 163 over a thermal tank set-up for which the mentioned time scale amounts to only 3min. 164

After the initial two-layer profile had been created, convection was initiated by supplying water with low salinity  $S_b$  from the elevated large reservoir to the porous tubes; the pressure difference due to the elevation is large enough to overcome the hydrodynamic resistance of the porous tubes and gives rise to a homogenous inflow velocity  $w_b$  of about  $2 \times 10^{-5} \text{m s}^{-1}$ . As will be detailed later, this inflow velocity is more than two orders of magnitude smaller than the convective velocity  $w_*$  and therefore has a negligible effect on the flow in terms of momentum. But the inflow of low density fluid has a strong effect in terms of buoyancy, the resulting surface buoyancy flux being

$$B = gw_{\rm b}[S(t) - S_{\rm b}] \tag{12}$$

where S(t) denotes the salinity in the mixed layer. Initially  $S(t = 0) = S_0$ , but during the course of the experiment *S* is diluted due to both the surface and the entrainment salinity flux, which entails that the surface buoyancy flux *B* decreases during the experiment. This effect will be accounted for in the analysis of the results. It is also important to note that the mixed-layer height h(t) increases not only due to entrainment  $w_e$  but also due to the surface mass flux  $w_b$ ,

$$\frac{\mathrm{d}}{\mathrm{d}t}h = w_{\mathrm{e}} + w_{\mathrm{b}}.\tag{13}$$

<sup>179</sup> So it is essential to very accurately measure  $w_b$  in the experiments because it directly controls <sup>180</sup> both the mixed-layer height and the surface buoyancy flux. We return to this particular issue <sup>181</sup> in the next section.

Table 1 gives an idea of the typical values we used in the experiments. Using Eq. 12 and 182 Eq. 5 one finds a surface buoyancy flux of  $B \approx 10^{-6} \text{ m}^2 \text{ s}^{-3}$  and a convective velocity scale 183 of  $w_* \approx 5 \times 10^{-3}$  m s<sup>-1</sup>, respectively, which indicates that the inflow velocity  $w_b$  is indeed 184 much smaller than the convective velocity:  $w_*/w_b \approx 250$ . The Reynolds number and Péclet 185 number based on  $w_*$  (see Eq. 8) are Re = 500 and  $Pe = 5 \times 10^5$ , respectively. The Reynolds 186 number is thus of comparable order as in the thermal convection tank of Deardorff et al. 187 (1980); the Péclet number is however two orders of magnitude larger in the saline set-up 188 due to the much lower diffusivity of salt compared to heat. We address these issues in more 189 detail in the Discussion. At this stage we merely point out that the Reynolds number does 190 not change much during an experiment. Indeed, whereas the increasing mixed-layer depth 191 tends to increase the Reynolds number during the experiments, this effect is largely offset by 192 the decreasing surface buoyancy flux resulting from the gradually decreasing mixed-layer 193 salinity concentration. The net effect is only a slight increase of the Reynolds number during 194 the course of the experiments. 195

The inversion strength is controlled by the salinity  $S_t$  in the upper layer for which we take values ranging between zero and 0.4%. The corresponding Richardson numbers follow from Eq. 8 and can be written as

$$Ri = \frac{\Delta b}{b_*} = \frac{w_*}{w_b} \frac{S - S_t}{S - S_b}.$$
(14)

<sup>199</sup> Since *S* decreases during an experiment, so the Richardson number changes. The interesting <sup>200</sup> exception is when  $S_t = S_b$  for which the Richardson number is nearly constant (because  $w_*$ <sup>201</sup> is nearly constant).

#### 202 3.2 Measuring entrainment with LIF

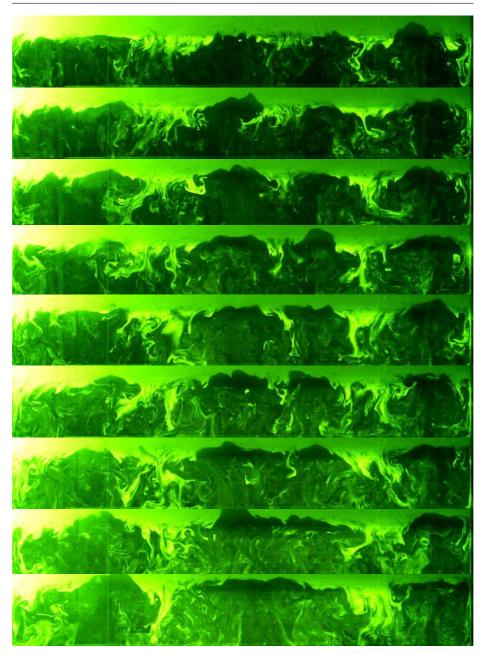
In order to find the entrainment rate  $w_e$  at given Richardson numbers, it would seem natural from Eq. 13 to try and directly measure dh/dt and subsequently subtract the measured value of  $w_b$ . But such a method fails for moderate to strong inversions because  $w_e$  is itself very small and is to be diagnosed from the difference between two relatively large terms, inevitably leading to inaccuracies. For this reason we focus herein on measuring entrainment fluxes rather than entrainment velocities. In the simplest form one can think of adding dye

only to the top layer and diagnosing entrainment from the increasing dye concentration in 209 the mixed layer. An example is given in Fig. 4. These images were not used in the analysis 210 of the results for reasons that will be detailed in the next section, but they nicely show a 211 number of features. One can, for example, clearly make out the cusp-shaped structures of 212 top-layer dye being entrained into the mixed layer; one also sees (in darker shades) the buoy-213 ant plumes that impinge on the inversion. Furthermore one can notice the gradual deepening 214 of the mixed layer and that the mixed layer is getting 'greener' due to the entrainment of the 215 dye from the overlaying layer. Such a setting was termed 'top-down diffusion' by Wyngaard 216 and Brost (1984); see also van Dop et al. (2005); Jonker et al. (1999). The advantage of such 217 a setting is that when the entrainment velocity  $w_e$  is very low, one can always 'boost' the 218 entrainment flux of the dye,  $-w_e(C_t - C)$ , by using larger dye concentrations  $C_t$  in the top 219 layer. In this way one can adjust the entrainment flux of the dye such that it can be measured 220 with desired accuracy. We use planar laser induced fluorescence (Ferrier et al. 1993; Snyder 221 et al. 2002; van Dop et al. 2005) for measuring the dye concentration in the mixed layer, 222 i.e. by using a fluorescent dye (disodium fluorescein C<sub>20</sub>H<sub>10</sub>O<sub>5</sub>Na<sub>2</sub>) that is illuminated by a 223 224 laser sheet.

We explored a variety of techniques to create a planar laser sheet, such as cylindrical 225 lenses, rotating polygon mirrors, parabolic mirrors, in an effort to create a homogeneous 226 light distribution over the vertical extent of the tank (0.5 m). The best performance in this 227 respect was obtained by directing the laser beam through an optic fibre, the end of which 228 was mounted on a scanning rail that moved up and down in a programmable fashion. The 229 produced narrow vertical laser sheet excited the fluorescent dye in the tank and the resulting 230 fluorescence image was captured by the digital camera located perpendicular to the sheet at 231 about 2 m distance. An optical filter in front of the camera was used to select the wavelengths 232 corresponding to fluorescence and filter out the wavelengths emitted by the Argon laser, 233 which might reach the camera through scattering 234

The camera was triggered by the position of the fibre on the transport rail, which enabled 235 us to obtain two types of images: (i) an image of the mixed layer only (type I), and (ii) an 236 image depicting the top layer and the mixed layer together (type II). The advantage of mak-237 ing two separate images is that the measurement of the mixed-layer concentrations in type I 238 images is not disturbed by the high concentrations in the upper layer with high fluorescent 239 intensities, because the camera shutter has already been closed by the time the beam illu-240 minates the top layer. This aspect was found to be very important: in some experiments the 241 top-layer dye concentrations were huge (as compared to the mixed-layer concentrations), 242 so when illuminated by the laser, fluorescent light emitted from the top layer can scatter in 243 the mixed layer; on the camera images it then appears as if the source is in the mixed layer, 244 which leads to an overestimation of the mixed-layer dye concentration. So, by closing the 245 camera shutter before the laser reached the top layer, this cross-talk problem was evaded in 246 image type I. 247

Image type II, on the other hand, provided useful qualitative information on the structure 248 of the inversion region and the interplay with the convective dynamics. Apart from these 249 qualitative insights, image type II provided important quantitative information on the inflow 250 velocity  $w_b$  because it allowed one to track the evolution of the height of the total water 251 column in time. To this end a floating device was placed on top of the water level (see Fig. 252 2), which is basically a simple piece of foam attached to a small vertical plate that blocks the 253 laser beam when it passes the floater, thus producing a clear demarcation of the water level 254 in the camera image. When the water level rises due to the inflow  $w_{\rm b}$ , the floater will rise at 255 the same speed. From the series of images obtained during the experiment the location of 256



**Fig. 4** Example of PLIF images of a top-down scalar at different instances ( $t = 14t_*$  to  $t = 18t_*$ ). As explained in the main text, because of the zero initial mixed-layer concentration ( $C_0 = 0$ ), these images are well suited for a qualitative (i.e. visual) analysis but not for a quantitative analysis of entrainment.

the laser-blocking floater is easily extracted, which yielded very accurate measurements of  $w_b$ .

To determine the actual dye concentrations from the digital camera images, a few cor-259 rections are required, such as deformation of the image and 'pixel vignetting' (Ferrier et al. 260 1993; Snyder et al. 2002). By photographing a grid consisting of 50 mm squares, located 261 in the tank at the position where the laser sheet normally would be, we concluded that the 262 deformation of the image was very minor and needed no correction. Vignetting (reduction 263 of the brightness in the image periphery compared to the image centre), however, was found 264 to be non-negligible and was subsequently corrected for. The resulting camera intensities 265 I(x,z) are related to the local dye concentration c(x,z) by the Lambert-Beer equation that 266 accounts for the attenuation of the laser beam intensity along its path 267

$$I(x,z) = \beta Q_0 c(x,z) \exp\left\{-\int_0^x \varepsilon c(\xi,z) d\xi - \eta x\right\}.$$
 (15)

Here c(x,z) is the fluorescein concentration at position (x,z),  $\varepsilon$  is the extinction coefficient 268 of fluorescein,  $\eta$  is the extinction coefficient of (non-purified) water,  $Q_0$  is the laser intensity 269 at entrance and  $\beta$  is a proportionality constant that, apart from camera related parameters, 270 includes the fluorescence efficiency of fluorescein (Ferrier et al. 1993). Unfortunately it is a 271 non-trivial matter to solve the inversion problem expressed by Eq. 15, which is required for 272 determining the concentration values c(x,z) from the measured pixel intensities I(x,z). But 273 this issue is not critical because we design the experiment in such a way that there is always 274 a mean dye concentration C in the mixed layer that is much larger than the fluctuations 275  $|c'(x,z)| \ll C$ . Equation 15 then simplifies to 276

$$I(x,z) = \beta Q_0 C \exp\left\{-(\varepsilon C + \eta)x\right\}.$$
(16)

Fitting a straight line to the logarithm of I(x, z) provides information on *C*, both through the slope and the abscissa resulting from the fit procedure. The values for the extinction parameters  $\varepsilon$  and  $\eta$  were determined in a calibration process entailed filling the tank consecutively with 30 known concentrations of fluorescein and recording the corresponding intensities.

Because the laser intensity may vary from experiment to experiment, it was found useful to take several images before initiating the convection. Because we always start with a precisely known dye concentration in the mixed layer we can diagnose in each experiment the proportionality factor  $\beta Q_0$  in Eq. 16. In principle this information is not needed because the slope of the abovementioned linear fit only requires the value of  $\varepsilon$  to give *C*. However we still calculate *C* in both ways (from the slope and the abscissa) because it provides a consistency check.

During the experiments images (type I and II) were taken every 6s, while all experiments 288 lasted longer than 1000s, i.e. at least  $50t_*$ . Another important experimental issue concerns 289 the spatial homogeneity of the surface fluxes. As mentioned, the porous tubes are covered 290 with small pebbles, but the finite size of the porous tubes (17 mm diameter) as well as their 201 orientation in one direction might in principle create a spatial pattern. To test whether the 292 surface fluxes were sufficiently homogeneous, we did a number of experiments in which 293 the dye was added to the reservoir tank rather than to the top layer -a so-called 'bottom-294 up diffusion' set-up. In these experiments the scanning rail was directed in the horizontal, 295 creating horizontal laser sheets while the camera was located above the tank. The observed 296 spatial distributions did not reveal structures with a preferred orientation in one direction or 297 the other. From looking at the images, for example, uninformed observers were unable to 298 determine the orientation of the tubes. 299

## 300 3.3 A mixed-layer model for the saline tank set-up

In order to better understand the results of the saline tank set-up, it proves very useful to con-301 sider a simplified model of the saline convection along the lines of the well-known mixed-302 layer model (Tennekes 1973; Vilà-Guerau de Arellano et al. 2004; van Driel and Jonker 303 2011) for the atmospheric convective boundary layer. This model needs to be slightly re-304 vised for the saline tank due to the mass-flow at the surface that influences the mixed-layer 305 height (Eq. 13) and due to the fact that the surface buoyancy flux (Eq. 12) is not constant. 306 But in all other respects we proceed along the same lines as in the standard mixed-layer 307 model. We also derive prognostic equations for a passive scalar (van Dop et al. 2005), which 308 enables us to compare the model results for the concentrations to the dye concentrations 309 observed in the tank experiments. 310

Within the mixed-layer assumptions the concentration of salinity in the mixed layer evolves according to the surface flux  $\phi_b(t) = -w_b[S(t) - S_b]$  and an entrainment flux  $\phi_e(t) = -w_e[S_t - S(t)]$ 

$$\frac{d}{dt}S = \frac{\phi_b - \phi_e}{h} = \frac{-w_b[S(t) - S_b] + w_e[S_t - S(t)]}{h(t)},$$
(17)

<sup>314</sup> while a similar equation holds for the passive scalar concentration

$$\frac{d}{dt}C = \frac{-w_b[C(t) - C_b] + w_e[C_t - C(t)]}{h(t)},$$
(18)

with the mixed-layer height developing according to Eq. 13. The model is closed as described in Sect. 2, i.e. by relating the entrainment rate  $w_e$  via a zero-order model to the entrainment flux by  $w_e = -B_e/\Delta b$ , while on the other hand expressing the entrainment flux

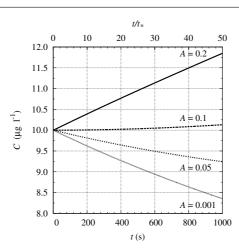
as a (not necessarily constant) fraction of the surface buoyancy flux  $B_e = -AB$ , leading to

$$w_{\rm e}(t) = w_{\rm b} A \frac{S(t) - S_{\rm b}}{S(t) - S_{\rm t}}.$$
 (19)

Upon providing initial conditions  $h(0) = h_0$ ,  $S(0) = S_0$ ,  $C(0) = C_0$  and boundary conditions 319  $S_{\rm b}$ ,  $S_{\rm t}$ ,  $C_{\rm b}$ ,  $C_{\rm t}$ ,  $w_{\rm b}$ , the equations can be numerically integrated once a value for A has been 320 prescribed. It seems as if we have adopted a  $Ri^{-1}$  closure, however, as discussed in Sect. 321 2, A may contain dependencies on Ri, Re and Pe. In our treatment elaborated in the next 322 section we modify A in order to seek the best correspondence with the experiments, i.e. A 323 can be considered as a fit parameter. Should a  $Ri^{-3/2}$  law better represent the entrainment 324 results of the saline tank set-up, then this will become apparent through the resulting best 325 fitting values of A, which in this example would then display a dependence  $A \sim Ri^{-1/2}$ . 326

The mixed-layer model, Eqs. 17–19, can be used to optimally design the entrainment 327 experiments. Since the most accurate measurements can be made when the dye concentra-328 tion in the mixed layer remains roughly constant (that is, within calibration range) during 329 an experiment, we aim to create a situation in which entrainment of higher dye concentra-330 tions is balanced by dilution from the surface flux. Equation 18 provides the experimental 331 settings such that C remains (reasonably) constant during the experiment. One notes from 332 Eq. 18 that it is not possible to keep C exactly constant, because S changes in time; however, 333 approximating  $S \approx S_0$  one can find settings such that C is stationary during the initial period. 334 Elaborating the entrainment closure (Eq. 19), and taking  $C = C_0$ ,  $C_b = 0$ , one finds that the 335 top-layer concentration  $C_t$  should then be chosen as 336

$$C_{\rm t} = C_0 \left[ 1 + \frac{S_0 - S_{\rm t}}{S_0 - S_{\rm b}} \frac{1}{\hat{A}} \right].$$
(20)



**Fig. 5** Evolution of the mixed-layer concentration *C* as predicted by the mixed-layer model for four different entrainment scenarios. The settings were chosen such that an entrainment ratio of A = 0.1 causes the mixed-layer concentration *C* to remain nearly constant during the experiment. The entrainment flux then balances the surface flux. More entrainment (A = 0.2) causes *C* to increase, whereas less entrainment (A < 0.1) entails that *C* decrease.

Note that  $C_t$  in Eq. 20 is based on a *presumed* value of the entrainment ratio, denoted by  $\hat{A}$ . 337 If during the experiment C(t) is found to increase, one can surmise that the entrainment flux 338 was larger than expected, i.e. that the assumed value of the ratio  $\hat{A}$  was too low. On the other 330 hand, if C(t) is found to decrease, the presumed value  $\hat{A}$  was too large. In either case one can 340 redo the experiment with initial/boundary conditions based on a refined hypothesis for the 341 entrainment ratio, until a satisfactory stationary state has been reached. To exemplify this 342 idea, we show the results of the mixed-layer equations Eqs. 17-19 for the typical settings 343 given in Table 1. Additionally we took  $S_t = 0.2\%$ , which corresponds to an initial Ri = 150. 344 Using a presumed value of  $\hat{A} = 0.1$ , we find  $C_t = 70 \mu g l^{-1}$ . Next we integrate the mixed-345 layer equations for four different 'entrainment scenarios' viz. A = 0.001, 0.05, 0.1 and A =346 0.2. Figure 5 shows what would happen with the mixed-layer concentration C(t) for the four 347 scenarios. When A = 0.1, i.e. equal to the presumed value  $\hat{A}$ , the dye concentration remains 348 reasonably constant – as designed. But if the entrainment is characterized by A = 0.2 (twice 349 the 'anticipated' value  $\hat{A}$ ), the concentration increases significantly during the experiment. 350 When A is lower than the anticipated value, concentrations show a marked decrease. In 351 both situations the deviations from the expected equilibrium concentrations are significant 352 enough to be captured by the experimental technique. And, as mentioned, based on the 353 results one can adapt the hypothesis for  $\hat{A}$ , and conduct a new experiment with modified top 354 concentration  $C_t$ . 355

#### 356 3.4 Extracting additional information

It is also possible to diagnose additional information from the concentration measurements, such as the mixed-layer salinity and the evolution of the Richardson and Reynolds number. To this end we follow a bulk mixing approach that can be best explained using Fig. 6. Starting with a concentration  $C_0$ , which is well mixed over a depth  $h_0$  (step 1 in Fig. 6), at

time t the mixed-layer height will have increased due to the mass inflow  $h_b = \int_0^t w_b dt$  and

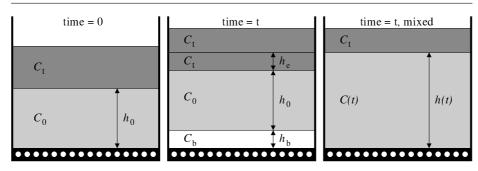


Fig. 6 Evolution of the mixed-layer height h(t) and mixed-layer concentration C(t) as a result of bottom-up (surface inflow) and top-down (entrainment) processes.

<sup>362</sup> due to entrainment  $h_e = \int_0^t w_e dt$ , i.e.

$$h(t) = h_0 + h_b(t) + h_e(t).$$
 (21)

The inflow is associated with a concentration  $C_{\rm b}$ , whereas entrainment infuses concentra-

tions  $C_t$  (step 2 in Fig. 6); the turbulent mixing of these contributions results in a new mixed-layer concentration (step 3 in Fig. 6)

$$C(t) = \frac{C_0 h_0 + C_b h_b(t) + C_t h_e(t)}{h_0 + h_b(t) + h_e(t)}.$$
(22)

In the experiment  $C_0$ ,  $C_t$  and  $h_0$  are known, and both C(t) and  $h_b(t)$  are carefully measured as outlined in Sect. 3.2; it is therefore possible to derive  $h_e(t)$  from Eq. 22,

$$h_{\rm e}(t) = \frac{[C(t) - C_{\rm b}]h_{\rm b}(t) + [C(t) - C_0]h_0}{C_{\rm t} - C(t)}.$$
(23)

Once  $h_e(t)$  has been determined one can also derive the salt concentration in the mixed layer by an equation equivalent to Eq. 22,

$$S(t) = \frac{S_0 h_0 + S_b h_b(t) + S_t h_e(t)}{h_0 + h_b(t) + h_e(t)}.$$
(24)

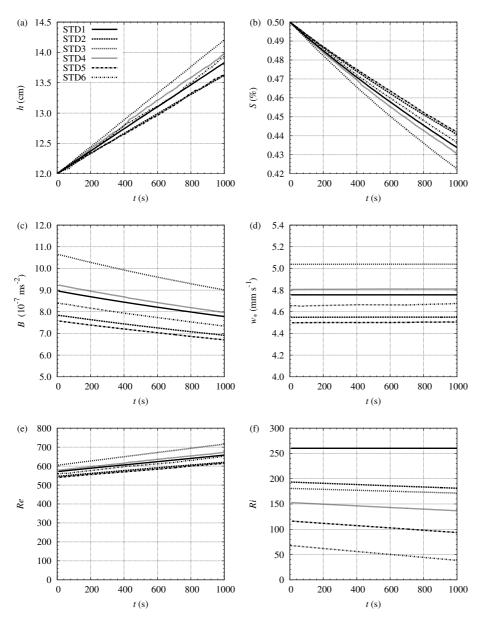
<sup>370</sup> It is therefore possible to calculate the evolution of the Richardson number (Eq. 14) during <sup>371</sup> the experiment, as well as the evolution of the surface buoyancy flux (Eq. 12), the convective <sup>372</sup> velocity scale  $w_*$  (Eq. 5), and the Reynolds number (Eq. 8).

Last but not least we can derive the entrainment ratio from its definition  $A = -B_e/B$ , in which  $B_e$  can be obtained from  $db/dt = (B - B_e)/h$ . Expressed in terms of salinity this amounts to determining

$$A = -\left[1 + \frac{h(t)}{w_{\rm b}(S - S_{\rm b})} \frac{\mathrm{d}S}{\mathrm{d}t}\right] \tag{25}$$

for which we need Eq. 24, Eq. 21 and Eq. 23, respectively. We tested this method on the time series of C(t) generated by the mixed-layer model (see, for example, Fig. 5) for different entrainment scenarios, and found excellent agreement between the diagnosed value of the entrainment ratio and the value of *A* that was used to generate the data. When applied to the real measurements, Eq. 25 produces significant scatter in the instantaneous values, which is why we report the average values of an experiment together with error bars to give an indication of the variability.





**Fig. 7** Evolution of the derived quantities in experiments STD1-6. See Table 2 for the particular settings in each experiment. (a) mixed-layer depth h(t); (b) mixed-layer salinity S(t); (c) surface buoyancy flux B(t); (d) convective velocity scale  $w_*(t)$ ; (e) Reynolds number; (f) Richardson number.

Exp.	<i>S</i> <sub>0</sub>	St	Sb	Ct	wb	$h_0$	$\langle Ri \rangle$	$\langle Re \rangle$
	(%)	(%)	(%)	$(\mu g l^{-1})$	$(m s^{-1})$	(m)		
STD1	0.50	0.00	0.00	210	$1.83 \cdot 10^{-5}$	0.12	260	619
STD2	0.50	0.16	0.00	140	$1.60 \cdot 10^{-5}$	0.12	187	583
STD3	0.50	0.11	0.00	164	$2.17 \cdot 10^{-5}$	0.12	176	660
STD4	0.50	0.20	0.00	70	$1.89 \cdot 10^{-5}$	0.12	145	625
STD5	0.50	0.30	0.00	85	$1.55 \cdot 10^{-5}$	0.12	105	578
STD6	0.50	0.38	0.00	50	$1.71 \cdot 10^{-5}$	0.12	53	605
LRE1	0.50	0.40	0.40	95	$1.89 \cdot 10^{-5}$	0.03	94	71
LRE2	0.50	0.34	0.40	85	$2.00 \cdot 10^{-5}$	0.03	160	73
LRE3	0.50	0.34	0.40	85	$1.96 \cdot 10^{-5}$	0.03	161	72
MRE1	0.11	0.00	0.00	200	$1.91 \cdot 10^{-5}$	0.12	152	379
HRE1	4.00	2.80	0.00	60	$1.89 \cdot 10^{-5}$	0.12	124	1261
HRE2	4.00	3.50	0.00	50	$1.89 \cdot 10^{-5}$	0.12	49	1200

**Table 2** An overview of the settings in the various experiments. In all experiments the initial fluorescein concentrations in the mixed layer is  $C(0) = C_0 = 10 \ \mu g \ l^{-1}$  and in the bottom layer  $C_b = 0$ . The indicated Richardson and Reynolds numbers are average values; their evolution during the experiments is shown in Fig. 7.

#### 383 4 Results

<sup>384</sup> Figure 7 shows the evolution of a number of key quantities during experiments STD1–

STD6, i.e. the evolution of the mixed-layer depth h(t), the salinity concentration S(t) in

the mixed layer, the surface buoyancy flux B(t), the convective velocity scale  $w_*(t)$ , the Reynolds number and the Richardson number. The figure shows that the mixed-layer salinity

Reynolds number and the Richardson number. The figure shows that the mixed-layer salinity decreases, and therefore also the surface buoyancy flux; the convective velocity scale  $w_*$ 

remains constant because the product of h and B is constant. The Reynolds number, Re =

remains constant because the product of *h* and *B* is constant. The Reynolds number,  $Re = w_*h(t)/v$ , therefore increases somewhat during an experiment. In experiments STD1–6 we

<sup>391</sup> chose to use the same value for the initial mixed-layer salinity,  $S_0 = 0.5\%$ , in order to set the

<sup>392</sup> Reynolds numbers at roughly the same value. The Richardson number was then controlled

<sup>393</sup> by the choice of the top-layer salinity  $S_t$ . Richardson numbers up to 260 were studied, and

several experiments (STD1,LRE1,MRE1) were designed such that the Richardson number

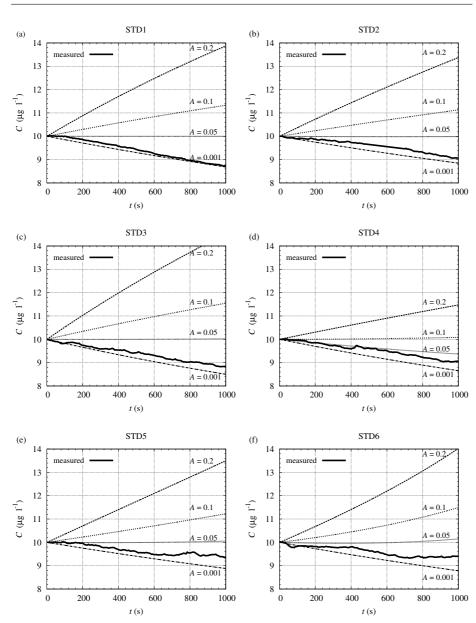
does not alter during an experiment; this was done by choosing  $S_t = S_b$ , see Eq. 14. In

the other experiments  $S_t \neq S_b$ , entailing that the Richardson number changes during the experiment.

In order to measure entrainment fluxes, an appreciable number of experiments was car-398 ried out following the principles outlined in Sect. 3.3, that is, the top layer was coloured 399 with dye concentration  $C_t$  based on a presumed entrainment ratio  $\hat{A}$  as described by Eq. 400 20. If during the experiment the *real* entrainment ratio A were equal to  $\hat{A}$ , the mixed-layer 401 concentrations C would be virtually constant because dilution by the surface flux is then 402 just compensated by the entrainment flux. We started out by choosing  $\hat{A} = 0.2$ , close to the 403 accepted value for atmospheric free convection. But invariably we were confronted with de-404 creasing fluorescein concentrations in the mixed layer, indicating that the actual entrainment 405 flux in the saline tank was significantly lower than anticipated. A new series of experiments 406 based on  $\hat{A} = 0.1$  yielded comparable results (see, for example, experiment STD4). Finally 407 we based the top-layer dye concentration in all experiments (except STD4) on  $\hat{A} = 0.05$ . 408

Figure 8 shows the measured fluorescein concentrations; as one can observe, in all experiments the mixed-layer concentrations *C decrease*. The inevitable conclusion is therefore that in all experiments the entrainment ratio was *lower* than A = 0.05. As an aid to the eye we

<sup>412</sup> have also indicated in the figure the predictions by the mixed-layer model (Eqs. 17–19) for



**Fig. 8** Evolution of the measured fluorescein concentration in the mixed layer (*C*) indicated by the solid line. See Table 2 for the particular settings in each experiment. Dotted lines show the evolution as predicted by the mixed-layer model (Eqs. 17–19) for four different entrainment scenarios:  $\hat{A} = 0.2, 0.1, 0.05$  and 0.001.

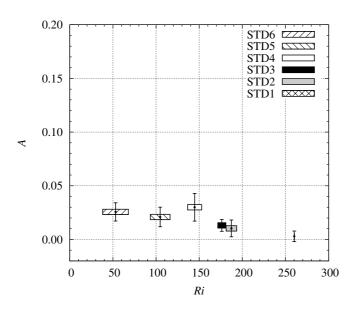


Fig. 9 Entrainment ratio A as a function of the Richardson number. The vertical bars indicate the error in the mean. The horizontal size of the box indicates the range in which the Richardson number varied during the experiment (see Table 2).

four different entrainment scenarios A = 0.2, 0.1, 0.05 and 0.001 (virtually a non-entraining situation). From the evolution of C(t) with respect to the cones delineated by the mixedlayer predictions, one can conclude that an entrainment ratio close to A = 0.02 would be most appropriate to represent the experimental results. That is more than ten times lower than A = 0.25, the entrainment ratio reported for a thermal convection tank set-ups, (Dear-

<sup>418</sup> dorff et al. 1980; Fernando and Little 1990).

From these data we have calculated the average value of the entrainment ratio by means 419 of Eq. 25 and plotted the result as a function of the average Richardson number in Fig. 420 9. Horizontal bars indicate the Richardson number range, whereas vertical bars indicate 421 the error in the mean as determined from the series of instantaneous values. Clearly there 422 is appreciable scatter, but the mean values display a clear signal and appear to be well in 423 agreement with the early conclusion drawn from Fig. 8 that  $A \approx 0.02$ . Figure 9 might suggest 424 that the flux ratio even decreases for larger Richardson numbers. Given the appreciable 425 scatter it is unclear whether this points to a  $E \sim Ri^{-3/2}$  entrainment law (Turner 1968), or 426 whether the data hint at a 'regime change' (near  $Ri \sim 150$ ) as purported by Kantha (1980). 427 But this issue seems of lesser importance as compared to the result that A has so low a value 428 in the entire Richardson number range - the entrainment fluxes are so small that it seems 429 hardly relevant that they become even lower for larger Richardson numbers. 430

Because of the large discrepancy between the accepted value of *A* and the saline convection tank results presented here, a number of additional tests was carried out to check the
validity of our measurement method. Since our data are primarily based on concentration
measurements, we performed the following tests (the list is not exhaustive):

435 1) Stability of the fluorescent dye. This was tested by setting-up and conducting an exper-

436 iment in the normal fashion, but without initiating convection ( $w_b = 0$ ). In the absence of

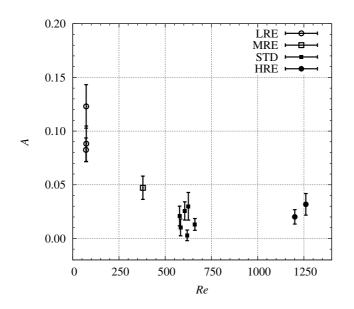


Fig. 10 Entrainment ratio A as a function of the (average) Reynolds number (Richardson numbers vary between experiments, see Table 2). Vertical bars indicate the error in the mean.

437 surface inflow and dynamics, the mixed-layer concentration must remain constant since, as

discussed in Sect. 3.1, molecular diffusion is very small in a saline tank set-up. Whether

<sup>439</sup> also the *measured* concentrations remained constant was tested because the fluorescence

efficacy of the dye might autonomously decay (due to aging, or overexposure by the laser
 sheet). However, the concentrations measurements were found to be satisfactorily constant
 within the duration of an experiment.

2) Study of a non-entraining case. In this test the top layer was not added, but convection

was initiated in the usual way resulting in a well-mixed turbulent layer, which grows due

to the surface mass inflow but not due to entrainment. The measured mixed-layer concen-

trations were found to decrease as expected on the basis of the surface dilution (which is known because  $w_b$  is measured). Also the reversed situation was tested where  $C_b > C_0$ , i.e.

<sup>448</sup> a 'bottom-up' case. Concentrations increased according to the expectations.

<sup>449</sup> 3) Direct determination of the entrainment velocity. The images captured during the exper-

 $_{450}$  iments were reanalyzed to determine the evolution of the mixed-layer height *h* directly by

<sup>451</sup> locating the concentration jump in the images (type II). The entrainment rate was subse-<sup>452</sup> quently derived from  $dh/dt - w_b$ . As argued in Sect. 3.2 such a method is prone to inaccu-

quently derived from  $dh/dt - w_b$ . As argued in Sect. 3.2 such a method is prone to inaccuracies, but we employed it nonetheless to look for potential large inconsistencies with the

454 concentration method. However, no such inconsistencies were detected.

455 These tests further strengthened the conclusion that the entrainment flux in our saline

456 convection tank set-up was about an order of magnitude lower than expected for penetrative

457 convection.

#### 458 **5** Discussion and conclusion

Because the saline convection tank was built so as to serve as a physical model for the atmo-459 spheric convective boundary layer, it is important how to interpret the present entrainment 460 results. Not only is there a discrepancy with previous convection tank experiments (Dear-461 dorff et al. 1980; Fernando and Little 1990) but also with recent atmospheric observations 462 463 which put A in the range 0.1—0.4 (Betts and Ball 1994; Angevine et al. 1998) and some-464 times even higher (Hägeli et al. 2000). Three items are important in this discussion: 1) the 465 two-layer set-up; 2) the mixed-layer approach to analyze and interpet the data, and 3) the potential influence of molecular properties of the working fluid, i.e. viscosity and diffusivity, 466 or in terms of dimensionless numbers, the influence of the Reynolds and the Péclet number. 467

#### 468 5.1 Two-layer set-up

The two-layer set-up, in particular the neutral top layer, differs from a linearly stably strat-469 ified upper layer in that a neutral top layer does not sustain buoyancy waves. But it is not 470 directly clear why the absence of waves would reduce entrainment; the argument (e.g. Stull 471 1976; Fedorovich and Mironov 1995) that part of the kinetic energy might be transported 472 through the interface to be transferred to wave energy in the upper layer, works in the oppo-473 site direction, i.e. absence of waves in the two-layer set-up leaves more energy for entrain-474 ment. Of course also in the two-layer set-up a stable interfacial layer is formed resulting 475 from the system dynamics itself, i.e. the interplay of convection-diffusion, creating some 476 inversion thickness. Indeed, in animations of the camera images travelling waves along the 477 interface were frequently observed, in particular for the higher Richardson numbers. But 478 these waves were not found to break and appeared unable to 'corrupt' the interface and 479 initiate appreciable entrainment. 480

#### 481 5.2 Mixed-layer analysis

In the current set-up entrainment information was retrieved by diagnosing the evolution 482 of the mean dye concentration in the mixing layer. To this end a mixed-layer model was 483 invoked that was adapted to the saline tank setting. The results could therefore be affected 484 if the mixed-layer approach is invalid or introduces a significant bias. In fact, the images 485 presented in Fig. 4 might question the validity of the mixed-layer view. We emphasize, 486 however, that these type of images, although visually appealing, were not used in the analysis 487 of entrainment because a variety of optical effects hamper a quantitative interpretation of the 488 image. For reasons detailed in Sect. 3.2, it was better to raise the mean concentration in the 489 mixed layer such that entrainment is balanced by surface dilution. The resulting, uniformly 490 green, images corroborate the validity of the mixed-layer view. In addition it is important 491 to note that, as mentioned in point 3 of the previous section, we also directly measured the 492 mixed-layer depth. The results turned out to be consistent with each other in the sense that 493 both methods show a very slowly deepening boundary layer. This implies that if the mixed-494 layer approach was inappropriate while in reality the layer was deepening much faster, then 495

this would have shown up in the direct measurements of the layer depth.

#### <sup>497</sup> 5.3 Reynolds number dependence

The next important issue is the influence of (geometrical) details and molecular fluid prop-498 erties. As mentioned earlier, the general idea is that these details should not matter when the 499 Reynolds number is large enough. Problem is that it is unclear a priori what 'large' means 500 for a certain flow configuration. For the present set-up both the geometry and Reynolds 501 number range were deliberately chosen to be comparable to the set-up of Deardorff et al. 502 (1980), which after all is used as a benchmark in a large number of studies. In order to study 503 the influence of the Reynolds number we conducted additional experiments with Reynolds 504 numbers ten times smaller (LRE), two times smaller (MRE), and two times larger (HRE) 505 than the STD experiments. Due to practical limitations it is not easy to change the Reynolds 506 number drastically. For instance, the high Reynolds number cases were realized by increas-507 ing the mixed-layer salinity  $S_0$ , whereas the low Reynolds number cases were realized by 508 lowering  $h_0$  and reducing the salinity contrast  $S_0 - S_b$ ; see Table 2. The results for the en-509 trainment ratio A as a function of Re are plotted in Fig. 10 for all experiments regardless 510 of the Richardson number. Interestingly, and perhaps unexpectedly, one observes that in the 511 figure that A decreases for increasing Re and then levels off. It is hard to speculate what 512 would happen for much larger Reynolds numbers, but as regards to Fig. 10 we cannot ob-513 serve a 'worrying' increasing trend in A. Of course the Reynolds numbers are much smaller 514 than is associated with atmospheric convection,  $Re \approx 10^8$ , based on  $h = 10^3$  m,  $w_* = 1$  m s<sup>-1</sup>. 515  $v \approx 10^{-5} \text{m}^2 \text{ s}^{-1}$ . But it is important to realize that if the Reynolds numbers in this study 516 are deemed not large enough, then the same conclusion must drawn for the thermal tank 517 set-up of Deardorff et al. (1980), which takes away an important assumption that A = 0.25518 for free convection. In addition, if one considers the Péclet number, one has  $Pe \approx 10^8$  for 519 the atmospheric situation,  $Pe \approx 10^6$  for the saline tank set-up and  $Pe \approx 10^4$  for a typical 520 thermal tank set-up. So from the perspective of the Péclet number the saline tank is closer to 521 atmospheric convection than a thermal tank. Turner (1968) suspected that the Péclet number 522 could be too low in a thermal set-up; see also the discussion on thermal convection tanks 523 in Turner (1973). However, Fernando and Little (1990) conducted a comprehensive study 524 on the impact of the Péclet number in a thermal set-up but could not detect any significant 525

<sup>526</sup> influence and confirmed the results of Deardorff et al. (1980).

#### 527 5.4 Prandtl number dependence

It is very disturbing that different – carefully executed – laboratory studies of the same 528 geophysical phenomenon give such different results. This holds in particular for experiments 529 on entrainment, see the review on laboratory experiments by Fernando (1991). But at least 530 one clear pattern seems to emerge, namely that laboratory experiments with low diffusivity 531 fluids (i.e. with a high Prandtl number) tend to entrain much less. If one looks closely at the 532 buoyancy flux profiles in the saline tank experiments of van Dop et al. (2005) (their figure 533 7), one notices quite small entrainment fluxes as well. The discrepancy between thermal 534 and saline arrangements was already observed by Turner (1968) for entrainment induced 535 by decaying forced convection (oscillating grid turbulence). Wolanski and Brush (1975) 536 extended Turner's study in a similar set-up but with different solvents (salt, sugar, clay, etc) 537 and found a strong dependence of the Prandtl number (actually Schmidt number) on the 538 entrainment rates. 539

In this context it is also important to consider the entrainment experiments of Sayler and
 Breidenthal (1998). Their experimental setting was different to the extent that the case of ra-

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diatively driven entrainment was studied mimicking stratocumulus clouds, yet the observed entrainment rates displayed a clear dependence on whether the stratification was applied by heat,  $w_e/w_* = 0.25Ri^{-1}$ , or by dextrose,  $w_e/w_* = 0.08Ri^{-1}$ . So the prefactor, i.e. the entrainment flux ratio *A*, revealed a strong Prandtl number dependence. Interpretation of this result is complicated by the fact that the turbulence in both cases was generated by thermal effects, while the stratification was applied by either heat or by salinity, so the effects of the Prandtl/Schmidt number are convoluted in this setting.

In the same vein we could summarize for free penetrative convection: heat  $w_e/w_* \approx$ 549  $0.25Ri^{-1}$  (Deardorff et al. 1980; Fernando and Little 1990), salt:  $w_e/w_* \approx 0.02Ri^{-1}$  (this 550 study and van Dop et al. 2005). But such a state of affairs is entirely unsatisfactorily when 551 it comes to the generalization to atmospheric cases, and, as such, it fails to sufficiently con-552 strain model predictions. An excellent example in this regard is the LES intercomparison 553 study targeted at the so-called 'smoke case' (Bretherton et al. 1999): noting in the exper-554 iments of Sayler and Breidenthal (1998) that A increased by a factor 6 when the Prandtl 555 number decreased from 1000 to 7, an *ad hoc* power law dependence  $A \sim Pr^{-\gamma}$  was invoked, 556 which was subsequently extrapolated to the atmospheric case of Pr = 0.7; this then yielded 557 A = 0.4 - a value much closer to the prevailing LES results. If, purely for the sake of the 558 argument, we apply the same procedure for the dry CBL based on A = 0.25 for Pr = 7 and 559 A = 0.02 for  $Pr = 10^3$ , we would obtain A = 0.8 for air. This value is clearly unacceptably 560

<sup>561</sup> high and underlines the problem of ambiguous data from laboratory experiments.

# 562 5.5 Outlook

Since the Prandtl number is the ratio between two molecular properties of the fluid, and since 563 Reynolds number similarity requires results to become independent of molecular properties, 564 it seems fair to conclude that laboratory experiments on (convective) entrainment are still 565 very much hampered by the low magnitude of the Reynolds number. There seem to be 566 two viable options to better understanding the reason for the different entrainment results 567 and resolve the impasse. First, one could make use of the newly available supercomput-568 ing resources and conduct direct numerical simulations (DNS) of the cases. Clearly, the 569 atmospheric Reynolds number is out of reach for DNS, but present resources allow for sig-570 nificantly larger Reynolds numbers than the original laboratory experiments (e.g. Mellado 571 2012; Jonker et al. 2013). 572

Second, one could attempt to reach much larger Reynolds numbers in the laboratory. 573 This can be done by using a different fluid, such as cryogenic helium gas, employed for high 574 Rayleigh number experiments (e.g. Niemela et al. 2000). Or one could try to significantly 575 upscale the water tank set-up. For a saline set-up, such an endeavour is challenging but not 576 inconceivable. Apart from increasing h, the surface buoyancy flux B can be increased by 577 increasing  $w_b$ . Limiting the inflow velocity to  $w_b = w_*/100$  in order not to disturb the flow, 578 one finds via Eq. 5 and Eq. 12 that  $w_* \sim h^{1/2}$  and hence  $Re \sim h^{3/2}$ . Together with some room 579 for increasing the mixed-layer salinity, it turns out that a swimming pool size set-up allows 580 for a Reynolds number hundred times larger than that achieved in the present study. Such an 581 experiment would be very worthwhile as it will shed light on the Reynolds number influence 582 on convective entrainment of a high Prandtl number fluid. This is an outstanding question, 583 not only from a basic fluid mechanics perspective, but also with respect to the oceanic mixed 584

<sup>585</sup> layer where salinity directly influences buoyancy.

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