

1 *Short title* [Lignin signature in humic acids, index for soil carbon stabilization]

2

3 **Pyrolytic appraisal of the lignin signature in soil humic acids: assessment of its**
4 **usefulness as carbon sequestration marker**

5

6 I. Miralles^a, A. Piedra-Buena^b, G. Almendros^b, F.J. González-Vila^c, J.A. González-Pérez^{c*}

7

8 ^a *Departamento de Desertificación y Geoecología, EEZA-CSIC, La Cañada de San Urbano,*
9 *E-04230 Almería, Spain,* ^b *Museo Nacional de Ciencias Naturales, MNCN-CSIC, Serrano*
10 *115B, E-28006 Madrid, Spain,* ^c *Instituto de Recursos Naturales y Agrobiología de Sevilla,*
11 *IRNAS-CSIC, Avda. Reina Mercedes, 10, E-41012 Seville, Spain.*

12

13 *) Corresponding author: J.A. González-Pérez E-mail: jag@irnase.csic.es

14

15 **Abstract**

16 Lignin markers in humic acids (HA, the alkali-soluble, acid-insoluble soil organic matter
17 fraction) molecular features are explored to assess the extent to which plant
18 biomacromolecules are progressively transformed by humification processes leading to
19 stable C-forms in soils. Humic acids extracted from a collection of mountain calcimorphic
20 soils from Sierra María-Los Vélez Natural Park (Southeastern Spain) under different use and
21 management practices were studied in detail by visible and infrared (FT-IR) spectroscopies
22 and analytical pyrolysis (Py-GC/MS). The HAs display a more or less marked lignin pattern
23 defined by characteristic methoxyphenol assemblages released after pyrolysis that are
24 associated to a typical infrared pattern including absorption frequencies bands at 1510, 1460,

25 1420, 1270, 1230 and 1030 cm^{-1} . This variability in the HA spectroscopic and pyrolytic
26 patterns was used as a source of molecular-level surrogates to establish the balance between
27 complementary mechanisms of soil C sequestration *i.e.*, a selective preservation of lignin
28 associated to raw organic matter and other plant-inherited macromolecules, or alternative
29 mechanisms involving microbial breakdown or plant precursors and its condensation with
30 microbial metabolites.

31 We found that HAs in which the lignin signature was comparatively less marked also show
32 high optical density values suggesting unsubstituted, condensed aromatic units and a chaotic
33 organic structure, pointing to the presence of highly resilient carbon forms. Upon analytical
34 pyrolysis, one group of HAs produced major yields of methoxyl-lacking aromatics
35 (alkylbenzenes and alkylphenols), and poor yields of alkyl compounds, which suggest
36 efficient cleavage of biomacromolecules and the occurrence of active microbial synthesis
37 and condensation processes. In fact, these HAs also displayed broadband IR spectra, and
38 visible spectra showing high optical density and polynuclear quinoid chromophors
39 considered of fungal origin. Other group of HAs yielded upon pyrolysis conspicuous series
40 of methoxyphenols and well-defined alkyl series (alkanes, alkenes and fatty acids). The IR
41 spectra also displayed clear lignin and amide bands, as well as intense 2920 cm^{-1} band and a
42 low optical density, indicative of a marked aliphatic character. This latter is interpreted as the
43 result of recent diagenetic alteration processes of young organic matter and suggests that C
44 sequestration mechanisms in these soils are mainly based on the stabilization of HAs from
45 plant biomacromolecules and aliphatic structures.

46 These differential lignin alteration patterns indicate that HAs are responsive to soil C
47 sequestration mechanisms, which in the studied soils seem to relay upon microtopographical
48 features rather than to changes in soil use and management.

49

50 **Keywords:** Lignin; humic acids; humification mechanisms; calcimorphic soils.

51 **1. Introduction**

52 Despite the crucial global implications of carbon (C) stored in soils and sediments [1], the
53 biogeochemical processes involved in C stabilization are not well understood [2]. The study
54 of the molecular structure and variability in soil organic matter (SOM) may help in
55 unravelling such stabilization processes as well as to infer resilience characteristics of
56 different SOM fractions [3, 4]. Although humification is an active process involving
57 biological cleavage of plant and microbial biomass followed by secondary condensation of
58 soluble products into humic substances [5], in some circumstances biodegradation is severely
59 hampered by climatic, biotic or mineralogical soil-forming factors leading to accumulation
60 of raw humus types [6]. In these cases the composition of the resulting humic substances
61 could be described as a dynamic heterogeneous mixture of relatively low molecular size
62 components associated via hydrophobic interactions and hydrogen bonds [7]. While at the
63 early humification stages such a supramolecular conformation could be stabilized mainly by
64 weak dispersive forces instead of covalent linkages [8], defining a conceptual model which
65 accounts for their essential role in providing and maintaining soil physical and chemical
66 quality [9], at advanced humification—maturation—stages the humic substances show an
67 outstanding intrinsic resilience and this is crucial to define soil biogeochemical quality of
68 soils temporarily behaving as active C sinks. In particular progressive organo-metallic
69 interactions and additional free-radical condensation of the three-dimensionally bridged
70 structure of humic substances in the course of progressive transformation—maturation—
71 stages may end in comparatively rigid condensed domains. In this situation most of the
72 humic structure could consist of a ‘megamolecule’ formed by a network of C–C and C–O
73 links, where discrete structural units can no longer be recognizable due to the similar stability
74 to chemical and biological degradation of all bonds involved in the whole structure [3]. Such

75 SOM advanced transformation stage is often found in continental Mediterranean semiarid
76 environments where abiotic factors *i.e.* contrasting temperature and moisture levels, sunlight
77 exposure and the historical effect of wildfires, may lead to soils with low SOM content but
78 highly stable and resistant to biodegradation (resilient).

79 In this line, several studies have pointed out the possibility to use the structural information
80 provided by the molecular characterization of SOM to differentiate between a) ecosystems
81 where soil C sequestration relies upon microbial mediated processes with an intense
82 reworking and abiotic condensation of precursors producing intrinsically resilient
83 macromolecular humic substances of chaotic structure and b) soils where the preservation of
84 raw organic matter prevails and depends on extrinsic factors—mainly organo-mineral
85 interactions—leading to a organic matter organization that is accessible to soil enzymes [10].

86 Adopting extreme positions about soil C sequestration mechanisms debate have frequently
87 led to hermeneutic controversies in the search of a unified theory justifying SOM
88 stabilization [11]. In particular, calcimorphic soils could be especially suitable to analyze
89 qualitative and quantitative features relevant in the dynamics of SOM. This is due to the fact
90 that these soils displays peculiar features associated both with e.g. microencapsulation
91 processes of particulate plant-inherited materials [12, 13] but also with active insolubilization
92 mechanisms of the mineral matrix related to the release of low molecular weight compounds
93 onto a Ca^{2+} saturated soil solution [14]. In fact, the prevailing limestone substrate in semiarid
94 Southern Spain's soils has been considered to play a role in the low structural variability in
95 the molecular structure of the HAs [12]. This situation demands the use of accurate analytical
96 techniques (*i.e.*, analytical pyrolysis) betraying environmental proxies (molecular markers)
97 responsive of the different sources of environmental variability reflected in the composition
98 of the SOM [15–17].

99 In the present study analytical pyrolysis (Py-GC/MS) together with visible and IR
100 spectroscopies, are used to study HAs molecular structure and to find compositional
101 descriptors informative about C stabilization processes in a variety of semiarid ecosystems
102 developed on calcimorphic soils.

103

104 **2. Material and methods**

105 *2.1. Soil sampling*

106 The area of study is located in the Natural Park Sierra María-Los Vélez (Almería,
107 Southeastern Spain) which includes a wide variety of semiarid ecosystems both seminatural
108 (forests and brushwood) and disturbed (almond-tree orchards and cereal fields) developed on
109 calcimorphic soils. The natural vegetation consists of pine forests (> 90 yr), oak forests,
110 reforested pine forests, brushlands, almond-tree orchards and cereal crops (Table 1) [18]. The
111 climate is Mediterranean-type, with typical continental features ranging from semiarid to
112 subhumid. Temperatures are 11.9–16.9 °C with a dry summer season; rainfall events are
113 intense and occasional. The geological substrate consists of sedimentary rocks (limestones,
114 marls and dolomites) and soils are Rendzic and Lithic Leptosols, Calcic, Petrocalcic and
115 Hypercalcic Chernozems, Kastanozems and Hypercalcic, Luvic and Petrocalcic Calcisols
116 [19].

117 Soil samples (*ca.* 500 g) were collected with a spade from the uppermost horizon (0–
118 10 cm) after litter removal. In the laboratory, composite samples (obtained by mixing three
119 subsamples taken *c.* 20 m apart in the field) were air-dried and sieved to fine earth (< 2 mm)
120 before analysis.

121 *2.1. Physical and-chemical analyses*

122 Soil pH was measured in a 1:1 soil:water suspension. Total carbonates were measured as
123 CaCO_3 with the Bernard calcimeter [20]. The soil water holding capacity was estimated at -
124 1.5 and -0.33 MPa in a pressure-membrane extractor [21]. Total nitrogen was determined by
125 micro-Kjeldahl digestion and soil C by wet oxidation using dichromate in acid medium
126 followed by redox titration [22]. Cation exchange capacity (CEC) and exchangeable ions Ca^{2+} ,
127 Mg^{2+} , K^+ , and Na^+ were measured after extraction with ammonium acetate solutions (1 mol L^{-1}
128 NH_4Ac at pH 7) [23].

129

130 *2.2. Soil organic matter fractionation*

131 The methods applied for the isolation and quantitative determination of the humus fractions
132 were based on standard procedures [14, 24]. The separation of the particulate, low density
133 fraction (floating organic particles not yet transformed into humic substances, which in some
134 cases may include some charcoal) was carried out by flotation using soil samples of 10 g
135 suspended in $2 \text{ mol L}^{-1} \text{H}_3\text{PO}_4$. After rotary stirring for 1 min, the floating soil fraction or
136 free organic matter was isolated by centrifuging the suspension and filtering, washed with
137 distilled water and analyzed for total C. The soil pellet remaining after centrifugation was
138 resuspended in $0.1 \text{ mol L}^{-1} \text{Na}_4\text{P}_2\text{O}_7$ (horizontal motion mechanical shaking for 3 h) and
139 centrifuged. This treatment was repeated up to 3 times followed by 2 additional extractions
140 with $0.1 \text{ mol L}^{-1} \text{NaOH}$; the dark brown extracts successively obtained (corresponding to the
141 total humic extract: HA + fulvic acid) were aggregated. Two aliquots were taken from this
142 extract, and precipitated with H_2SO_4 (1:1 by vol.) for further determination of the amounts of
143 the acid-soluble fulvic acid and the precipitated HA fraction. The soil residue after the
144 alkaline extraction was washed with distilled water and desiccated at $40 \text{ }^\circ\text{C}$. The C
145 concentration in this residue corresponded to the total humin.

146

147 *2.3. Preparative isolation and purification of the HA fraction*

148 Qualitative isolation and purification (de-ashing) of the HAs was performed by precipitating
149 the total humic extract with 6 mol L⁻¹ HCl to pH= 2, centrifuging, redissolving the acid
150 insoluble HA in 0.5 mol L⁻¹ NaOH and high-speed centrifuging at 43500 *g*. The
151 centrifugation pellet (particulate organic matter and clay minerals) was discarded and the
152 brown supernatant sodium humate was reprecipitated with HCl and centrifuged. Finally the HA
153 in the gel state and acid pH was dialyzed in distilled water using cellophane bags (Visking®
154 dialysis tubing, molecular weight cutoff 12,000–14,000 Da; pore diameter *ca.* 25 Å,
155 Medicell) and desiccated at 40 °C.

156 *2.4. Characterization of the HA fraction*

157 The HAs were studied by pyrolysis-gas chromatography/mass spectrometry (Py-GC/MS),
158 and by visible and IR spectroscopies. The optical density, that is considered as a surrogate of
159 the aromaticity of HAs [25], was measured in solutions of 66.6 mg L⁻¹ C in 0.02 mol L⁻¹
160 NaOH. Second-derivative spectra were acquired with a Shimadzu UV-240 OPI-2
161 spectrophotometer. The IR spectra of the HAs were obtained with a Shimadzu FTIR-8400 PC
162 using KBr dishes with 2.00 mg HA. In order to assist the visual inspection the broadband
163 profiles in the IR spectra from HAs, a resolution enhancement algorithm was applied based
164 in subtracting the raw spectrum (sized to 640 data points) from multiple (×60) of its second
165 derivative and further application of smoothing [26, 27].

166 Pyrolysis–gas chromatography–mass spectrometry (Py–GC/MS) was carried out using a
167 double-shot pyrolyzer device PY2020iD (Frontier Lab Ltd., Fukushima, Japan) attached to a
168 GC/MS system Agilent 6890 [28]. Samples of 1–2 mg in weight were pyrolyzed in small
169 crucible capsules introduced for 1 min into the micro-furnace preheated at 500 °C. The
170 evolved gasses were directly injected into the GC/MS for analysis. The GC was equipped
171 with a low-to-mid polarity fused silica capillary column DB-1701 (J&W Scientific) (30 m ×

172 250 $\mu\text{m} \times 0.25 \mu\text{m}$ film thickness), the oven temperature was held at 50 $^{\circ}\text{C min}^{-1}$, and then
173 increased up to 100 $^{\circ}\text{C}$ at 30 $^{\circ}\text{C min}^{-1}$, from 100 to 300 $^{\circ}\text{C}$ at 10 $^{\circ}\text{C min}^{-1}$, and stabilized at
174 300 $^{\circ}\text{C}$ for 10 min using a heating rate of 20 $^{\circ}\text{C min}^{-1}$. The carrier gas used was helium at a
175 flow of 1 mL min^{-1} . The detector was an Agilent 5973 mass selective detector and mass
176 spectra were acquired with a 70 eV ionizing energy. Compound assignment was achieved *via*
177 single-ion monitoring for various homologous series, *via* low-resolution mass spectrometry
178 and comparison with published and stored data (NIST and Wiley libraries).

179 2.5. Statistical analyses

180 Data treatments were performed with the Statistica package [29]. The joint classification of
181 soil and environmental variables in addition to variables corresponding to molecular
182 characteristics of the SOM was carried out by a multidimensional scaling method [30].

183

184 3. Results

185 3.1. Spectroscopic characterization of the HAs

186 The HAs chosen for Py-GC/MS studies were selected through a previous exploratory
187 analysis by visible (UV-Vis) and infrared (FT-IR) spectroscopy [19]. The results are
188 summarized in Table 2 and Fig. 1 depicting the spectroscopic characteristics of two selected
189 samples showing contrasting levels of residual lignin (i.e., high (19) and weak (16) lignin
190 signature). The high lignin signature sample 19 displays IR conspicuous peaks at 1510, 1460,
191 1420, 1270, 1230 and 1030 cm^{-1} , just coinciding with a weak intensity of the peaks revealed
192 in the second derivative (as valleys) in the visible spectrum at *ca.* 620, 570 and 530 cm^{-1} ,
193 typically interpreted as a biomarker feature of specific fungal metabolism [31]. The opposed
194 situation is observed in sample 16, the HA with less marked lignin pattern and high optical
195 density. Due to the suspected importance of the more or less defined lignin signature

196 observed in the resolution-enhanced FT-IR spectra, samples were ranked from high to low
197 similarity to lignin and labelled with an ordinal (IR Lignin) used for data treatments as
198 shown in Table 2.

199 *3.1. Pyrolytic characterization of the HAs*

200 Up to 250 different compounds could be detected in the pyrolysates of the studied samples
201 and their relative yields with an indication of the possible precursor compounds [32] is
202 shown in Table 3, where the relative abundance of the groups of compounds is coded with
203 symbols indicating the levels of the semiquantitative percentages related to the total
204 chromatographic area.

205 Methoxyphenols have been considered as an index for the persistence of plant-derived lignin
206 in the less transformed structural domains of the HAs [33, 34]. Thus, accumulation of
207 methoxyphenols would point to comparatively early humification processes with a weak
208 structural alteration of macromolecular lignin, not necessarily requiring its complete
209 breakdown into low molecular weight compounds.

210 When the HA samples are classified in terms of the more or less marked lignin pattern as
211 seen in the IR spectra, it was possible to classify the samples in a gradient ranging between
212 two extreme categories as previously illustrated in Fig. 1. The group of HAs showing ‘weak
213 lignin pattern’ was characterized by major peaks of N-containing and carbohydrate-derived
214 pyrolysis compounds, and poor yields of alkyl compounds (e.g., HA samples coded as
215 M,N,O,P (samples #12, 4, 2, and 16), indicative of microbial metabolites accumulation and a
216 more or less efficient biodegradation of plant-derived HA precursors. The major pyrolytic
217 products in these samples consisted of non-methoxylated aromatic compounds (mainly
218 alkylbenzenes), i.e., compounds that have been classically interpreted as typical pyrolytic
219 proxies for matured HAs from terrestrial sources [35]. Despite the low potential of these

220 alkyl-substituted aromatic compounds as source indicator compounds, and to the fact that
221 some might be derived from rearrangements of unsaturated aliphatic compounds during
222 pyrolysis known to occur mainly in the presence of mineral catalysts [36], it is clear that both
223 the stoichiometry of the pyrolysis compounds assemblages as well as the structure of the
224 major fragments in this group of samples (Figs. 2 and 3) clearly differ from those expected
225 from biomass constituents, as it would be the case for soil HAs for which C sequestration
226 mechanisms could be assimilated to a selective preservation of plant and microbial
227 constituents.

228 On the other side, the HA group showing a 'conspicuous lignin pattern' display a well-
229 preserved methoxyphenol signature (i.e., guaiacol, syringol and their methyl-, ethyl-, vinyl-,
230 propenyl- and acetyl- derivatives) and comparatively lower amounts of alkyl and N-derived
231 pyrolysis products (HAs coded as A,B,C,D,E,F, corresponding to samples # 19, 10, 20, 9, 14
232 and 3).

233 Intermediate patterns between these two extreme situations were observed in HAs samples
234 coded as G,H,I,J,K,L, corresponding to HAs # 11, 8, 6, 15, 17, 7, respectively illustrating the
235 above-suggested simultaneous active humification processes contributing to SOM
236 stabilization which also coincides with the occurrence of the mixed vegetation in these sites.

237 In some cases, the group of HAs with features pointing to a slow biodegradation of
238 the original sources of SOM, *i.e.* the 'conspicuous lignin pattern', yielded aromatic pyrolysis
239 products of relatively high molecular weight (e.g., biphenyls, compounds 149, 173, 126, 201)
240 and well-defined, wide homologues series of alkanes, α,ω -alkanediene and fatty acids. An
241 outstanding systematic feature, particularly in HAs from soils developed under natural pine
242 forest, was the release after pyrolysis of polyaromatic compounds (Fig. 2) such as
243 phenanthrenes, retene and diterpene resin acids (e.g., dehydroabietic acid), which coincided
244 with previous pyrolytic descriptors for pine soils in Mediterranean sites [4, 37].

245 The above differences between HA characteristics are in agreement with the previous
246 exploratory analysis using only visible and IR spectroscopies, which also indicated a series
247 of HAs with featureless IR spectra, high optical density (e.g., at 465 nm) and resolved peaks
248 in the 2nd derivative visible spectra characteristic of polynuclear chromophors of fungal
249 origin [31], which coincided with the ‘weak lignin pattern’ group of samples (Fig. 2).
250 Another group of HAs presenting a comparatively marked aliphatic character and low
251 molecular weight (intense 2920 cm⁻¹ IR band, low optical density and IR spectra displaying
252 typical lignin and amide bands), could be ascribed to group of samples with the ‘conspicuous
253 methoxyphenol pattern’ betrayed by analytical pyrolysis.

254

255 **4. Discussion**

256 The results from Py-GC/MS suggested a variety of structural features in active
257 biogeochemical scenarios which could be ordered as a gradient of progressive humification
258 or SOM maturation [38]. This could be operatively interpreted as the results of a balance
259 between non-excludent soil C sequestration processes, *i.e.* HAs with lignin domains at early
260 alteration stages *vs.* HAs formed after heavy structural rearrangement of humic precursors
261 not necessarily with a macromolecular nature.

262 In fact, in one extreme situation, soil HAs with spectroscopic and Py-GC/MS patterns
263 indicative of heterogeneous composition suggest efficient microbial reworking of the SOM
264 precursors, through humification mechanisms associating products from both plant or
265 microbial synthesis, as well as secondary molecules and mixtures of oligomers released
266 during litter biodegradation. The Ca²⁺-saturated medium would favour the insolubilization of
267 these humic precursors to be progressively arranged into supramacromolecular mixtures. The
268 broadband IR spectra of these HAs suggest few repeating structural units, and its high optical
269 density is compatible with an advanced diagenetic ‘maturation’, *i.e.*, selective biodegradation

270 of labile HA moieties and free-radical-induced cross-linking between components that are
271 being randomly incorporated into the HA systems [3].

272 In other extreme situation, the characteristics of the HAs point to the preservation of
273 comparatively young, lignoprotein-type substances, as betrayed by correlated spectroscopic
274 and Py-GC/MS patterns suggesting that diagenetic stabilization of plant biomacromolecules
275 and aliphatic structures are predominant C-sequestration mechanisms in the corresponding
276 soils. Irrespectively to the more or less conspicuous lignin signature in the HAs, the sample
277 groups did not significantly differ in terms of classical stoichiometric indices based on the
278 yields of methoxyphenols (e.g., the classical syringyl-to-guaiacyl ratio) which in many cases
279 lead to accurately distinguish between land use or vegetation cover [39]. This is not
280 unexpected in our case due to the large heterogeneity in terms of vegetation, soil use and
281 HAs degree of maturity.

282 The above features are summarized in Fig. 4 where a variety of spectroscopic, pyrolytic and
283 soil general characteristics with environmental relevance are classified by multidimensional
284 scaling. This nonlinear mapping statistical procedure yields a reduced-dimensionality plot
285 where the operative taxonomical units (in this case variables) are represented as points in the
286 space defined by two dimensions calculated in a way in which the distances in the space are
287 optimized to most closely fit the values of a similarity index (in this case the 1-Pearson
288 correlation index) between the processed variables.

289 In this plot it is evident the association between the 'lignin signature in HAs' with the above
290 indicated pyrolytic yields of methoxyphenols and aliphatic compounds, but also with soil
291 physical and microtopographical features classically associated to moisture levels and
292 accumulation of raw SOM (e.g., FOM). The other cluster of variables included
293 characteristics typical of condensed, black-coloured HAs including fungal quinoid
294 metabolites (E_{465} , Vis 620) in addition to non-methoxylated mono- or polycyclic pyrolysis

295 products. These HA characteristics appear correlated with physical and topographical factors
296 pointing to more conspicuous semiarid features in presumably flat, not intensely eroded areas
297 where clay accumulation and seasonal desiccation would presumably contribute to reach
298 advanced maturation of HAs. In fact, these soil-forming factors are classically described as
299 positive for the humification, *i.e.*, formation of recalcitrant HA-clay complexes and periodic
300 desiccation enhancing SOM insolubilization and condensation processes [14].

301 Additional statistical treatments mainly correspondence- and discriminant analyses (not
302 shown) failed in evidencing a substantial influence of vegetation types, soil use or direct
303 anthropogenic impact, a situation that has been previously described for similar calcimorphic
304 semiarid mountain ecosystems [12] and attributed to the homogeneizing or ‘buffering’
305 physical and physiochemical effect of limestone (calcium saturation) on the different organic
306 matter forms (precipitation, encapsulation...). In a similar way, all unsupervised automated
307 classification of the HAs from our soils suggested that the continuous ‘gradient’ between
308 contrasting ‘C-sequestration pathways’ observed in our data, was quite independent as
309 regards to local features of soil use and vegetation and the major source of variability being
310 local-scale due to geomorphological factors [40] which are not currently recorded in most
311 studies on SOM dynamics, but that could play a relevant role in most calcium-saturated soils
312 where humification processes are not controlled by wide contrasts in soil reaction.

313

314 **5. Conclusions**

315 The analysis of the lignin-derived pyrolytic molecular assemblages in HAs suggests a
316 series of surrogate indicators of SOM quality based on the variable influence of soil C
317 stabilization mechanisms. This was the case with the relative yields of methoxyphenols
318 (which in this study paralleled the intensity of the lignin pattern in the IR spectra, and were
319 negatively related to the E_{465} optical density in the visible spectra) but also with the N-

320 containing, carbohydrate-derived and alkyl compounds. In general the above HA features
321 could be useful for rapid discrimination of the prevailing humification processes influencing
322 C-balance in calcimorphic soils . In this study, SOM quality (i.e., its most advanced stages of
323 transformation) could be assimilated to the extent to which the HAs accumulated in the soil
324 differ in its molecular composition as regards to the biomacromolecules from microbial and
325 plant sources. For instance, the incorporation of specific biomarkers such as conifer resin
326 constituents into the HA structure could represent a valid proxy for raw SOM preservation in
327 forest ecosystems with comparatively low biogeochemical performance. On the opposite,
328 conspicuous concentration of perylenequinone fungal chromophors in HAs (as seen in the
329 2nd derivative visible spectra) would be pointing to the occurrence of intense microbial
330 reworking of SOM leading to chaotic HA structures (broadband IR spectral profiles) and
331 where the lignin spectroscopic signature is no longer evident by spectroscopic or pyrolytic
332 approaches.

333

334 **Acknowledgements**

335 The authors wish to thank the Spanish Ministry of Science and Innovation by CTM2005-
336 04739 / CGL2008-04926 / CARBORAR CGL2011-27493 / GEOFIRE CGL2012-38655-
337 C04-01 research projects. Dr. Ana Piedra Buena has been contracted by the CCMA-CSIC via
338 I3P Program supported by the European Social Fund. Dr. Isabel Miralles received
339 postdoctoral fellowships Juan de la Cierva 2008-39669 and a Marie Curie Intra-European
340 Fellowship (FP7-PEOPLE-2013-IEF, Proposal No. 623393)

341

342

343

344 **References**

- 345 [1] R. Lal, Soil carbon sequestration impacts on global climate change and food security,
346 Science 304 (2004) 1623–1627.
- 347
- 348 [2] M. von Lützow, I. Kögel-Knabner, K. Ekschmitt, E. Matzner, G. Guggenberger, B.
349 Marschner, H. Flessa, Stabilization of organic matter in temperate soils: Mechanisms
350 and their relevance under different soil conditions—A review, Eur. J. Soil. Sci. 57
351 (2006) 426–445.
- 352
- 353 [3] G. Almendros, Humic substances, In: Ward Chesworth (ed) Kluwer Encyclopedia of Soil
354 Science. Springer, Dordrecht 2008, pp. 97–99.
- 355
- 356 [4] J. Schellekens, G.G. Barberá, P. Buurman, G. Pérez-Jordà, A. Martínez-Cortizas, Soil
357 organic matter dynamics in Mediterranean A-horizons—The use of analytical pyrolysis
358 to ascertain land-use history, J. Anal. Appl. Pyrol. 104 (2013) 287–298.
- 359
- 360 [5] F.J. Stevenson, Biochemistry of the formation of humic substances, In: Humus
361 Chemistry, Wiley & Sons, Inc. New York 1982, pp. 195–220.
- 362
- 363 [6] M.H.B. Hayes, R.S. Swift, Genesis, isolation, composition and structures of soil humic
364 substances. Soil Colloids and Their Associations in Aggregates, In: De Boodt MF,
365 Hayes MHB and Herbillon A (eds) NATO ASI Series. Series B: Physics 215 (1990)
366 245–305.
- 367

- 368 [7] R. Sutton, G. Sposito, Molecular structure in soil humic substances: The new view,
369 Environ. Sci. Technol. 39 (2005) 9009–9015.
- 370
- 371 [8] A. Piccolo, The supramolecular structure of humic substances, Soil Sci. 166 (2001) 810–
372 832.
- 373
- 374 [9] A. Piccolo, The supramolecular structure of humic substances: A novel understanding of
375 humus chemistry and implications in soil science (Review), Adv. Agron. 75 (2002) 57–
376 134.
- 377
- 378 [10] G. Almendros, Z. Hernández, H. Knicker, J.A. González-Pérez, A. Piedra-Buena, F.J.
379 González-Vila, Humic substances as drivers of carbon flow and source of
380 biogeochemical proxies for climate change. MOLTER Conference 'SOM 2010', Organic
381 matter stabilization and ecosystem functions, p.67. Presqu'île de Giens (Côte d'Azur,
382 France) 19-23 September 2010. Available in the web at:
383 [http://www.bioforsk.no/ikbViewer/Content/83713/Proceedings%20SOM2010%281%29](http://www.bioforsk.no/ikbViewer/Content/83713/Proceedings%20SOM2010%281%29.pdf)
384 [.pdf](http://www.bioforsk.no/ikbViewer/Content/83713/Proceedings%20SOM2010%281%29.pdf) (Last consulted 27/010/2014)
- 385
- 386 [11] M.W. Schmidt, M.S. Torn, S. Abiven, T. Dittmar, G. Guggenberger, I.A. Janssens, M.
387 Kleber, I. Kögel-Knabner, J. Lehmann, D.A. Manning, P. Nannipieri, D.P. Rasse, S.
388 Weiner, S.E. Trumbore, Persistence of soil organic matter as an ecosystem property,
389 Nature. 478 (2011) 49–56.
- 390
- 391 [12] C. Oyonarte, A. Perez-Pujalte, G. Delgado, R. Delgado, G. Almendros, Factors
392 affecting soil organic matter turnover in a Mediterranean ecosystems from Sierra de

- 393 Gador (Spain): An analytical approach, *Commun. Soil Sci. Plant. Anal.* 25 (1994)
394 1929–1945.
395
- 396 [13] S. Derenne, K. Knicker, Chemical structure and preservation processes of organic
397 matter in soils and sediments, *Org. Geochem.* 31 (2000) 607–608.
398
- 399 [14] P. Duchaufour, F. Jacquin, Comparaison des processus d’humification dans les
400 principaux types d’humus forestiers, *Bulletin Association Française pour l’Etude des*
401 *Sols* 1 (1975) 29–36.
402
- 403 [15] T.I. Stuczynski, G.W. McCarthy, J.B. Reeves, R.J. Wright, Use of pyrolysis GC/MS for
404 assessing changes in soil organic matter quality, *Soil. Sci.* 162 (1997) 97–105.
405
- 406 [16] P.F. van Bergen, I.D. Bull, P.R. Poulton, R.P. Evershed, Organic geochemical studies
407 of soils from the Rothamsted Classical Experiments—I. Total lipid extracts, solvent
408 insoluble residues and humic acids from Broadbalk Wilderness, *Org. Geochem.* 26
409 (1997) 117–135.
410
- 411 [17] P. Buurman, P.F. van Bergen, A.G. Jongmans, E.L. Meijer, B. Duran, B. van Lagen,
412 Spatial and temporal variation in podzol organic matter studied by pyrolysis-gas
413 chromatography/mass spectrometry and micromorphology, *Eur. J. Soil. Sci.* 56 (2005)
414 253–270.
415
- 416 [18] S. Rivas Martínez. Memoria del Mapa de Series de Vegetación de España 1:400.000.
417 ICONA, Madrid, 1987.

418

419 [19] I. Miralles, R. Ortega, M. Sánchez-Marañón, M. Soriano, G. Almendros, Assessment of
420 biogeochemical trends in soil organic matter sequestration in Mediterranean
421 calcimorphic mountain soils (Almería, Southern Spain), 324 Soil Biol. Biochem. 39
422 (2007) 2459–2470.

423

424 [20] CSIC, Métodos Analíticos de la Estación Experimental del Zaidín. CSIC, Granada
425 1969, Spain.

426

427 [21] L.A. Richards, Diagnosis and Improvement of Saline and Alkaline Soils, United States
428 Salinity Laboratory Staff. Agricultural Handbook No 60. United States Department of
429 Agriculture 1954, 160 p.

430

431 [22] D.V. Nelson, L.E. Sommers, Total carbon, organic carbon and organic matter, In: Page
432 AL, Miller RH, Keeney DR (eds) Methods of Soil Analysis: Part 2, Chemical and
433 Microbiological Properties, 2nd ed. American Society of Agronomy, Madison 1982,
434 WI, pp. 539–579.

435

436 [23] M.E. Sumner, W.P. Miller, Cation exchange capacity and exchange coefficients, In:
437 Sparks D.L. (ed) Methods of Soil Analysis, Part 3. Chemical Methods. Soil Science
438 Society of America 1996, Book series No. 5.

439

440 [24] B. Dabin, Etude d'une méthode d'extraction de la matière humique du sol, Science du
441 Sol 1 (1971) 47–63.

442

- 443 [25] S.J. Traina, J. Novak, N.E. Smeck, An ultraviolet absorbance method of estimating the
444 percent aromatic carbon content of humic acids, *J. Environ. Qual.* 19 (1990) 151–153.
445
- 446 [26] A. Rosenfeld, A.C. Kak, *Digital Picture Processing*, vol 1. 2nd ed. Academic Press, NY
447 1982.
448
- 449 [27] G. Almendros, J. Sanz, A structural study of alkyl polymers in soil after perborate
450 degradation of humin, *Geoderma* 53 (1992) 79–95.
451
- 452 [28] C.P. Assis, J.A. González-Pérez, J.M. de la Rosa, I. Jucksch, E.D. Mendonça, F.J.
453 González- Vila, Analytical pyrolysis of humic substances from a Latosol (Typic
454 Hapludox) under different land uses in Minas Gerais, Brazil, *J. Anal. Appl. Pyr.* 93
455 (2012) 120–128.
456
- 457 [29] StatSoft Inc., *STATISTICA Data Analysis Software System*, version 7.1. (2005)
458 www.statsoft.com
459
- 460 [30] J.B. Kruskal, Multidimensional scaling by optimising goodness of fit to a nonmetric
461 hypothesis, *Psychometrika* 29 (1964) 1–27.
462
- 463 [31] K. Kumada, H.M. Hurst, Green humic acid and its possible origin as a fungal
464 metabolite, *Nature* 214 (1967) 631–633.
465
- 466 [32] D.F.W. Naafs. What are humic substances? A molecular approach to the study of
467 organic matter in acid soils. PhD Thesis. Universiteit Utrecht, 2004.

468

469 [33] F. Martín, C. Saíz-Jiménez, F.J. González-Vila, Pyrolysis-gas chromatography-mass
470 spectrometry of lignins, *Holzforschung* 33 (1979) 210–212.

471

472 [34] P. Tinoco, G. Almendros, F.J. González-Vila, Impact of the vegetation on the lignin
473 pyrolytic signature of soil humic acids from Mediterranean soils, *J. Anal. Appl. Pyrol.*
474 64 (2002) 407–420.

475

476 [35] H.R. Schulten, M. Schnitzer, A state of the art structural concept for humic substances,
477 *Naturwissenschaften*. 80 (1993) 29–30.

478

479 [36] C. Saíz-Jiménez, Analytical pyrolysis of humic substances: Pitfalls, limitations, and
480 possible solutions. *Environ. Sci. Technol.* 28 (1994) 1773–1780.

481

482 [37] N.T. Jiménez-Morillo, J.A. González-Pérez, A. Jordán, L.M. Zavala, J.M. de la Rosa,
483 M. Jiménez-González, F.J. González-Vila. Organic matter fractions controlling soil
484 water repellency in sandy soils from the Doñana National Park (SW Spain). *Land*
485 *Degrad. Dev.* DOI: 10.1002/ldr.2314 (First published online: 5 AUG 2014).

486

487 [38] I. Miralles, R. Ortega, G. Almendros, M. Sánchez-Marañón, M. Soriano, Soil quality
488 and organic carbon ratios in mountain agroecosystems of South-East Spain, *Geoderma*
489 150 (2009) 120–128.

490

491 [39] C.N. Jex, G.H. Pate, A.J. Blyth, R.G.M. Spencer, P.J. Hernes, S.J. Khan, A. Baker,
492 Lignin biogeochemistry: from modern processes to Quaternary Archives. Quaternary
493 Sci. Rev. 87 (2014) 46–59.

494

495 [40] A. Piedra-Buena, I. Miralles, R. Ortega, Z. Hernández, G. Almendros, Molecular
496 proxies responsive to the effect of microtopographical constraints on soil C
497 sequestration processes in mountain Mediterranean ecosystems (Almería, Spain). Soil
498 Geography: New Horizons. 16-20 Noviembre. Huatulco, Santa Cruz, Oaxaca, México
499 (2009).

500

501

502

503

504

505

505 **FIGURE CAPTIONS**

506 Fig. 1. Spectroscopic patterns of humic acids from semiarid calcimorphic soils representative
507 of extreme levels of residual lignin (left: sample 19 with high lignin signature; right sample
508 16 with low lignin signature). Top: IR spectra in the 2000–700 cm^{-1} range, superimposed to
509 the corresponding resolution-enhanced IR spectra. Below: visible spectra and their 2nd
510 derivatives.

511

512 Fig. 2. Relative amounts of different group of pyrolysis products in humic acids from
513 semiarid mountain calcimorphic soils, classified by lignin pattern group as originally defined
514 by the intensity of the lignin pattern in the IR resolution-enhanced IR spectra (Fig. 1).

515

516 Fig. 3. Average values and confidence limits of relative amounts of the pyrolysis products in
517 humic acids from semiarid mountain calcimorphic soils classified by compound group.

518

519 Fig. 4. Automatic classification by multidimensional scaling (final stress = 0.175) of
520 variables corresponding to HA characteristics (solid circles) together with soil analytical
521 properties and environmental factors with a bearing on the humification process (void
522 circles), using the 1-Pearson index as similarity criterion. The plot shows two extreme
523 situations of variable characteristics: a) humic acids presumably derived from diagenetic
524 transformation of lignins preserved in soil; b) variables associated to HAs in advanced
525 transformation stages where lignin signature is not evident. Variables in the middle of the
526 scatterdiagram show no strong trends as regards transformation of lignin in the studied soils.

527 Variables obtained with a Geographical Information System: GSR_E: global solar radiation
528 at equinox; GSR_S: global solar radiation at summer solstice; GSR_W: global solar radiation

529 at winter solstice; HS_E: hours of insolation at equinox; HS_S: hours of insolation at
530 summer solstice; HS_W: hours of insolation at winter solstice; LSF: length slope factor; R:
531 contributing area; SPlanC: slope plan curvature; SProfileC: slope profile curvature; W:
532 wetness index.

533 Soil physical and chemical variables and humic acid characteristics (solid circles):

534 C: Total soil C; E₄₆₅: optical density of the HA at 465 nm; FOM: concentration of free
535 organic matter; HA: concentration of soil humic acid; IR Lignin: intensity of the lignin
536 pattern in the IR spectra; Sand, Clay: granulometric fractions; Tot Por: Total porosity; Vis
537 620: absorption at 620 nm in the 2nd derivative visible spectrum.

538 Total abundances of the main groups of pyrolysis compounds are shown with their formulas

539

540

541 **Table 1** General characteristics of soils in mountain calcimorphic semiarid ecosystems

Sample/ code ^a	Soil type	Vegetation ^b	Geological substrate	Total sand (2–0.02 mm) g kg ⁻¹	Total clay (<0.002 mm) g kg ⁻¹	WHC g cm ⁻³	pH (H ₂ O)	CaCO ₃ g kg ⁻¹	Soil C g kg ⁻¹	C/N	CEC cmol _c kg ⁻¹
2 / O	Rendzic Leptosol	Brush encroached	Clastic limestone	391	237	140	8.3	528	18.6	11.5	21.5
3 / F	Hypocalcic Calcisol	Relictual oak	Detritic limestones and marls	327	397	283	7.7	278	81.2	18.8	41.9
4 / N	Hypercalcic Calcisol	Reforested pine	Detritic limestones	597	189	117	8.0	673	31.3	13.9	25.0
6 / I	Hypercalcic Calcisol	Climacic pine	Clastic limestones	414	292	193	8.2	313	32.2	16.1	30.7
7 / L	Hypercalcic Petric Calcisol	Cereal	Detritic sediments	572	173	118	8.5	588	13.2	12.4	12.7
8 / H	Petric Calcisol	Reafforested pine forest (<i>Pinus halepensis</i>)	Marls and limestones	513	108	59	8.4	940	13.2	16.0	9.2
9 / D	Calcic Chernozem	Climacic pine	Detritic limestones	257	272	341	7.9	360	85.4	29.8	53.9
10 / B	Calcic Chernozem	Climacic pine	Detritic limestones and dolomies	344	324	267	8.1	544	74.4	37.7	7.9
11 / G	Calcaric Rendzic Leptosol	Alpine brush	Limestones	110	490	412	7.7	46	107.4	10.3	50.4
12 / M	Gleyic Hypocalcic Calcisol	Orchard	Alluvial calcic marls	69	668	209	8.3	439	6.9	8.1	24.8
14 / E	Mollic Calcaric Cambisol	Relictual oak	Limestones	74	554	252	7.6	5	50.3	14.3	37.6
15 / J	Hypercalcic Petric Calcisol	Chaparral-like brushland	Limestones	161	475	290	8.0	294	49.0	13.5	35.9
16 / P	Calcaric Rendzic Leptosol	Chaparral-like brushland	Limestones	64	516	325	7.9	42	52.9	8.0	37.1
17 / K	Hypercalcic Calcisol	Cereal	Limestones	331	293	161	8.2	674	23.4	9.7	17.1
19 / A	Mollic Calcaric Cambisol	Climacic pine	Detritic limestone material	508	309	521	7.3	248	195.3	15.1	46.2
20 / C	Calcaric Rendic Leptosol	Alpine brush	Limestones	203	482	238	7.9	380	25.0	3.9	26.9

542 ^a One-character code used in Table 3 to refer the contribution of lignin in the HA structure543 ^b Reforested pine: *Pinus halepensis*; Bush encroached cleared site: *Stipa tenacissima*, *Lygeum spartum*, *Genista scorpius*, *Artemisia* sp.; Relictual oak:
544 *Quercus ilex* ssp. *rotundifolia*; Orchard: almond trees; Climacic pine: *Pinus halepensis*; Climacic pine: *Pinus nigra* (9, 10), *Pinus halepensis* (19); Alpine
545 brush: *Juniperus oxycedrus*, *Vella spinosa*, *Erinacea antillis*, *Quercus coccifera* (11), *Vella spinosa*, *Erinacea antillis*, *Lygeum spartum* (20); Chaparral-like
546 brushlands: *Quercus ilex* ssp. *rotundifolia*, *Juniperus phoenicia*. WHC: water holding capacity; CEC: cation exchange capacity.

Table 2 Soil organic matter characteristics and spectroscopic properties of humic acids from mountain calcimorphic semiarid soils

Sample/ code ^a	FOM (C g 100 g soil ⁻¹)	HA (C g 100 g soil ⁻¹)	E ₄₆₅	E 620	IR Lignin
2 / O	1.2	21.3	1.23	0.0024	5
3 / F	2.9	25.5	0.56	0.0016	14
4 / N	0.6	54.3	0.66	0.0018	3
6 / I	2.0	57.7	0.71	0.0019	13
7 / L	2.5	52.0	0.97	0.0018	12
8 / H	2.4	51.9	0.63	0.0021	11
9 / D	1.6	10.0	0.86	0.0020	10
10 / B	1.0	10.5	0.73	0.0019	9
11 / G	6.5	89.7	0.84	0.0020	6
12 / M	0.9	63.2	1.26	0.0020	2
14 / E	4.3	23.8	0.62	0.0009	8
15 / J	4.3	12.2	0.95	0.0021	4
16 / P	1.4	2.7	1.10	0.0019	1
17 / K	1.5	37.6	0.87	0.0021	7
19 / A	4.0	13.5	0.52	0.0015	15
20 / C	2.5	7.4	0.64	0.0028	16

^a One-character code used in Table 3 to refer the contribution of lignin in the HA structure.

FOM: Free organic matter, HA: Humic acid; E₄₆₅: Optical density values at 625 nm in the visible range in absorption units; E 620: Intensity of the valley at 620 nm in the second derivative spectrum in absorption units; IR Lignin: relative intensity of the lignin pattern in the IR spectra.

Compound	Lignin pattern type			Compound	Lignin pattern type			Compound	Lignin pattern type		
	Clear ABCDEF	Average GHIJKL	Weak MNOP		Clear ABCDEF	Average GHIJKL	Weak MNOP		Clear ABCDEF	Average GHIJKL	Weak MNOP
142 Propylguaiaicol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	176 Guaiacylacetone I [Ps]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	210 Acetylsyringol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
143 C ₁₁ -Alkene [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	177 Dibenzofuran [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	211 C ₁₅ -Alkene I [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
144 Benzonitrile [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	178 C ₃ -Naphthalene I [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	212 C ₁₅ -Alkene II [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
145 Hydroxybenzonitrile [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	179 C ₃ -Naphthalene II [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	213 Phenanthrene [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
146 Trimethylindene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	180 Isocyanonaphthalene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	214 C ₁₅ iso-Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
147 Ethylcatechol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	181 C ₁₂ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	215 C ₁₅ anteiso-Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
148 C ₁₁ -Alkene [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	182 Hydroxyquinoline [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	216 Syringic acid	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
149 Biphenyl [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	183 Vanillic acid [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	217 C ₁₅ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
150 Methylindole [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	184 Vinylsyringol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	218 Dimethoxyphenol [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
151 Vanillin [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	185 C ₃ -Naphthalene III [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	219 C ₁₅ -Alkyl nitrile [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
152 C ₁₁ -Alkene II [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	186 C ₁₃ -Alkene [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	220 C ₁₆ -Alkene I [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
153 C ₁₁ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	187 Isoquinoleidone [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	221 Methylphenanthrene I [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
154 Alkene C ₁₁ [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	188 C ₃ -Naphthalene IV [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	222 C ₁₆ -Alkene II [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
155 Methylbenzofurandione [Ps]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	189 Fluorene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	223 C _{16:1} -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
156 Eugenol II [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	190 C ₁₃ -Alkene I [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	224 C ₁₆ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
157 C ₂ -Naphthalene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	191 C ₁₃ -Alkene II [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	225 C ₁₇ iso-Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
158 Trimethylbenzaldehyde I [Ps]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	192 C ₃ -Naphthalene V [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	226 C ₂ -Phenanthrene [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
159 C ₂ -Naphthalene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	193 Methoxyeugenol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	227 C ₁₇ anteiso-Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
160 Methylsyringol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	194 Phenylbenzenamine [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	228 C ₁₇ -Alkene [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
161 Eugenol III [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	195 C ₁₃ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	229 C ₁₇ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
162 C ₄ -Phenol II [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	196 Methylnaphthalenol	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	230 Octadecanenitrile [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
163 Hydroxyphenylethanone [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	197 Phenoxyphenol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	231 Hexadecanamide [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
164 Propenylguaiaicol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	198 Methoxydimethylindole [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	232 Dimethoxyphenanthrene [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
165 Coumaric acid [Ps]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	199 C ₁₀ -Alkybenzene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	233 C ₁₈ -Alkene [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
166 C ₂ -Naphthalene III [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	200 Propenylsyringol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	234 C _{18:1} -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
167 Phenylpyridine [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	201 Dimethylbiphenyl [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	235 C ₁₈ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
168 Isoindole [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	202 C ₁₄ -Alkene I [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	236 Trimethylphenanthrene [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
169 Isoindoleidone [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	203 C ₁₄ -Alkene II [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	237 Retene [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
170 Acetoguaiacone [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	204 Camazulene I [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	238 Dehydroabietane [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
171 C ₁₂ -Alkene I [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	205 C ₁₄ -Alkene III [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	239 Dehydroabietic acid [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
172 C ₁₂ -Alkene II [Al]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	206 Acetylsyringol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	240 Abietic acid [Tp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○
173 Methylbiphenyl [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	207 Camazulene II [Ar]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○				
174 Dimethylindoleidone [Pp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	208 C _{14:1} -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○				
175 Ethylsyringol [Lg]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○	209 C ₁₄ -Fatty acid [Lp]	○ ○ ○ ○ ○	○ ○ ○ ○ ○	○ ○ ○ ○ ○				

^a Sample sets: Lignin pattern, based in the observation of the IR spectra: "Clear"= ABCDEF: 19, 10, 20, 9, 14, 3, respectively; 'Average' = GHIJKL: 11, 8, 6, 15, 17, 7 respectively; 'weak' = MNOP: 12, 4, 2, 16, respectively (Figure 2).

Total abundance referred to total volatile compounds: ○ = 0 %; ○ = 0–2 %; ● = >2 %.

Origin: Lg: lignin; Ps: carbohydrate; Lp: lipid; Ar: unspecific aromatic (methoxyl-lacking); Pp: peptides; Tp: terpenoid. Roman numbers indicate different isomers.

Figure 1

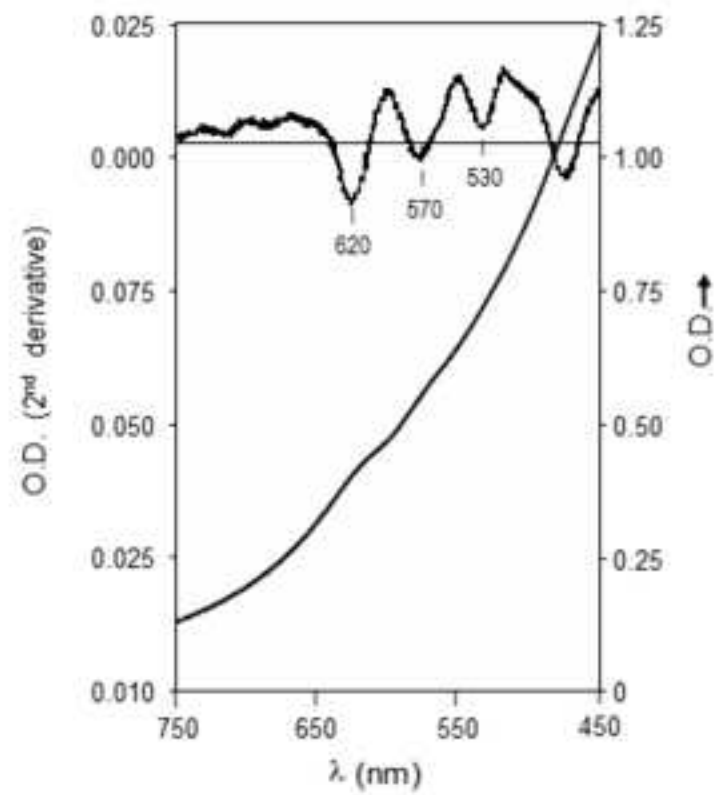
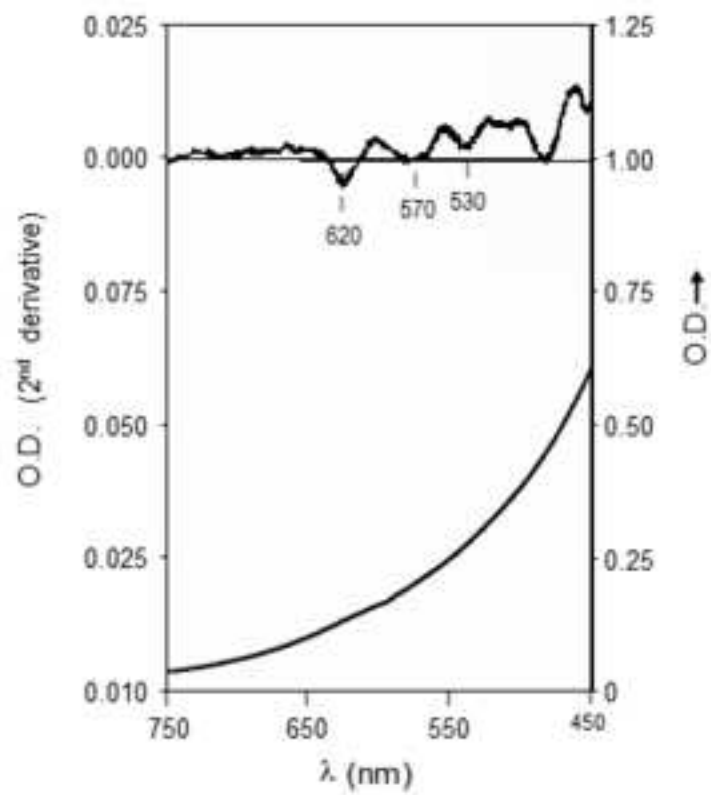
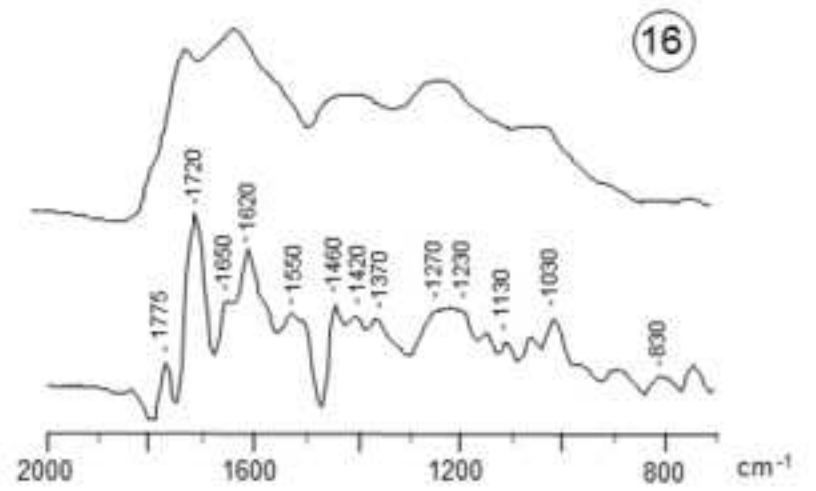
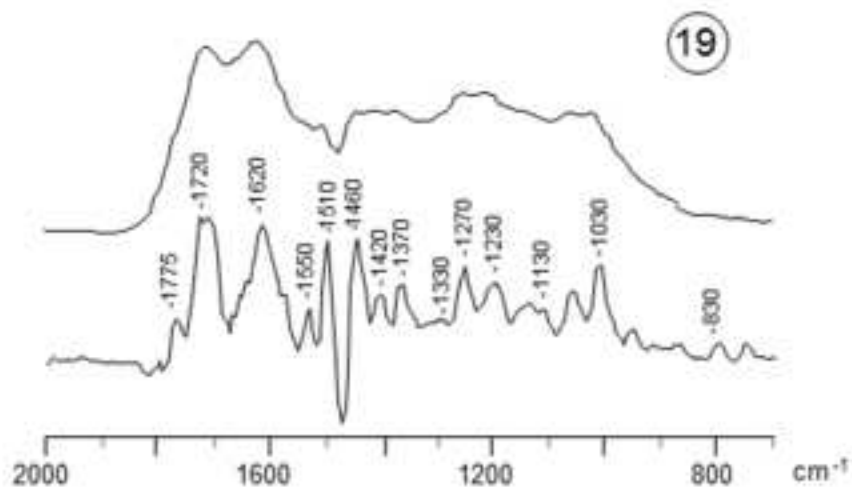
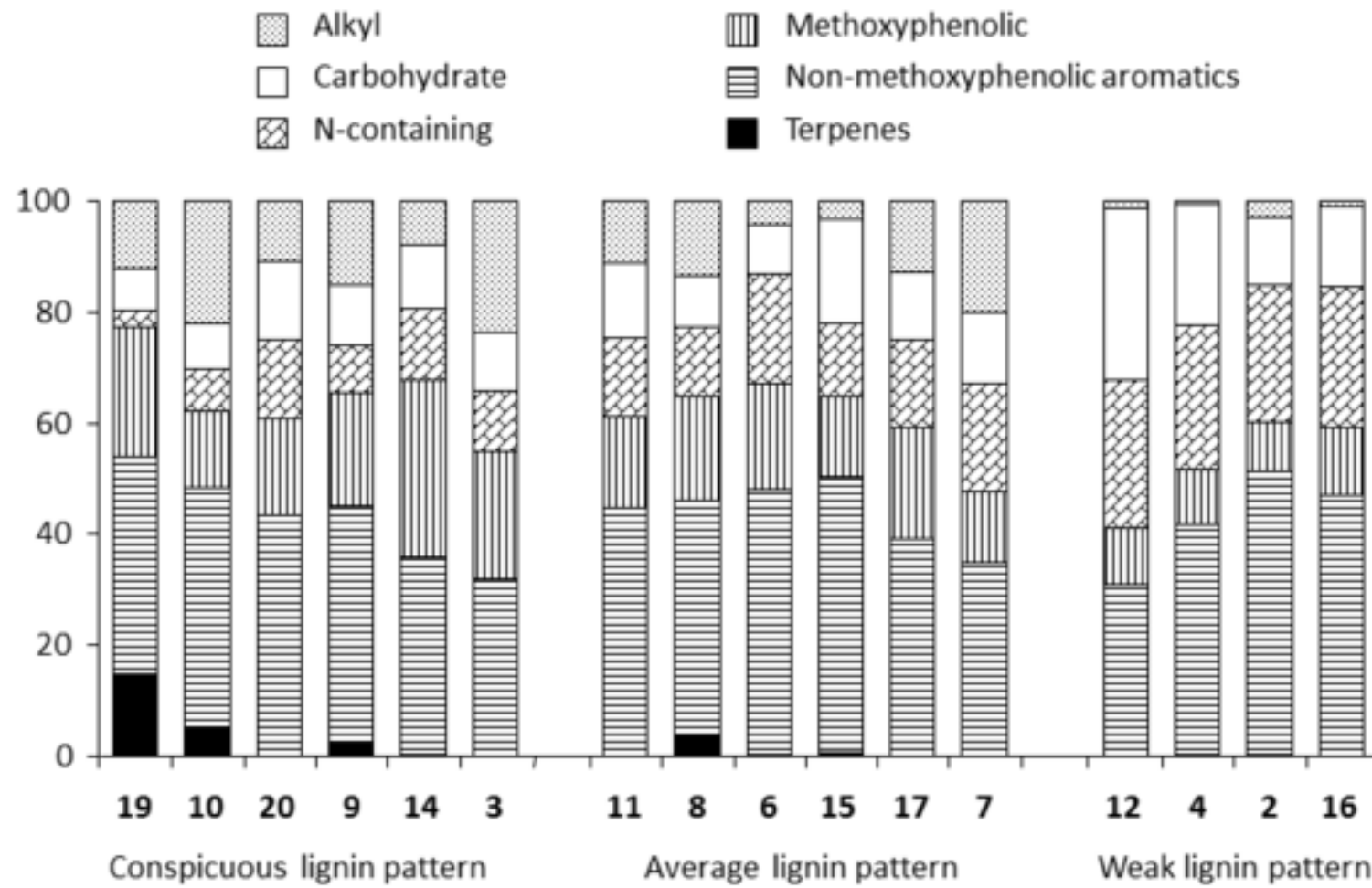


Figure 2



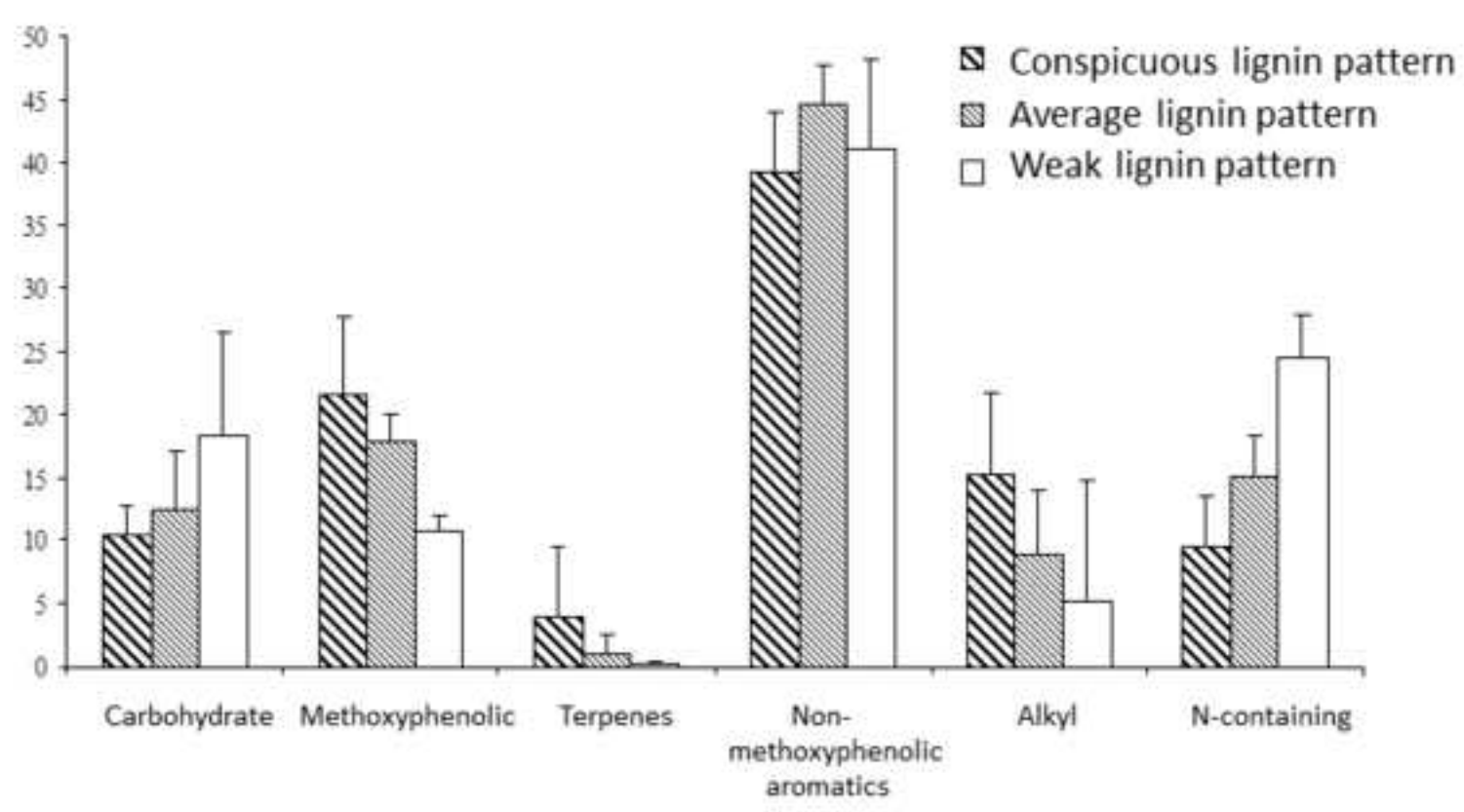
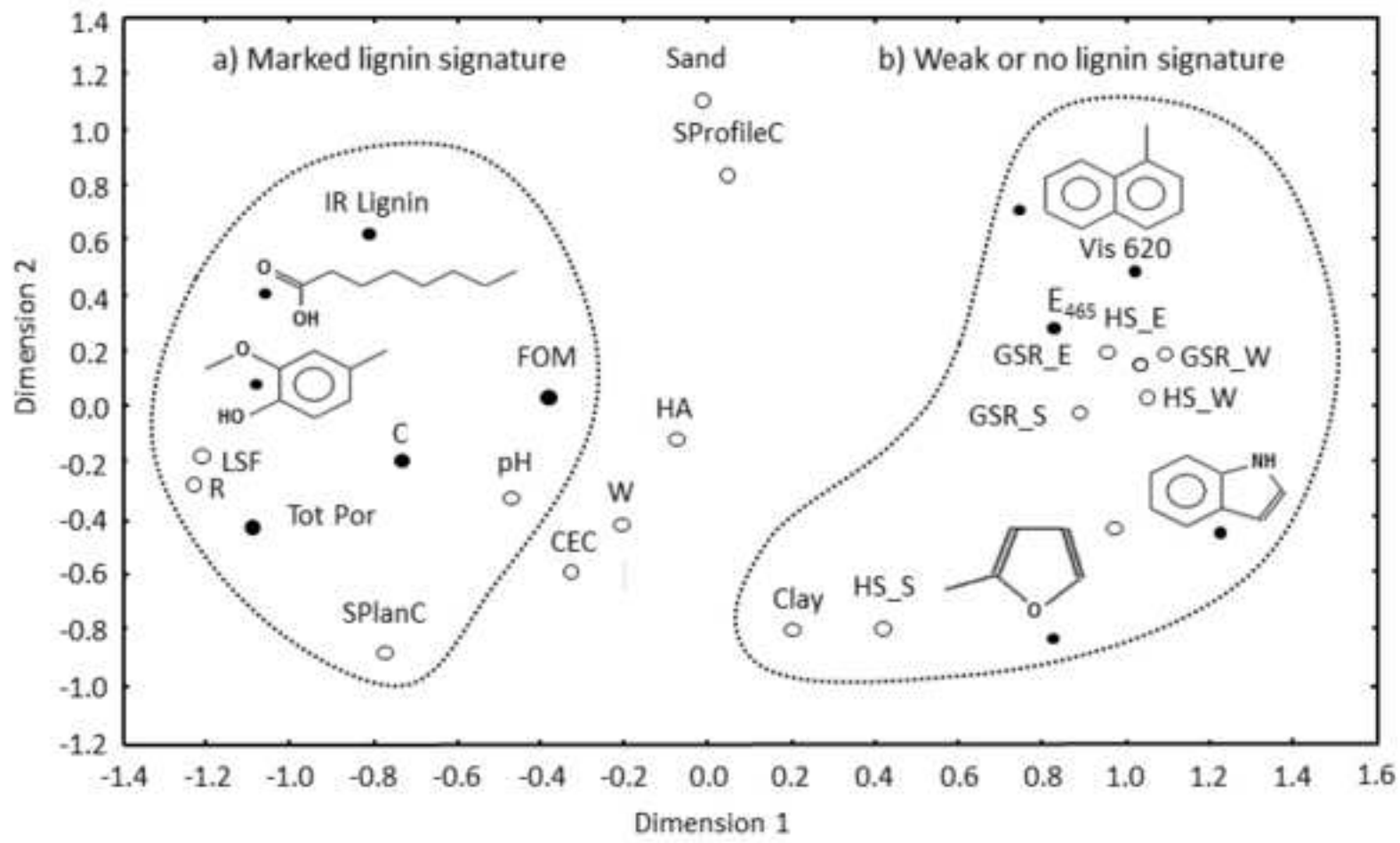


Figure 4



Highlights

Humic acids were studied by UV/VIS, FT-IR and Py-GC/MS

Pyrolytically-released methoxyphenols paralleled spectroscopic markers of lignin

Lignin signature in humic acids as an index for soil carbon stabilization

Accepted Manuscript