

A BRIEF SUMMARY OF ORDOVICIAN CONODONT FAUNAS FROM THE IBERIAN PENINSULA

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INTRODUCTION

Ordovician conodont studies in the Iberian Peninsula were initiated by Fuganti and Serpagli (1968), who recognized 21 morphospecies included in 15 morphogenera in the Upper Ordovician Urbana Limestone from a single locality in the Central Iberian Zone. Two years later Boersma (in Hartevelt, 1970) identified several morphotaxa in the Upper Ordovician Estana Formation of the Central Pyrenees. In the type section of the Upper Ordovician Cystoid Limestone of the Eastern Iberian Cordillera, Carls (1975) recognised 31 conodont morphotaxa. These pioneer findings were followed by the contributions of Kolb (1978), Hafenrichter (1979), Robert (1980), Robardet (1982) and Sanz (1988), who increased the number of taxa and localities with Katian conodonts, mostly attributed to the *Amorphognathus ordovicicus* Zone.

For twenty years, our knowledge on Ordovician conodonts came only from the single ubiquitous limestone unit that occurs in the upper part of many Iberian successions. Nonetheless, these are predominantly composed of terrigenous rocks (shales, siltstones and sandstones) which were deposited at high Gondwanan paleolatitudes near the South Pole (Gutiérrez-Marco et al., 2002, 2004). Then, some of these clastic deposits (siltstones, shales and storm-induced coquinoïd lenses, sometimes calcareous) were also sampled for conodonts: while siltstones and shales produced only fragmentary specimens, bioclastic beds in tempestites yielded usually fragmentary, but recognisable, elements.

EARLY TO EARLY-MID ORDOVICIAN CONODONTS

The oldest Ordovician conodont assemblage from the Iberian Peninsula was found by Sarmiento and Gutiérrez-Marco (1999) near Adamuz (Fig. 1, loc. 17) in the tectonically-complex boundary area between

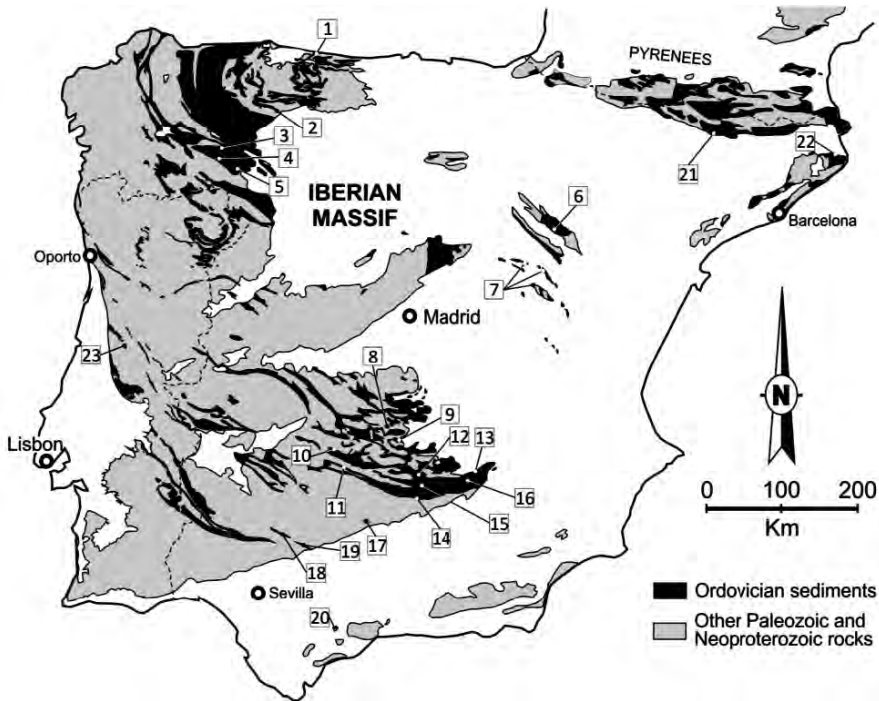


Figure 1. Map of the main Ordovician outcrops in the Iberian Massif, with locations of mentioned conodont localities in Spain and Portugal. 1, Sueve and Ordovician tunnel; 2, Portilla de Luna; 3, La Aquiana; 4, Casaio; 5, Rozadais-Truchas; 6, Fombuena; 7, Aragoncilo, Ojos Negros and Checa areas; 8, Porzuna; 9, Corral de Calatrava; 10, north of Chillón; 11, Alamillo; 12, Calzada de Calatrava; 13, Villamanrique-Terrinches; 14, El Centenillo; 15, Huertezuelas and Viso del Marqués; 16, Aldeaquemada; 17, Adamuz; 18, Cazalla de la Sierra; 19, Constantina; 20, Ardales; 21, Els Castells; 22, Les Gavarres; 22, Buçaco.

the Central Iberian and Ossa-Morena zones. Conodonts were obtained from a few decalcified (iron-rich) limestone cobbles embedded as olistholits in a Mississippian olisthostrome unit from the Guadalmellato domain, Los Pedroches Basin (Cózar et al., 2004; Armendáriz Dufur, 2009 and references therein). Besides some indeterminate brachiopod and trilobite fragments, the transported limestone contains a diverse and partially reworked conodont assemblage that includes a mixture of Lower and Middle Ordovician taxa. Sarmiento and Gutiérrez-Marco (1999) emphasized the importance of the reworked conodonts to infer the existence of unknown units with Early Ordovician limestones in an area exclusively characterized by siliciclastic deposits at that epoch. The existence of the so-called “phantom formations” (Branson and Mehl, 1940) can be assumed to have occurred in some original areas located westwards, that were tectonically juxtaposed to the southern border of the Central Iberian Zone during the Variscan collision. These would have been partially eroded and resedimented in a typical syn-orogenic foredeep basin during the Mississippian, as can be reconstructed by the presence of a continuous record of Early Ordovician to Devonian fossiliferous olistoliths.

Most of the recovered conodont elements from the older limestone cobbles are poorly preserved, preventing an accurate taxonomic identification. Preliminary studies reveal the presence of *Cordylodus* sp., *Paltodus* cf. *deltifer* (Lindström), *Paltodus* cf. *subaequalis* Pander, *Drepanodus* cf. *arcuatus* Pander, *Drepanodus* spp. *Scolopodus striatus* Pander, *Teridontus* spp., *Acanthodus*? sp., *Acodus* spp.,

Hammannodus sp., *Protopanderodus*? sp., *Gothodus* cf. *costulatus* (Lindström), *Baltoniodus* cf. *triangularis* (Lindström), and *Baltoniodus* sp., beside other coniform elements yet to be identified (Pl. 1, fig. 1-5). The Lower Tremadocian *Paltodus deltiifer* Zone can be inferred by the presence of *Paltodus* cf. *deltifer* (Lindström) and *Hammannodus* sp., but the record of *Cordylodus*? sp. with robust elements of *Acanthodus*? sp. and *Teridontus*? sp. does not exclude the existence of an even older reworked conodont fauna in the assemblage. We propose here a correlation of these “phantom limestone formations” with the Upper Cambrian to Lower Ordovician sequence of the southern Montagne Noire (France), which occasionally incorporates thin limestone beds formed in temperate waters of the Gondwanan margin (Álvaro et al., 2003). This may be confirmed by the Spanish record of *Hammannodus*, a genus so far only known from the Saint Chinian Formation of the southern Montagne Noire (Serpagli et al., 2007).

The younger elements recorded in the mixed Early Ordovician conodont sample from Adamuz may compose an assemblage of elements doubtfully related with *Baltoniodus triangularis* (Lindström), *Baltoniodus* sp., *Drepanodus arcuatus* Pander, *Scolopodus striatus* Pander and possible representatives of the genera *Drepanoistodus* and *Protopanderodus*, among other taxa. This association can be tentatively assigned to the *Baltoniodus triangularis* Zone, broadly representative of Lower Dapingian strata (lowermost Middle Ordovician), which must be simultaneous with the sedimentation of the local limestone also bearing the previously mentioned reworked conodonts.

The only Floian conodonts known from the Iberian Peninsula occur in a lingulid shell-bed located near the top of the Barrios Formation, a local equivalent of the Armorican Quartzite in the Cantabrian Zone of northern Spain (Fig. 1, loc. 1). The assemblage is currently being studied, but preliminary data indicate the presence of coniform elements of the genera *Drepanodus*, *Drepanoistodus* and *Protopanderodus* (Gutiérrez-Marco and Bernárdez, 2003).

MID TO EARLY-LATE ORDOVICIAN CONODONTS

Conodont assemblages representative of the late Darrivilian–early Sandbian interval are assigned to the Dobrotivian stage in Mediterranean regional chronostratigraphy (Gutiérrez-Marco et al., 2008; Bergström et al., 2009), and were reported by Sarmiento et al. (1995a). These occur in thin lenses of calcareous coquinas intercalated in sandy tempestites, or interbedded with quartzite and micaceous shale alternations. All of them belong to a thick siliciclastic group dominated by dark shales and sandy tempestites, broadly known in central Spain as the “Tristani Beds” (San José et al., 1992, with previous references). These have similar counterparts in the Iberian Cordillera and also in the French Armorican Massif, where related conodonts were described by Lindström et al. (1974: age reviewed by Lindström, 1976).

A stratigraphically older conodont assemblage (early Dobrotivian) was recognized in the Central Iberian Zone, as occurring near the top of the El Caño Formation (loc. CC-II), southwest of Calzada de Calatrava (Ciudad Real) (Fig. 1, no. 9). Some other stratigraphically younger conodonts (late Dobrotivian) come from the lower part of the Botella Quartzite in a nearby outcrop (loc. VM-X), and from the La Cierva Quartzite (loc. POR-V) near Porzuna (Ciudad Real) (Fig. 1, no. 8). Similar Dobrotivian conodonts from the Iberian Cordillera have been identified near Fombuena (Zaragoza) in a calcareous intercalation occurring in the middle part of the Sierra Member of the Castillejo Formation (FB-II: Fig. 1, loc. 6).

Most of the conodont taxa recorded from these localities were identified in open nomenclature due the scarcity of elements and their poor preservation. The presence of *Amorphognathus* aff. *inaequalis*

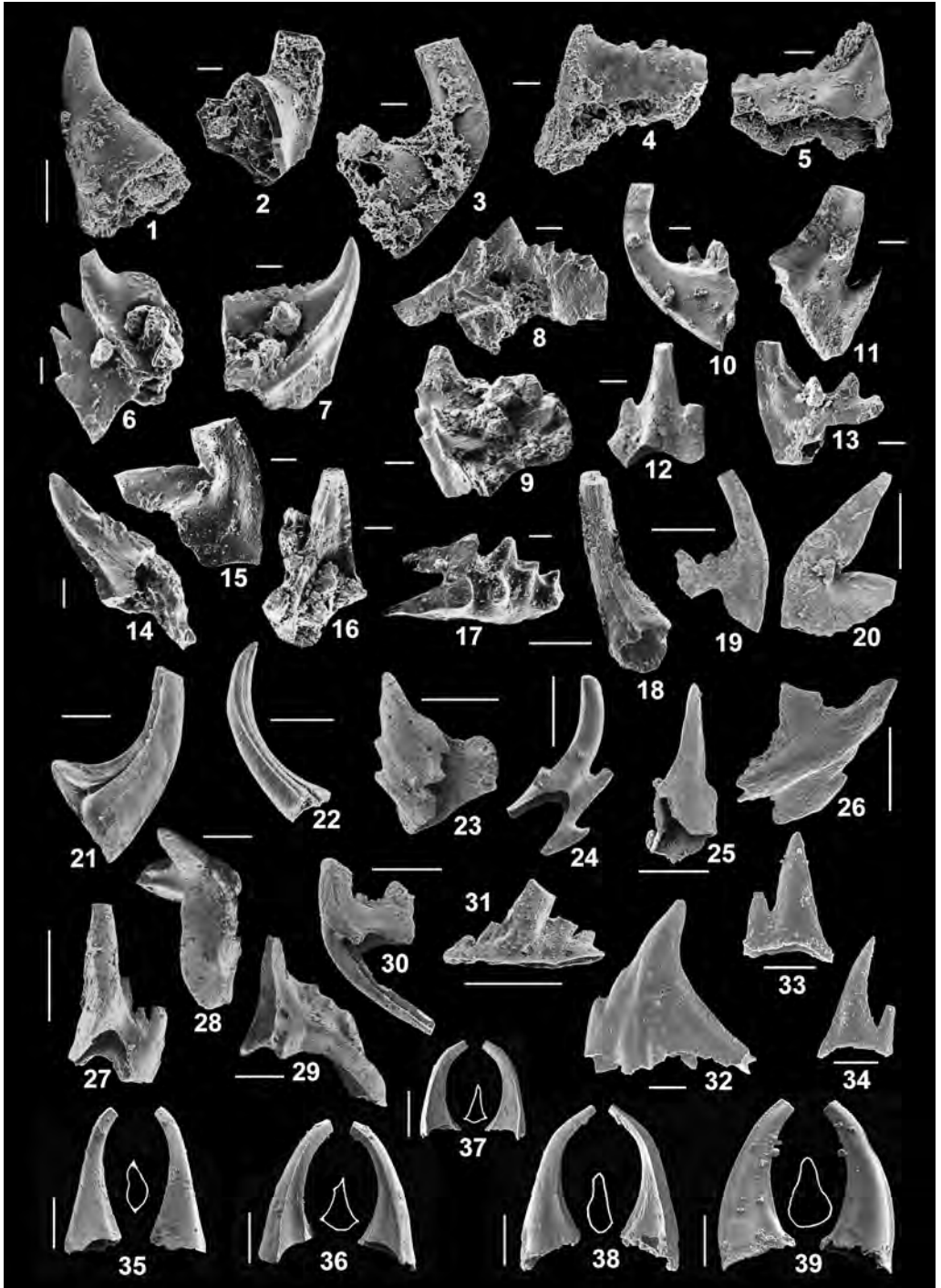
Rhodes, *Amorphognathus* sp., *Baltoniodus* aff. *variabilis* (Bergström), *Coelocerotodus* sp., *Complexodus* sp., *Drepanoistodus* sp., *Icriodella* cf. *praecox* Lindström, Racheboeuf and Henry, *Plectodina* cf. *flexa* Rhodes, among other indeterminate specimens (Pl. 1, figs. 6-16), suggests a biostratigraphical interval below or within the basal part of the *Amorphognathus tvaerensis* Zone. The existence of some reworked conodonts from older conodont biozones occurring in the assemblages from both the Central Iberian and the Eastern Iberian Cordillera localities cannot be ruled out. However, the exact dating of these sequences is firmly established by a combination of graptolite and chitinozoan data, as well by the local trilobite and brachiopod biozones (see Gutiérrez-Marco et al., 2002).

Bedding plane assemblages of lower Dobrotivian conodonts occur in shales from the Sueve Formation of the Cantabrian Zone (Fig. 1, loc. 1). A preliminary identification of these (Gutiérrez-Marco and Bernárdez, 2003) reveals several forms of *Drepanoistodus*, *Panderodus* and *Semiacontiodus*, that are being presently studied.

LATE ORDOVICIAN CONODONTS

The presence of mid Berounian (= late Sandbian to early Katian) conodonts in the Central Iberian Zone was first mentioned by Sarmiento (1993) as coming from an ironstone level at the base of the Cantera Shales near the El Centenillo, Jaén (Fig. 1, no. 14). Only *Amorphognathus* sp., *Panderodus* sp. and *Icriodella* sp. (Pl. 1, figs. 17-18) were identified from this level. This finding allows a chronostratigraphical correlation with equivalent levels of the Piedra del Tormo Member of the Fombuena Formation, Eastern Iberian Cordillera, where Kolb (1978) previously identified and illustrated one fragment of *Icriodella* sp. The mid Berounian age of these conodonts was provided by their association with index species of brachiopod and trilobite biozones.

Plate 1. Some Ordovician conodonts from Spain. 1-3, Early Ordovician reworked specimens from Adamuz; 3-4, Early Mid Ordovician autochthonous assemblage from Adamuz; 6-16, Dobrotivian (late Darriwilian-early Sandbian) beds from Central Iberian localities; 17-18, mid Katian specimens (?*Amorphognathus superbus* Zone) from the Central Iberian Zone; 19-34, mid-late Katian conodonts (*Amorphognathus ordovicicus* Zone) from Central Iberian localities; 35-39, Hirnantian assemblage (upper *A. ordovicicus* Zone) from the Malaguide Complex.— 1, *Teridontus*? sp. ADZ-OI-9875; 2, *Drepanodus*? sp. ADZ-OI-9359; 3, *Cordylodus*? sp. ADZ-OI-9859; 4-5, *Baltoniodus* cf. *triangularis* (Lindström) [4, Pb element ADZ-OI-9374; 5, Pb element ADZ-OI-9375]; 6, *Complexodus*? sp. Pb element POR-V-134A; 7, *Icriodella* aff. *praecox* Lindström, Racheboeuf and Henry. S element POR-V-0834; 8-9, *Amorphognathus* aff. *inaequalis* Rhodes [8, Pa element POR-V-091A; 9, M element POR-V-093A]; 10, *Plectodina* sp. Sa element, CC-II-164A; 11-13, *Plectodina* cf. *flexa* (Rhodes) [11, Pa fragmentary element POR-V-163A; 12, Sa element POR-V-162A; 13, Sc element POR-V-161A]; 14-15, *Baltoniodus* aff. *variabilis* (Bergström) [14, Sa element POR-V-103A; 15, M element POR-V-108A]; 16, *Baltoniodus* sp. Sa element POR-V-104A; 17, *Icriodella* cf. *superba* Rhodes, fragmentary Pa element LC-IV-109; 18, *Panderodus* sp. LC-IV-102; 19-20, *Hamarodus europaeus* (Serpagli) [19, Sc element HZ-IA/6-6567L; 20, M element HZ-IA/7-2735G]; 21, *Scabbardella altipes* (Henningsmoen), a element HZ-IA/6-2547F; 22, *Panderodus gracilis* (Branson and Mehl), ?graciliform element RN-X-6564L; 23-26 and 32, *Saggittodontina robusta* Knüpfel [23, Pb element CO-B/5-2073F; 24, Sb element HZ-IA/6-1708D; 25, indeterminate element HZ-IB/III-1709D; 26, Pb? element HZ-IA/6-1711D; 32, Pb element CT-III/7-1]; 27-30, *Amorphognathus ordovicicus* Branson and Mehl [27, M element HZ-IB/VIII-2743G; 28, Pb element (left) HZ-IA/6-885H; 29, Pb element (right) HZ-IA/6-951H; 30, Sb element HZ-IA/6-965H]; 31, *Eocarniodus gracilis* (Rhodes), CO-A/8-2293F; 33-34, *Istorinus erectus* (Knüpfel) [33, CT-I/1-36; 34, CS-VIII/A-6-7]; 35-39, *Walliserodus amplissimus* (Serpagli), elements in inner and outer lateral views [35, a element O3A-69-58; 36, b element O3A-69-60; 37, c element O3A-69-50; 38, d element O3A-69-40; 39, e element O3A-69-57]. Scale bars=100 µm (figs. 1-34) and 200 µm (figs. 35-39)



Amorphognathus superbus Zone

In the Ossa-Morena Zone, conodonts that are attributed with doubts to the *A. superbus* Zone were obtained from a redeposited horizon occurring in the basal part of the Pelmatozoan Limestone Formation from the northern flank of the Valle syncline, west of Cazalla de la Sierra, Seville (Robardet et al. 1998, Sarmiento et al., 2000c). The conodont assemblage includes few but well preserved elements of *Icriodella* sp., *Amorphognathus* aff. *complicatus* Rhodes, *Plectodina* sp. and *Aphelognathus*? sp.

In the Central Iberian Zone, Del Moral (2004, 2007) described and illustrated conodonts of the *Amorphognathus superbus* Zone coming from the upper beds of the Bancos Mixtos Formation in the sections of Corral de Calatrava, Huertezuelas and Viso del Marqués (Fig. 1, loc. 15). The conodont record includes *Amorphognathus superbus* (Rhodes), *Sagittodontina robusta* Knüpfer, *S. cf. robusta*, *Dichodella*? sp., *Icriodella superba* Rhodes and *I. cf. superba*.

Amorphognathus ordovicicus Zone

Pioneering conodont studies developed in Spain before 1990 (see Introduction) were centered in a single Ordovician limestone deposit found throughout the Iberian Peninsula, and that was correctly attributed to the middle to upper Katian (Ka3-4) *Amorphognathus ordovicicus* Zone, and referred to the Kalodvorian stage of the Mediterranean regional scale.

The common assemblage from this Zone in Spain (Pl. 1, figs. 19-34) is represented by *Amorphognathus ordovicicus* Branson and Mehl, *Eocarniodus gracilis* (Rhodes), *Hamarodus europaeus* (Serpagli), *Istorinus erectus* Knüpfer, *Panderodus gracilis* (Branson and Mehl), *Sagittodontina robusta* Knüpfer, *Scabbardella altipes* (Henningsmoen), and several species identified in open nomenclature belonging to the genera *Drepanoistodus*, *Nordiodus*, *Panderodus*, *Protopanderodus*, *Pseudooneotodus* and *Walliserodus*?

In NW Spain, records of this conodont zone occur in an unnamed limestone near Portilla de Luna (Fig. 1, loc. 2) in the Cantabrian Zone (Del Moral, 2003; Del Moral et al., 2003), as well as in the La Aquiana and Casaio formations of the Ollo de Sapó domain of the northern Central Iberian Zone (Fig. 1, locs. 3-5), also from limestone pebbles redeposited in the Hirnantian shales of the Rozadais Formation (Sarmiento, 1993; Sarmiento et al., 1999).

In the Iberian Cordillera conodonts of the *A. ordovicicus* Zone were identified in the Cystoid Limestone (Fig. 1, loc. 6) and Ojos Negros formations (Fig. 1, loc. 7) by Sarmiento (2002) and Del Moral (2007), as well as from limestone pebbles and dropstones incorporated into the Hirnantian Orea Formation.

In the Central Iberian Zone, the Urbana Formation is a characteristic and widespread limestone unit, in spite of its lensoid character, the scarcity of its outcrops and its reduced and highly variable thickness. More than twenty sections belonging of twelve localities, mainly in the southern Central Iberian Zone (Fig. 1, locs. 9-16), were carefully sampled for conodonts since 1990, which constitutes a very important increase in our knowledge of Katian conodont faunas in the Iberian Peninsula (Sarmiento, 1990, 1993; Sarmiento et al., 2000b; Del Moral, 2002a, 2002b, 2007; Del Moral and Sarmiento, 2008) (Pl. 1, figs. 19-34). In the Portuguese part of the Central Iberian Zone, conodonts from the *A. ordovicicus* Zone have been also recorded in the Poiaras Member of the Ferradosa Formation, Serra do Buçaco (Sarmiento et al., 2000a, 2001).

In the southern border of the Central Iberian Zone, a few conodont elements obtained from rare Katian pebbles in a Mississippian olisthostrome from the Adamuz area (Córdoba: Fig. 1, loc. 17) were assigned, with doubts, to the *A. ordovicicus* Zone.

In the Ossa Morena Zone, conodonts indicative of the *A. ordovicicus* Zone have been recorded in the Pelmatozoan Limestone Formation of the Valle and Cerrón del Hornillo synclines, northern Seville province (Fig. 1, loc. 18-19) by Sarmiento (1993) and Sarmiento et al. (2008).

In NE Spain, conodonts of the *A. ordovicicus* Zone have been identified in the Pyrenees (Fig. 1, loc. 21) and in the Catalanian Coastal Ranges (Fig. 1, loc. 22), occurring in the El Baell Formation of the Freser valley (Sanz-López and Sarmiento, 1995) and in the Madremanya limestones of Les Gavarres massif (Sarmiento et al., 1995b), respectively.

All the mentioned Late Ordovician conodont occurrences of the *A. ordovicicus* Zone, with the exception of that from the Pyrenees and others with very low number of specimens, have been ascribed to the Mediterranean Province of the North Atlantic Realm, by the presence of the typical genera *Sagittodontina* and *Istorinus* (Sweet and Bergström, 1984).

Recently, a peculiar conodont fauna dominated by simple cones of *Walliserodus amplissimus* (Serpagli) and *Scabbardella altipes* was found in the Malaguide Complex of the Betic Cordillera (Fig. 1. loc. 20; Pl. 1, figs. 35-39), and referred to the extension of the *A. ordovicicus* Zone into the Hirnantian stage. This conodont association significantly differs in composition from older assemblages (Ka2-3) of the same Zone occurring in the remaining areas of the Iberian Peninsula, but it is very similar to the assemblage recorded from the upper levels of the Uggwa limestones of the Carnic Alps, thus suggesting close palaeogeographical relationships between the Malaguide Complex and the Alps (Rodríguez-Cañero et al., 2010). This Betic fauna represents the youngest Ordovician conodont record in SW Europe and the first Ordovician conodonts found in the Western Mediterranean Alpine Orogen.

CONCLUSIONS

Ordovician conodonts from the Iberian Peninsula account for the following record: early Tremadocian *Paltodus deltifer* Zone, the early Dapingian *Baltoniodus triangularis* Zone, some Floian forms are under study, successive poorly known assemblages from the latest Darriwilian to early Sandbian, and late Sandbian to early Katian intervals, the Katian *Amorphognathus superbus* Zone and the mid-late Katian to Hirnantian *Amorphognathus ordovicicus* Zone. The very sporadic and incomplete record of conodont associations previous to those from the *A. ordovicicus* Zone is common in all areas placed at high latitudes near the south polar margin of the Gondwanan continent. Nevertheless, these occurrences can be used for regional correlations and can provide palaeogeographical inferences for this characteristic domain of the North Atlantic Conodont Realm. CAI data derived from Ordovician conodonts occurring in several places of the Iberian Massif were summarized by Sarmiento and García-López (1996) and Sarmiento et al. (1999).

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