Influence of the hydrodynamic conditions on the accessibility of *Aristeus antennatus* and other demersal species to the deep water trawl fishery off the Balearic Islands (western Mediterranean)

Angel Amores^{*1}, Lucía Rueda², Sebastià Monserrat^{1,3}, Beatriz Guijarro², Catalina Pasqual¹, and Enric Massutí²

¹Instituto Mediterráneo de Estudios Avanzados, IMEDEA (UIB-CSIC). Palma de Mallorca, Spain

²Instituto Español de Oceanografía. Centre Oceanogràfic de les Balears. Moll de Ponent s/n, 07015 Palma de Mallorca, Spain
³Departament de Física. Universitat de les Illes Balears (UIB). Palma de Mallorca, Spain

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*aamores@imedea.uib-csic.es

Abstract

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Monthly catches per unit of effort (CPUE) of adult red shrimp (Aristeus antennatus), reported in the deep water bottom trawl fishery developed on the Sóller fishing ground off northern Mallorca (Western Mediterranean), and the mean ocean surface vorticity in the surrounding areas are compared between 2000 and 2010. A good correlation is found between the rises in the surrounding surface vorticity and the drops in the CPUE of the adult red shrimp. This correlation could be explained by assuming that most of the surface vorticity episodes could reach the bottom, increasing the seabed velocities and producing sediment resuspension, which could affect the near bottom water turbidity. A. antennatus would respond to this increased turbidity disappearing from the fishing grounds, probably moving downwards to the deeper waters. This massive displacement of red shrimp specimens away from the fishing grounds would consequently decrease their accessibility to fishing exploitation. Similar although more intense responses, have been observed during the downslope shelf dense water current episodes that occurred in a submarine canyon, northeast of the Iberian peninsula. The proposed mechanism suggesting how the surface vorticity observed can affect the bottom sediments is investigated using a year-long moored near-bottom current meter and a sediment trap moored near the fishing grounds.

The relationship between vorticity and catches is also explored for fish species (Galeus melastomus, Micromesistius poutassou, Phycis blennoides) 22 and other crustacean (Geryon longipes and Nephrops norvegicus), consid-23 ered as by-catch of the deep water fishery in the area. Results appear to 24 support the suggestion that the water turbidity generated by the vorticity 25 episodes is significant enough to affect the dynamics of the demersal species. 26

27 1 Introduction

The decapod crustacean red shrimp, Aristeus antennatus (Risso, 1816), a demersal 28 species distributed throughout the Mediterranean and the north-eastern Atlantic, 29 from Portugal to the Cabo Verde Islands [Arrobas and Ribeiro-Cascalho, 1987], 30 mainly occurs in the muddy bottoms of the slope, between 400 and at least 3300 31 m [Sardà et al., 2004]. This species is one of the most valuable deep-water fishing 32 resources in the western and central basins of the Mediterranean, remaining at a 33 low level of exploitation in the eastern basin [Papaconstantinou and Kapiris, 2001] 34 and revealing important bathymetric migrations [Relini et al., 2000]. However, 35 despite its wide bathymetric distribution, it is mainly exploited between 400 and 36 800 m depth, and is the target species of the well-developed deep water bottom 37 trawl fishery on the western basin slope [Sardà et al., 2003]. 38

The trawl fleet operating off the Balearic Islands (western Mediterranean) is characterized by its versatility, which is determined by the specific dynamics of the resources, among other factors (e.g. sea conditions and fish market). Bottom trawlers not only target different species, they also change the fishing location at a given time of the year, as well as the fishing tactics during the same fishing trip. Palmer et al. [2009] defined four fishing tactics in this fishery, related to the exploitation of different bathymetric strata and target species.

The annual catches of the red shrimp in the Balearic Islands are estimated to be 46 around 100-200 t, which represents 10% of the landings and 40% of the earnings 47 in the trawl fishery [Guijarro et al., 2012]. Sóller, one of the most important 48 fishing grounds for red shrimp around the Balearic Islands, is situated North 49 of Mallorca (solid black line area in Fig. 1), where an important part of the 50 island fleet is concentrated during the summer months [Moranta et al., 2008], 51 when catches of large specimens occur. The red shrimp population in this fishing 52 ground shows important seasonal variations throughout the year (such as the high 53 abundance of juveniles recruiting to the fishing grounds in autumn-winter and the 54 high abundance of large spawning females during the summer), compared with 55 the other nearby fishing grounds, south of Mallorca [Guijarro et al., 2008]. 56

The Sóller fishing ground is located on the island slope, in a well known very active area, with numerous eddies normally generated by some instabilities of the Northern current or the Balearic current (Fig. 1), particularly more intense during winter (October-March; Amores et al. [2013]). These eddies, clearly visible on satellite images, have been known to reach the deeper waters, and their effects are usually felt down to the seabed, where their velocities may increase to several
times those of the mean currents measured in the zone [Amores et al., 2013]. These
strong bottom currents of the order of 25 cm/s are known to produce sediment
resuspension which, in turn, may generate additional cross slope turbidity currents
[Thomson et al., 2010].

In the western Mediterranean, the red shrimp distribution, and its accessibility 67 to fishing exploitation, has been shown to be mainly influenced by geomorphol-68 ogy [Sardà et al., 1994, 1997] and hydrodynamics [Bombace, 1975; Demestre and 69 Martín, 1993; Ghidalia and Bourgois, 1961; Guijarro et al., 2008; Relini and Re-70 lini, 1987; Sardà et al., 2009]. These last factors are probably linked to regional 71 and large-scale climatic patterns [Carbonell et al., 1999; Massutí et al., 2008; May-72 nou, 2008]. In a recent study, Company et al. [2008] revealed that the downslope 73 shelf dense water current events into submarine canyons, along the whole north-74 ern Catalan margin, strongly affected the red shrimp landings. This downslope 75 shelf dense water current events is one of the main processes contributing to the 76 shelf-deep ocean exchange [Ivanov et al., 2004], enhancing organic-matter flux and 77 deposition, increasing suspended particulate matter concentrations and transport 78 of organic matter from coastal zones to the deep ocean [Bosley et al., 2004; Canals 79 et al., 2006; Company et al., 2008]. In the northern Catalan margin, it exerts a 80 negative effect on the catches of red shrimp and a positive effect for recruitment, 81 due to the transportation of the particulate organic matter. The increase of sus-82 pended particulate matter also appears to be related to the abundance of other 83 crustacean species such as pandalids and penaeid [Lin et al., 1992; Puig et al., 84 2001] and to an enhance of benthic productivity and biodiversity inside canyon 85 habitats [Rowe et al., 1982; Schlacher et al., 2007; Vetter et al., 2010]. In addition 86 to downslope shelf dense water current, mesoscale eddies have also been reported 87 to be responsible of transport of shelf sediments to the deep ocean, resuspension 88 of bottom sediments creating turbidity layers and formation of sediment plumes around their periphery [Washburn et al., 1993]. The influence of vorticity (as 90 indicator of eddy development) on catchability of marine species has been mostly 91 addressed for pelagic organisms such as tuna fisheries [Hyder et al., 2009; Kai 92 and Marsac, 2010; Ramos et al., 1996; Zainuddin et al., 2006]. However, the ef-93 fect of such physical processes has also been explored for benthic species, which 94 are also linked to variables that describe water column properties and structures 95 [Beentjesa and Renwick, 2001; Palamara et al., 2012]. 96

The objective of this work is to analyze the possible links between the presence 97 of eddies (which will be quantified by their associated surface vorticity) affecting 98 the Sóller fishing ground and the red shrimp yields of the deep water trawl fishery 99 developed in the area. This relationship is also explored for other demensal species 100 frequently caught by the deep water bottom trawl fishery developed in the area 101 [Guijarro and Massutí, 2006], which consist of three fishes (Galeus melastomus, 102 Micromesistius poutassou and Phycis blennoides) and two decapod crustaceans 103 (Geryon longipes and Nephrops norvegicus), with the objective of discussing their 104 different responses in relation to their living habits. 105

A year-long near-bottom current meter and a sediment trap moored near the fishing grounds are used to infer the mechanism to explain how the surface vorticity observed can affect the bottom sediments and, in turn, the red shrimp yields.

¹⁰⁹ 2 Data and Methods

110 2.1 Catches

Daily time series of the landings from the bottom trawl fleet have been obtained 111 from the official sale bills of OP Mallorca Mar, the fishery producer organization 112 of Mallorca, between 2000 and 2010 (both years included). Each daily sale bill 11: was assigned to one fishing tactic (FT) or a combination of them following the 114 methodology described by Palmer et al. [2009]. Landings were standardized to 111 CPUEs (catches per unit of effort), referred to as kilograms caught per day and 11 (boat. For A. antennatus, only catches obtained from the middle slope fishing 117 tactic, developed between 600 and 800 m depth, have been considered, because 118 this is the target species for this FT. Moreover, the daily sale bills distinguished 11 red shrimp catches into two size categories (small and large) up to year 2004, 120 and three categories (small, medium and large) from 2004 to the present day. 121 According to Guijarro et al. [2008], two different categories were defined in order 122 to homogenize the available data, small (including individuals with a carapace 123 length <32mm) and medium-large (adults, with a carapace length ≥32 mm). For 124 this analysis, only those of the medium-large sized category, mainly adult females, 125 were considered. Juveniles are not taken into account for two reasons: 126

The fishing fleet mainly targets large individuals (adults) due to their higher
 commercial value. This fact would surely provoke a bias when trying to
 relate juvenile catches with abundances

Adult and juvenile red shrimps present a clear different bathymetric distribution. Adult individuals are mainly located at the 500-800 m range, where the fishing fleet is developed. But the highest concentrations of juveniles are situated deeper than 1000 m Sardà et al. [2003], where the bottom trawl fishery is forbidden. So juvenile catches do not properly reflect the juvenile population abundances.

From the entire fleet that currently operates in Mallorca, only five boats regularly fish in the zone of interest (Sóller) throughout the year (other boats fishing in this area only in summer are not considered). Among these five boats, exclusively two devote most of their efforts to red shrimp fishery along the middle slope and they were the only ones finally considered for the analysis. Finally, as a direct response to hydrodynamics changes in a daily basis is not expected, a monthly average was calculated according to the daily time series. We intentionally filtered out the high frequency variations in order to compute an integrated response in a longer time scale and take into account for the gaps in the data (that occurred, for example, on weekends or bad weather days).

Regarding the other species considered in this work (*G. melastomus, M. poutas*sou, *P. blennoides, G. longipes* and *N. norvegicus*), all the boats and slope fishing tactics have been considered because, unlike *A. antennatus*, they are the 'bycatch' species of the deep water trawl fishery, and therefore, their abundance in daily landings is not as frequent as that of the red shrimp. The final time series for each species were also averaged monthly as the CPUEs in terms of kg of catch per day per boat.

153 2.2 Hydrodynamic Data

154 2.2.1 Satellite images

We estimated the relative vorticity ζ (from now on referred to as only vorticity) from the daily Sea Surface Height (SSH) satellite images with a map spacing of $1/8^{\circ} \times 1/8^{\circ}$, obtained from the merged satellite AVISO products available at http://www.aviso.oceanobs.com. The vorticity is calculated as the curl of the velocity field, but we only retain the third component as it represents the vorticity of a horizontal field. By considering the hydrostatic and homogeneous fluid, the final expression of ζ is:

$$\zeta = \frac{g}{f} \cdot \nabla^2 SSH \tag{1}$$

where g is the gravity acceleration, f the Coriolis parameter and ∇^2 the horizontal Laplacian.

After computing the daily vorticity fields, their absolute value was taken be-164 cause both, cyclonic ($\zeta > 0$) and anticyclonic episodes ($\zeta < 0$), were expected to 165 have the same effect on the seabed velocities and sediment resuspension. Next, 160 we computed the spatial average in the dashed rectangle, as shown in Fig. 1. The 167 choice of the area is somewhat arbitrary. The size should be significantly greater 168 than the fishing ground dimension because the eddy sizes are significantly greater, 169 and their horizontal influence is not known. The best results are found when the 170 area is selected to be large enough to include the Northern and the Balearic cur-171

rents, potentially the eddy generators. Finally, the daily time series was averaged
on a monthly basis in order to have the same time step as the time series of the
catches.

175 2.2.2 Moorings

A mooring was deployed north-west of Mallorca $(39^{\circ}49.682' \text{ N} - 02^{\circ}12.778' \text{ E})$ 176 star in Fig. 1) between November 2009 and February 2011. Located around 900 17 m depth in the Mallorca slope, it had four CTD Seabird 37 (300, 500, 700 and 17 900 m) and two current meters Nortek Aquadopp (500 and 900 m). Moreover, 179 the mooring had also a sediment trap placed 30 m above the bottom. The CTD 180 sampling rate was 10 minutes, while the current meters recorded one value every 181 30 minutes. The sediment trap had a sampling interval of 10 days and it had 12 18: bottles. The combination of sampling rate and number of bottles made necessary 18 a maintenance of the mooring every four months. 184

All the instruments operated perfectly during the entire period, except for the 500 and 900 m CTDs, which ran out of batteries in mid-December 2010 and mid-January 2011, respectively. The sediment trap worked well too, but the unavailability of boat lead to no recorded data between July 5 and September 21, 2010.

Sediment trap samples were wet-sieved through a 1 mm nylon mesh in order to retain the largest organisms. Swimmers smaller than 1 mm were manually removed under a dissecting microscope using fine tweezers. Finally, the sample was freeze-dried and weighed to calculate the Total Mass Flux (TMF).

Despite we did not have any direct measurement of turbidity, the Nortek Aquadopp current meters give us an estimation of it throughout the backscattering of the particles used to compute the velocity (as suggested byLohrmann [2001]). Acoustic backscattering has been used as turbidity surrogate in different references such as Thomson et al. [2010].

2.3 Statistical Analysis

Quantification of the similarity between surface vorticity and time series of catches was performed with the correlation function. If we define V and C as the monthly anomalies (time series after subtracting the mean value) for vorticity and catches, respectively, the correlation between these two series is calculated as:

$$\rho_{VC}(lag) = \frac{1}{(N-1-|lag|)\cdot\sqrt{\sigma_{VV}\cdot\sigma_{CC}}} \begin{cases} \sum_{i=|lag|+1}^{N} V_i \cdot C_{i-|lag|} & \text{if } lag < 0\\ \sum_{i=|lag|+1}^{N} V_{i-|lag|} \cdot C_i & \text{if } lag \ge 0 \end{cases}$$
(2)

with N being the series length, σ_{VV} the covariance of V and σ_{CC} the covariance of C.

²⁰⁶ The significance level may be obtained with:

$$sig(lag) = \frac{T_q(0.99, N - |lag| - 2)}{\sqrt{N - |lag| - 2 + [T_q(0.99, N - |lag| - 2)]^2}}$$
(3)

where $T_q(0.99, D)$ is the t-student distribution with a significance of 99% and D degrees of freedom.

²⁰⁹ 3 Results and Discussion

A visual inspection of the monthly average vorticity and *A. antennatus* CPUE time series appears to suggest that any increase in vorticity generally causes a decrease in the CPUEs, although a decrease in vorticity does not cause an increase in CPUEs (Fig. 2a). In fact, a negative correlation (i.e. both series are in antiphase) between time series has been found (Fig. 2b).

In the following we will give several evidences supporting the fact that surface 21 ! vorticity affects red shrimp availability by modifying near bottom turbidity in the 21 6 fishing grounds. If we assume that any increase in vorticity affects the CPUE of the 21 red shrimp by producing a near bottom turbidity, which in turn would decrease 21 the resource availability, we could not expect that the opposite, a decrease in 21 9 vorticity, would immediately produce an increase in the CPUEs. Water turbidity 220 would persist for some period of time in the area after the end of any vorticity 221 episode and then the red shrimp response would be delayed. Therefore, the two 222 series have been modified, trying to take into account the different, although 22: expected, CPUE response to the increase and decrease in vorticity. Vorticity and 224 CPUE time derivatives have been computed. As we want to highlight the part 225 when the vorticity increases (positive derivative) and when de CPUE decreases 220 (negative derivative), the negative vorticity and positive CPUE derivatives were 227 artificially set at zero. Therefore, only the increases in vorticity and decreases in 228 CPUE are considered. Furthermore, the sign of the CPUE derivatives has also 229 been reversed for better visualization (Fig. 3a). 230

Vorticity and A. antennatus CPUE's derivative series show quite a similar 231 pattern. The zones where the series have been forced to read zero coincide, and 232 almost any increase in the vorticity derivative corresponds with an increase in the 23 reversed catches derivative. The correlation between both series at lag zero, now 234 positive due to the sign change in the CPUE's derivatives mentioned, is slightly 23! larger than before, as expected, reaching a value of 0.48 (Fig. 3b). This result 236 strongly supports a relationship between the increases in the surrounding absolute 237 surface vorticity and decreases in the adult A. antennatus availability in the fishing 23 grounds. 239

The suggested mechanism explaining the relationship observed is now supported analyzing some surface vorticity episodes recorded when the mooring was deployed in the area. During 2010, at least three of these episodes produced some footprint in the instruments deployed in the mooring line.

The increase in the absolute value of the surface vorticity is commonly caused 244 by the presence of an eddy, such as the one shown in Fig. 4. This particular eddy 245 remained in the area between mid-November and mid-December 2010 and was 24 0 studied in detail by Amores et al. [2013]. This eddy was clearly reflected in the 24 currents registered at 500 and 900 m depth. A significant velocity increase was 24 8 measured at both depths (episode 3 in Fig. 5a and 5b). Velocities showed spikes 24 9 reaching up to 26 cm/s at 900 m, where the mean current during the whole year 250 was computed to be around 5 cm/s. This eddy also affected the current direction, 251 causing a complete reversal in the currents at 500 m (Fig. 5c) and a down slope 252 gvre at 900 m (Fig. 5d). 25:

This particular eddy clearly reached down to the bottom, and the recorded gyres and velocity increases could easily have caused the material resuspension. This hypothesis is supported by three indirect measurements:

the increase in the total flux mass (TFM) recorded by the moored sediment
 trap at the time of the eddy (Fig. 6)

259 2. the increase of the acoustic backscattering during the episode (Fig. 7a)

3. the clear down slope gyre at 900 m which could be related to a near bottom
turbidity current (Fig 5d).

The eddy shown in Fig. 4 was not the only one recorded when the mooring was deployed. Another eddy, which occurred between mid-January and March 2010, was also measured by the mooring (episode 1 in Fig. 5). This eddy too reached the bottom, although with a weaker footprint in velocity. However, the TFM still reached similar values to those observed during the December 2010 eddy and an increase in backscattering is also observed.

Still another gyre was observed between June and July 2010; however, it was 268 only noticeable at 500 m (episode 2 in Fig. 5). Its effect at 900 m was weak. 269 Even so, a TFM peak was also measured by the sediment trap, although much 270 weaker than during the other two eddies (Fig. 6). Backscattering did not show 271 any significant increase during this episode. This could be explained by the steep 272 slope of the area. Even if the eddy was not energetic enough to reach down to 273 900 m depth, where the mooring was deployed, it could still affect the bottom 274 at shallower depths and the resuspension of material could have reached deeper 27! waters, causing the increased TFM that is recorded in the sediment trap. 276

From the above described episodes, it has been observed that the surface eddies 277 exert some degree of influence on the seabed dynamics and that might increase 278 the water turbidity near the bottom and affect the availability of A. antennatus 279 in the fishing grounds. This mechanism can be better visualized by restricting the 280 time interval to a shorter period of time of some of the data shown in Fig. 2, for 281 the period, 2006-2010 (Fig. 8). During these years, the vorticity episodes are time 282 spaced enough to allow an almost complete recovery towards a normal situation 283 after any episode, before the arrival of the next one. The mechanism suggested 284 are then better observed in the data. A vorticity increase (dark gray bands in 285 Fig. 8) would trigger an increase in the re-suspended material and would force 286 A. antennatus to move away from the fishing ground, probably towards greater 28 depths, leading to a decrease in the catches of this species. This scenario would 288 remain unchanged until the eddy effects disappear and the sediments once again 289 precipitate to the sea floor (soft gray bands in Fig. 8). Once all the sediments 290 are completely settled down, the water conditions would become suitable to allow 291 the individuals to return to the depths where they can be caught (white bands in 293 Fig. 8). Four episodes appear to follow one after another in the period shown. 293

Company et al. [2008] described a similar but stronger phenomenon in the 294 submarine canyon system off the north-eastern Iberian Peninsula. They found a 29! correlation between the strong currents associated with intense downslope shelf 290 dense water current events and the disappearance of A. antennatus from its fish-297 ing grounds, exerting a negative effect on the catches reporting and a temporary 29 collapse of its fishery. An increase in the mortality rated after exposure to high 299 turbidity has also been detected for Penaeid shrimps at juvenile and adult stages 300 [Lin et al., 1992]. Both downslope shelf dense water current events and mesoscale 301 eddies have been reported to enhance organic-matter flux and deposition, increas-302 ing suspending particulate matter concentrations with the transport of organic 303 matter from coastal zones to the deep ocean or by resuspension of bottom sed-304 iments [Bosley et al., 2004; Canals et al., 2006; Washburn et al., 1993]. Life 305 forms as diverse as phytoplankton, protozoans, crustaceans, fish, sea snakes, ma-306 rine mammals and birds are found to alter their distributions in the presence of 307 such flow patterns [Owen, 1981], which can be responsible for enhancing benthic 308 productivity and biodiversity inside canyon habitats [Rowe et al., 1982; Schlacher 309 et al., 2007; Vetter et al., 2010]. The presence of a significant amount of suspended 31 sediment has also been related to the higher occurrence of juveniles and females 311

of the deep-water pandalid shrimp species, genus Plesionika [Puig et al., 2001] 31 2 and the regions where the intermediate nepheloid layers detach from the seabed 31 3 have been defined as potential deep-water nursery habitats for these species. The 314 transportation of the particulate organic matter associated with the downslope 31 ! shelf dense water current appears to be positive for recruitment of red shrimp 31 (in the north-eastern Ibearian Peninsula [Company et al., 2008], with a positive 317 increase in the landings 3-5 years after these events, preceded by an increase in 318 the number of juveniles. However, this last effect has not been detected off the 31 9 Balearic Islands, probably because of the slight difference in the feeding strate-320 gies of the species between the north-eastern Iberian Peninsula and the Balearic 321 Islands. A. antennatus has a highly varied diet, being among the mega-benthic 323 species mainly preving on the benthos in the deep Mediterranean [Cartes, 1994; 323 Cartes and Carrassón, 2004]. However, benthic preys are particularly significant 324 off the north-eastern Iberian Peninsula, where the submarine canyons enhance 325 such types of food availability. Conversely, the trophic webs off the Balearic Is-326 lands show an impoverishment of the benthos biomass and depend more directly 321 on food of planktonic origin, enhancing the consumption of micro-nektonic preys 328 [Cartes et al., 2008; Maynou and Cartes, 2000]. In this sense, the positive effects 329 of downslope shelf dense water current and sediment resuspension in the long 330 term should be also more marked off north-eastern Iberian Peninsula than off the 331 Balearic Islands. 332

Although the mechanism suggested is somehow speculative because we have to 333 rely on indirect data (backscattering) to deduce bottom turbidity, we have shown 334 several evidences that the presence of enough energetic eddies may cause bottom 335 turbidity increases in the fishing ground. Still another indirect evidence supporting 336 the mechanism suggested comes from the analysis of other demersal species caught 337 in the same region that also appear to be related to vorticity changes in the area. 338 The correlations found for each species (Fig. 9) are different from those observed 339 for A. antennatus but consistent with the suggested mechanism where vorticity 34 0 affects the seabed by increasing the bottom water turbidity. 341

Other decapod crustaceans, Geryon longipes and Nephrops norvegicus, the more sedentary and benthic species, show a significant positive correlation to the vorticity events at around 0.4 and 0.5, respectively. These two species are closely connected to the bottom, as reflected by their feeding behavior and biological characteristics. G. longipes preys on a broad range of benthic invertebrates [Cartes, 1993] and N. norvegicus shows a scavenging activity [Cristo and Cartes, 1998]
and has been related to the sediment characteristics [Maynou and Sardà, 1997].
Unlike A. antennatus, it is likely that these two species may take some advantage
of the re-suspended matter.

Galeus melastomus, which has more mobility than the previously considered
epi-benthic species, feeds on the mesopelagic preys with the occasional occurrence
of benthic feeding activity and scavenging in the adult phase [Fanelli et al., 2009].
This species showed a lower correlation (0.33) with the vorticity time series.

Finally, *Micromesistius poutassou* and *Phycis blennoides*, the two benthopelagic teleosts with greater capacity of movement above the bottom, are expected to be less affected by bottom water turbidity. In fact, no significant correlations between the CPUEs and vorticity have been found.

359 4 Summary and Conclusions

A reasonable good negative correlation is noted between the monthly CPUE of 360 the adult A. antennatus bottom trawl yields in the fishing grounds off northern 361 Mallorca and the mean surface vorticity in the surrounding area. We have shown 363 that the eddies causing the vorticity events may reach the bottom, increasing 363 the current velocities, which in turn would trigger sediment resuspension and 364 increased bottom water turbidity. Such a change in the water conditions would 365 force adult A. antennatus individuals to move away from the fishing ground, 366 probably downwards, to greater depths. This proposed mechanism is similar, 36 although lower in magnitude, to the one suggested by Company et al. [2008] in 36 the downslope shelf dense water current events of the submarine canyons in the 369 northern Catalan margin, off the north-eastern Iberian Peninsula. In the Balearic 370 Islands, where these geomorphological structures do not exist and where there is 371 no river runoff, the eddies would be the triggering factor. 372

Other deep water demersal species found along with the catch of the red shrimp fishery, possessing different behavior and feeding habits, exhibit different responses to these events, but all of them are consistent with the eddy generation near bottom velocities increase and bottom water turbidity.

A final hypothesis could also be suggested from the results obtained. The 377 seasonal migration of most of the fishing fleet of Mallorca, targeting the red shrimp 378 in the Sóller fishing grounds during the summer has been explained by the highest 379 abundance of large spawning females in this area during this season [Guijarro 380 et al., 2008], similar to other areas off the north-eastern Iberian Peninsula [Sardà 381 et al., 1994, 1997]. In light of this, the absence of these large aggregations in 382 the Sóller fishing grounds during the rest of the year could be related to the 38 particular behavior of the species. However, it is worth noting that, according to 384 Amores et al. [2013], the vorticity episodes are much more intense off northern 38! Mallorca during the winter time (October to March) than in the summer. This 38 fact could be an additional factor explaining the decrease in the availability of 387 the red shrimp to fishing exploitation during these months. However, off southern 38 Mallorca, yields from the bottom trawl fishery targeting to red shrimp remain 389 more stable [Guijarro et al., 2008]. 390

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Figure Captions

563

Fig. 1. Map of the studied area in the western Mediterranean. The unbroken line encloses Sóller fishing grounds where *Aristeus antennatus* is exploited and the broken line corresponds to the zone where the time series of vorticity has been calculated. Mooring location is indicated by a star.

568

Fig. 2. a) Monthly averaged time series of vorticity (black) and Aristeus antennatus CPUE's (green). b) Correlation between these two series, showing the maximum correlation (-0.35) around lag 0. Black line represents the 95% confidence level.

573

Fig. 3. a) Derivative time series of vorticity (black) and Aristeus antennatus CPUE's (green). In the series of the vorticity derivative, the negative values have been fixed at 0, while the positive values of the derivative CPUE's series have been set at 0. Notice that the last one has suffered a change of sign. b) Correlation between these two series show the maximum value (0.48) at lag 0. Black line represents the 95% confidence level.

580

Fig. 4. Sea Surface Height (SSH) image from December 1, 2010. It shows an eddy
in the region analyzed. The star shows the mooring position.

583

Fig. 5. 24h low-pass filtered speed series of 500 (a) and 900 (b) m depth current 584 meters of the mooring for the whole recorded period. Blue indicates the low speed 58 values degrading to red, which indicates the high values. (c) and (d) are the 58 progressive vector diagrams for 500 and 900 m depth, respectively. The different 587 colors coincide temporally with the speed time series. Enclosed areas represent 584 moments where an eddy is present in the zone. Ellipse number 1 highlight an eddy 589 which is strongly present at 500 m depth, but weakly present at 900 m; number 590 2 ellipse shows an eddy which reached to 500 m depth, although not right up to 591 900 m; and number 3 ellipse illustrates an eddy which arrived strongly to the 500 592 and 900 m depths. Note that in the PVDs the ratio between the scales of x and 593 y axis is 2:1. 594

595

Fig. 6. Total Flux Mass (TFM) collected by the sediment trap during the whole

sampling time. The gap in the data is due to the unavailability of ship for carrying
out the mooring maintenance. The dashed ellipses show the increment of TFM
due to the eddies reported in Fig. 5.

600

Fig. 7. Acoustic backscattering (a) and speed (b) measured by the 900 m current meter during the third episode.

603

Fig. 8. Zoom of the Fig. 2, where the effect of the vorticity (blue) on the Aristeus antennatus CPUE's (green) can be seen. The colored bands indicate the amount of particles that would be re-suspended.

607

⁶⁰⁸ Fig. 9. Time series of CPUE's for five demersal species (by catch) from the deep

water trawl fishery and its correlation with the absolute value of surface vorticity.

•10 Figures

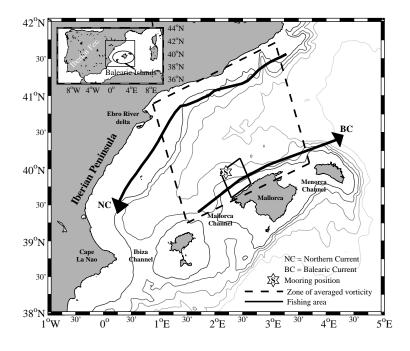


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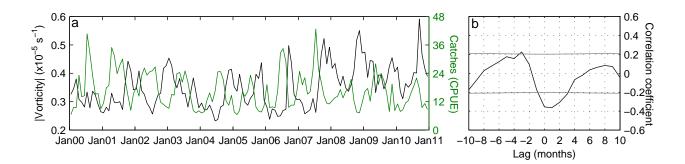


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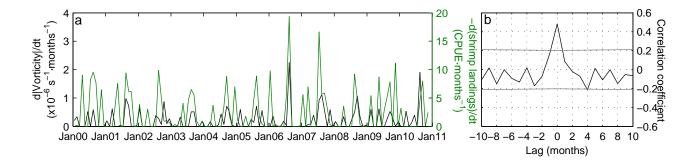


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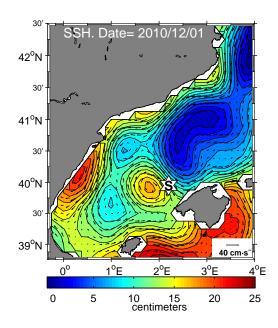


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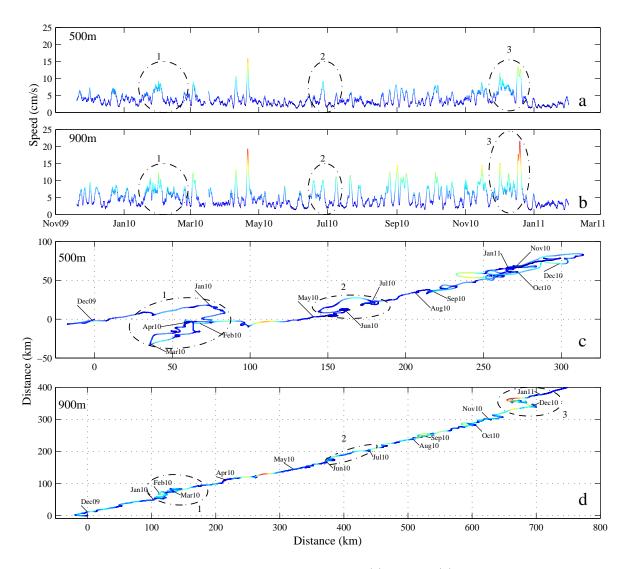


Figure 5: 24h low-pass filtered speed series of 500 (a) and 900 (b) m depth current meters of the mooring for the whole recorded period. Blue indicates the low speed values degrading to red, which indicates the high values. (c) and (d) are the progressive vector diagrams for 500 and 900 m depth, respectively. The different colors coincide temporally with the speed time series. Enclosed areas represent moments where an eddy is present in the zone. Ellipse number 1 highlight an eddy which is strongly present at 500 m depth, but weakly present at 900 m; number 2 ellipse shows an eddy which reached to 500 m depth, although not right up to 900 m; and number 3 ellipse illustrates an eddy which arrived strongly to the 500 and 900 m depths. Note that in the PVDs the ratio between the scales of x and y axis is 2:1.

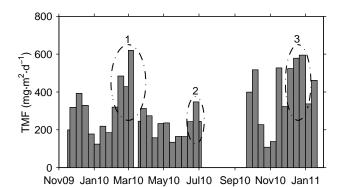


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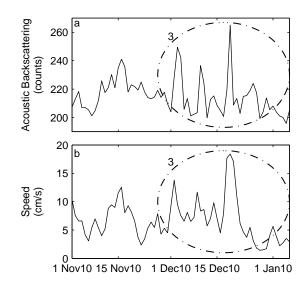


Figure 7: Acoustic backscattering (a) and speed (b) measured by the 900 m current meter during the third episode.

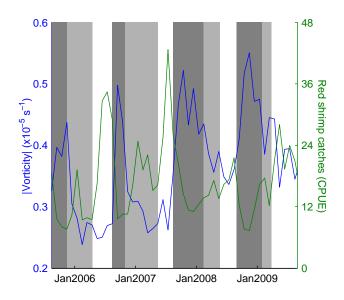


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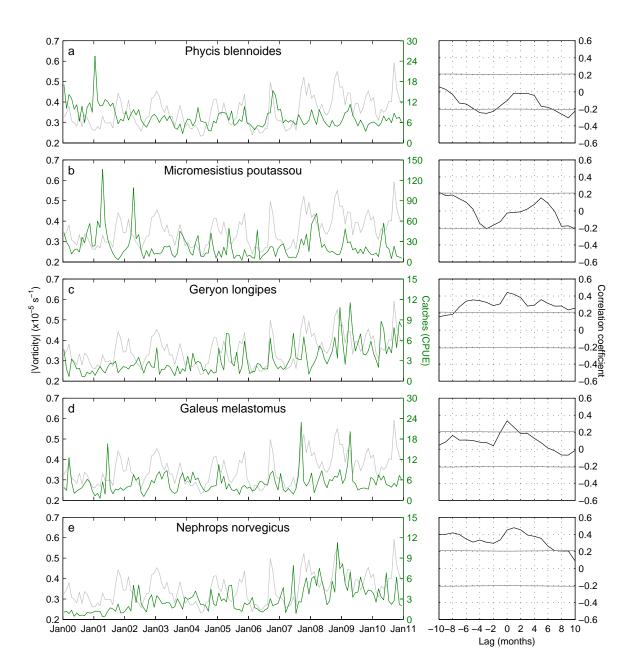


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