Heat Alteration of the Blue Pigment Aerinite: Application to Sixena's Romanesque Frescoes

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INTRODUCTION

Aerinite is a bluish fibrous inosilicate mineral associated with the alteration of mafic igneous rocks. Previous work has provided detailed information about the mineralogy of aerinite, including its chemical composition and crystal structure (Azambre & Monchoux, 1988; Rius *et al.*, 2004; Rius *et al.*, 2009). Aerinite is particularly interesting as it was used between the XI-XV centuries in many Romanesque fresco wall paintings of Catalonia, Aragon, and Andorra (Palet & de Andrés, 1992).

From the point of view of art conservation, studies on heat-induced mineral pigment alterations may provide useful information regarding the original state of a painting damaged by a fire event (Rickerby, 1991). This type of studies may be relevant for the knowledge and preservation of damaged paintings and even for their restoration. A characteristic example of such alterations can be found in the Romanesque wall paintings from the Chapter House of the Monastery of Santa Maria de Sixena (Huesca, Aragón, Spain). These frescoes, painted between 1196 and 1208, were highly damaged after the fire of 1936, with a high loss of pictorial surface and important color paintings alteration. The were transferred to canvas and are now exhibited at the Museu Nacional d'Art de Catalunya (MNAC), in Barcelona (see Fig. 1).

Aerinite has been identified as one of the blue pigments employed in Sixena's paintings (Palet & de Andrés, 1992). In the frescoes exhibited at MNAC, there exists a relatively large, blue-colored section that was not strongly affected by fire. Spectroscopic infrared transmission measurements confirm that aerinite was employed in this composition. For the preservation and complete understanding of these paintings, it would be desirable to determine which sections of the fire-damaged frescoes were originally blue. For this purpose, the thermal behavior of aerinite should be well understood.



fig 1. Paintings from the Chapter House in Sixena, on display at MNAC. ©MNAC – Museu Nacional d'Art de Catalunya. Barcelona. Photo by: Calveras / Mérida / Sagristà

The aim of the present work is to investigate the heat-induced alterations of the mineral pigment aerinite. The chromatic alterations of aerinite samples are monitored as a function of temperature and x-ray diffraction (XRD) is employed to investigate the heat induced mineralogical transformations of this mineral pigment.

SAMPLES AND METHODS

Aerinite specimens from the Estopinyà area (Huesca, Aragon, Spain) were subject to 1 h thermal annealing at the following temperatures: 100°, 200°, 400°, 600°, 800°, and 1000°C. Digital images of the annealed samples were obtained with conventional photographic techniques in order to record color alterations. For the XRD measurements, all samples were carefully hand-ground. XRD scans were recorded using a Bruker-AXS (Siemens) D5005 powder diffractometer equipped with a 2.2 kW sealed Cu x-ray source, a graphite monochromator to filter out the Cu Kß radiation, and a Nal(TI) scintillation detector. The scans were obtained between 4° and 60° (in 2θ) with a 0.05° step size and long integration times (6 seconds per step).

RESULTS AND DISCUSSION

The chemical alteration induced by heating of the mineral aerinite has an important bearing on the color of the annealed samples (Table I). At the lowest annealing temperatures, the blue color of aerinite is barely altered. At 400°C the samples acquire a green hue, which tends to be yellow-brown above 600 °C. At the highest temperature considered in this work (1000°C), the samples become red-brown.

Remarkably, the colors of the aerinite samples annealed at temperatures ranging from 600°C up to 1000°C can be easily observed in-situ in the firedamaged paintings from Sixena on display at MNAC. In particular, sections of the paintings that might be expected to be blue (following iconographic reasons) resemble very closely to the aerinite samples annealed at the higher temperatures.

The structural formula of aerinite as obtained from Syncrotron Powder Diffraction measurements on samples from the Estopinyà area is $Ca_{5.1}Na_{0.5}(Fe^{3+}AIFe^{2+}_{1.7}Mg_{0.3})(AI_{5.1}Mg_{0.7})$ $[Si_{12}O_{36}(OH)_{12}H][(CO_3)_{1.2}(H_2O)_{12}]$ (Rius et al., 2004). Note however that, owing to the chemical variability of aerinite, different structural results are obtained on specimens from other locations (Rius et al., 2009). On account of its chemical composition, it can be expected that upon heating, aerinite will suffer a loss of H₂O, OH and CO₃ groups at low or medium temperatures. At sufficiently elevated temperatures. the reorganization and recrystallization of the remaining atomic species may be expected to yield Ca, Fe, Mg, and Na

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polvo, Pigmento mineral azul	mineral pigment
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Temperature	not annealed	100 ℃	200 °C	400 °C	00 °C	800 °C	1000 ℃
Color	blue	blue	bluish	greenish	yellowish- brown-	light brown	brownish-red
Phases detected by X-ray diffraction	aerinite other phases	aerinite other phases	aerinite other phases	other phases	other phases	augite other phases	plagioclase, hematite, other phases

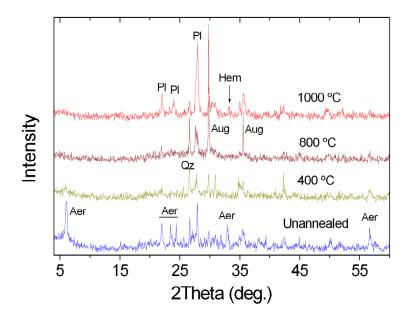
Table 1. Results of the heating tests on aerinite as a function of annealing temperature. 'Other phases' refers to minority, accompanying minerals like quartz or diopside that are observed in all the XRD scans, also in the non-annealed aerinite sample.

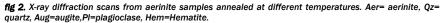
aluminosilicates.

Fig. 2 shows selected XRD scans of nonannealed aerinite together with aerinite samples annealed at different temperatures. The XRD spectrum of the untreated aerinite sample displays peaks from this phase as well as from other minerals present in the powder: quartz, feldspars (microcline and (close to plagioclase), augite the diopside end-member) and mesolite. The latter, which is a tectosilicate mineral associated to aerinite, is no longer detected in the annealed material.

With increasing temperature (Fig. 2), the intensity of the XRD reflections from aerinite is lowered, while the peaks from quartz and augite remain basically unchanged. At 600°C, no signal from aerinite is detected. At 800 °C, the XRD peaks augite, with general formula Ca(Fe,Mg)Si₂O₆, are sizably more intense. We attribute this observation to the partial recrystallization of the atomic species that remain after the degradation of aerinite. It should be noted that the temperature required for the growth of the Mg-rich and Fe-rich end members of the augite solid

solution (diopside and hedenbergite) is in the 800-950°C range (Deer et al., 1992). In the XRD scans of the sample annealed at 1000°C, strong peaks from Ca-rich albite also emerge (see Fig. 2). Note that the Ca-rich (ordered) low albites are expected to crystallize at around 1000-1100°C (Deer et al., 1992). Hematite peaks also show up in the XRD spectra of the sample annealed up to 1000°C. This result is consistent with the red-brown color acquired by this sample. While at 800 °C the XRD patterns of augite are closer to those of hedenbergite (Fe-rich end of augite) and no hematite reflections are observed, at 1000°C the peaks assigned to augite are closer to those of diopside (Mg-rich end of augite). This suggests that at such elevated temperatures Fe is fully oxidized, giving rise to the hematite phase, while the rest of species are incorporated into the felsic (plagioclase) feldspar. In contrast, at lower temperatures (800°C) Fe, and also the rest of cations from the original aerinite (Ca, Mg, Na), seem to be mainly incorporated in the hedenbergite phase. This result is compatible with the brownish color displayed by the sample annealed at 800°C.





CONCLUSIONS

The blue pigment aerinite, widely employed in Romanesque paintings in the region of the southern Pyrenees, undergoes important heat-induced color alterations. The color acquired by aerinite samples after thermal annealing is found in-situ in the frescoes of Sixena, which were badly damaged several decades ago. We have used XRD to monitor the mineralogical alterations of the pigment aerinite upon heating. The XRD data suggest that aerinite is totally degraded above 400°C. At 800°C recrystallization yields Fe-rich augite. At 1000°C Fe is mainly found as hematite, whereas the rest of cationic species give rise to (felsic) plagioclase feldspars.

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